Is the forced response of precipitation timescale-dependent?



Synopsis

Issue:

- The pattern effect leads to time-dependence of climate sensitivity via the multiple response timescales of TOA net radiative fluxes. Since the global-mean precipitation is in balance with the atmospheric radiative cooling (=TOA + surface fluxes) plus sensible heat flux, is its response to forcing timescale-dependent too?
- How does precipitation change in response to abrupt 4xCO2 forcing across different timescales: years 1-20, 21-150, and 151-1000?

Approach:

- Investigate with LongRunMIP simulations (Rugenstein et al., 2019)
- Abrupt 4xCO2 forcing, 1000+ year simulations, and pre-industrial control

Findings:

- The hydrological sensitivity has a less consistent trend across time scales compared to climate sensitivity.
- Timescale dependence of climate sensitivity can be overestimated when calculated with OLS regression for 1 ensemble member because of differing bias across timescales.

Regression methods applied to TOA radiation

Slope estimates are not always unbiased; Ordinary Least Squares (OLS) regression assumes the x-variable is known exactly (Gregory et al., 2020 discuss implications for climate sensitivity) and there is no correlation between x- and y- internal variability. Can these biases affect different timescales differently?



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The Pattern Effect: Coupling of SST Patterns, Radiative Feedbacks, and Climate Sensitivity Workshop, May 10 - May 13, 2022

Spatial Patterns in Hydrological Sensitivity Hydrologic Sensitivity $HS = \frac{L\Delta P}{\Delta T}$ 2.65 2.23 Several regions Years 21-150 have reversals in Years 1-20 sign of the precip response from years 1-20 to years 21-150 2.03 Between years 21-150 and 151-1000, the sign of the precip response Consistently generally remains Freen indicates decreasing for 4 $|m| > |m_{21-150}|,$ the same, but out of 12 models **Red** indicates magnitude of $|\mathbf{m}| < |\mathbf{m}_{21-150}|.$ moistening and drying both decrease Beyond 150 years, the magnitude of HS decreases for most models Smaller Smaller magnitude for **11** magnitude for 6 out of 12 models **References and Acknowledgements** out of 12 models 1-20 21-150 151-1000 Time Period (Years) Kao (senior thesis; in review) / Kao and Pendergrass (in prep.) References • York. "Least squares fitting of a straight line with correlated errors." Earth and Planetary Science Letters (1968) • Mahon. "The New "York" regression: Application of an improved statistical method to geochemistry." International Geology Review (1996) • Rugenstein et al. "LongRunMIP: motivation and design for a large collection of millennial-length AOGCM simulations." Bulletin of the American Meteorological Society (2019) • Gregory et al. "How accurately can the climate sensitivity to CO2 be estimated from historical climate change?" *Climate Dynamics* (2020) \mathbf{X} • Payton et al. "Overlapping confidence intervals or standard error intervals: what do they mean in terms of statistical significance?." Journal of Insect Science 3.1 (2003). Acknowledgement Thanks to everyone involved in running and coordinating the LongRunMIP simulations and for generously sharing them with us, to Jonathan sensitivity seems to be downwelling Gregory for insightful discussions about the regression technique, and to Mike Town for bringing the "New" York method to our attention and $\frac{\Delta R_{sfc}}{\Delta T} \quad \frac{\Delta LW \downarrow}{\Delta T} \quad \frac{\Delta SW \downarrow}{\Delta T} - \frac{\Delta LW \uparrow}{\Delta T} - \frac{\Delta SW \uparrow}{\Delta T}$ $\frac{\Delta P}{\Delta T} - \frac{\Delta R_{TOA}}{\Delta T} \quad \frac{\Delta R_{sfc}}{\Delta T} - \frac{\Delta SH}{\Delta T}$ shortwave radiative flux at surface for simplifying the Python function for calculating it. The original release of the "New" York regression Python code is publicly available at

github.com/LLNL/MahonFitting/











