Issues in identifying atmospheric and SST response to the **Oyashio Extension Front variability** R Hall¹, A Czaja^{1,} G. Danabasoglu², C Deser², C Frankignoul³, Y-O Kwon³¹ Imperial College, ²NCAR, ³ WHOI

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Regression of OEI against SST, DJFM, for ERA5 and



The OEI indices for September-March, 1981-2016 Derived from different SST datasets

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Ol mean February poleward SST gradient

The SST gradient is complex in the OE region

Conclusions

- Issues with front identification where SST gradients are complex
- Implications for interpretation of air-sea interaction results
- An ensemble approach can identify robust features
- GMPE is a possible way forward for deriving the indices:

February latitude of maximum SST gradient, 1982-2016 ERA5 46 44 latitude 6 38 36 153.25 153.75 155.25 155.75 156.75 157.25 157.75 158.75 160.75 166.25 169.75 54.25 154.75 156.25 159.25 159.75 160.25 161.75 162.25 164.25 165.75 167.75 164.75 167.25 169.25 158.2 161.2 162.7 163.2 168.2 163.7 165.2 166.7 168.7

Internal wave motions, precipitation variance, and spectral kinetic energy cascades in state-of-the-art high-resolution ocean and coupled ocean-atmosphere models Brian Arbic, University of Michigan, arbic@umich.edu

The impacts of small-scale motions on air/sea coupling can be studied with high-resolution models, satellite missions, and in-situ measurements. Here we use high-resolution models to examine three topics of interest for satellite missions and models focused on air-sea interaction and mesoscale processes.

1) Simulation of near-surface internal gravity wave KE

Internal wave kinetic energy (KE) will be a source of "noise" in satellite missions such as Odysea that will focus on near-surface mesoscale KE. Models can be used to characterize near-surface internal wave KE. But, how accurate are these models? And, are they impacted by coupling to the atmosphere?

Arbic et al. (2022; JGR-Oceans; <u>https://doi.org/10.1029/2022JC018551</u>) compared the near-surface KE in two state-of-the-art global high-resolution ocean-only simulations:

-HYbrid Coordinate Ocean Model (HYCOM) simulations run by the US Navy

–Massachusetts Institute of Technology general circulation model (MITgcm) simulations run by NASA

with observations from undrogued and drogued drifters, in low-frequency (< 0.5 cpd), near-inertial, semi-diurnal, and diurnal bands.

In a follow-on result (right), we have found that zonally averaged near-inertial KE in coupled ocean-atmosphere simulations with frequent coupling to the atmosphere lie much closer to drifter observations than ocean-only simulations with infrequent coupling. Thus an added benefit of high-resolution coupled ocean-atmosphere models, arising from the frequent coupling, is more accurate near-inertial motions.



2) High-frequency precipitation variance

Most precipitation variance is in sub-daily time scales, and low-resolution climate models do not capture this high-frequency activity (Covey et al., 2018).

Light et al. (2022; Climate Dynamics, <u>https://doi.org/10.1007/s00382-022-06257-6</u>) demonstrated that models with high-resolution in the atmosphere produce more high-frequency (subdaily) precipitation variance, as in observations, and that the ocean model resolution also matters, especially when the atmosphere model resolution is already high. The plot below shows high-frequency precipitation variance in the US Atlantic coast region. Higher-resolution models have greater high-frequency precipitation variance and ocean model resolution also matters.



3) Will the Odysea mission be able to measure spectral kinetic energy cascades?

Spectral KE fluxes $\Pi(K)$ measure the direction of spectral KE cascade as a function of wavenumber K. Spectral KE fluxes computed from altimetry (Scott et al. 2005; Scott and Arbic, 2007, Arbic et al. 2013, and others) apply the geostrophic assumption to sea surface heights. The proposed Odysea mission would enable the computation of spectral KE fluxes directly from velocities.

We computed spectral KE flux $\Pi(K)$ from 1731 hourly snapshots of coupled MITgcm ocean/GEOS atmosphere in Kuroshio region, 2020-01-19 to 2020-03-29.

Also computed from 90 snapshots of "Odysea samples" of this region during the same time period, with and without noise, prepared by Alexander Wineteer of NASA JPL.



Ocean small-scale fronts in the Northwestern Tropical Atlantic: Assessment from the EUREC4A-OA/ATOMIC field experiment

Solange Coadou, Sabrina Speich, Sebastiaan Swart, Chelle Gentemann, Dongxiao Zhang and Johannes Karstensen





Relate maximum velocity vertical shear

Detailed picture of the fronts at the meso and submesoscale,

revealing particularly strong gradients in the freshwater plume area

Increasing T compensation at the submesoscale, but strong salinity gradients remain uncompensated

Strong salinity fronts are associated with the rise of the maximum velocity vertical shear level, which indicates a shallower MLD

Fronts associated to the fresh water signal are coincident with a sharp increase of the maximum velocity vertical shear level, related to a shallow MLD

Upper ocean (6 - 100 m)



<u>Red shading:</u> velocity shear computed from the saildrone ADCP, <u>color dots:</u> Depth of the maximum velocity shear, <u>black</u> <u>line:</u> Salinity time series, <u>black dash line</u>: indicates eddy crossing, <u>blue line:</u> distance between the saildrone and the eddy center

Investigating long-term variability and trends of Atlantic Tropical Instability Waves with multidecadal satellite observations and mooring data





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Recommendations:

- Mesoscale TIWs are important components of the tropical heat and freshwater budgets and yet their net effect is not well understood
- Maintaining (or better expanding) in-situ observations of current velocities at established equatorial mooring sites (i.e., PIRATA buoys)
- Validation and comparison of (future) satellite missions or gridded surface current products with near-surface moored velocities



The Mesoscale Ocean Contribution to Sea Surface Temperature Distribution

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What are we doing?

• Developing a framework by which the contributions of different types of mesoscale dynamics to surface variability can be quantified, focusing on the mixed layer temperature budget (i.e. SST variability)

Temperature Budget

$$\frac{\partial T}{\partial t} = -\nabla \cdot \mathbf{u}T - \nabla_h^2 (A_h \nabla_h^2 T) + \frac{\partial}{\partial z} \kappa \left(\frac{\partial T}{\partial z} - \Gamma\right) + \frac{1}{\rho C_\rho} Q_{\text{net}}$$

Decompose advection term u (3D velocity) and T (temperature) using two-dimensional spectral filters, in 0.1° forced ocean CESM simulation





What is the effect of the mesoscale on temperature tendency?

These are for a 5-day mean (2009 Jan 1-5). Low-pass spatial filter also applied after tendency computation to smooth patterns



Temperature tendency – all advection

Work in progress/planned

- Extend this analysis to guantify seasonal/interannual/decadal timescale contributions of mesoscale ocean dynamics to SST
- Adapt budget and filtering code for use with Arakawa C-grid models (e.g., MITgcm)
- Thanks to NASA Physical Oceanography for support, and Ben Johnson and Frank Bryan for running the model simulation and assisting with obtaining output.

Feel free to contact me at andrewdelman@ucla.edu

What can synthetic aperture radar (SAR) filaments tell us about North Atlantic right whales and their prey?





An Observational Study of the Dependence of Ocean Surface Filaments on Wind Speed

Rick Danielson, Hui Shen, Jing Tao, Will Perrie, Fisheries and Oceans Canada

Please see https://arxiv.org/abs/2302.01533 or reach out rick.danielson@dfo-mpo.gc.ca

 $SAR \quad C = t + \epsilon + \epsilon_C$ $ERA \quad U = \alpha_U + \beta_U t + \epsilon + \epsilon_U$ $Var(C) = \sigma_t^2 + \sigma_\epsilon^2 + \sigma_C^2$ $Var(U) = \beta_U^2 \sigma_t^2 + \sigma_\epsilon^2 + \sigma_U^2$ $Cov(C, U) = \beta_U \sigma_t^2 + \sigma_\epsilon^2$

- 0.8 a) define specific – ERA5 c) scale SAR measurements wind contrast to of interest speed reduce-wind **b)** ask how these enhance-ocear measurements 6 ms⁻¹ dependence are associated
- Pearson correlation $\rho = Cov(C, U) / \sqrt{Var(C)Var(U)}$ with linear $(\beta_U \sigma_t^2)$ and nonlinear (σ_{ε}^2) components
- **Distance correlation** is a measure of nonlinear/nonmonotonic dependence (Székely et al. 2007, Székely and Rizzo 2009) that is comparable to Pearson correlation (Edelmann et al. 2021).



Investigation of the surface and subsurface salinity wakes of tropical cyclones



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0.2

0.3

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Sea Surface Salinity wakes (psu) of tropical cyclones (TC) using Argo floats



Using SAR to diagnose boundary layer stratification change at the ocean sub-to-mesoscale D.Vandemark, R. Foster, J. Stopa, C. Wang, B. Chapron, A. Mouche

Workshop question 4:

What are key analysis approaches and diagnostics of air-sea coupling which could guide modeling and observational strategies, including time and space scale dependencies and specific variables and processes?

New 'Routine' Observational Tool:

- Satellite SAR detects boundary layer coherent structure (CS) type
- CS relates to state of surface layer stratification (unstable-to-stable)
- Spatial scales of 1-1000 km including seasonal mapping capability

Possible Discussion question:

Could such stratification information provide a new useful diagnostic for model/data comparisons and improved flux prediction?

Implies model BL flux parameterization and predictions change with SST gradients, stability, and CS... including LES.



20 x 20 km Sentinel 1 SAR "imagettes"

Research supported by NASA Physical Oceanography grants: NNX17AH17G and 80NSSC20K0822

Example global multi-year mapping to illustrate SAR detected diversity in surface layer stratification (Ri) (Stopa et al. 2022)





CS information tied to Ri holds even when sub-20km variability exists (atm and ocean fronts)





Fig. S8 Variations due to SAR image homogeneity. (a) Image count of each WS/MC class separated for homogeneous and heterogeneous cases, i.e., whether or not the coherent structures coexist with atmospheric fronts, rains or cold pools. (b) comparison of -Ri between homogeneous and heterogeneous for each WS/MC class.

Vandemark slide 2

Using scatterometer observations to correct for persistent biases in modelled ocean surface winds



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- The ocean surface wind plays a key role at the atmosphere-ocean interface
- Persistent biases are found between scatterometer observations and ECMWF NWP model winds
- Lack of small-scale variability in ECMWF model winds
- Use temporally-averaged differences between scatterometer observations and collocated NWP model winds to correct for persistent local NWP wind vector biases:

$$SC(i, j, t_f) = \frac{1}{M} \sum_{m=1}^{M} u_{10s}^{scat}(i, j, t) - u_{10s}^{NWP}(i, j, t)$$

Metop-A ASCAT – ECMWF ERA5 [2018]



Application 1: Bias-corrected wind products for the Copernicus Marine Service

Products providing ECMWF surface wind and stress fields, corrected for persistent biases using scatterometer observations (hourly and monthly temporal resolution, 0.125° and 0.25° horizontal grid)



Application 2: Using machine learning to correct NWP surface wind biases

Machine learning setups are explored to find functional relationships between oceanic and atmospheric variables and the persistent NWP biases.

Predictions can be used to correct ocean forcing forecasts or NWP biases for periods before scatterometers existed.



Application 3: Bias-adjusted scatterometer observations for NWP data assimilation

The large systematic errors in NWP model winds inhibit effective assimilation of scatterometer winds. Using bias-adjusted scatterometer observations is expected to be beneficial for data assimilation.



Characteristics and Trends of the Campbell Plateau Meander in the Southern Ocean: 1993-2020 xinlong.liu@utas.edu.au

Xinlong Liu (IMAS); Dr. Amelie Meyer (IMAS); Dr. Chris Chapman (CSIRO)



Figure 1: The Campbell Plateau meander's mean positions over three different decades. Adapted from Liu et al. (in prep).

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al. in prep).

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Results & Implications

Xinlong Liu



Discussions

- Downstream Section significantly moving northward:
 - Hypotheses: 1. Changes in the stability properties of the downstream jet [Youngs et al. (2017); Barthel et al. (2022)];
 Interaction between the South Pacific gyre and Antarctic Circumpolar Current jets [Roemmich et al. 2007].
- Increasing width:
 - Consistent with Gille (1994) & Shao et al. (2015).
 - Hypothesis: Changes in the meander's volume transport.
- Increasing speed:
 - Consistent with recent Southern Ocean studies [e.g., Shi et al. (2021)].
 - Hypotheses: 1. Increased wind forcing; 2. Increased buoyancy forcing [Shi et al. (2021); Peng et al. [2022)].

Figure 2: Monthly time series and corresponding linear trends over the 1993-2020 period of the Campbell Plateau meander's position, width and speed. From Liu et al . (in prep).

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Regional Coupled System Development at the UK Met Office

PML Plymouth Marine



UK Environmental Prediction research



Ensembles

Climate:

al. 2021)

ERSEM



Met Office Some highlights











