# **3D** reconstruction of upper ocean dynamics in the Nordic and Beaufort Seas. Assessment of the Surface Quasi-Geostrophic Approach

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### **Motivation**

Polar regions have been changing since the 1990s; Better estimates of 3D ocean circulation are required to improve our understanding of polar dynamics and to better comprehend the impact of the changes caused by climate change. Our objective is to see if remotely sensed variables may be used to reconstruct 3D ocean dynamics in Arctic and sub-Arctic Seas.

We have assessed the capability of Surface Quasi-Geostrophy (eSQG) to reconstruct the three-dimensional (3D) dynamics in two key areas of the Arctic

### Results

### Surface currents in Nordic Seas



SSB is not able to reconstruct key surface mesoscale currents seen in SSH in the Nordic and Beaufort Seas.

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#### Ocean: the Nordic and the Beaufort Seas using TOPAZ4 outputs.



### Surface Quasi-Geostrophic model

SQG model allows to derive the stream function at each depth from sea surface height *(η)*, surface from the buoyancy (b), or from surface velocities (v)through the effective SQG relations (Isern *et. al, 2008*).

**Reconstruction from SSH :**  $\hat{\psi}_{\eta}(\vec{k},z) = \exp(n_0kz)\hat{\psi}_s(\vec{k})$ Streamfunction **Reconstruction from SSB:** Surface Density (SSS & SST)



Figure 3: Example of surface currents for the 28th August in Nordic Seas.

a) reconstruction from SSH, b) reconstruction from SSB

### **Section along Fram Strait** (Latitude = 77.5°N)





Figure 4:

Currents along Fram Strait for the 28th August. a) Reconstruction from SSH, b) reconstruction from SSB, c) reconstruction from SSV, d) TOPAZ4 geostrophic velocity, e) TOPAZ4 total velocity.



#### **Computation of Prandtl ratio**

sigma [kg/m<sup>3</sup> N<sup>2</sup> [rad<sup>2</sup>/sec<sup>2</sup>] ×10<sup>-</sup> Summer months 2.5 higher exhibit -100 -100 stratification, which → JAN → FEB -200 associated to IS -MAR APR freshwater inflows E -300 E -300 Depth [ JUN from ice melting and hth 100-700 → JAN -- FEB AUG warmer temperatures. -OCT MAY ►NOV -500 -500 -JUL AUG Ocean mixing causes -SEP -600 -600 --NOV lower stratification in -- DEC -700 -700 the spring and winter Figure 2: Mean monthly profiles of Brunt Vaisala months.

frequency and density in Nordic Seas

The daily-mean Prandtl ratio  $(n_0)$  was calculated by dividing the mean Brunt-Vaisala frequency in the first 100 meters by the mean Coriolis frequency.

Longitude [°]

Ocean 3D dynamics are reconstructed from surface buoyancy, surface height and surface velocities and compared to model geostrophic current and model total current. Better results are achieved reconstructing 3D dynamics from SSH and from surface velocities.

from SSB



Figure 5: Vorticity correlations.

a) Reconstruction from SSH / Model geostrophic velocity, b) reconstruction from SSB / Model geostrophic velocity, c) reconstruction from SSV / Model total velocity.

Vorticity correlations between reconstruction form SSH and model geostrophic current, show good agreement (corr.>0.8) up to 400 meters. Vorticity correlations between reconstruction from SSV and model total currents, exhibit fairly good agreement (corr.>0.6) up to 200 meters. Reconstructions are better in the winter and spring in all locations than in the summer and fall, when the water column is less stratified and the

### Summary and conclusions

• Surface Quasi-Geostrophic (SQG) theory allows 3D dynamics in the Arctic Ocean to be reconstructed using only surface information of SSH or surface velocities, but it does not allow 3D dynamics to be reconstructed using only sea surface buoyancy. • Improved 3D reconstructions are obtained during the winter and spring months when Brunt-Vaisala frequency is the lowest. • The results of the research encourage us to use future remotely sensed SWOT and CRISTAL high-resolution sea surface height (SSH) and Seastar and WaCM direct measurements of ocean surface currents in polar areas to reconstruct 3D dynamics.

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