Projections of the Transient State-Dependency of Climate Feedbacks

Robbin Bastiaansen (r.bastiaansen@uu.nl), Henk A. Dijkstra and Anna S. von der Heydt

Summary

Transient and state-dependent behaviour of climate feedbacks is important for accurate future projections. However, it is hard to track these via linear regressions of climate feedback contributions to changes in global temperature. Hence, here we introduce a new multivariate climate feedback framework that does take into account the transient state dependencies of climate feedback. Using the new framework, changes in feedback processes can be analyzed per time scale and temporal evolution can be tracked. Further, within the framework it is possible to create transient and equilibrium projections of (the spatial patterns of) climate feedbacks for all sort of emission scenarios. Illustrated here on CESM2.

A new climate feedback framework

In the linear regime, the response to forcing of any observable climate feedback contributions included - can be split into M climate modes as

$$R_j(t) = \sum_{m=1}^M \beta_m^{[R_j]} \mathcal{M}_m^g(t)$$

where $\beta_m^{[R_j]}$ contains all observable-dependent information and $\mathcal{M}_m^g(t)$ contains all time and forcing information.

Feedback strength can then be computed per mode as

$$\mu_j^m \coloneqq \frac{\beta_m^{[R]}}{\beta_m^{[T]}}$$

and, for any given forcing scenario, the instantaneous feedback strength can be computed as

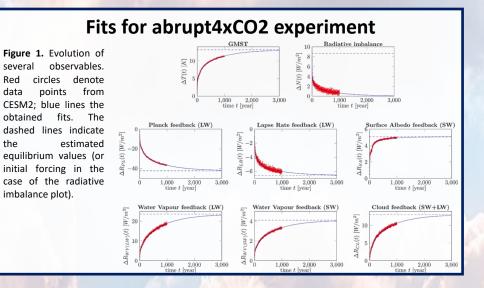
$$\lambda_j^{inst}(t) \coloneqq \frac{\frac{d}{dt}\Delta R_j}{\frac{d}{dt}\Delta T}$$

Procedure for CESM2

- 1. Compute feedback contributions $\Delta R_i(t)$ using radiative kernel approach
- 2. Use abrupt4xCO2 experiment to find parameters using nonlinear fitting
- 3. Compute feedback strengths & make projections

Remarks:

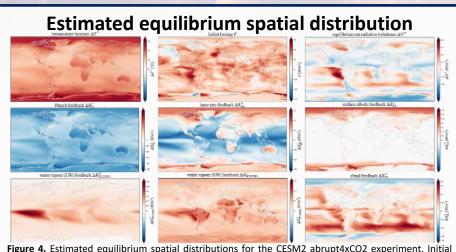
- Projections can also be made for other forcing scenarios. If applied on an ensemble of runs Linear Response Theory suggest correctness of these projections.
- Method can be easily extended to patterning feedback spatial of contributions



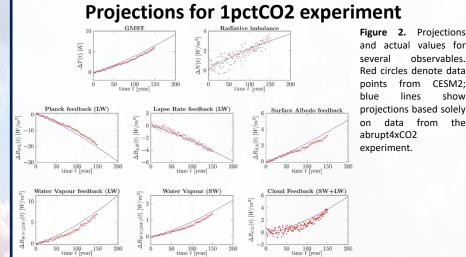
Values for climate feedback strength per mode

	Mode 1	Mode 2	Mode 3	Equilibrium
τ_m	4.5 (± 0.1)	$127 (\pm 3.8)$	$889 (\pm 50)$	—
λ_m	$-1.28~(\pm 0.08)$	$-0.38~(\pm 0.03)$	$-0.37~(\pm 0.02)$	$-0.66 (\pm 0.03)$
Planck (LW) Lapse Rate (LW)	$-3.16 (\pm 0.02)$ $-0.73 (\pm 0.03)$	$-3.24 (\pm 0.02)$ $-0.50 (\pm 0.03)$	$-3.23 (\pm 0.01)$ $-0.32 (\pm 0.03)$	$-3.21 (\pm 0.05)$ $-0.50 (\pm 0.01)$
Surface Albedo (SW) Water Vapour (LW)	$+0.62 (\pm 0.04)$ +0.97 (± 0.03)	$+0.56 (\pm 0.02)$ +1.38 (± 0.02)	$+0.08 (\pm 0.10)$ +2.71 (± 0.01)	$+0.39 (\pm 0.01)$ +1.79 (± 0.04)
Water Vapour (SW) Clouds (SW + LW)	$+0.21 (\pm 0.09) +0.27 (\pm 0.36)$	$+0.26 (\pm 0.05)$ +1.19 (± 0.02)	$+0.43 (\pm 0.02)$ +1.43 (± 0.01)	$+0.31 (\pm 0.01)$ +1.00 (± 0.03)
sum	-1.82 (± 0.37)	-0.36 (± 0.07)	+1.09 (± 0.11)	-0.22 (± 0.08)
residue	+0.54 (± 0.38)	$-0.02~(\pm 0.08)$	$-1.46~(\pm 0.11)$	$-0.43 (\pm 0.08)$

Table 1. τ_m denotes the time scale in years; the other values have units $Wm^{-2}K^{-1}$. Plusminus values indicate



forcing is derived directly from data; the other estimates are derived via the fitted spatial modes



Evolution of feedback strength over time



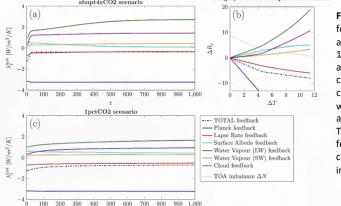


Figure 3. Instantaneous feedback strengths for (a) abrupt4xCO2 and (c) 1pctCO2 experiments, and (b) Gregory plot for climate feedback contributions against warming in the abrupt4xCO2 experiment. The 1pctCO2 panel shows feedback strengths for continued 1% yearly CO2 increase.

CESM2.

show

Discussion

- Long-term evolution and state-depency of climate feedbacks can be captured
- Multivariate climate projections can be made, even for other forcing scenarios
- Results on CESM2 align with previous research for the faster modes
- Results for mode 3 consider longer time scales than normally and show a larger increase in mainly water vapour and cloud feedbacks
- It seems a feedback might be overlooked that is relevant on long time scales

