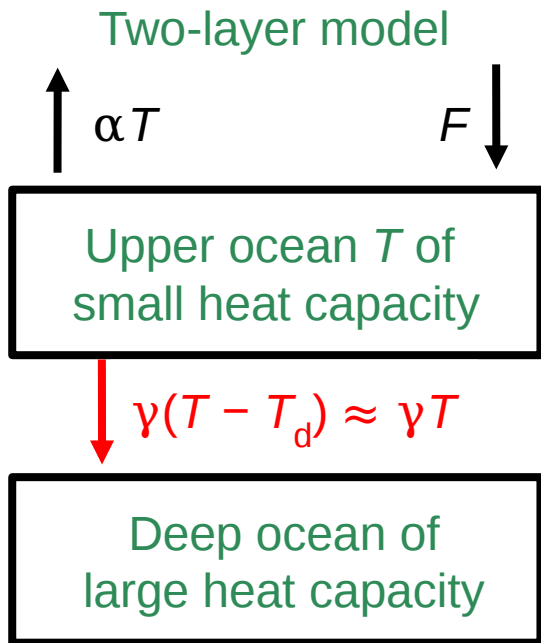


A new conceptual model of global ocean heat uptake and transient climate response

Jonathan Gregory *et al.*

See end of poster for list of authors



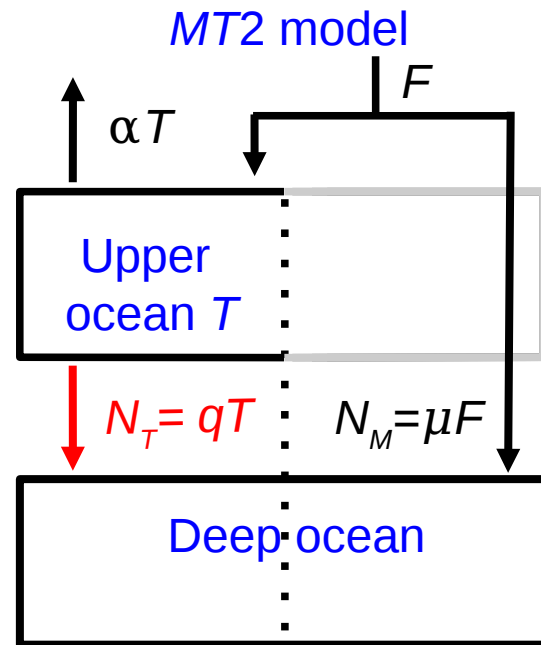
γ is AOGCM-specific.

Net downward TOA flux
 $= N = F - \alpha T$
 $=$ Heat flux into ocean

Ocean heat uptake efficiency
 $\text{OHUE} \equiv N \div T$

$\leftarrow N = \gamma T$
 $\Rightarrow N$ is small if T is small,
 and N correlates with γ
 Neither is true in CMIP.

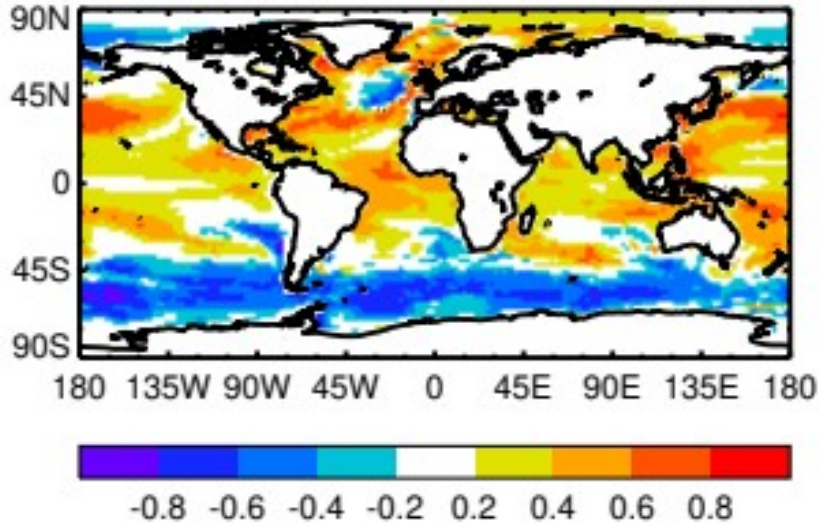
$N = qT + \mu F \rightarrow$
 CMIP T anticorrelates with μ ,
 so $\text{OHUE} = q + \mu F / T$
 $q(t)$ is AOGCM-independent,
 is larger with stronger AMOC. μ correlates with AMOC.



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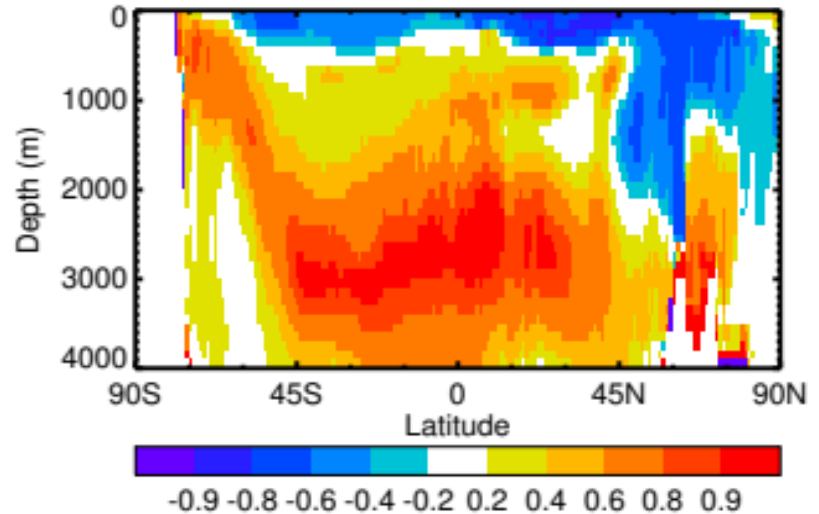
Interpretation and implications of the *MT2* model

Correlation of Δ SST with global-mean N_T



N_T occurs at low latitude

Correlation of passive heat with AMOC



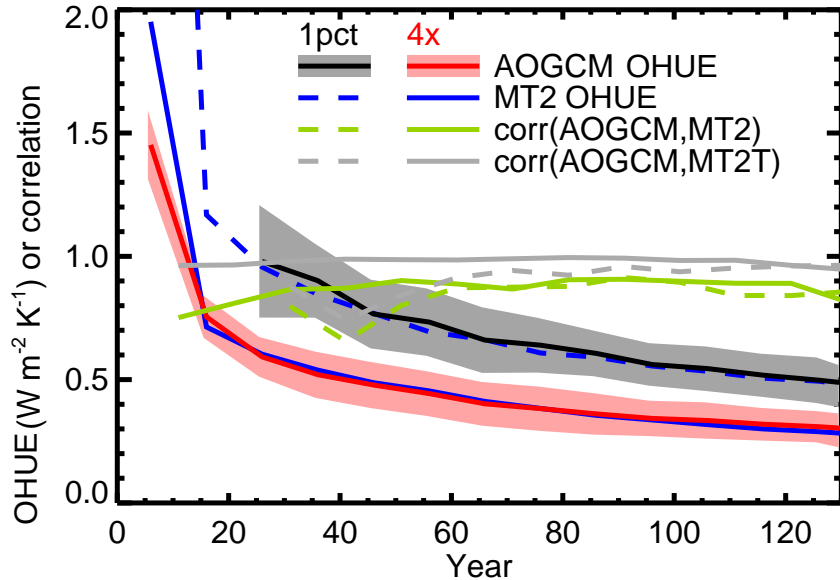
N_M occurs at high latitude

AOGCMs with stronger AMOC have smaller T because α is larger (EffCS smaller) and heat is removed more efficiently from the surface (but **not by** the AMOC).

$$F - \alpha T = N = qT + \mu F \Rightarrow T = F(1 - \mu)/(\alpha + q)$$

$\Rightarrow T$ determined almost entirely by $\alpha \Rightarrow T$ correlates with EffCS (Grose *et al.*)

MT2 model for time-dependent ocean heat uptake efficiency



$N_M(t) = \mu F(t)$, where μ is constant in time and AOGCM-specific, $\mu = s_0(M - M_0)$, $s_0 = 0.5\% \text{ Sv}^{-1}$, $M_0 = -10.2 \text{ Sv}$.

$N_T(t) = q(t) T(t)$ is described by the two-layer model, with AOGCM-independent heat capacities and $\gamma = 0.47 \text{ W m}^{-2} \text{ K}^{-1}$.
 $\text{OHUE } N/T = (\mu\alpha + q)/(1 - \mu)$
 decreases in time because α and q do.

MT2 reproduces OHUE well in AOGCM mean (black, red and blue) and individual AOGCMs (green); MT2T is better (grey, AOGCM-specific deep-layer thickness).

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