

Non-monotonic feedback dependence on CO₂ due to a North Atlantic pattern effect

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Abstract

We explore the effective climate sensitivity S_G with abrupt CO₂ forcing experiments, spanning the range 2×, 3×, 4×, 5×, 6×, 7×, and 8×CO₂, using the CESM Large Ensemble model configuration (Kay et al., 2015). We find that S_G is a non-monotonic function of CO₂, decreasing between 3× and 4×CO₂, and then increasing at larger CO₂. We attribute this non-monotonicity to the negative feedbacks in the North Atlantic which stem from cooling in the North Atlantic due to AMOC collapse. To isolate the importance of how the North Atlantic cooling pattern affects the net radiative feedback, we run atmosphere-only simulations of the same model with prescribed sea surface temperatures (SSTs) taken from 1) the fully coupled runs and 2) different SST patterns.

Key Points

- We find a non-monotonic response in Effective Climate Sensitivity across a range of abrupt $n \times \text{CO}_2$ forcing experiments with a minimum at 4×CO₂
- We attribute this non-monotonicity to changes in radiative feedbacks over the North Atlantic, caused by a surface cooling in that region associated with the collapse of the AMOC

Non-monotonicity of the Effective Climate Sensitivity due to radiative feedbacks, not radiative forcing

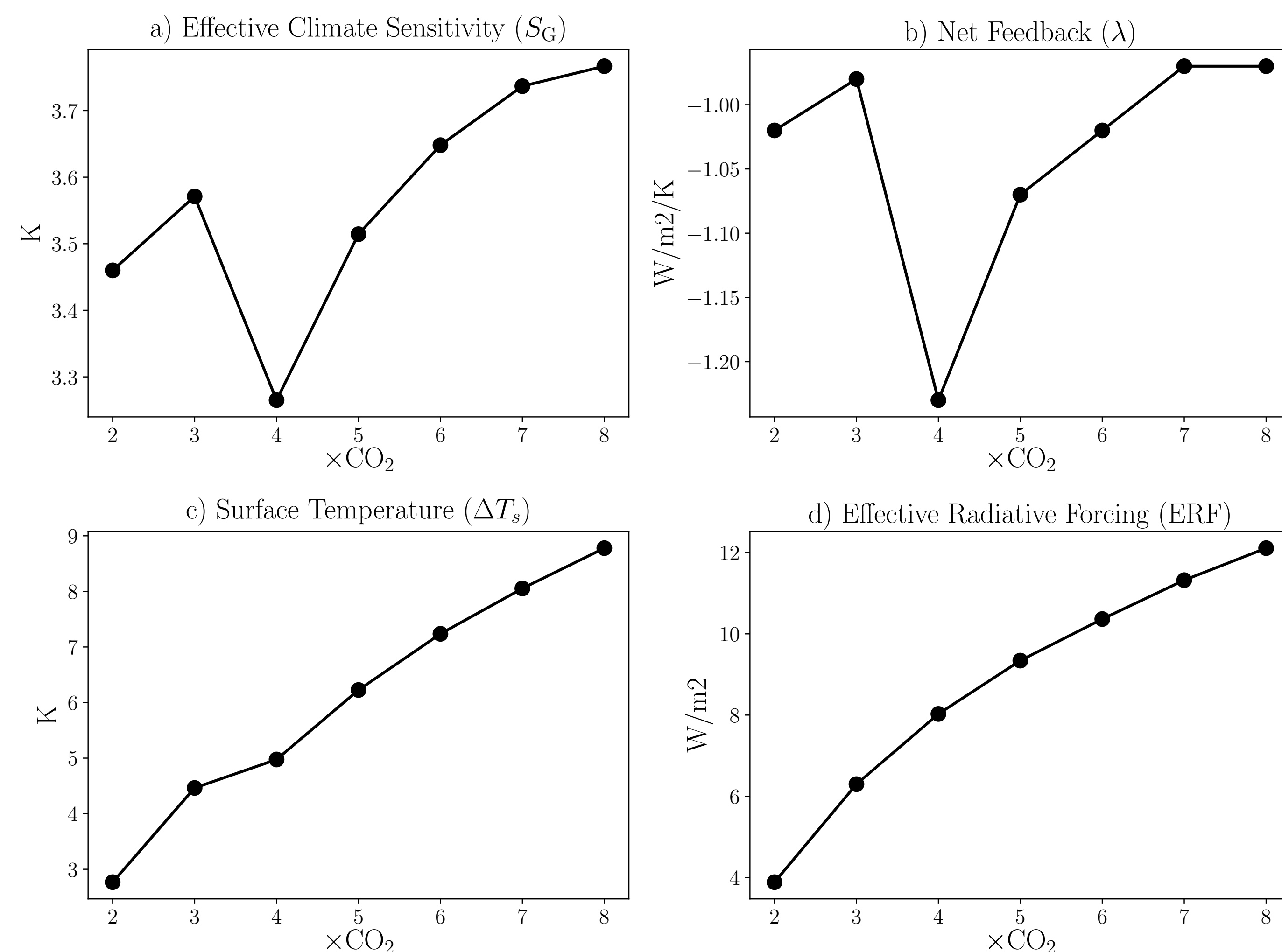


Figure 1: **a)** Effective Climate Sensitivity (S_G), and **b)** net feedback parameter (λ) from the 150 year Gregory regression of abrupt $n \times \text{CO}_2$ runs. **c)** Global mean surface temperature response, and **d)** effective radiative forcing (ERF) from 30-year fixed SST runs (Forster et al., 2016).

- We find a non-monotonic response in Effective Climate Sensitivity (Fig. 1a) with a minimum at 4×CO₂
- The net feedbacks (Fig. 1b) show same non-monotonicity, and the effective radiative forcing (Fig. 1d) does not
- We **hypothesize** that either λ responds non-monotonically due to 1) increased global-mean temperature or 2) to a different sea surface temperature pattern

Non-monotonicity in λ is due to SST pattern, not global mean surface temperature increase

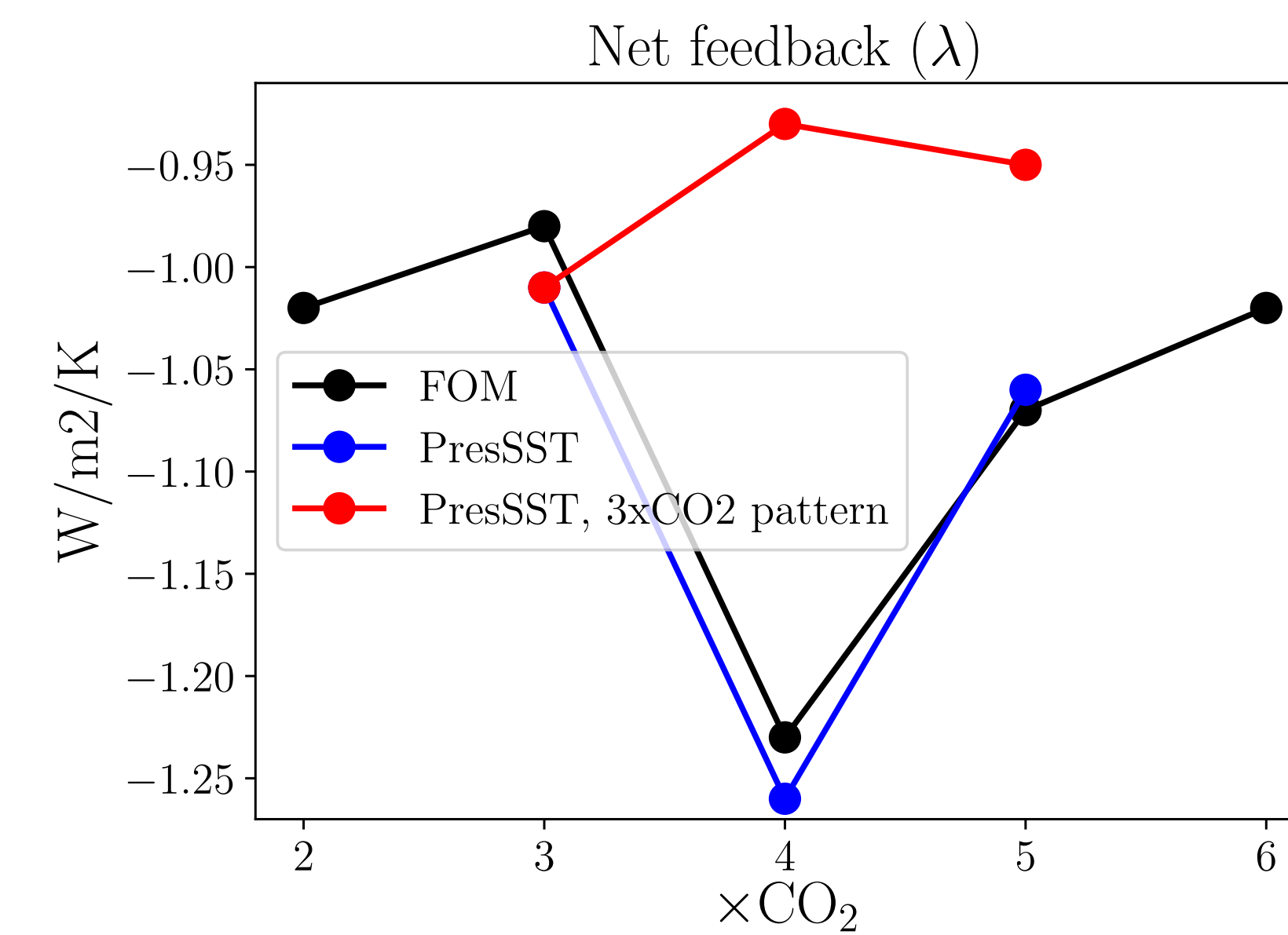


Figure 2: Net feedback parameter λ (black) from 150 year Gregory regression with fully coupled runs (FOM), AMIP runs with prescribed SSTs from fully coupled runs (blue), and AMIP runs with prescribed SSTs with 3×CO₂ warming pattern (red).

- The AMIP runs with prescribed SSTs (blue in Fig. 2) can fully reproduce λ from the fully coupled runs (black)
- We get rid of the non-monotonicity (red in Fig. 2) when we repeat the 4× and 5×CO₂ FOM experiments with AMIP runs with same global mean surface warming as in FOM but SST pattern from 3×CO₂

Cooling SST pattern in North Atlantic coincides with λ non-monotonicity

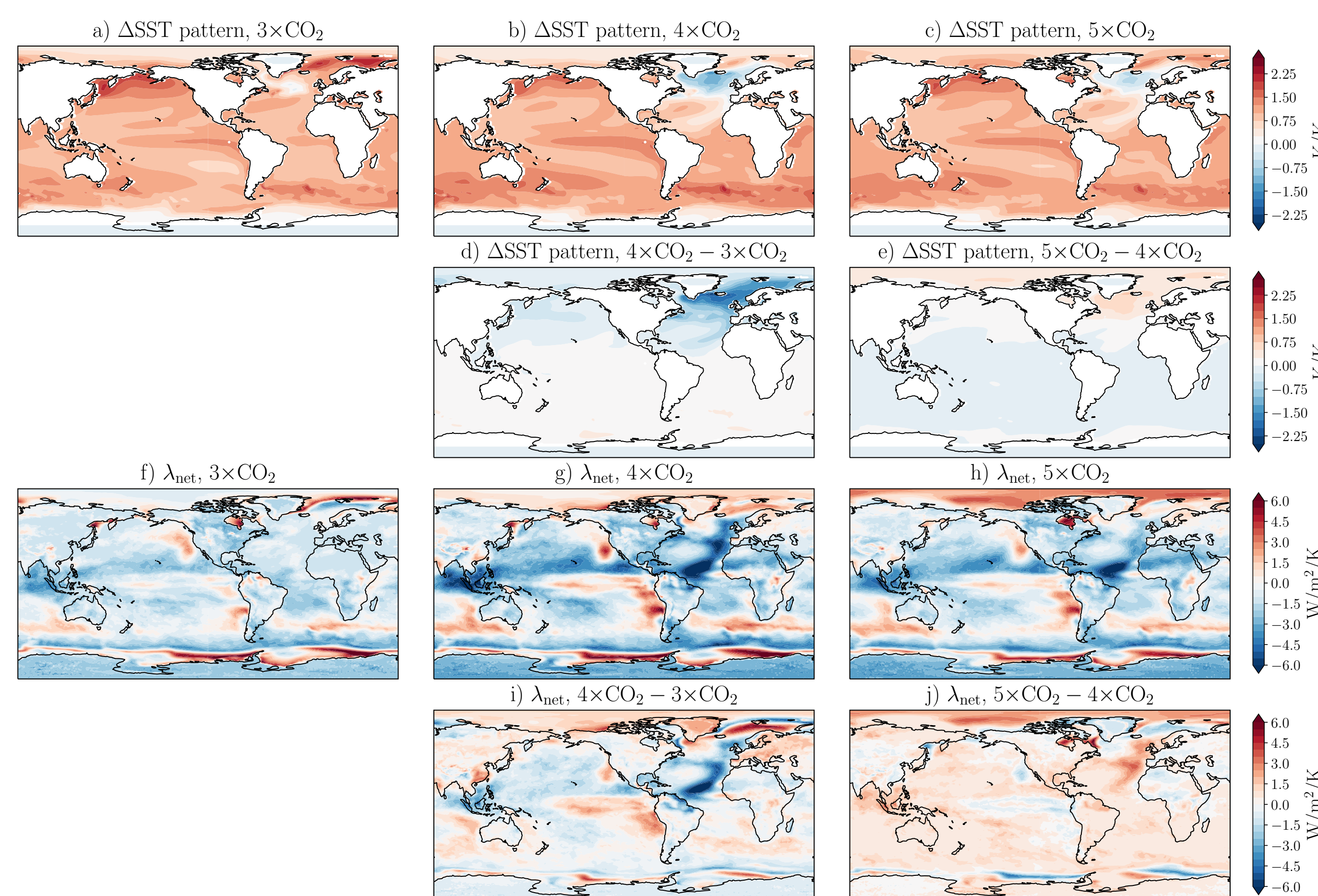


Figure 3: SST pattern in **a)** 3×CO₂, **b)** 4×CO₂, and **c)** 5×CO₂. The difference between 4× and 3×CO₂, and 5× and 4×CO₂ are shown in **d)** and **e)**, respectively. Figures **f-j)** show λ for the same CO₂ experiments.

- Cooling SST pattern in the North Atlantic between 3× and 4×CO₂ (Fig. 3d) coincides with a more negative λ (Fig. 3i) for the same region
- Similarly, “warming” SST pattern in North Atlantic between 4× and 5×CO₂ (Fig. 3e) coincides with a more positive λ (Fig. 3j)

North Atlantic cooling causes λ non-monotonicity

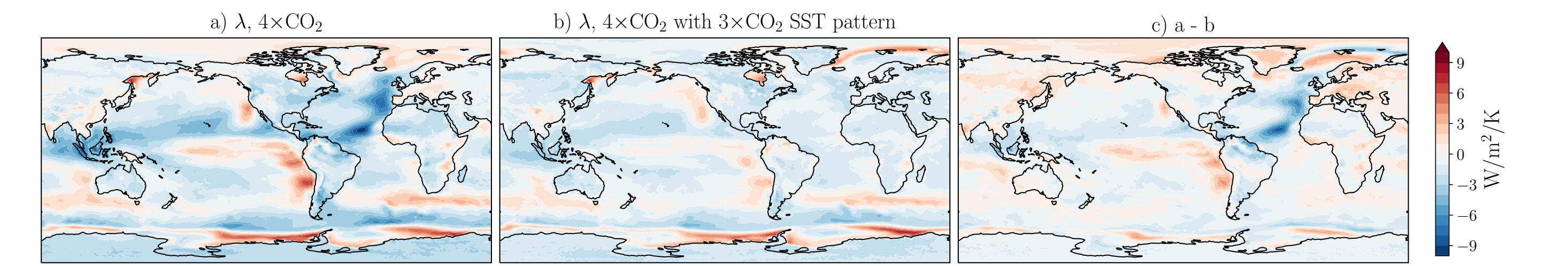


Figure 4: Net feedbacks from 4×CO₂ AMIP runs with **a)** prescribed SSTs from fully coupled runs, and **b)** prescribed global mean SSTs from fully coupled runs and 3×CO₂ warming pattern. The difference is shown in **c)**.

- Negative λ over the North Atlantic (Fig. 4a) disappears when we re-do the 4×CO₂ run with the same global mean warming but different SST pattern taken from 3×CO₂ (Fig. 4b, same as red dot at 4×CO₂ in Fig. 2)

LR and SW cloud feedbacks are most responsible

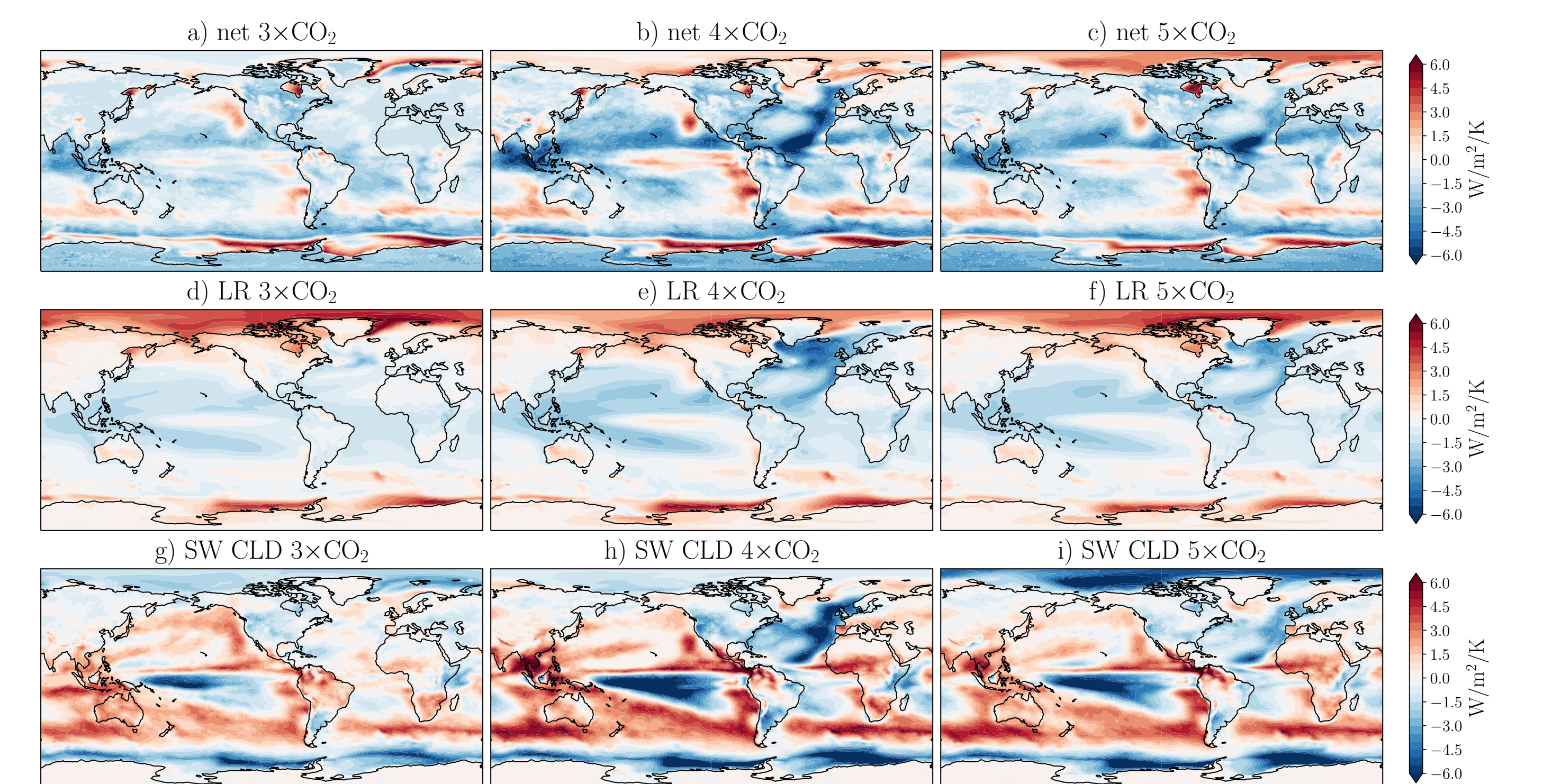


Figure 5: **a-c)** Net feedbacks λ , **d-f)** lapse rate, and **g-i)** shortwave cloud feedback. Left column shows 3×CO₂, middle shows 4×CO₂, and right column shows 5×CO₂.

AMOC collapse coincides with North Atlantic cooling

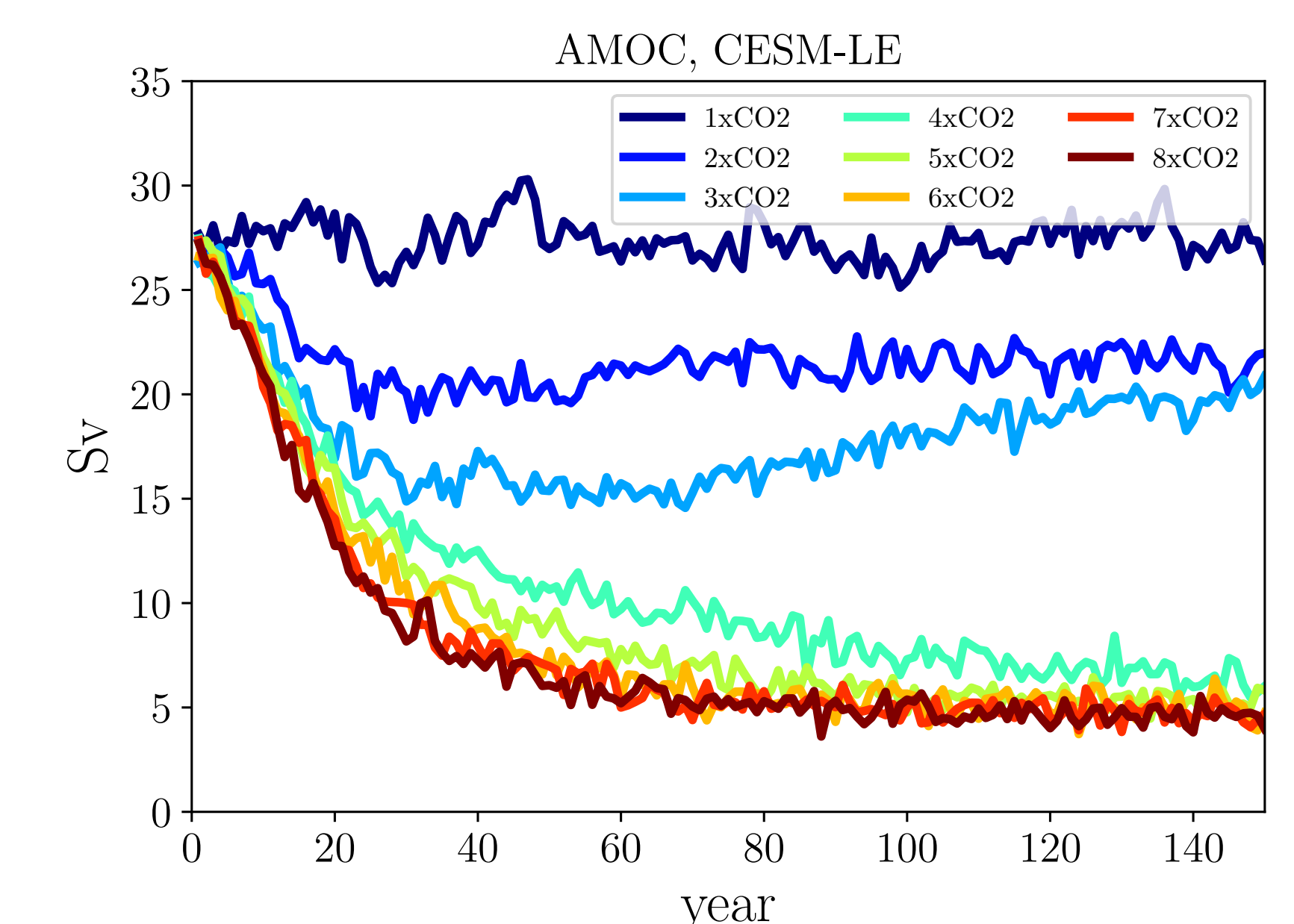


Figure 6: The evolution of the Atlantic Meridional Overturning Circulation.

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