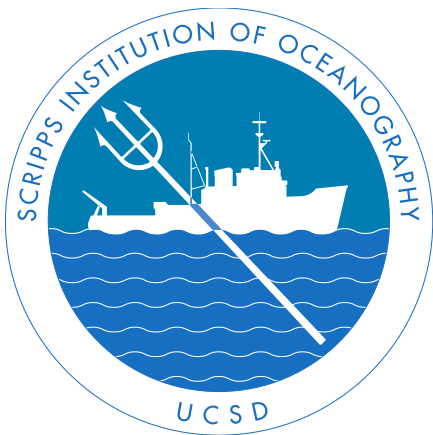


# Rethinking the AMOC stability in climate models

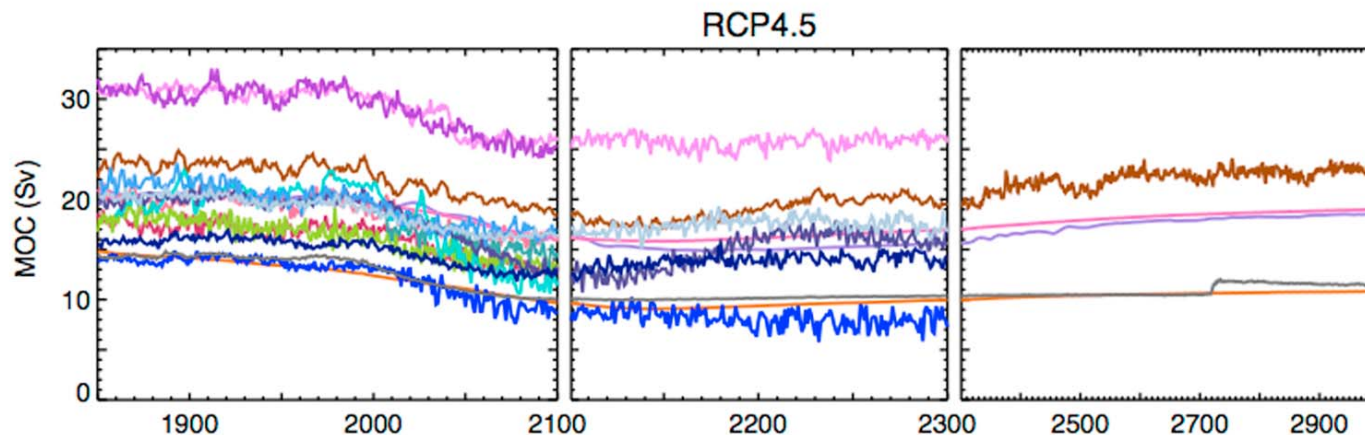
Wei Liu, Zhengyu Liu, Shang-Ping Xie,  
Esther Brady, Aixue Hu and Jiang Zhu

October 20, 2016  
US CLIVAR Webinar



# Background and motivation

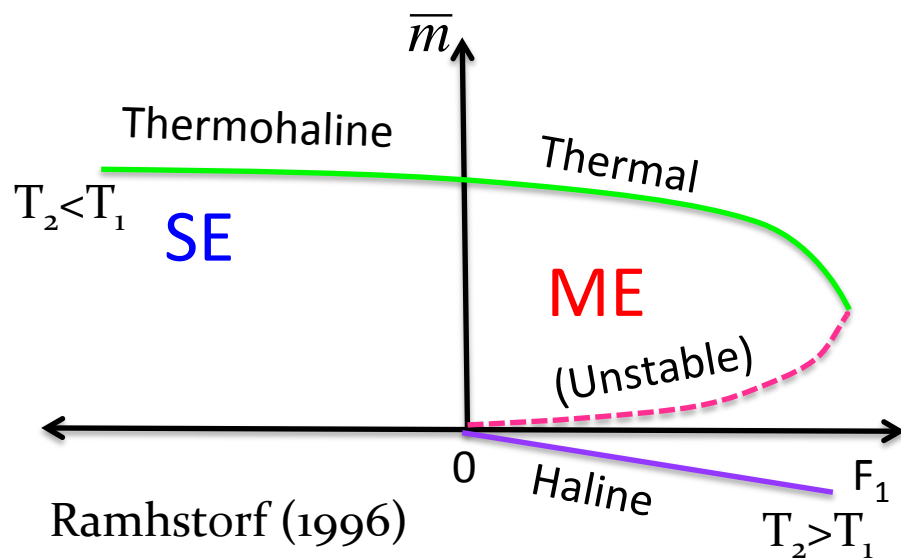
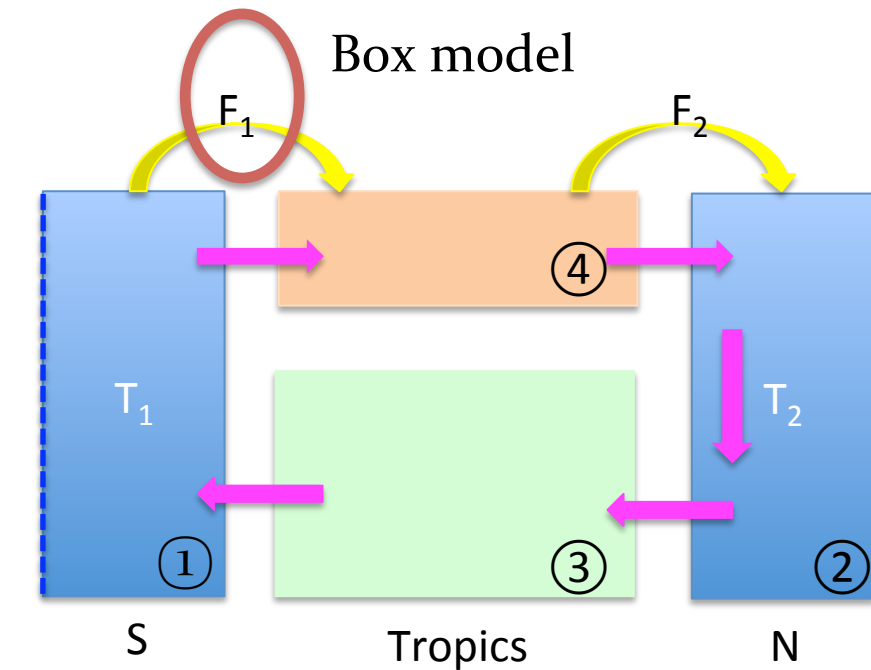
- An abruptly collapsed AMOC can cause abrupt climate change. This AMOC change is considered to be associated with multiple equilibria (ME) of the AMOC.
- Nevertheless climate models mostly simulate AMOCs with Single equilibrium (SE).
- Example 1: A strong and long lasting freshwater forcing is needed to keep an collapsed AMOC during Heinrich Event 1 (Liu et al. 2009).
- Example 2: The AMOC recovers in pulse-like hosing experiments (e.g., Stouffer et al. 2006).
- Also, the CMIP5 model projection shows a moderately weakened but not collapsed AMOC till 2300 (e.g., Weaver et al. 2012).



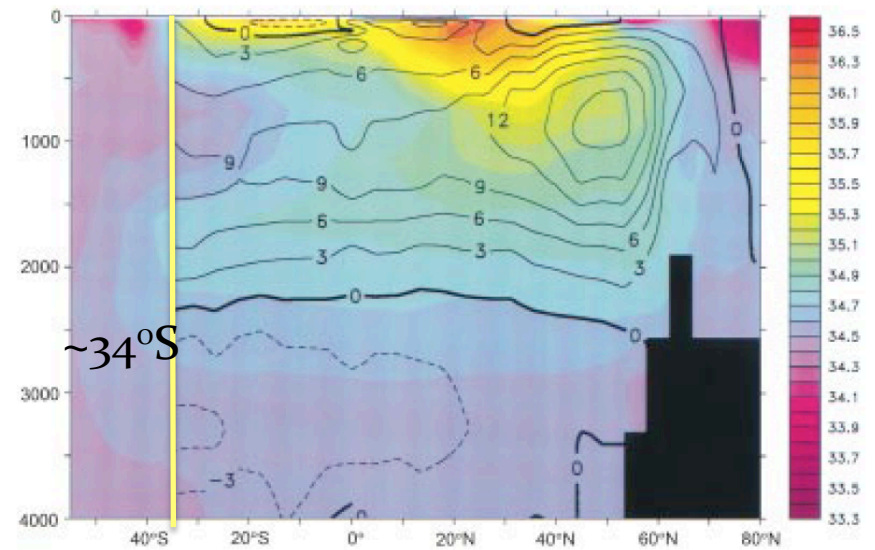
## Scientific questions

- a. What is the indicator of the AMOC stability in fully coupled climate models?
- b. Why do most state-of-art climate models fail to obtain AMOCs with ME?
- c. How does this AMOC stability bias affect model projection?

## a. The AMOC stability indicator



**Climate model**



$$[E - P - R] = M_{ov} + M_{az} + M_{dif} + M_{BS}$$

FWT by MOC

FWT by gyre

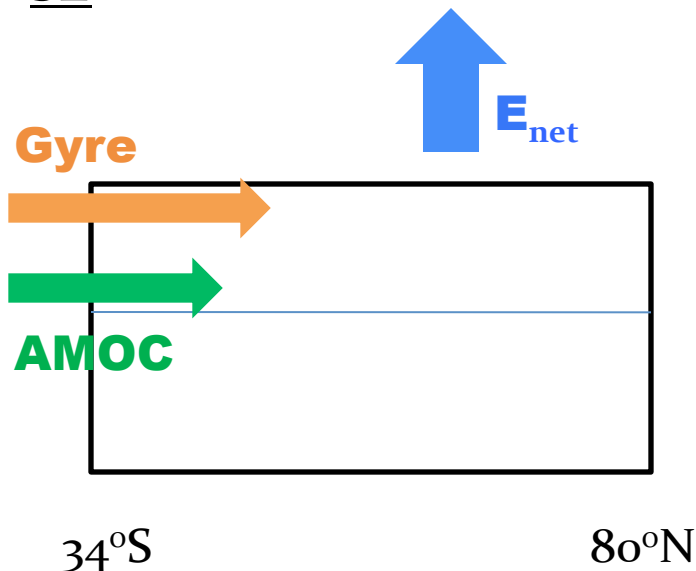
$$M_{ov} = -\frac{1}{S_0} \int dz \bar{v}(z) [\langle S(z) \rangle - S_0]$$

**Indicator**

## a. The AMOC stability indicator

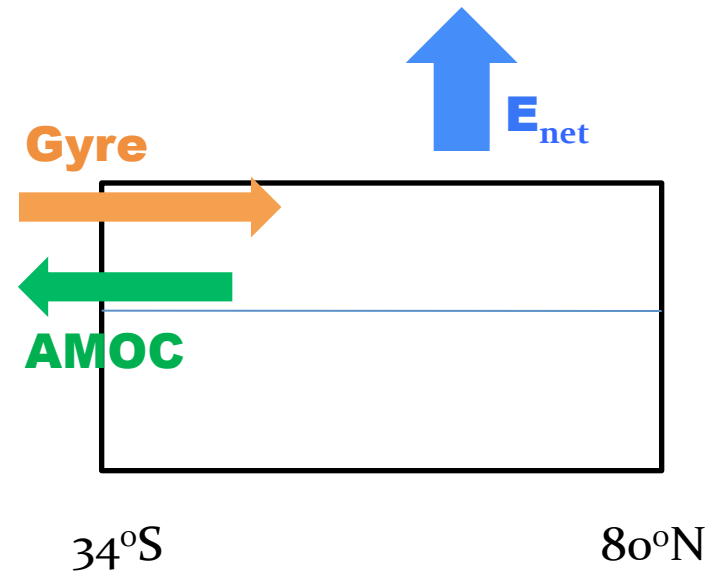
Why does  $M_{ov}$  act as an AMOC stability indicator?

SE



$M_{ov} > 0$ , freshwater import,  
AMOC recovery,  
SE(monostable)

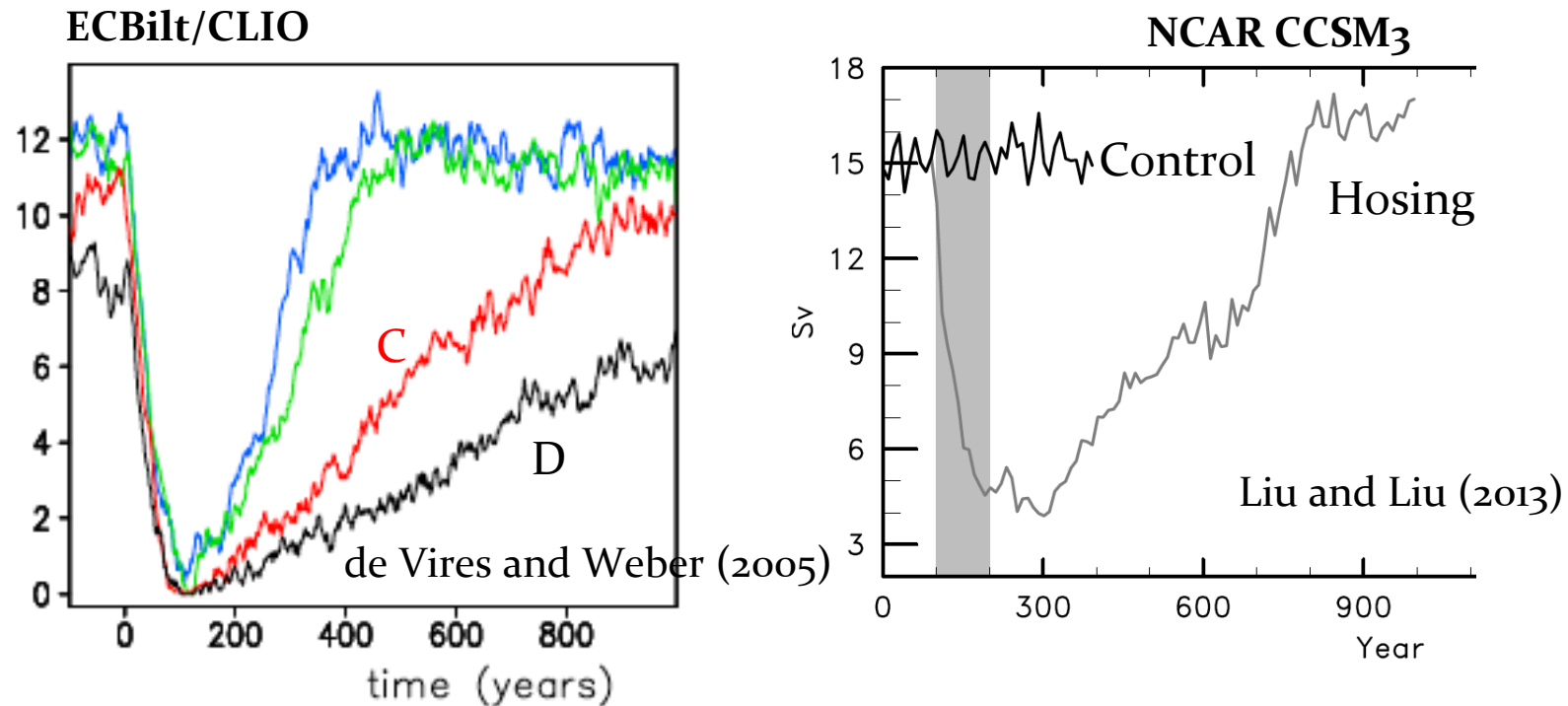
ME



$M_{ov} < 0$ , freshwater export,  
No AMOC recovery  
ME(bistable)

## a. The AMOC stability indicator

$M_{ov}$  may not work



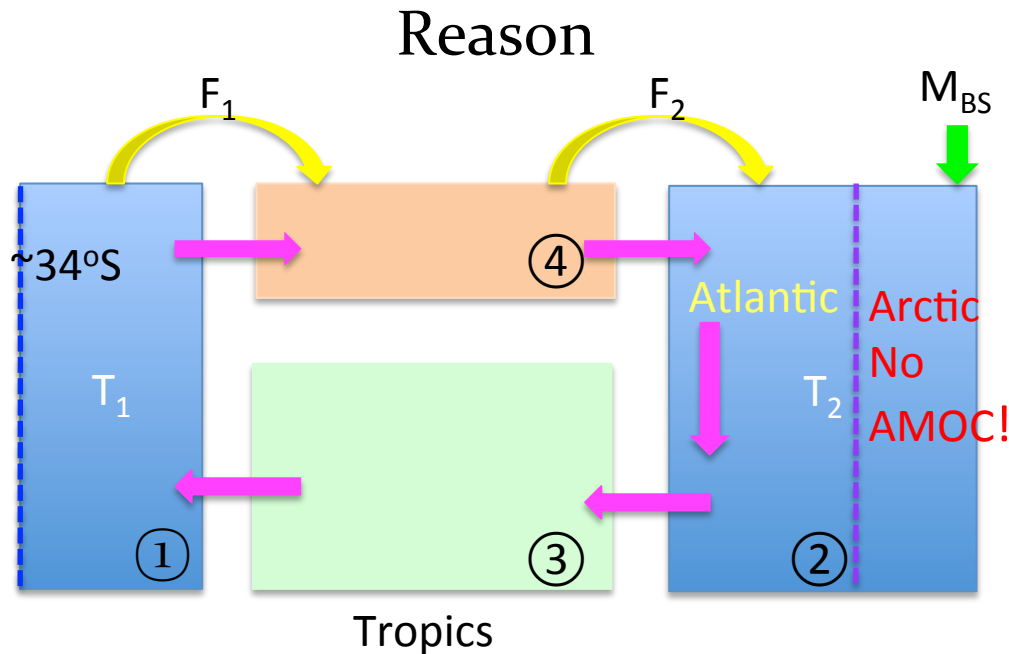
The ECBilt/CLIO Runs C & D and CCSM<sub>3</sub> PD control run show negative  $M_{ov}$  (FW exports at  $\sim 34^\circ\text{S}$ ) but AMOCs in a SE regime.

An alternative indicator? (Dijkstra 2007)

$$\Sigma(\theta_s, \theta_n) = M_{ov}(\theta_s) - M_{ov}(\theta_n).$$

$$\Theta_s = 35^\circ\text{S}; \Theta_n = 60^\circ\text{N}$$

## a. The AMOC stability indicator

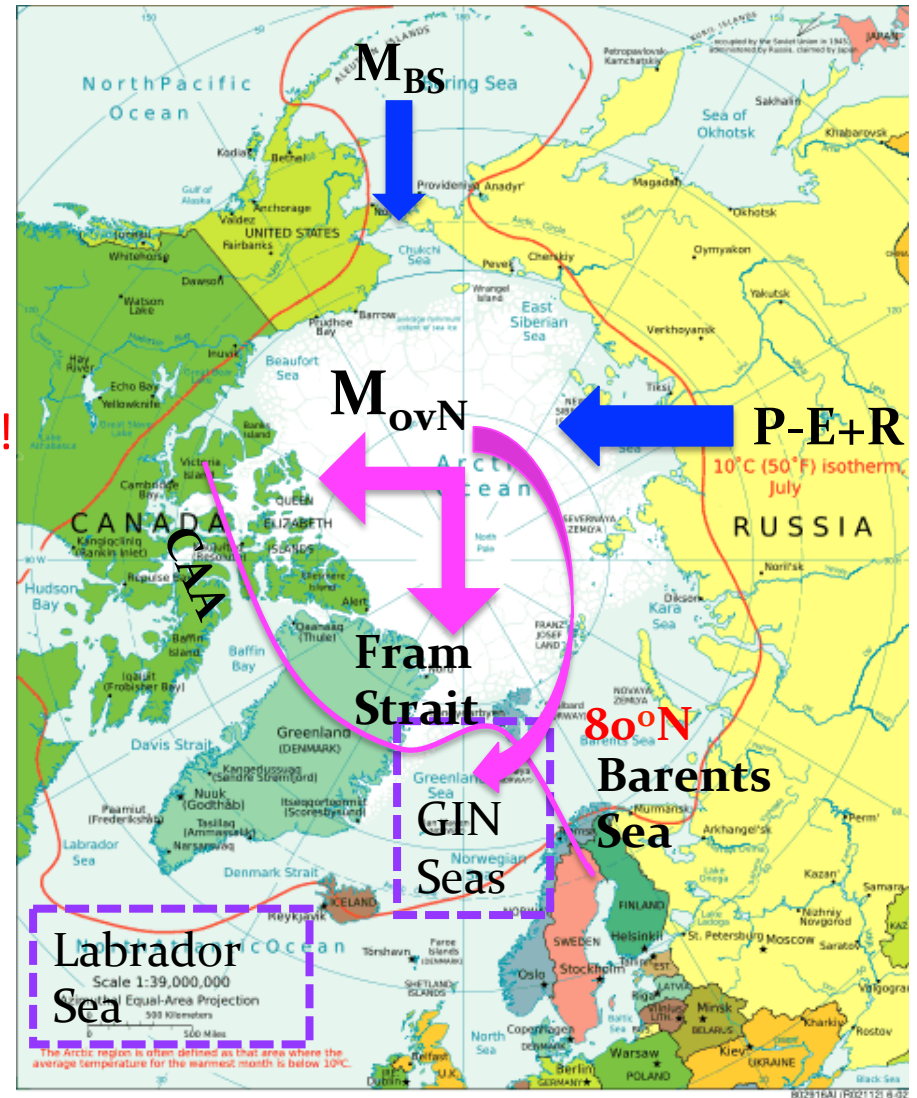


A refined indicator  
(Liu and Liu 2013,2014)

$$\Delta M_{ov} = M_{ovS} - M_{ovN}$$

$\Theta_s \sim 34^\circ\text{S}$ : the boarder b/ the Atlantic & SO

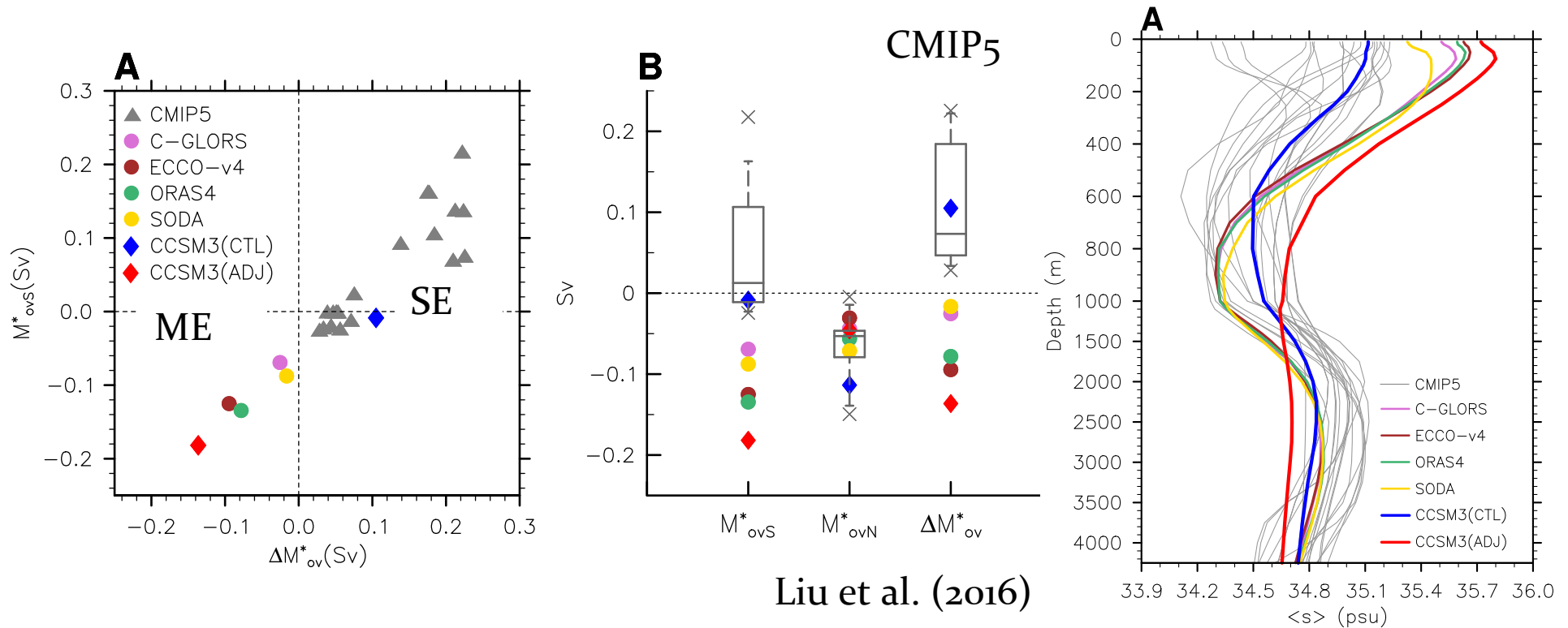
$\Theta_n \sim 80^\circ\text{N}$ : the boarder b/ the Atlantic & Arctic



$\Delta M_{ov} < 0$ , FW divergence, AMOC ME

$\Delta M_{ov} > 0$ , FW convergence, AMOC SE

## b. The AMOC stability bias in climate models



Similar results can also be found in the CMIP3 simulations (Liu et al. 2014)

- Observations suggest a FW divergence ( $\Delta M_{ov} < 0$ ) and an AMOC with ME.
- Climate models show a FW convergence ( $\Delta M_{ov} > 0$ ) and an AMOC with SE.
- This AMOC stability bias mainly comes from the southern boundary and is related to a fresh bias in the upper ocean of the South Atlantic.

## b. The AMOC stability bias in climate models

How to correct this bias? Flux adjustment (CCSM<sub>3</sub>)

Restore sfc heat flux  $H_{res} = (SST - SST_{obs})/\tau_T$

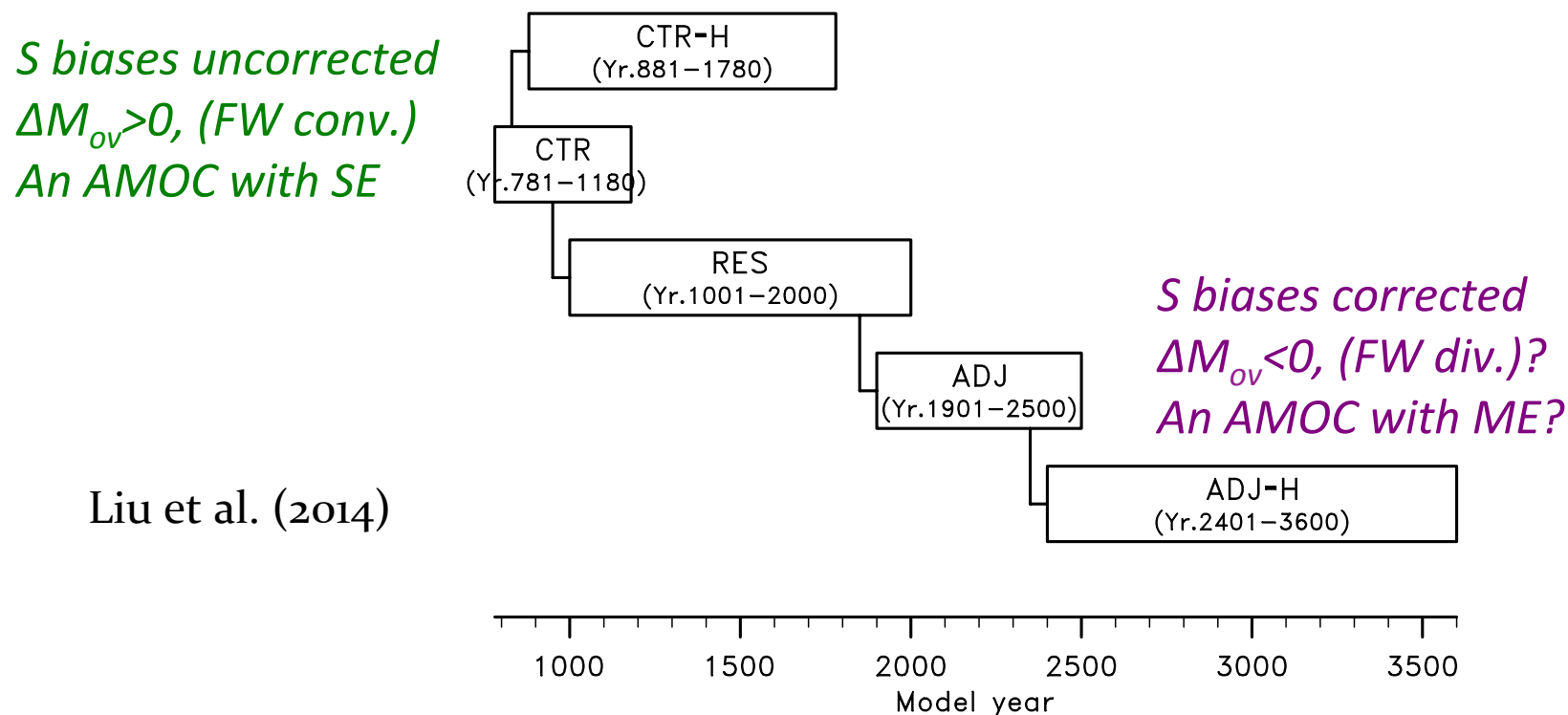
Restore sfc freshwater flux  $F_{res} = (SSS - SSS_{obs})/\tau_S$

$\xrightarrow{\text{equilibrium}}$

$H_{adj} = H_{res}$

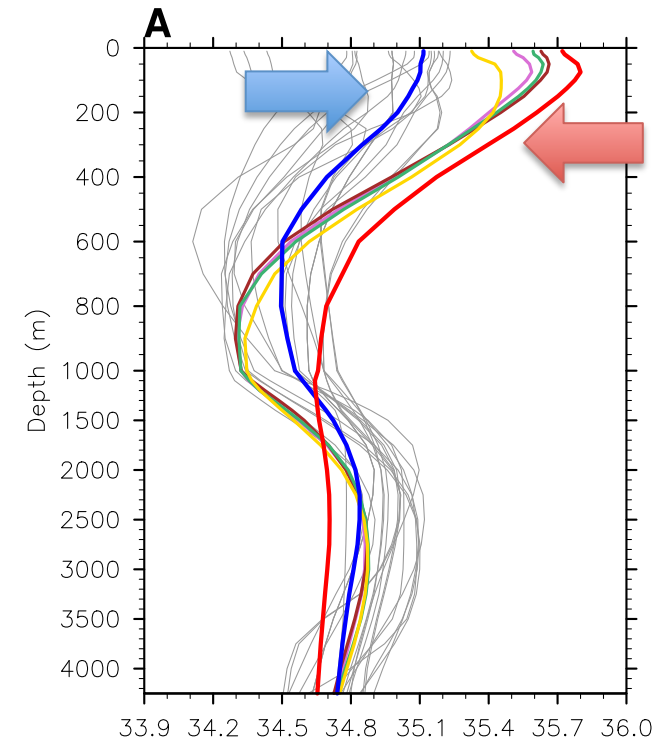
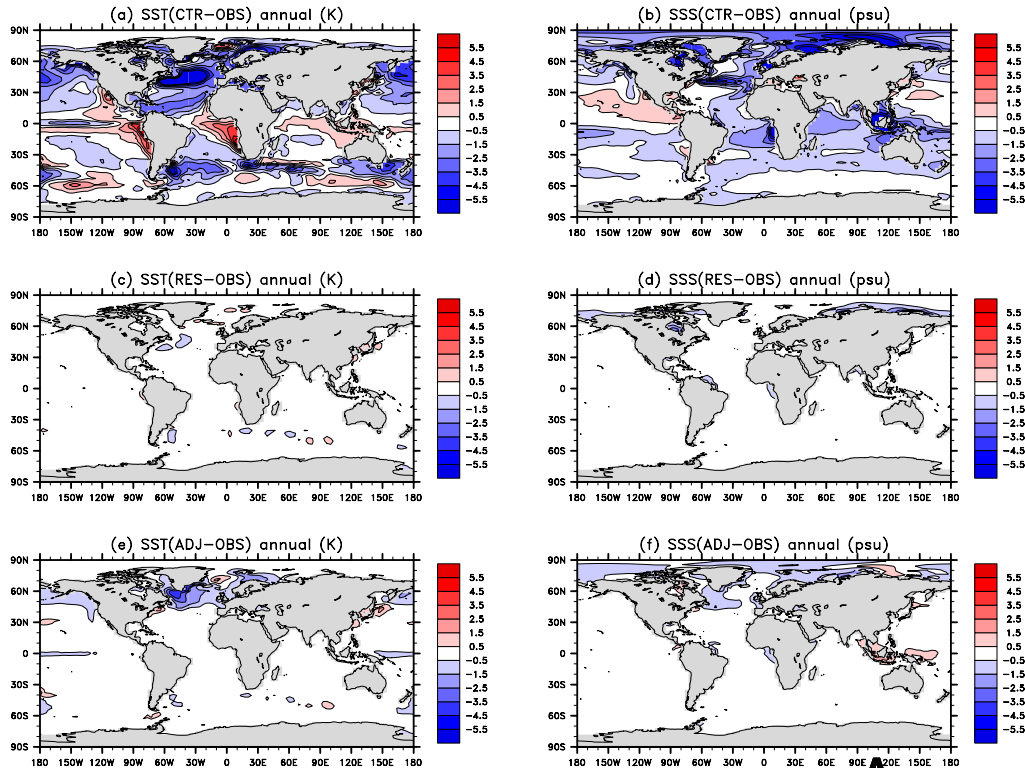
$F_{adj} = F_{res}$

Experiment Chart

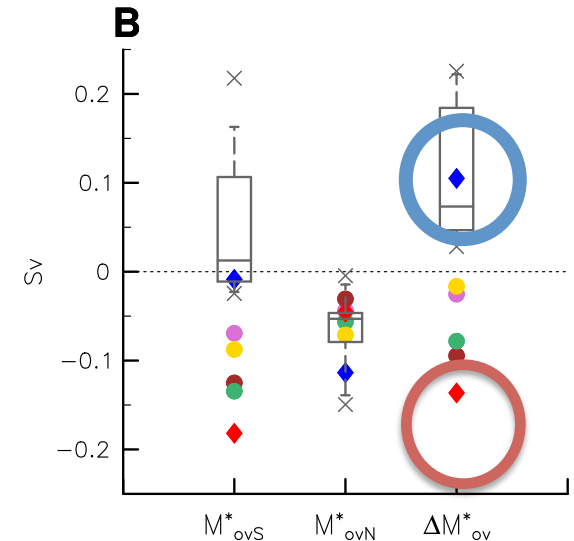
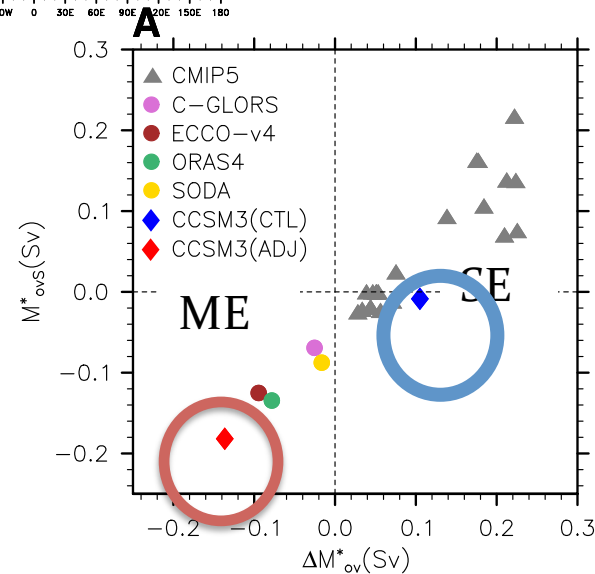


Liu et al. (2014)

## b. The AMOC stability bias in climate models

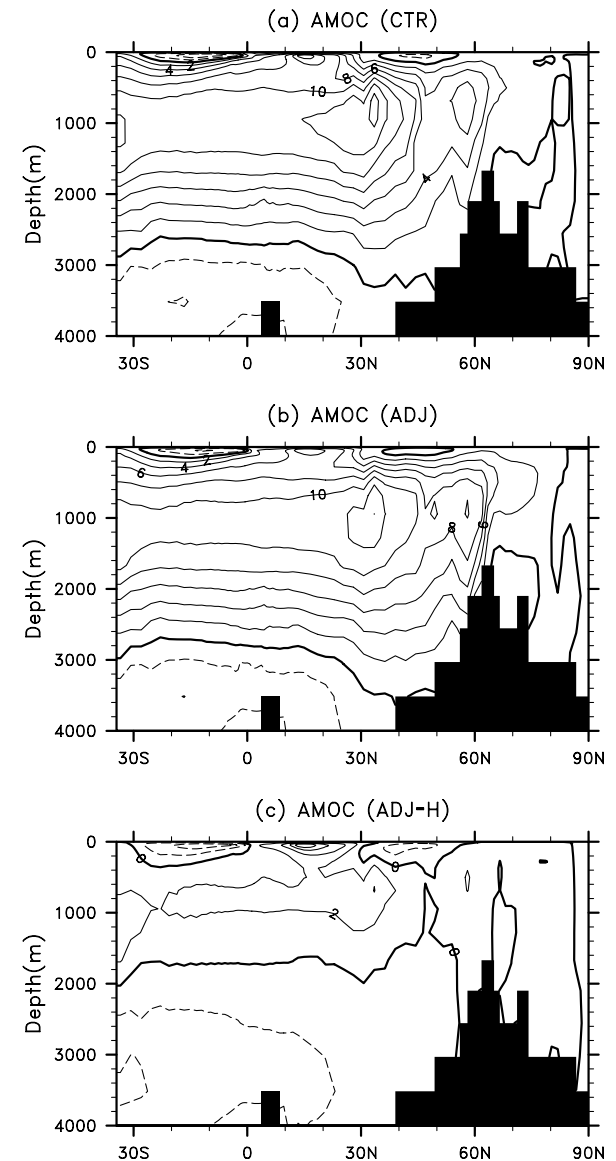
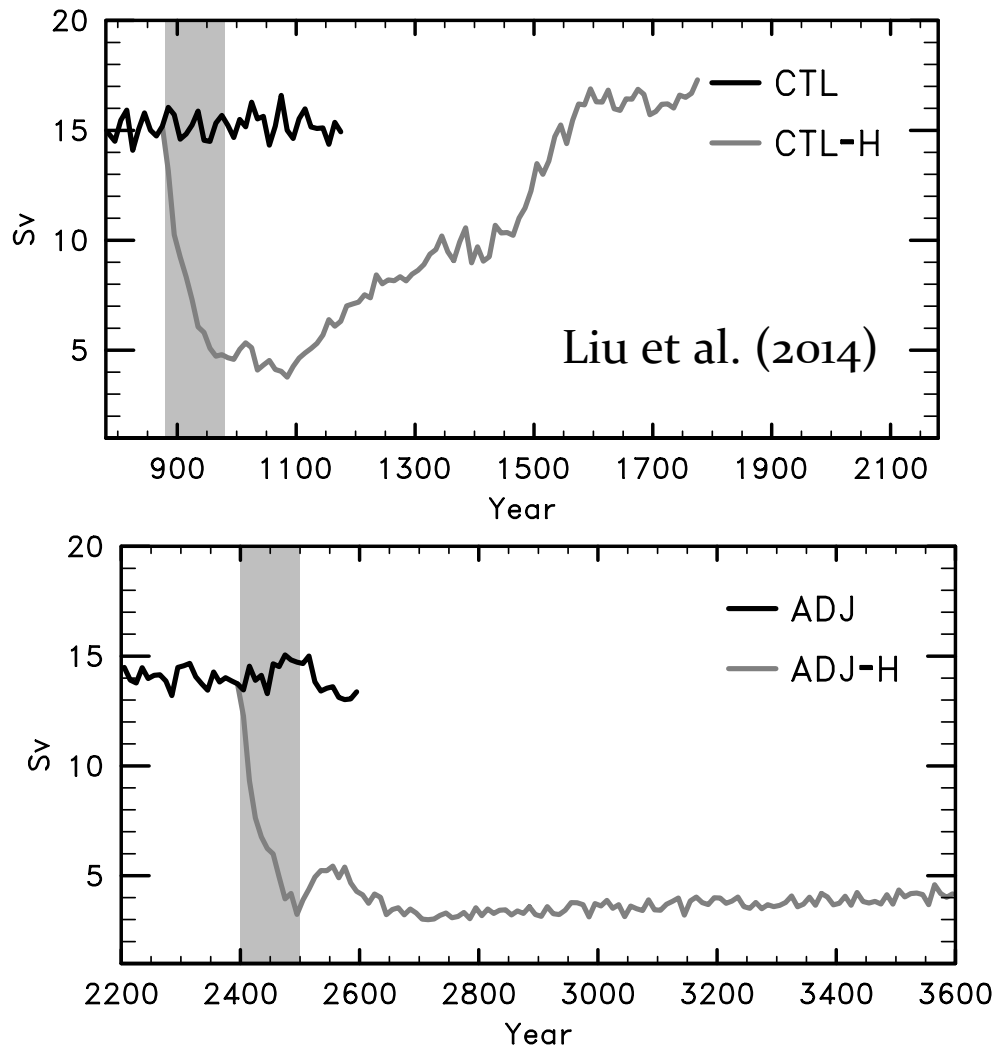


The flux adjustment corrects the salinity bias and tunes the model AMOC into a ME regime to be consistent with observations.



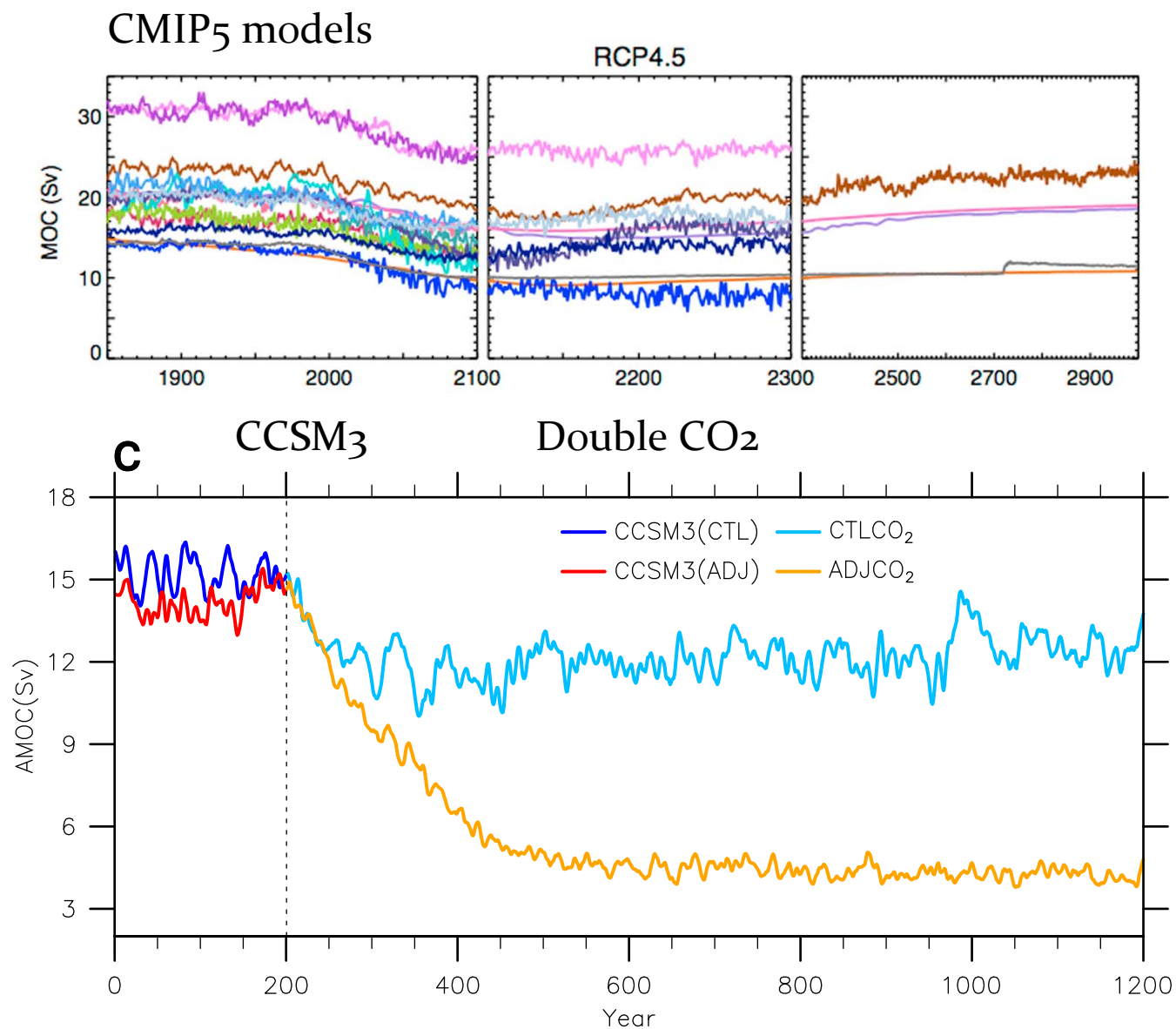
## b. The AMOC stability bias in climate models

Test: 1Sv, 100-yr pulse-like hosing experiment



How will this bias change future projection by climate models?

## c. Effects of the AMOC stability bias on future projection

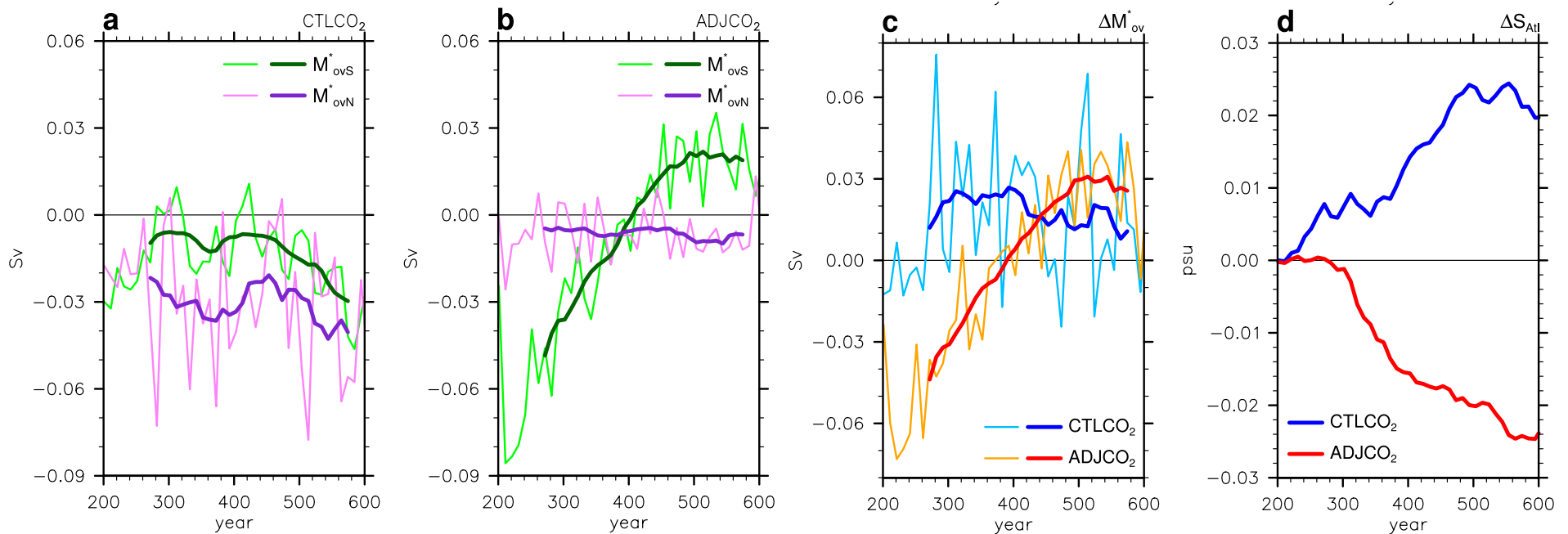


Liu et al. (2016)

## c. Effects of the AMOC stability bias on future projection

Different processes during years 250-500  
(50-250 years after CO<sub>2</sub> doubling)

Liu et al. (2016)

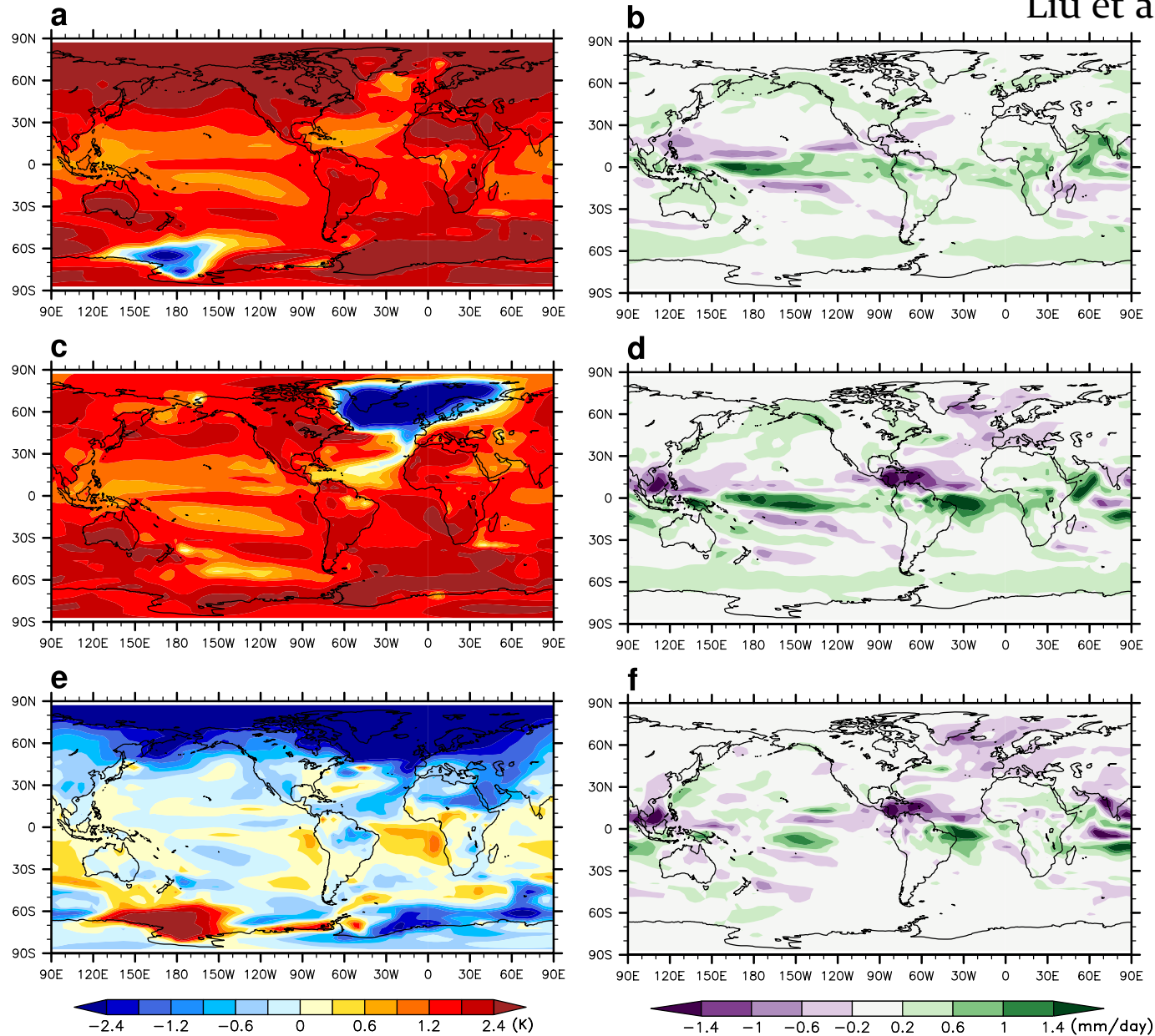


- The initial weakening of the AMOC in the ADJCO<sub>2</sub> (CTLCO<sub>2</sub>) causes a decline of freshwater divergence (convergence) in the Atlantic.
- This change freshens (salinifies) the Atlantic, inhibits (promotes) deep convection and deep-water formation, and finally leads to a collapse (partial recovery) of the AMOC.

## c. Effects of the AMOC stability bias on future projection

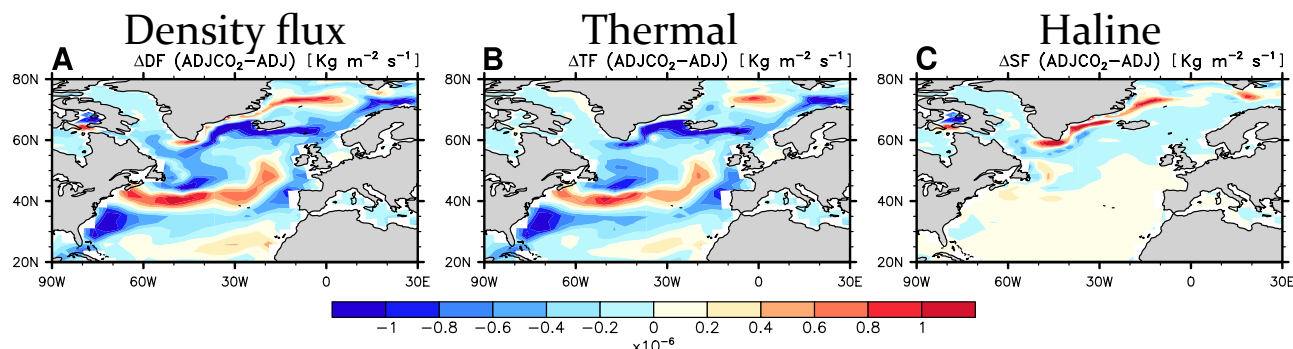
300 years after CO<sub>2</sub> doubling

Liu et al. (2016)

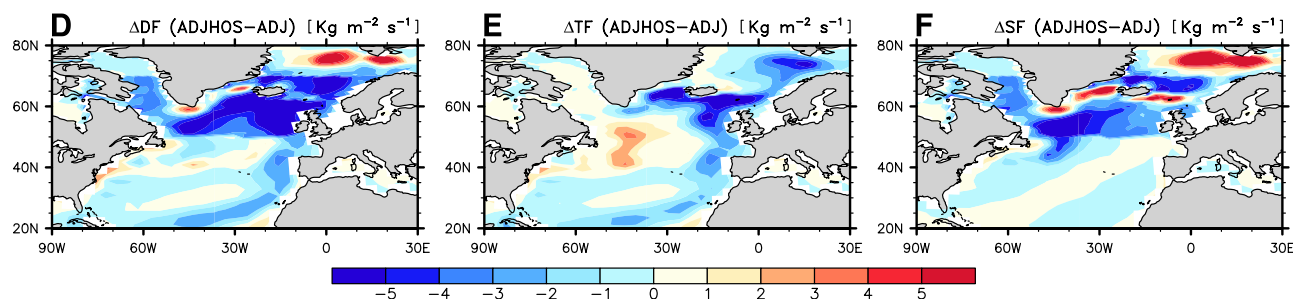


## c. Effects of the AMOC stability bias on future projection

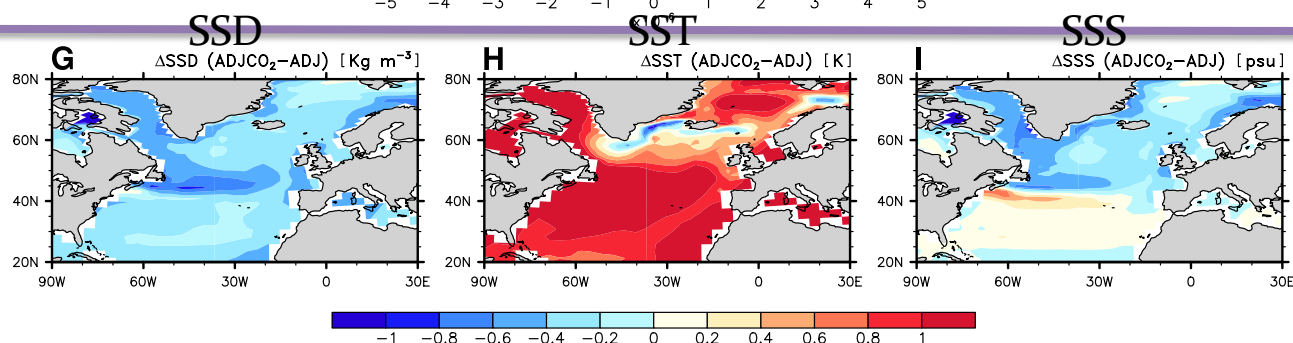
Double CO<sub>2</sub>



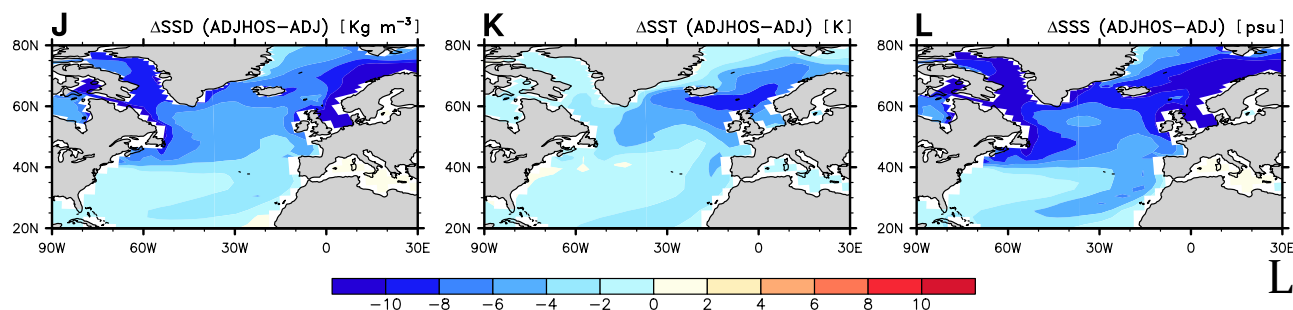
1Sv Hosing



Double CO<sub>2</sub>



1Sv Hosing



Liu et al. (2016)

# Conclusion and discussions

- A diagnostic indicator  $\Delta M_{ov}$  is proposed to monitor the AMOC stability. A negative (positive)  $\Delta M_{ov}$  indicates that the AMOC is in a ME (SE) regime.
- Observations suggest that modern AMOC is in a ME regime, whereas climate models simulate AMOCs in a SE regime.
- This AMOC stability bias is primarily related to a salinity bias in the upper ocean of the South Atlantic and can remarkably change future projection by climate models.
- How to solve this problem using physically improved model?
- Address the double ITCZ issue?
- High resolution model to resolve Agulhas leakage and improve the stratification at  $\sim 34^\circ\text{S}$ ?