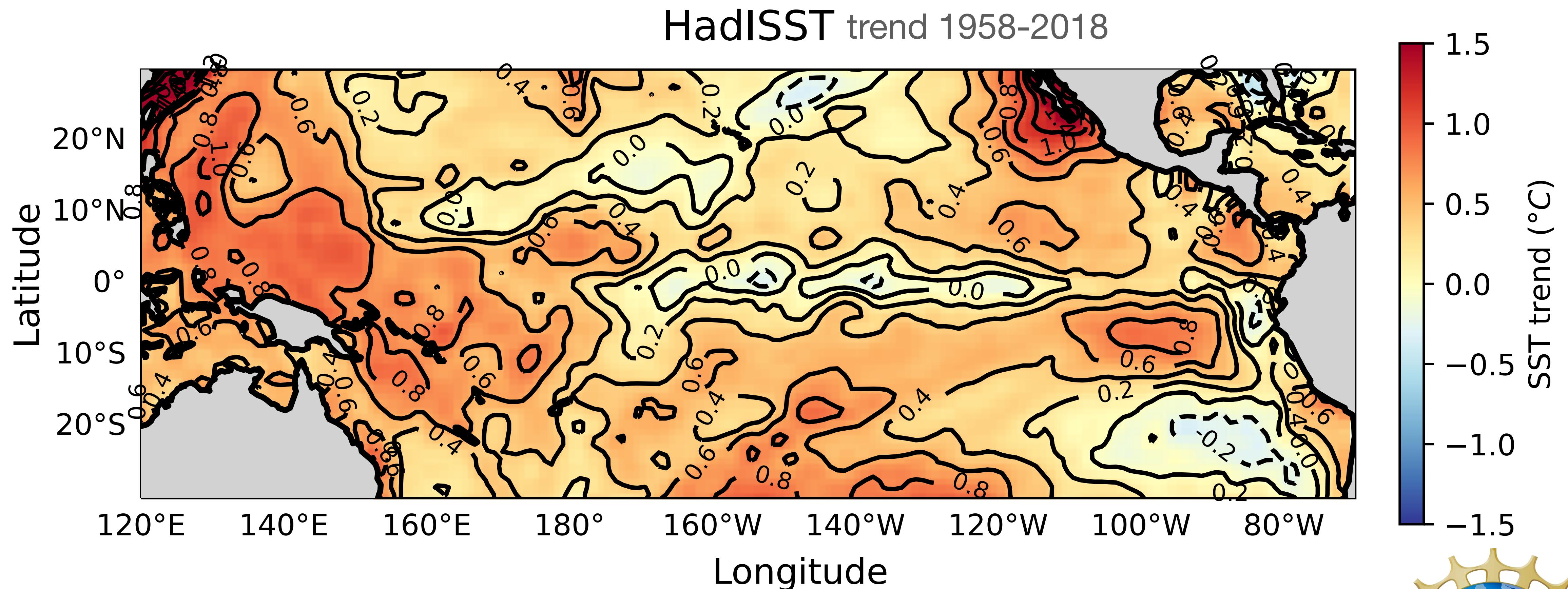


Tropical Pacific zonal SST gradient trends: biases and mechanisms



Richard Seager, Naomi Henderson and Mark Cane
Lamont Doherty Earth Observatory of Columbia University, Palisades, NY
supported primarily by NSF Physical Oceanography program

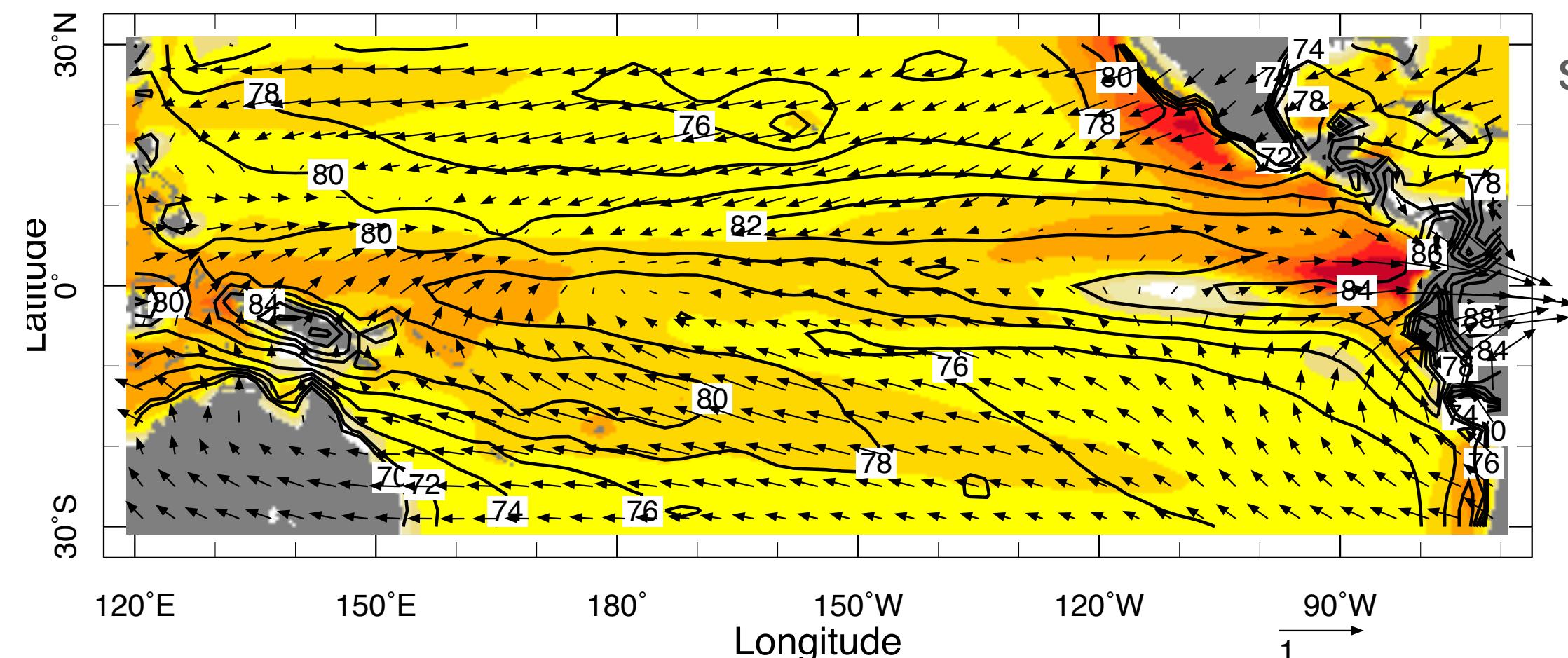
US CLIVAR, Seattle, August 2023



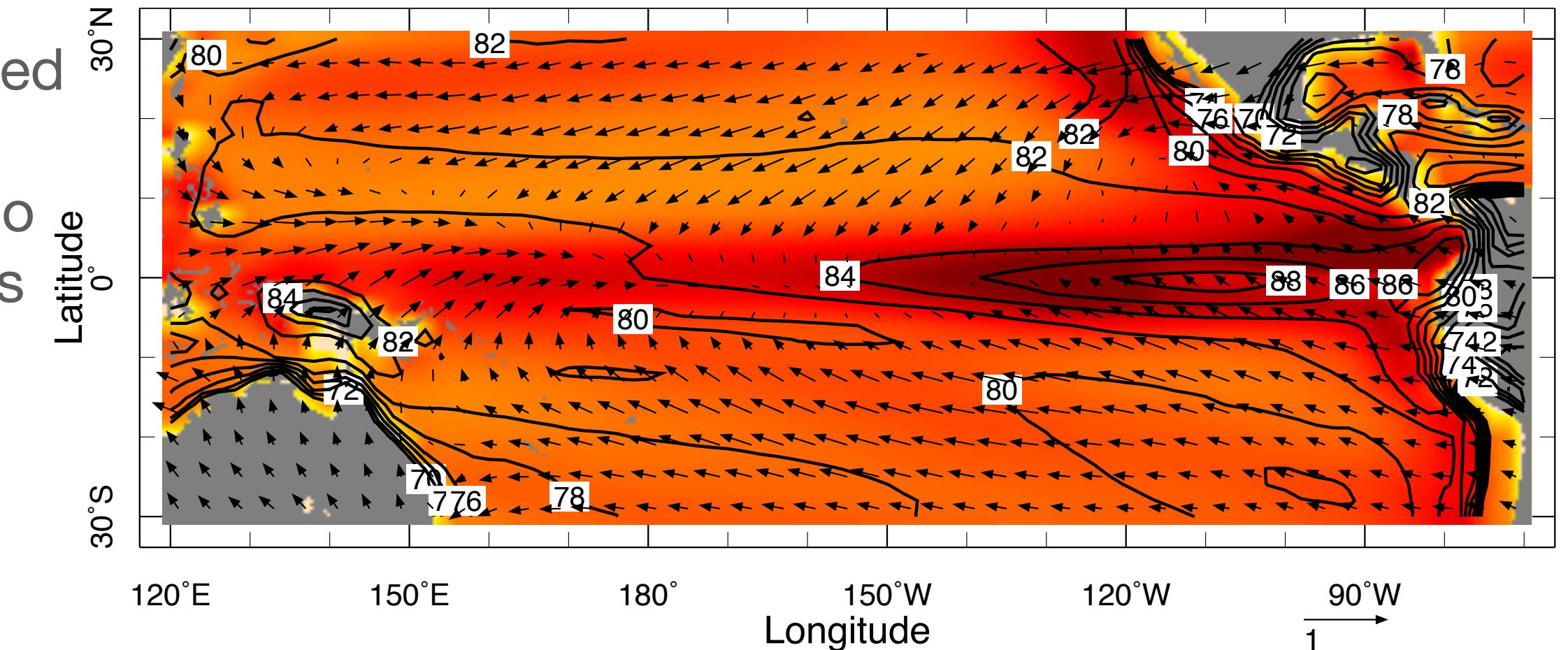
Seager et al. (2019) argued the observed trend was the forced response to rising GHGs and CMIP5 models get this wrong due to tropical Pacific mean state biases

SST trend(colors), 10m wind trend(vectors) and climatological relative humidity(contours)

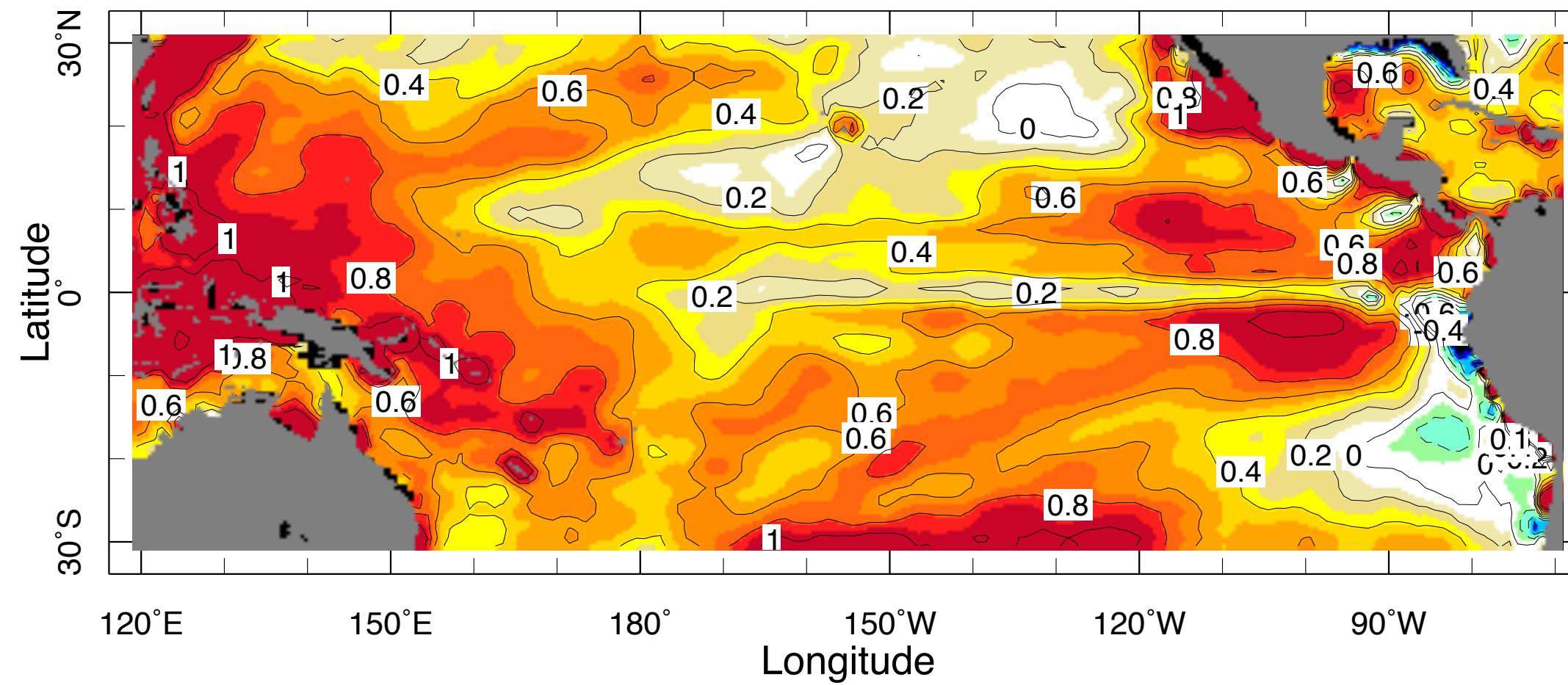
a) CM-ECMWF world



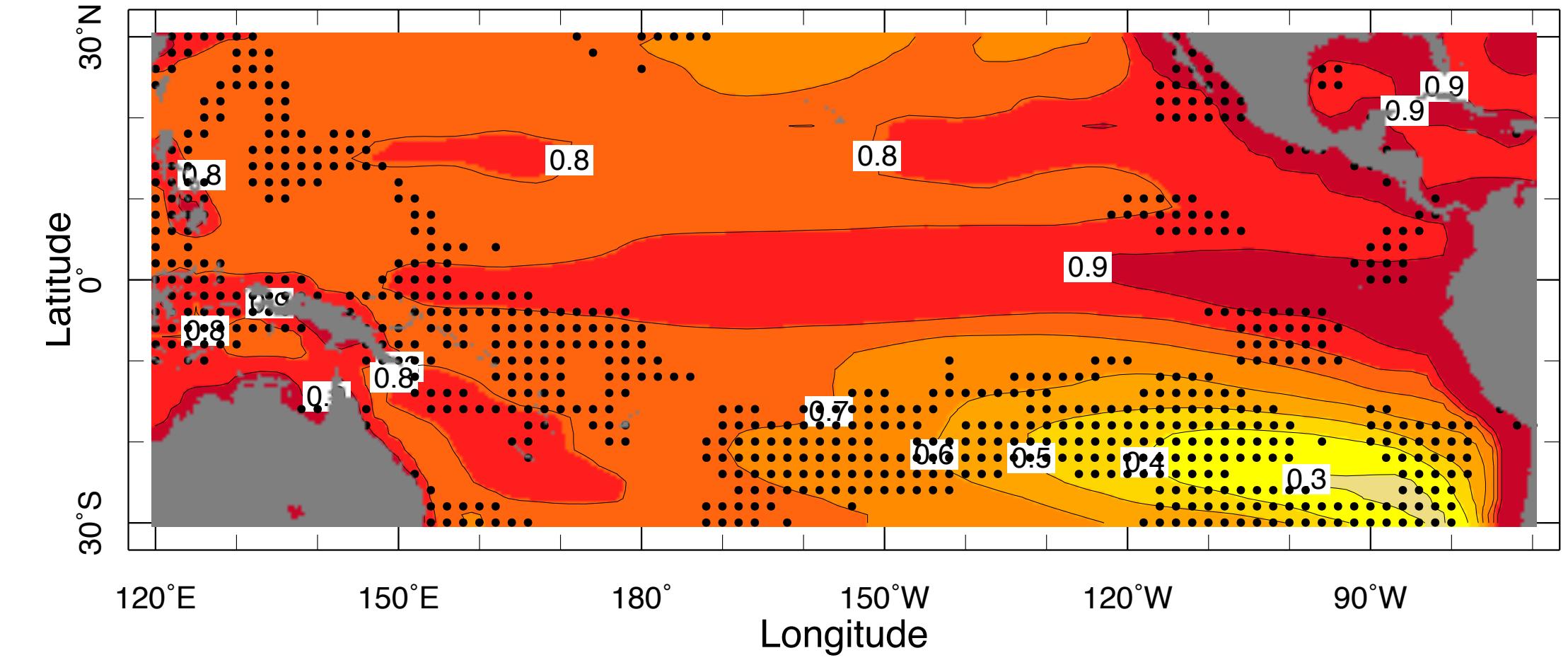
d) CM-CMIP5 world



a) ECMWF/ORAs4



c) CMIP5: historical + RCP8.5



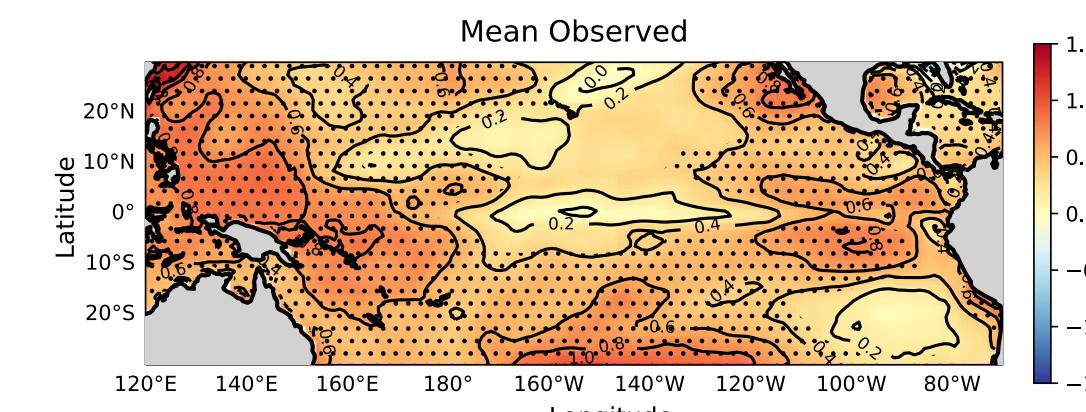
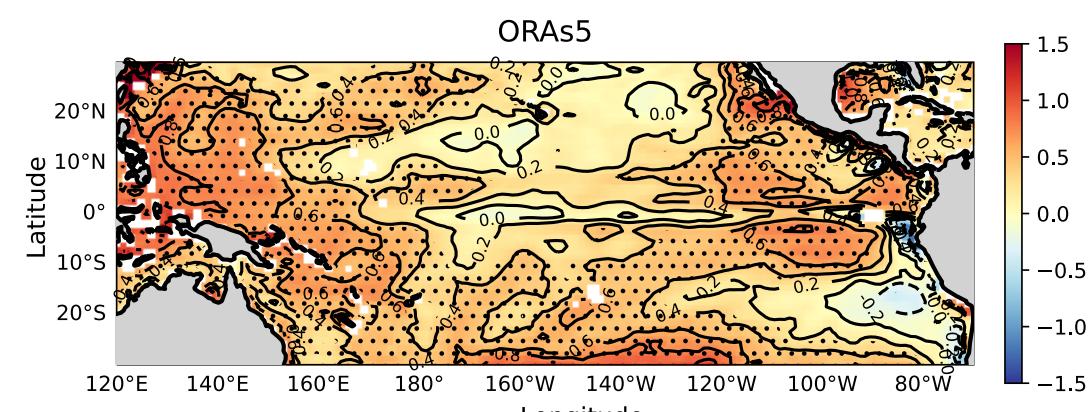
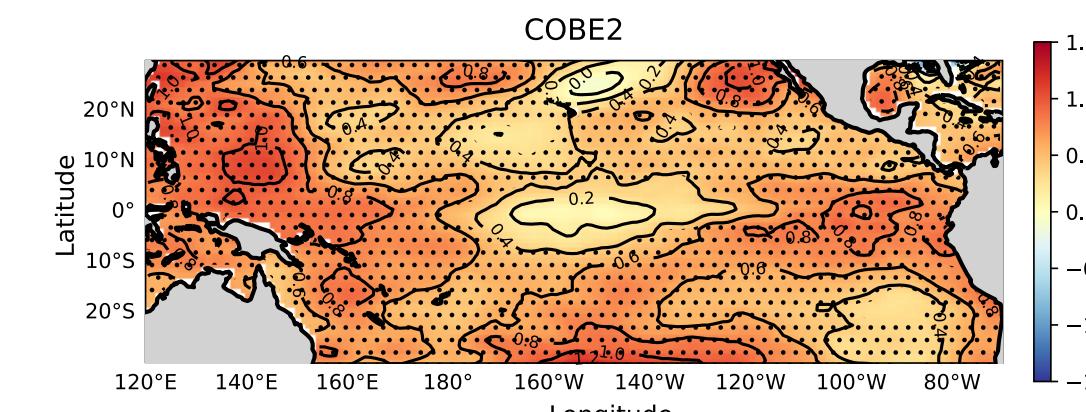
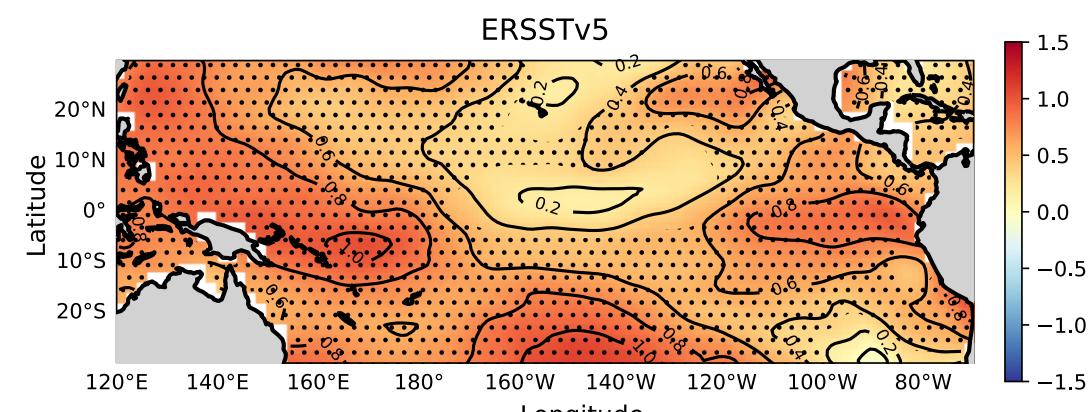
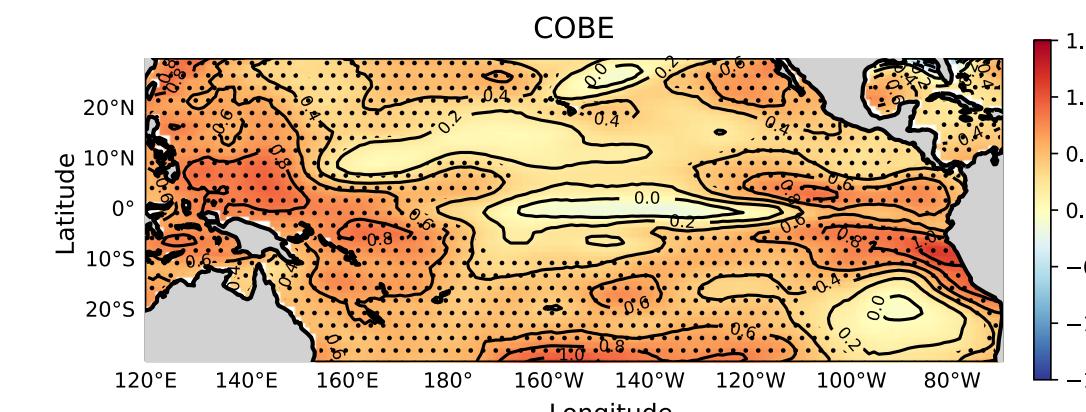
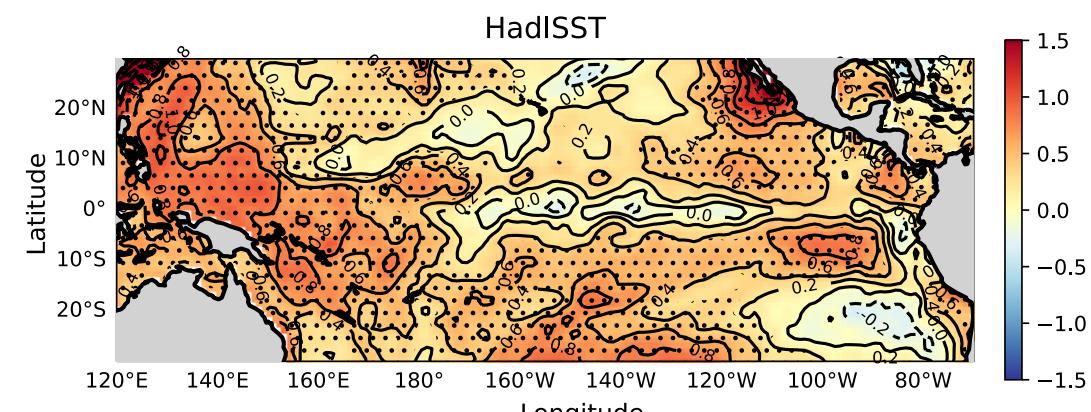
Watanabe et al. (2020) and Olonscheck et al. (2020) argued instead that the observed trend is unlikely but due to natural variability

simple coupled
model
responses to
rising GHGs

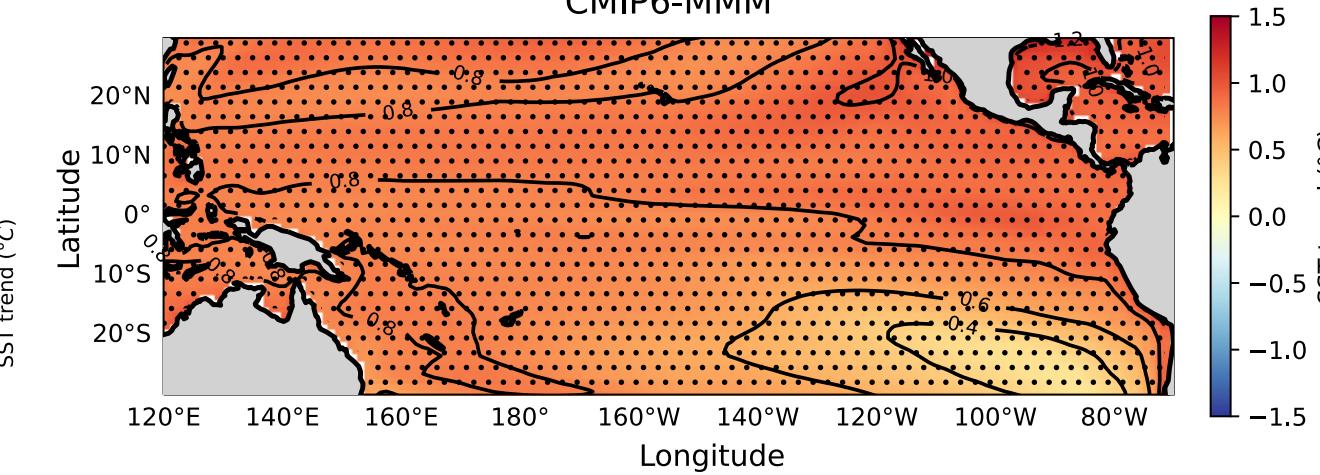
Observed and CMIP6 and Large Ensemble SST trends over past 60 years

CMIP6 and LENS SST Trends, 1958-2018

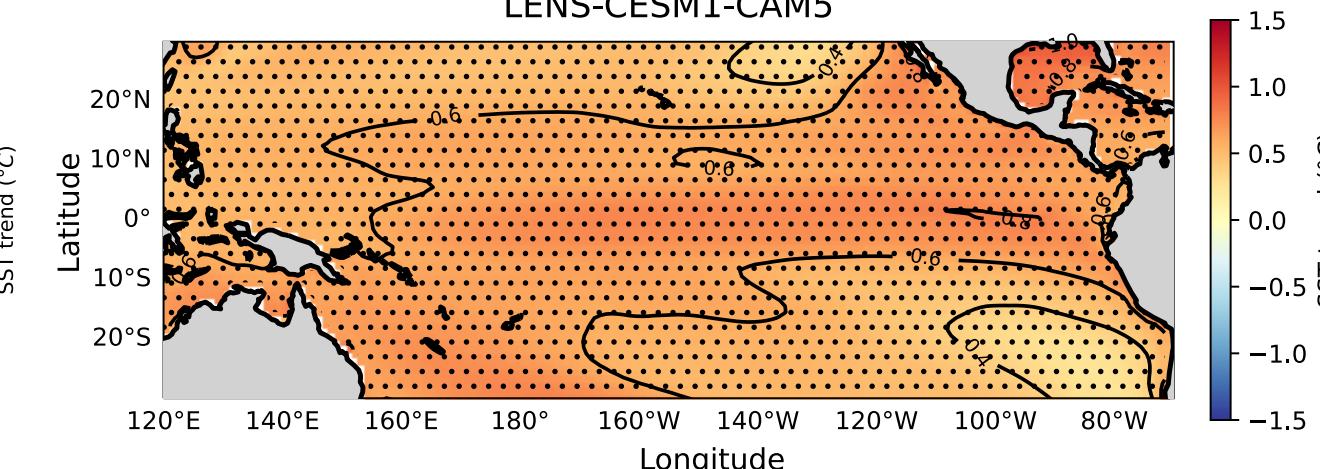
Observed SST Trends, 1958-2018



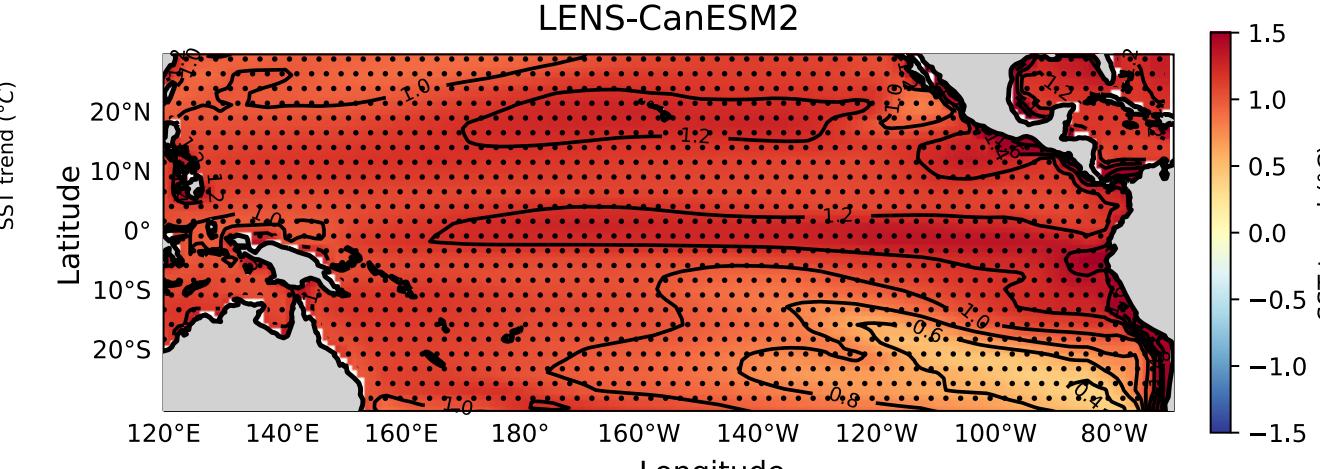
CMIP6-MMM



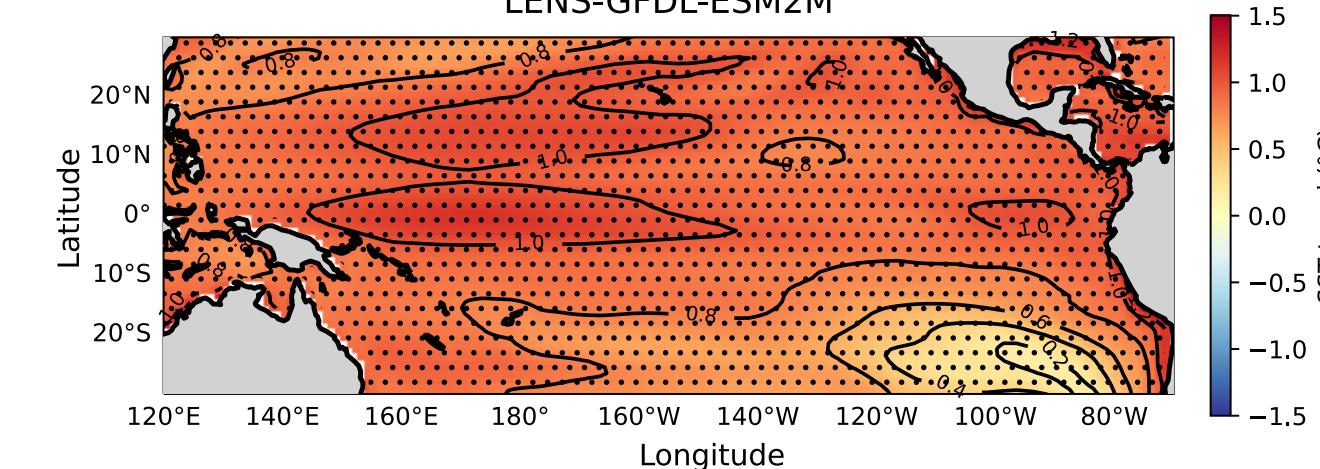
LENS-CESM1-CAM5



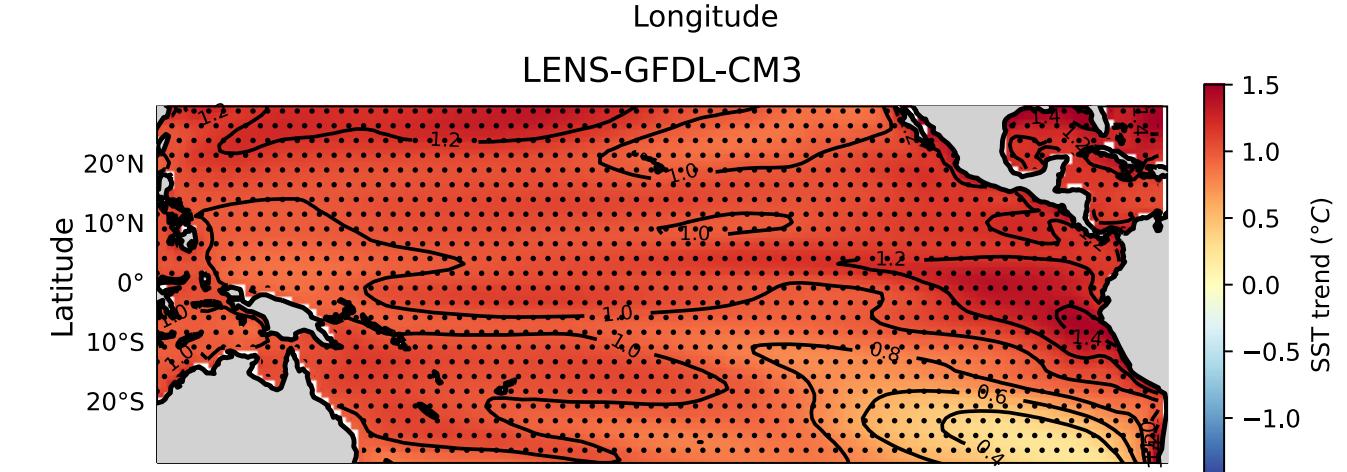
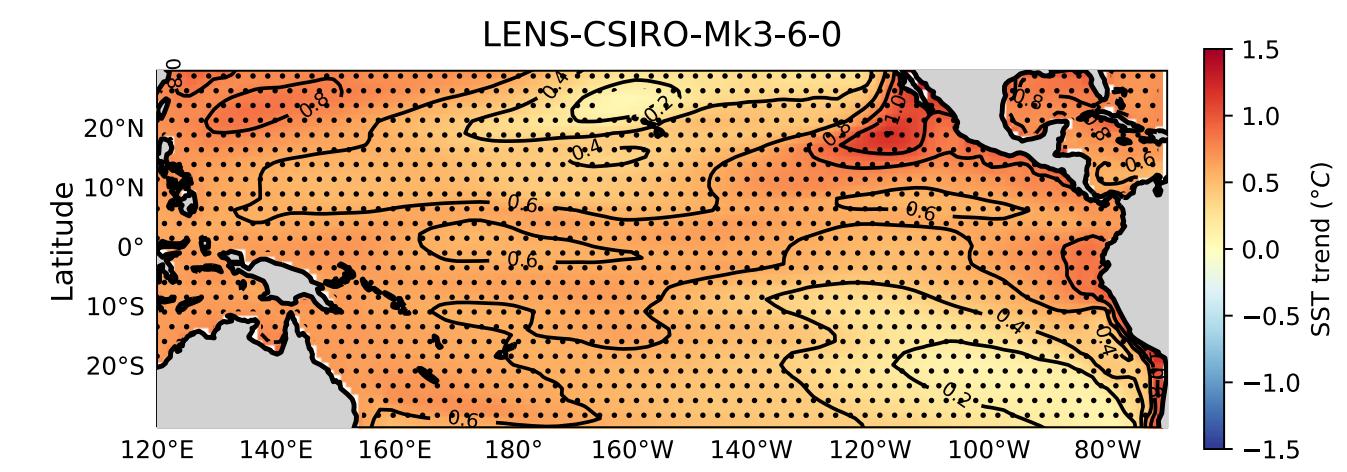
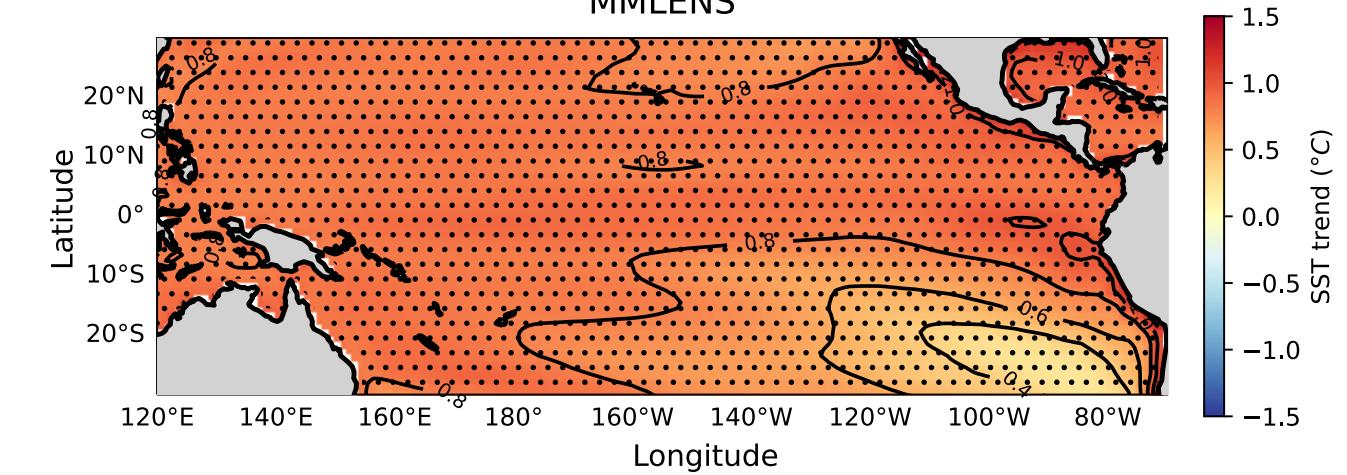
LENS-CanESM2



LENS-GFDL-ESM2M



LENS-GFDL-CM3



In 5 SST datasets and their average there is a lack of cold tongue warming. This has been seen back to Cane et al. (1997) for the 1900-1993 period.

In contrast, CMIP6 and Large Ensembles have cold tongue warming

Observed, CMIP6 and LENS SST Trend Indices, 1958-2018

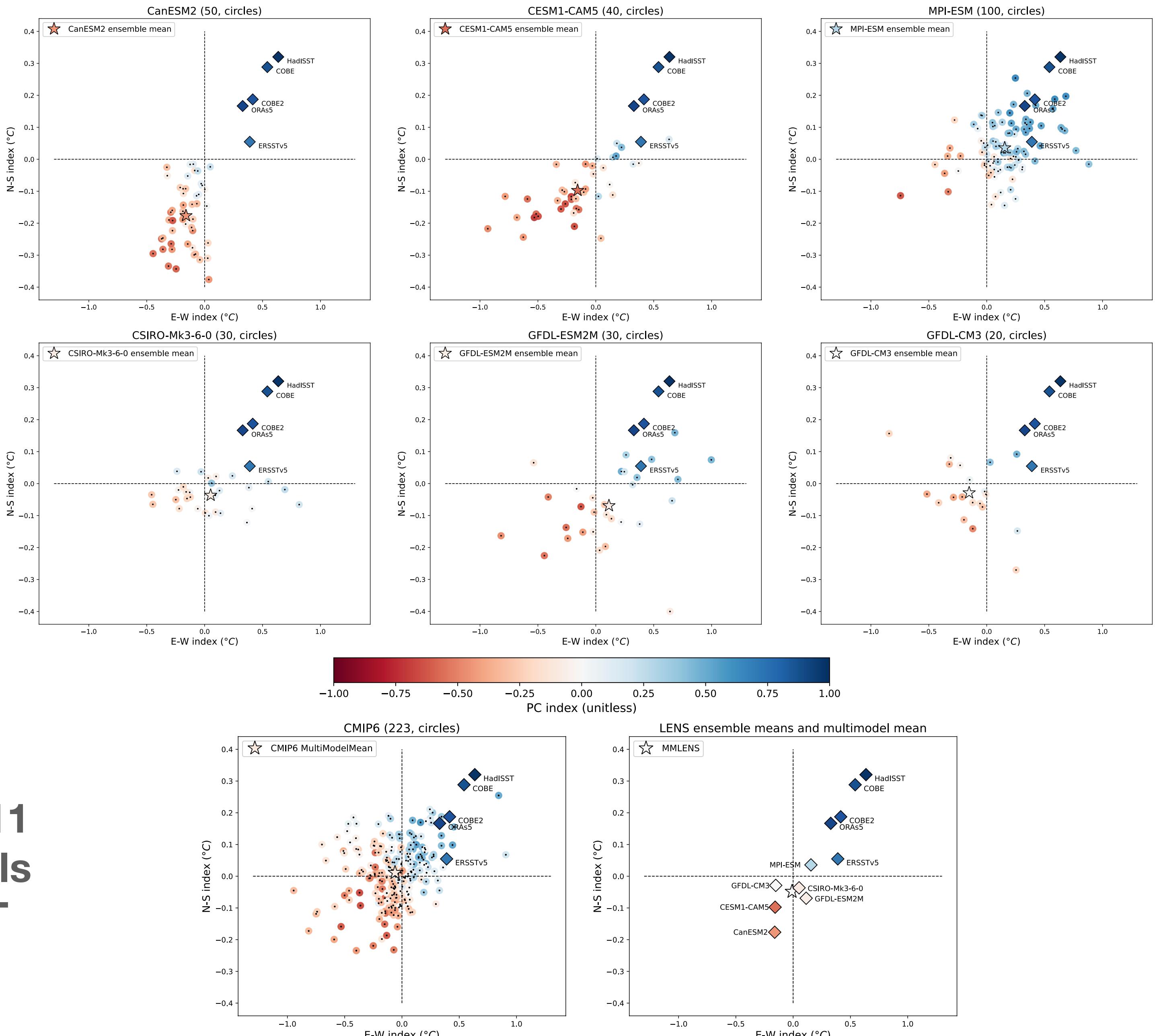
The observed SST trend is characterized by:

An enhanced west-east SST gradient (x-axis)

An enhanced cooling of the cold tongue w.r.t. waters immediately to the north and south (y-axis)

Also measure model trends with a pattern correlation to HadISST observed (color)

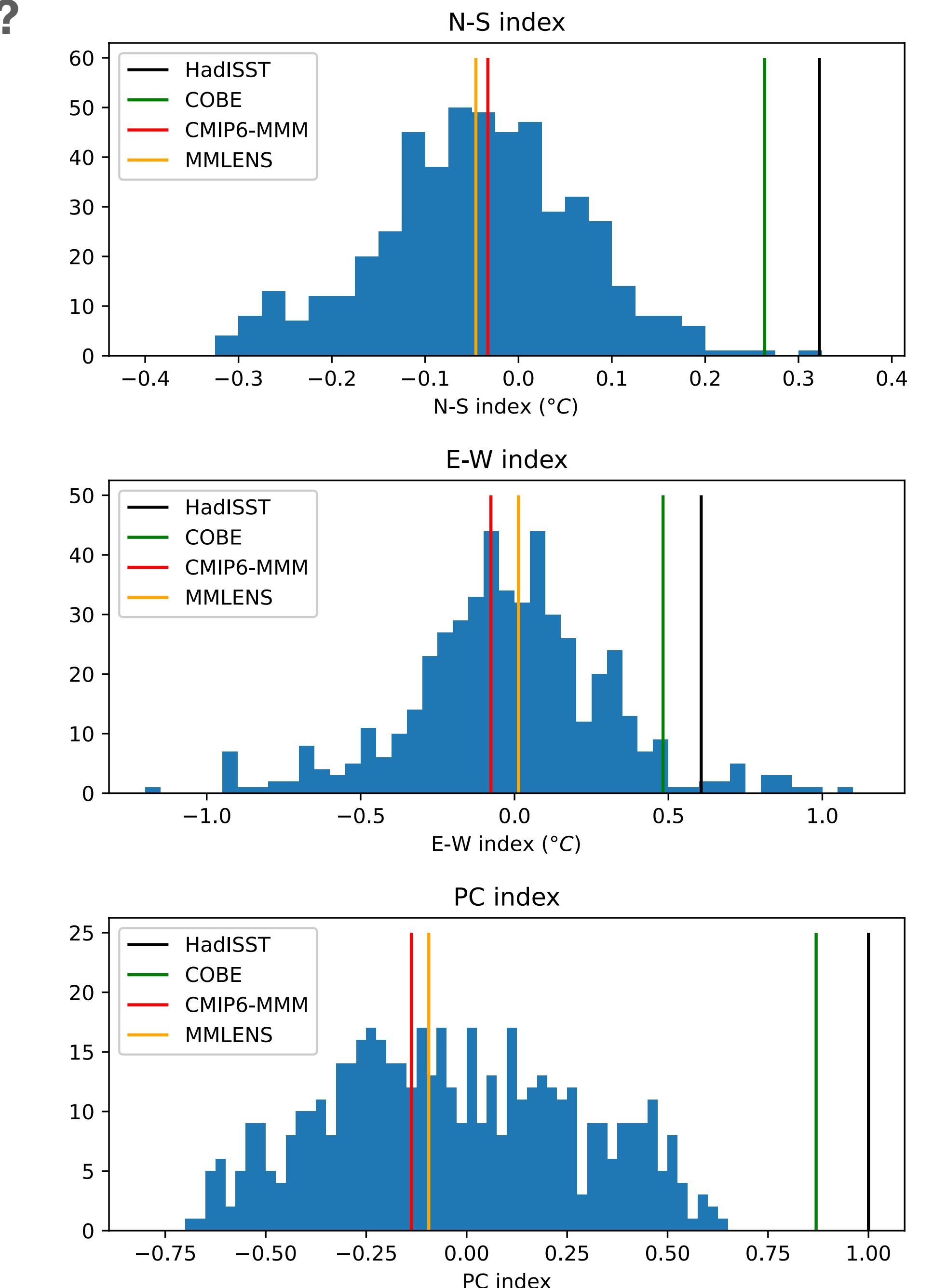
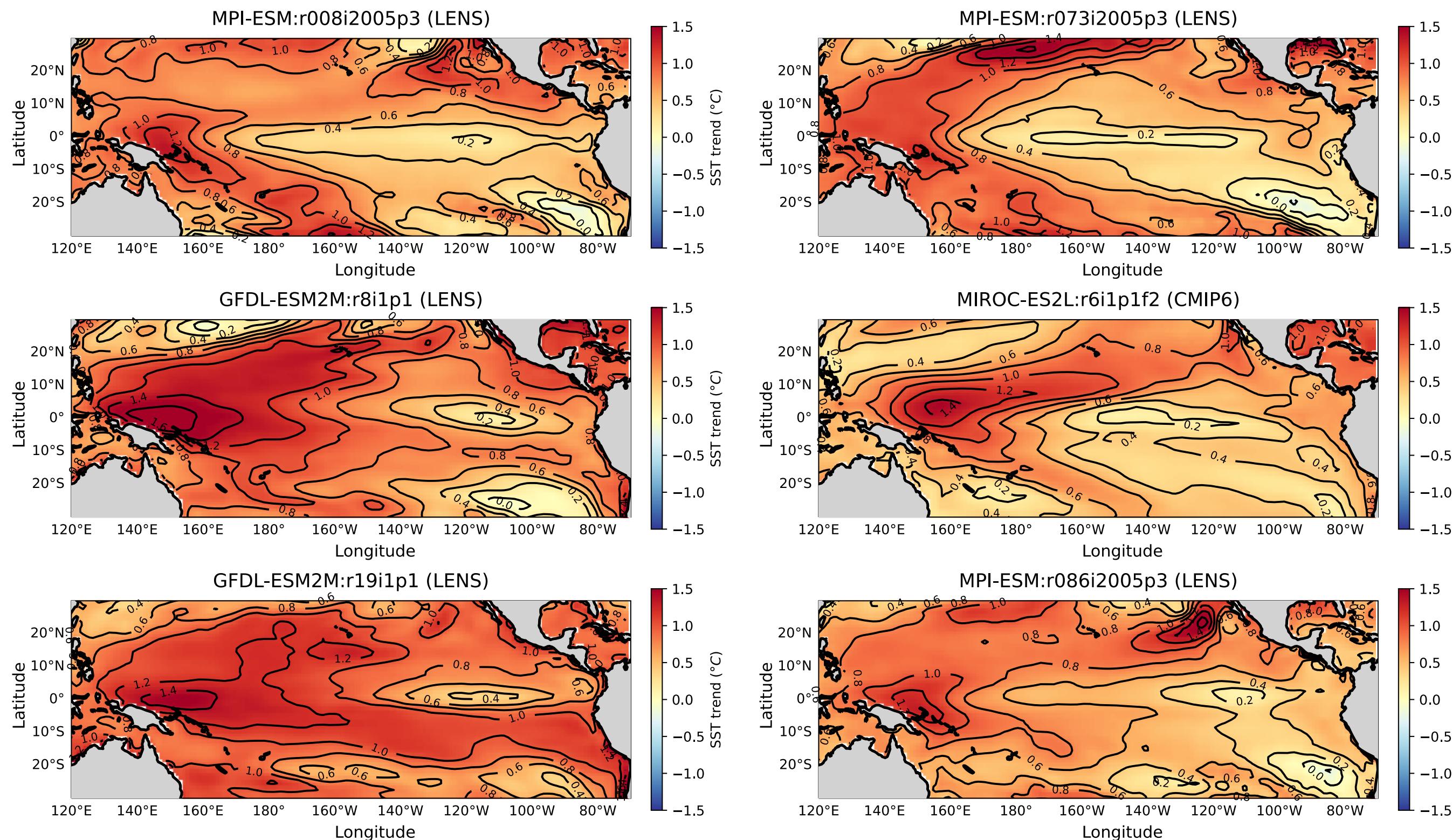
There is almost no overlap between 511 runs from 45 CMIP6 and 6 LENS models and the observed tropical Pacific SST trends



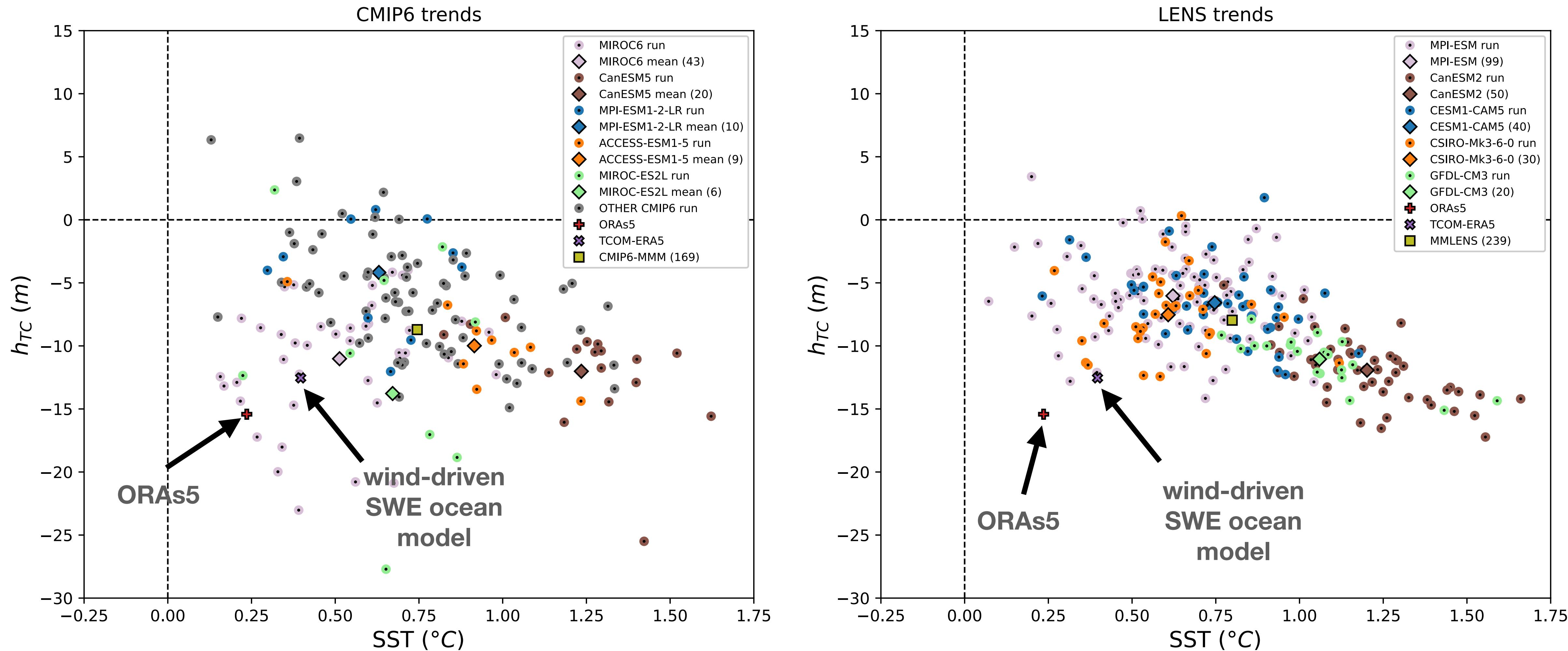
Could the observed trend be explained by natural variability?

Only a handful of 511 CMIP6+LENS model runs come even close to matching observations

CMIP6 and LENS Trends Closest to Observations



SST and Thermocline Depth Trends in NINO3.4 region

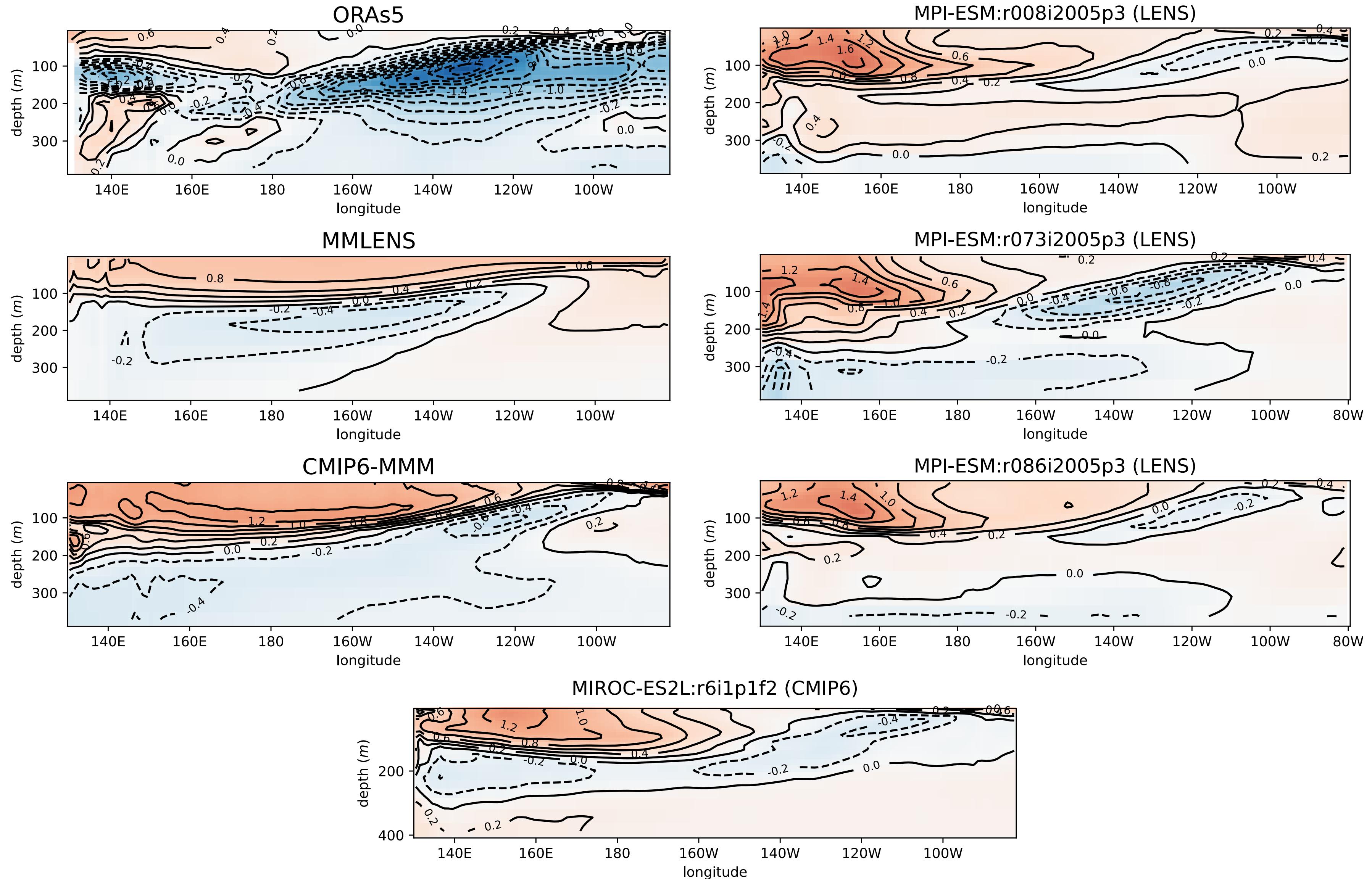


“observed” combination of weak warming and thermocline shoaling in the NINO3.4 region
exceedingly rare in CMIP6+LENS models

Upper Equatorial Ocean Temperature Trends, 1958-2018

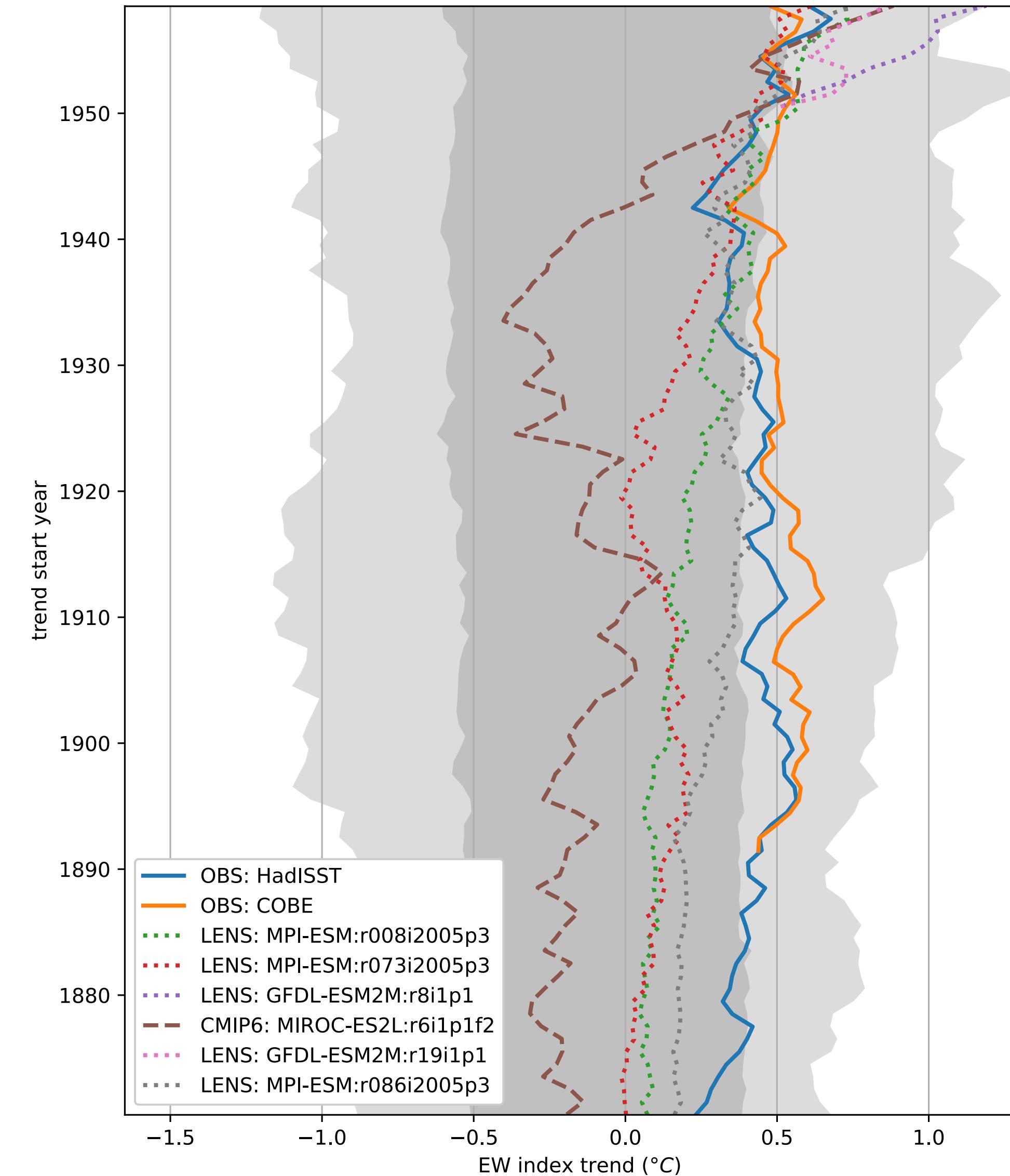
Strong dynamical
shoaling of the
equatorial thermocline
seen in ocean
reanalysis (and data).

Missing in CMIP6 and
LENS models, other
than a handful of runs
that weakly mimic



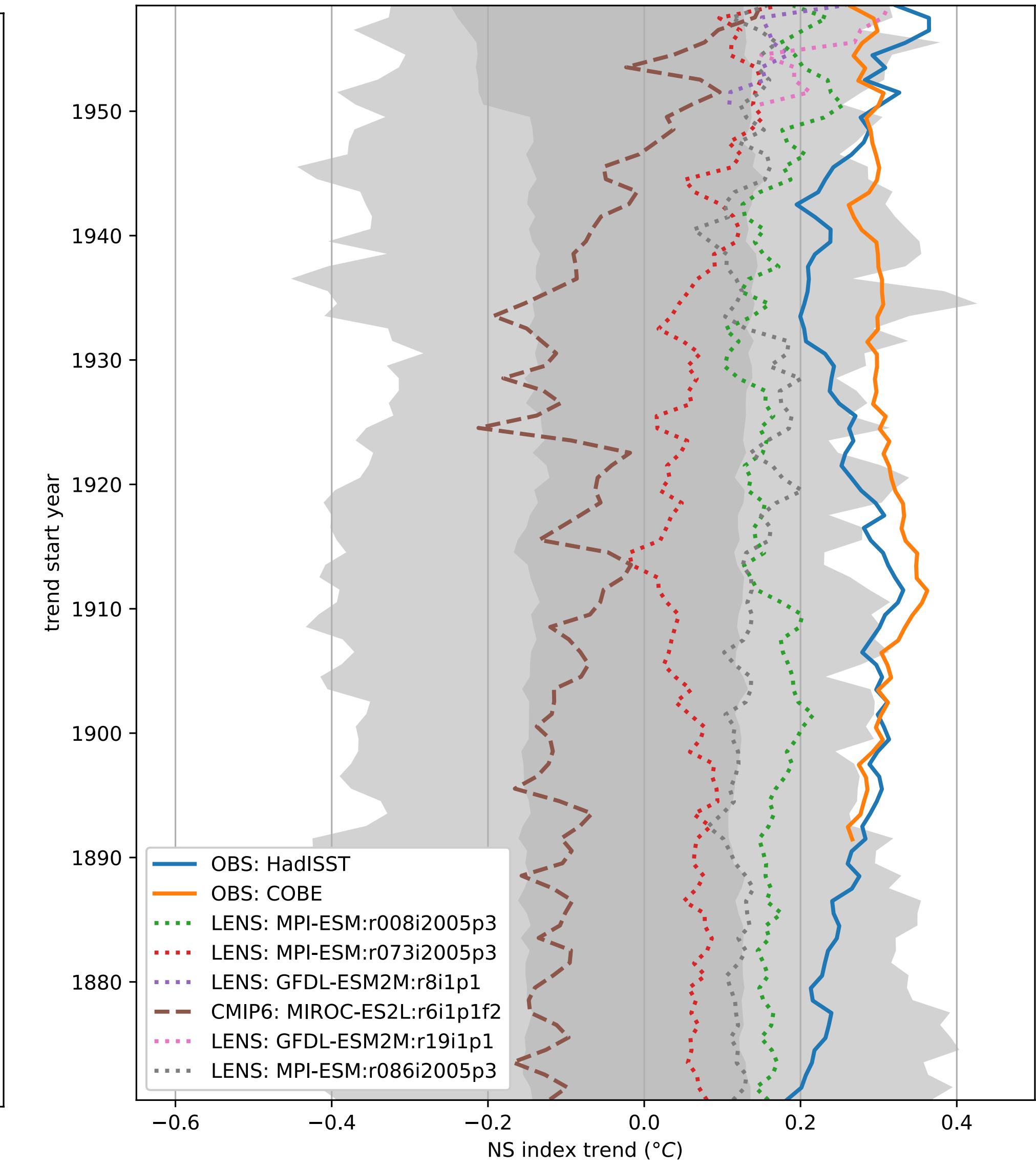
Observed E-W
and N-S gradient
trends have been
very persistent
over time -
suggestive of a
forced response

All Trends to year 2018 in the EW index
in Observations and Closest 6 Model Runs



No model runs
capture this
with possible
exception of
two MPI runs

All Trends to year 2018 in the NS index



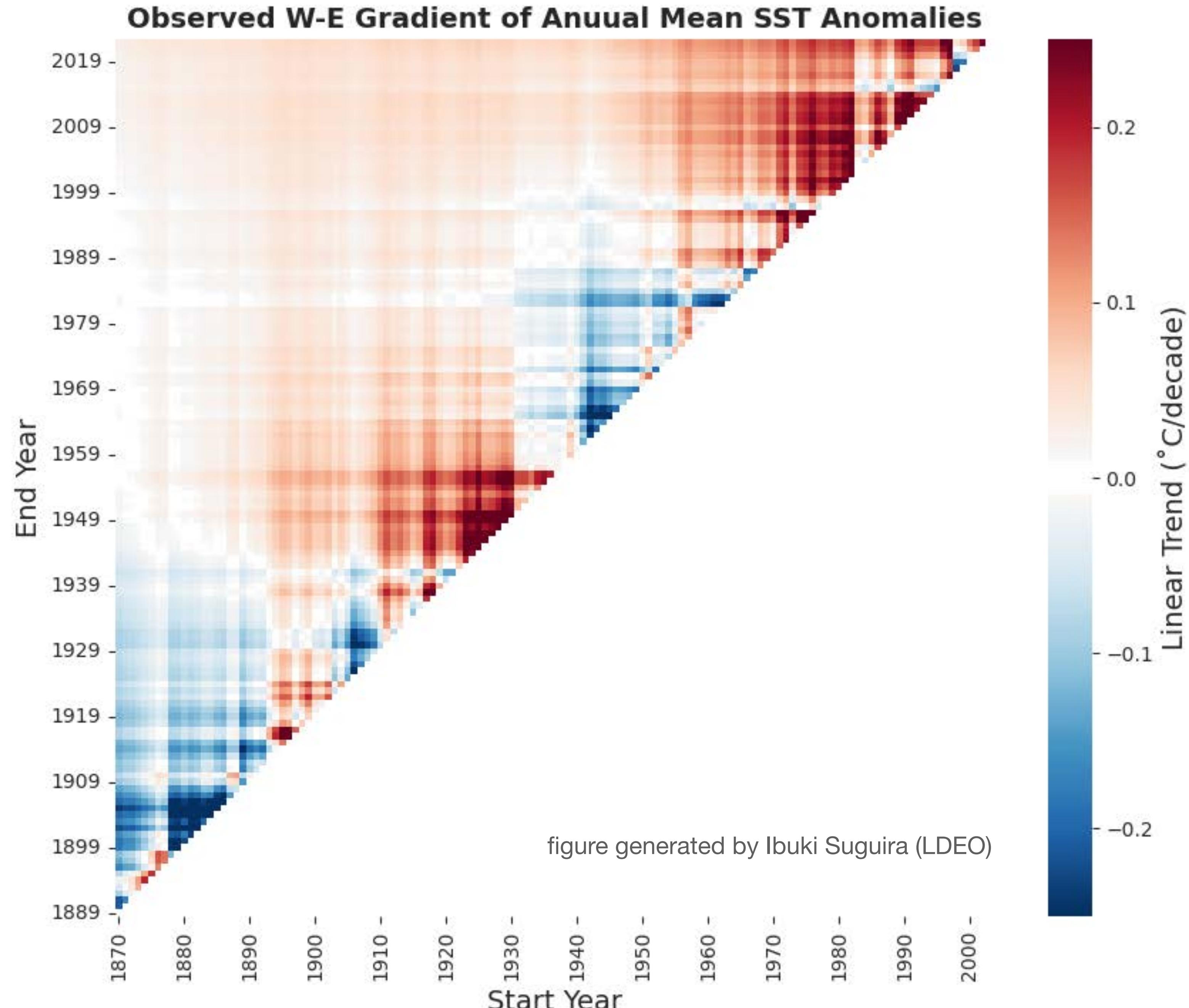
SST gradient trends over all start and end dates for periods 30 years and longer

The longer the period, the weaker the trend

The later the end date, the more likely the trend is positive

The longest trends are weakly positive

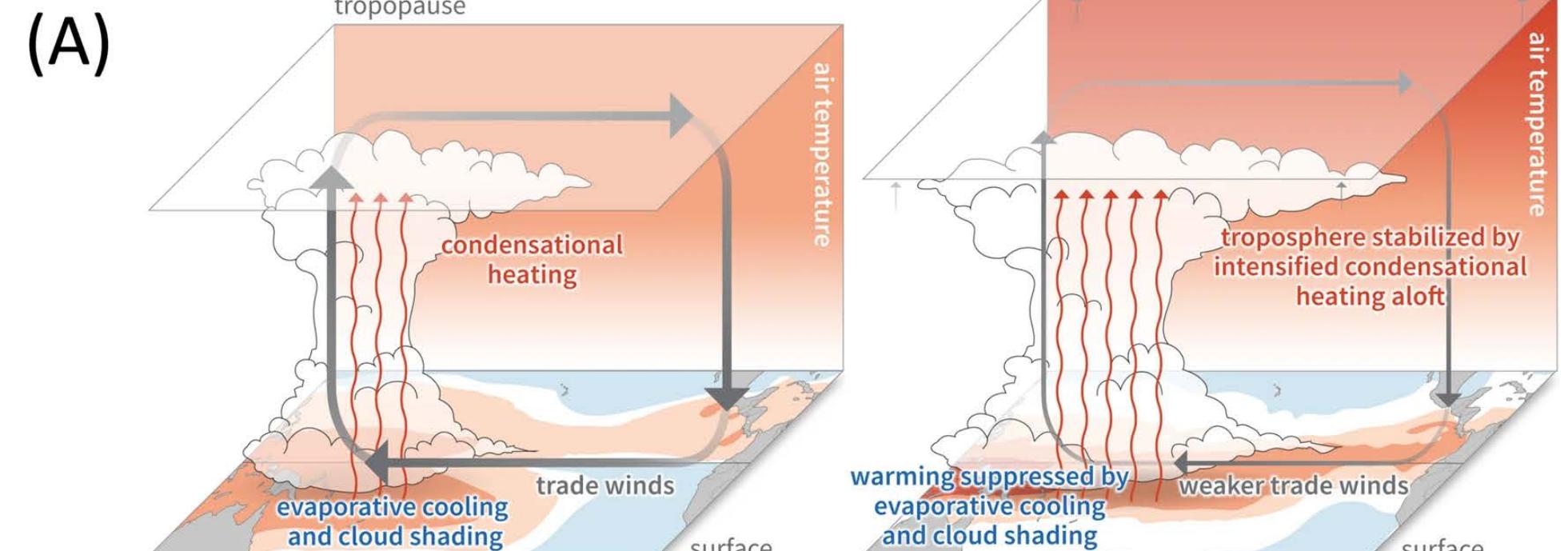
Given that (1) longer the period the more able to identify a forced signal from the variability, (2) the recent strength of the forcing is this consistent with a forced strengthening gradient trend?



(A) processes that can cause a reduced zonal gradient

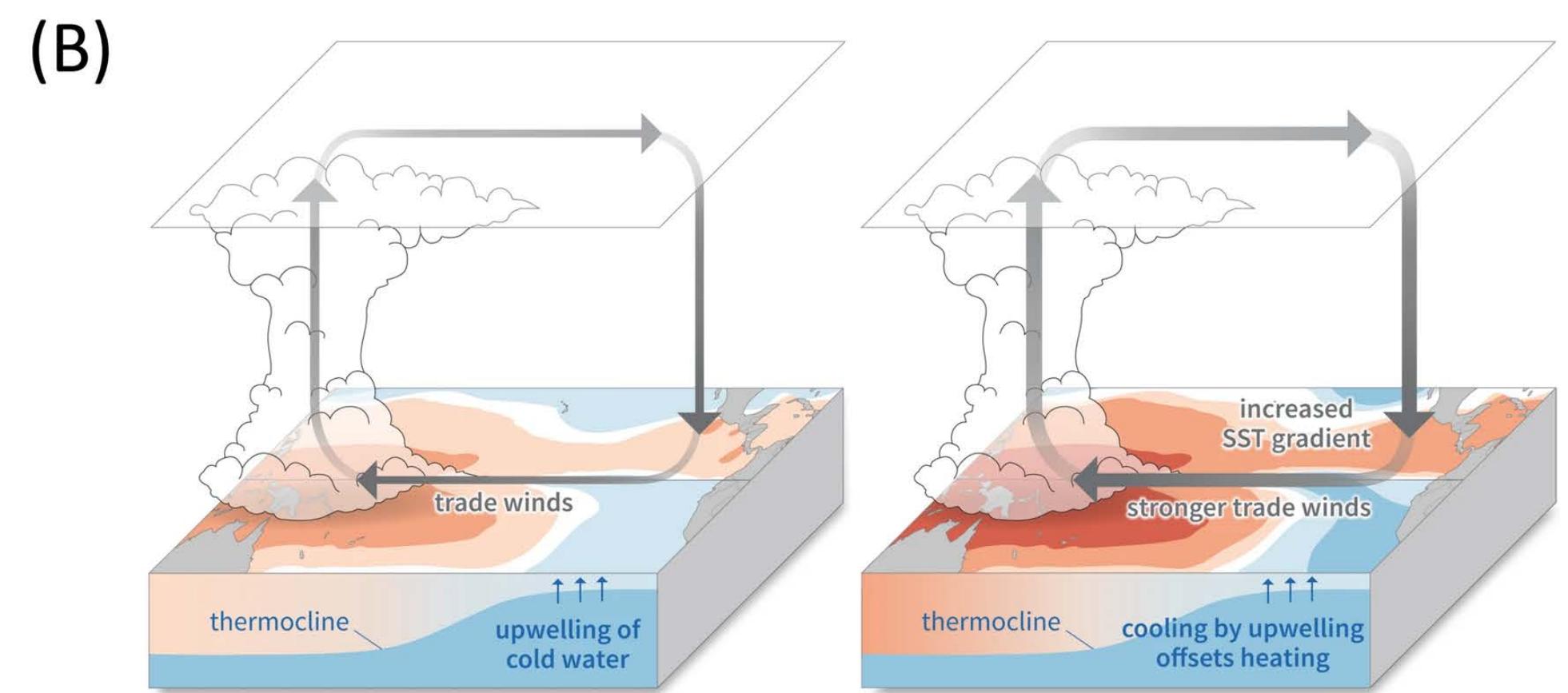
Enhanced thermal stratification in atmosphere causes a weaker Walker circulation

Cooler SST must warm more for same LH' to balanced enhanced GHG forcing



(B) coupled processes that cause an enhanced zonal gradient

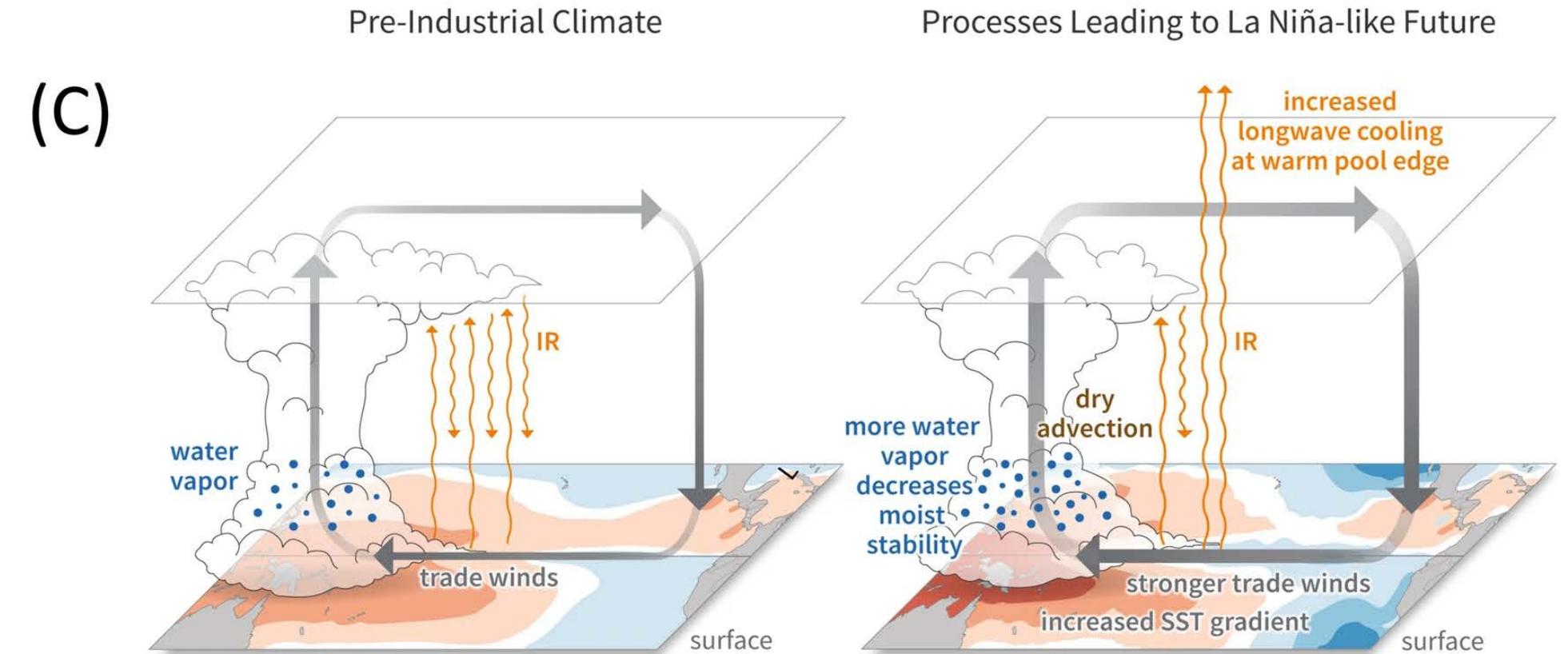
GHG induced warming in east is offset by enhanced cooling due to vertical advection, Walker circulation strengthens, thermocline shoals in east cooling SST



(C) enhanced gradient from atmospheric processes

More moisture causes reduced Gross Moist Stability and a stronger Walker circulation.

Dry zonal advection narrows warm pool convection and cools central Pacific

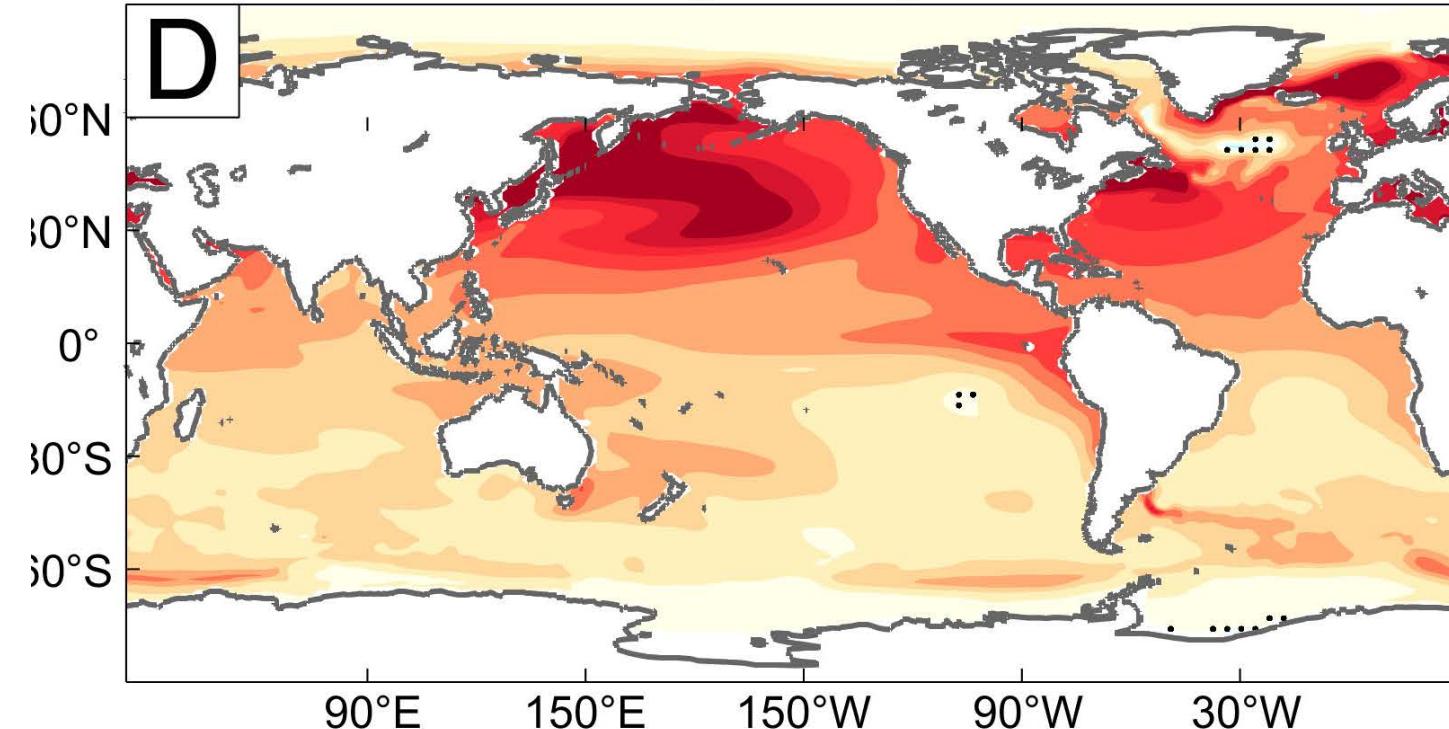


Another hypothesis ... the equatorial Pacific cooling is a response to Southern Ocean cooling

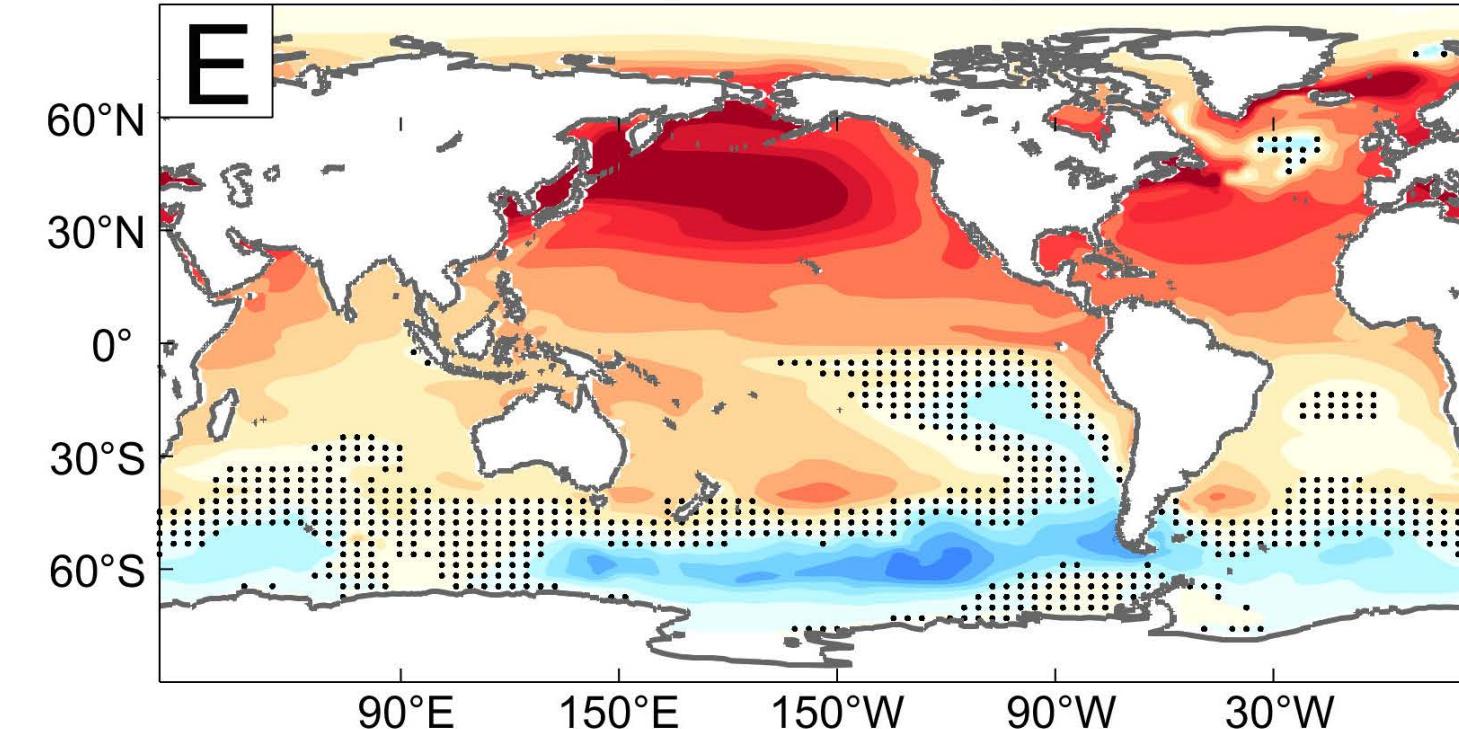
SST trends 1979 to 2013

Kang et al. (2023 PNAS)

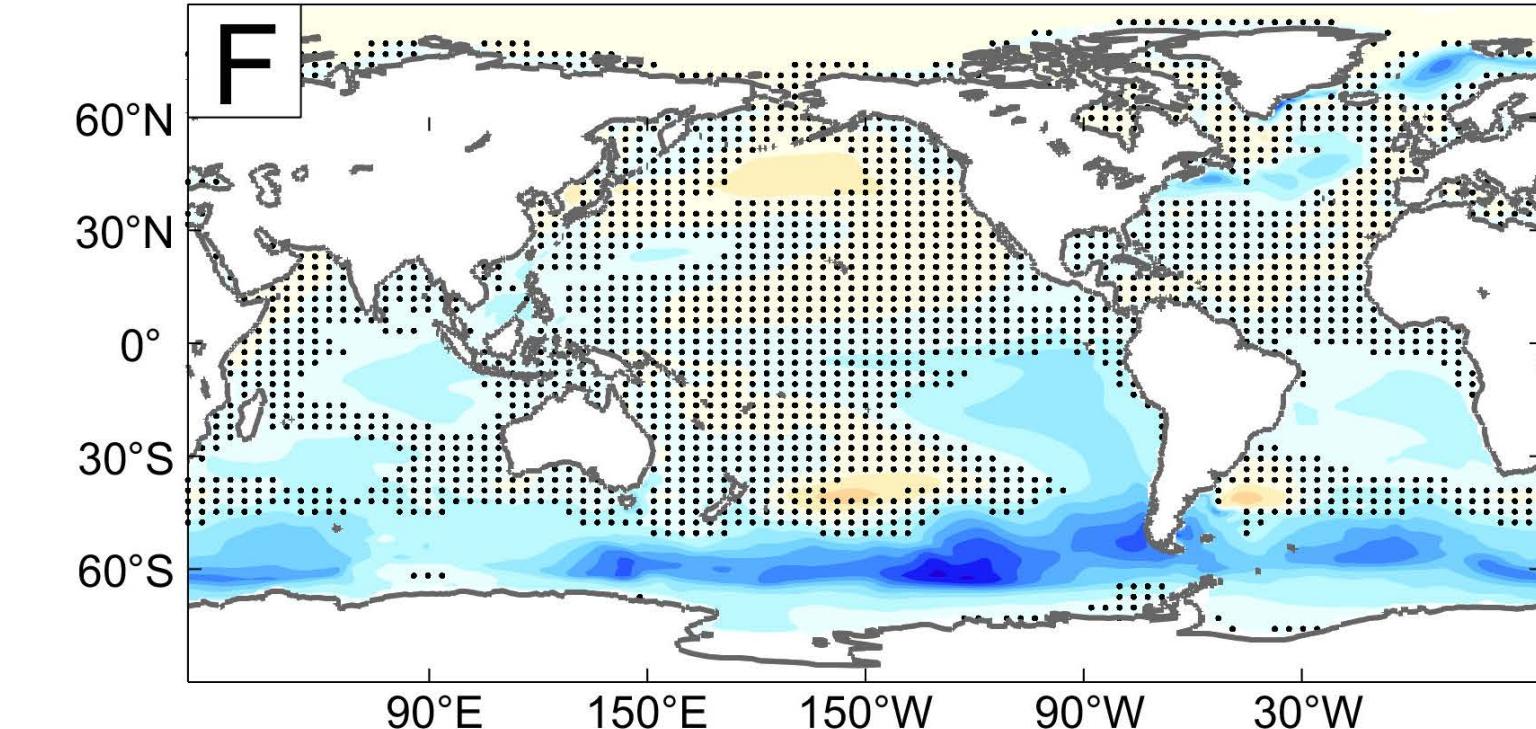
fully coupled CESM2
[HIST2]



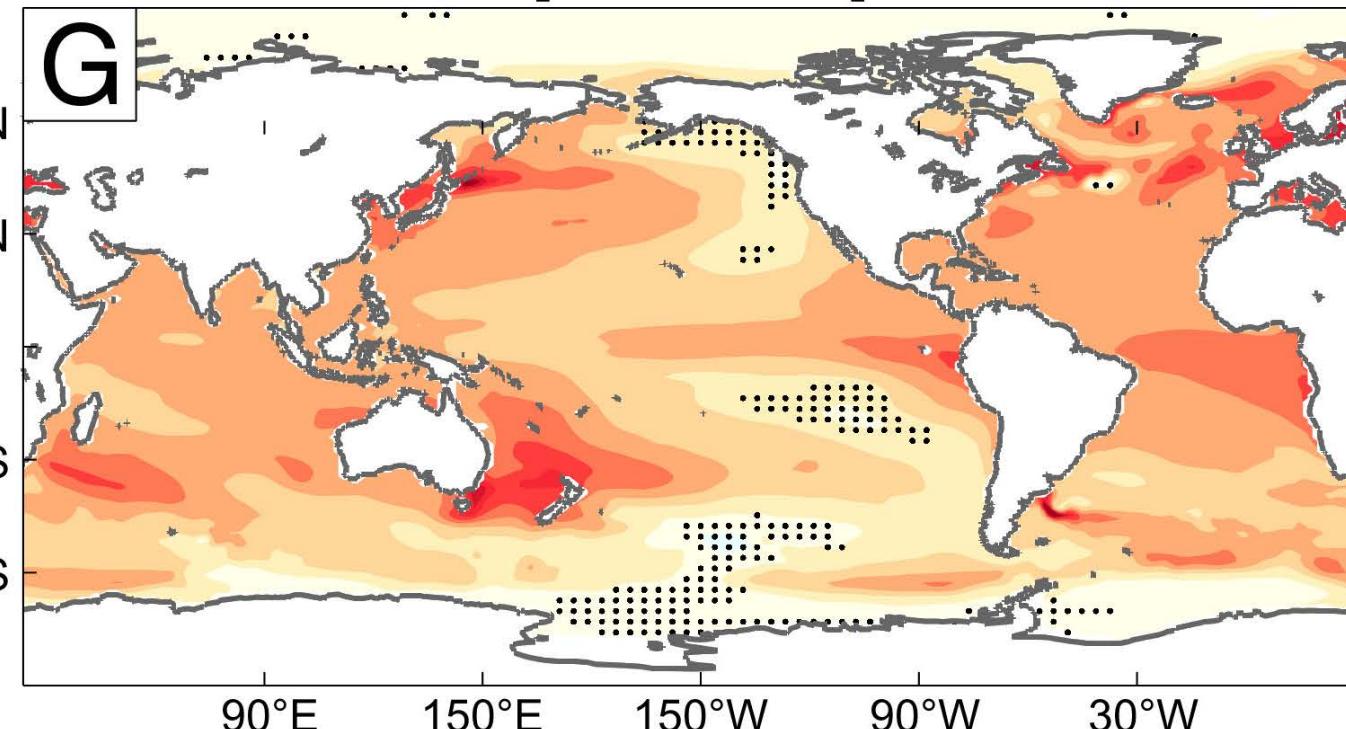
Southern Ocean pacemaker CESM2
[SOPACE2]



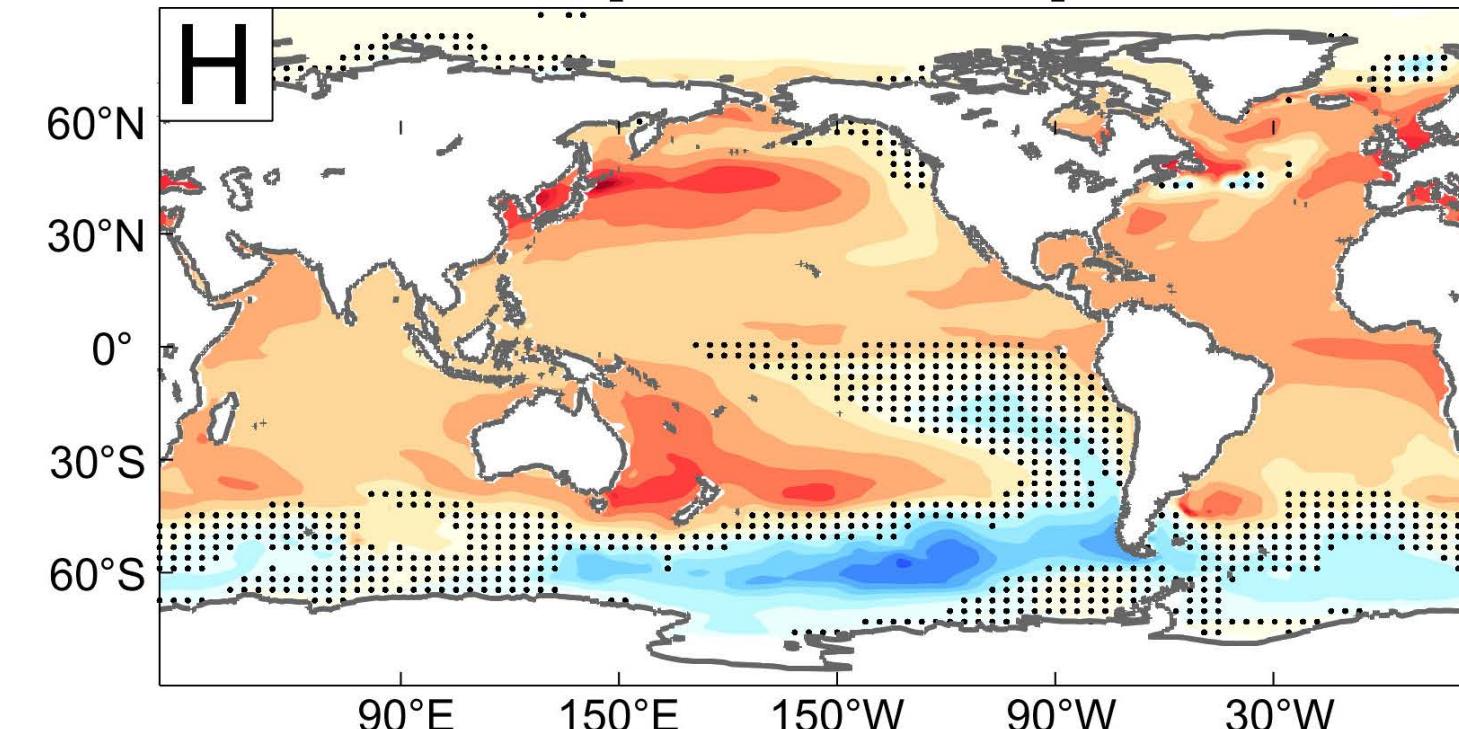
difference (above) obs (below)
SO-driven2



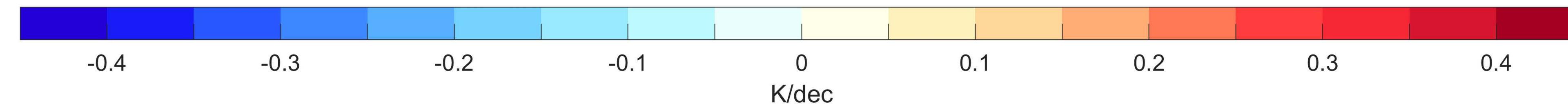
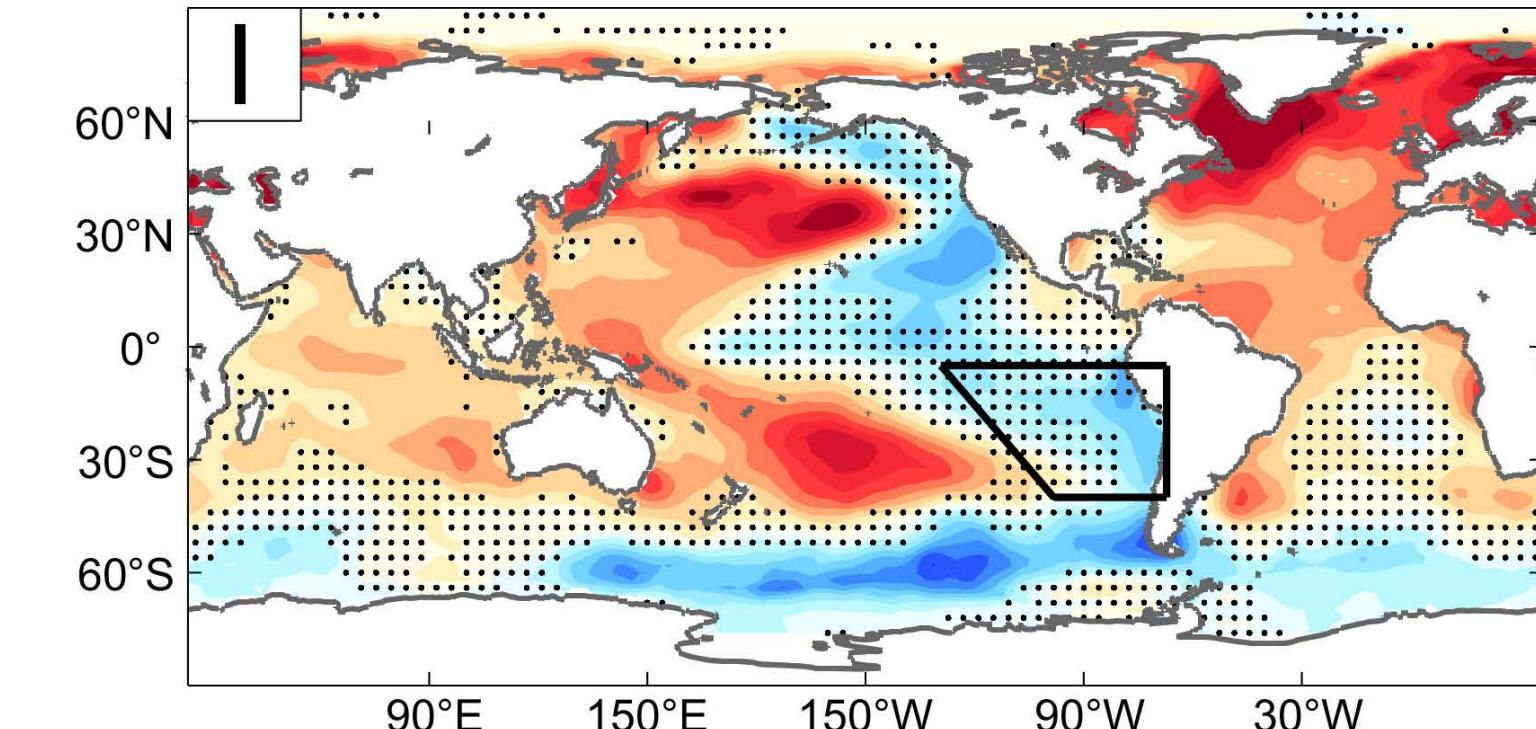
[HIST2-C5]



[SOPACE2-C5]



ERSSTv5



SO cooling → stronger SE Pacific trades → cooler SSTs → more low level clouds → cooler SSTs

Conclusions

Observations are consistent in indicating an enhancing east-west SST gradient across the equatorial Pacific over the past decades and century

Climate models from CMIP3,5,6 indicate that radiative forcing either has, or soon will, reduce this gradient

Some individual runs of some CMIP6 climate models do strengthen the gradient and consistently over past century but almost none with a pattern like that observed

Perhaps, the observed record is dominated by internal variability - an about one in a hundred chance according to models

Or - more likely in our view - the observed record is the forced response and CMIP models get this wrong due to chronic tropical Pacific mean state biases (or remote influences)

Either way, the difference between observed and CMIP modeled SST trends matters for regional hydroclimate, tropical cyclones, drought risk and more (but that's another talk)