



Spacebased Observations of Atlantic Ocean Meridional Transport Preliminary Results

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- Heat Transport
- Water Transport

WORLD CLIMATE PROGRAMME

RESEARCH • APPLICATION • IMPACT • DATA

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WORLD CLIMATE RESEARCH PROGRAMME

REPORT OF THE JSC/CCCO

'CAGE' EXPERIMENT: A FEASIBILITY STUDY

WCP - 22

MAY 1982

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Meridional Heat Transport (MHT)

Conservation of heat

$$\frac{\partial H}{\partial t} + \nabla \cdot \zeta = SW - LW - LH - SH$$

By Green's theorem

$$MHT(\theta) = \int_{\theta}^{\theta_0} \int_{x_1}^{x_2} \left(\frac{\partial H}{\partial t} - SW + LW + LH + SH \right) dx dy$$

H: Heat content

ζ : Horizontal heat flux

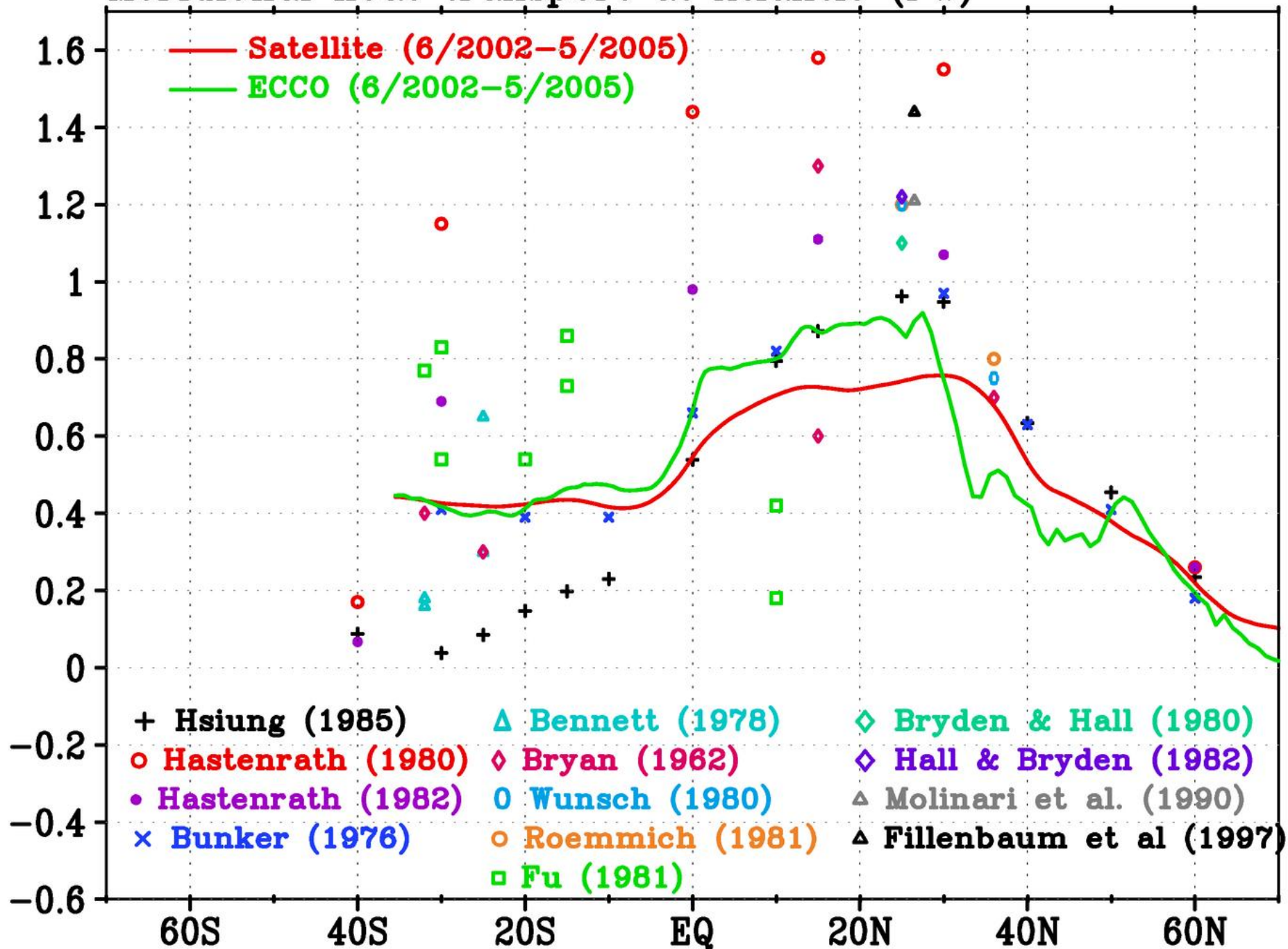
SW: Short wave radiation

LW: Long wave radiation

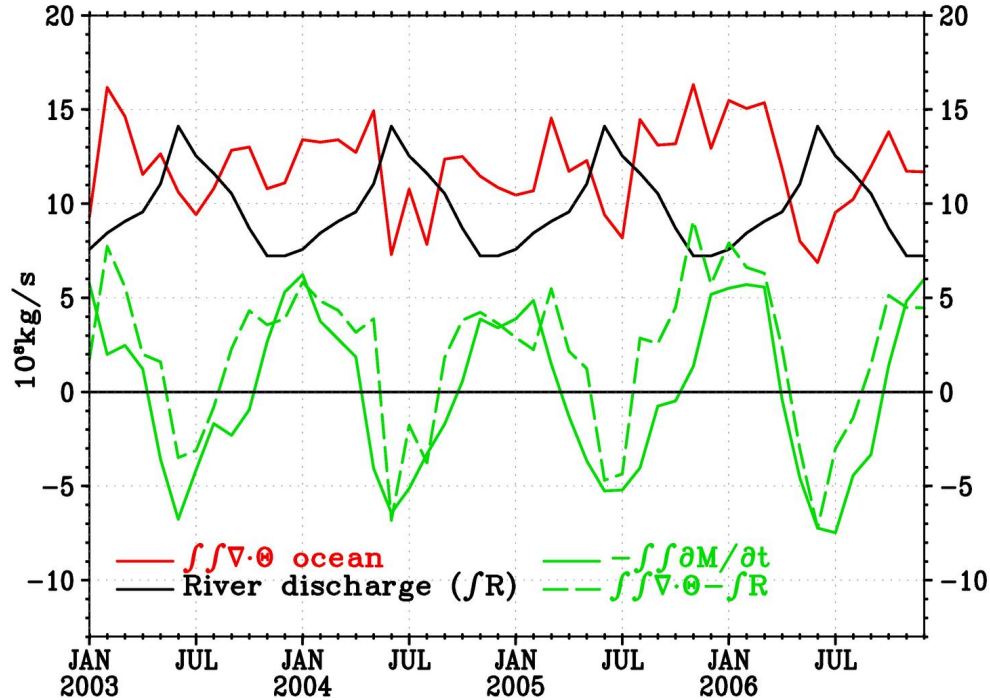
LH: Latent heat

SH: Sensible heat

Meridional heat transport at Atlantic (PW)



- **H can be derived from Argo**
- **Coverage is not sufficient**
- **Space-observation may be a solution**

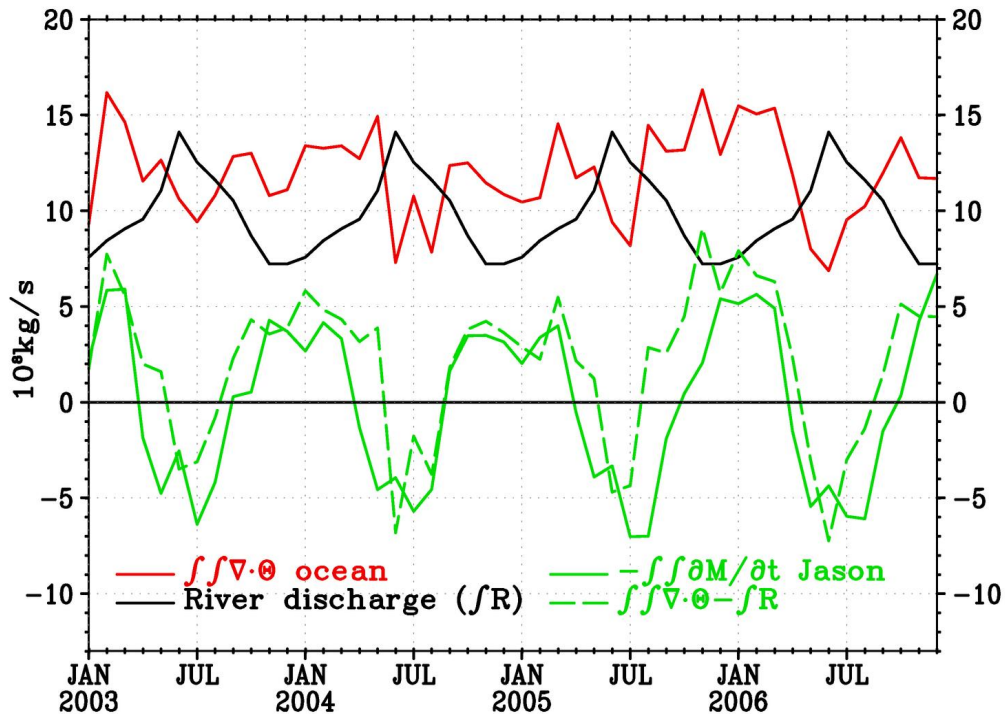


Red - divergence of water vapor transport integrated over depth of the atmosphere

Black - sum of climatological river discharge across all coastline

Green - loss rate of water stored in all oceans from GRACE

---- difference between fresh water flux and river input



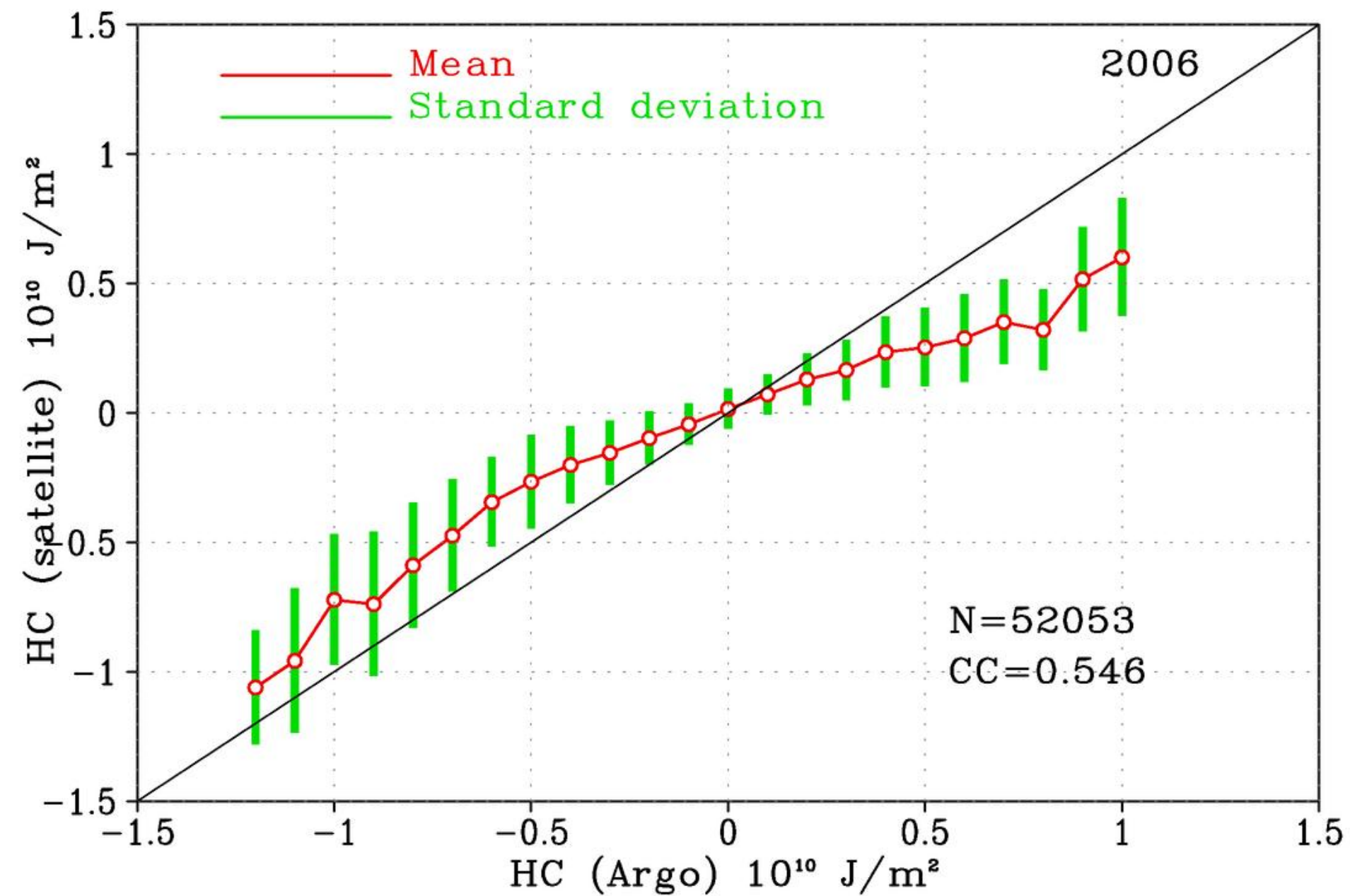
Green - subtracting climatological steric change from altimeter

$$\eta = \frac{P_a - P_b}{g \rho_0} + \frac{\langle \rho \rangle}{\rho_0}$$

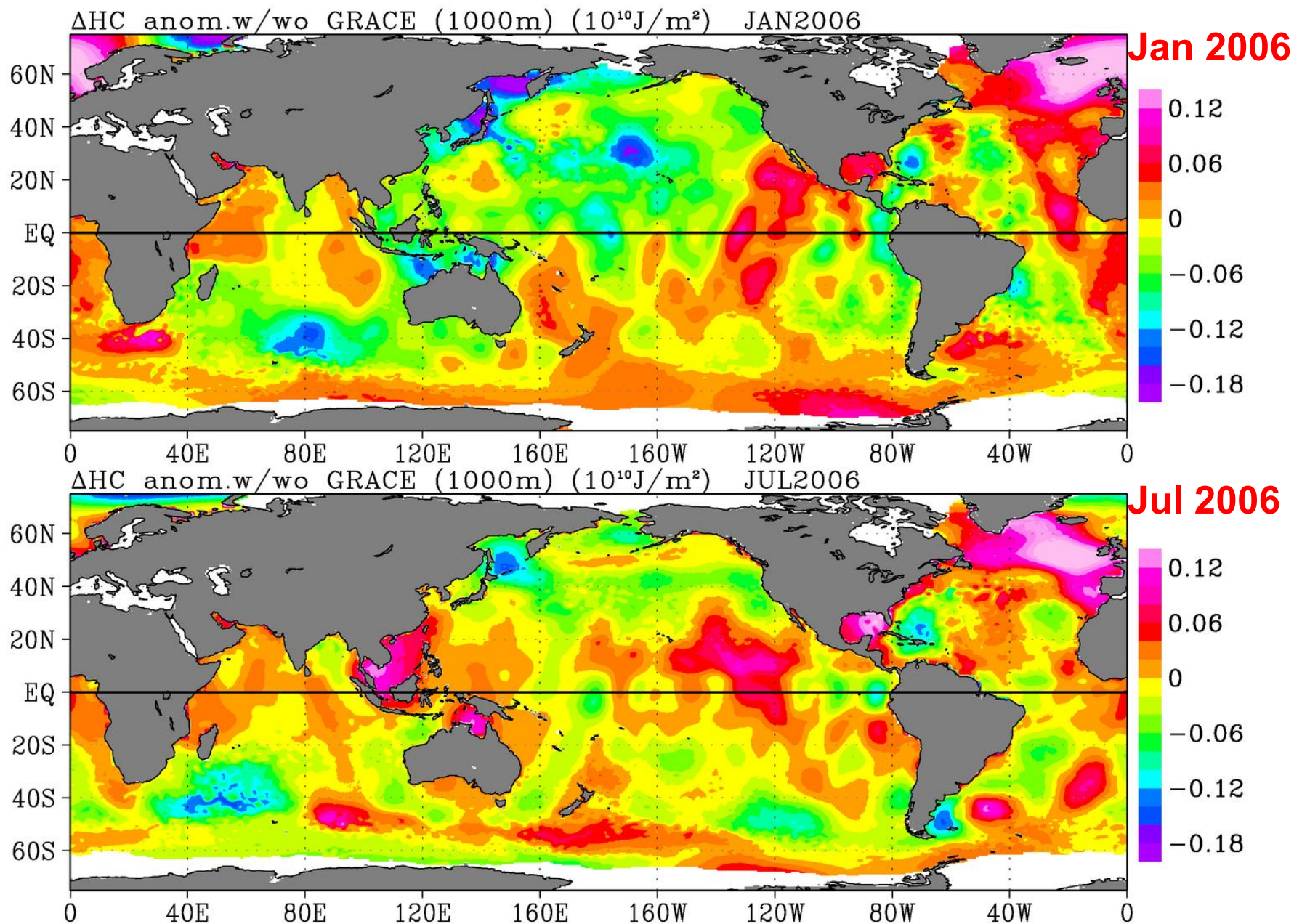
Sea level Mass change Steric change

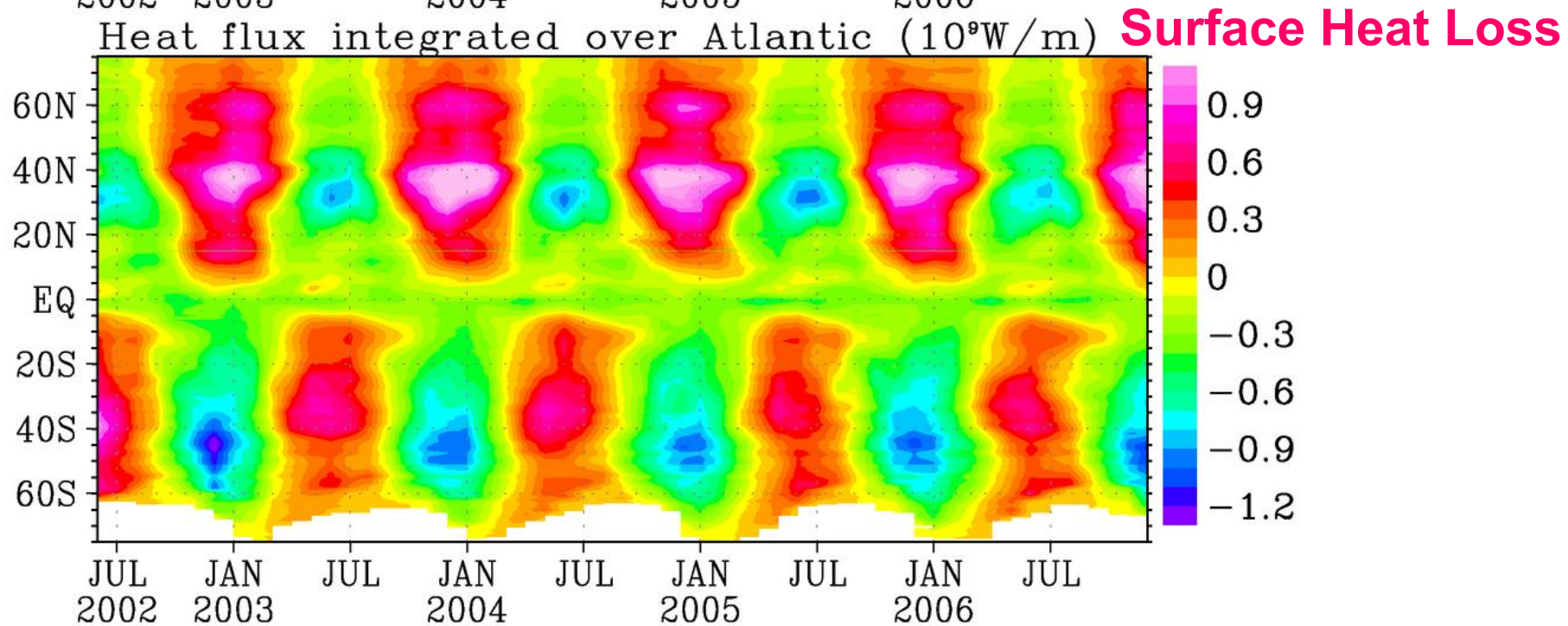
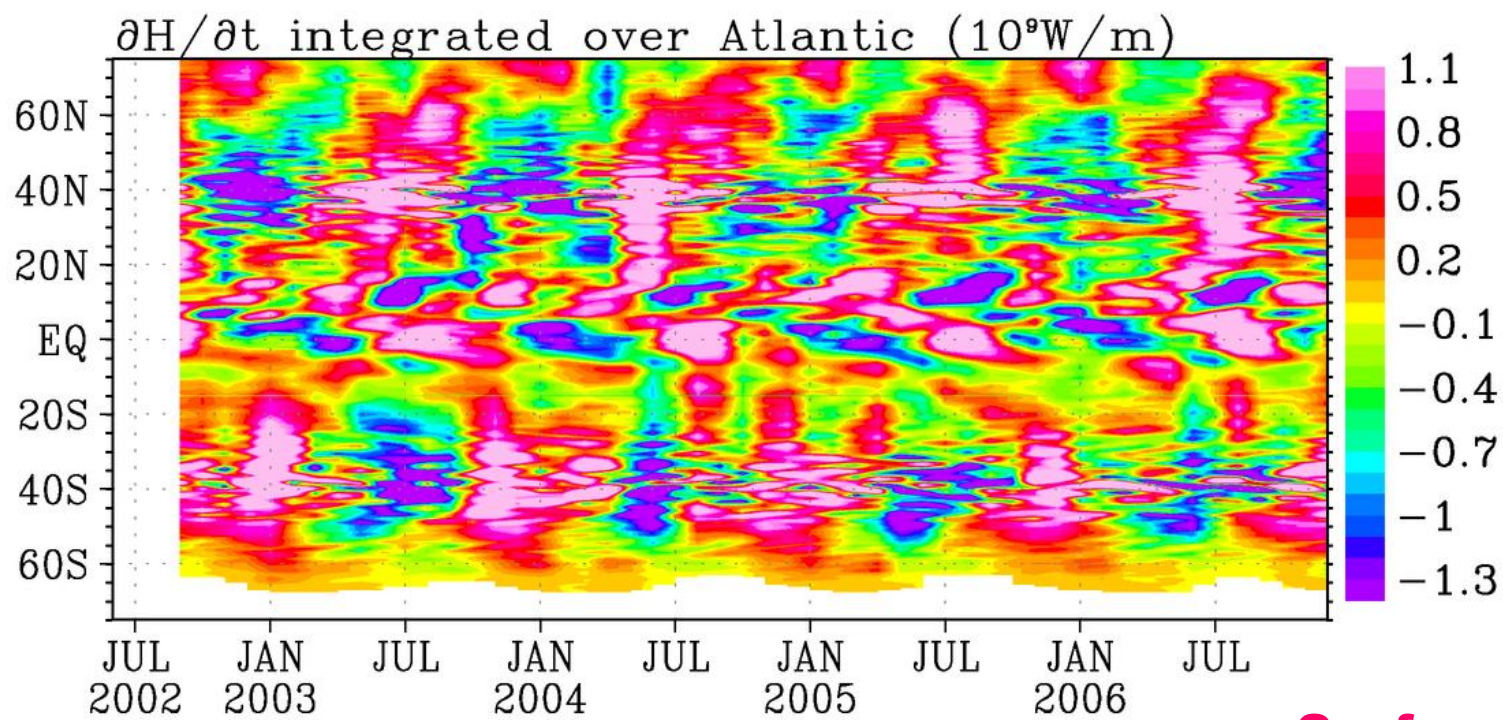
$$H' = \frac{c_p \rho_0}{\alpha} \eta'_s$$

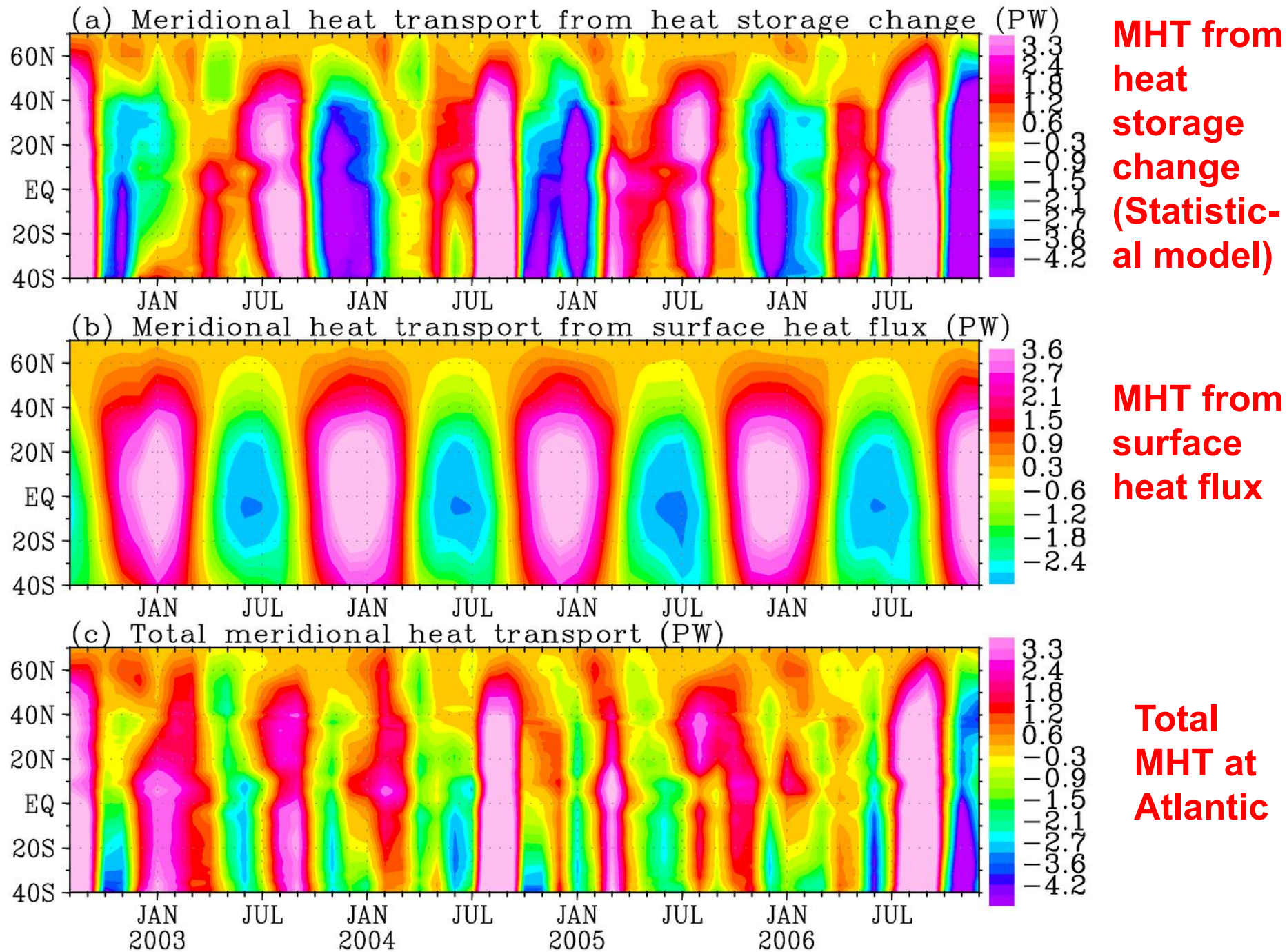
α - Thermal expansion coefficient



Impact of GRACE in deriving heat content







Meridional Water Transport (MWT)

Conservation of water mass

$$\frac{\partial M}{\partial t} + \nabla \cdot \psi = P - E$$

By Green's theorem

$$MWT(\theta) = \int_{\theta}^{\theta_0} \int_{x_1}^{x_2} \left(\frac{\partial M}{\partial t} + E - P - R \right) dx dy$$

Ekman water transport

$$EWT(\theta) = \int_{x_1}^{x_2} -\frac{\tau_x}{\rho f} dx$$

P: Precipitation

E: Evaporation

Ψ : Horizontal mass flux

R: River discharge

τ_x : Zonal stress

HYDROLOGIC BALANCE

$$\frac{\partial W}{\partial t} + \nabla \cdot \Theta = E - P$$

$$\Theta = \frac{1}{g} \int_0^{p_0} q U dp$$

$$W = \frac{1}{g} \int_0^{p_0} q dp$$

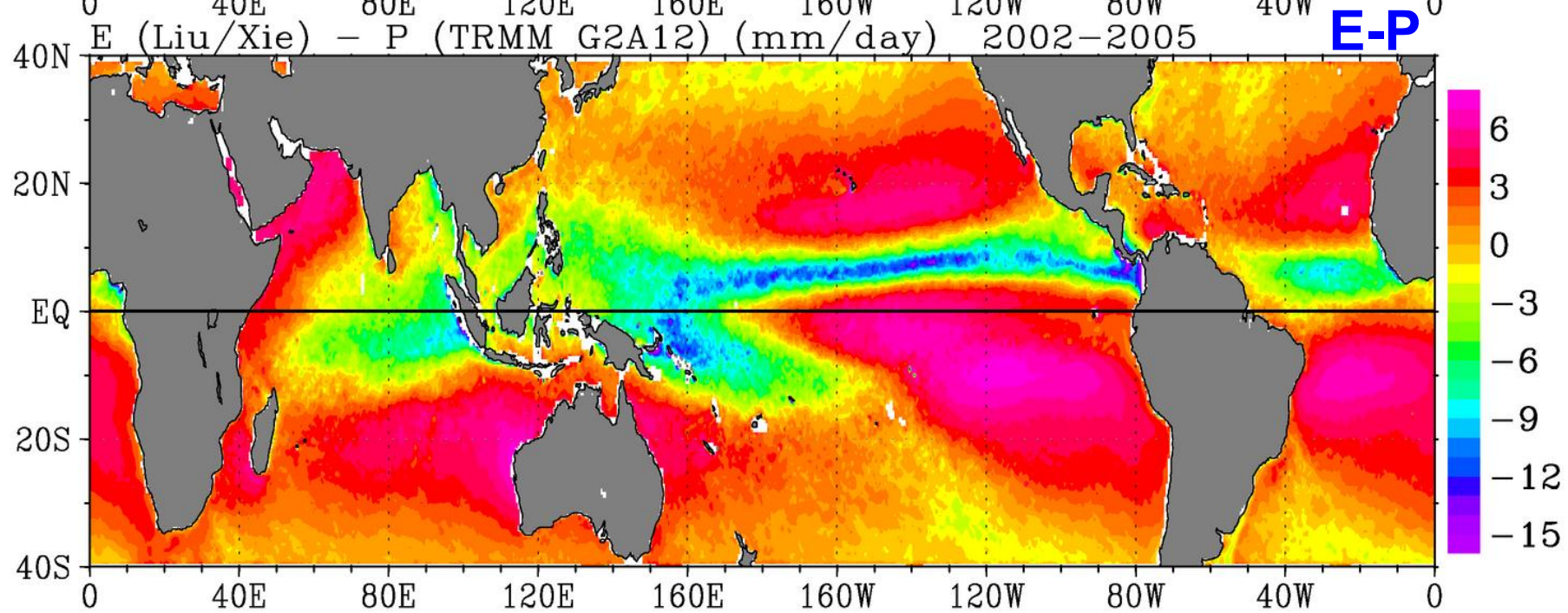
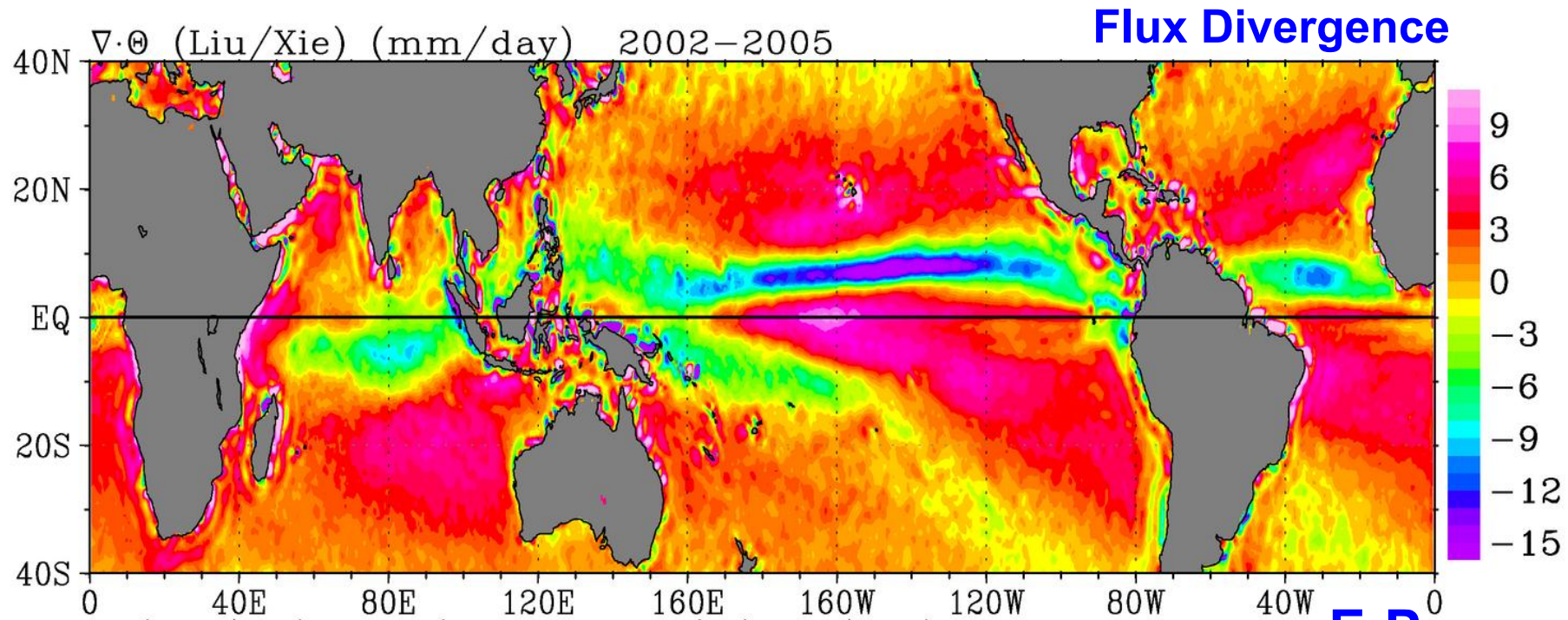
$$\Theta = U_e W$$

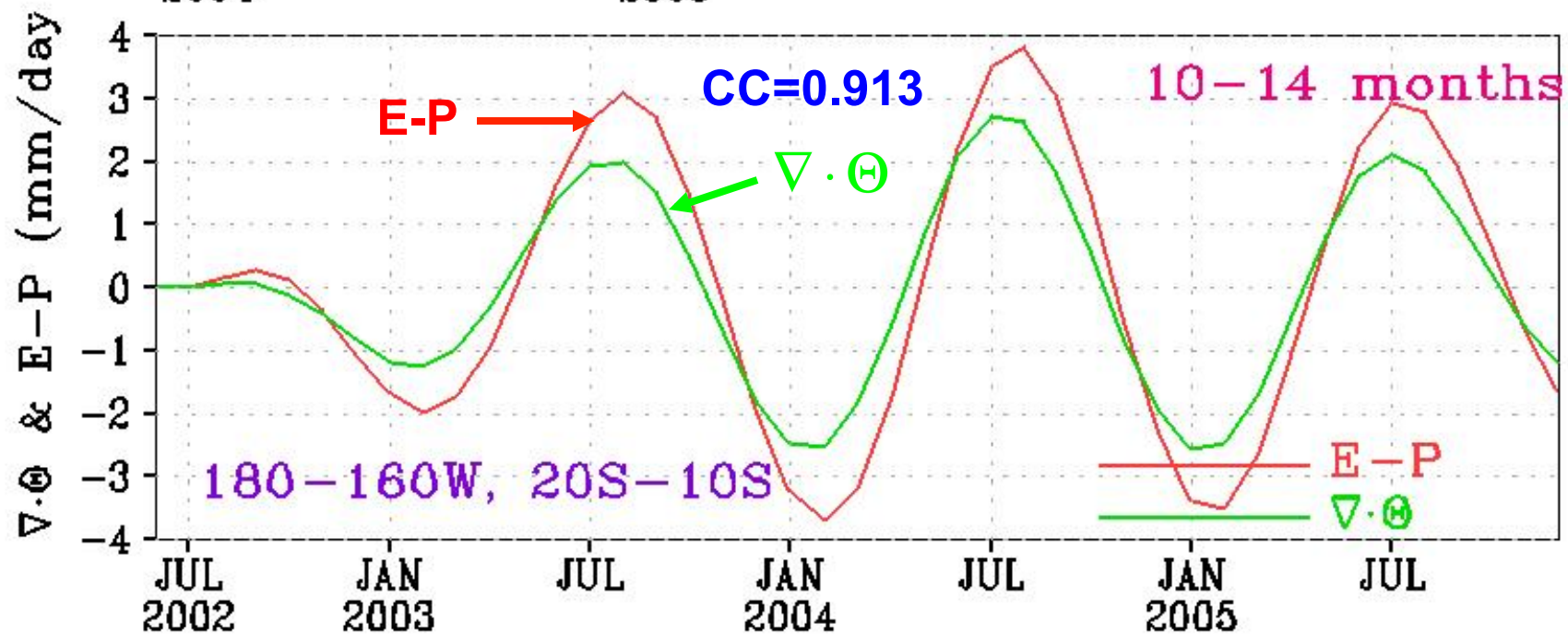
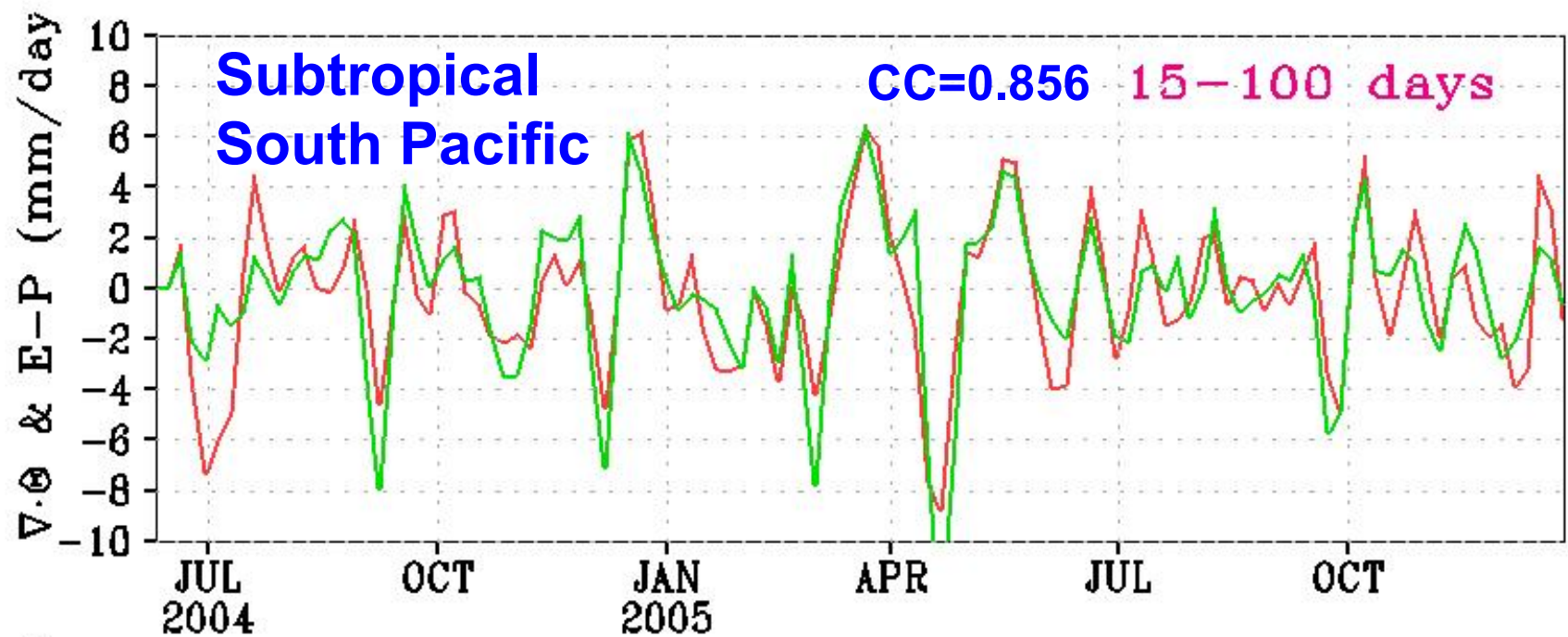
$U_e = f(U_s)$ Liu (1993)-polynomial

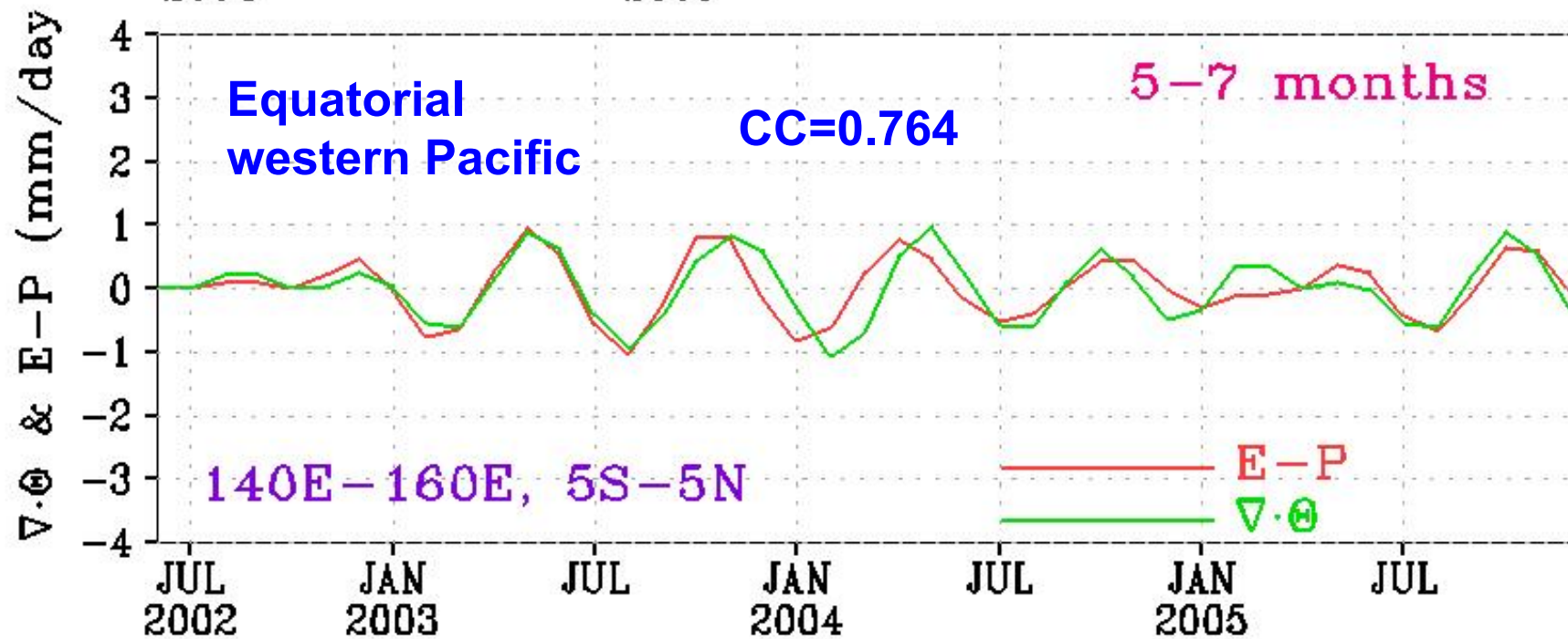
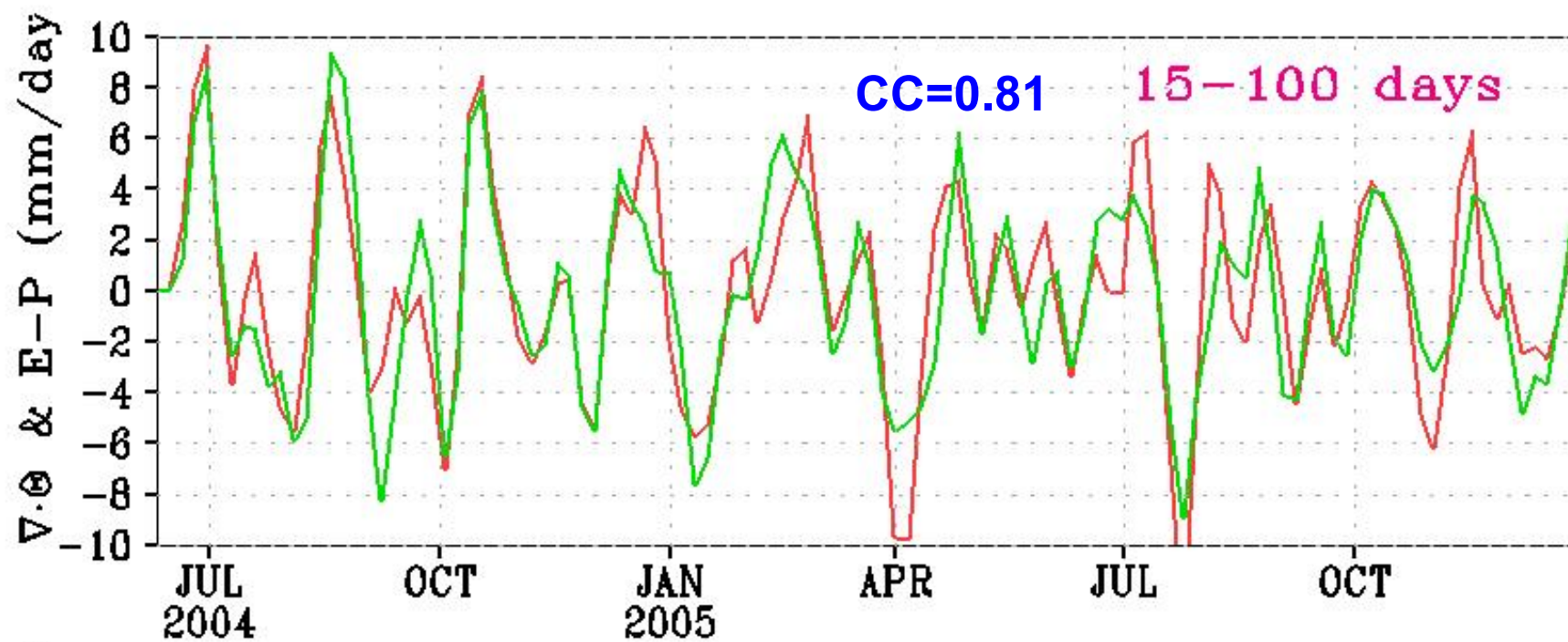
Liu & Tang (2005) - Neural Network

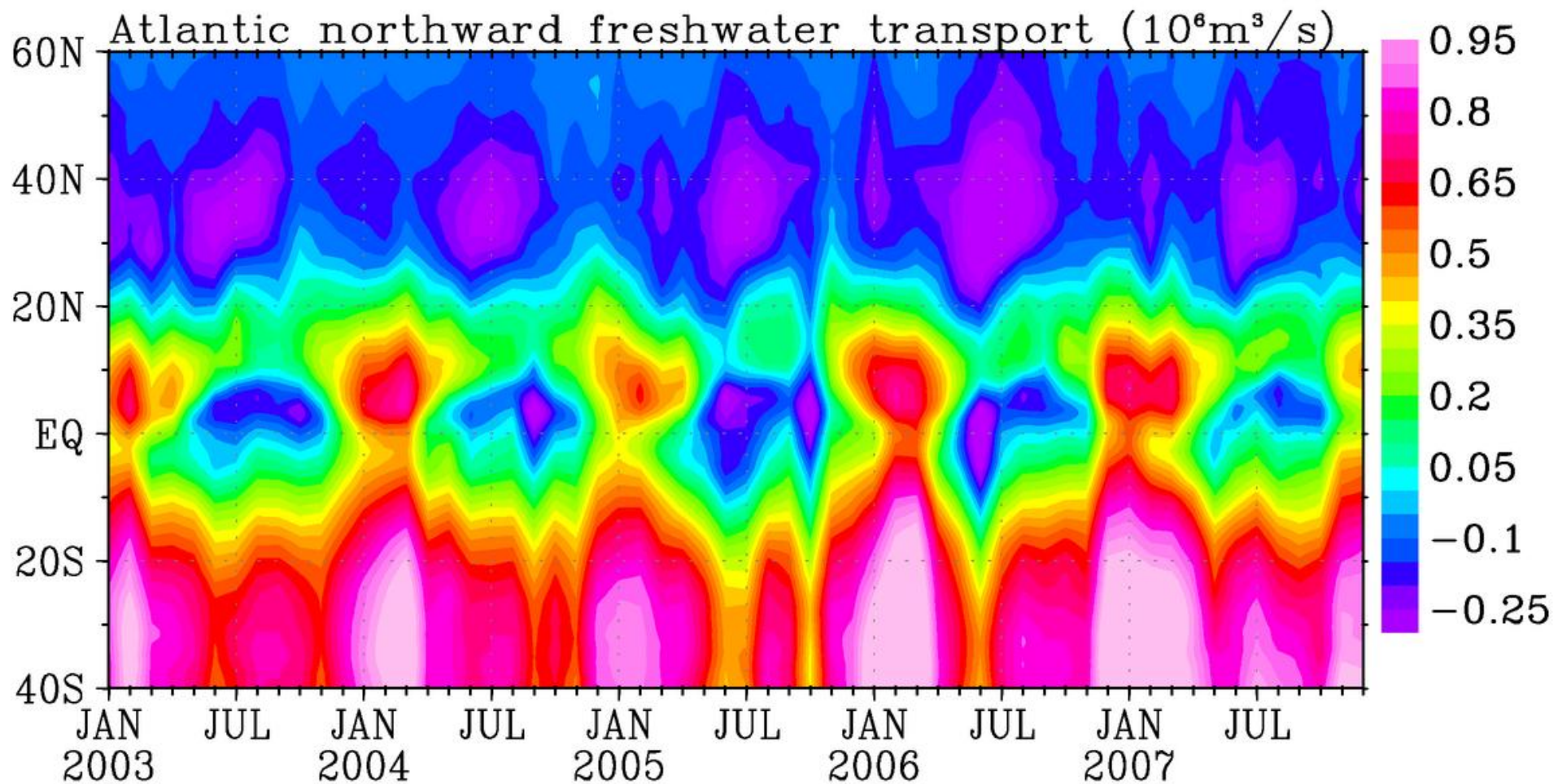
$U_e = U_{850mb}$ Heta & Mitsuta (1993)

Both U_s & U_{850mb} Xie et al. (2007) - SVR

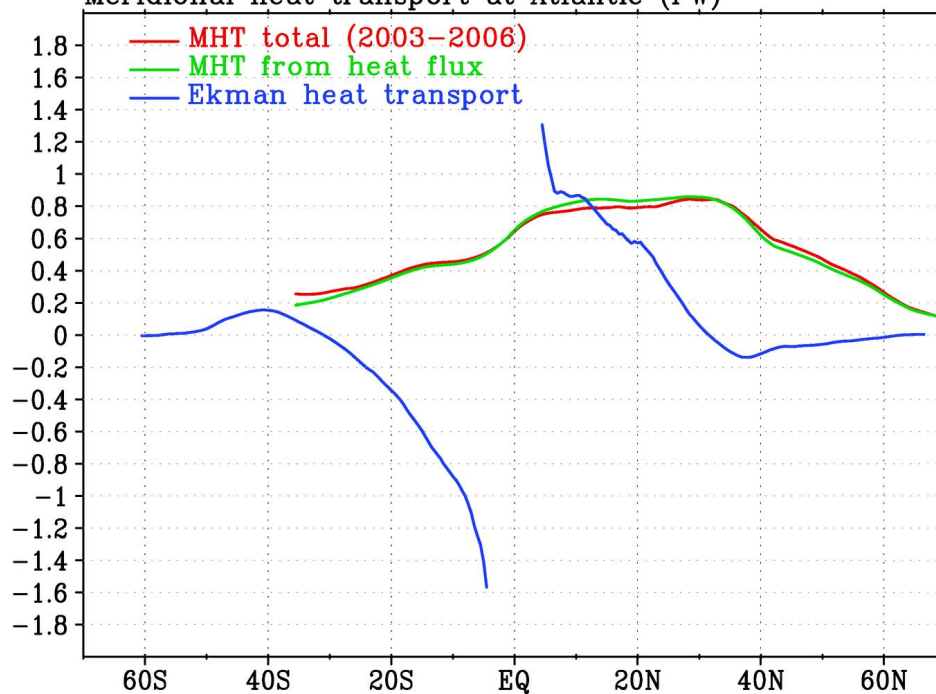








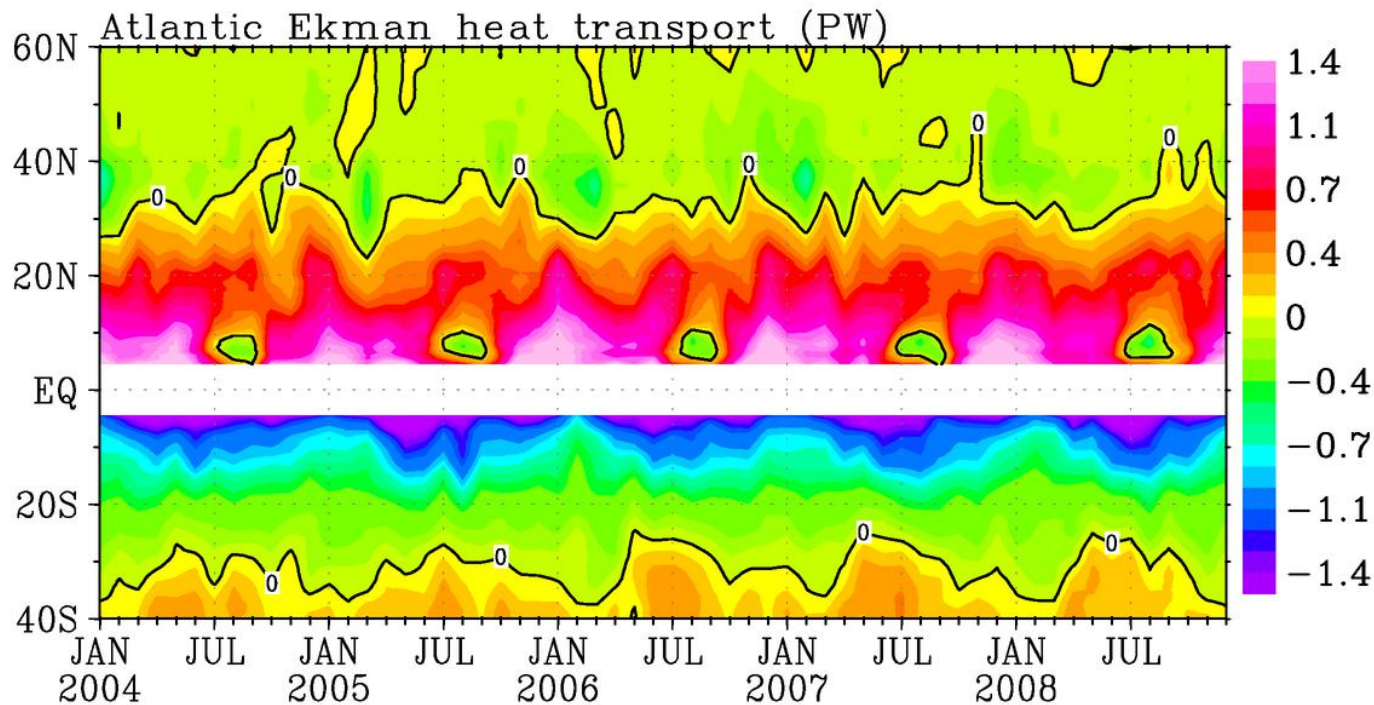
Meridional heat transport at Atlantic (PW)



$$\text{EHT}(\theta) = -c_p \int_{x_1}^{x_2} \frac{\tau_x}{f} (T_e - \bar{T}) dx$$

T_e : Potential temp. of Ekman layer

\bar{T} : Mean potential temp. of the water column



Summary

- ❑ Spacebased data provide almost continuous spatial and temporal coverages for a decade
- ❑ Reality checks are needed
- ❑ What is the relation between surface Ekman transport in the total meridional transport

backup

Ocean Balance

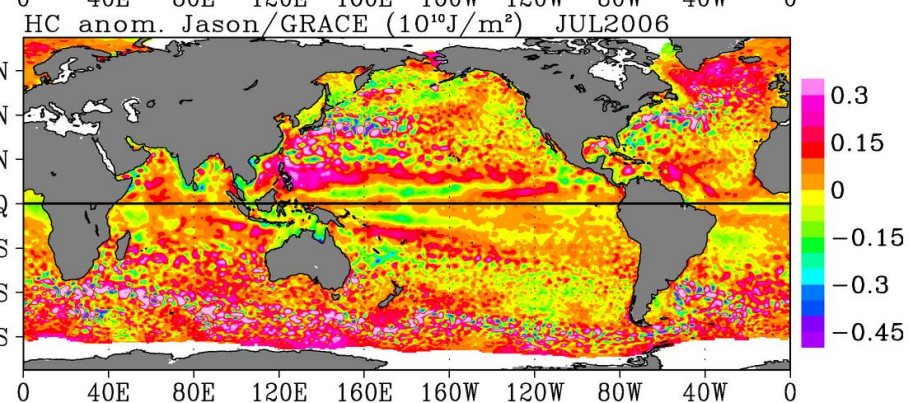
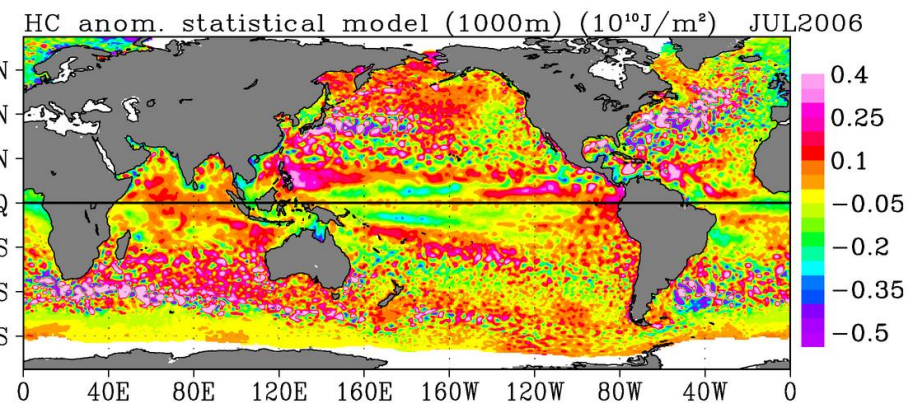
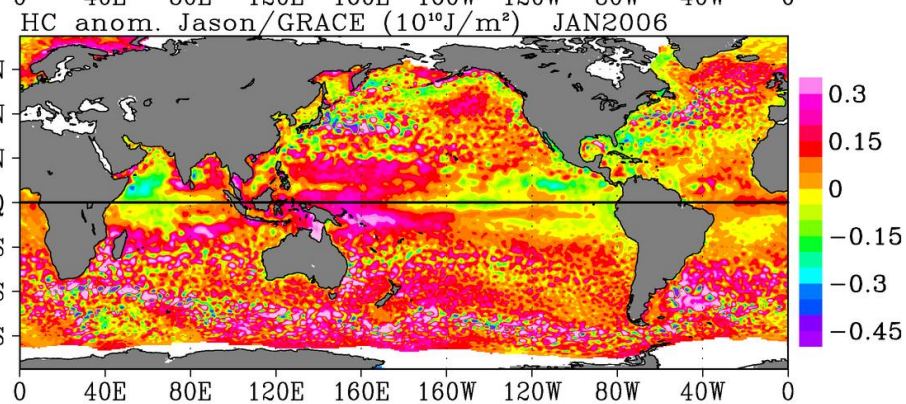
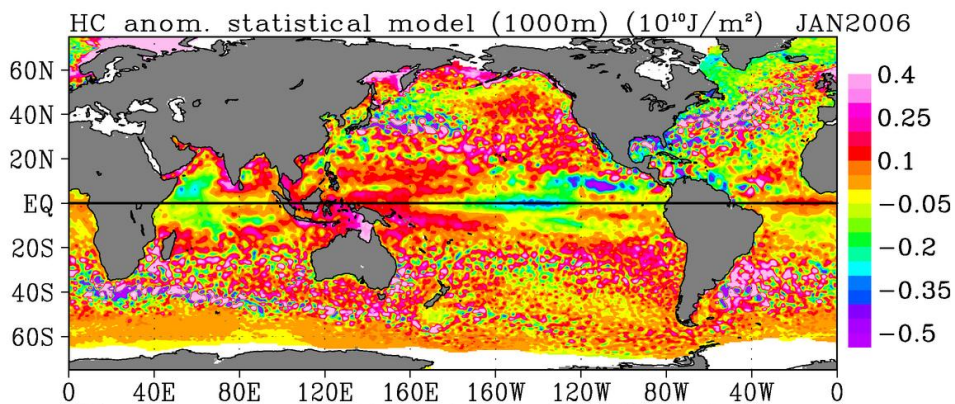
$$\iint \frac{\partial M}{\partial t} = \int R - \iiint \nabla \cdot \Theta$$

M could come from GRACE
or subtracting climatological steric change from
altimeter

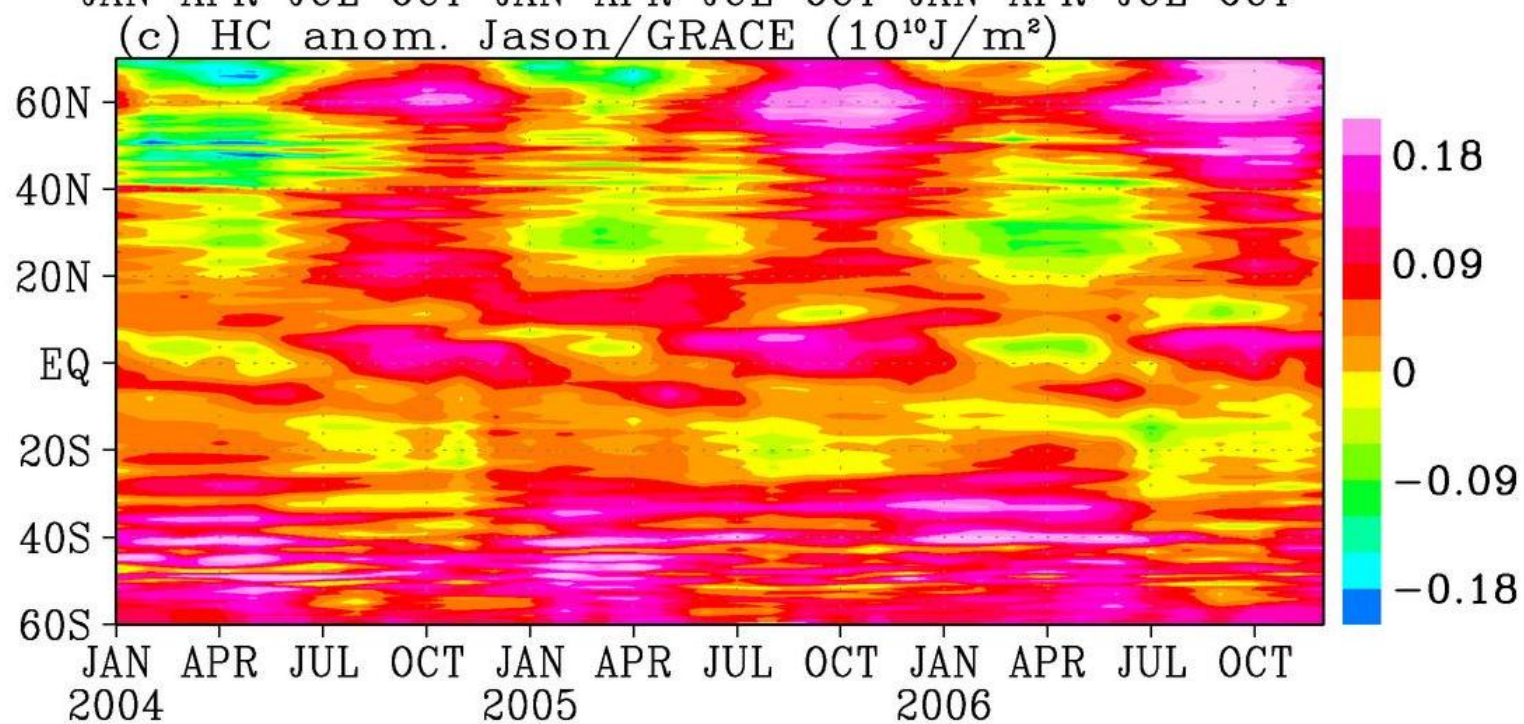
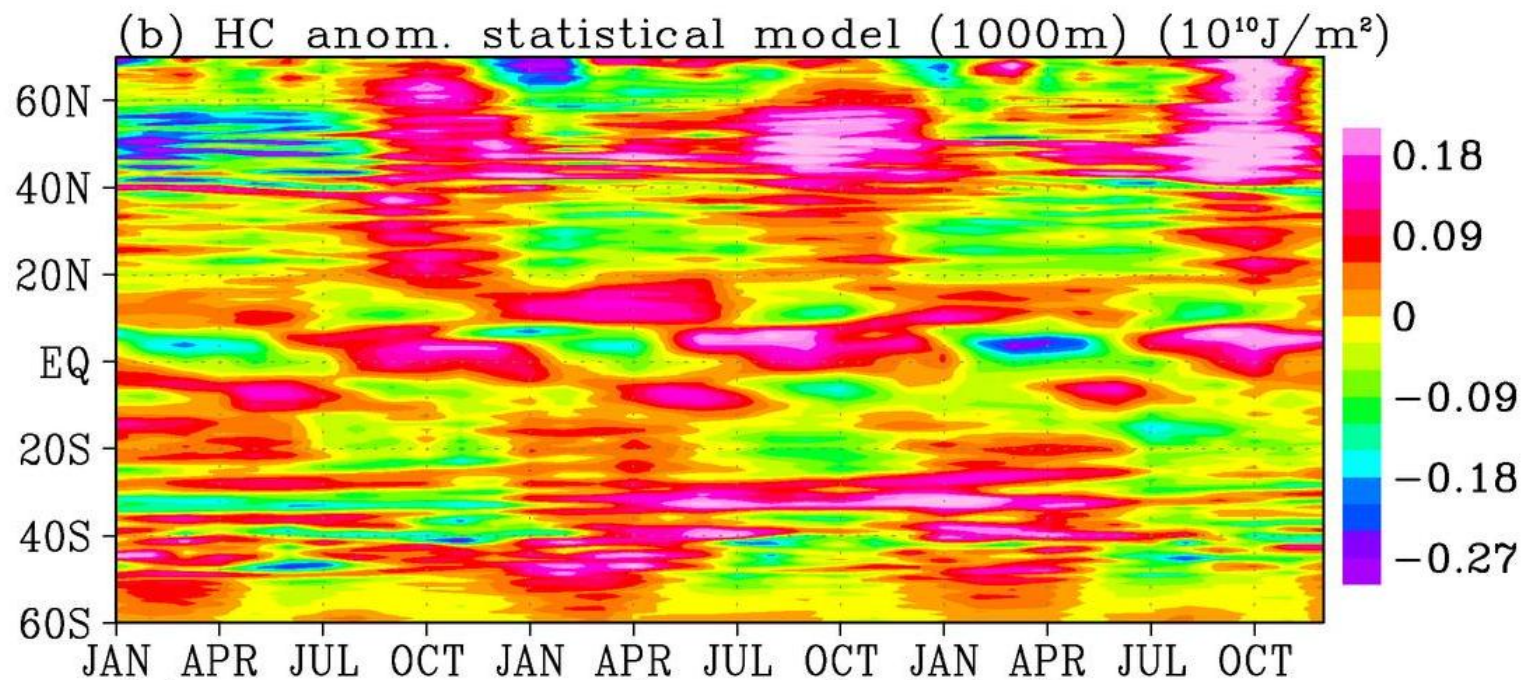
HC Jan 2006

Statistical model

HC Jul 2006



Jason/GRACE



Thermal expansion coeff. from Argo

