

# Observational constraints on climate sensitivity derived from the 1971-2017 global energy budget



Jonathan Chenal, Benoit Meyssignac, Aurélien Ribes, Robin Guillaume-Castel

✉ jonathan.chenal@legos.obs-mip.fr – jonathanchenal@yahoo.fr

LEGOS, Université de Toulouse - ENPC

## Summary

### Approach

- 1 Regression on **observational data** from 1971 to 2017
- 2 Include all sources of **observational uncertainty**
- 3 Quantify and include the **pattern effect**

### Results

- 1 We find an  $CO_2\text{effCS}$  of **5.5 K** with a 5-95% interval of 2.4-35.6 K
- 2  **$CO_2\text{effCS}$  below 2.4 K is inconsistent with the observed energy budget**
- 3 The upper bound is not constrained by observations
- 4 The main observational sources of uncertainty are 1. the **aerosol forcing** and 2. the **earth energy balance** estimate from the ocean heat content

### Comparison with previous estimates

- 1 Our lower bound estimate is **1.2 K** above Lewis and Curry (2018) and **0.4 K** above Sherwood et al. (2020)
- 2 The difference with Sherwood et al. is explained by the high ocean heat uptake of  $+0.2W/m^2$  they use as the reference state in 1860
- 3 The further difference with Lewis and Curry is explained by the fact that, in addition, they ignore the pattern effect and use gaussian aerosol forcing from AR5

### References

Bellouin et al. (2020). Bounding global aerosol radiative forcing of climate change. *Reviews of Geophysics*, 58(1), e2019RG000660.  
 Chenal et al. (2022). Observational constraint on the climate sensitivity to atmospheric CO<sub>2</sub> concentrations changes derived from the 1971-2017 global energy budget. *J. Clim.*, in press.  
 Cheng et al. (2017). Improved estimates of ocean heat content from 1960 to 2015. *Science Advances*, 3(3), e1601545.  
 Cowtan & Way (2014). Coverage bias in the HadCRUT4 temperature series and its impact on recent temperature trends. *Quarterly Journal of the Royal Meteorological Society*, 140(683), 1935-1944.  
 Forster et al. (2021). The Earth's energy budget, climate feedbacks, and climate sensitivity. In *Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press. In Press.  
 Gregory et al. (2020). How accurately can the climate sensitivity to CO<sub>2</sub> be estimated from historical climate change?. *Climate Dynamics*, 54(1), 129-157.  
 IPCC (2014). *Climate change 2014: synthesis report. Contribution of working groups I, II, III to the Fifth assessment report of Intergovernmental Panel on Climate Change*. C. W. Team, R. K. Pachauri, and L. A. Meyer, Eds., IPCC, Geneva, Switzerland.  
 Ishii et al. (2017). Accuracy of global upper ocean heat content estimation expected from present observational data sets. *Sola*, 13, 163-167.  
 Lewis and Curry (2018). The impact of recent forcing and ocean heat uptake data on estimates of climate sensitivity. *Journal of Climate*, 31(15), 6051-6071.  
 Richardson et al. (2021). Reconciled climate response estimates from climate models and the energy budget of Earth. *Nature Climate Change*, 6(10), 931-935.  
 Sherwood et al. (2020). An assessment of Earth's climate sensitivity using multiple lines of evidence. *Reviews of Geophysics*, 58(4), e2019RG000678.

### Earth Energy Budget

On a global scale, the Earth Energy budget can be written with the **forcing/feedback framework**, assuming a linear radiative response with surface temperature:

$$N = F + \lambda T$$

The **equilibrium climate sensitivity (ECS)** is the equilibrium surface temperature increase induced by an abrupt doubling in atmospheric CO<sub>2</sub>.

$$ECS = -\frac{F_{2\times}}{\lambda}$$

ECS is reached only at equilibrium: it cannot be directly measured in the real world and requires extremely long runs to be simulated.

### Effective Climate sensitivity

In practice the climate sensitivity is derived from the transient regime and is called effective climate sensitivity ( $effCS$ ).

$$effCS = -F_{2\times} \frac{T}{N - F}$$

histeffCS	$CO_2\text{effCS}$
from the historical energy budget	from 150yr of abrupt $2\times CO_2$

$CO_2\text{effCS}$  is a precise estimate of the ECS (although biased by ~17%).

**The aim of this study is to derive an observational constraint on  $CO_2\text{effCS}$**

### The Pattern Effect

The radiative response of the Earth depends on the geographic pattern in surface air temperature. (Sherwood et al. 2020; Gregory et al. 2020)

This effect is called **the pattern effect**. It arises from:

- ▶ Mix of radiative forcings
- ▶ Lag-dependent responses to forcings
- ▶ Unforced variability.

The pattern effect leads to apparent time variations in  $\lambda$  and thus:

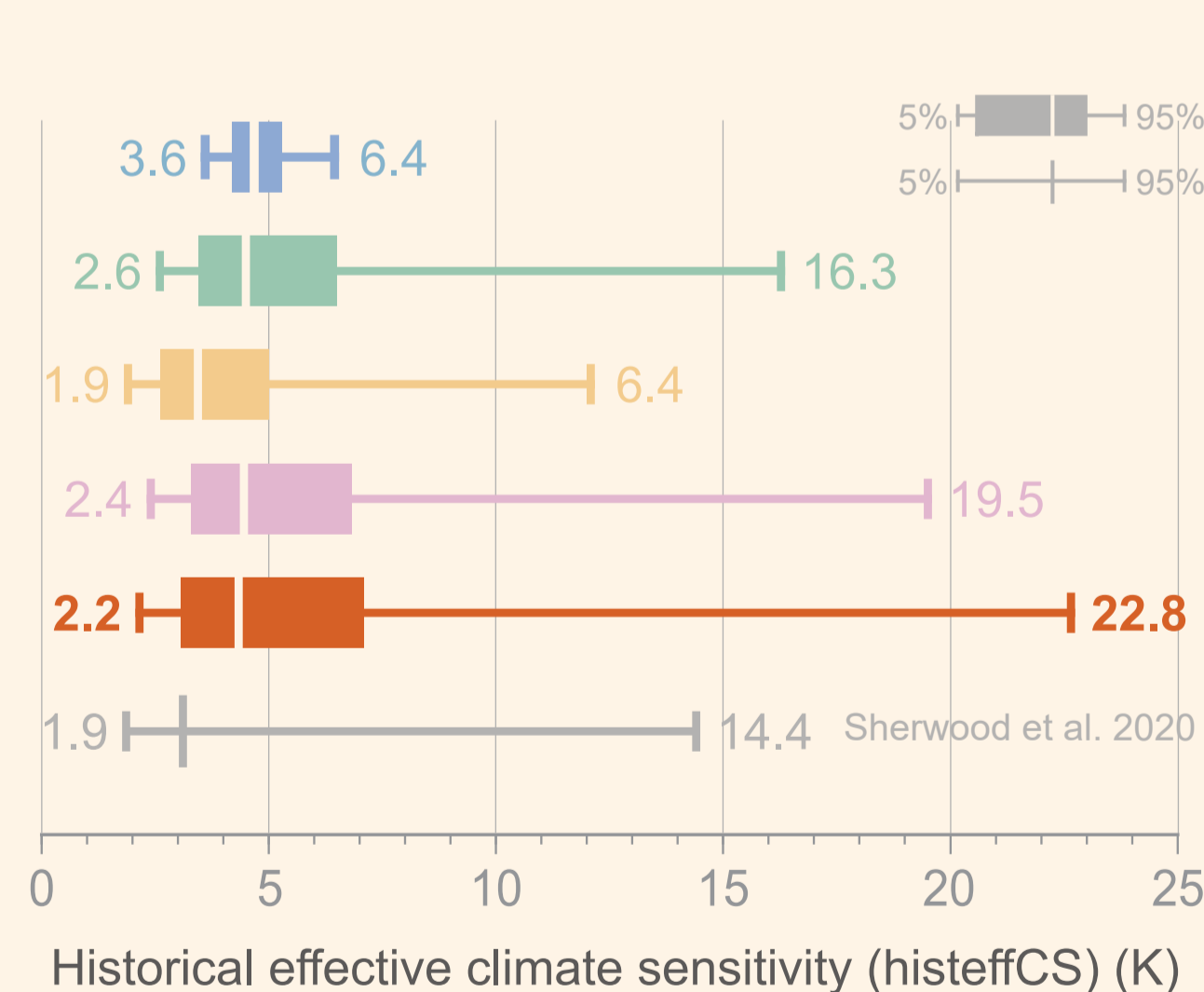
$$histeffCS \neq CO_2\text{effCS}$$

### Observational constraints on $CO_2\text{effCS}$

- Compute  $histeffCS$  using a regression on observations of the energy budget from 1971-2017:
  - ▶ **F non aerosol ( $F_{NA}$ )** from Sherwood et al. 2020
  - ▶ **F aerosol ( $F_{AER}$ )** from Bellouin et al. 2020
  - ▶ **T** from the Cowtan and Way 2014 corrected for the surface bias due to satellite data (Richardson et al. 2016)
  - ▶ **N** from Ocean heat content estimates derived from optimal interpolation of ocean in situ data (Ishii et al. 2017, Cheng et al. 2017)
- Include all sources of observational uncertainty
- Quantify the pattern effect from CMIP6:
  - ▶ Internal variability i.v. in  $\lambda$
  - ▶ A distance between  $histeffCS$  and  $CO_2\text{effCS}$

## Main results

### histeffCS and associated uncertainties

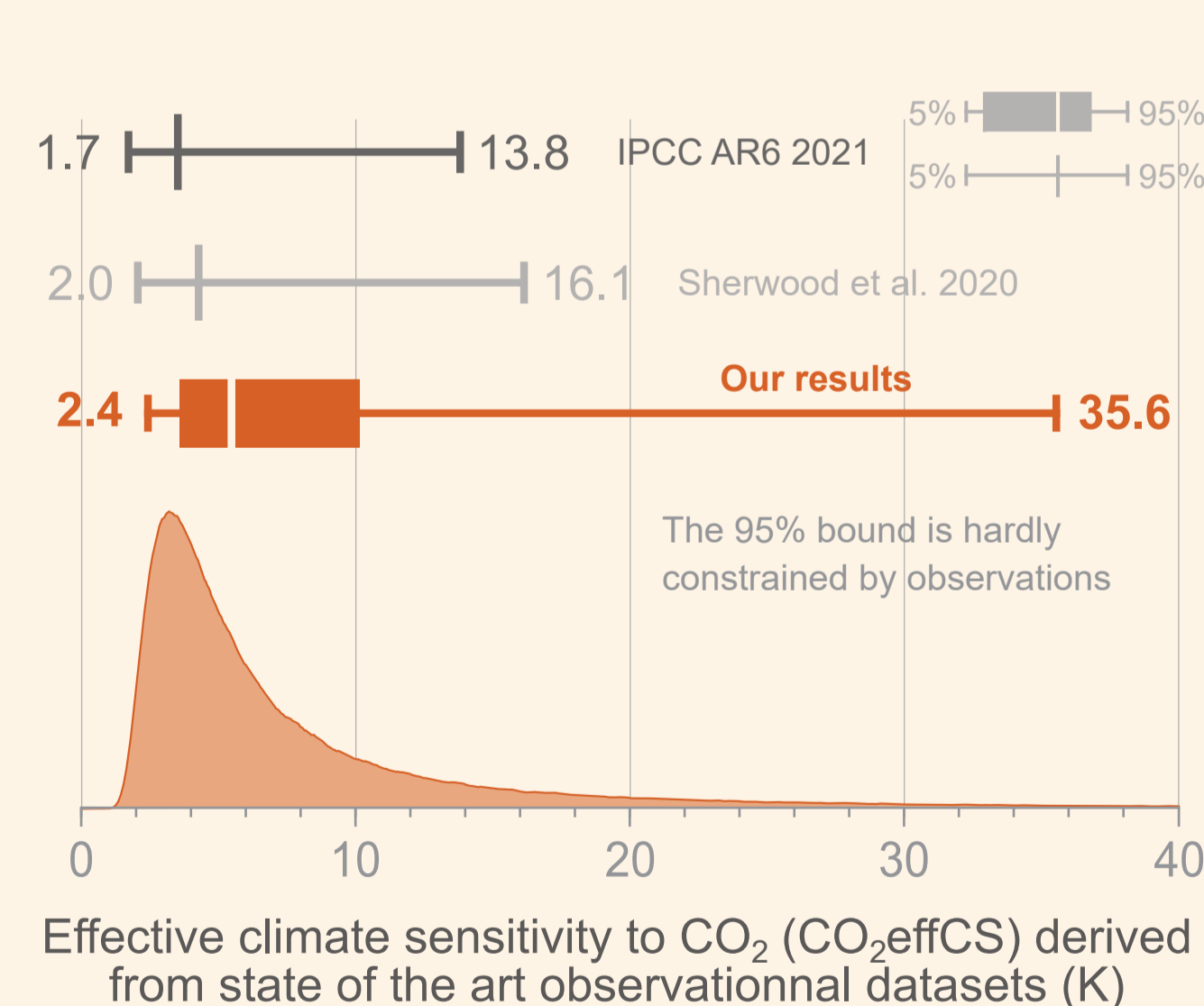


Uncertainties included:

- i.v. (Internal variability)
- N (Ocean heat content)
- $F_{AER}$  (Aerosol forcing)
- $F_{NA}$  (Non-aerosol forcing)

**The major sources of uncertainty in  $histeffCS$  are first the aerosol forcing and then the ocean heat content**

### $CO_2\text{effCS}$



**The observed 1971-2017 Earth energy budget is inconsistent with a  $CO_2\text{effCS}$  lower than 2.4°C**

### Comparison with previous studies

The two most recent studies estimating climate sensitivity from observational data are **Lewis and Curry (2018)** and **Sherwood et al. (2020)**

$CO_2\text{effCS}$ estimates:	5%	Med.	95%
<b>Lewis &amp; Curry (2018)</b>	1.2	1.7	2.7
<b>Sherwood et al. (2020)</b>	2.0	4.3	16.1
<b>Chenal et al. (2022)</b>	2.4	5.5	35.6

L&C18 use the IPCC AR5 gaussian aerosol forcing and do not correct for the **pattern effect**

Both L&C18 and S20 use a state difference method with  $N(\sim 1860) = +0.2 W/m^2$  but it is an uncertain value (probably  $< 0?$ ). The regression approach removes this dependence

The higher bound difference between our estimate and S20 is insignificant