Implications of the pattern effect for climate sensitivity and projections

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First Annual US CLIVAR Pattern Effect Workshop

Courtesy of NASA's Earth Observatory

The pattern effect and ECS



 $\Delta T = 0.75 \pm 0.2 \,^{\circ}\text{C}$

 $\Delta N = 0.65 \pm 0.27 \text{ Wm}^{-2}$

 $\Delta R = 2.3 \pm 1 \ \text{Wm}^{-2}$

(years 2000-2009 relative to 1860-1879)

Global-mean energy budget:

$$\Delta N = \lambda \Delta T + \Delta R$$

$$\lambda = -\frac{\Delta R - \Delta N}{\Delta T}$$

$$\mathrm{ECS} = -\frac{\Delta R_{2\times}}{\lambda}$$







 ΔT = 0.75 ± 0.2 °C

 ΔN = 0.65 ± 0.27 Wm⁻²

 ΔR = 2.3 ± 1 Wm⁻²

IPCC AR5 in 2013:

"Equilibrium climate sensitivity is likely in the range 1.5°C to 4.5°C (*high confidence*)... No best estimate for equilibrium climate sensitivity can now be given because of a lack of agreement on values across assessed lines of evidence and studies."

(Because other lines of evidence suggested ECS around 3°C)



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The view from 2021 (IPCC AR6)

p < 10% 6 6 Equilibrium climate sensitivity (°C) 5 5 Very likely: 2-5°C SAR AR5 Charney FAR AR4 TAR 4 Likely: 2.5-4°C 4 Likely: p > 66% AR6 3 3 Best estimate: 3°C 2 2 AR6 combines evidence from: Process understanding Instrumental record 1 Primarily model evidence Paleoclimates p < 5%Emergent constraints Also considers instrumental record and paleoclimates 1980 1990 2000 2010 2020 2030 Year of assessment

(b) Equilibrium climate sensitivity (°C) assessed in AR6 and simulated by CMIP6 ESMs



(a) Evolution of equilibrium climate sensitivity assessments from Charney to AR6

λ'

Observed sea-surface temperature trend over 1870-2019





2
1
0 SST trend (°C per century)
-1



Global radiative feedback in response to local warming patch in CAM4 (Dong et al. 2019)



λ'

Observed sea-surface temperature trend over 1870-2019







The radiative feedback over the historical record is different (more negative) than that for the equilibrium response to CO₂ owing to the pattern effect

Thus, historical warming provides only an estimate of the effective climate sensitivity (EffCS) which may differ from ECS

EffCS =
$$-\frac{\Delta R_{2\times}}{\lambda}$$

ECS = $-\frac{\Delta R_{2\times}}{\lambda + \lambda'}$



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Process understanding		ŀ	—					۰ı	
Instrumental record		I	-	ĺ.					
Paleoclimates		Į			-			1	
Emergent constraints		I			1				
Combined assessment		1	+						
CMIP6 ESMs	>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>>								
	1	1	1	1		1	1	1	1
0	1	2	3	4	5	6	7	8	9
Best estimate range or value							Likely range or limit		

- Can we produce better reconstructions of historical SSTs and sea ice (with uncertainty quantification), particularly for the 1800s reference period?
- Why do climate models generally fail to replicate observed patterns of warming (particularly since ~1980)?
- How confident are we in the models' radiative response to SST changes?
- Can we place observational constraints on the historical pattern effect?

• Fundamental issue: delayed warming (or cooling) has occurred preferentially in regions of most positive feedbacks, hiding potentially-high ECS from us; can we estimate what the radiative response (and thus ECS) will be to warming in these regions?

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SST trend (°C/century)

CMIP5 sea-surface temperature trend after CO₂ quadrupling (years 1-150) Pliocene sea-surface temperature anomaly (PRISM)

Pliocene SST reconstructions

Pliocene sea-surface temperature anomaly (Burls and Fedorov 2017)







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- Pliocene pattern isn't accurate; we need better methods to reconstruct SSTs and quantify their uncertainty (e.g., Jess Tierney's data assimilation work)

Does the pattern effect matter for transient warming?

Transient warming is highly correlated with ECS



Transient warming is highly correlated with ECS



(Sherwood et al. 2020)

Transient warming is highly correlated with ECS

CMIP5 1%/yr CO₂ ramping simulations



Transient warming is more highly correlated with EffCS



Transient warming and EffCS with freshwater forcing



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(Work of Yue Dong and Shaina Sadai)



Correlation between ECS and transient warming over recent decades has been proposed as a strong emergent constraint on ECS (e.g., Jiménezde-la Cuesta and Mauritsen 2019; Nijsse et al. 2020; Tokarska et al. 2020; Winton et al. 2020)

But shouldn't it be EffCS (rather than ECS) that controls transient warming??



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Transient warming and EffCS with freshwater forcing



(Work of Yue Dong and Andrew Pauling)



Outstanding questions

 How will the pattern of warming evolve in the future, and on what timescale? (depends on mechanisms driving observed patterns, which we don't currently know)

 Fundamental issue: multiple potential mechanisms project onto same pattern of SST response (ENSO/PDO dynamics), yet all have different future evolutions. As summarized by Tim Andrews, candidate mechanisms are:

- internal variability (originating in tropical Pacific and/or Southern Ocean?)
- non-CO₂ forcing (ozone depletion, Southern Ocean freshwater forcing, tropospheric or stratospheric aerosols?)
- role of teleconnections (from Southern Ocean or from Atlantic Ocean)
- response to CO₂ forcing (delayed E Pacific warming or nonlinear ENSO mechanisms)















By 2060, the spread of internal variability in the coupled model is ~1.3°C. This increases to ~2.0°C when considering SST pattern uncertainty

Thank you Maria and Cristi!



