

# A brief overview of the air-sea fluxes of heat, momentum, and gas tracers at oceanic mesoscales

**Lucas Laurindo**

*University of Miami Cooperative Institute for Marine &  
Atmospheric Studies (CIMAS)*



# Introduction

- The surface fluxes are fundamental indicators of the coupling between the atmosphere and the ocean.
- In climate dynamics, a traditional view is that surface flux variability predominantly arises from the action of intrinsic atmospheric processes.
- However, mesoscale current and SST features can induce significant surface flux anomalies, with feedback to both oceanic and atmospheric circulations.
- Here, we'll revise the knowledge on the variability of air-sea fluxes of heat, momentum, and gas at the oceanic mesoscales.

## Turbulent heat flux (THF) variability

- Sensible and latent heat flux bulk formulae:

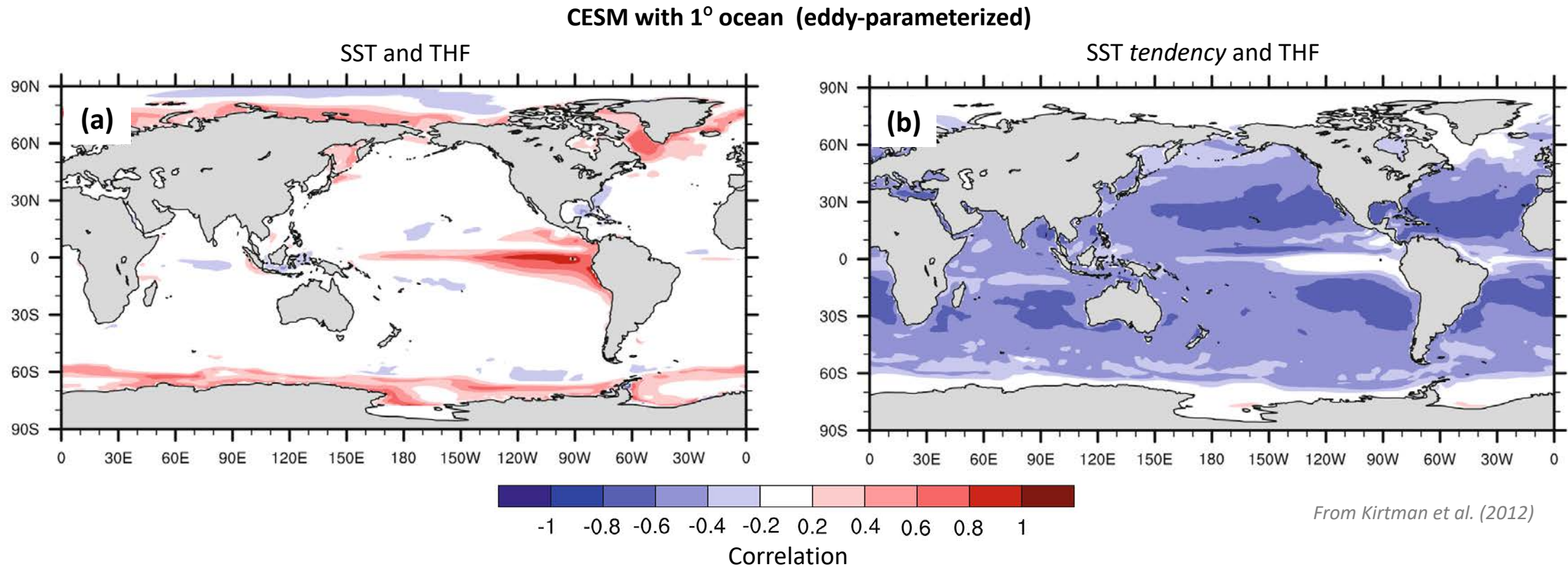
$$Q_s = \rho c_p c_s w (t_o - t_a)$$

$$Q_l = \rho L_e c_l w (q_o - q_a)$$

- Are proportional to the near-surface wind speed and the temperature and humidity contrast between the ocean and the atmosphere.
- Here, I'll define the heat fluxes as positive when out of the ocean.

## Turbulent heat flux (THF) variability

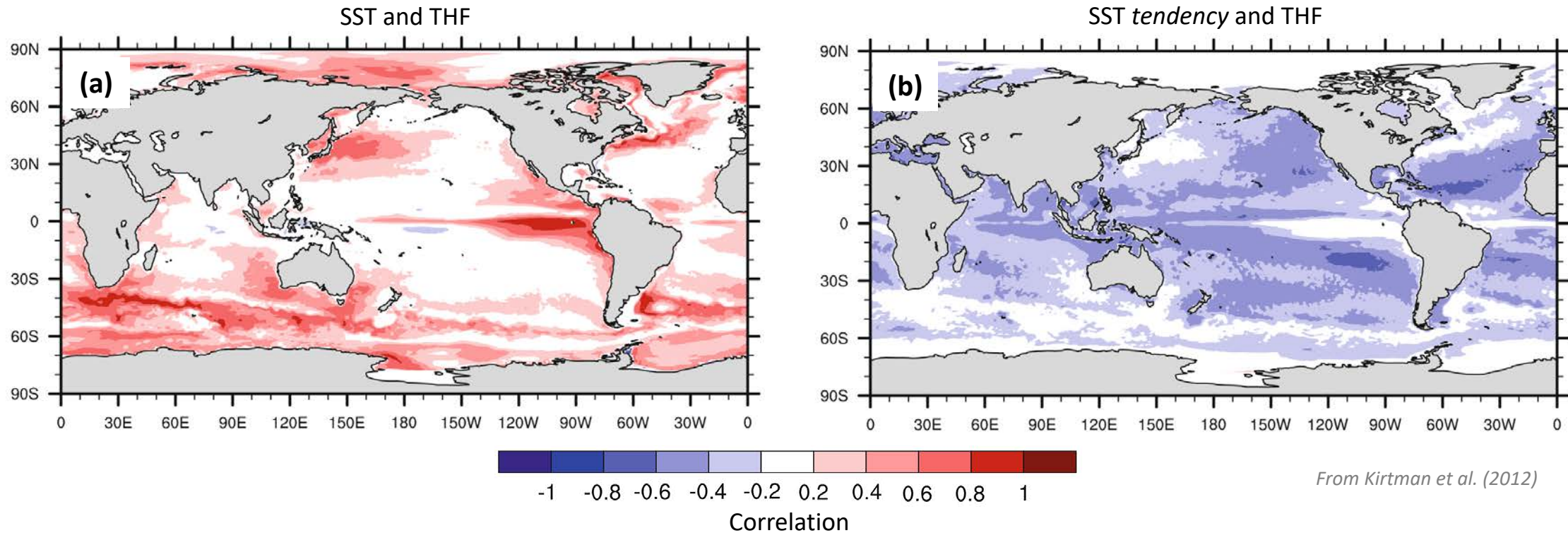
- The stochastic climate model theory of Hasselmann (1976) successfully predicted the SST power spectrum and SST/THF lag-correlations at midlatitude oceanic regions away from strong currents.



## Influence of mesoscale ocean dynamics

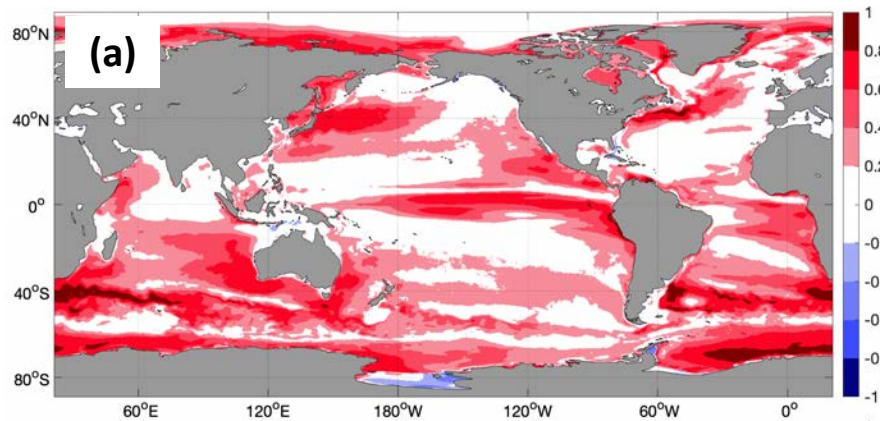
- Observations and eddy-resolving coupled model simulations produce positive SST/THF correlations and near-zero SST tendency/THF correlations at energetic current systems, such as the Kuroshio Current and Gulf Stream.

CESM with 1/10° ocean (eddy-resolving)



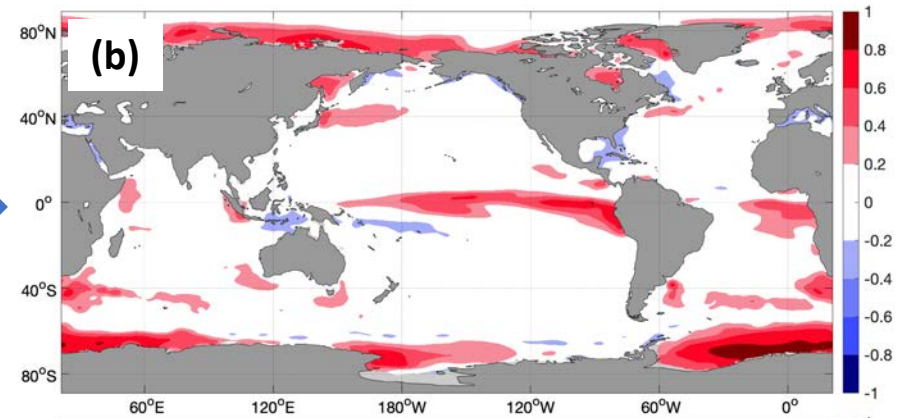
# The correlations depend on the spatial scale

SST and THF correlation

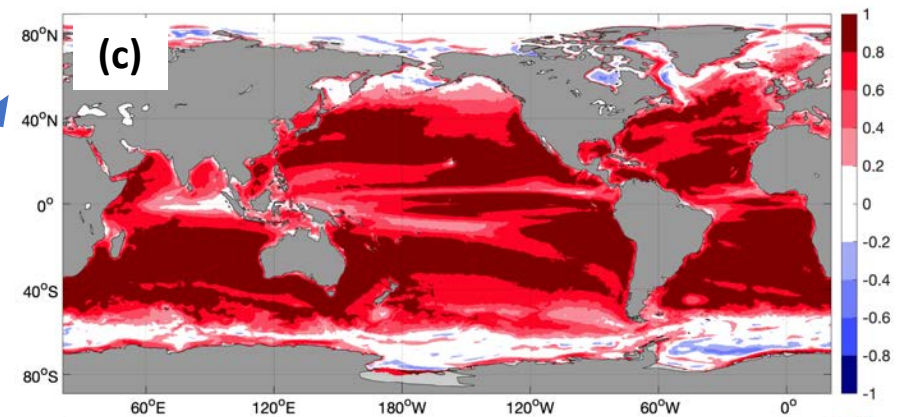


*iHESP CESM HR preindustrial control run.*

Large-scales ( $> 10^\circ$ )



Oceanic mesoscales ( $< 10^\circ$ )

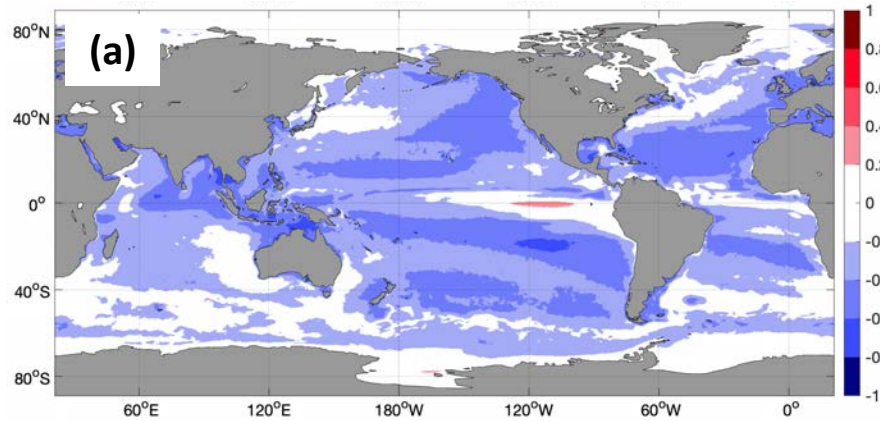


*Based on Bishop et al. (2017) and Small et al. (2019, 2020).*

- The atmosphere drives the SST and THF variability at large spatial scales.
- The ocean processes dominate the variability at the mesoscales.

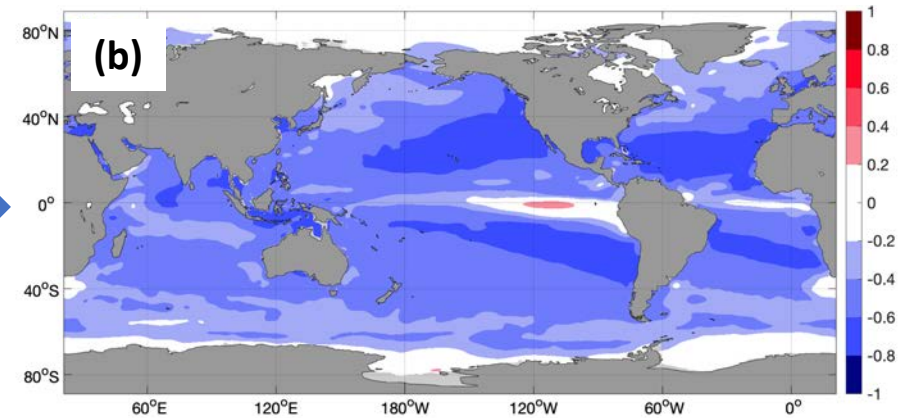
# The correlations depend on the spatial scale

SST tendency and THF correlation

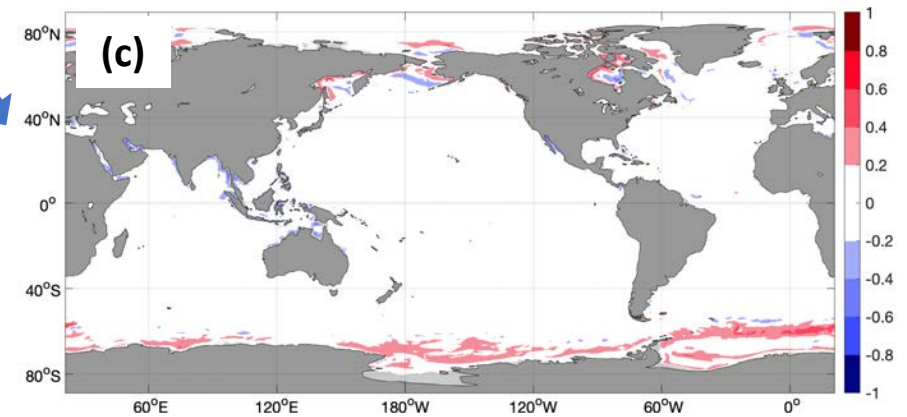


*iHESP CESM HR preindustrial control run.*

Large-scales ( $> 10^\circ$ )



Oceanic mesoscales ( $< 10^\circ$ )

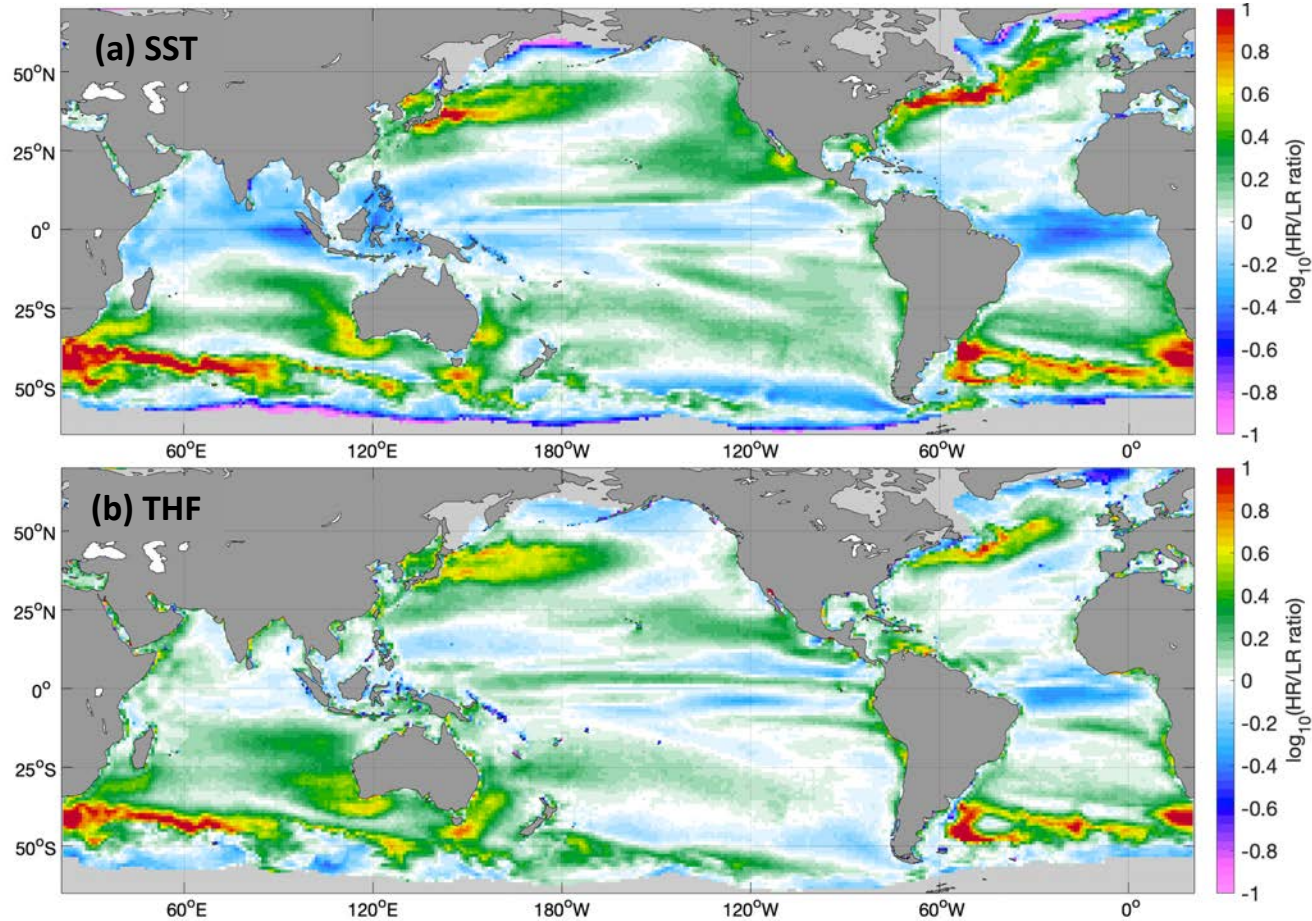


*Based on Bishop et al. (2017) and Small et al. (2019, 2020).*

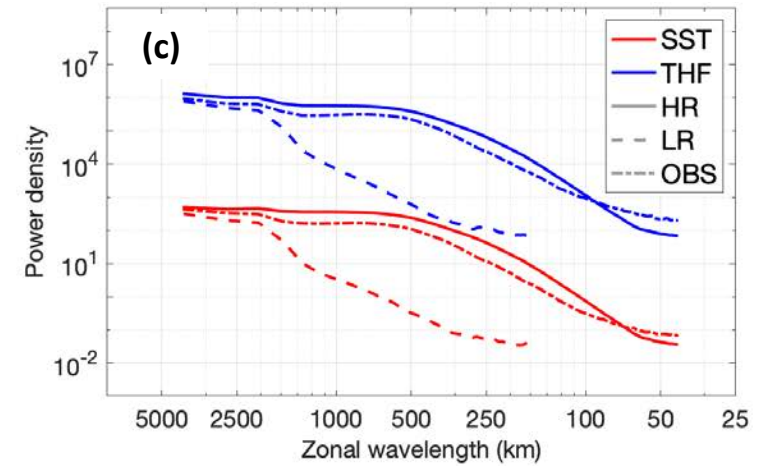
- The atmosphere drives the SST and THF variability at large spatial scales.
- The ocean processes dominate the variability at the mesoscales.

# SST and turbulent heat flux variances

HR/LR monthly variance ratio



SST and THF power spectra



*Spectra for the Agulhas Return Current*

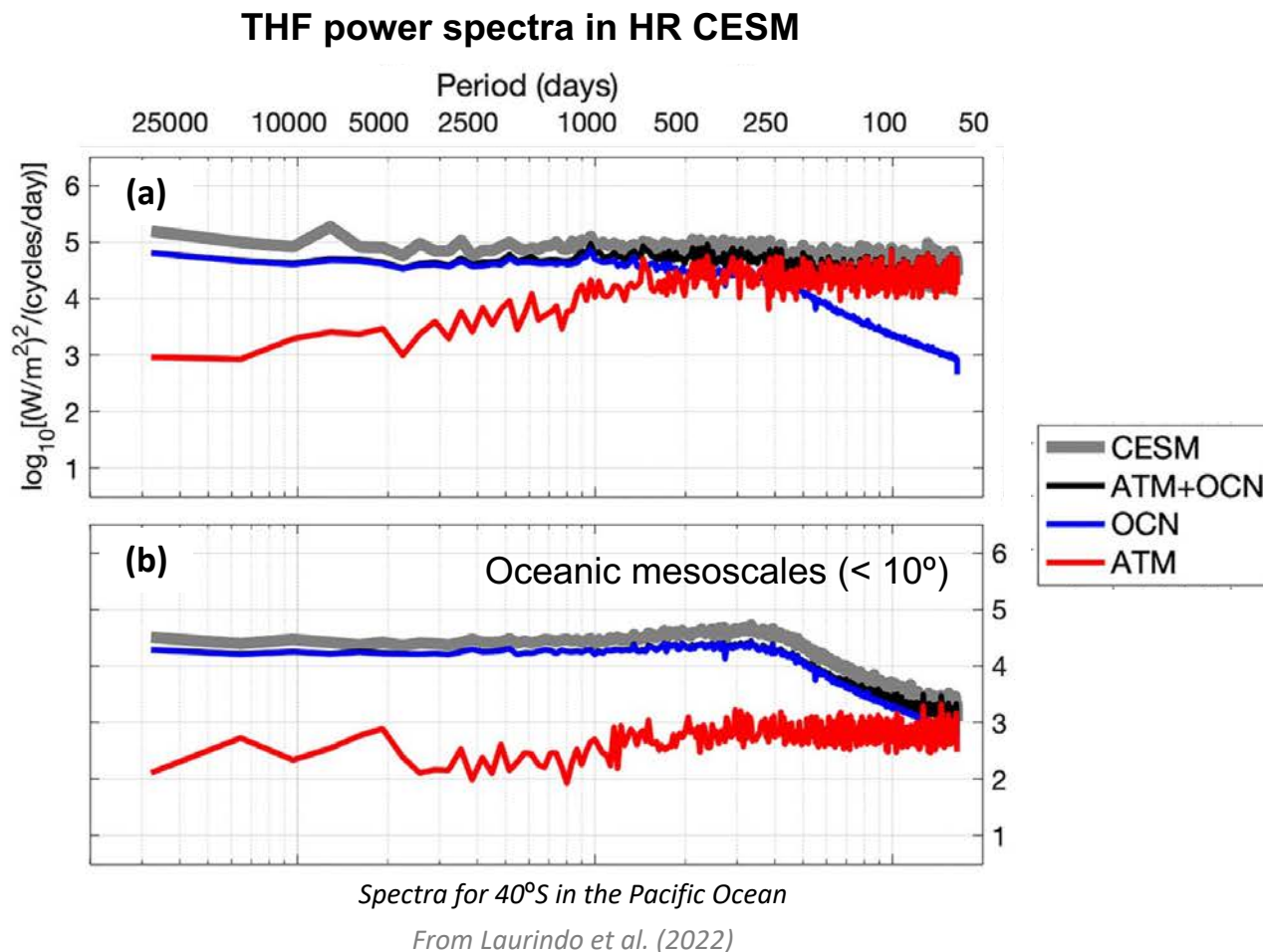
*From Chang et al. (2020)*



# Role of the ocean and atmosphere across different timescales

- Turbulent heat flux variability at monthly timescales is driven by weather.
- At longer timescales, variability is forced by ocean dynamics.

(e.g., Bishop et al. 2017, Sun and Wu 2021, Laurindo et al. 2022)



# Horizontal momentum flux (wind stress)

- Wind stress bulk formulation:

$$\boldsymbol{\tau} = \rho_{air} c_d |\mathbf{w} - \mathbf{u}| (\mathbf{w} - \mathbf{u})$$

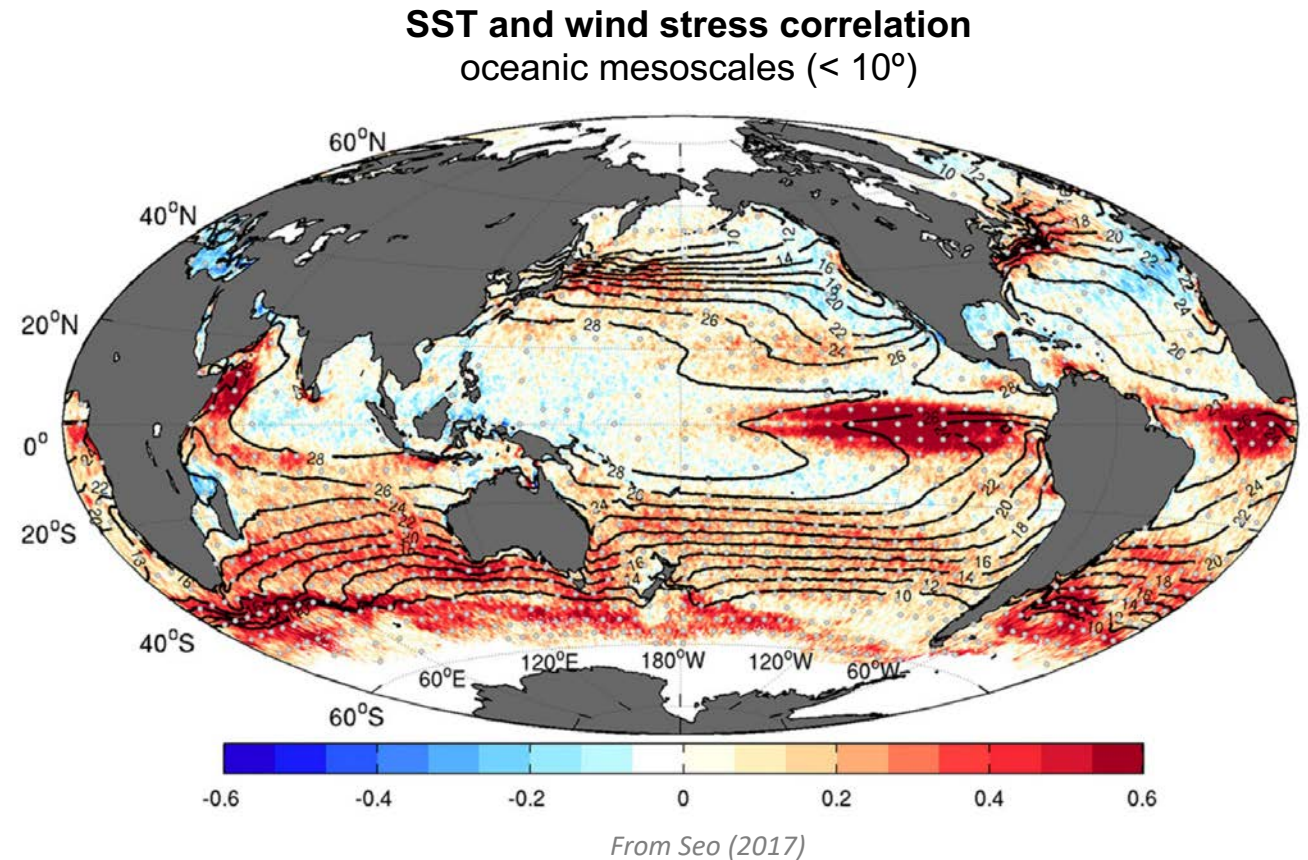
Diagram illustrating the bulk formulation of wind stress. The vector  $\mathbf{w}$  is identified as 10-m winds, and the vector  $\mathbf{u}$  is identified as surface ocean currents. The relationship between  $\mathbf{w}$  and the wind speed  $\mathbf{w}_{LS}$  and SST  $\mathbf{w}_{SST}$  is given by:

$$\mathbf{w} = \mathbf{w}_{LS} + \mathbf{w}_{SST}$$

- Influence of SST:
  - *Modifies turbulent mixing, drag, and pressure gradients within the marine atmospheric boundary layer (MABL).*
  - *Ultimately accelerates the near-surface winds moving from cool to warm SST and decelerates winds moving from warm to cool SST.*
- Influence of surface ocean currents:
  - *Wind stress acts relative to the ocean surface, which is in motion.*

# Characterizing the SST-induced response in MABL

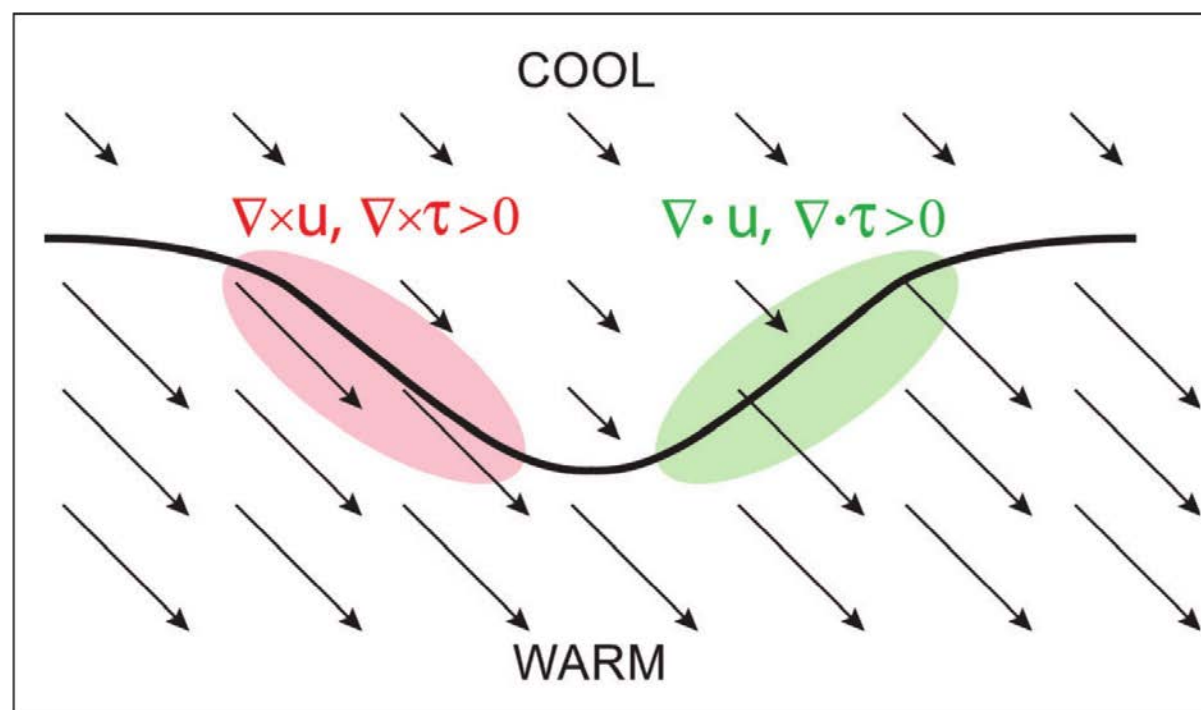
- It is usually characterized by linear regressions between SST and 10-m wind speed/wind stress.
- Data needs to be high-pass filtered at scales of about 1000 km or smaller.
- Data is also frequently smoothed in time or averaged over a few weeks.



## Characterizing the SST-induced response

- Positive correlations also develop between:
  - The downwind component of the SST gradient and the wind curl/wind stress curl.
  - The cross-wind component of the SST gradient and the wind stress divergence.
- Correlations between the Laplacian of SST and the Laplacian of sea-level pressure are also frequently used.

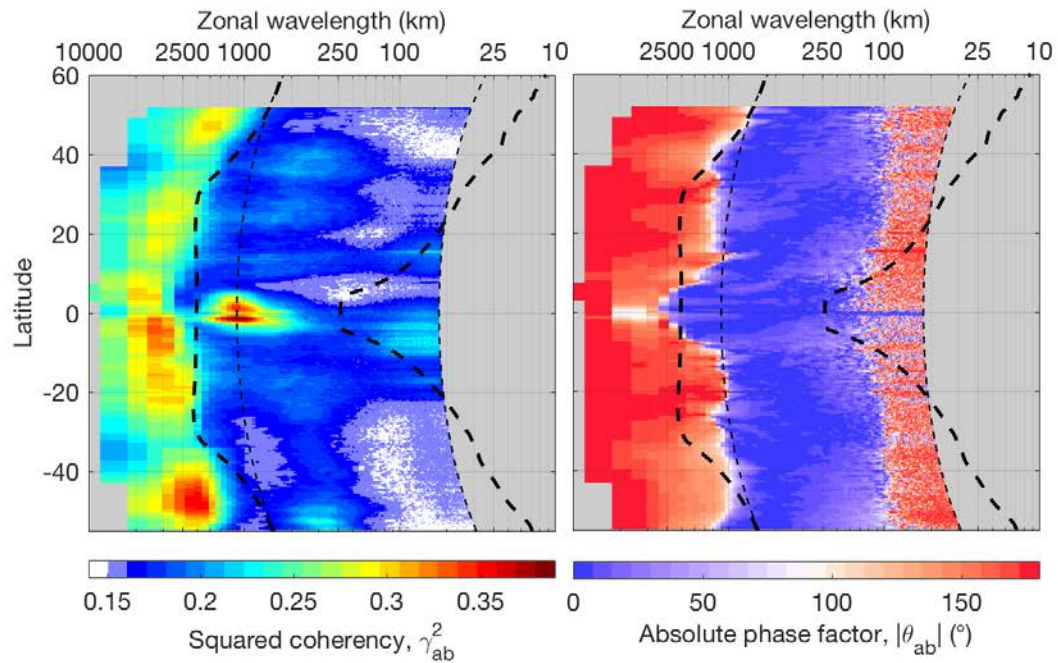
Schematic illustration of winds blowing across a meandering SST front



From Chelton et al. (2010)

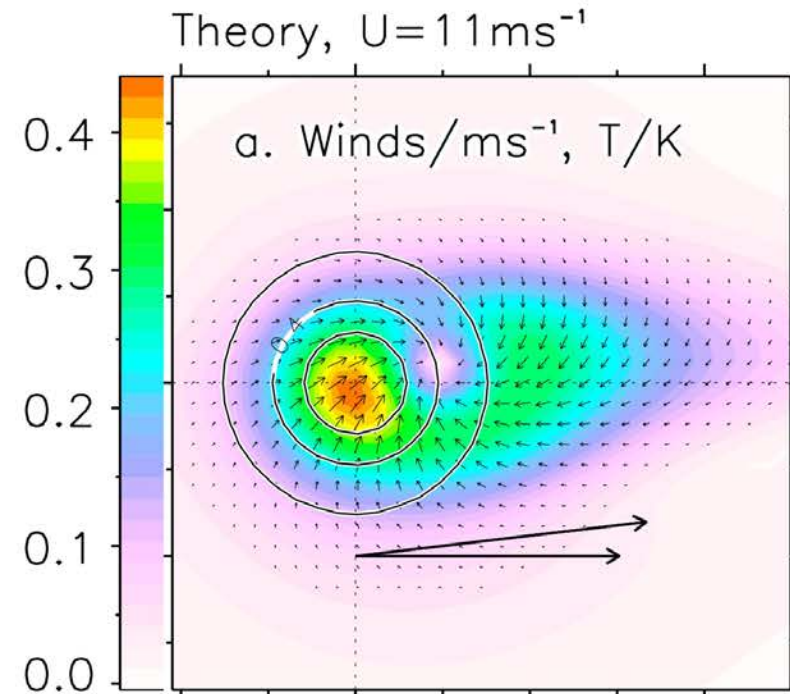
# Spectral methods

- Zonal wavenumber coherence and phase factor between SST and 10-m wind speed, Pacific Ocean.



From Laurindo et al. (2019)

- Surface wind response to a Gaussian monopole SST anomaly of 1 K.
- Computed using spectral transfer functions derived from the linear MABL model of Schneider and Qiu (2015).



From Schneider (2020)

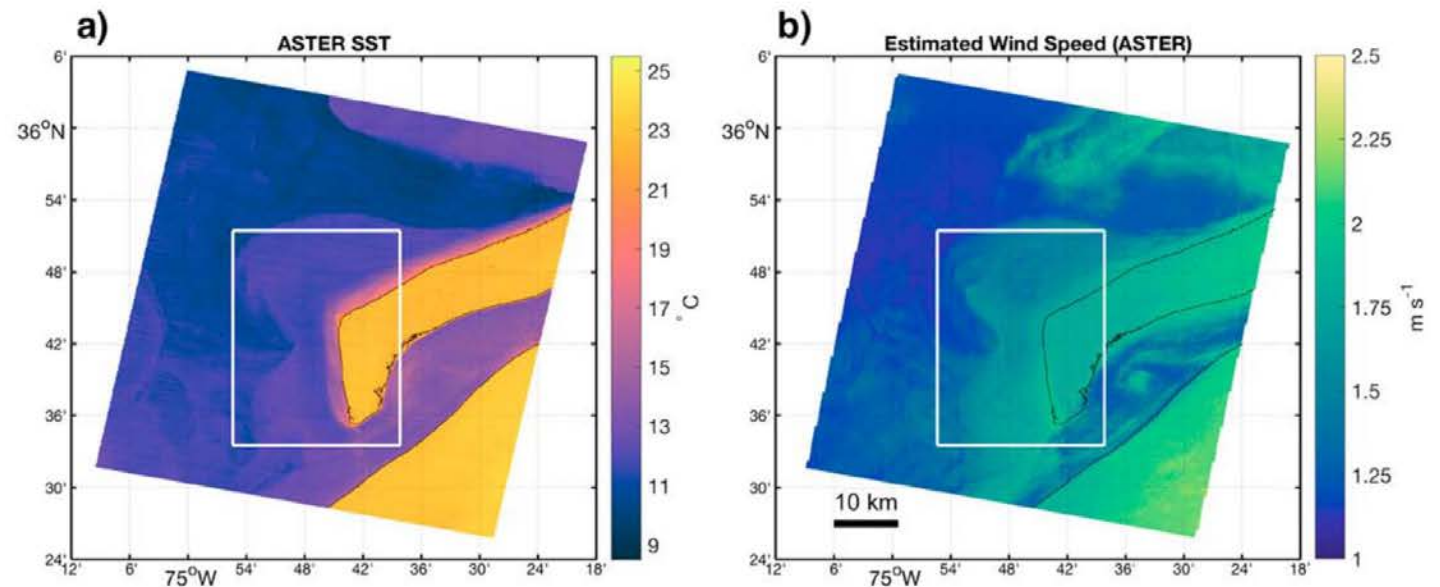
## Submesoscale SST-driven response

- Papers on the influence of SST in the MABL over submesoscale ranges (1-10 km) are beginning to appear in literature.

- Meroni et al. (2022): standard metrics may be insensitive to the atmospheric response to SST forcing due to the influence of advection.

- *Suggests using the linear relationship between wind divergence and the SST second derivative.*

- *SST and wind speed fields over a submesoscale filament associated with the Gulf Stream from the Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER)*



From Gaube et al. (2019)

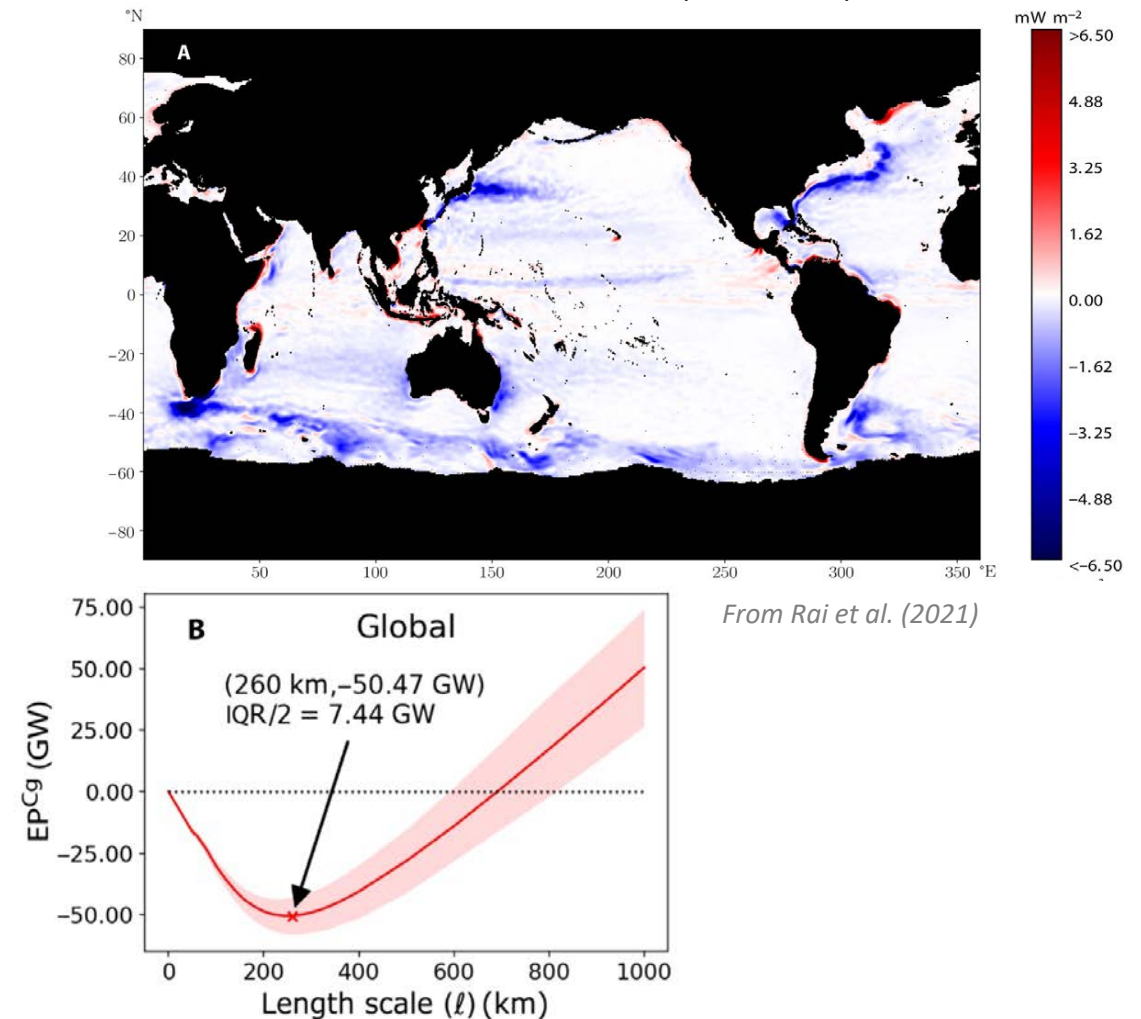
# Influence of surface ocean currents

- Wind stress bulk formulation:

$$\boldsymbol{\tau} = \rho_{air} c_d |\mathbf{w} - \mathbf{u}| (\mathbf{w} - \mathbf{u})$$

- Accounting for the influence of surface currents ( $\mathbf{u}$ ) is usually a small correction to wind stress.
- However, it greatly impacts wind stress curl and eddy energetics.
- It also affects the low-level wind shear.

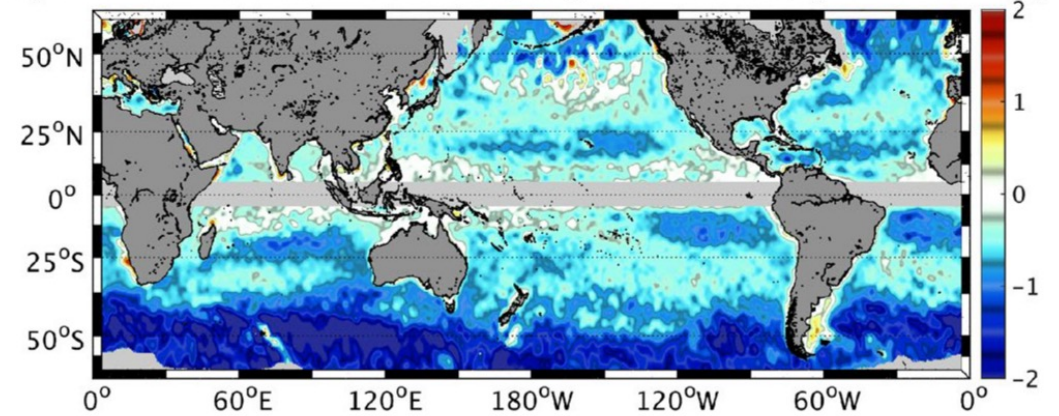
Covariance btw geostrophic currents and wind stress  
oceanic mesoscales (< 250 km)



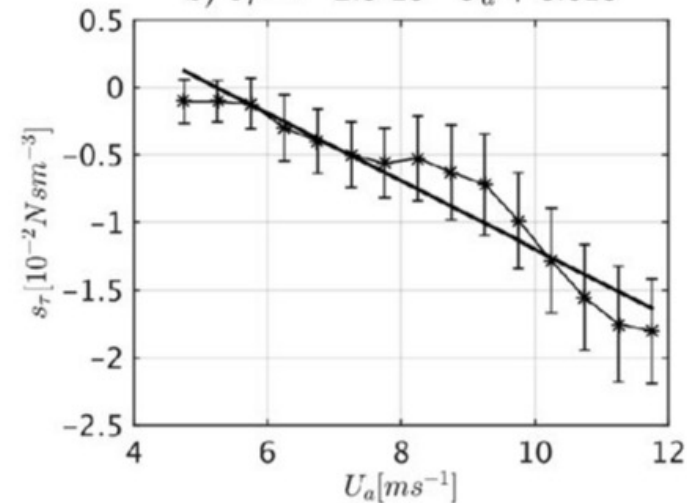
## Characterizing the current-induced response

- Separating the SST-driven modifications of the wind stress and derived quantities from those induced by surface currents is not trivial.
- Renault et al. (2017): the current-driven response can be characterized via the linear regression coefficient between the wind stress curl and surface current curl ( $s_\tau$ ).

a) Observed Current-Stress Coupling Coefficient  $s_\tau$  [ $10^{-2} \text{ N s m}^{-3}$ ]



b)  $s_\tau = -2.5 \cdot 10^{-3} U_a + 0.013$



From Renault et al. (2017)



## Air-sea flux of gas tracers

- The air-sea flux of a gas  $x$  can be given by:

$$F_x = k(C - C_{eq})$$

- where:

- $k$ : gas transfer velocity.
- $C$ : gas concentration at seawater.
- $C_{eq}$ : gas concentration in equilibrium with the atmosphere

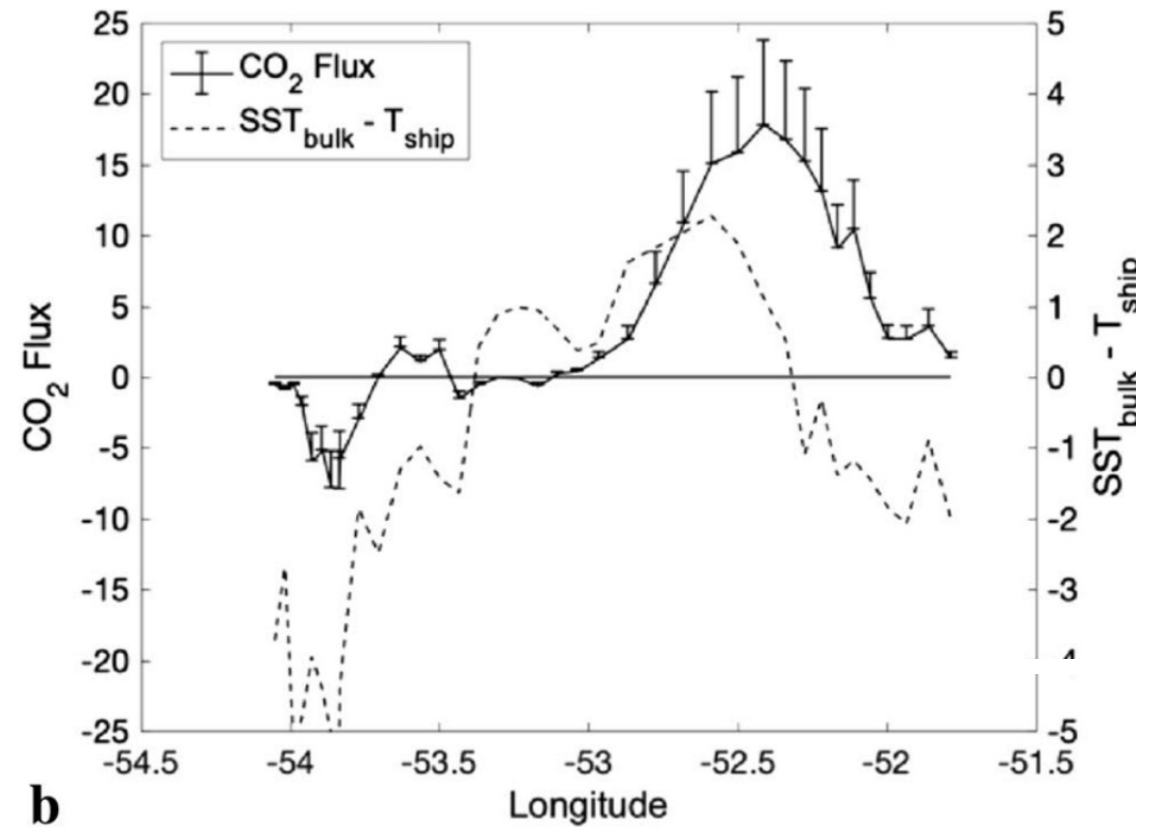
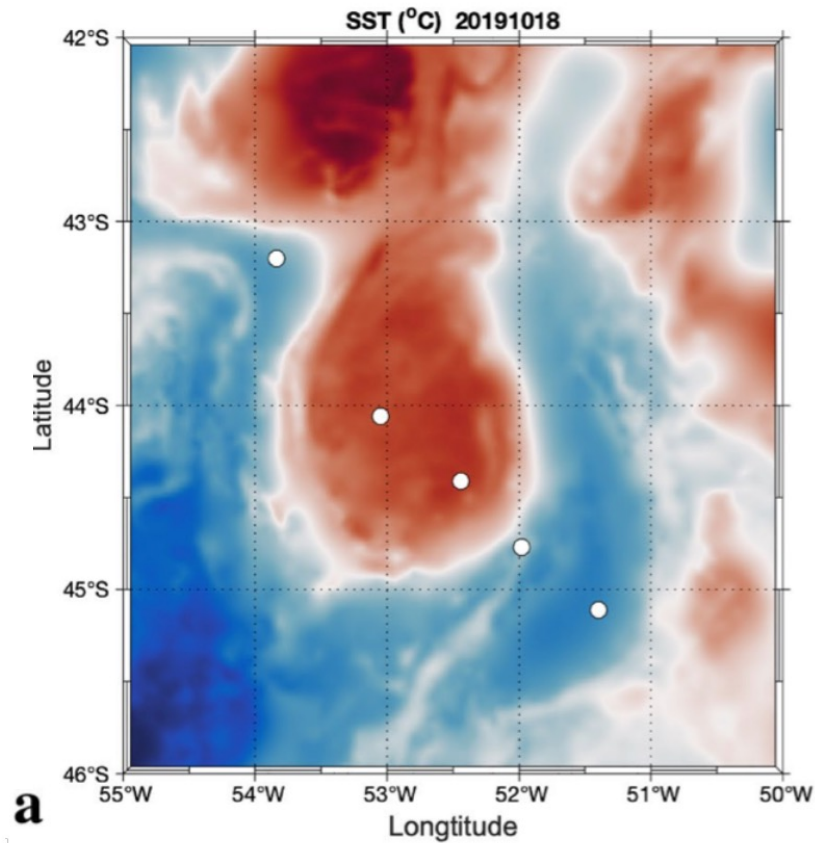
- Mesoscales can impact  $F_x$  via  $k$  or  $C_{eq}$ :

- *Parameters vary nonlinearly with wind speed and depend on sea state*

- Mesoscales can also impact  $C$ :

- *Influence on biological sources and sinks of gas tracers.*

# Air-sea flux of gas tracers



From Seo et al. (2023), reproduced from Pezzi et al. (2013)

# Thank you!

[llaurindo@earth.miami.edu](mailto:llaurindo@earth.miami.edu)