

# Examining state dependence of the cloud feedback using a perturbed parameter ensemble

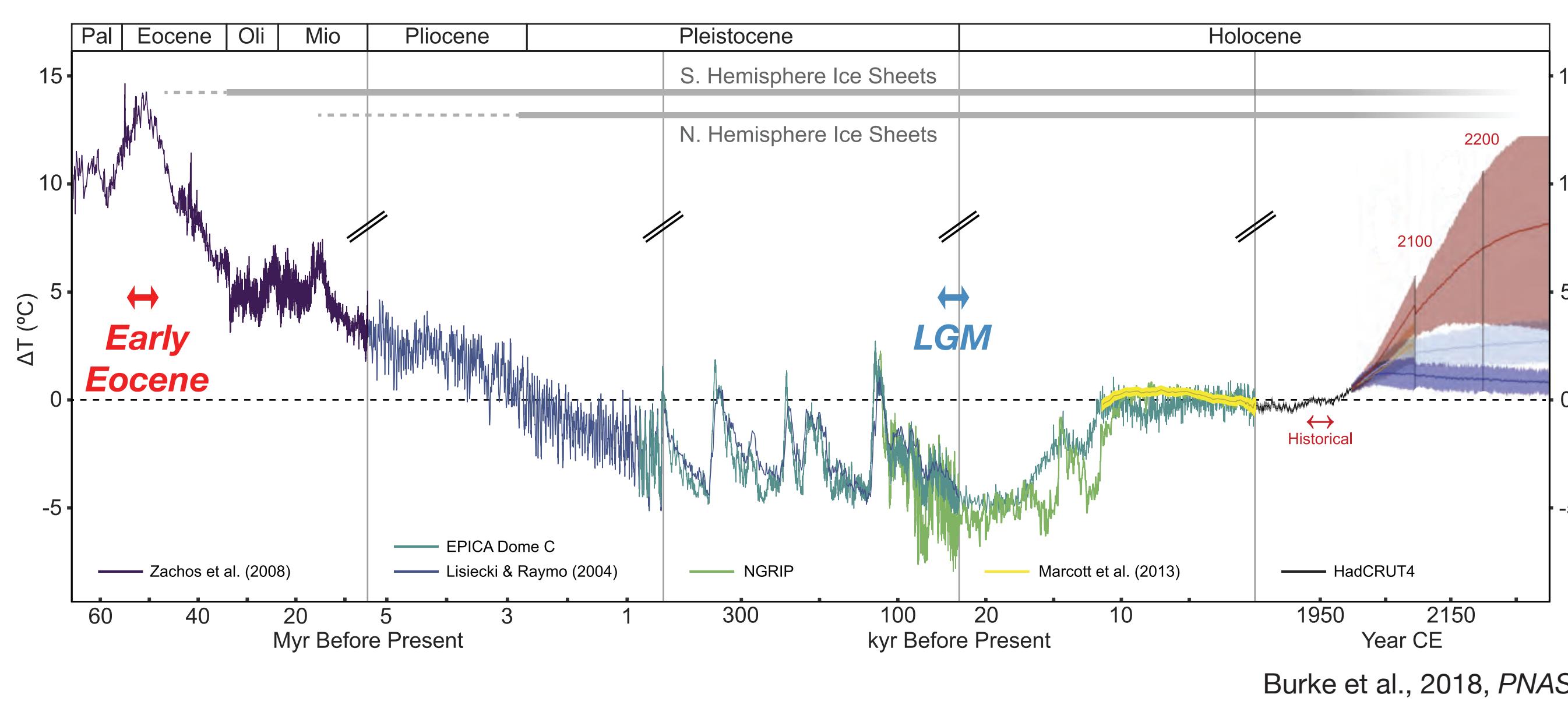
Jiang Zhu (jiangzhu@ucar.edu), Bette L. Otto-Bliesner, Esther C. Brady, Andrew Gettelman, Trude Eidhammer

National Center for Atmospheric Research

NCAR  
CGD



## I. Motivation

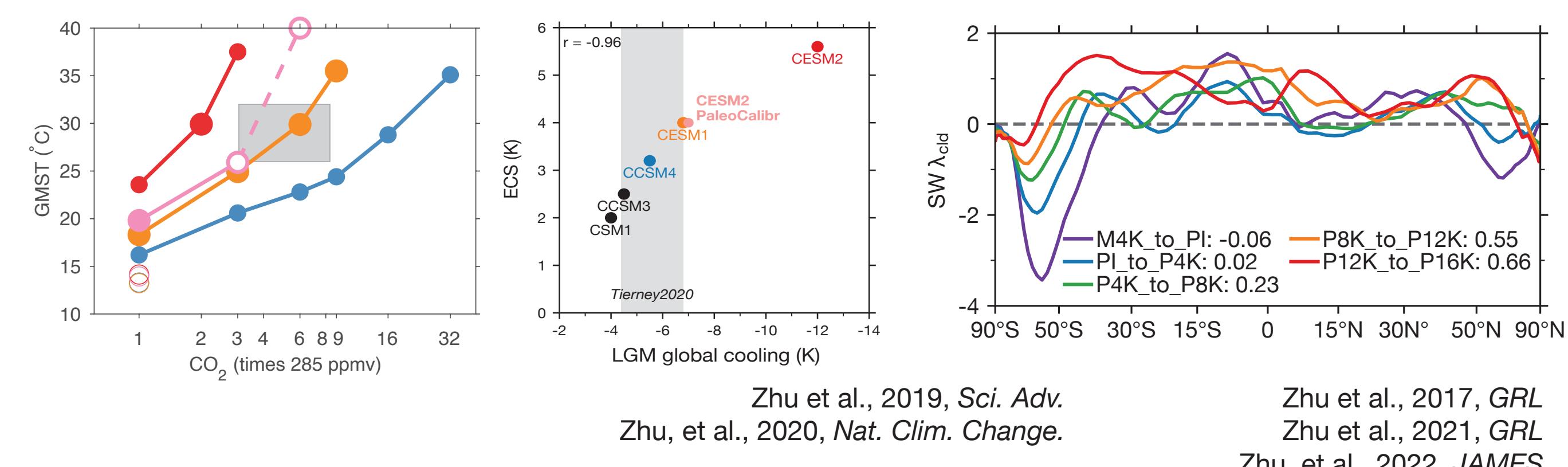


**1. Past climates inform our future; understanding of the state and forcing dependent response is critical.**

- How does the cloud feedback depend on climate states? Are there non-linear tipping points?
- How does the cloud feedback interact with temperature and its spatial pattern?

**2. Past climates inform the equilibrium climate sensitivity (ECS) in models.**

- Community Earth System Model v2 (CESM2) fails to simulate the Last Glacial Maximum (LGM) & the early Eocene (EECO); its high ECS (>5°C) is unrealistic
- An LGM-calibrated CESM2 still fails to simulate the Early Eocene
- Which processes/parameters control the nonlinear cloud feedback ( $\lambda_{\text{cl}}^*$ )?



## 2. Approach: perturbed parameter ensemble (PPE)

**1. Community Atmosphere Model v6 (CAM6) in ~2°**

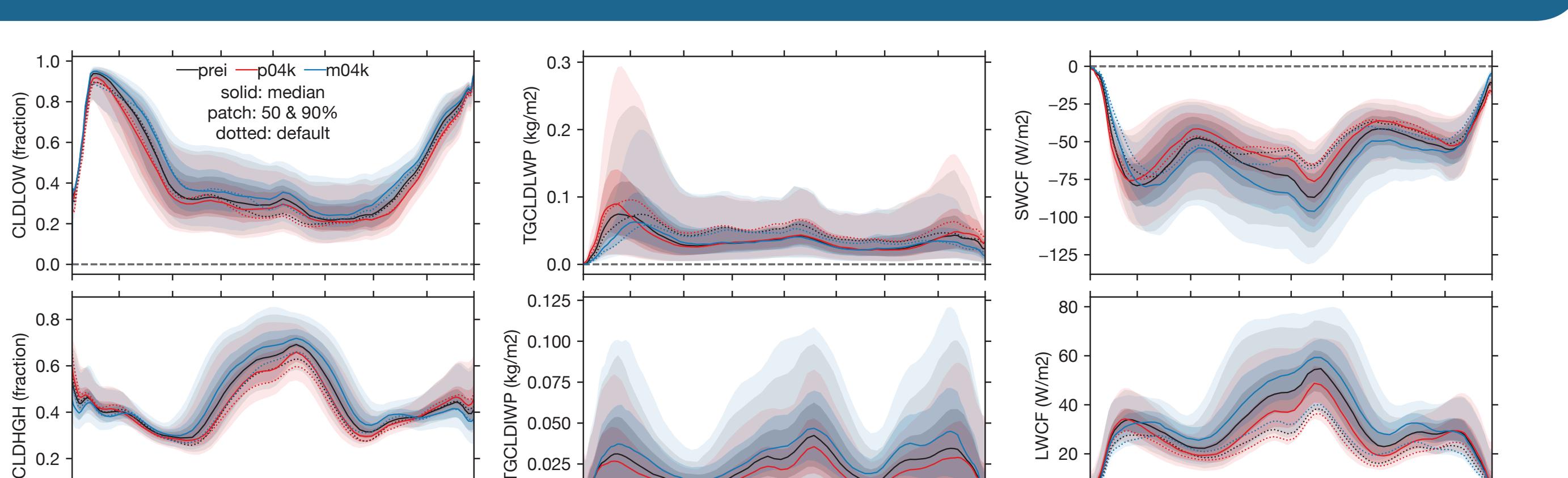
- PaleoCalibr fixes: no cloud-ice-number-limiter (nimax); shorter microphysical  $\Delta t$

**2. PPE setup**

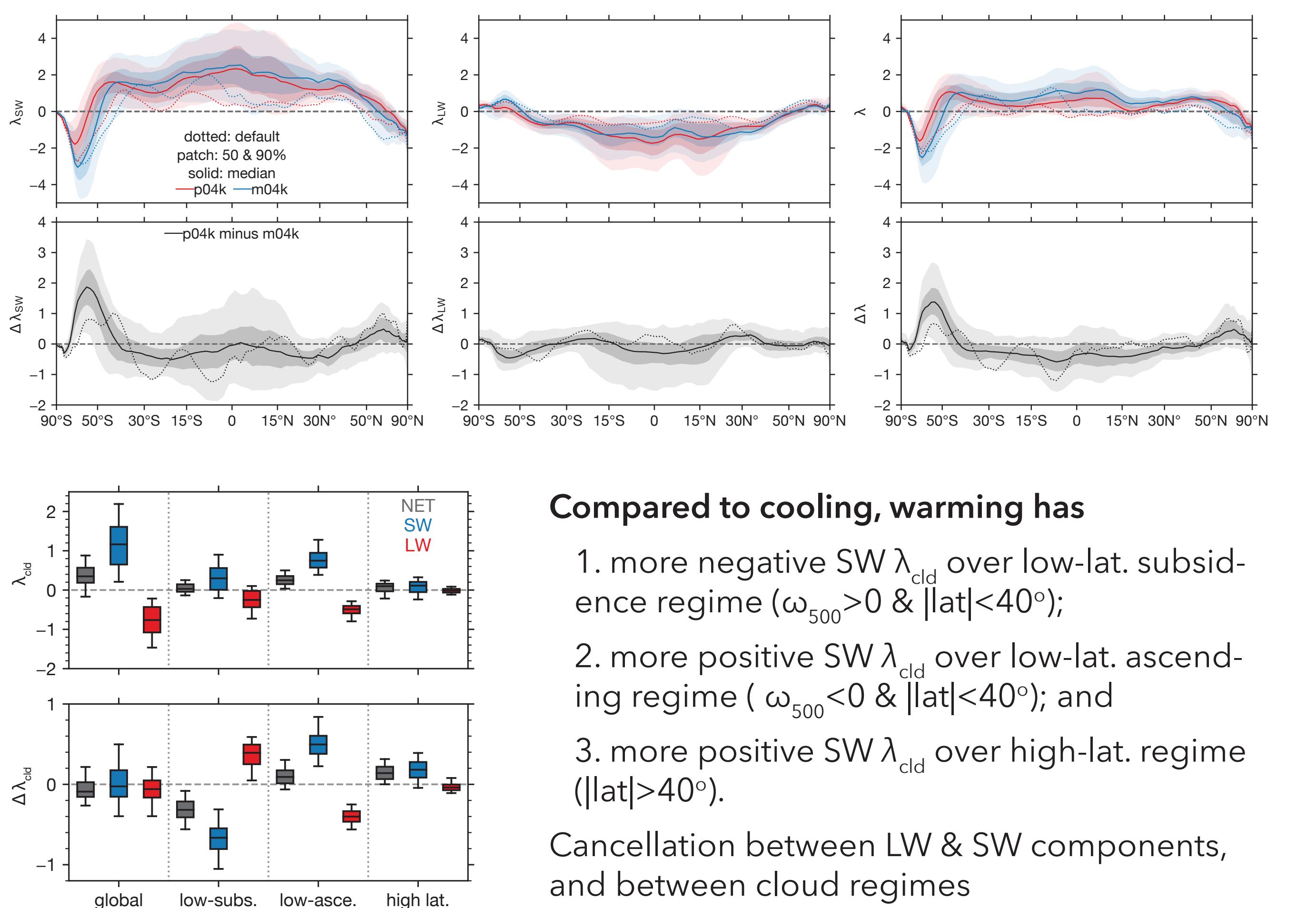
- 45 parameters in convection & cloud parameterizations (see below for a subset)
- 250 ensemble simulations (5 years each) from Latin Hypercube sampling
- Preindustrial SST (prei) and SST+4K (p04k) & -4K (m04k) with uniform change
- Future plans: SST+8K, +12K, +16K, & patterned warming

Physics	Parameter Name	Description	Default	Min	Max
CLUBB	clubb_c1	Low Skewness in C1 Skw.	1	0.4	3
	clubb_C2rt	Damping on scalar variances	1	0.2	2
	clubb_C8	Coeff. #1 in C8 Skewness Equation	4.2	1	7
	clubb_c11b	High Skewness in C11 Skw	0.35	0.2	0.8
	clubb_c14	Constant for u'2 & v'2 terms	2.2	0.4	4
MG3	micro_mg_accre_enhan_fact	Accretion enhancing factor	1	0.1	10
	micro_mg_berg_eff_factor	Bergeron efficiency factor	1	0.1	1
	micro_mg_autocon_lwp_exp	KK2000 LWP exponent	2.47	2.1	3.3
	micro_mg_dcs	Autoconversion size threshold ice-snow	2.00E-04	5.00E-05	1.00E-03
	micro_mg_vtrmi_factor	Ice fall speed scaling	1	0.2	5
ZM	zmconv_dmpdz	Parcel fractional mass entrainment rate	-1.00E-03	-2.00E-03	-2.00E-04
	cldfrc_dp1	Parameter for deep convection cloud fraction	0.1	0.05	0.25
	zmconv_tiedke_add	Convective parcel temperature perturbation	0.5	0	2
Aerosol	zmconv_capelmt	Triggering threshold for ZM convection	70	35	350
	microp_aero_wsub_scale	Subgrid velocity for liquid activation scaling	1	0.1	5

## 3. Results: asymmetrical response to warming/cooling



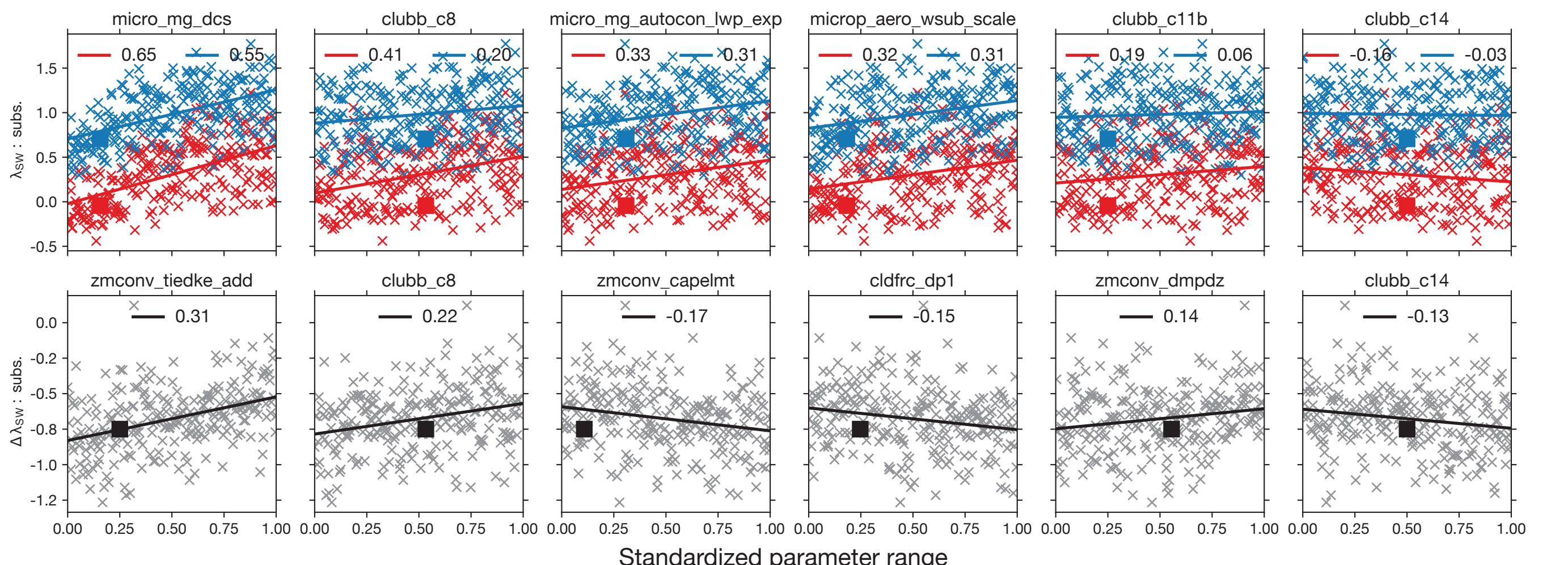
PPE generates a wide range of cloud states & feedbacks



### Compared to cooling, warming has

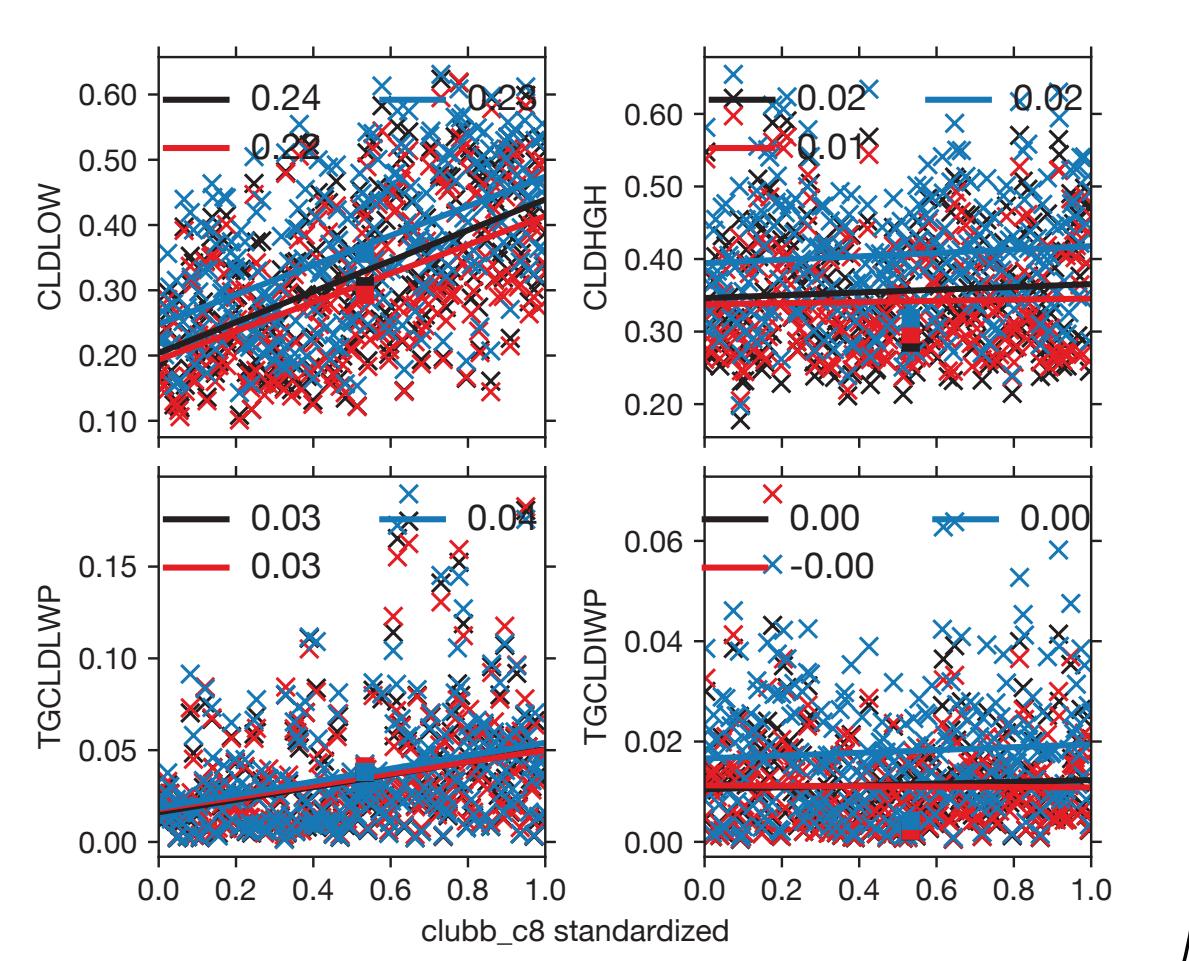
1. more negative SW  $\lambda_{\text{cl}}$  over low-lat. subsidence regime ( $\omega_{500} > 0$  &  $|\text{lat}| < 40^\circ$ );
2. more positive SW  $\lambda_{\text{cl}}$  over low-lat. ascending regime ( $\omega_{500} < 0$  &  $|\text{lat}| < 40^\circ$ ); and
3. more positive SW  $\lambda_{\text{cl}}$  over high-lat. regime ( $|\text{lat}| > 40^\circ$ ).

Cancellation between LW & SW components, and between cloud regimes

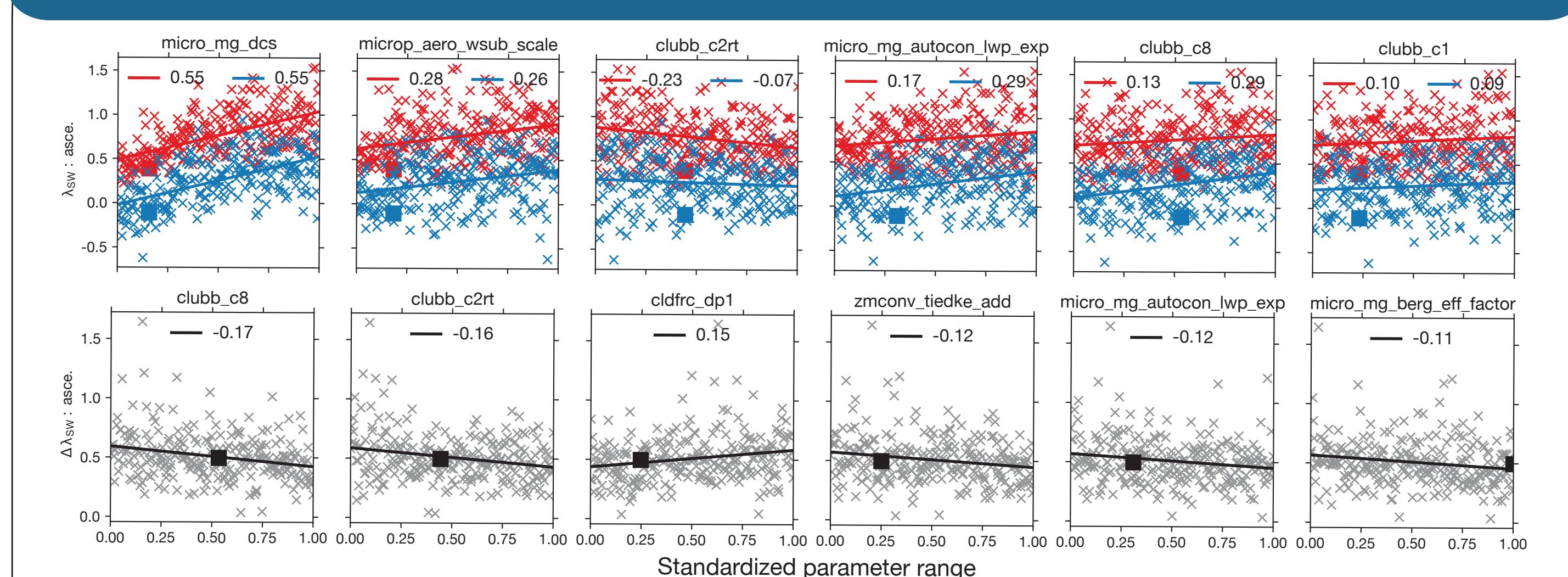


Over the low-lat. subsidence regime, moist turbulence & deep convection have the strongest impact on the SW  $\lambda_{\text{cl}}$  nonlinearity

1. Role of the background stratocumulus (sc) fraction: larger clubb\_c8 → stronger damping on  $w^3$  → clouds more sc-like → stronger  $\lambda_{\text{cl}}$
2. Role of the background low-tropospheric stability?

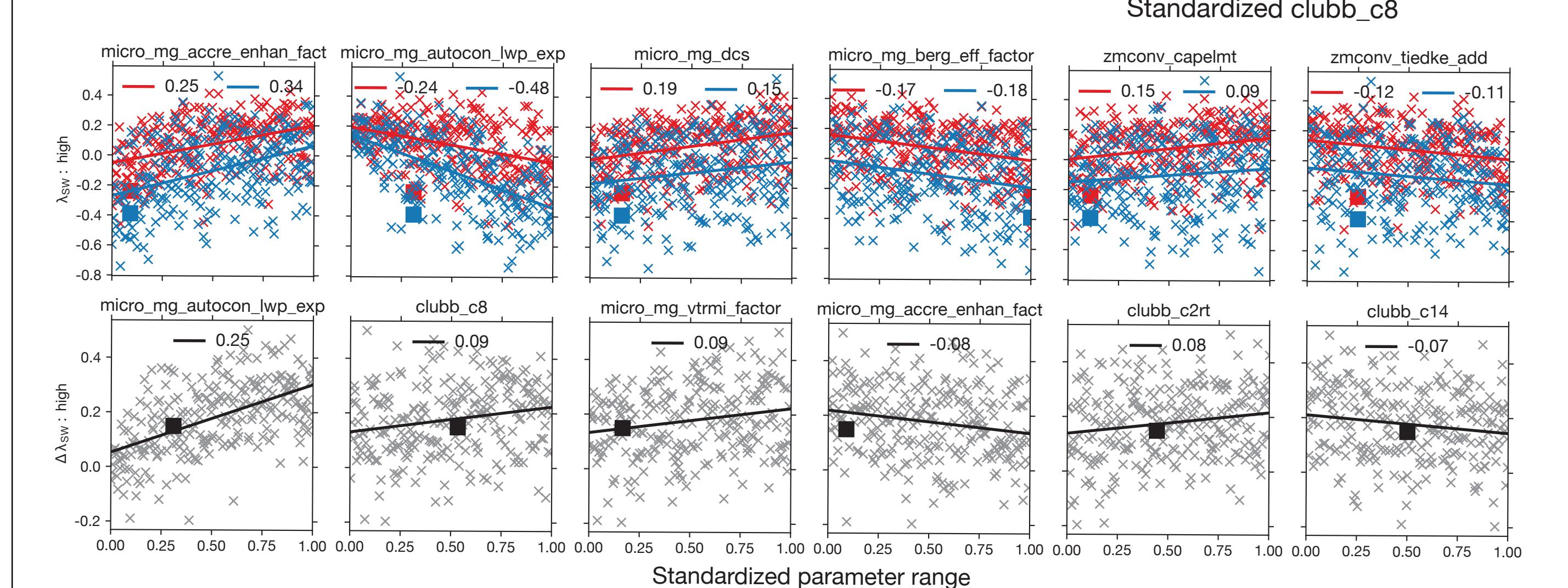


## Results: continued



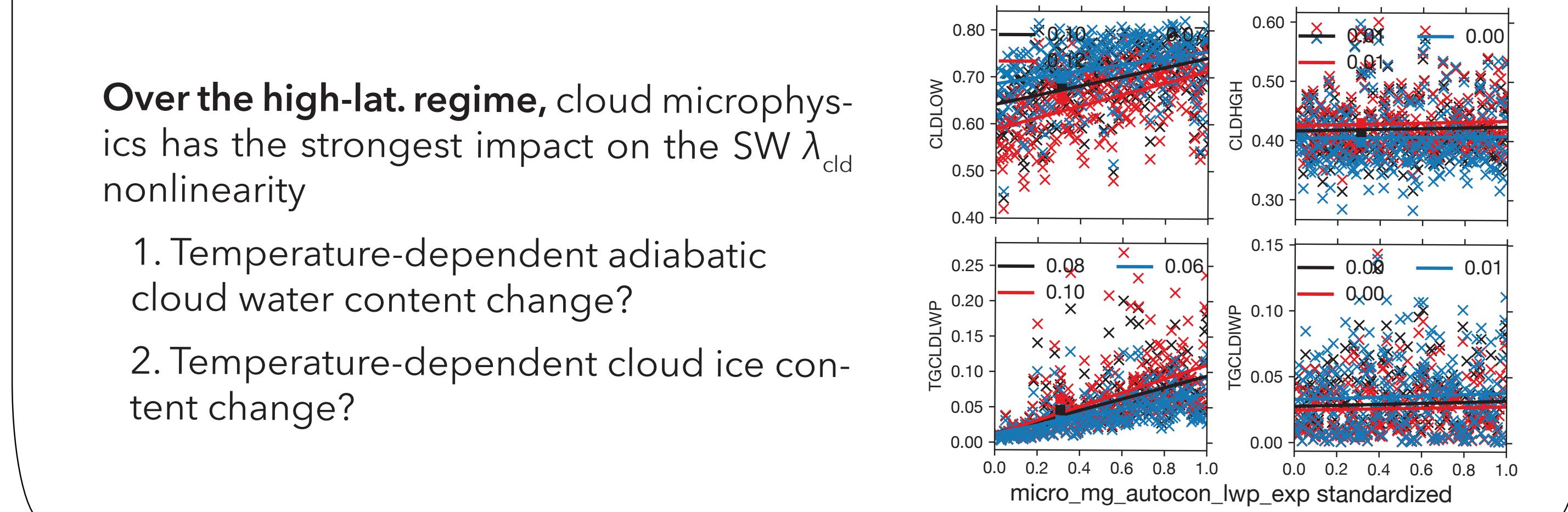
Over the low-lat. ascending regime, moist turbulence & deep convection have the strongest impact on the SW  $\lambda_{\text{cl}}$  nonlinearity

1. Temperature-dependent surface moisture flux change?



Over the high-lat. regime, cloud microphysics has the strongest impact on the SW  $\lambda_{\text{cl}}$  nonlinearity

1. Temperature-dependent adiabatic cloud water content change?
2. Temperature-dependent cloud ice content change?



## 4. Preliminary summary

- Within the model parameter space, the cloud feedback exhibits strong and robust state dependence, i.e., asymmetrical response to warming/cooling
- Mechanisms are cloud regime specific, and may include the dependence of stratocumulus fraction, low-tropospheric stability, moisture flux/content, cloud ice content on the background temperature.