



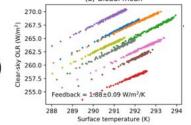
Historical period – the checkered past?

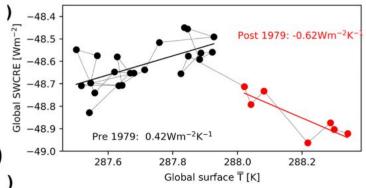


OLR = f([T]) -> functional relation between mean surface temperature and all components of the climate state

vector. Bold! (Surprisingly ok!)

For CO₂/H₂O gas phase ok (Zhang/Jeevanjee/Fueglistaler; GRL, 2021)





Clouds? With historical SSTs (Fueglistaler & Silvers, JGR, 2021)

Coupled ocean-atmosphere GCMs show monotonic time-varying "sensitivity" (-> ocean, heat uptake).

Q: Is what we see with historical period SSTs the same as in coupled GCMS?

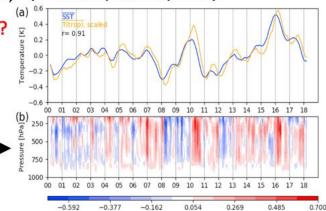
Patterns: How many?

Fueglistaler 2019: Case for 2 modes: Mean, and difference between

Tropical mean and SSTs in atmospheric deep convective regions.

Atm. T departure from moist adiabatic scaling with mean SST







1880

1900

1920

1980

2000

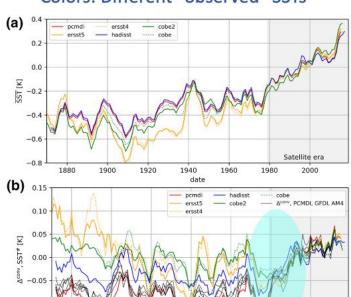
Historical period – the checkered past?

SST (K)



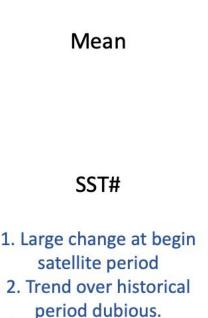
- 2 Modes: mean and "convective SST minus mean" (SST#; or: variance, width of PDF).
- -> Very tropic-centric view of the world, explains AMIP historical global cloud radiative effect surprisingly well.
- -> Reduced dimensionality to 2 parameters -> timeseries. "Peculiar Trajectory" (Fueglistaler & Silvers, 2021)

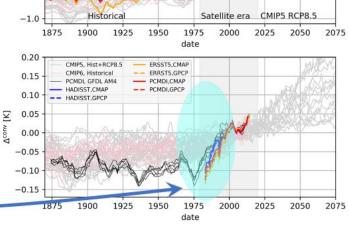
Colors: Different "observed" SSTs



Strong: Obs./AMIP <-> faded: Coupled GCMs

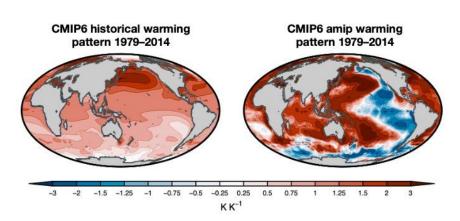
CMIP6, Historical





1) SST → Radiation

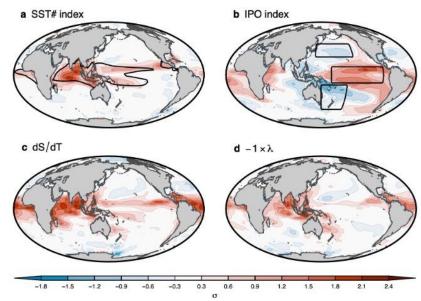
What modes of SST variability are relevant to the pattern effect?



The warm pool matters. The IPO is mostly radiatively neutral.

See also poster by Erik Schaffernicht

Regression onto CMIP6 historical warming patterns



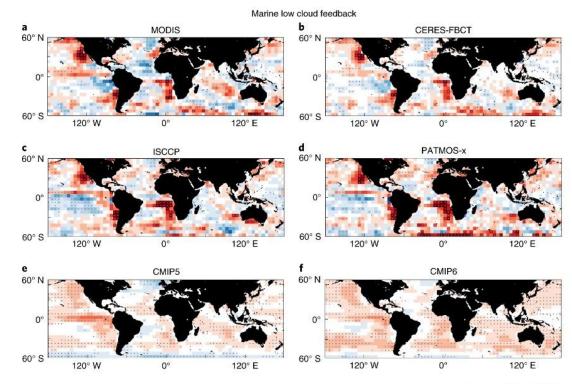
2) Radiation → SST

What are the implications of incorrect low-cloud feedbacks for coupled climate variability?

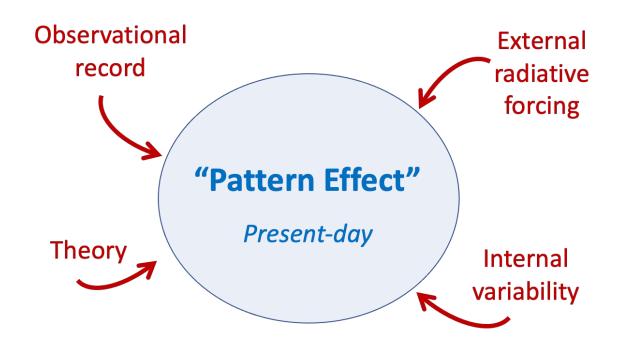
AOGCMs underestimate positive low cloud feedback in stratocumulus regions.

Likely important for coupled variability and teleconnection patterns.

Also for the forced response?

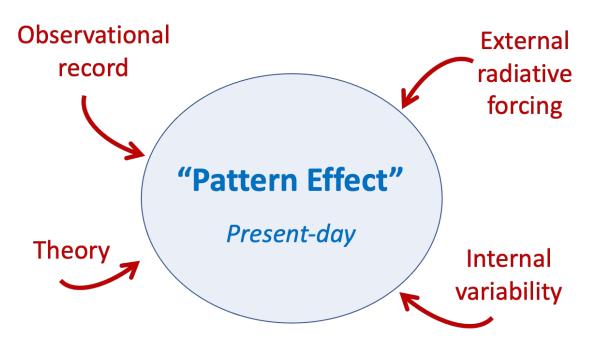


Myers et al. 2021



Short, sparse, inhomogenous.

Discrepancies remain even after 1979. Need for proxy records to expand our time horizon.

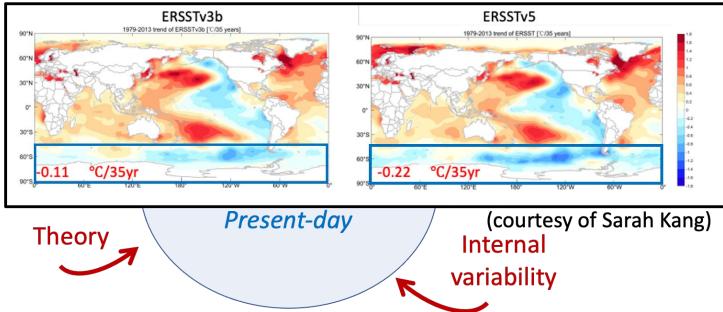


SST Trends: 1979-2013

Short, sparse, inhomogenous.

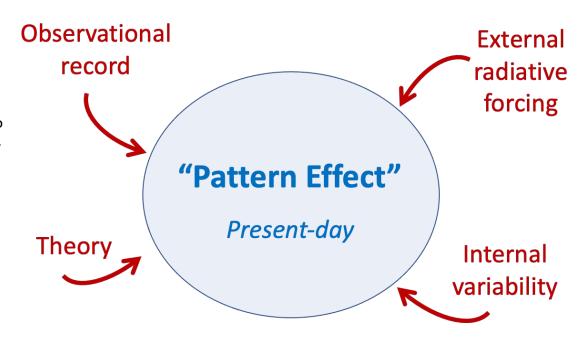
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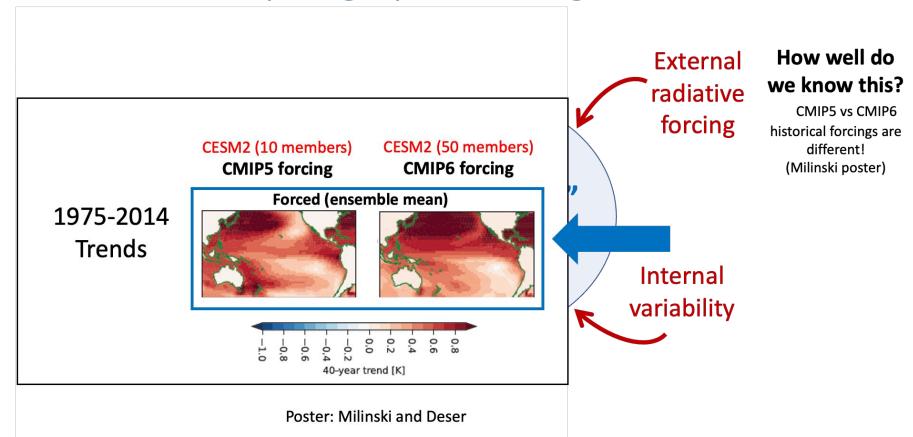
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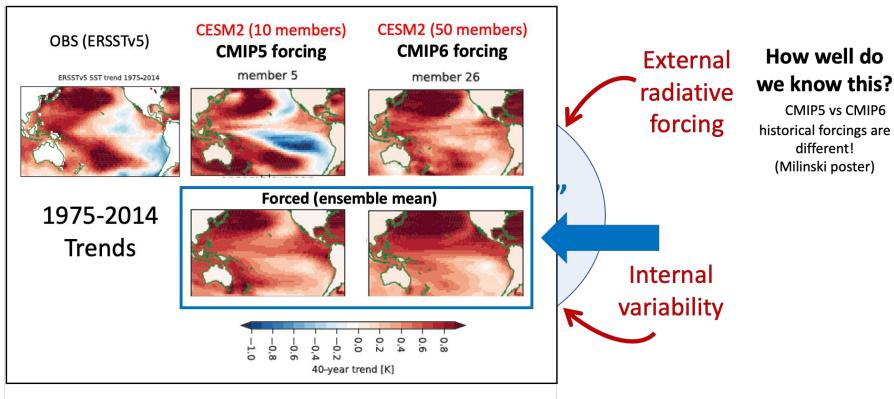
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How well do we know this?

CMIP5 vs CMIP6 historical forcings are different! (Milinski poster)

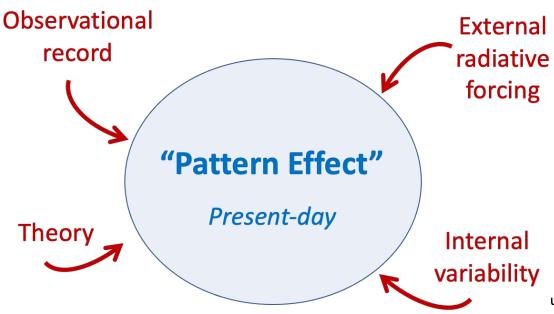




Poster: Milinski and Deser

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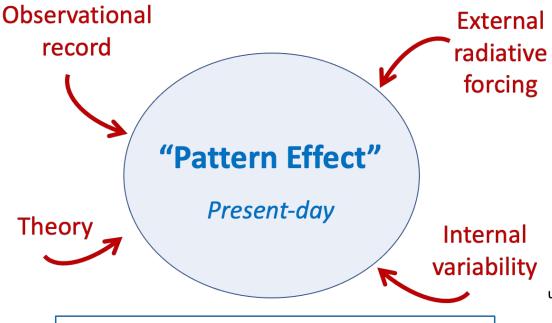
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Need for model Large Ensembles.

But do models underestimate magnitudes, and are they missing / misrepresenting key physical processes (EUC, Antarctic meltwater, aerosol-cloud interactions)?

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Discrepancies remain even after 1979. Need for proxy records to expand our time horizon.



Need for empirical approaches to separating forced and internal variability using observations alone; "Observational" Large Ensembles.

(tested within a model LE framework; eg Wills et al. 2020).

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Short, sparse, inhomogenous.

Discrepancies remain even after 1979. Need for proxy records to expand our time horizon. Observational External record radiative forcing "Pattern Effect" **Present-day** Theory Internal variability

Need for better understanding of the underlying physics.

Transient vs. equilibrium responses.

Need for empirical approaches to separating forced and internal variability using observations alone; "Observational" Large Ensembles.

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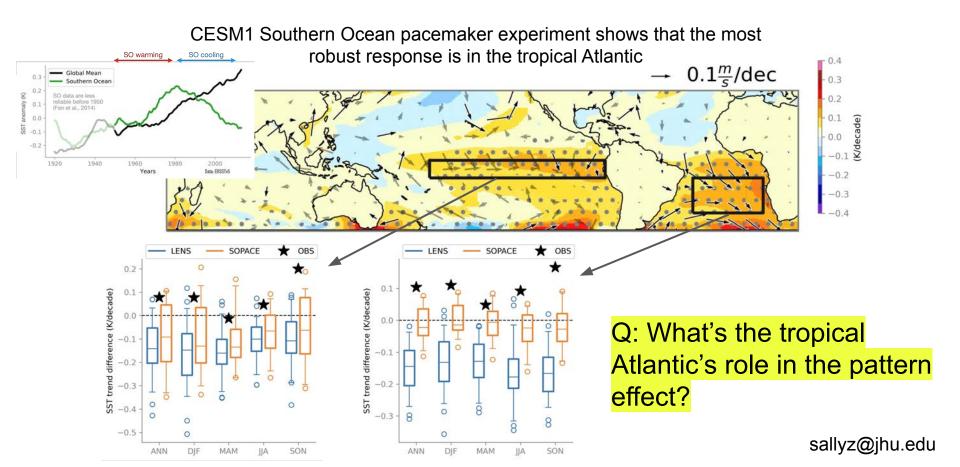
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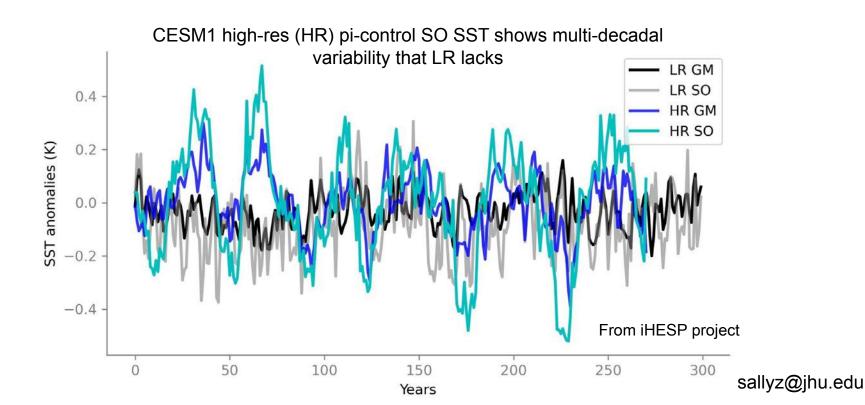
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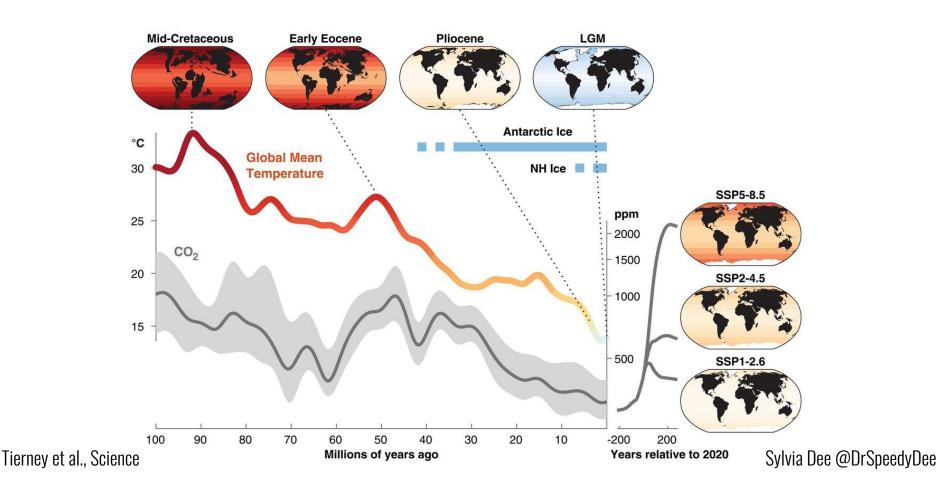
Q: How do "observed" Southern Ocean SST trends affect the tropical SST?



Q: How is the pattern effect affected by model resolution (or by processes that are missing in current GCMs)?

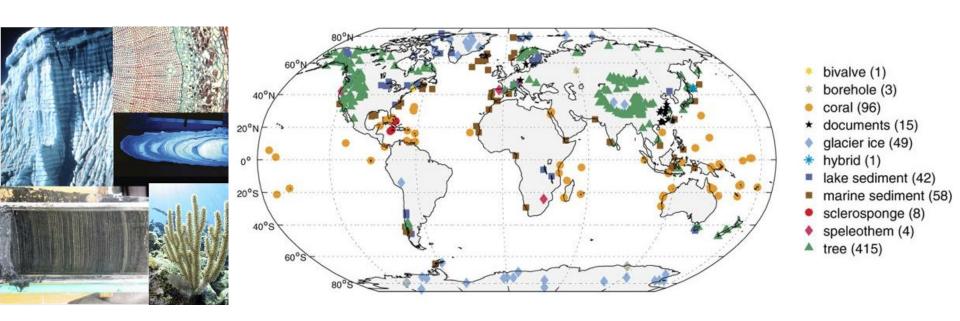


Long-Term and Paleoclimate Perspectives



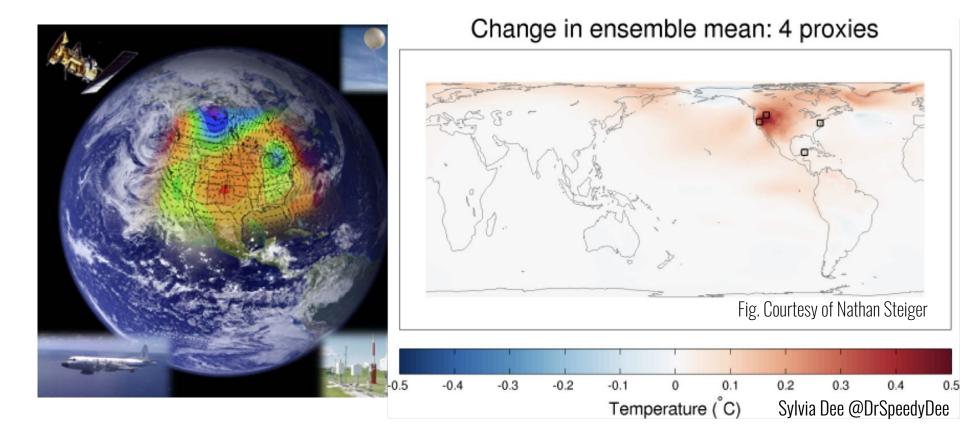
CONSTRAINTS ON THE PATTERN EFFECT:

Individual and Networks of Paleoclimate Reconstructions



CONSTRAINTS ON THE PATTERN EFFECT:

Paleoclimate Data Assimilation Reconstructions

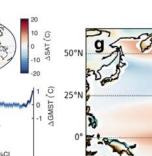


CONSTRAINTS ON THE PATTERN EFFECT:

Paleoclimate (Data Assimilation) Reconstructions

New Paleoclimate Data Assimilation products provide a larger sample size LMR: Last Millennium Reanalysis (Hakim et al., 2016), PHYDA: Paleo Hydrodynamics Data Assimilation product (Steiger et al., 2018)

COLD (LGM) 🐽



120°E

MEDIUM (LM) 😎



A new view of the Eocene greenhouse world from paleoclimate data assimilation

Show affiliations

Tierney, J. E.; Zhu, J.; King, J.; Li, M.; Malevich, S. B.; Poulsen, C. J.; Ridgwell, A.; Hakim, G. J.; Tardif, R.; Kump, L. R.

The Eocene, as the most recent greenhouse world, offers valuable lessons about how the climate system behaves under high levels of CO₂ (ca. 1000 ppm). However, our understanding of Eocene climate dynamics is fundamentally limited by uncertain, and spatially and temporally restricted, paleoclimate data. Conversely, climate model

(Sea Surface) Temperature
Precipitation, Geopotential Height
Palmer Drought-Severity Index Luo & Dee et al., 2022

140°W

170°E

Osman et al., 2021

18 ka - 2k

CONSTRAINTS ON THE PATTERN EFFECT: Paleoclimate

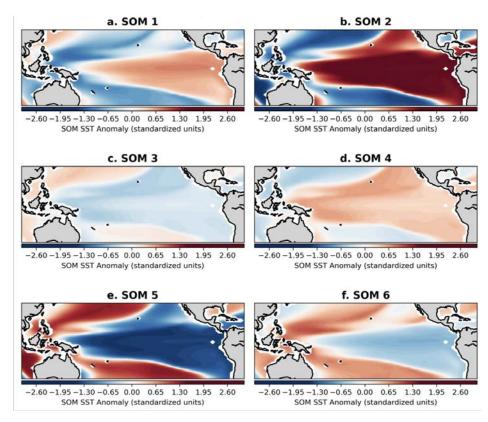


Ulla's talk: "How can we address the model differences?" We can use paleoclimate data.

DISCUSSION Q1: What are the best methods for extracting information on the pattern effect from the paleoclimate record?

Issues: sparse data in space and time; time scale of DA products; model priors in DA

DISCUSSION Q2: Wish list for the paleoclimate modeling/data-model comparison community?



Luo AGU 2021 talk, Luo et al., in prep

Sylvia Dee @DrSpeedyDee

DISCUSSION: THE PATTERN EFFECT IN THE HISTORICAL RECORD