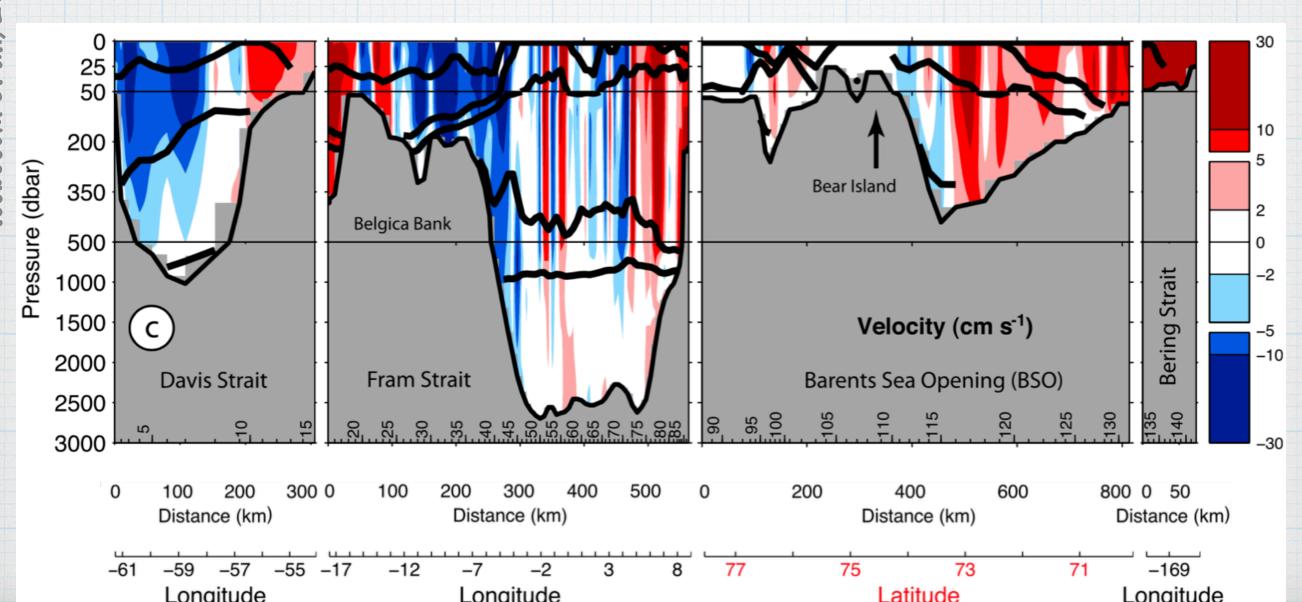
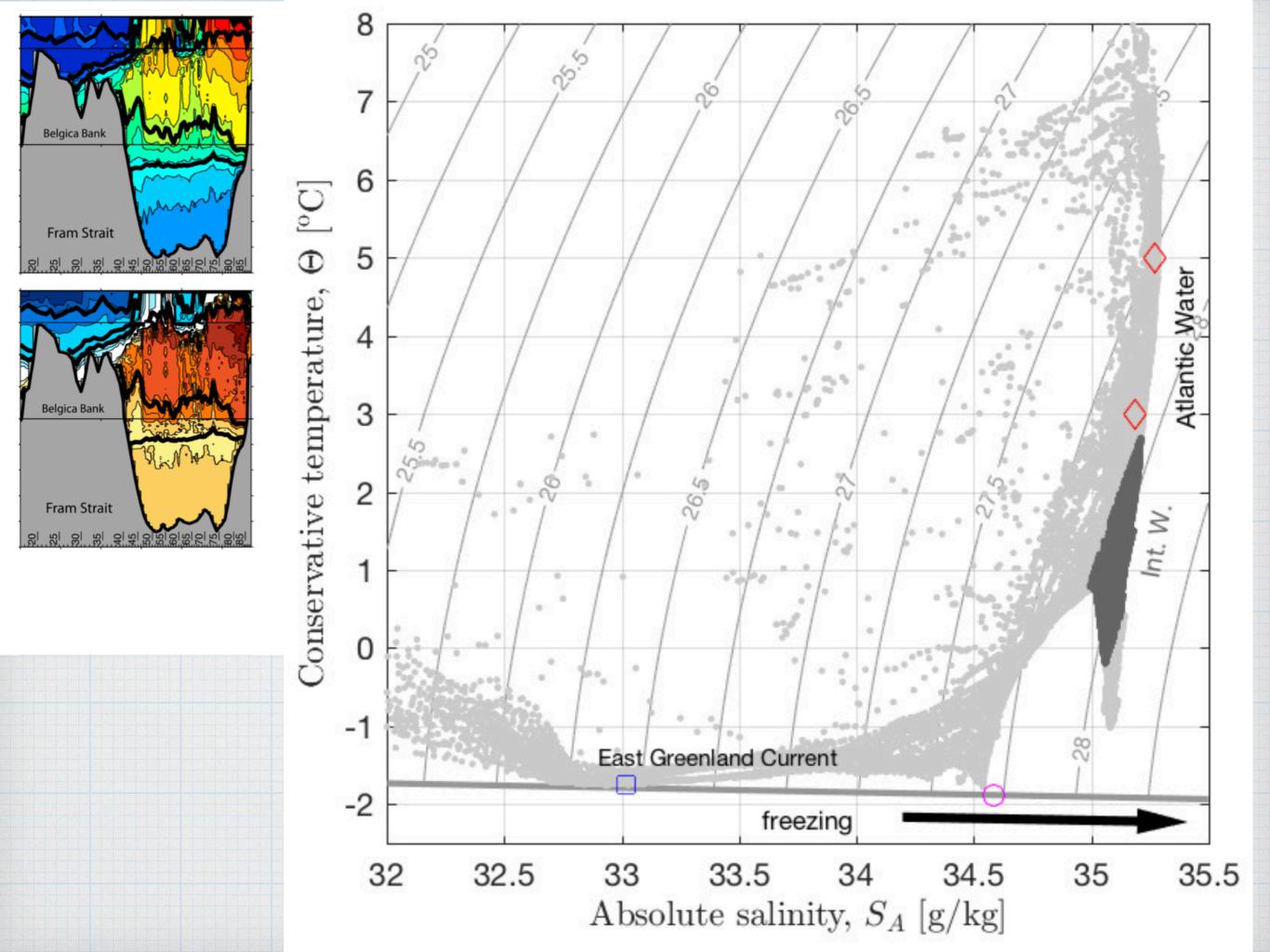


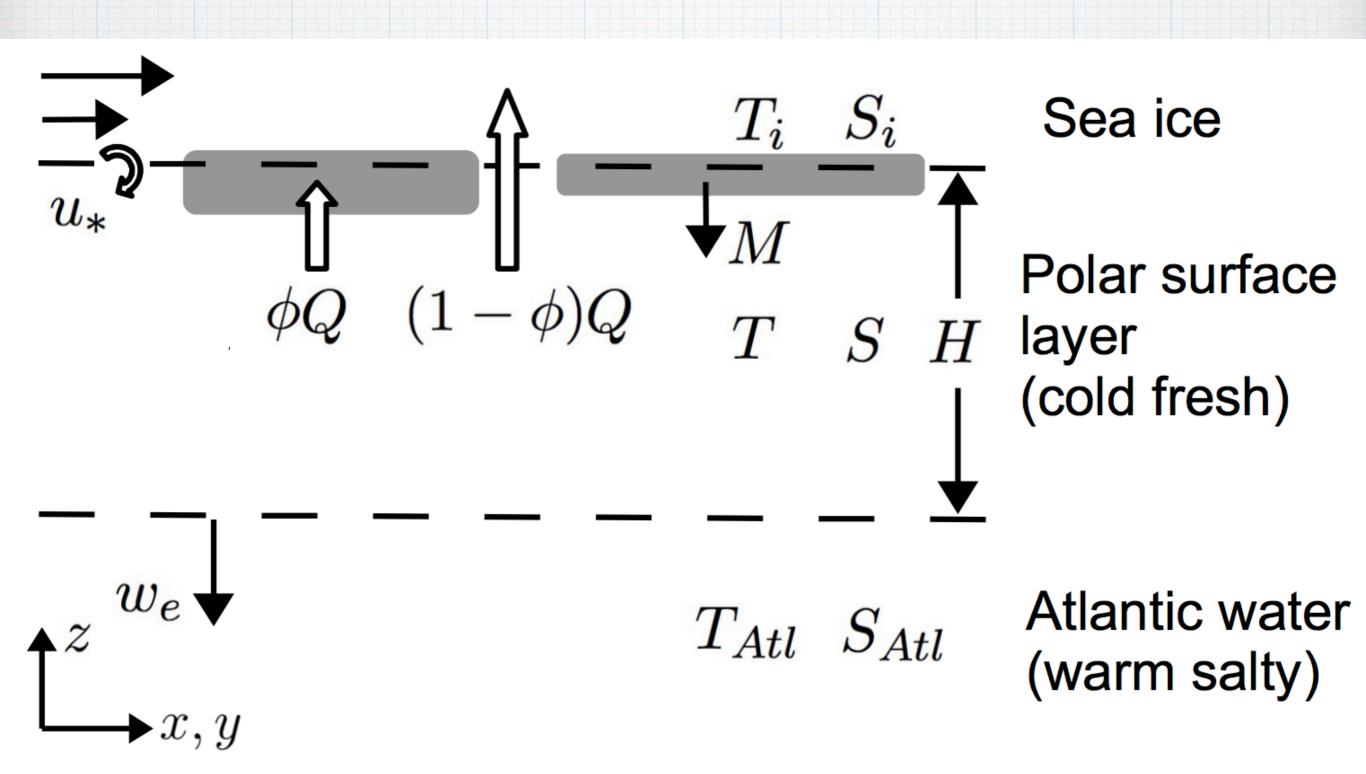
Conceptual Constraints on Polar MOCs

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Simple conceptual upper ocean model of Atlantic Water melting sea ice, e.g., in the Eurasian Basin



Fluxes of heat and freshwater

$$T_{i} S_{i}$$
 Sea ice
$$\phi Q (1-\phi)Q T S H$$
 Polar surface layer (cold fresh)
$$T_{Atl} S_{Atl} S_{Atl}$$
 Atlantic water (warm salty)

$$Q = Q_a + \rho_i M_i [L + c(T - T_f) + c_i (T_f - T_i)]$$

$$\mathcal{F} = -\frac{\rho_i}{\rho} M_i (S - S_i)$$

$$M = (\rho_i/\rho)M_i$$

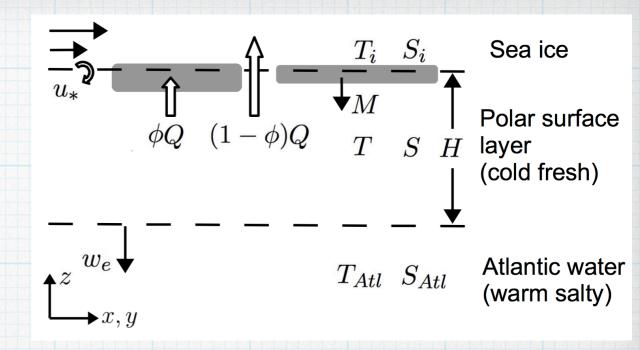
$$M = \phi \frac{Q}{\rho \left[L + c(T - T_f) + c_i(T_f - T_i) \right]}$$

Evolution equations:

$$\frac{dH}{dt} = w_e + M,$$

$$H\frac{dT}{dt} = w_e \Delta T - \frac{Q}{\rho c},$$

$$H\frac{dS}{dt} = w_e \Delta S - M(S - S_i)$$

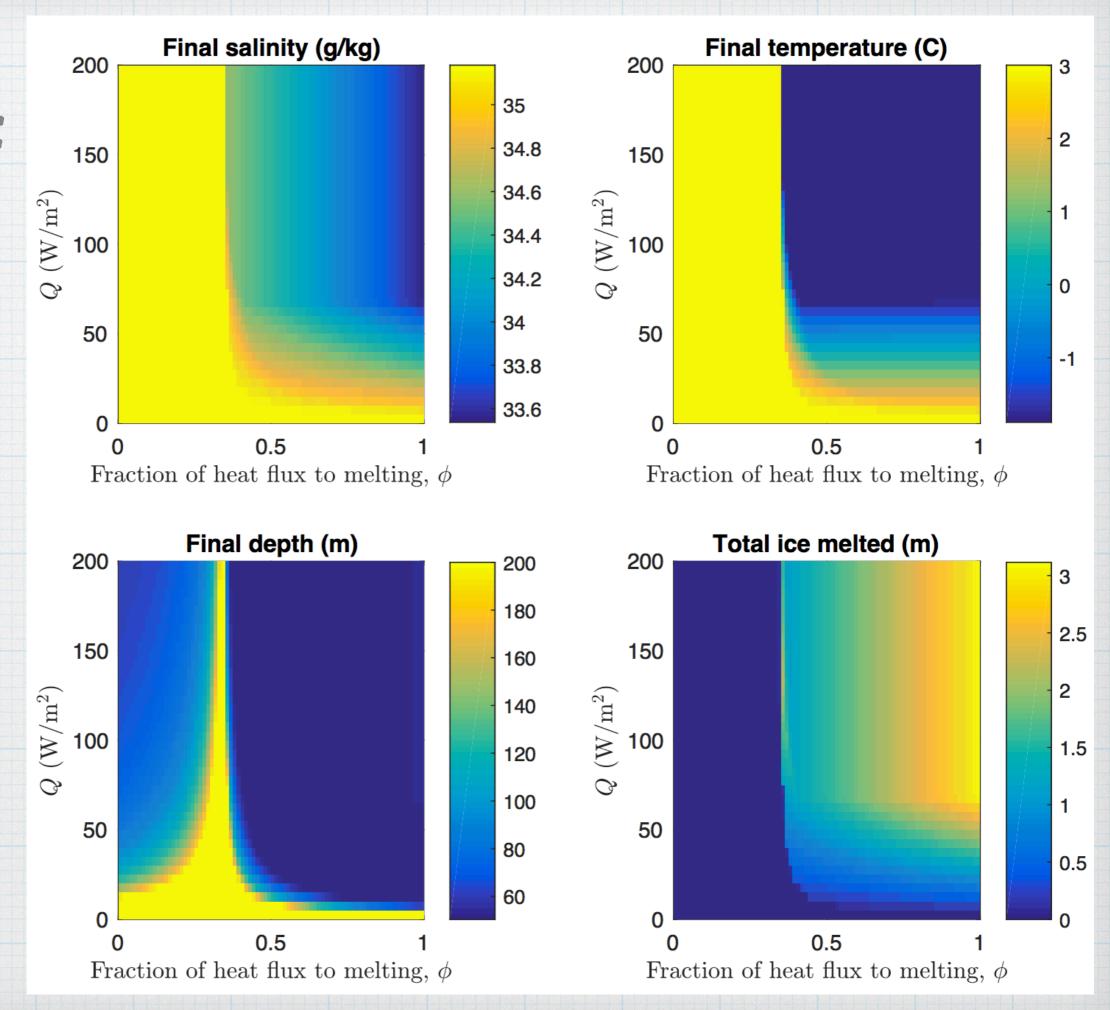


$$\Delta \rho = \rho(T_{Atl}, S_{Atl}) - \rho(T, S)$$
 $w_e = \frac{\rho}{g\Delta\rho} \left(\frac{2m_0u_*^3}{H} - \epsilon B\right)$
 $\epsilon = \begin{cases} 1 \text{ for } B \ge 0, \\ 0.05 \text{ otherwise.} \end{cases}$

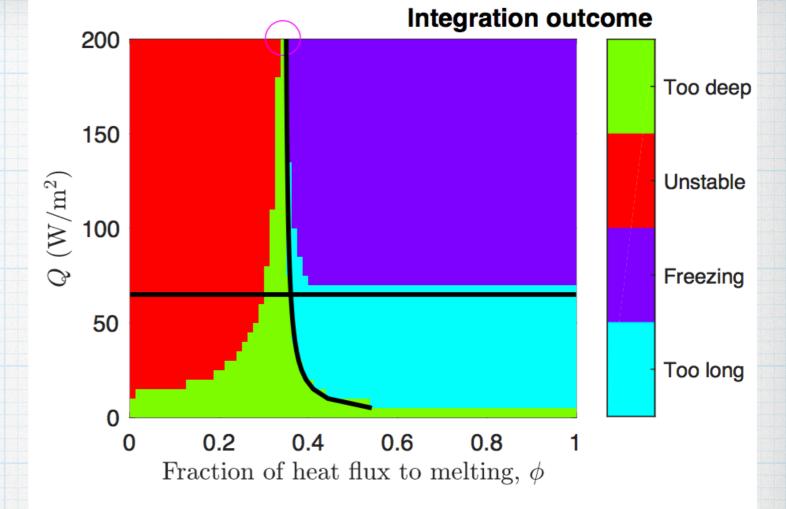
 $B = g \left| \frac{\alpha Q}{\rho c} + \beta M \left(S - S_i \right) \right|$

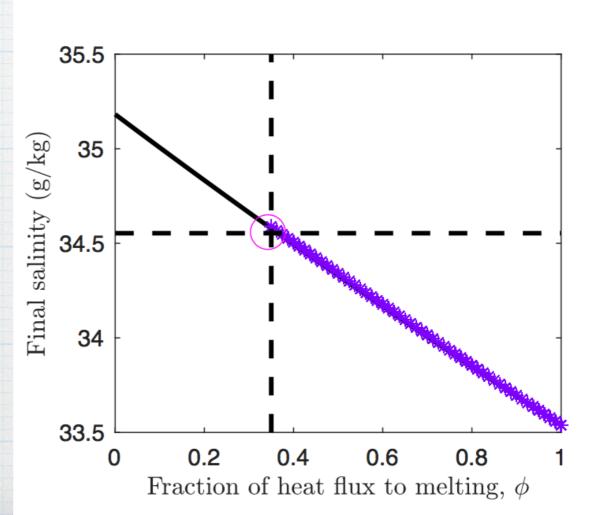
Solutions:

CFor many (phi, Q) pairs]

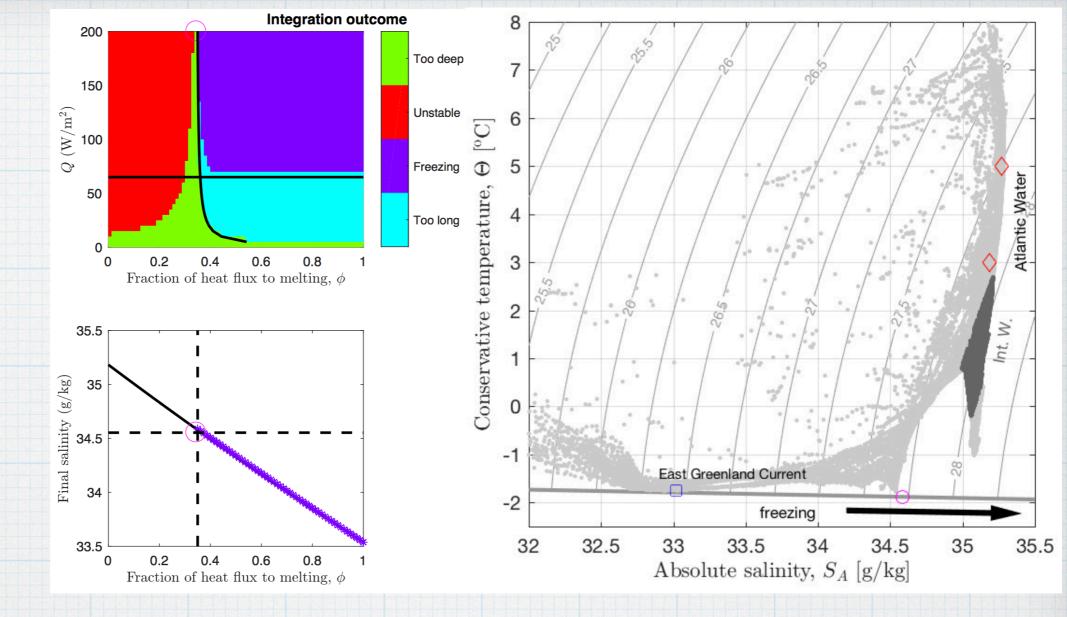


Solution categories:





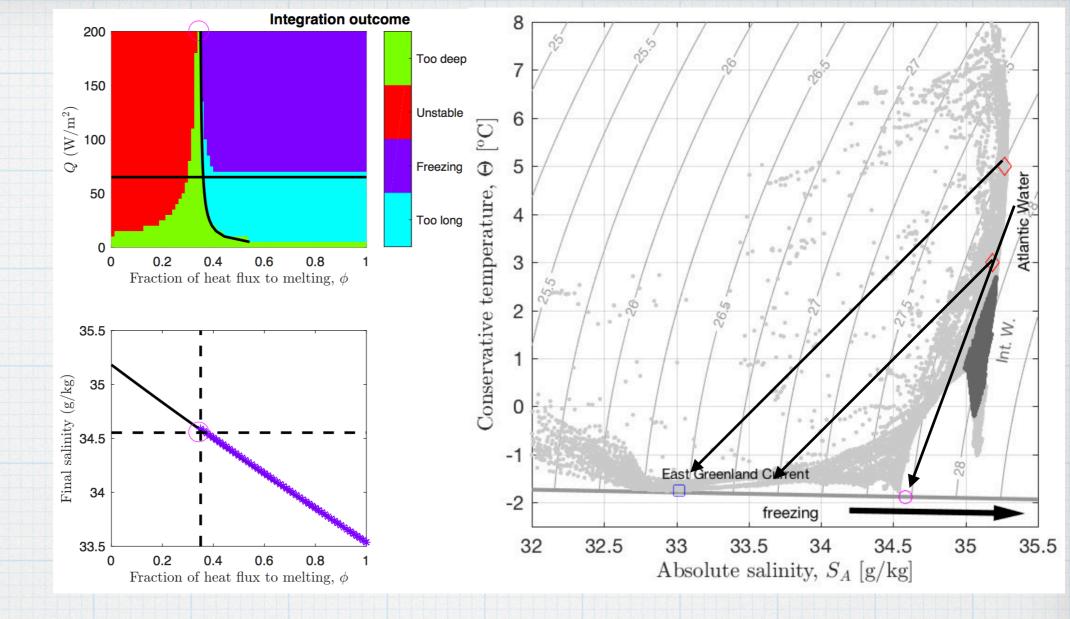
Properties of "EGC" water



$$S_{1} = \frac{\left[L + c_{i} \left(T_{f} - T_{i}\right)\right] S_{Atl} + \phi c \left(T_{Atl} - T_{f}\right) S_{i}}{\left[L + c_{i} \left(T_{f} - T_{i}\right)\right] + \phi c \left(T_{Atl} - T_{f}\right)} \approx \frac{L S_{Atl}}{L + \phi c \left(T_{Atl} - T_{f}\right)}$$

$$\phi_{min} = \frac{-\alpha \left[L + c(T - T_f) + c_i(T_f - T_i) \right]}{\beta c \left(S - S_i \right)} \approx \frac{-\alpha L}{\beta c S}$$

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Remarks

- * Conceptual model for properties of EGC (fresh polar outflow) water that fits the data well.
- * Comparison with Bert Rudels' formula. Similar fit to data, but different rationale.
- * Importance of partial melting of remotely-formed sea ice.
- * What sets the properties of the intermediate water?