

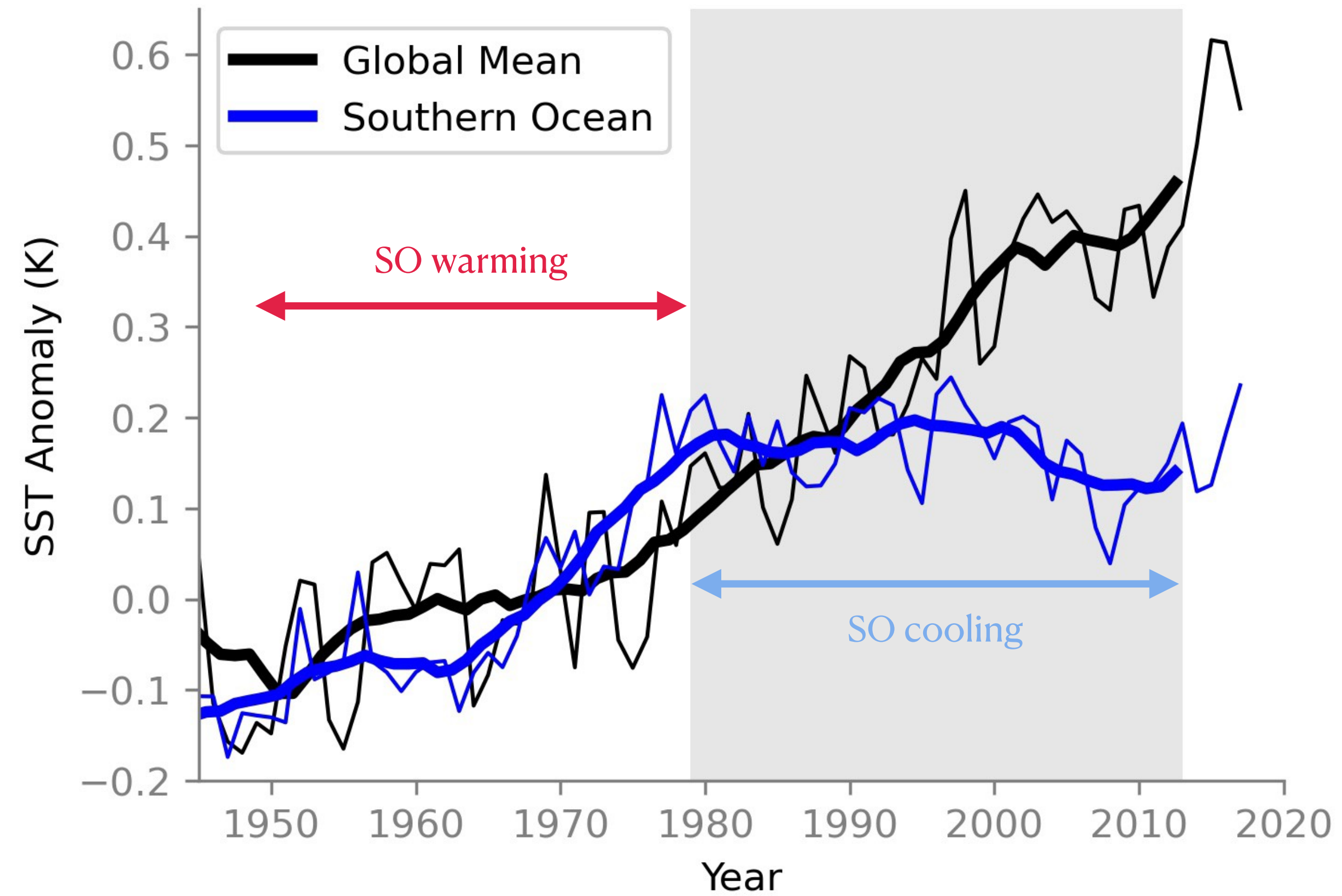
Global Impacts of Recent Southern Ocean Cooling

Xiyue (Sally) Zhang

**Department of Earth and Planetary Sciences
Johns Hopkins University**

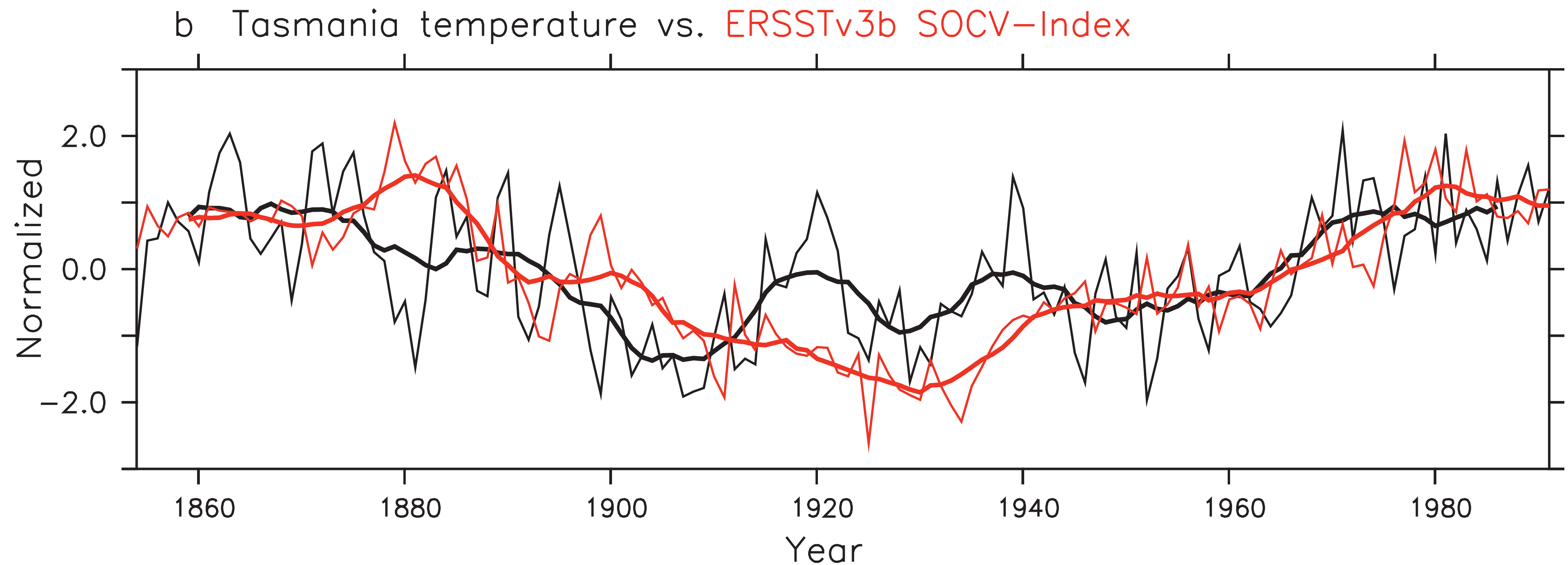
Clara Deser (NCAR), Sarah Kang (UNIST), Yue Yu (SIO)

The Southern Ocean (SO) surface has cooled since late 1970s



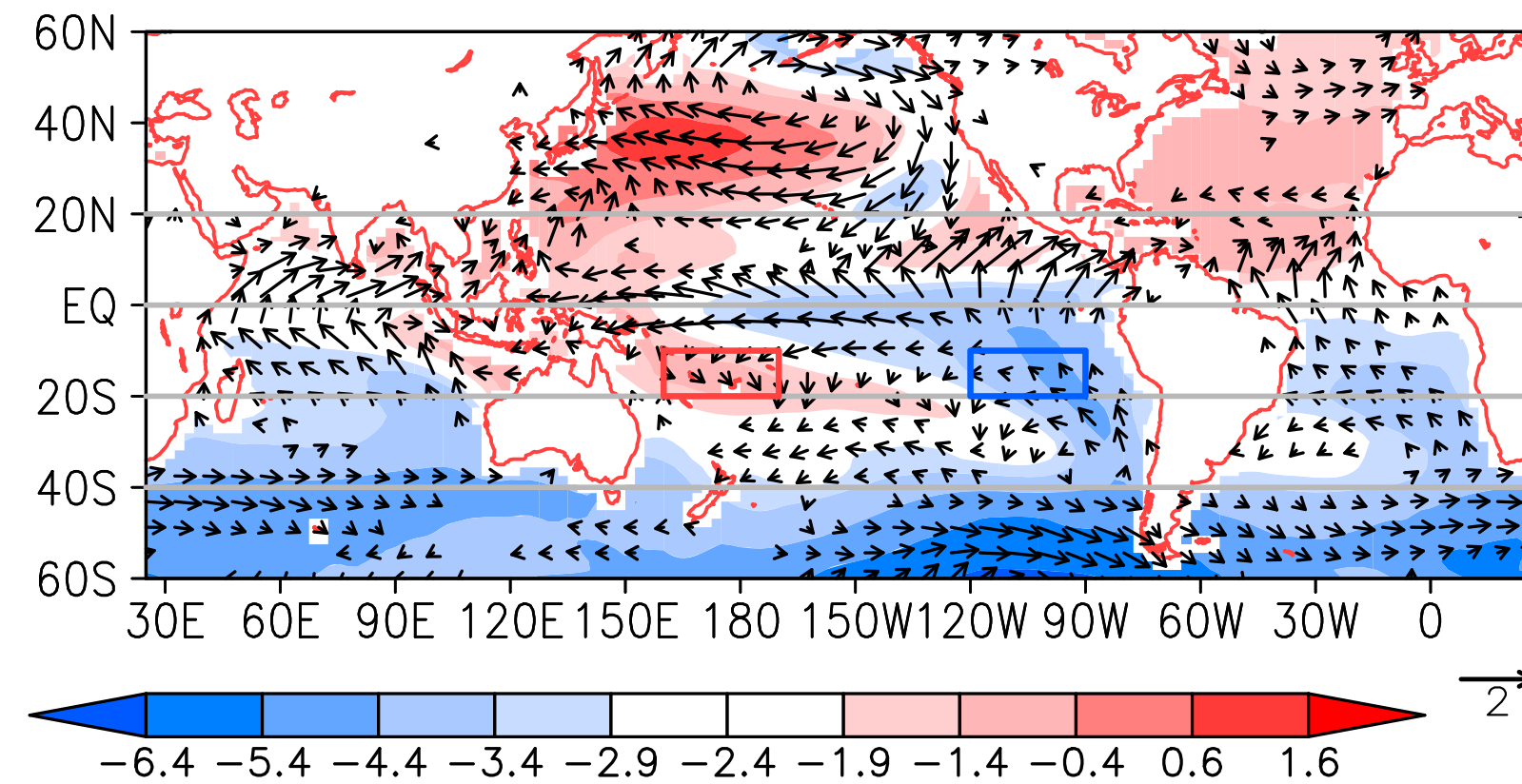
Paleoclimate proxy shows SO SST multi-decadal variability

With tree ring records from Tasmania



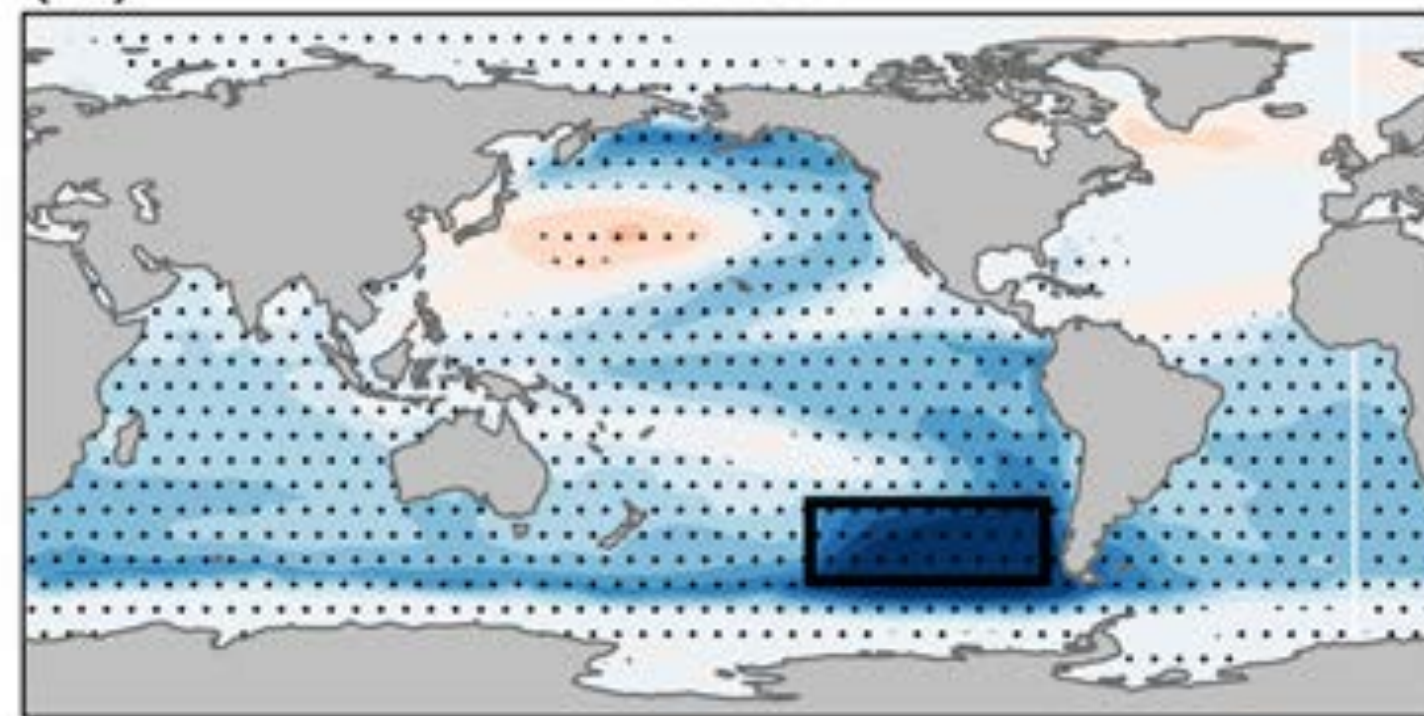
Previous studies suggest a robust response due to SH extratropical forcing

Hwang et al. (2017)



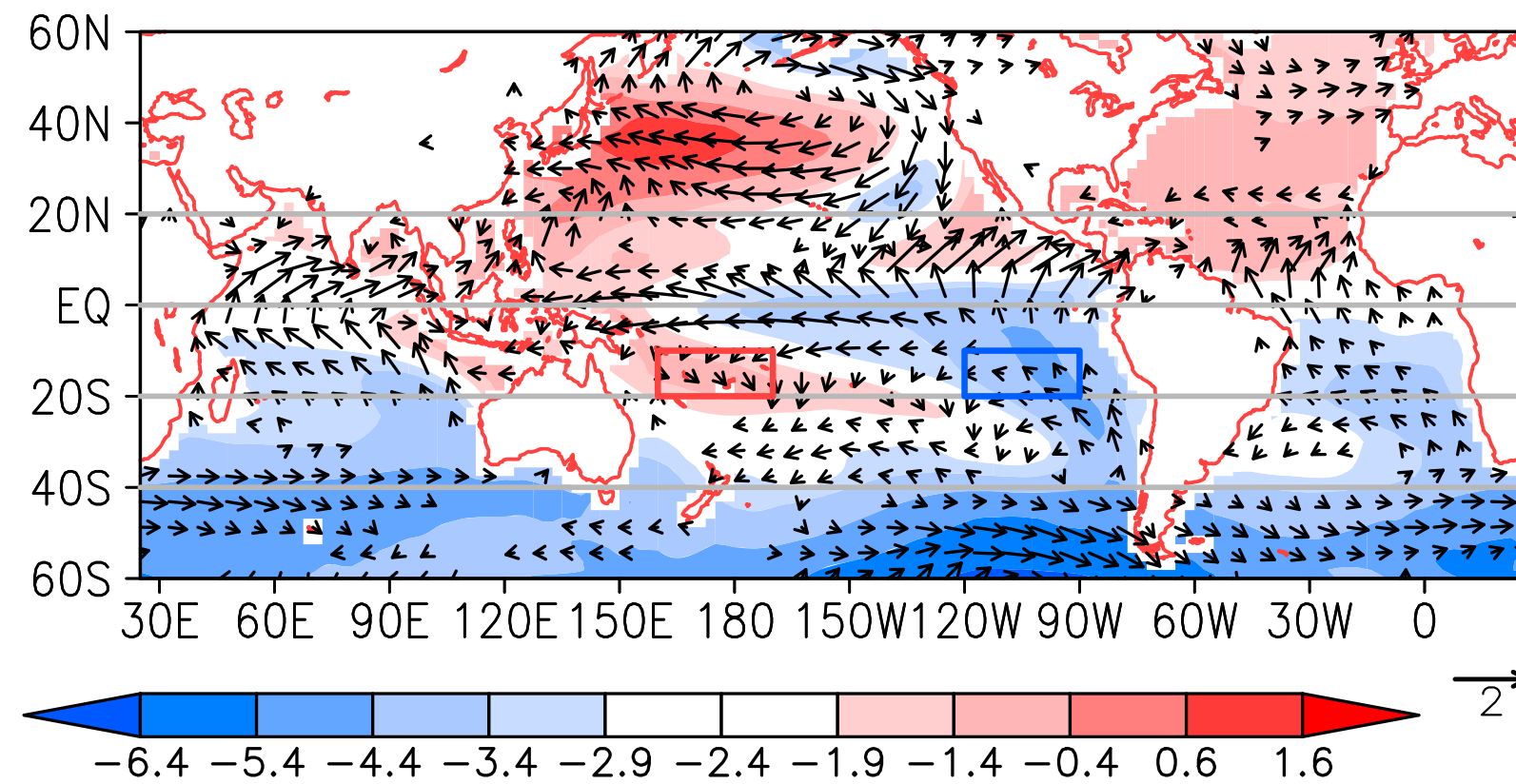
Slab
ocean

(b) Δ SST Dong et al. (2022)



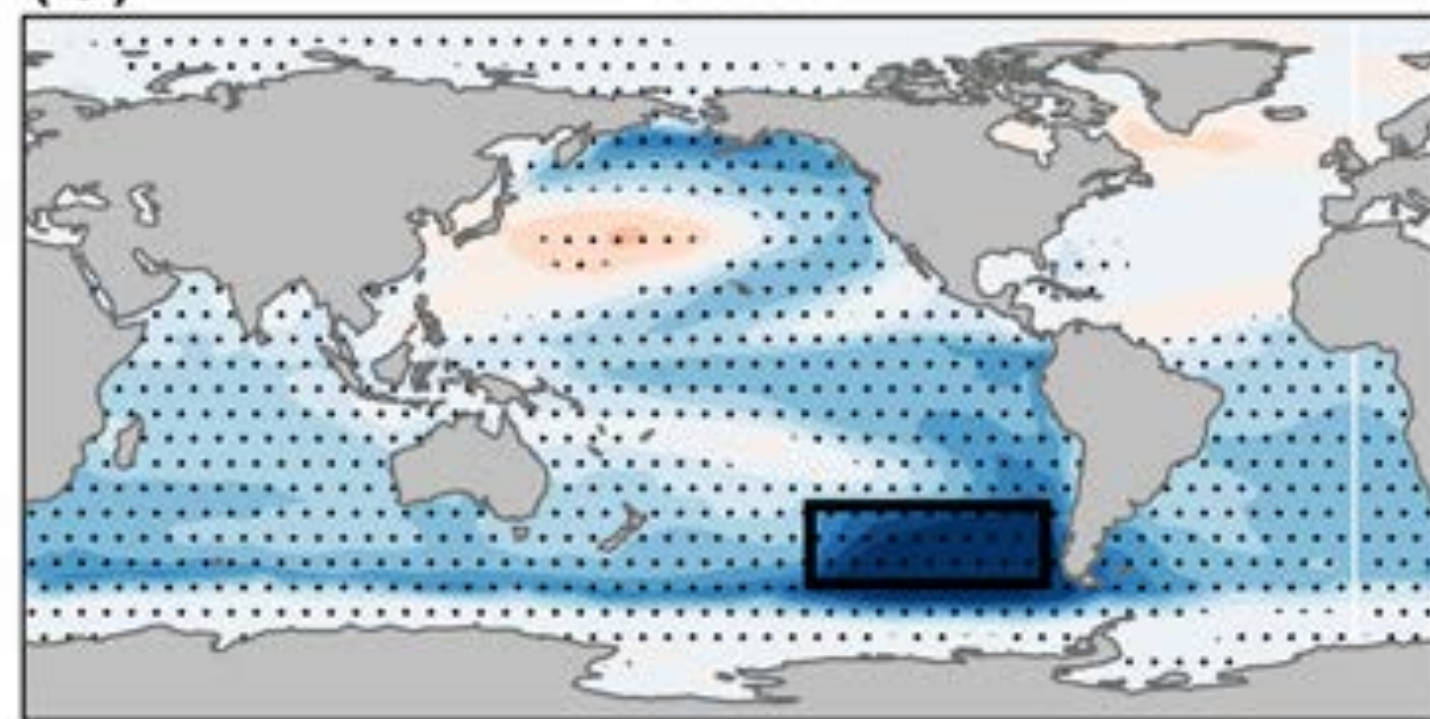
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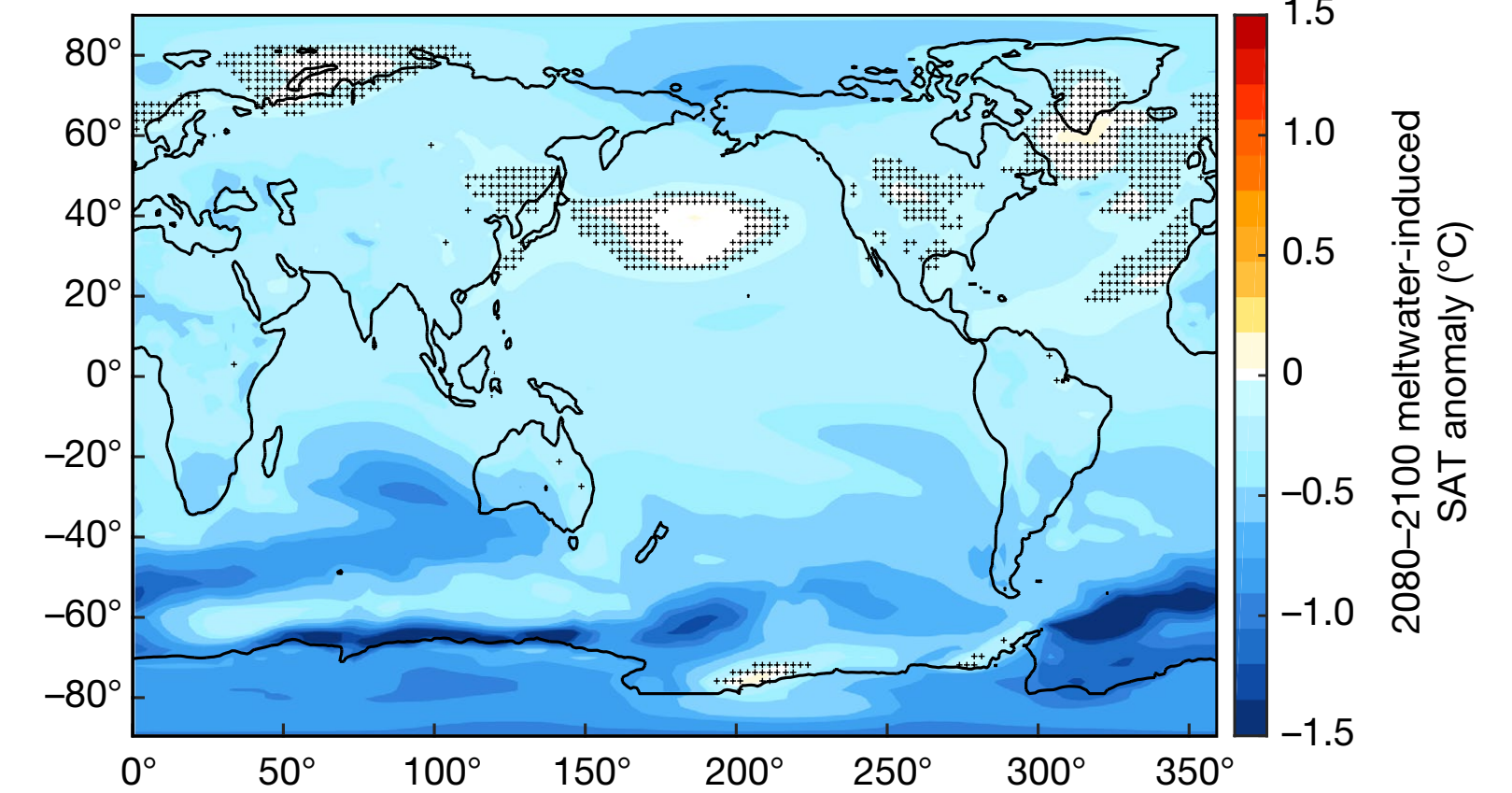


Slab
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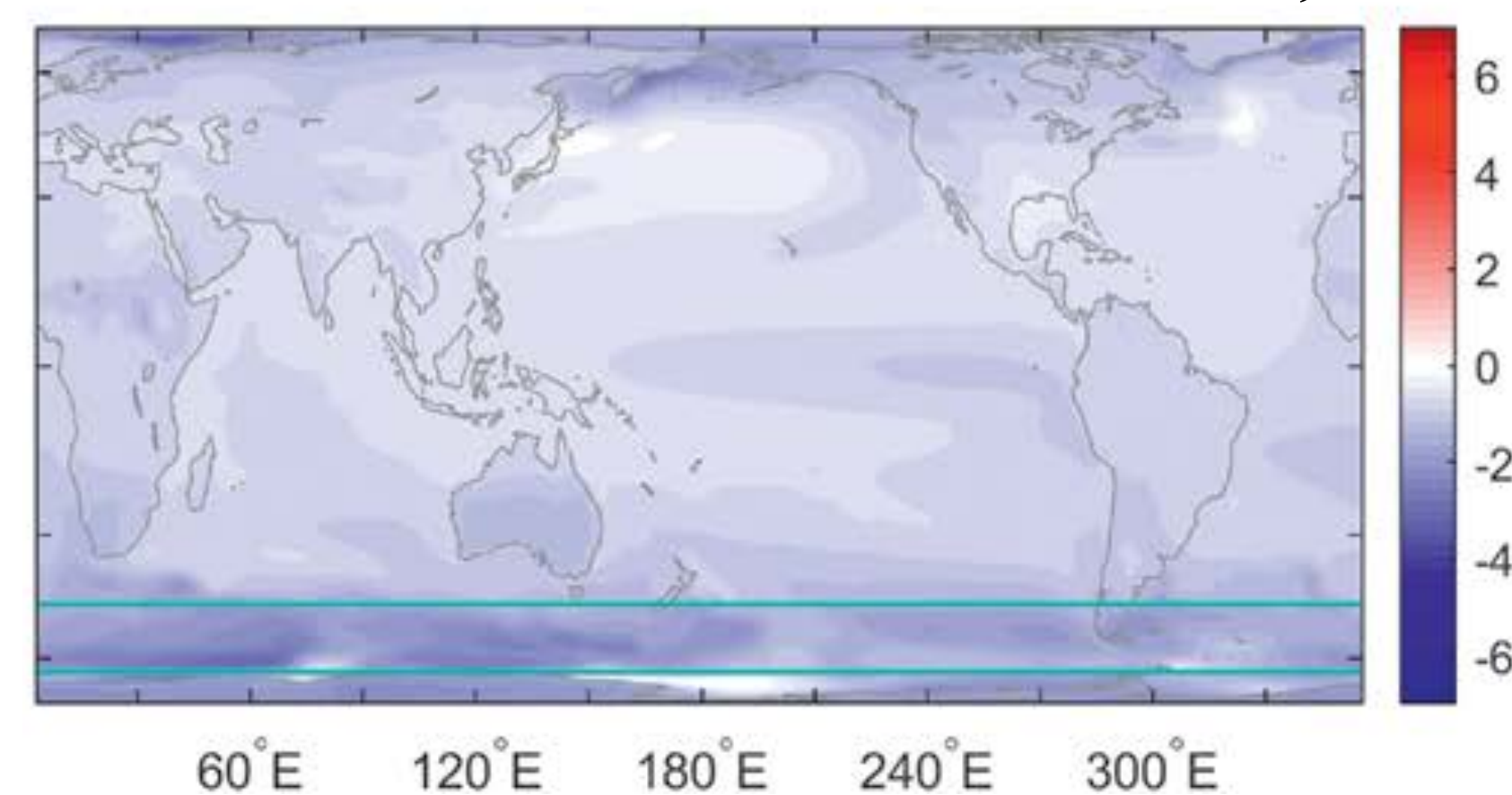


Bronse laer et al. (2018)



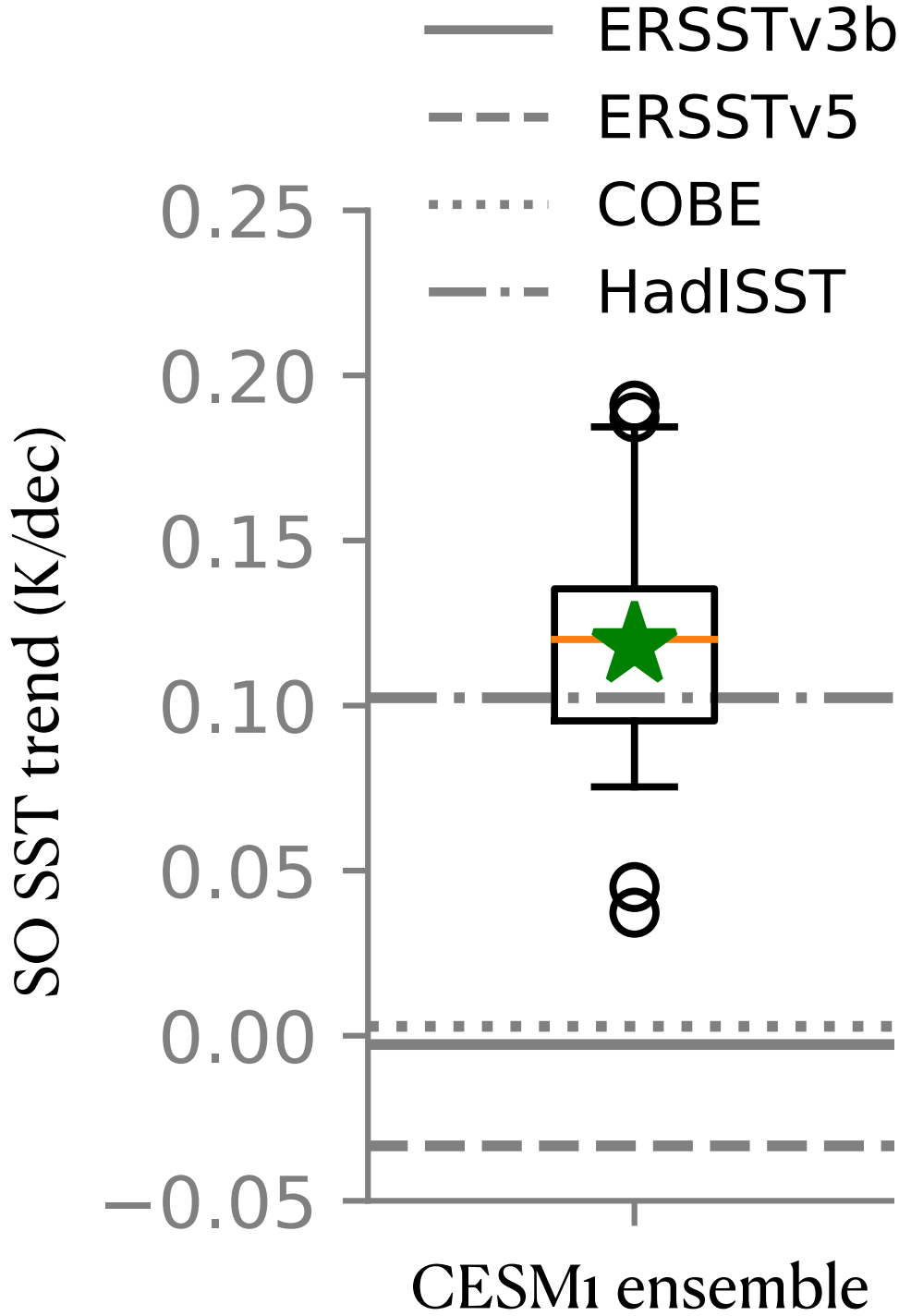
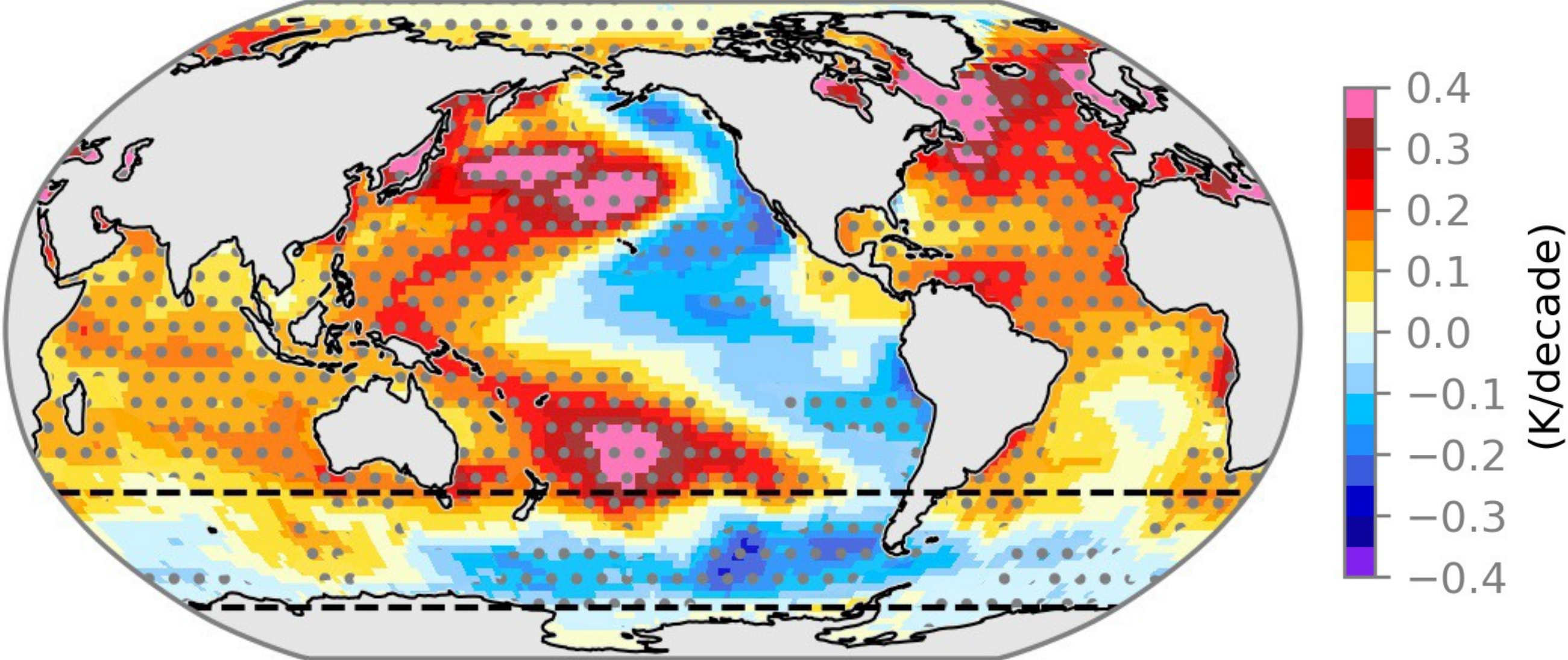
Dynamic
ocean

Kang et al. (2019)



Climate models are unable to capture observed SO SST trends

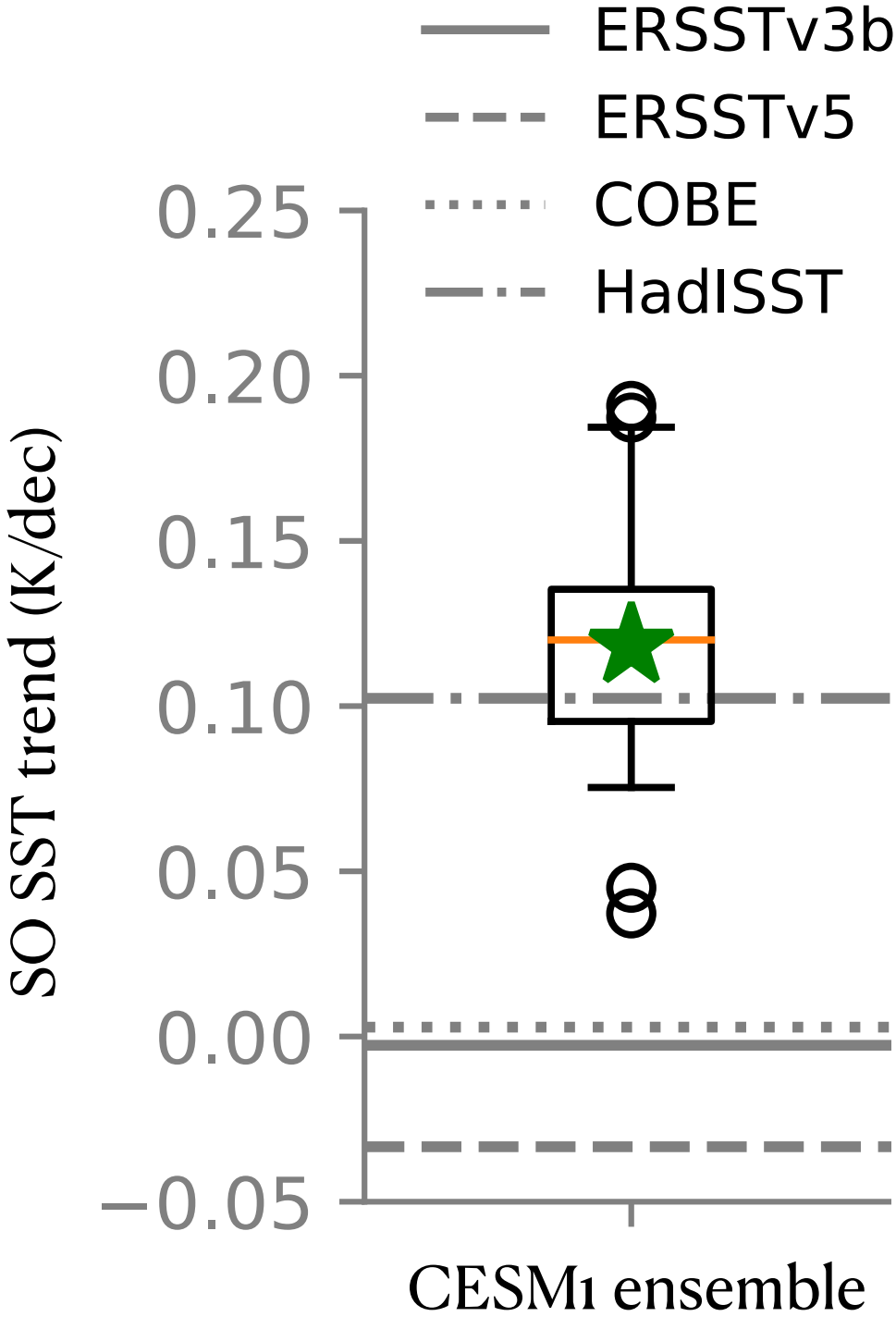
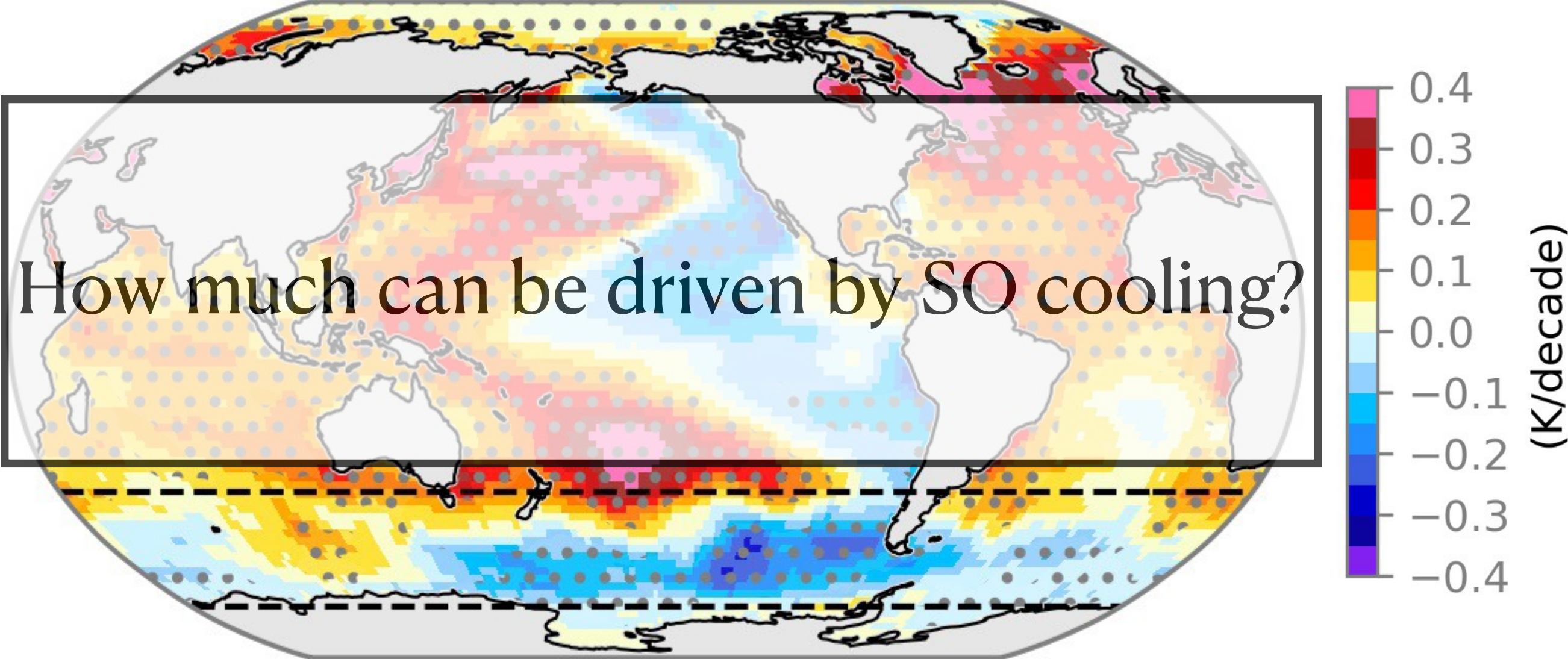
Observed trends 1979-2013



Stippling shows significant trends at 95% level
Data: ERSSTv3b; Zhang et al. (2021)

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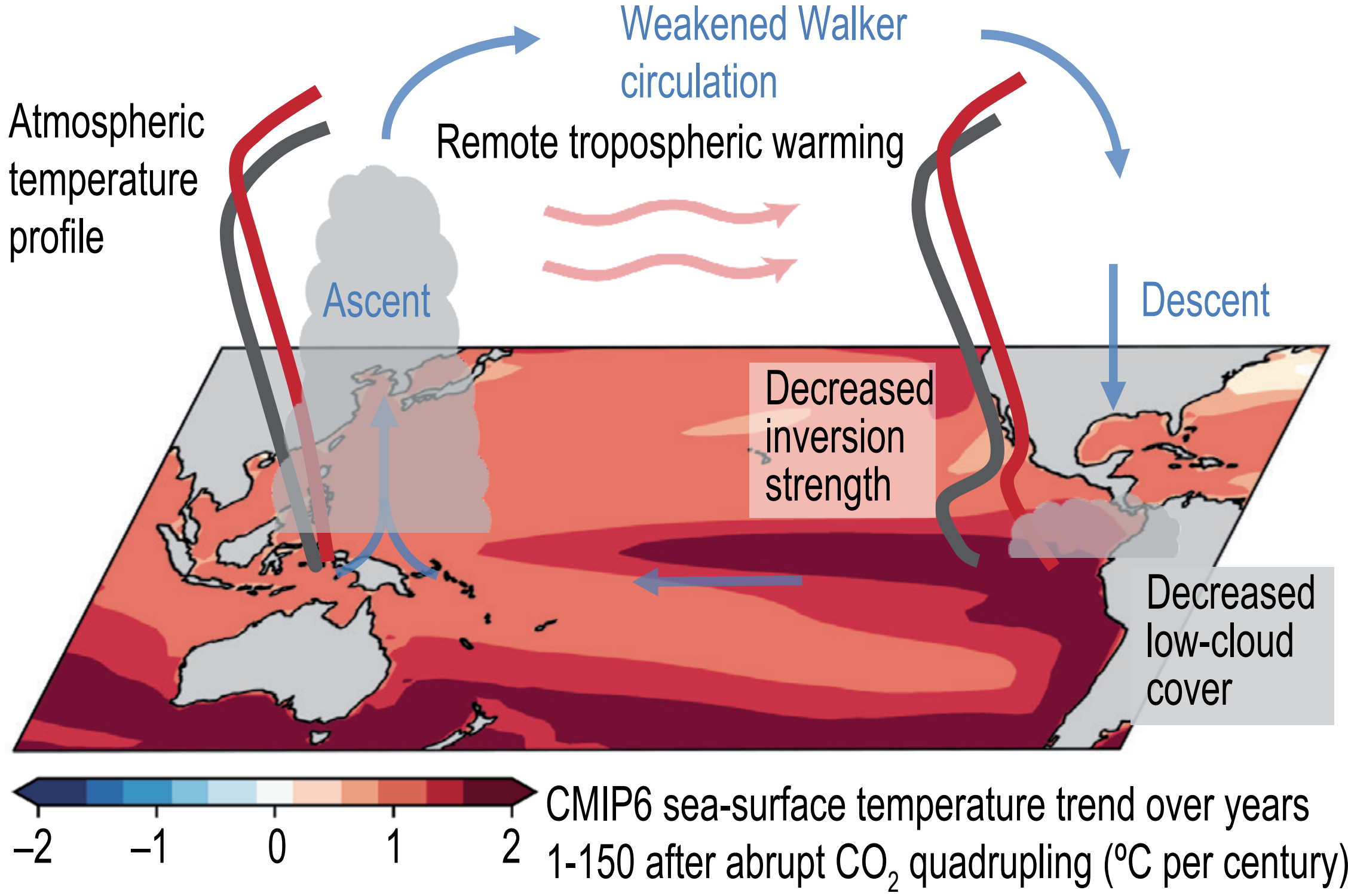
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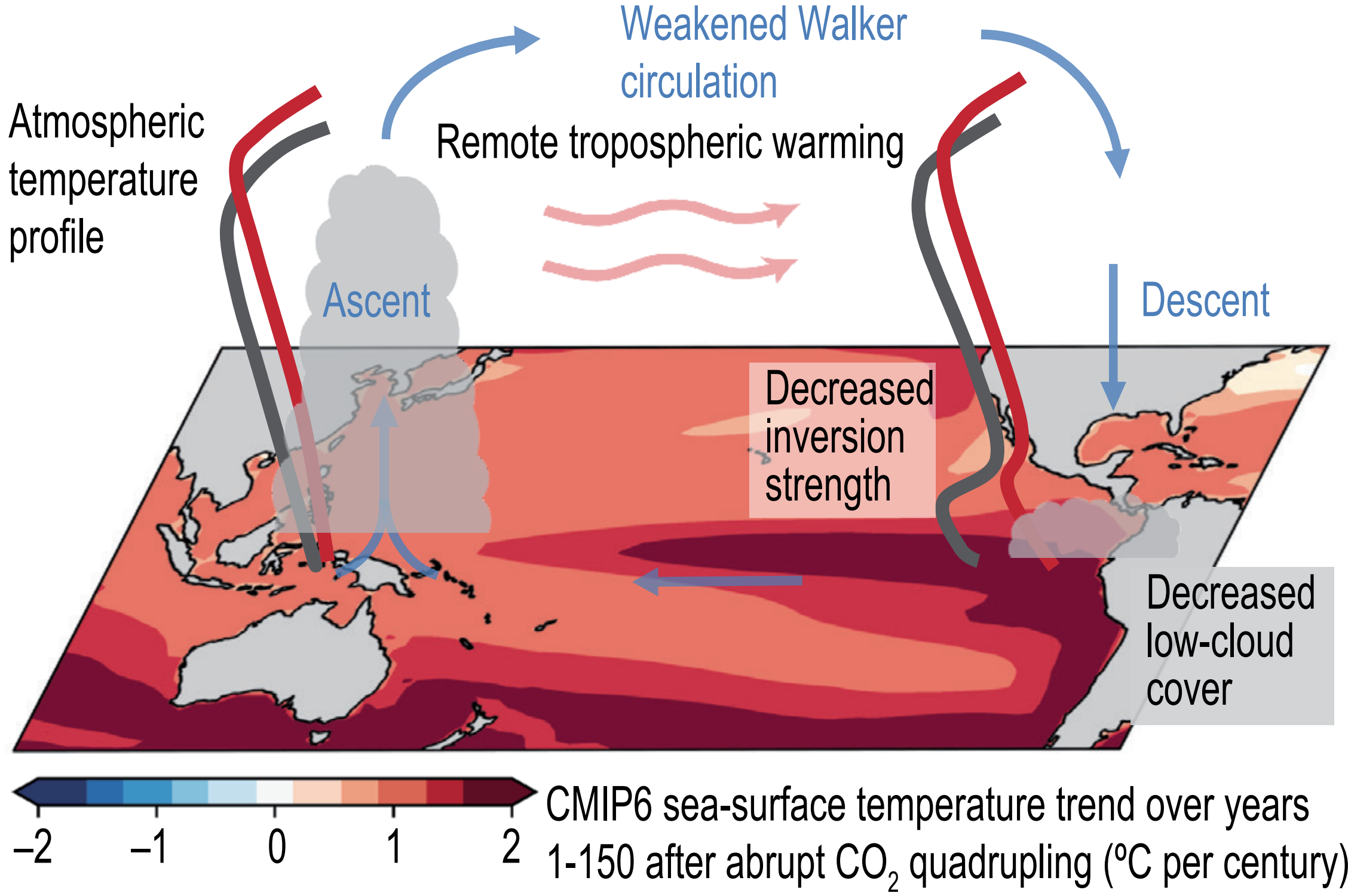
SO-induced teleconnection can affect the tropical warming patterns and climate sensitivity

Projected warming pattern



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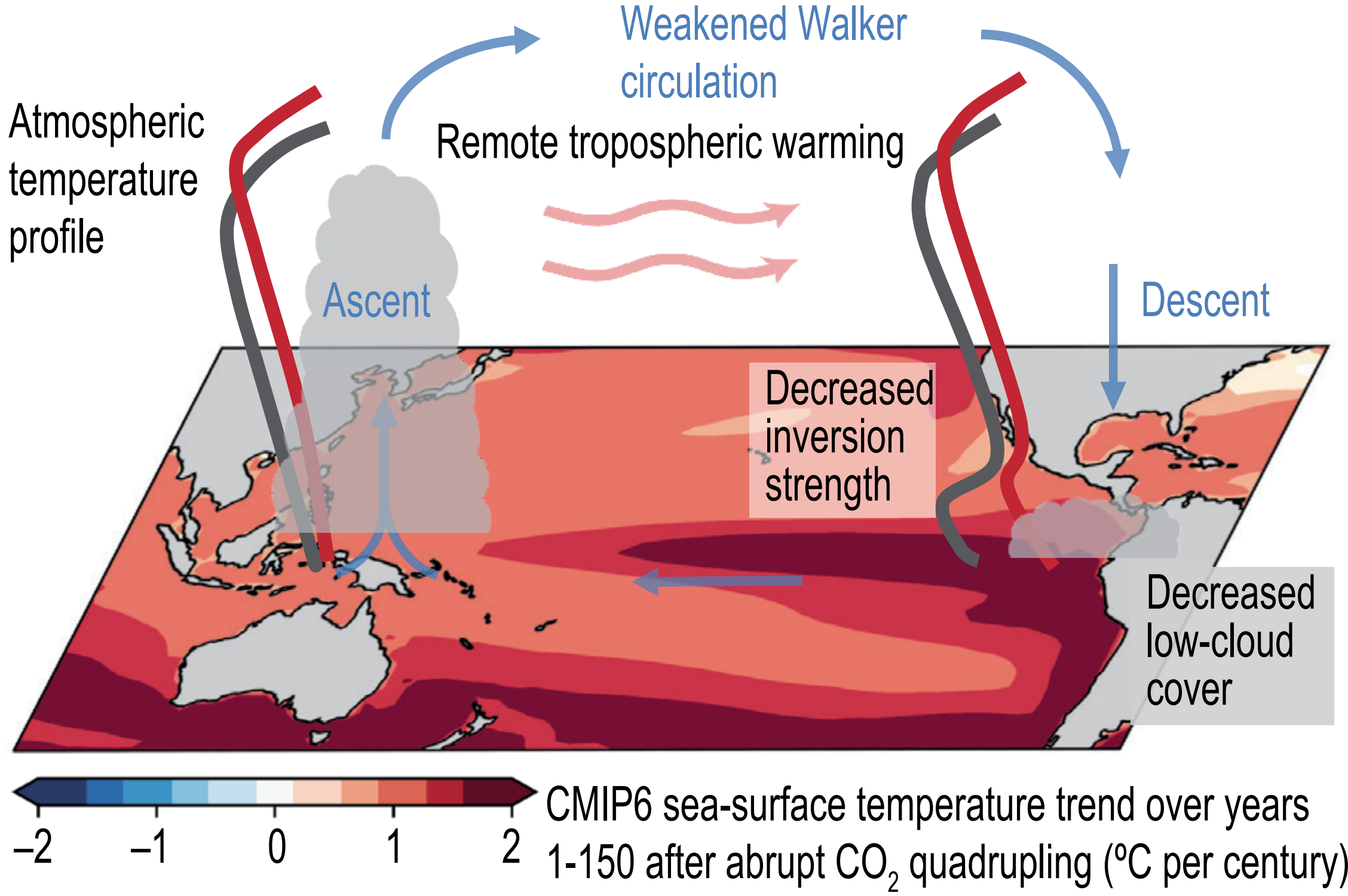
Projected warming pattern



Positive Feedback

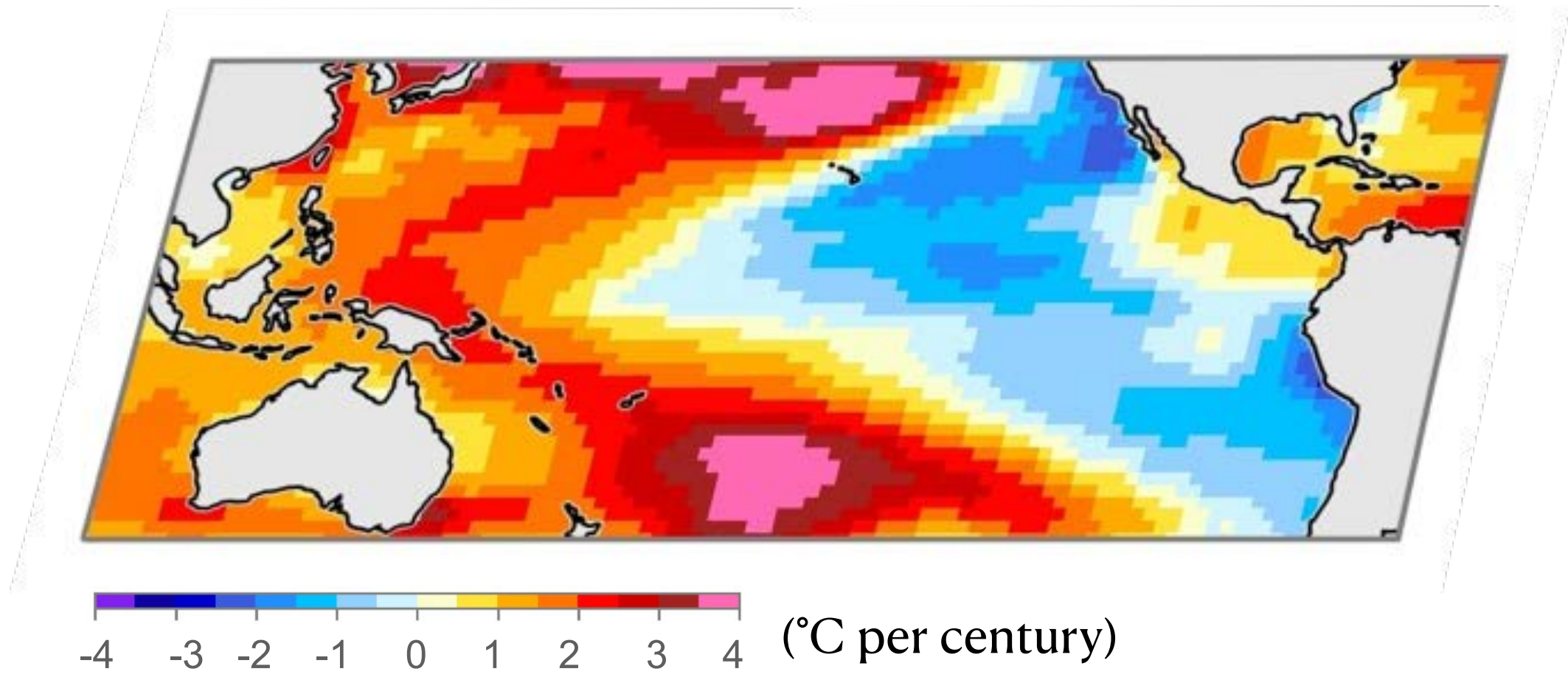
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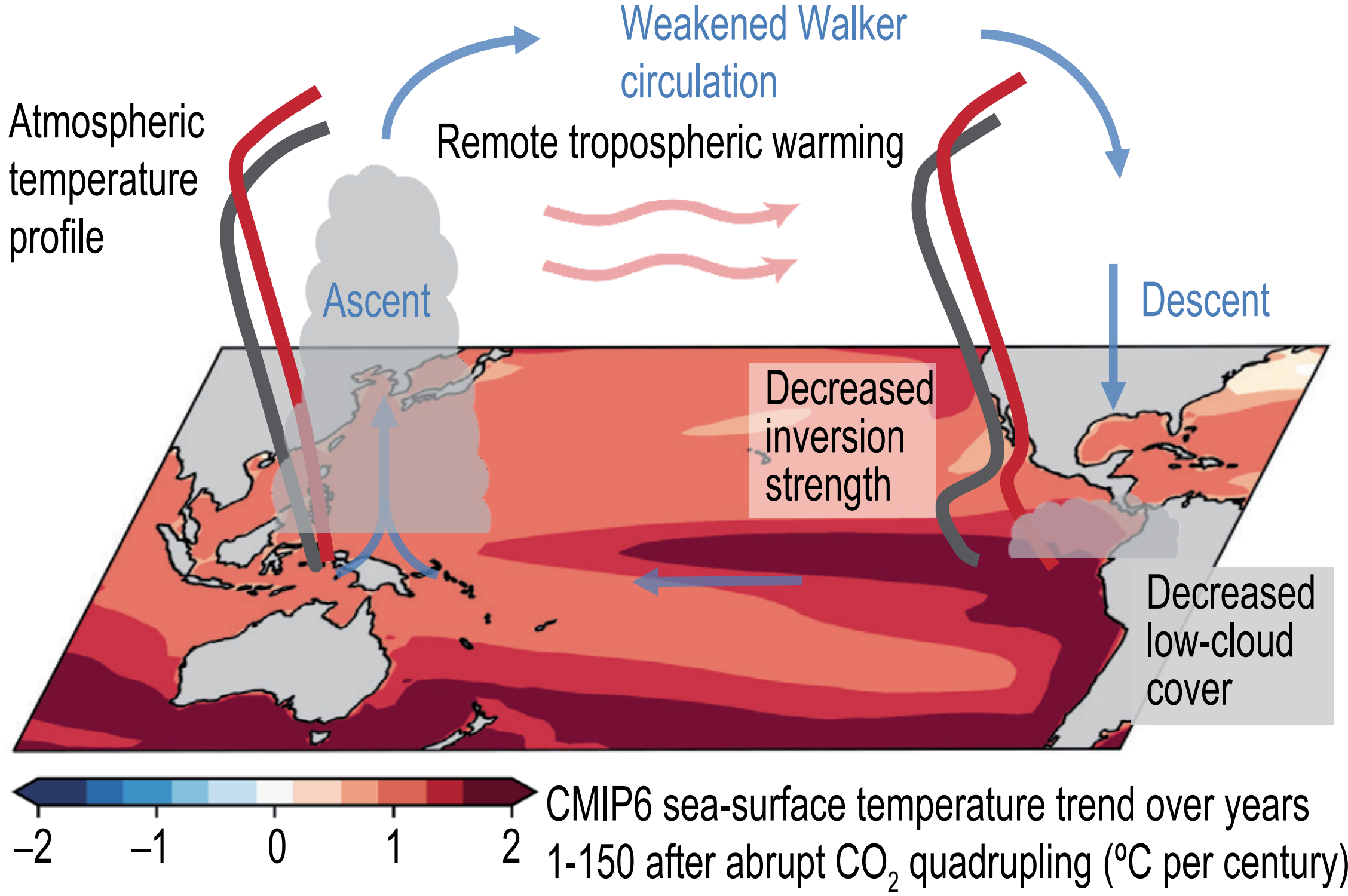
Positive Feedback

Observed pattern 1979-2013



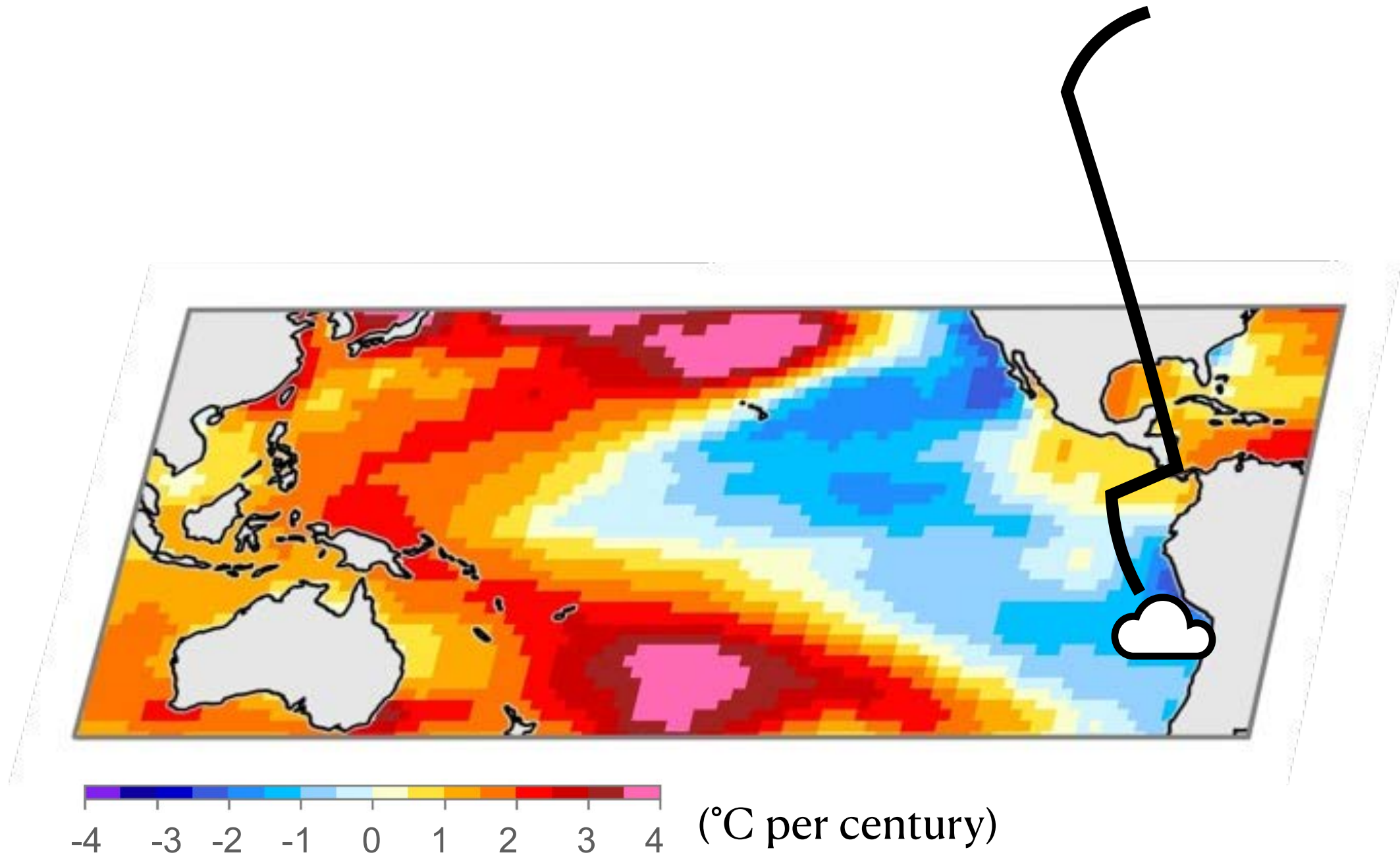
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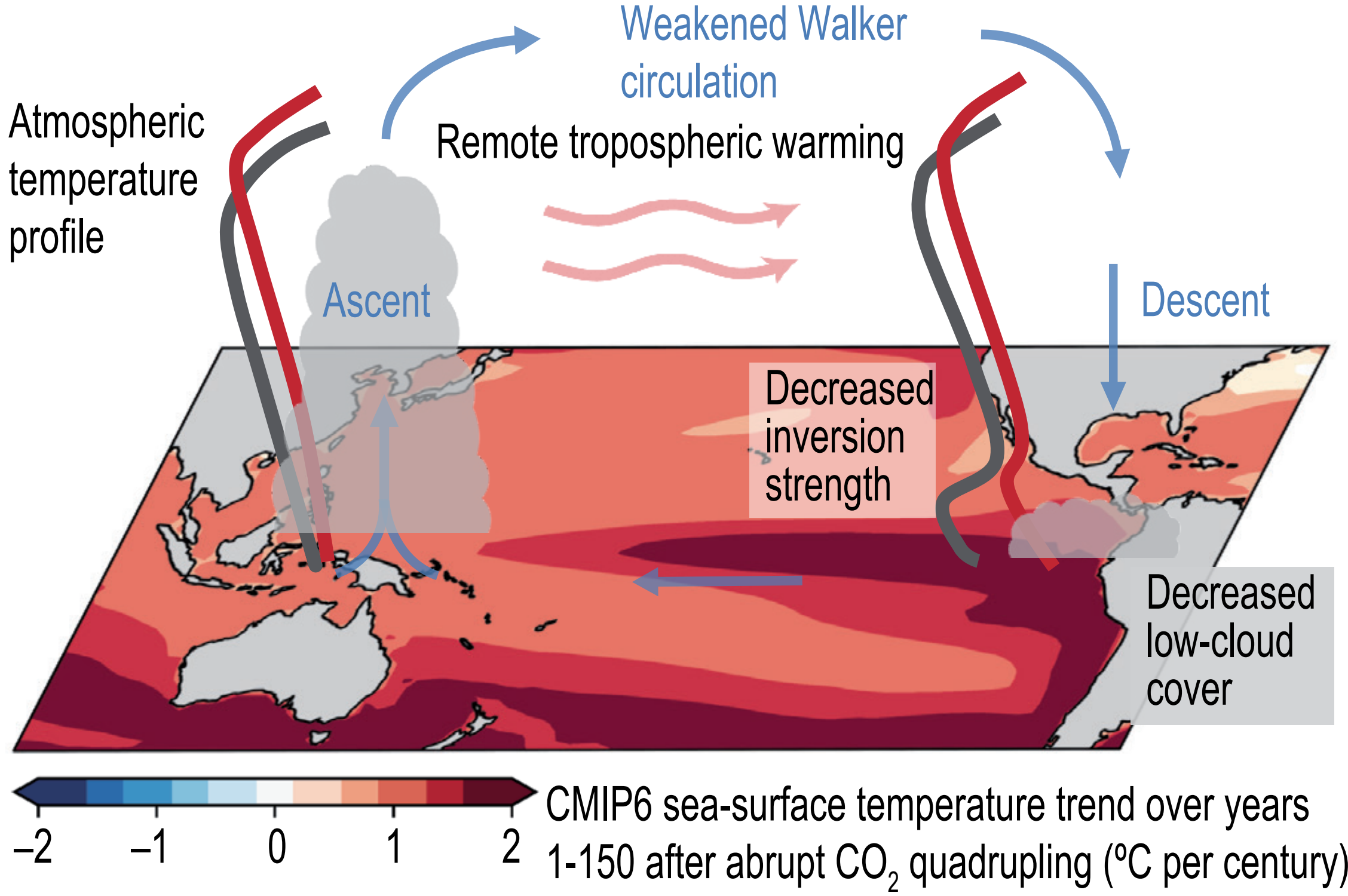
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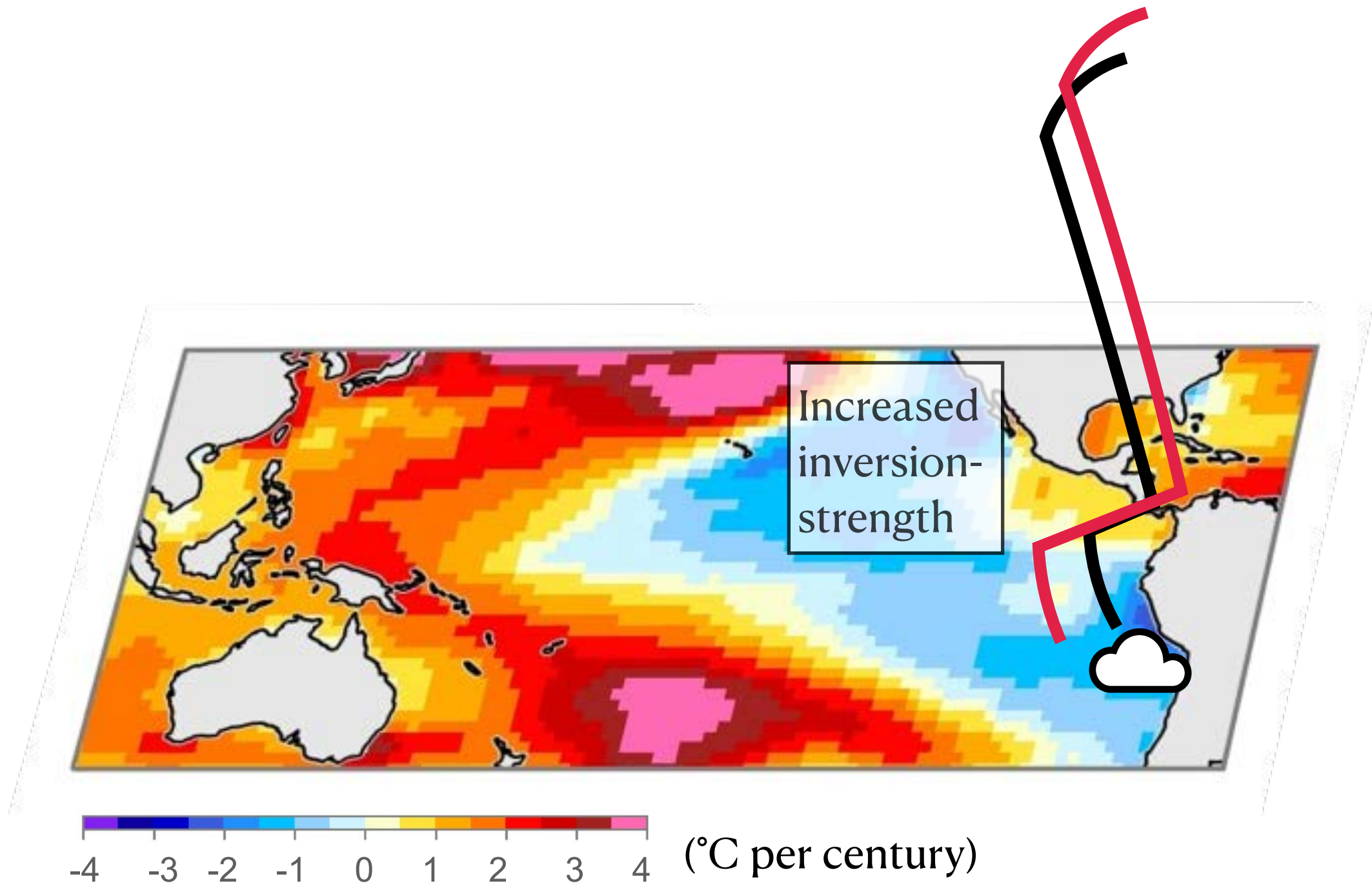
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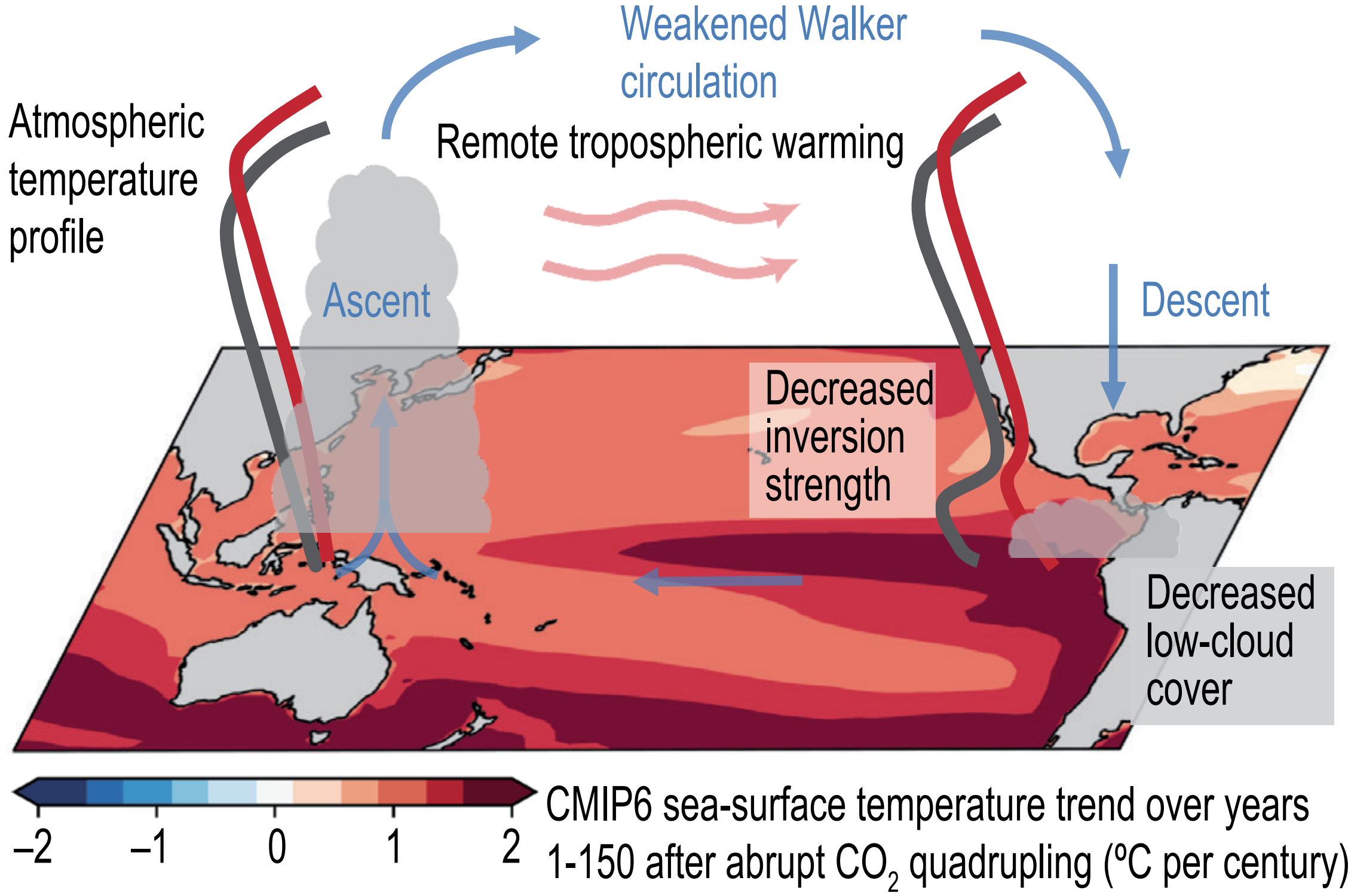
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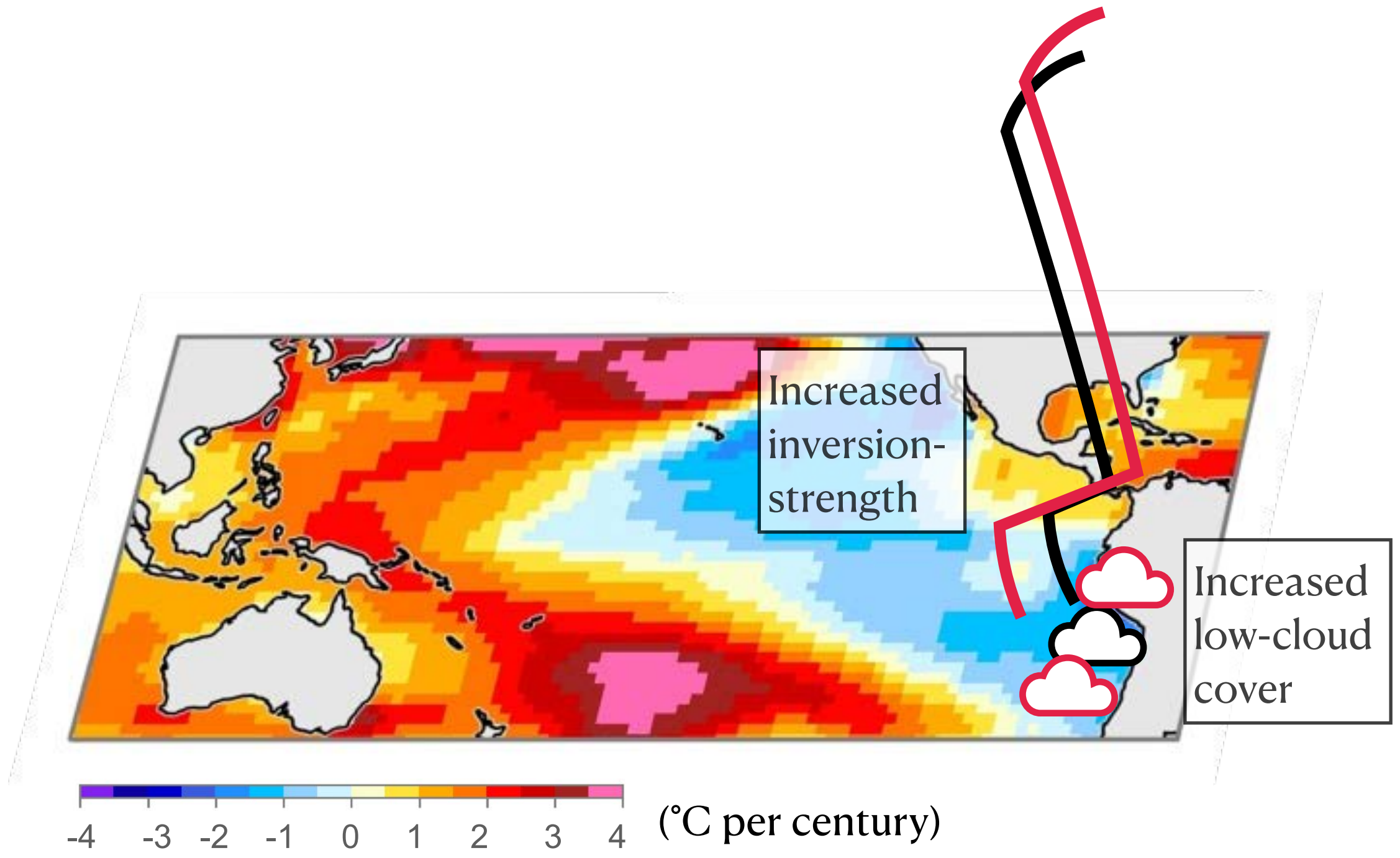
SO-induced teleconnection can affect the tropical warming patterns and climate sensitivity

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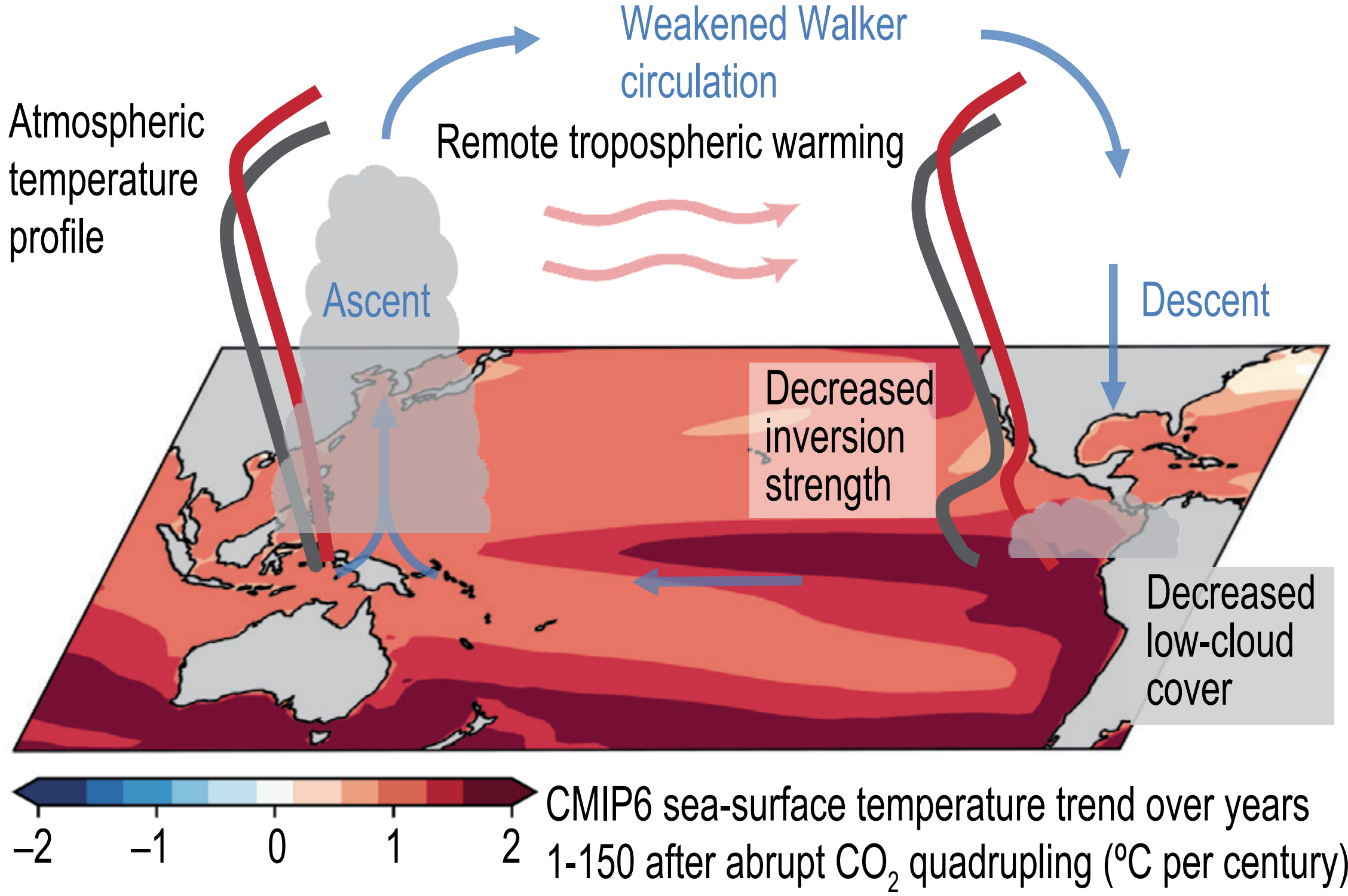
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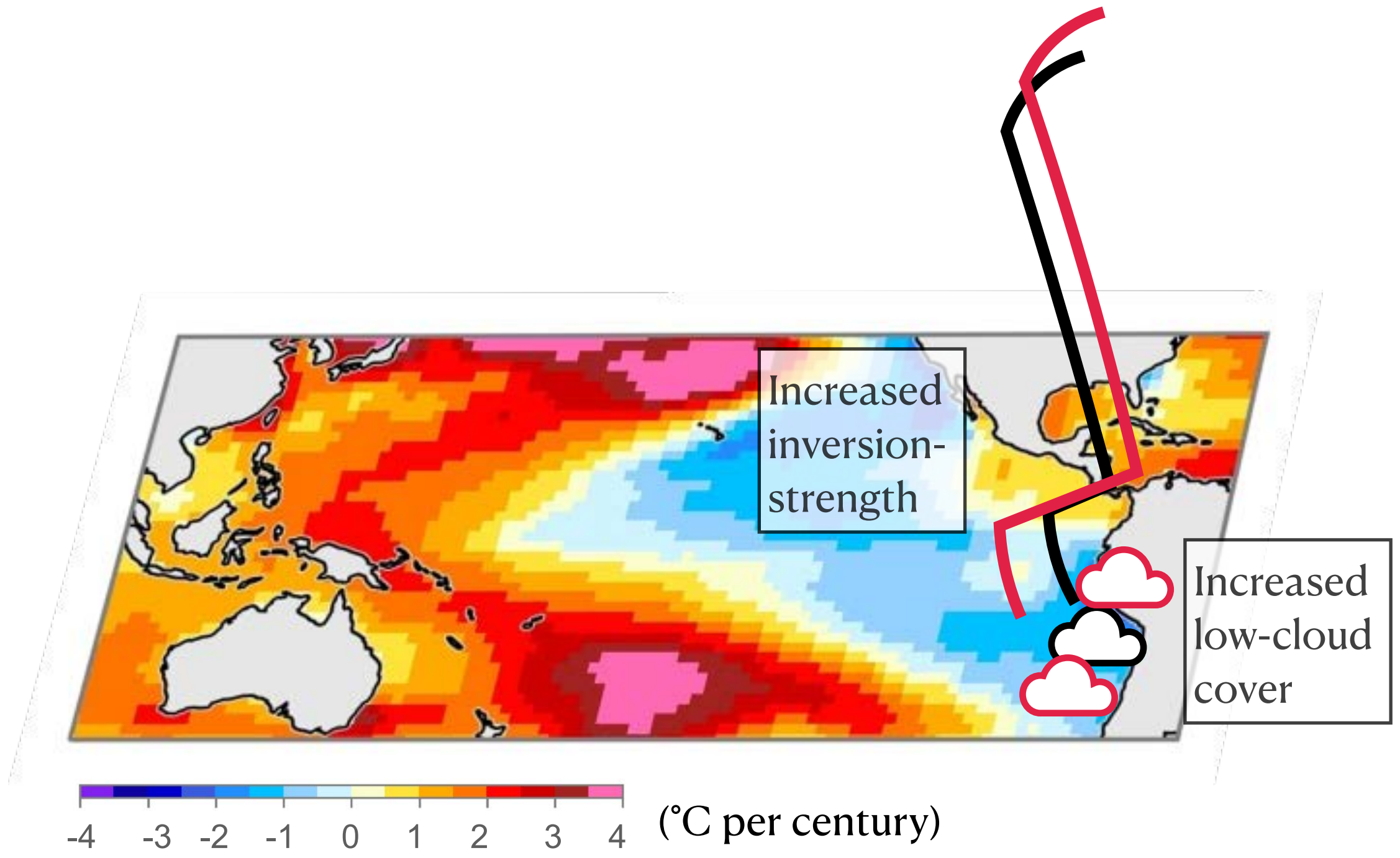
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Positive Feedback

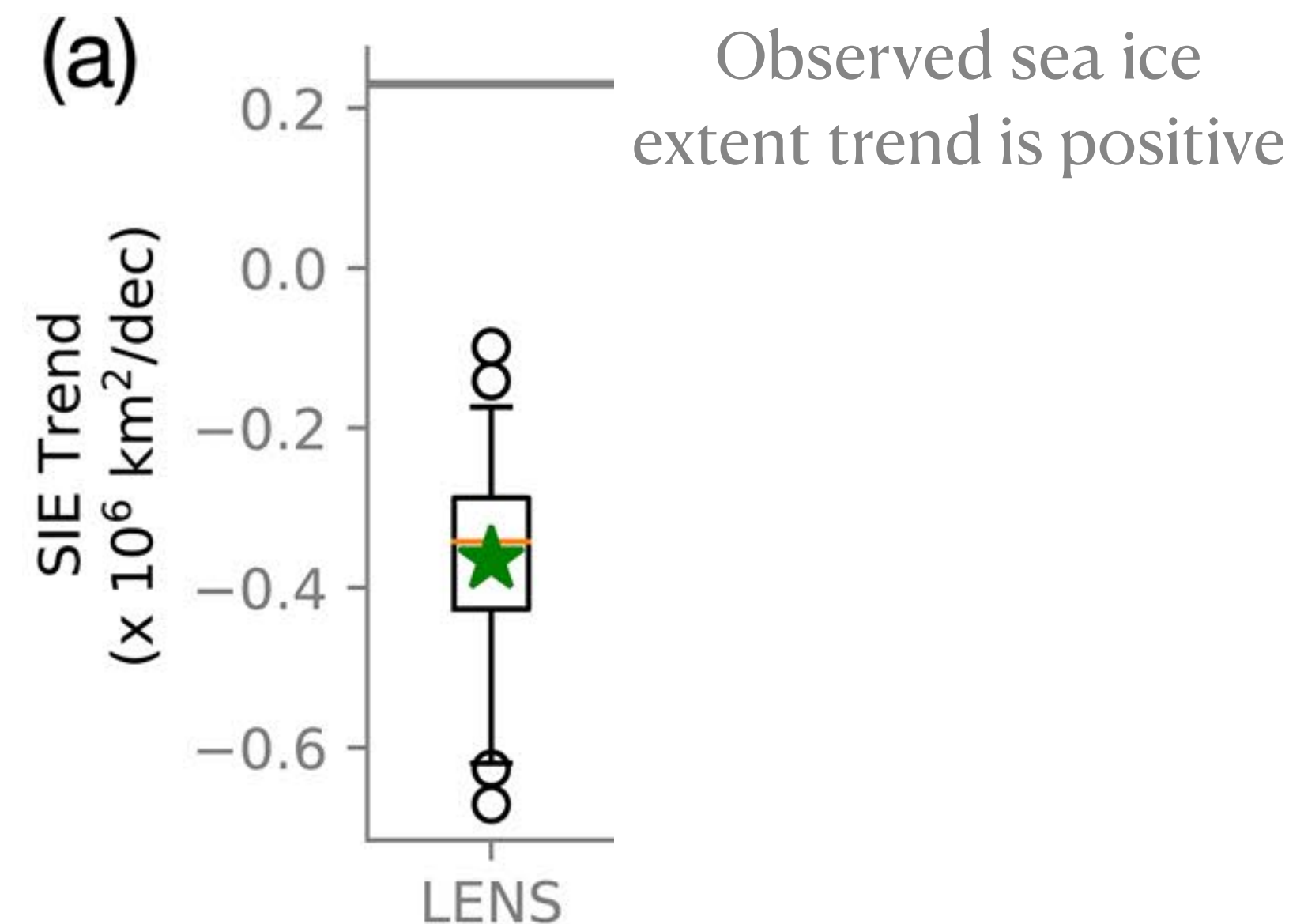
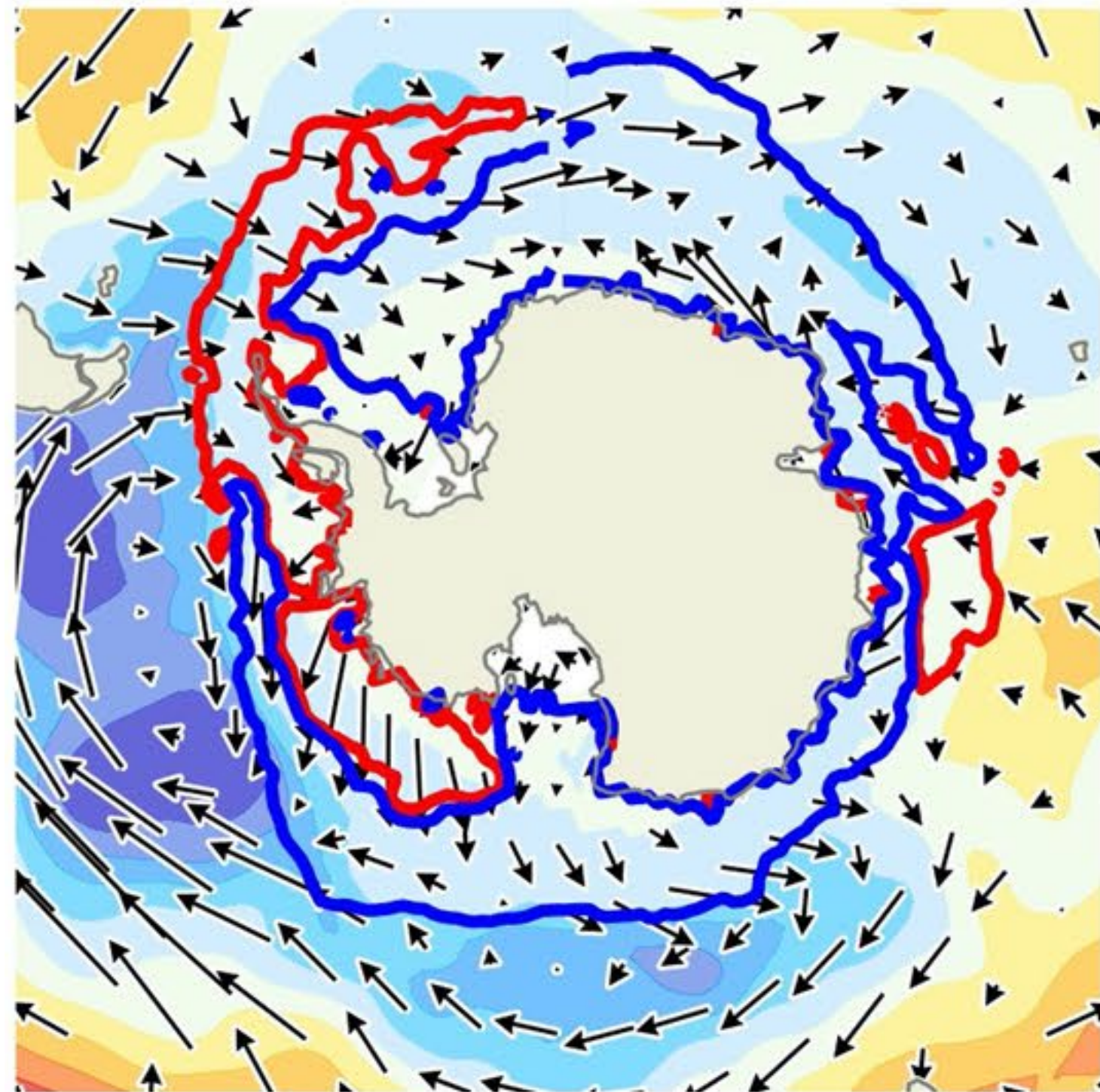
Observed pattern 1979-2013



Negative Feedback

SO surface cooling may explain model's inability to simulate observed Antarctic sea ice expansion

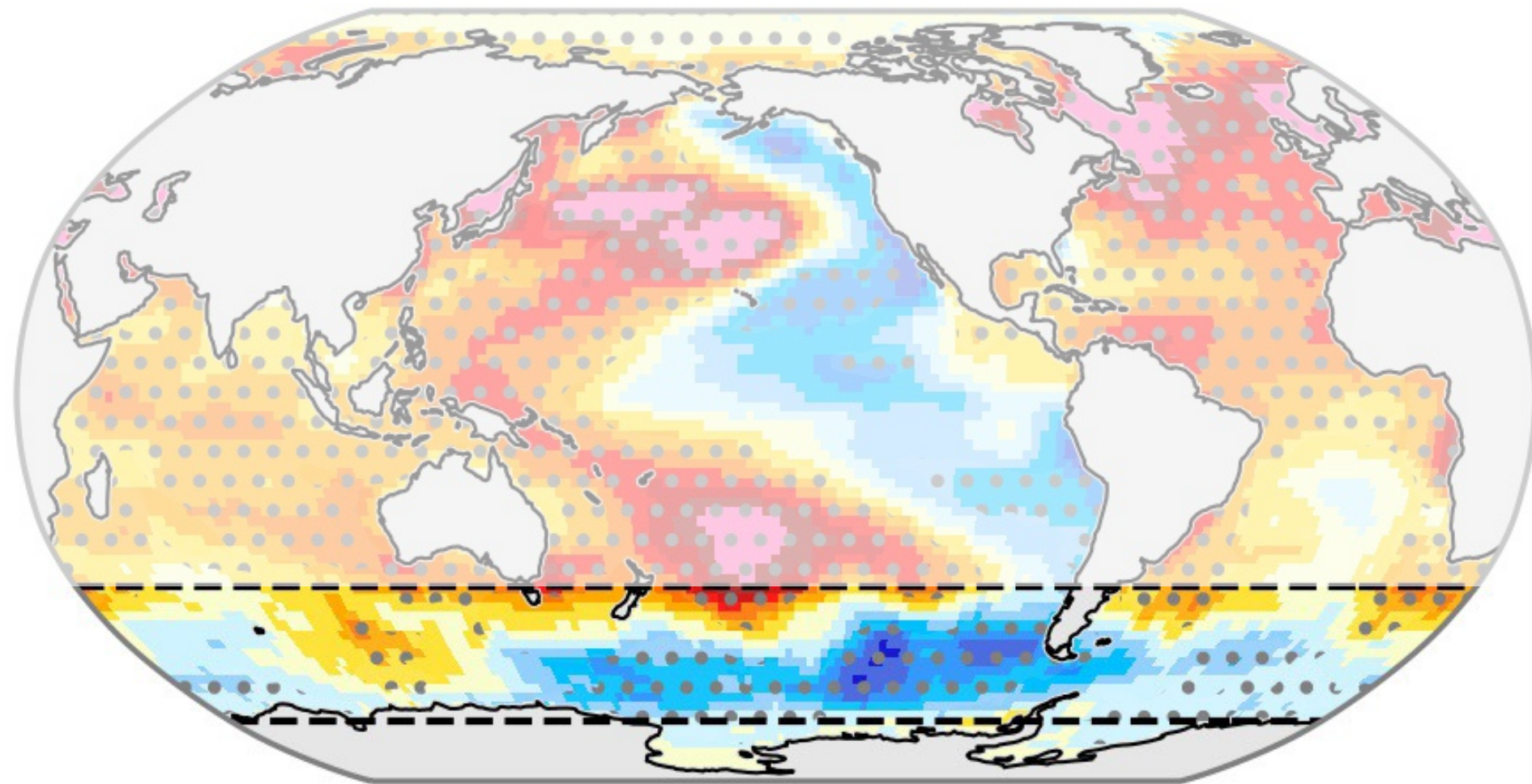
Observed SST (colors), wind (vectors), and sea ice (contours) trends



Using Southern Ocean pacemaker experiment (SOPACE) to study SO teleconnection

By including observed SO SST evolution in historical simulations

ERSSTv3b 1979-2013 trends

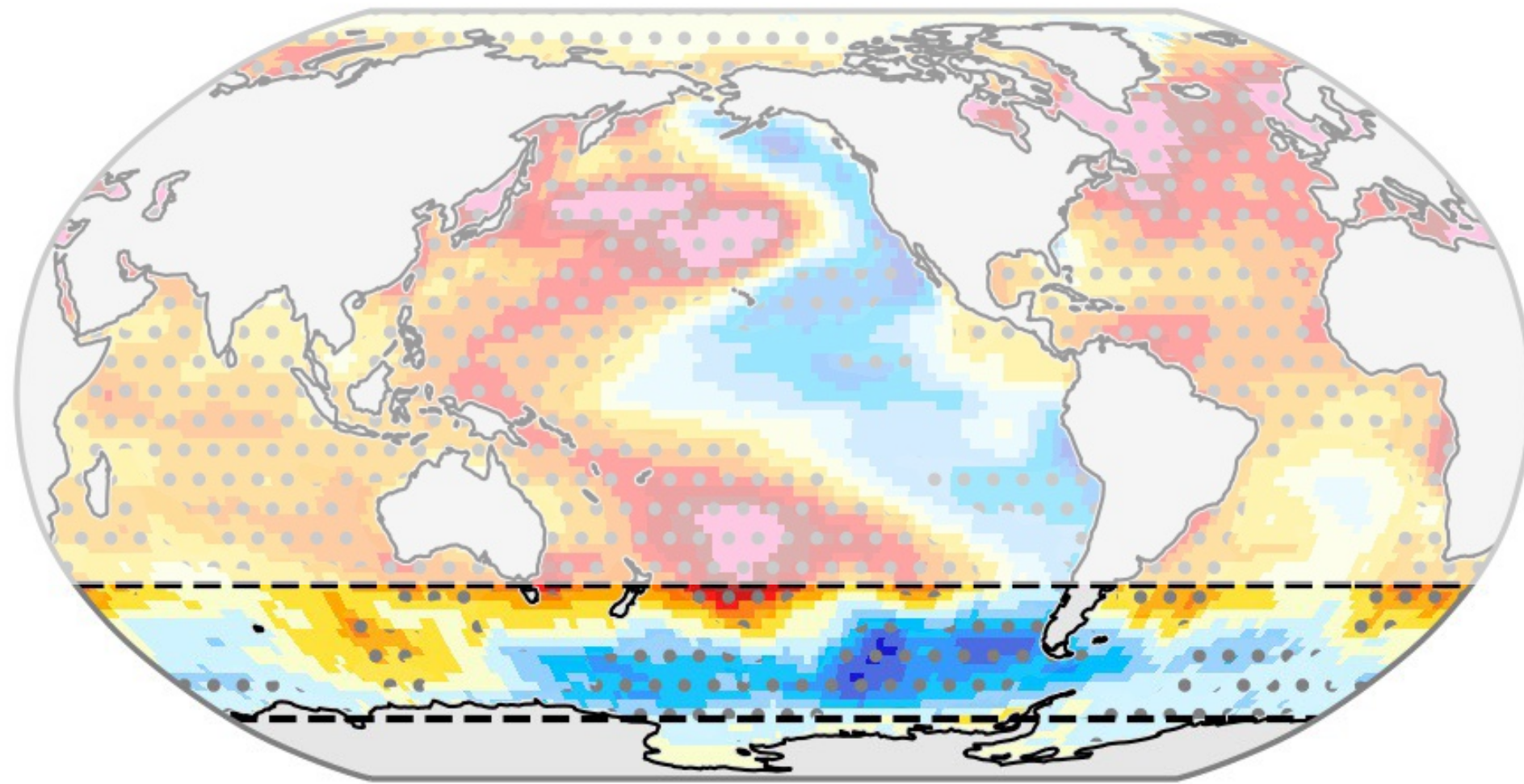


We nudge SST at each grid point in SO to **observed monthly anomaly** + **CESM1 mean state**.

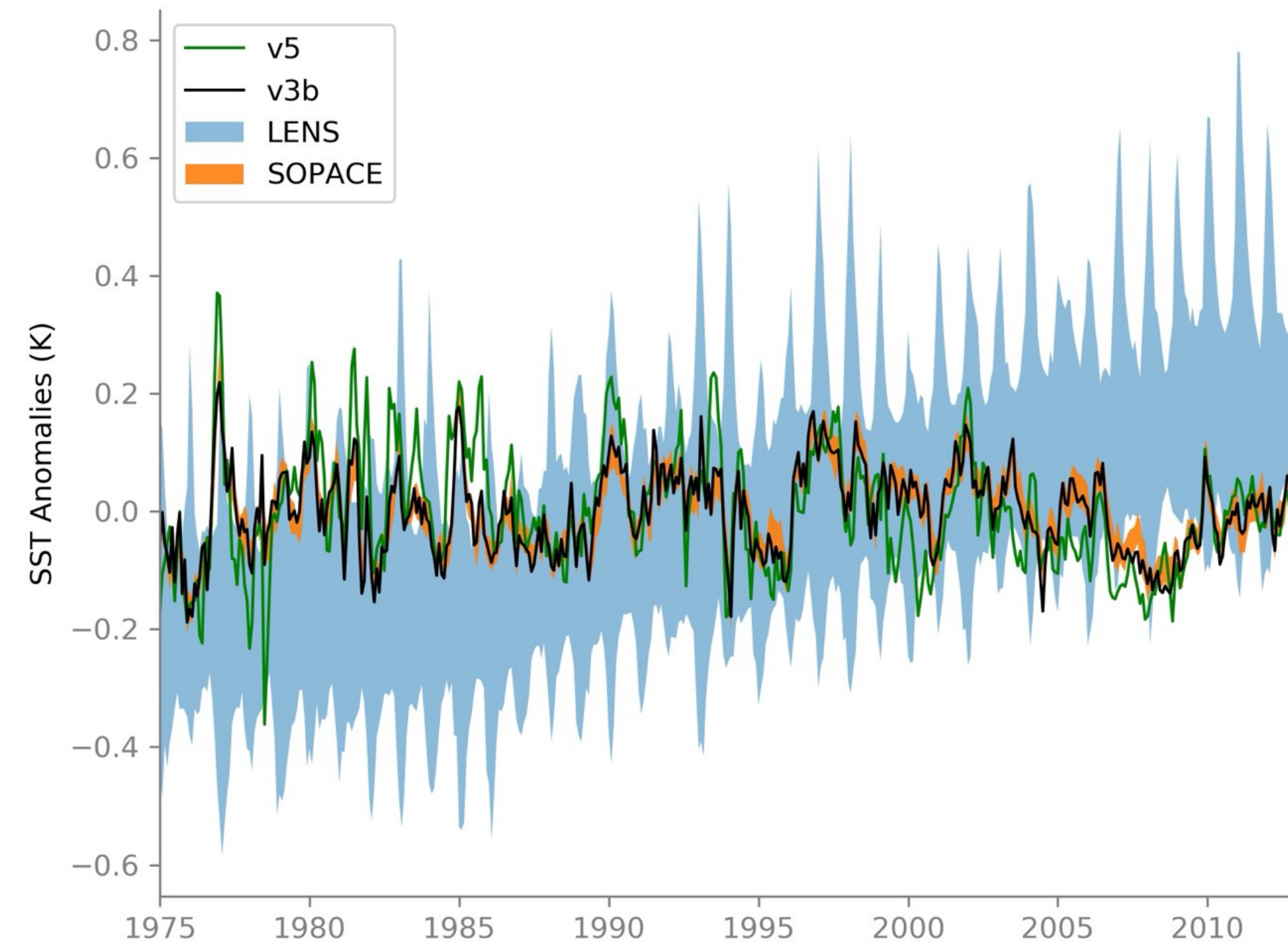
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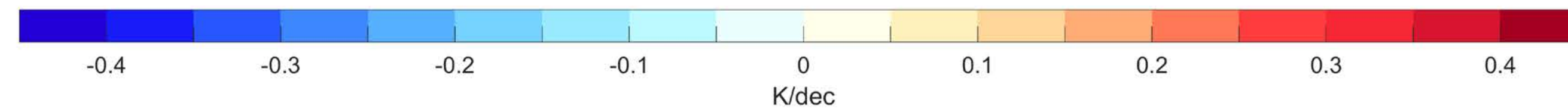
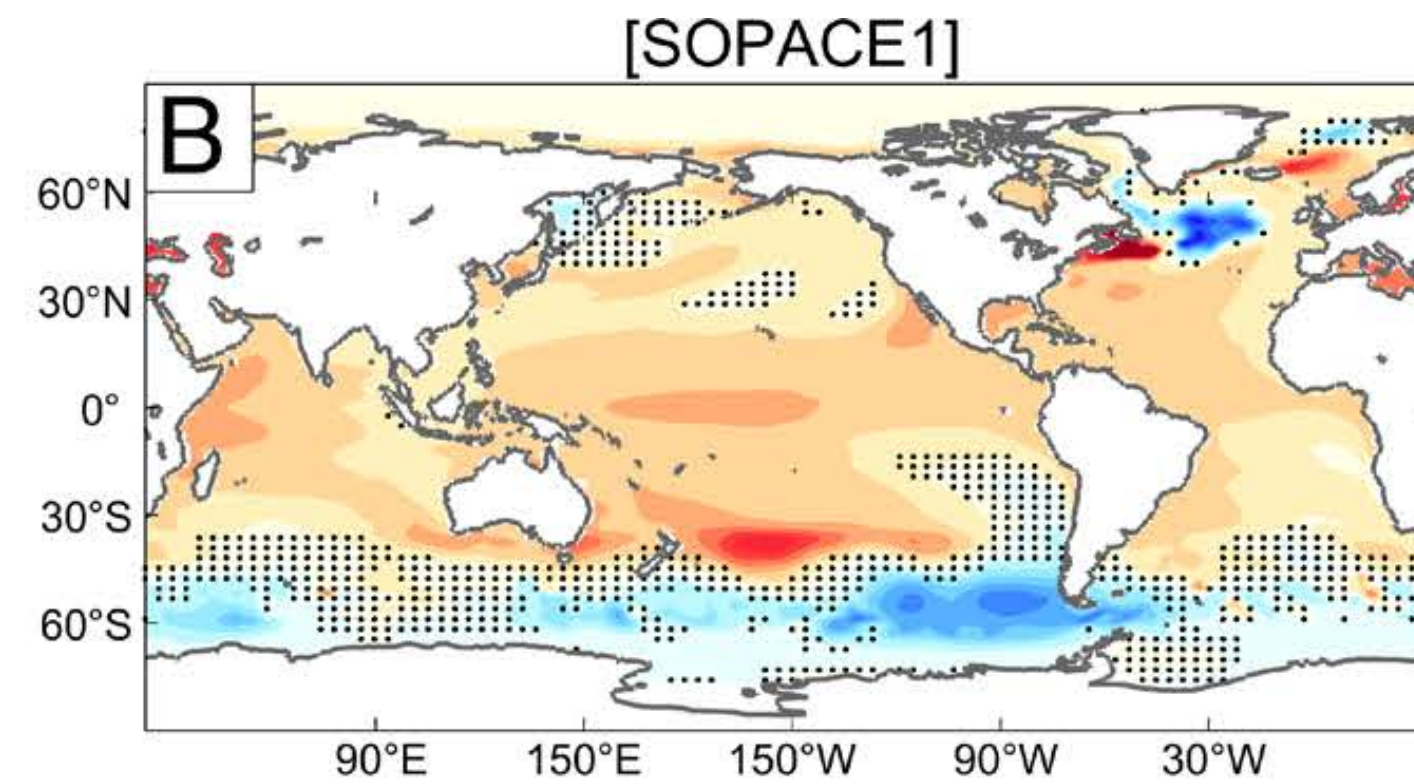
SO averaged SST anomalies



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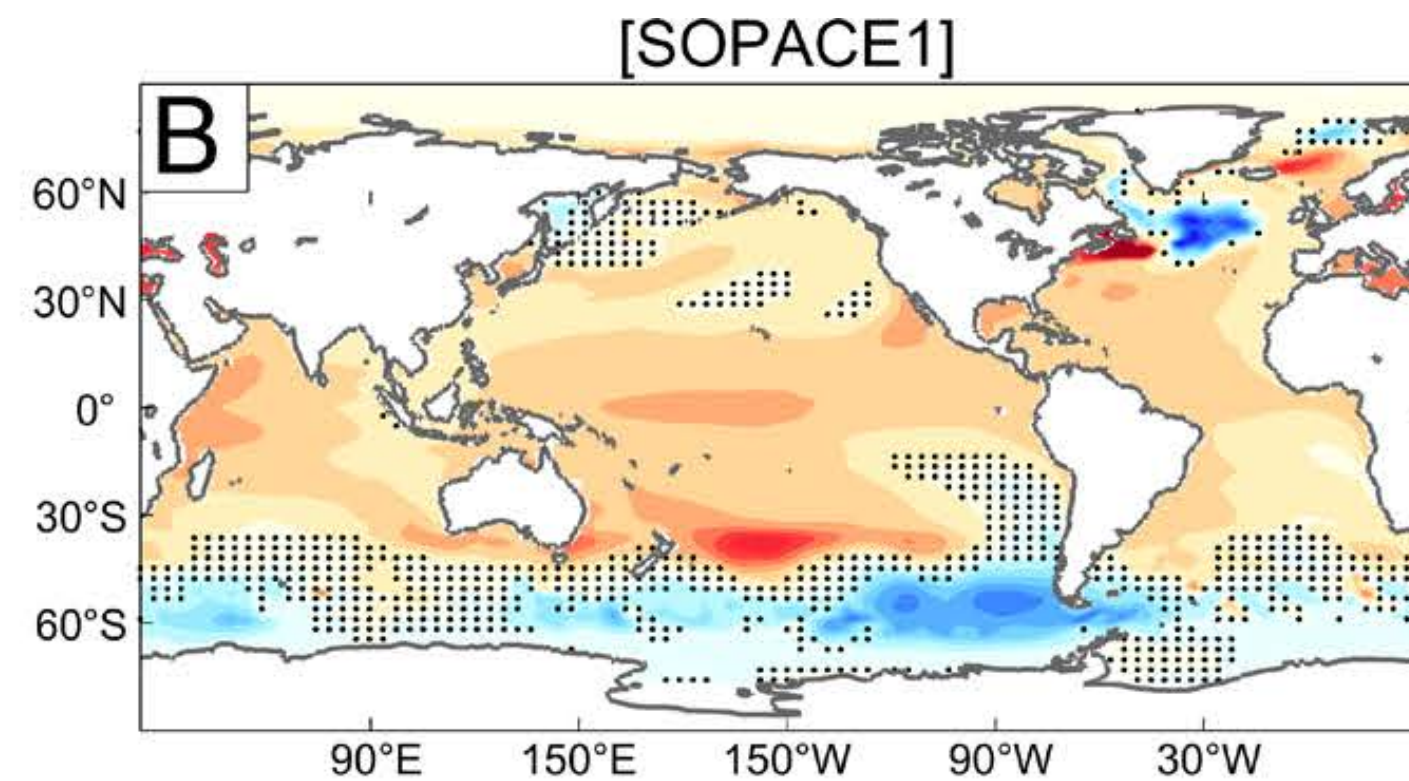
Global SST response to observed SO cooling

Results from CESM1

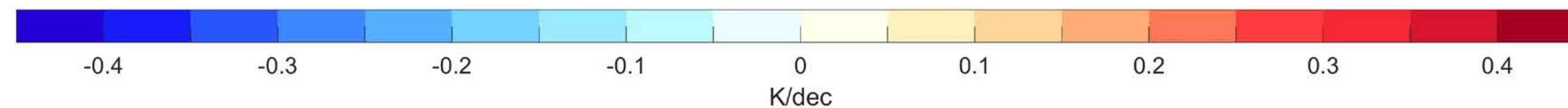


Global SST response to observed SO cooling

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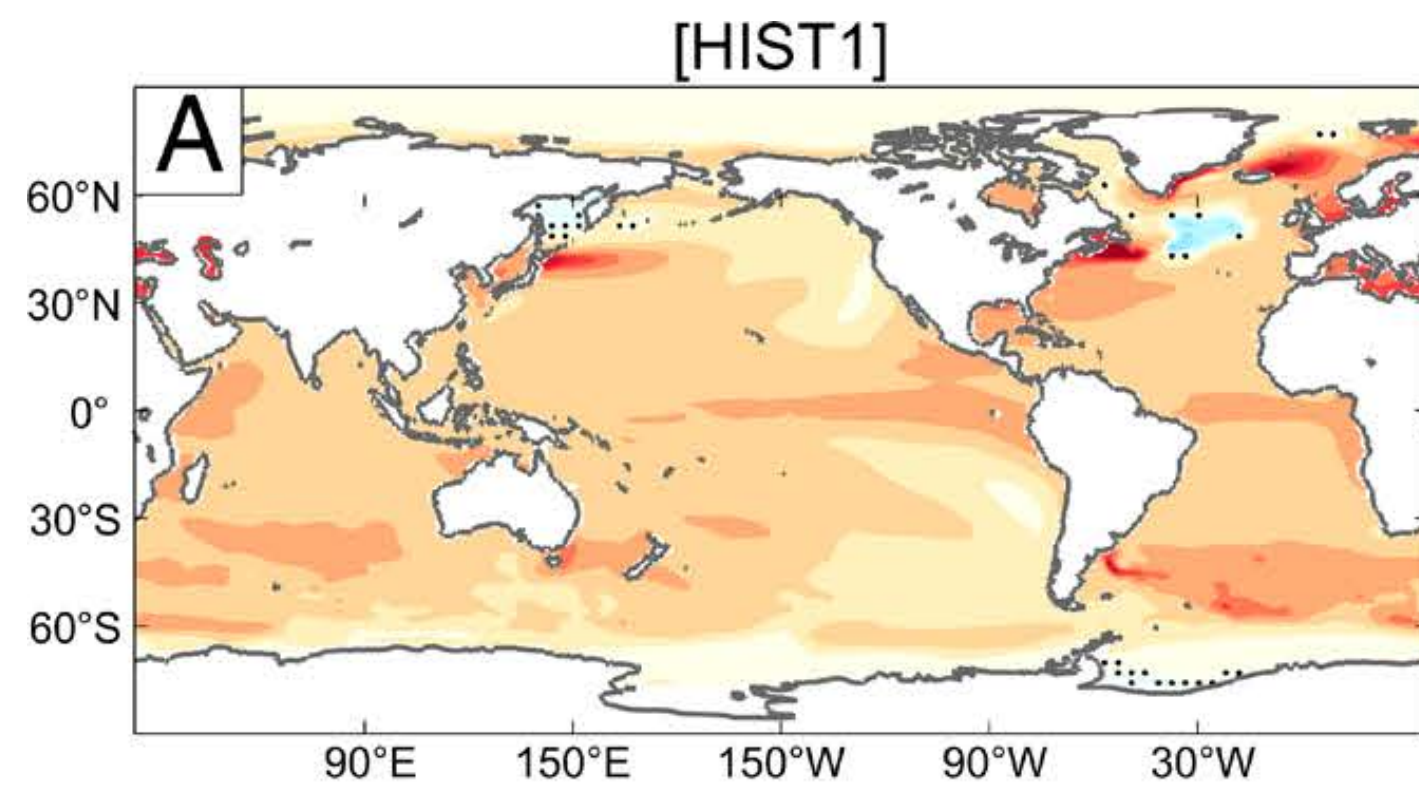


Radiatively-forced
response + SO-
driven response

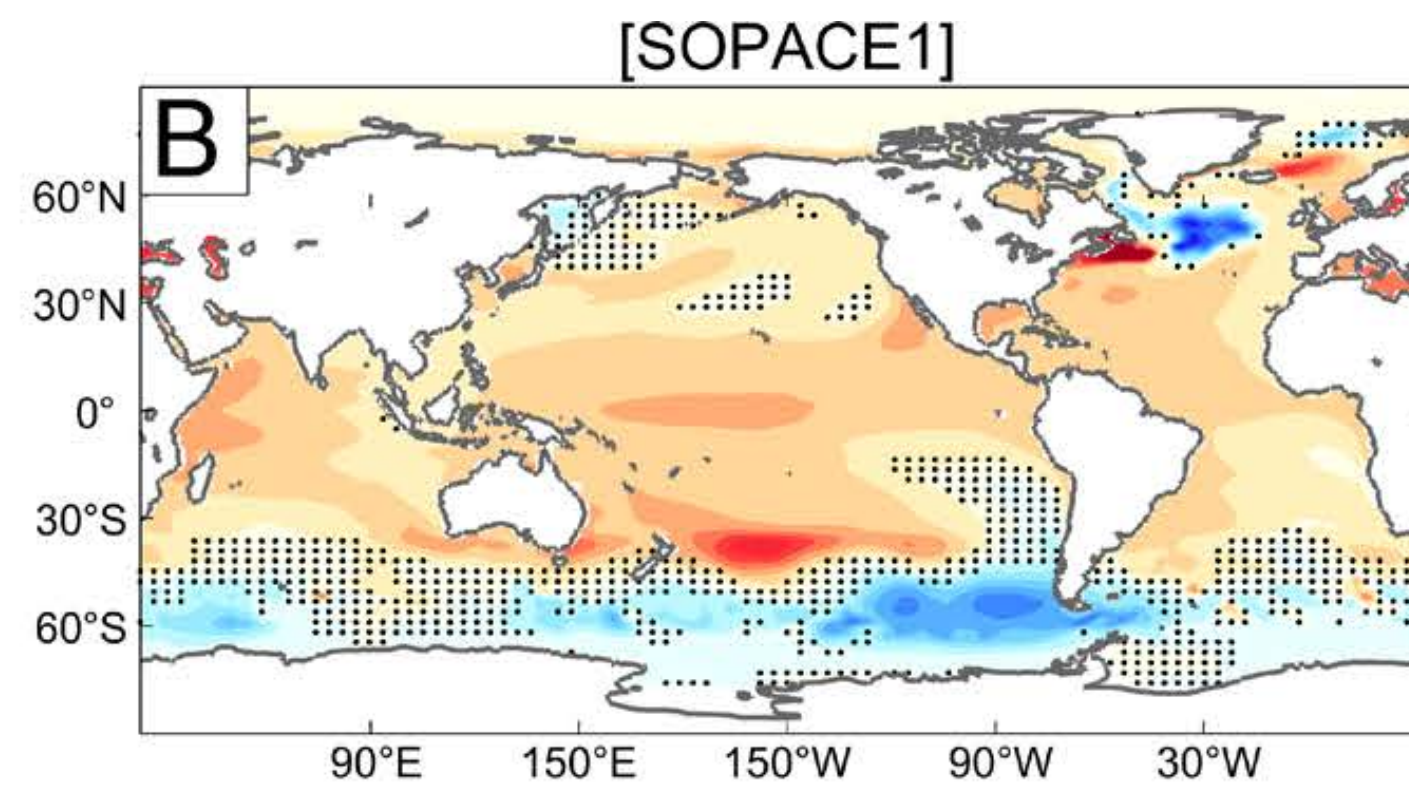


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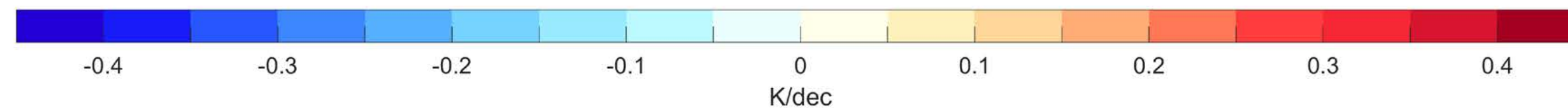
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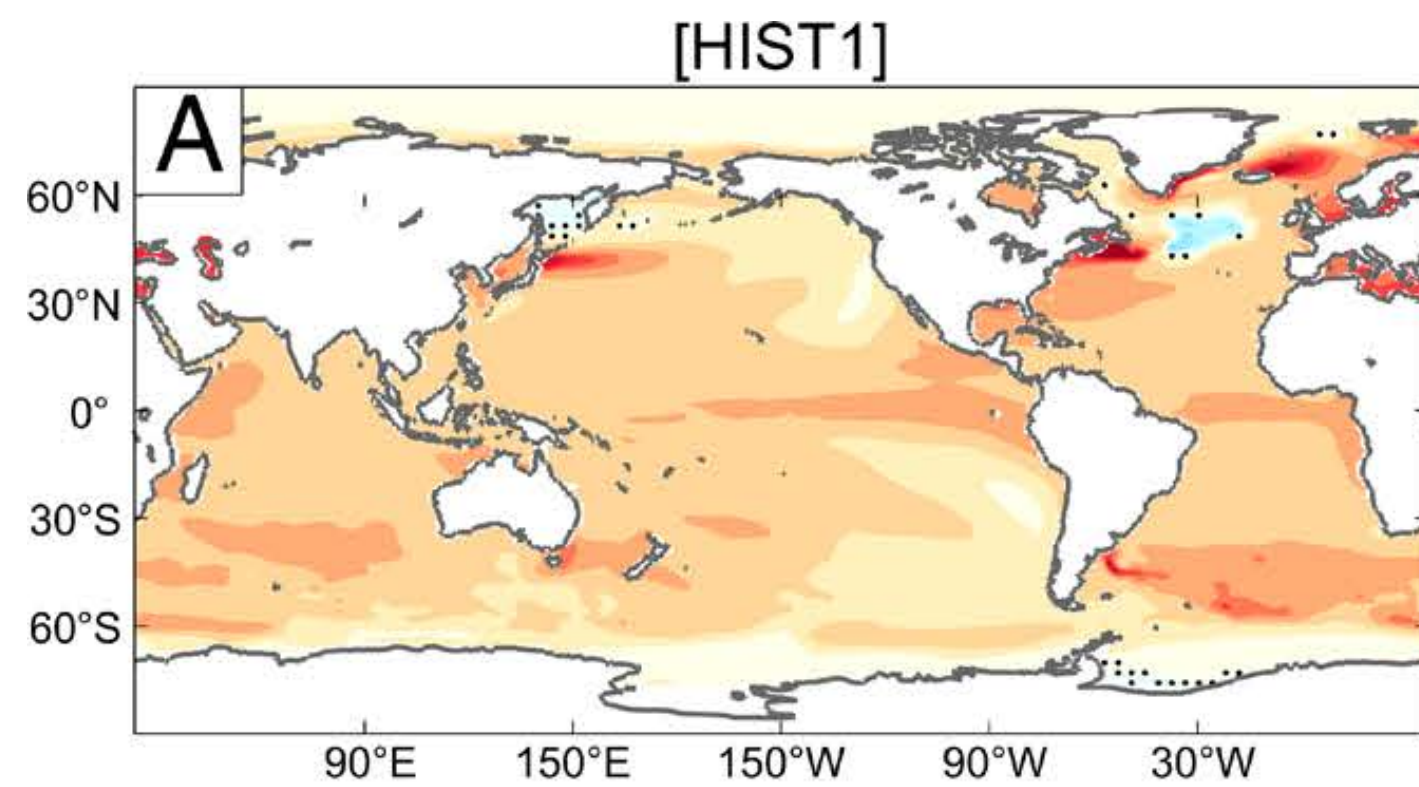


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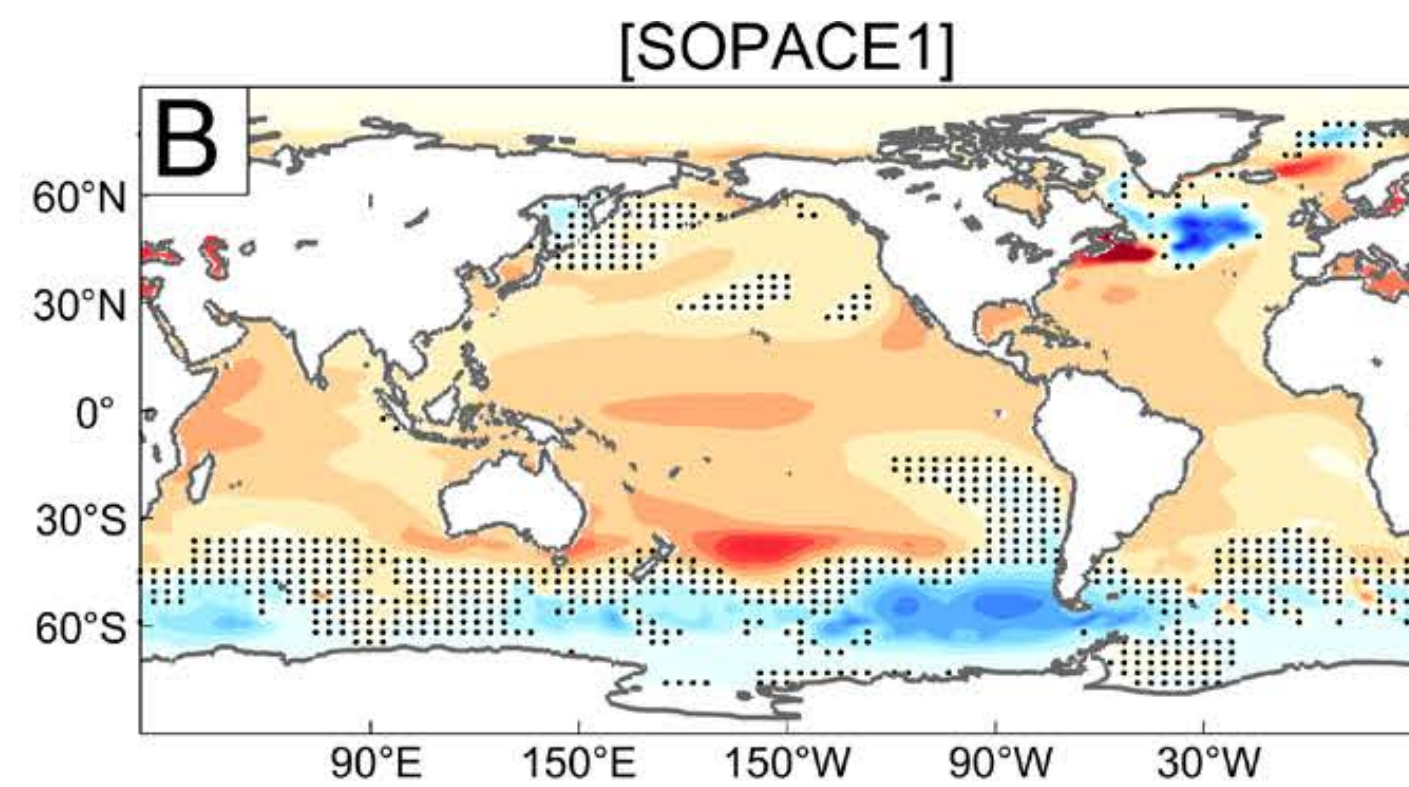


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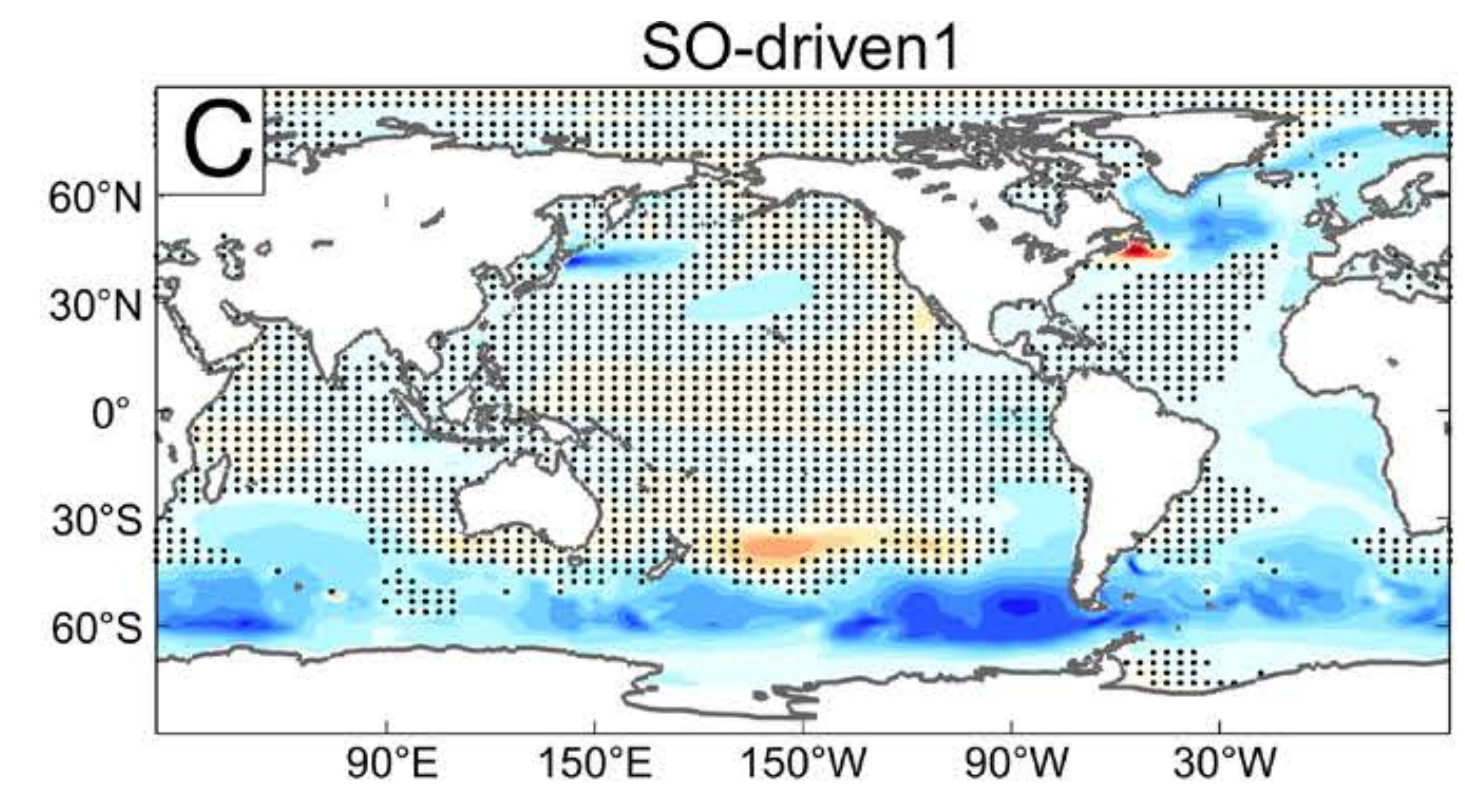
Results from CESM1



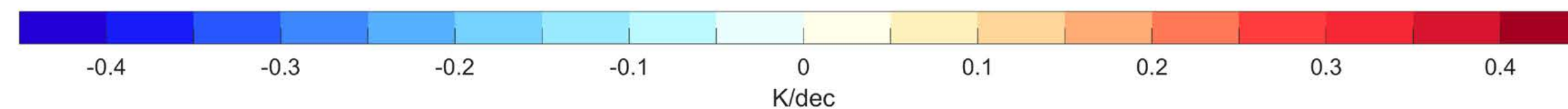
Radiatively-forced
response



Radiatively-forced
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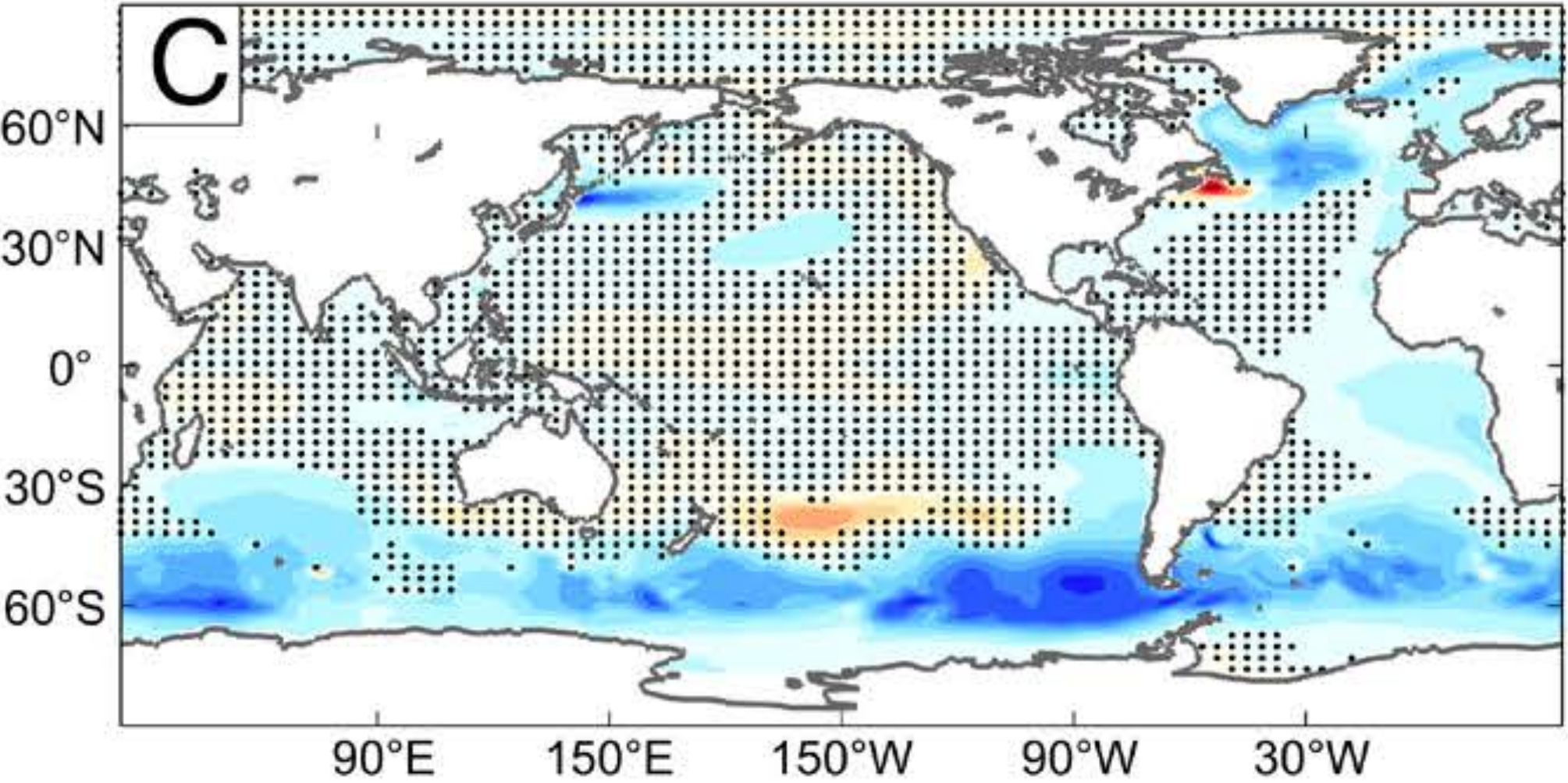


SO-driven
response

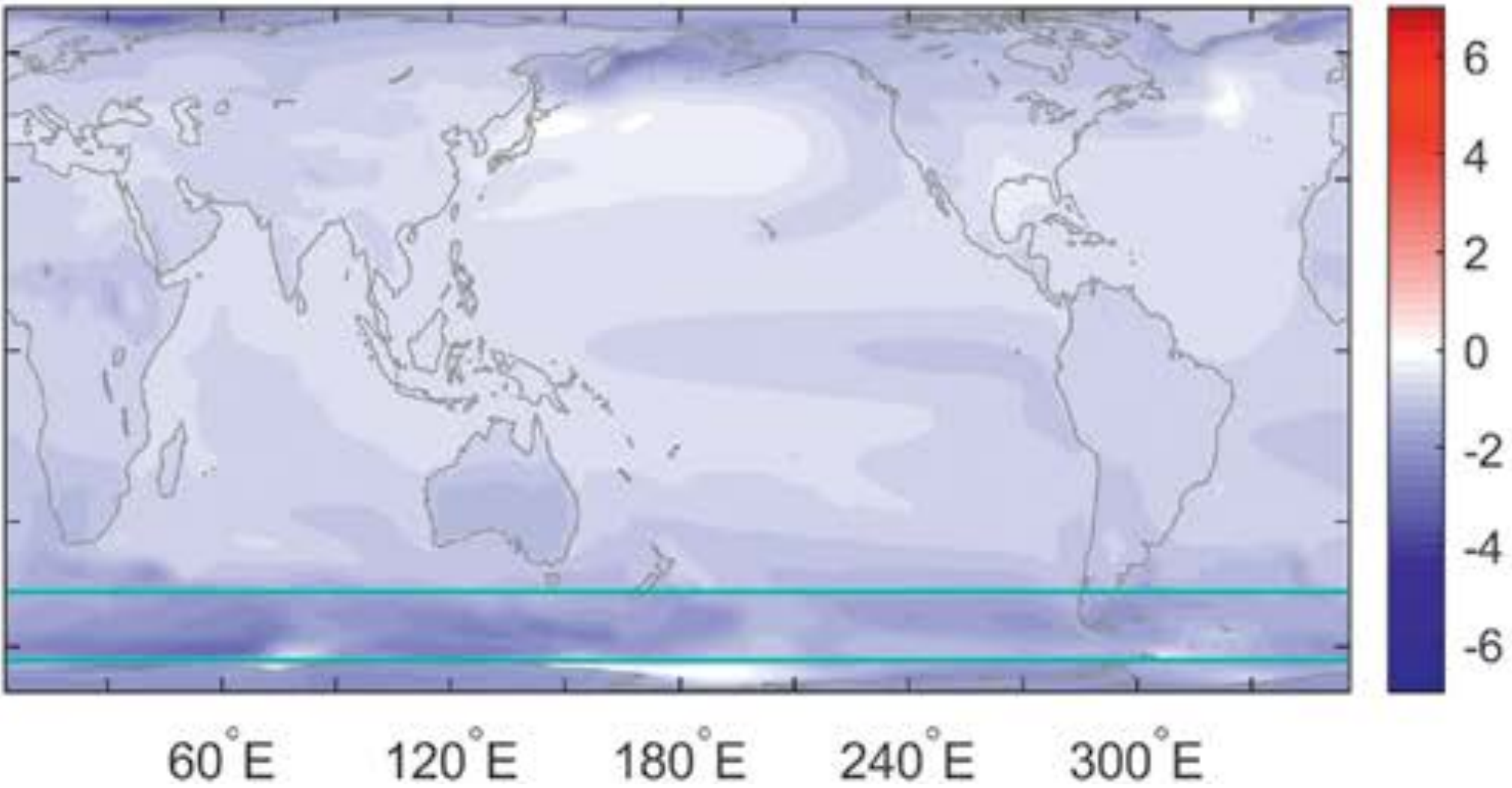


Why is the response so weak in the tropical Pacific?

Pacemaker experiment



Idealized experiment



CESM1 has a subtropical low cloud bias in the Pacific

Subtropical stratocumulus cloud feedback west of South America

CMIP version	Model name	Institution	CF _{WSA}
CMIP6	CESM2-WACCM	NCAR, USA	11.95
CMIP6	CESM2	NCAR, USA	11.36
CMIP6	CESM2-FV2	NCAR, USA	9.31
CMIP6	NorESM2-MM	NCC, Norway	8.49
CMIP6	CESM2-WACCM-FV2	NCAR, USA	8.15
ISCCP-FH & OISST 2 (observational data)			7.51
CERES-EBAF 4.1 & OISST 2 (observational data)			7.30
CMIP5	IPSL-CM5B-LR	IPSL, France	7.11
CMIP6	NorESM2-LM	NCC, Norway	6.84
CMIP6	KIOST-ESM	KIOST, South Korea	6.82
CMIP5	ACCESS1-0	CSIRO and BOM, Australia	6.32
CMIP5	MIROC-ESM	AORI, NIES, and JAMSTEC, Japan	6.10
CMIP5	CSIRO-Mk3-6-0	CSIRO and BOM, Australia	6.09
CMIP5	MIROC-ESM-CHEM	AORI, NIES, and JAMSTEC, Japan	6.09
CMIP6	GFDL-CM4	GFDL, USA	5.92
...			
CMIP5	CESM1-FASTCHEM	NCAR, USA	3.57
CMIP6	BCC-ESM1	BCC, china	3.45
CMIP6	AWI-CM-1-1-MR	AWI, Germany	3.45
CMIP5	CCSM4	NCAR, USA	3.34
CMIP5	CESM1-BGC	NCAR, USA	3.27
CMIP6	EC-Earth3-Veg	EC-Earth consortium	3.19
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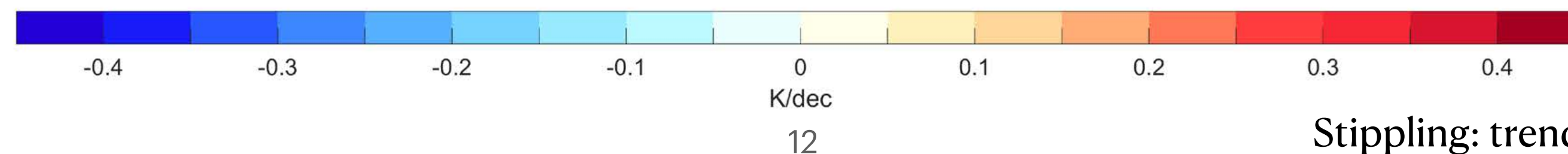
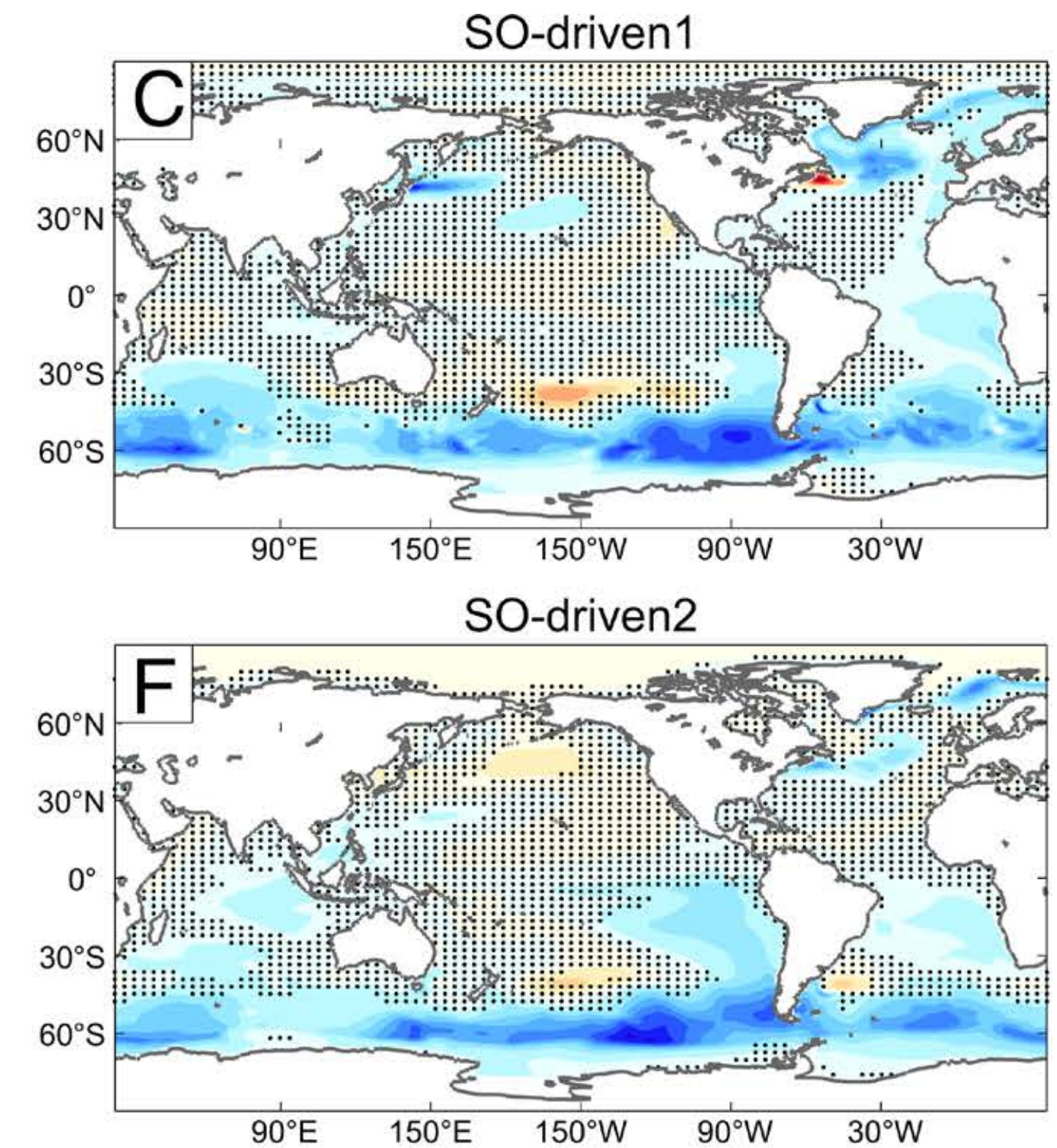
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- CESM1's stratocumulus cloud feedback is too weak
- CESM2 has much stronger stratocumulus cloud feedback
- Will SOPACE with CESM2 show a stronger tropical Pacific response?

Global SST response to observed SO cooling

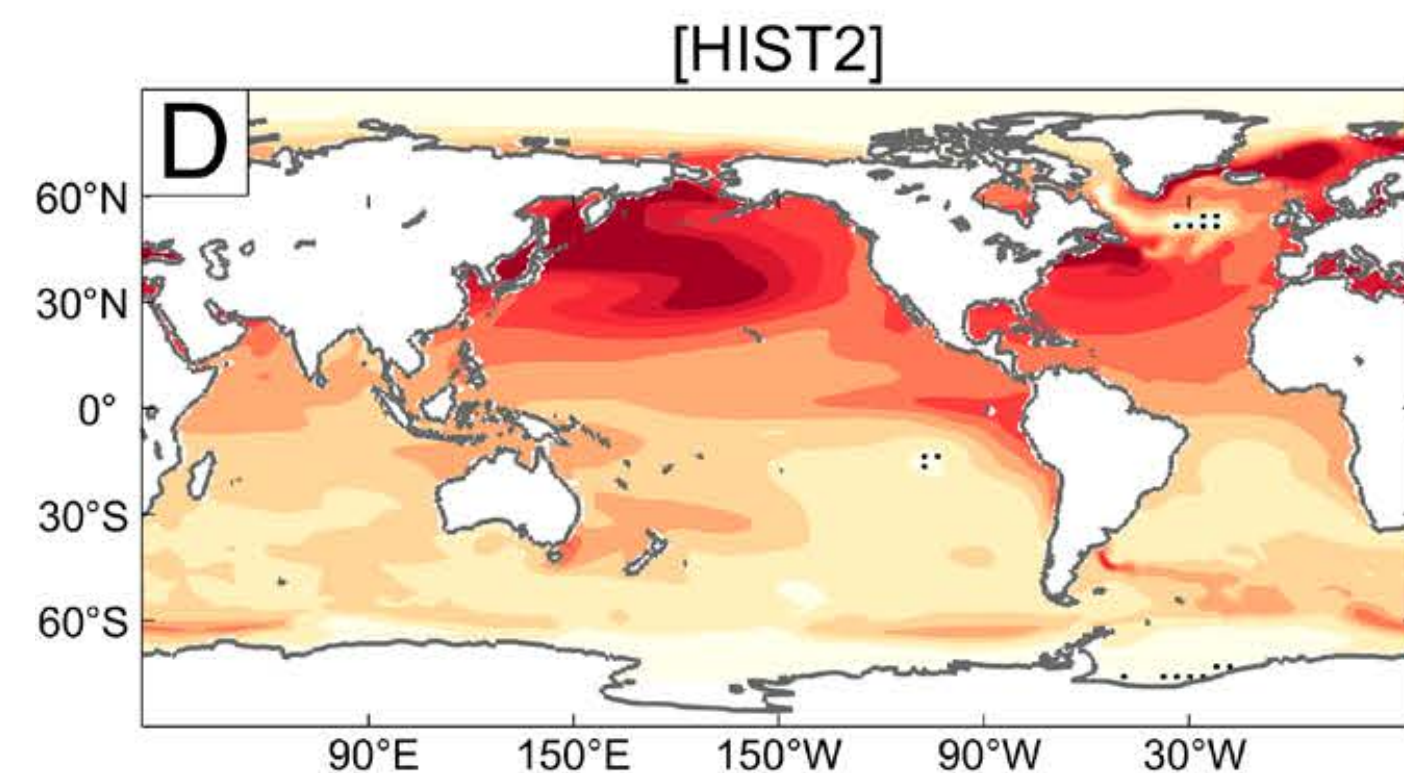
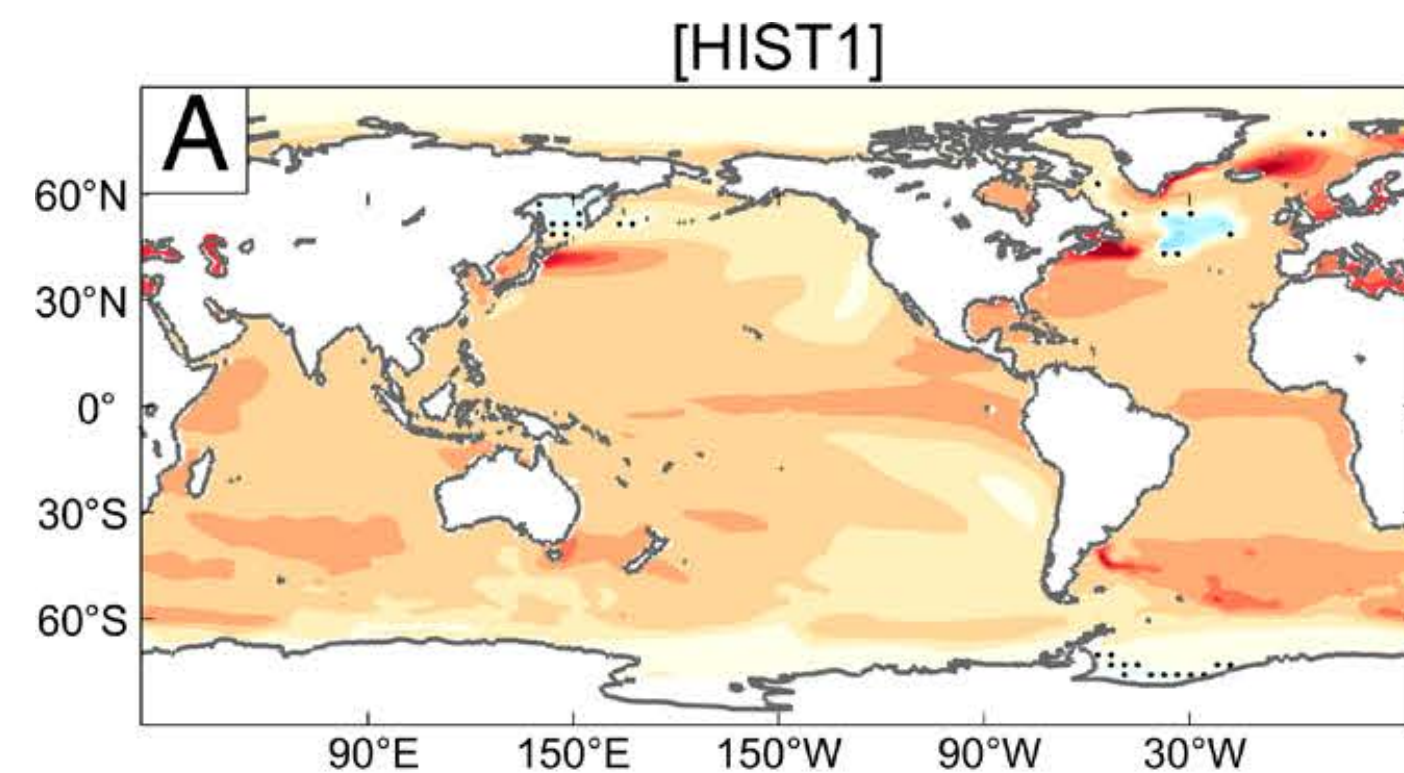
CESM1 vs CESM2



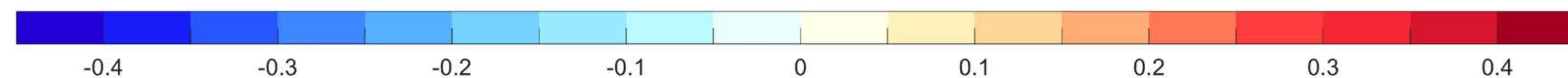
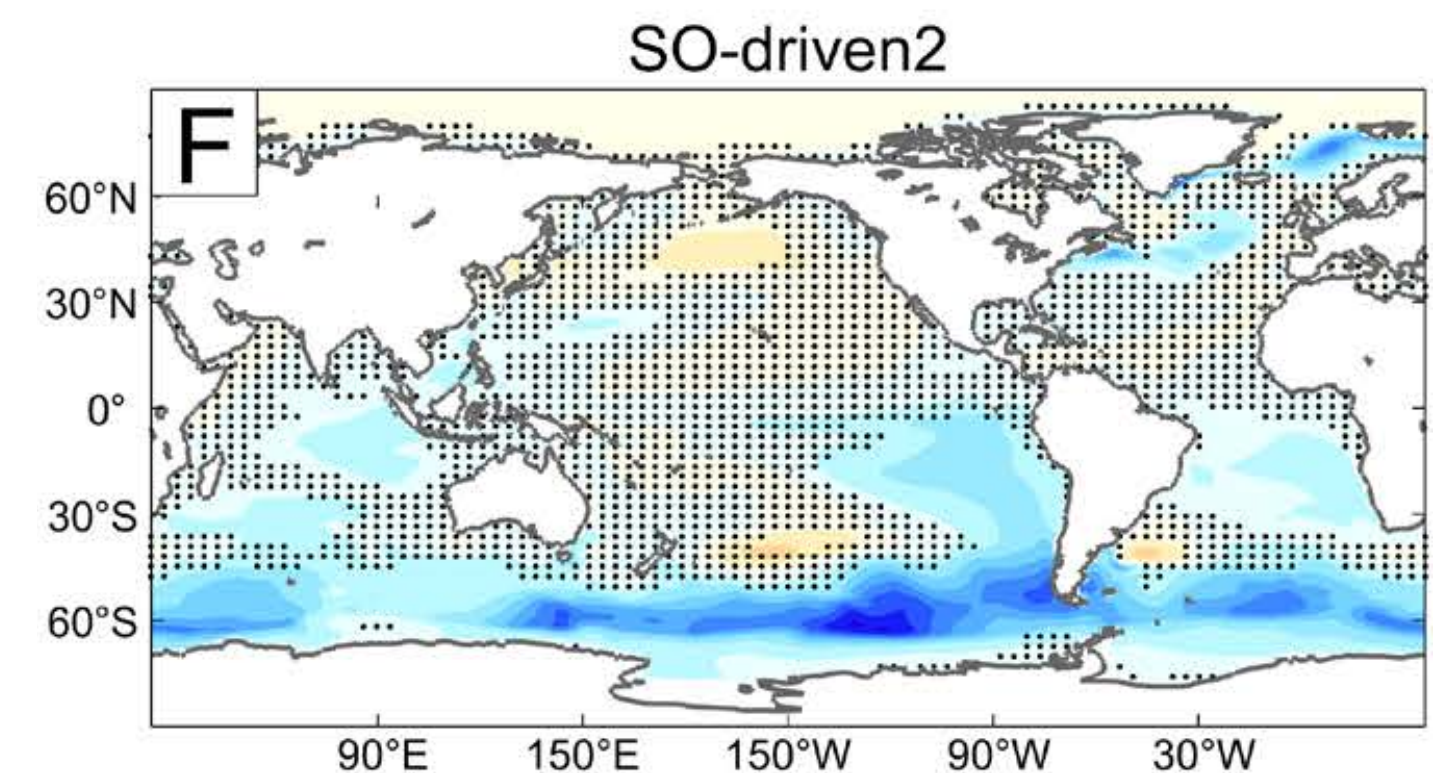
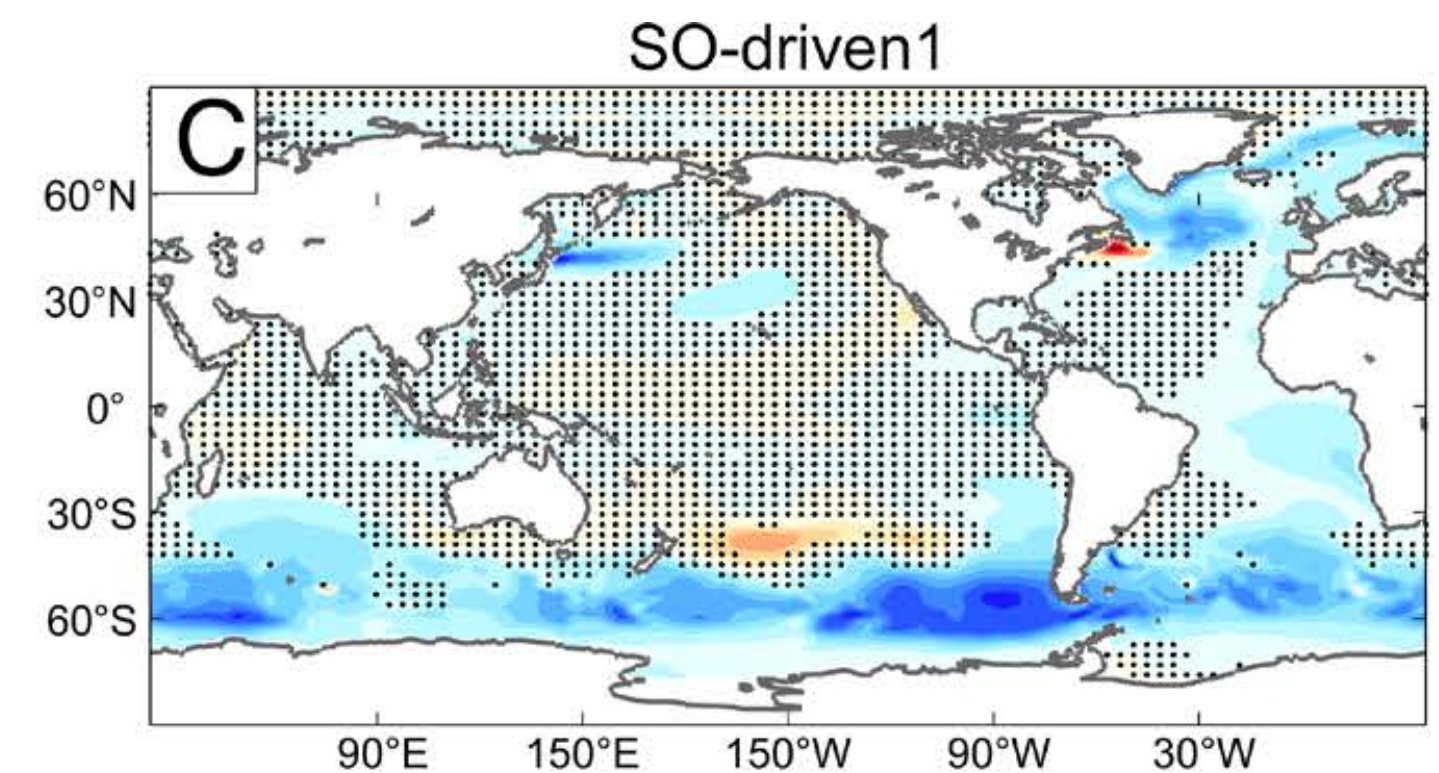
Stippling: trends NOT significant at 95% confidence level
Zhang et al. (2021); Kang et al. (2023)

Global SST response to observed SO cooling

CESM1 vs CESM2



Radiatively-forced only



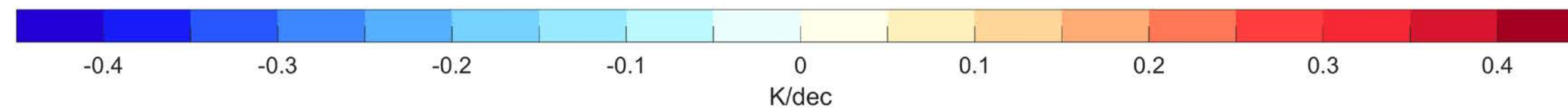
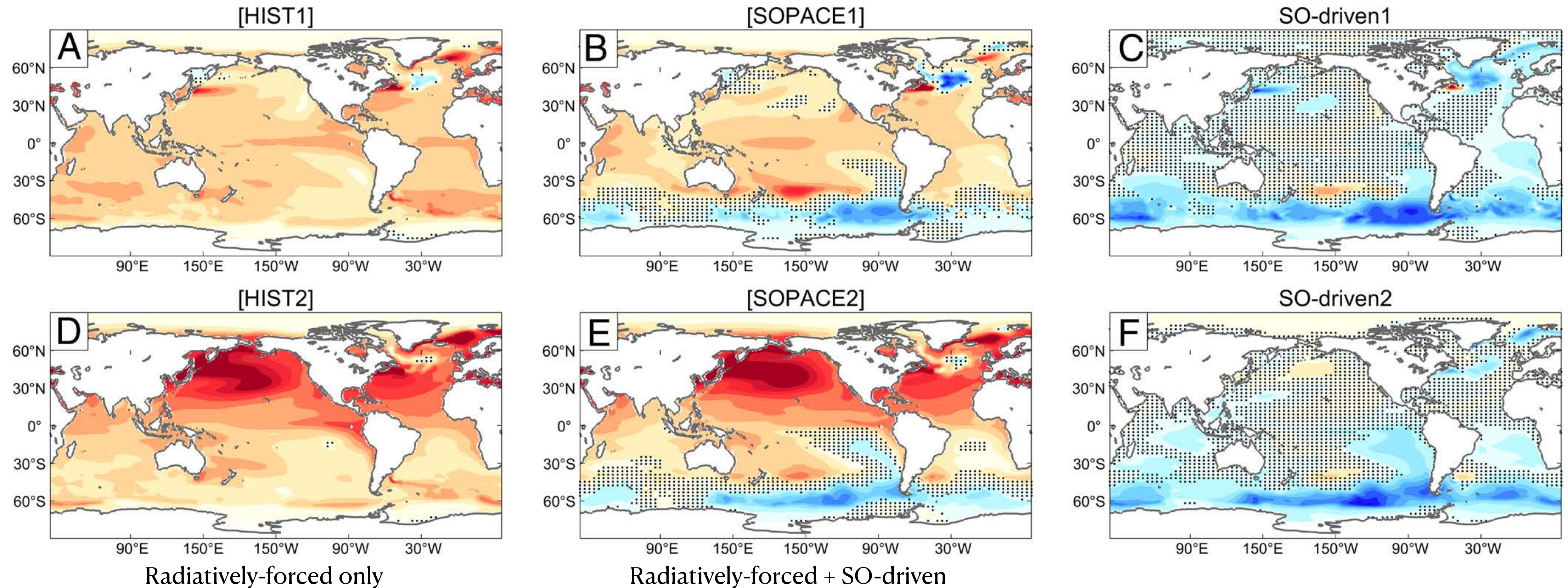
K/dec

12

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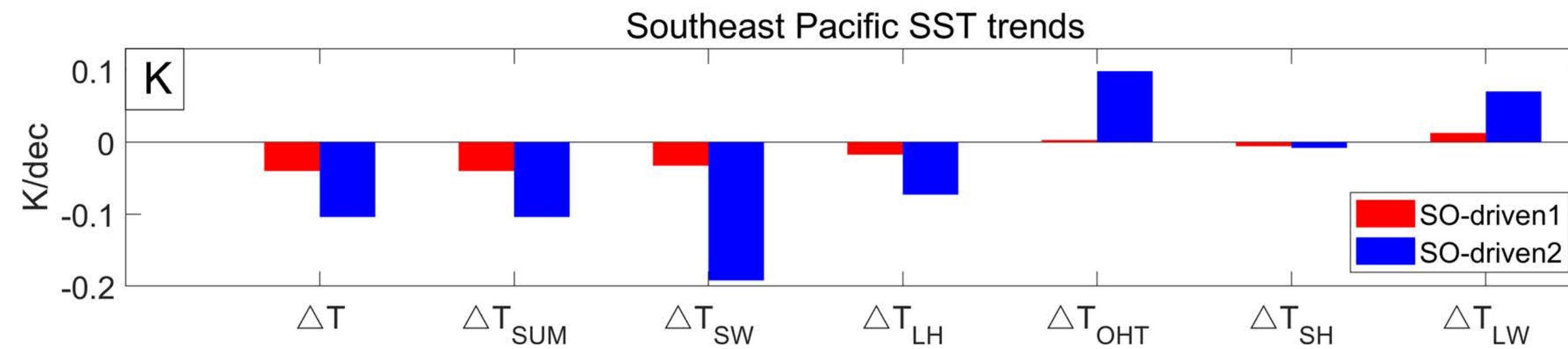
Global SST response to observed SO cooling

CESM1 vs CESM2



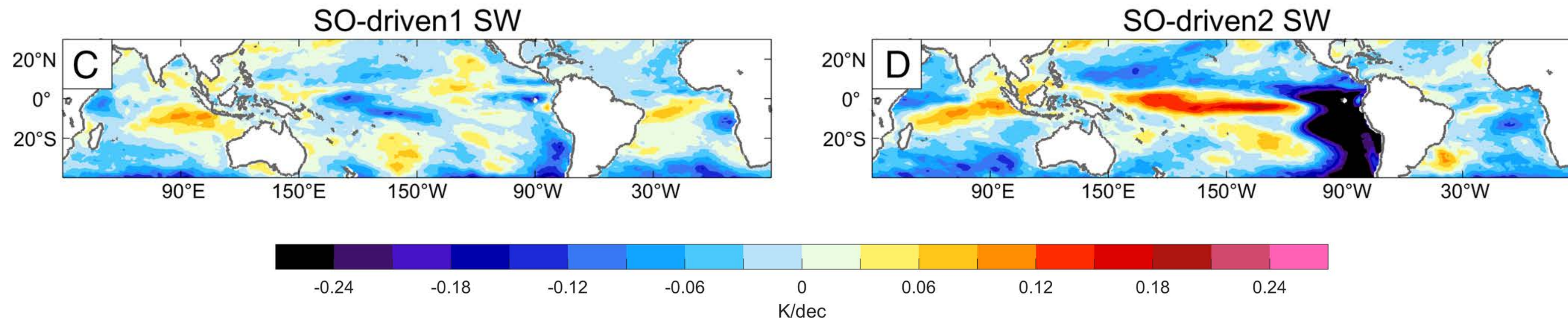
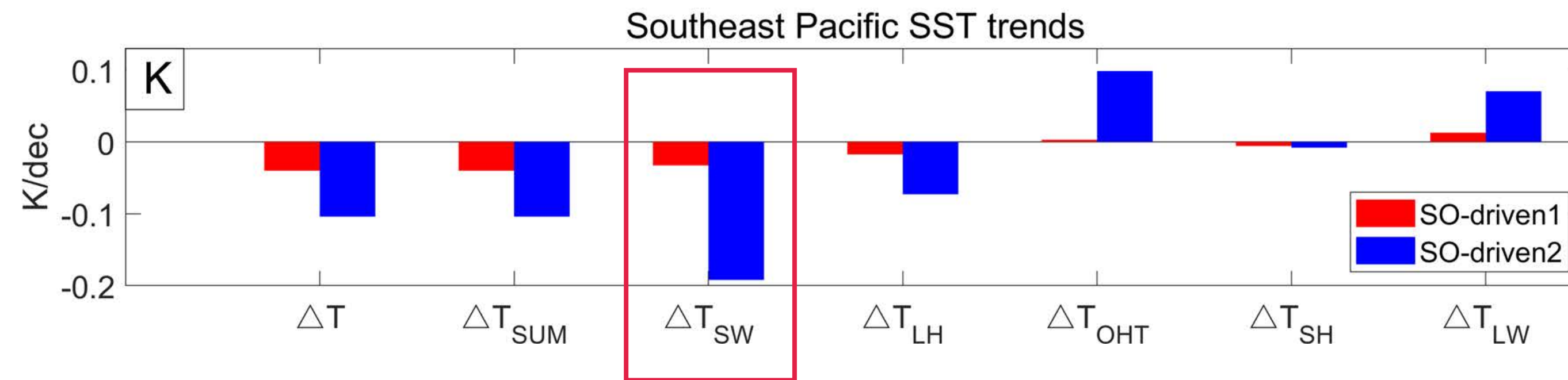
Decomposition of SST trends via surface energy budget

$$\rho c_p H \frac{\partial T}{\partial t} = SW + LW - LH - SH + OHT$$

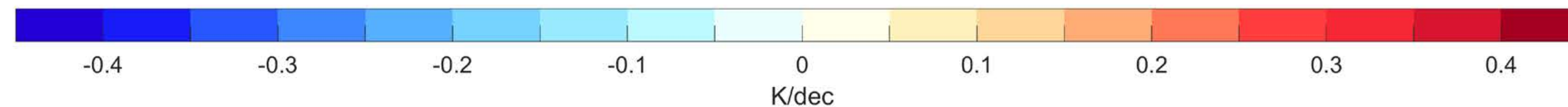
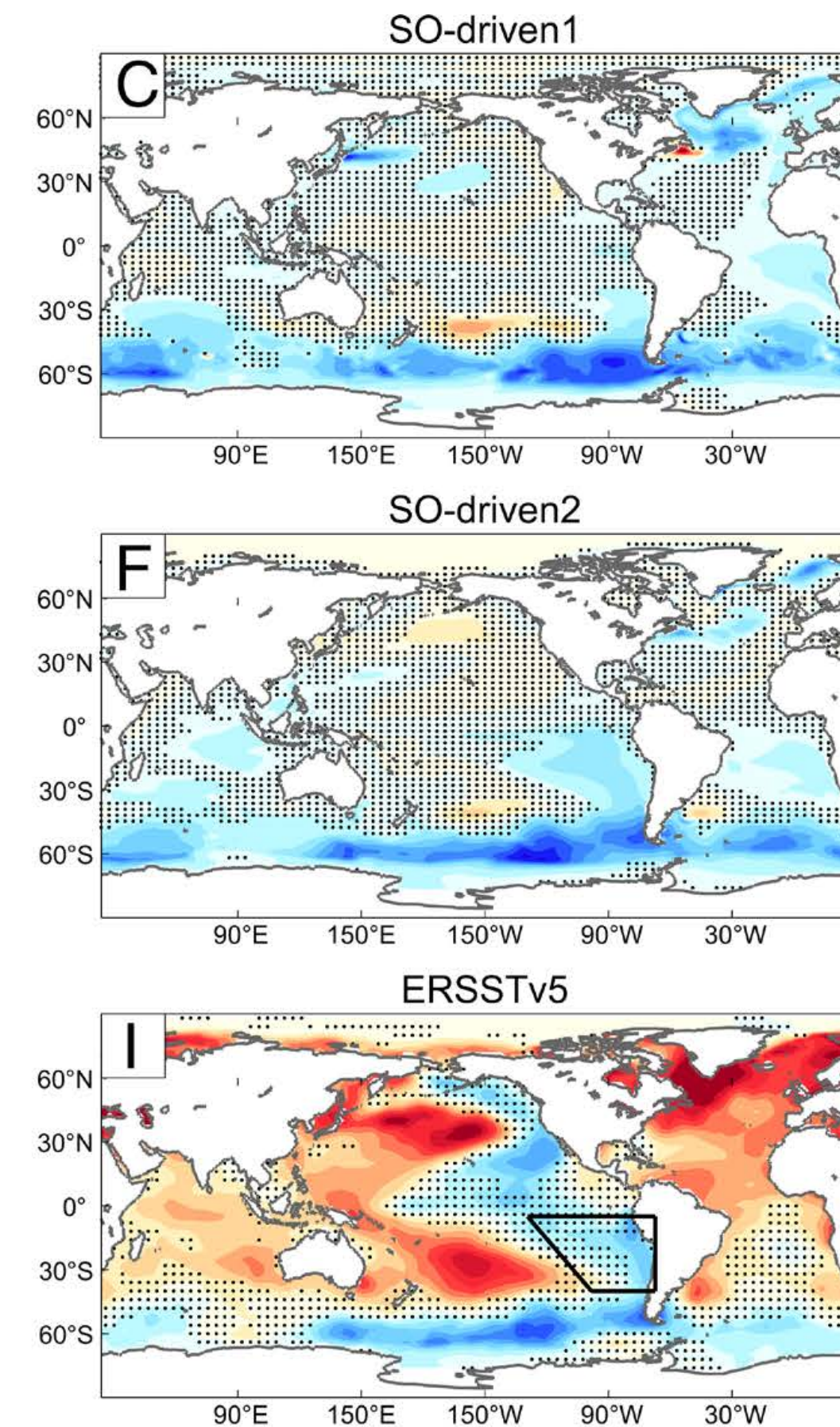


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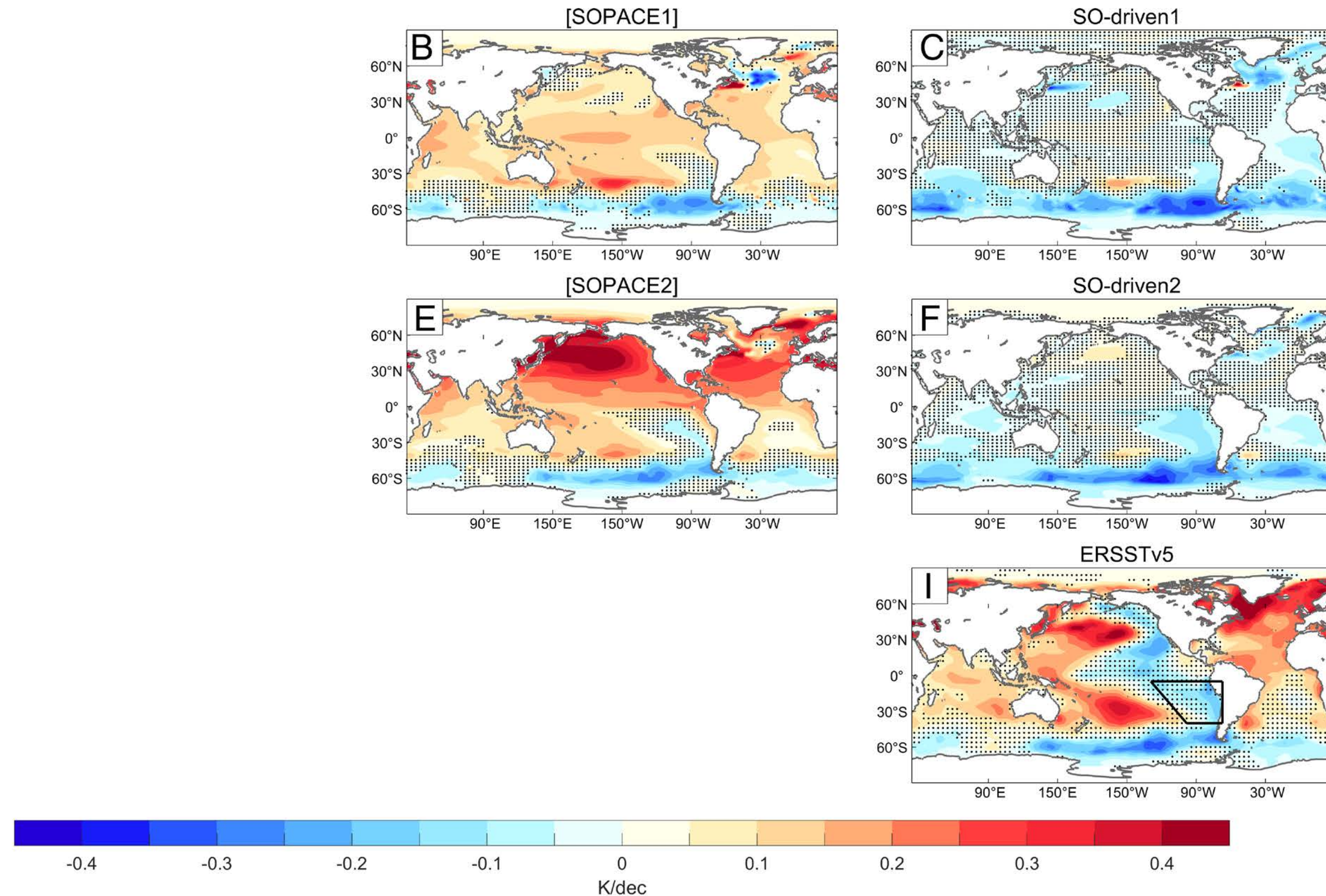
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Comparison with observations: Differences in radiatively-forced response

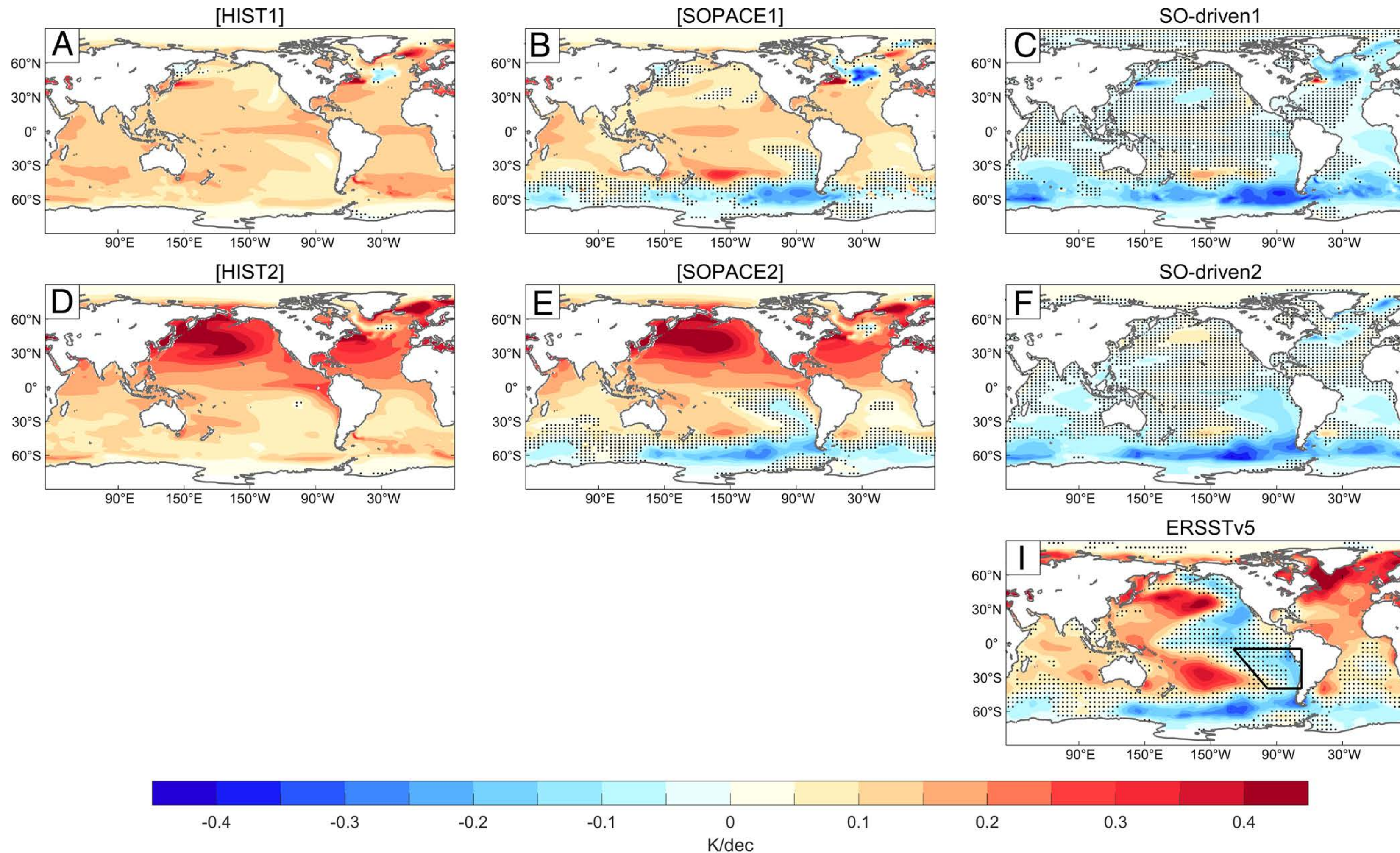


Comparison with observations: Differences in radiatively-forced response



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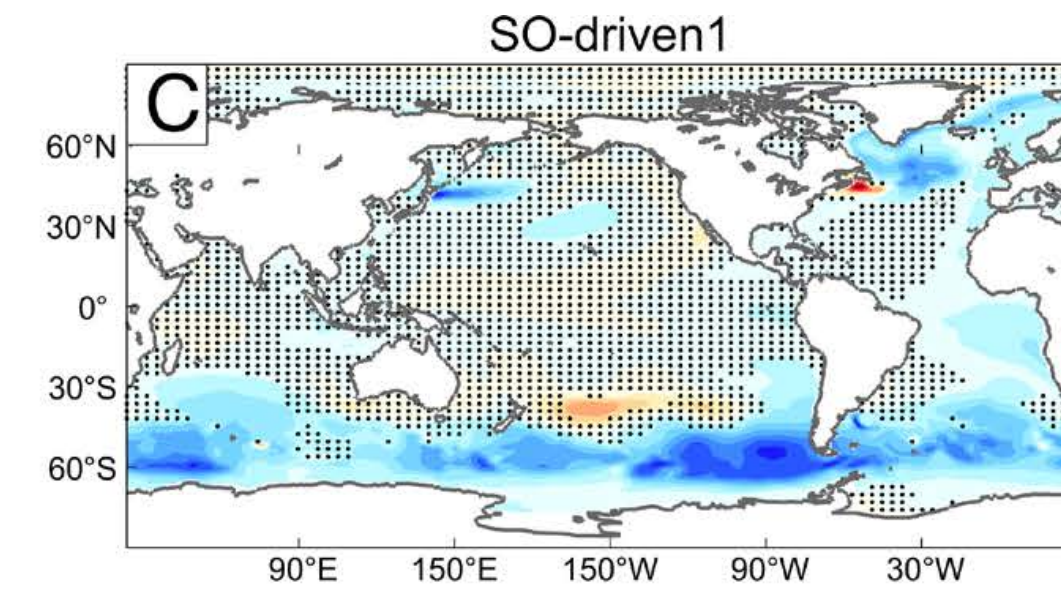
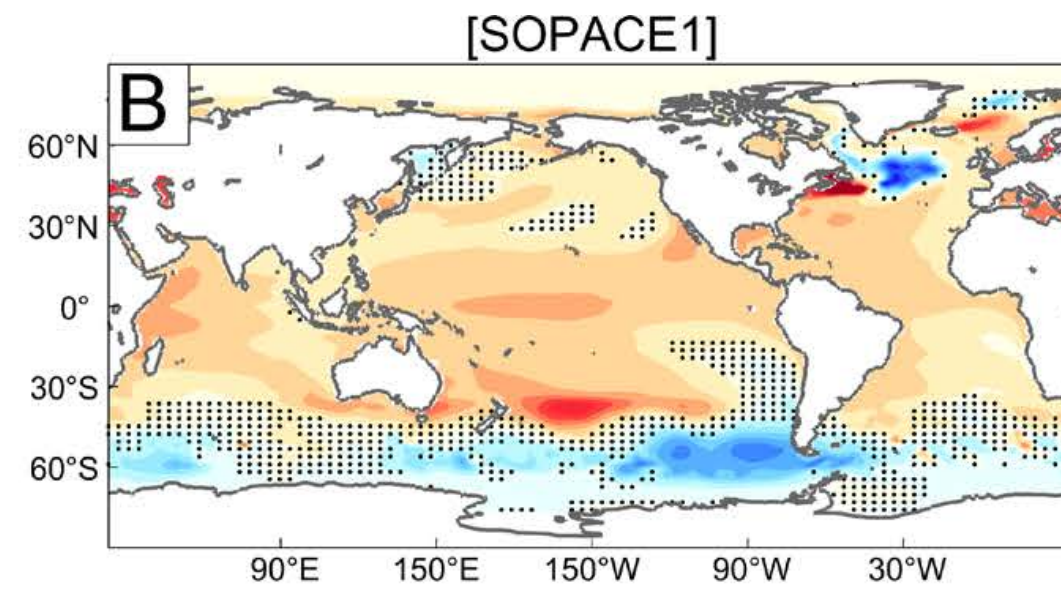
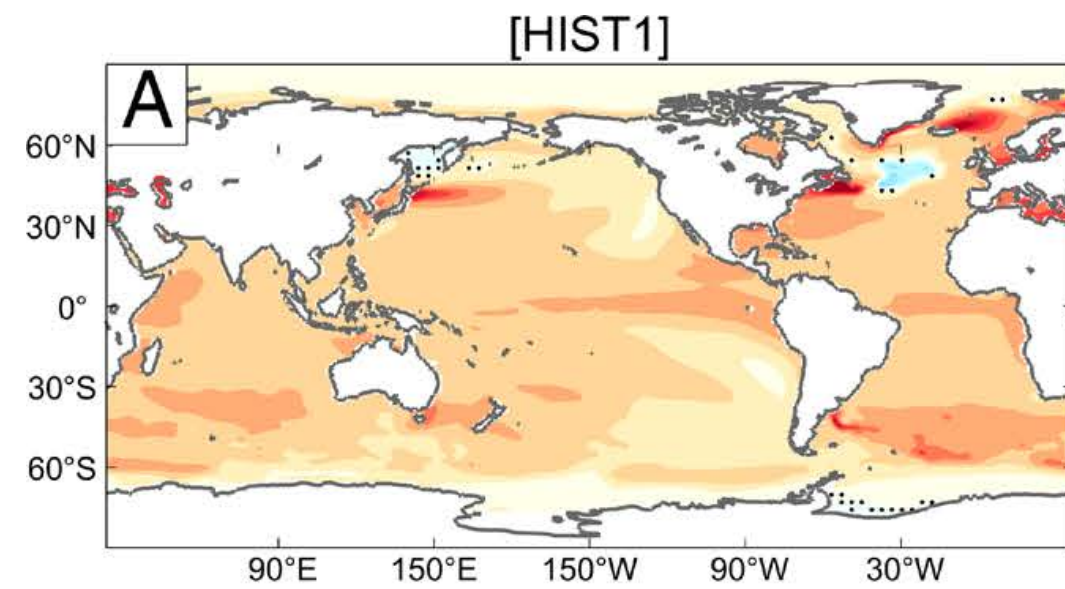
Comparison with observations: Differences in radiatively-forced response



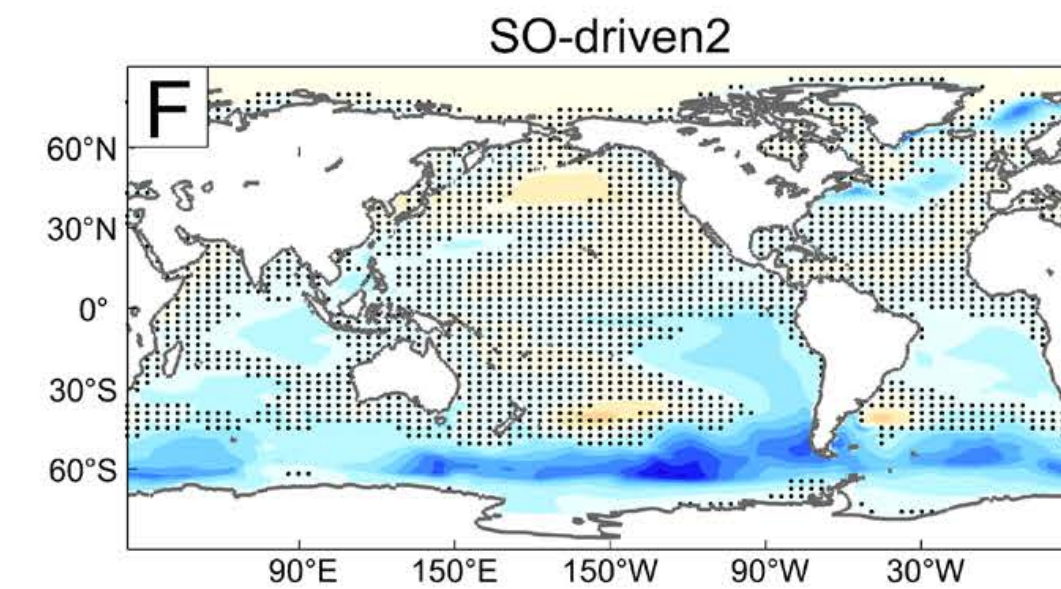
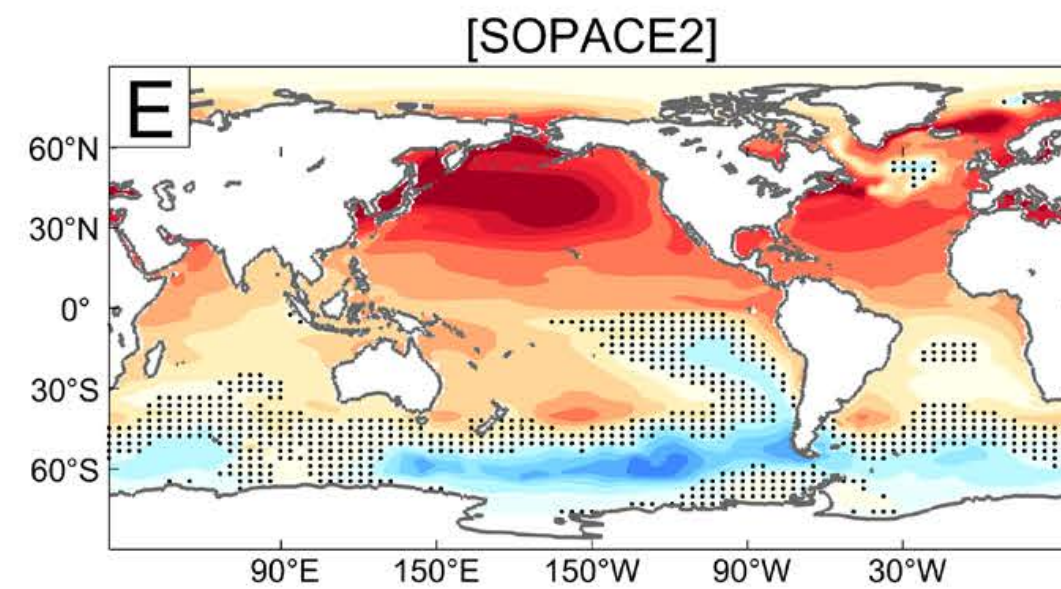
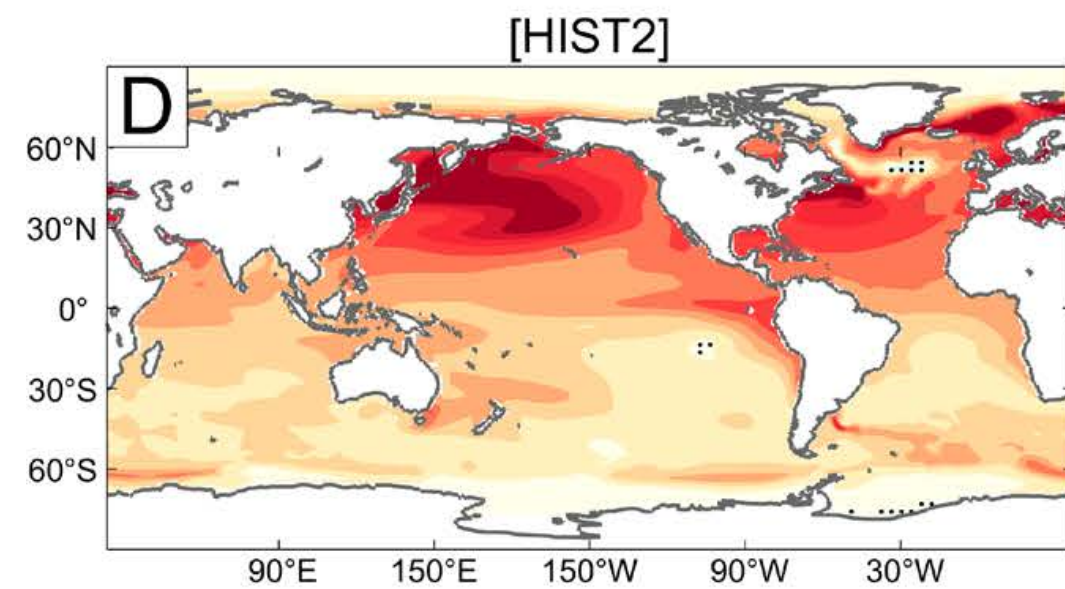
Stippling: trends NOT significant at 95% confidence level
Zhang et al. (2021); Kang et al. (2023)

Comparison with observations: Differences in radiatively-forced response

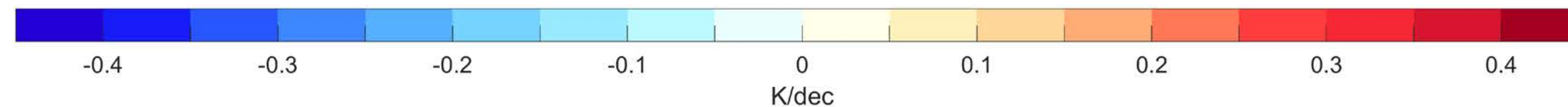
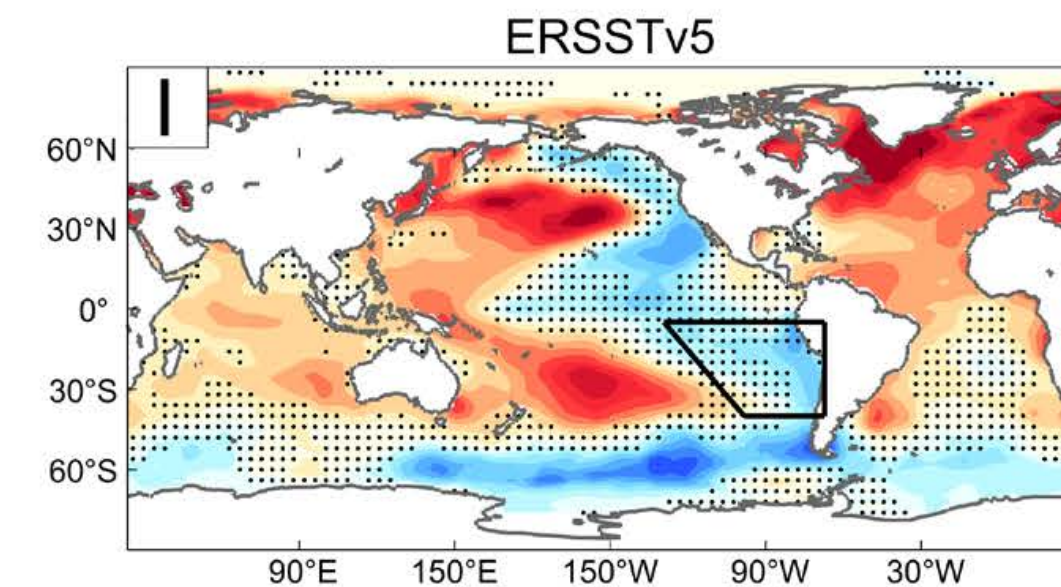
CESM1 + CMIP5 forcing



CESM2 + CMIP6 forcing



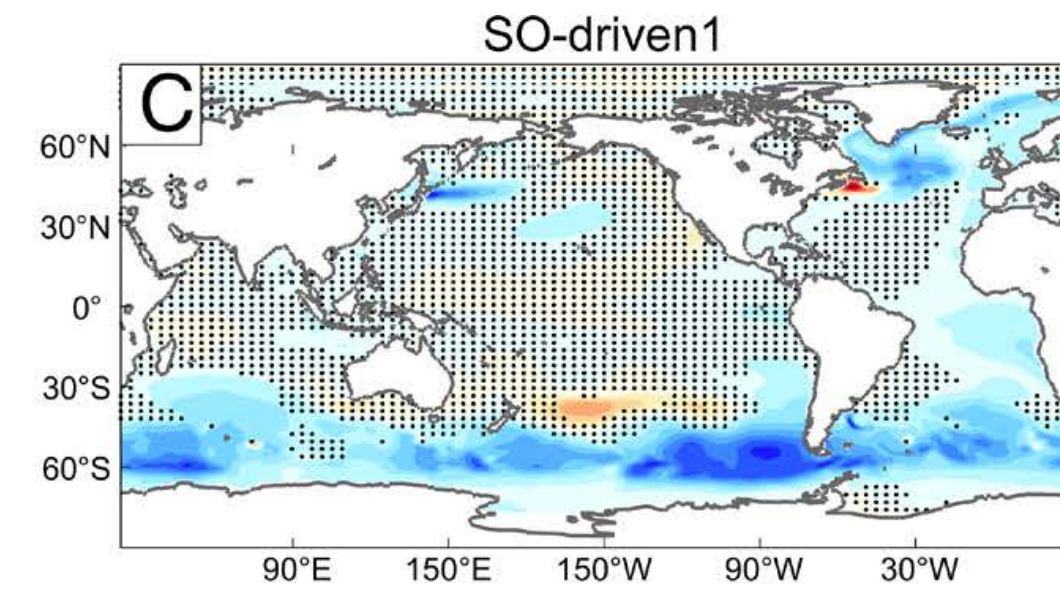
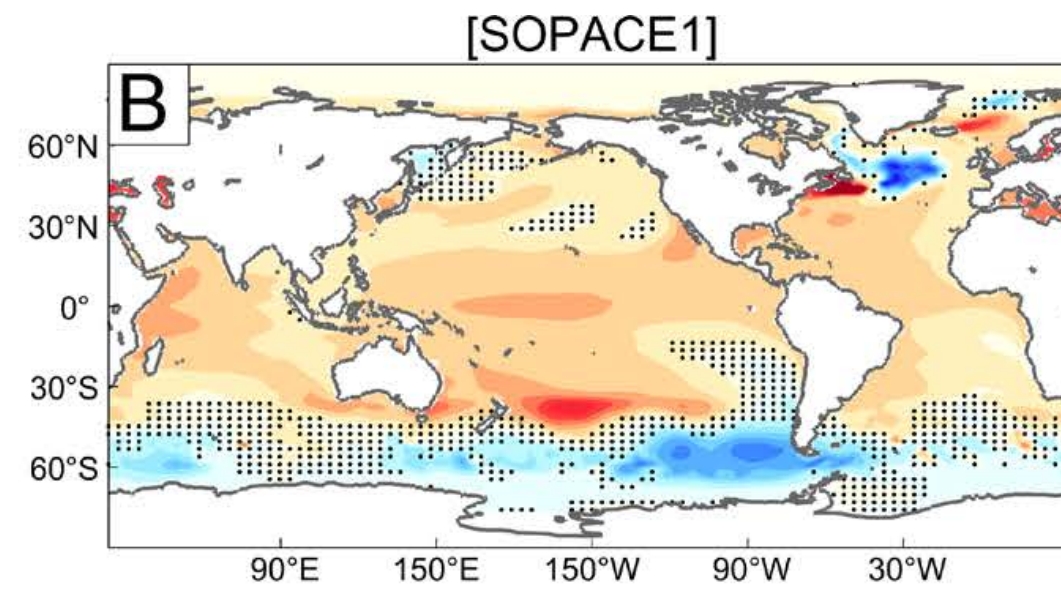
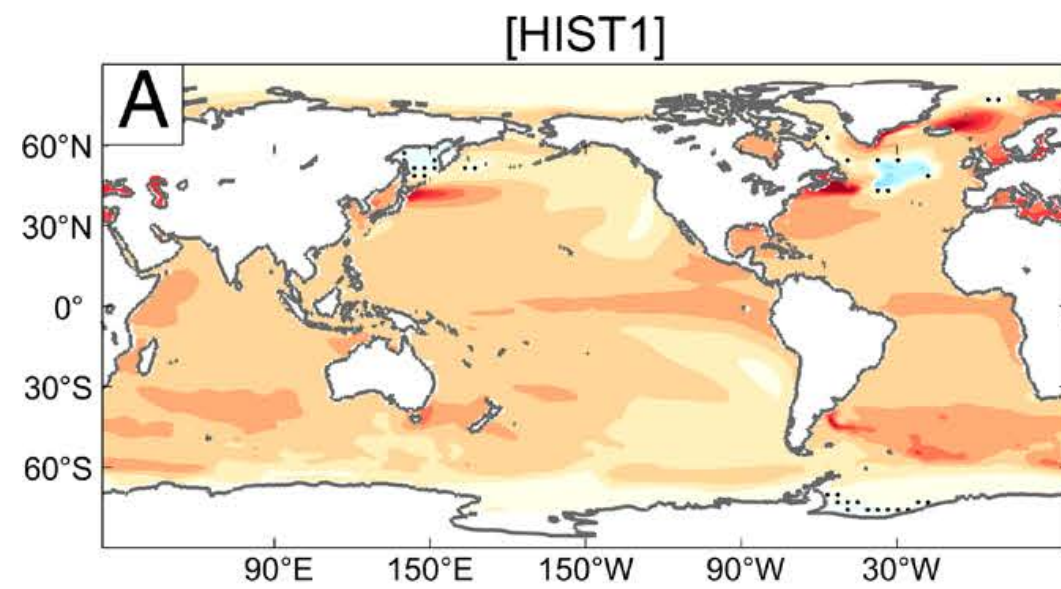
CESM2 + CMIP5 forcing



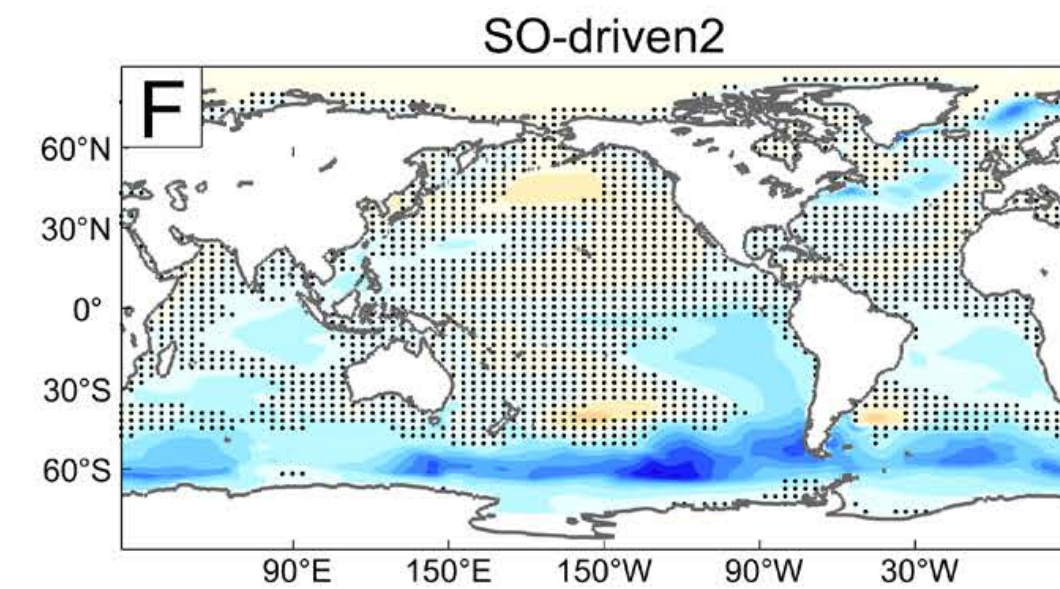
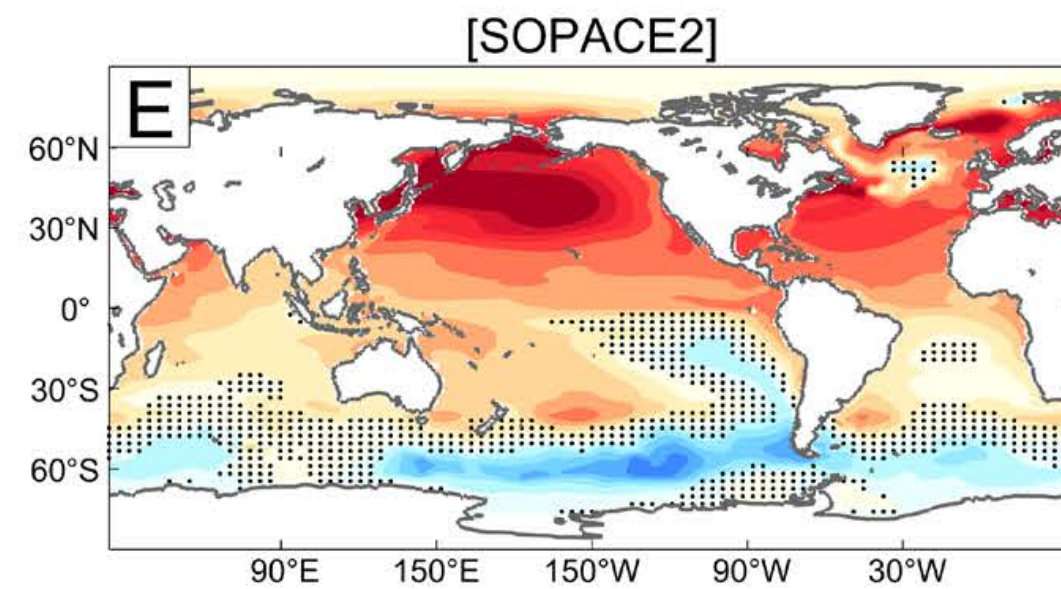
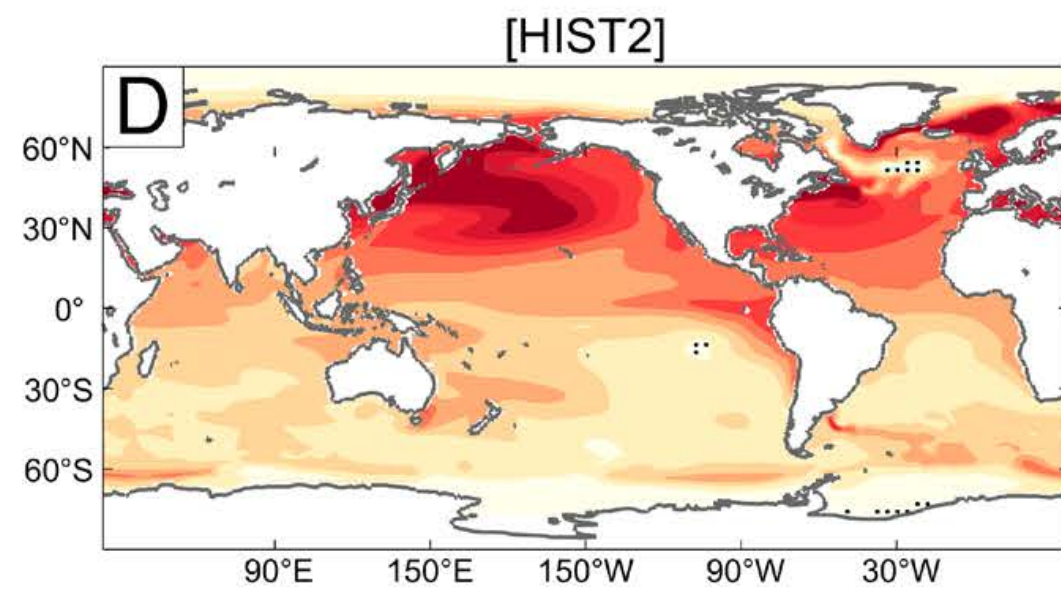
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Comparison with observations: Differences in radiatively-forced response

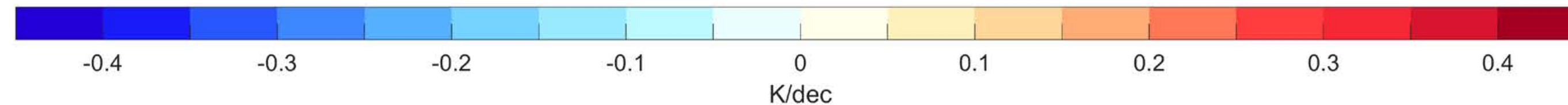
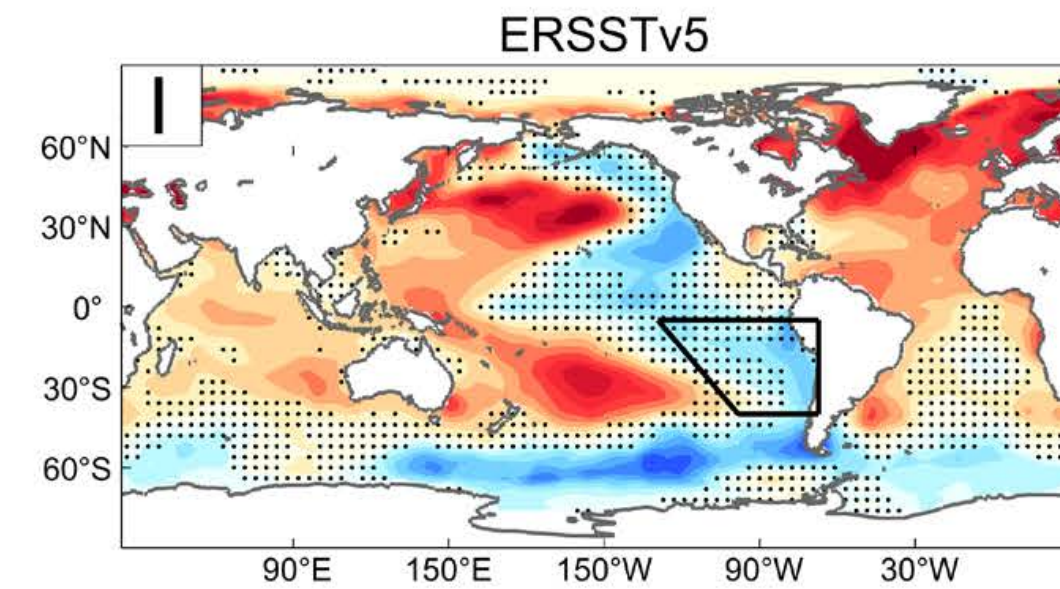
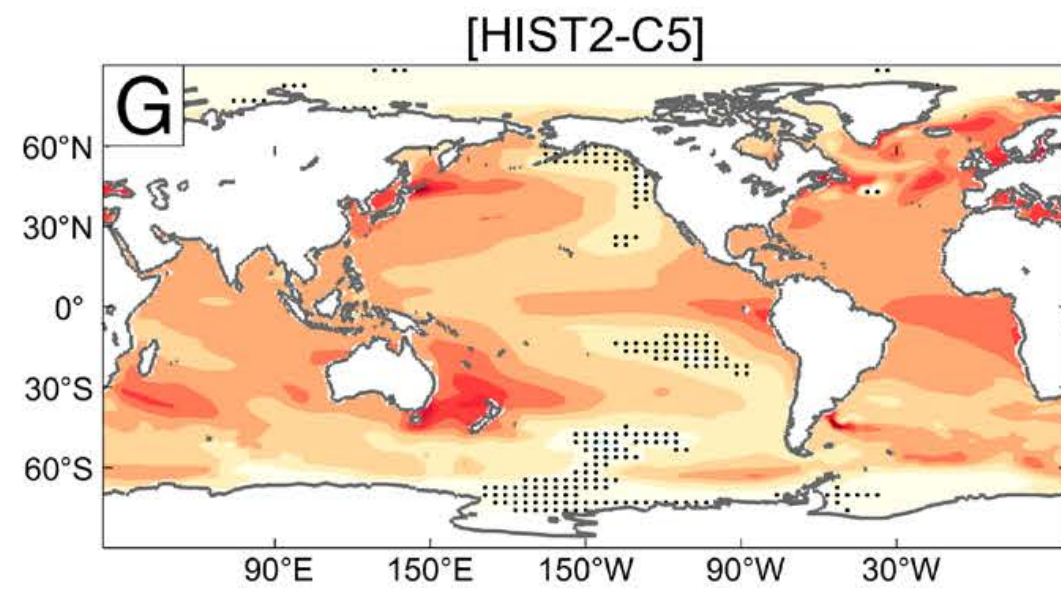
CESM1 + CMIP5 forcing



CESM2 + CMIP6 forcing

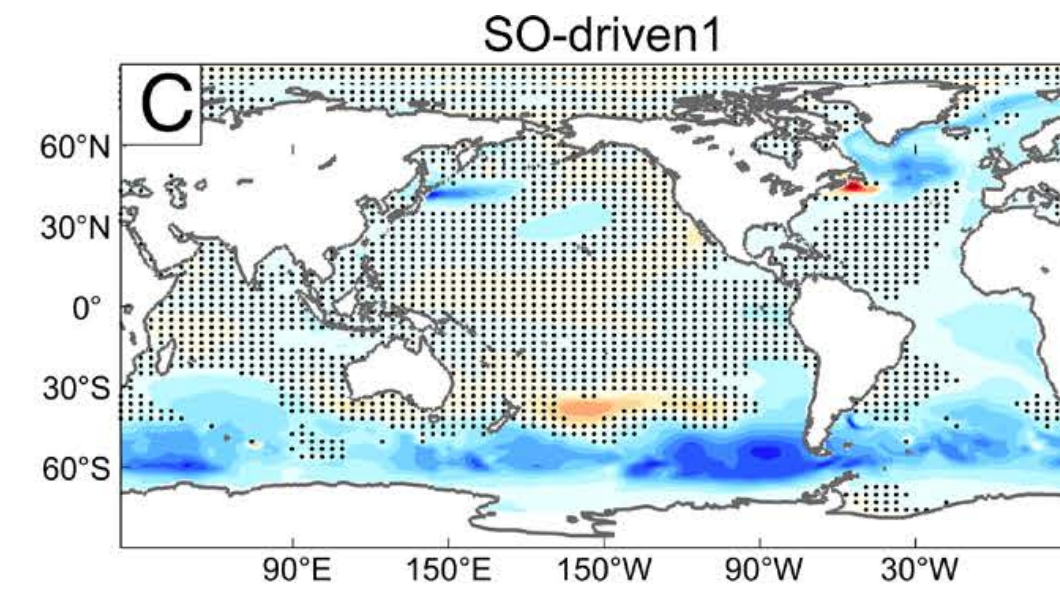
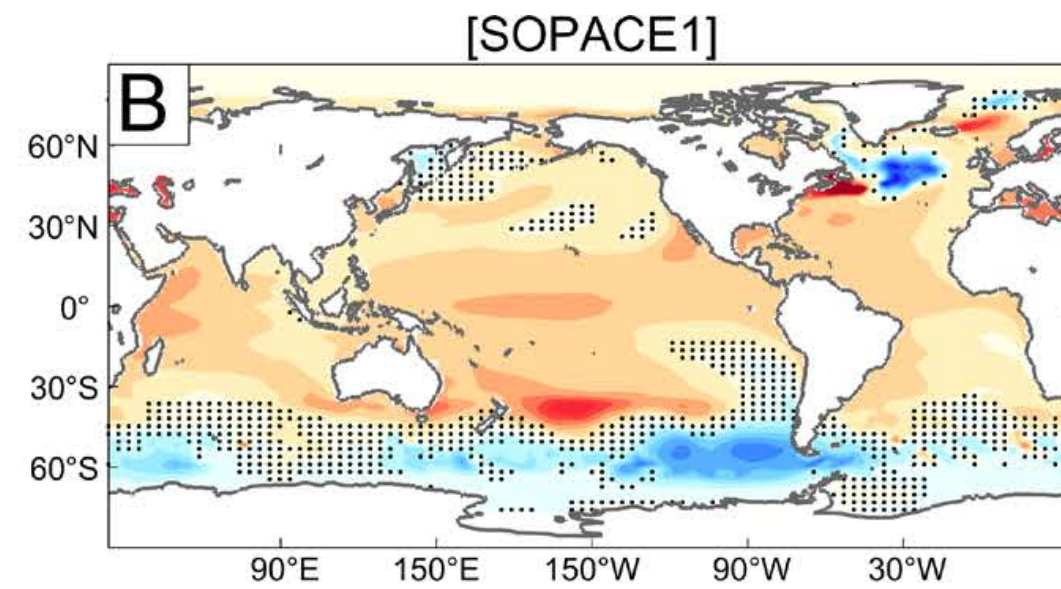
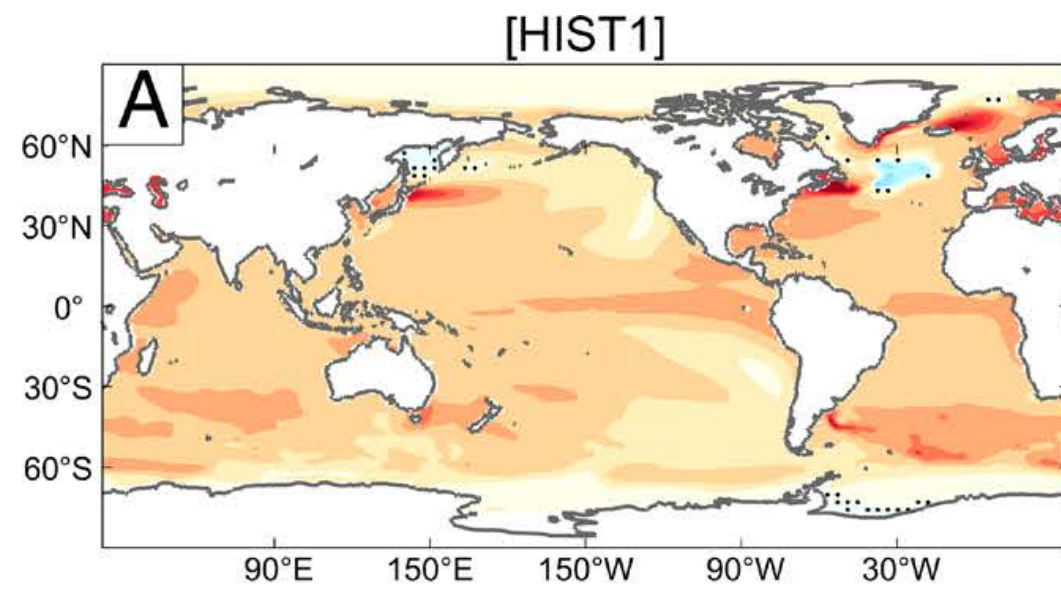


CESM2 + CMIP5 forcing

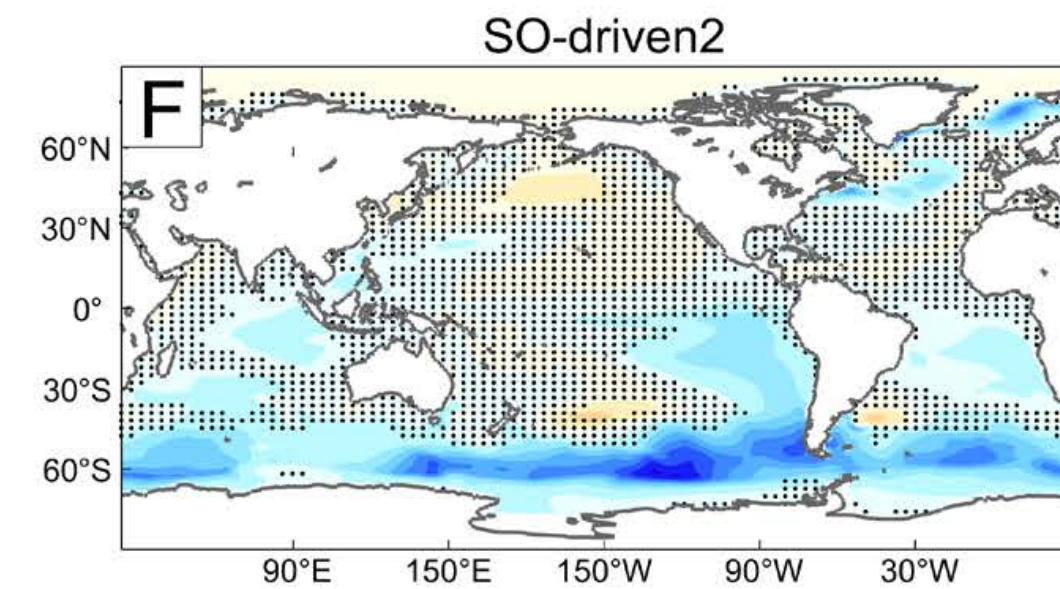
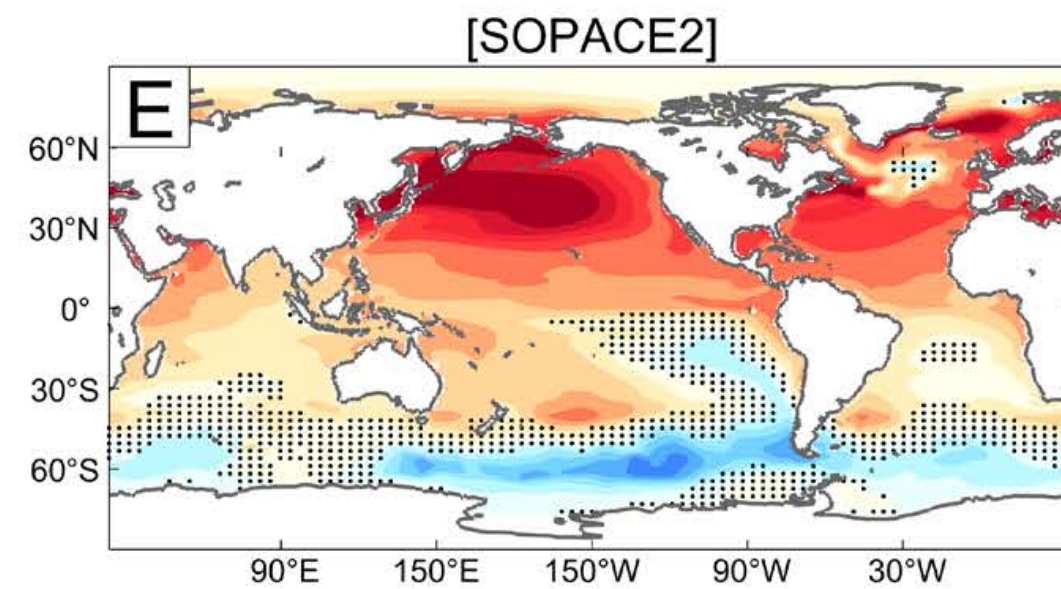
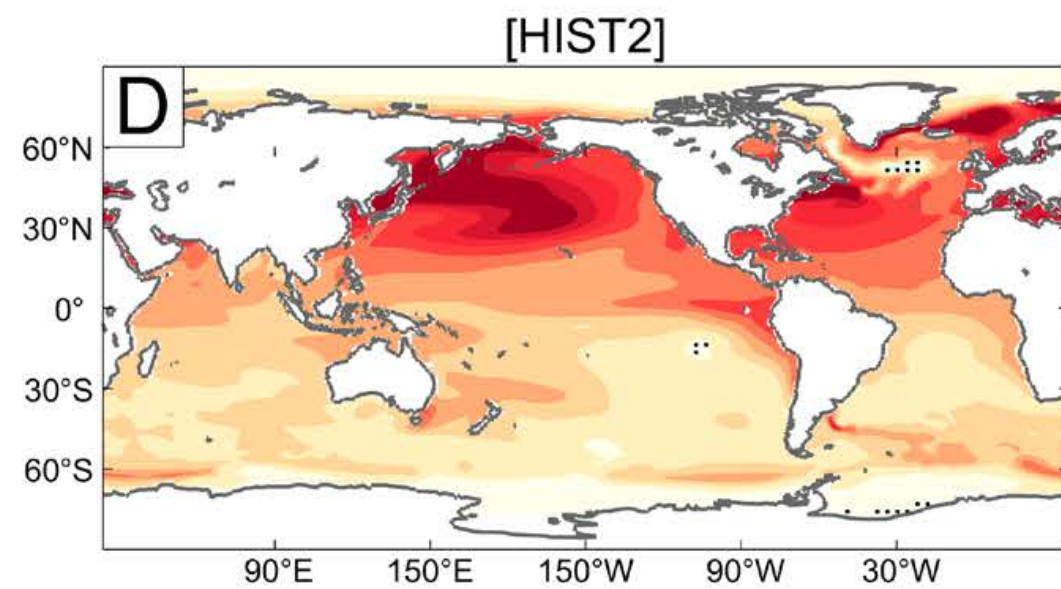


Comparison with observations: Differences in radiatively-forced response

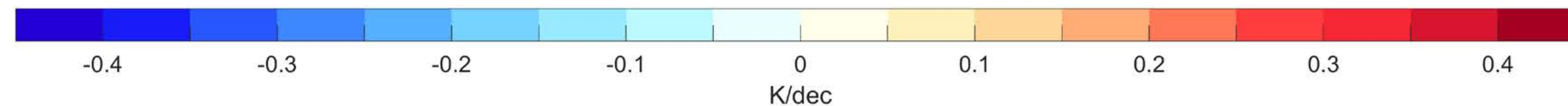
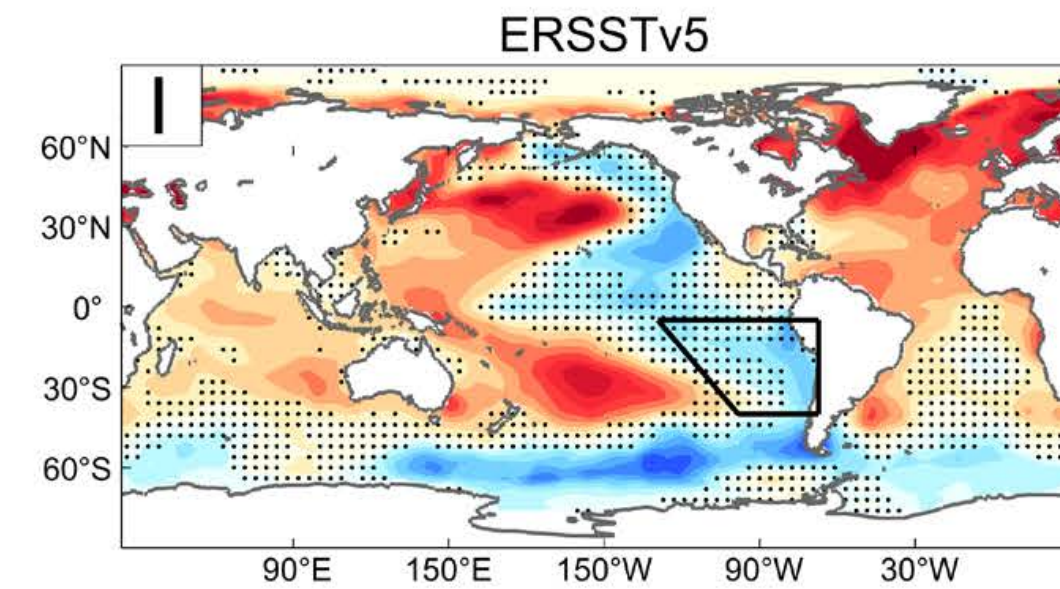
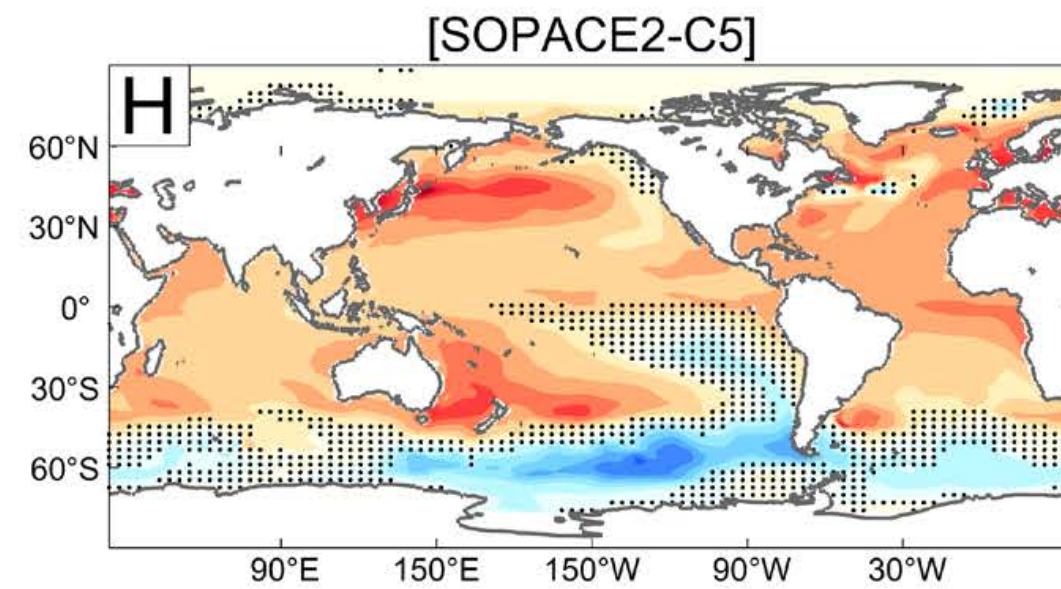
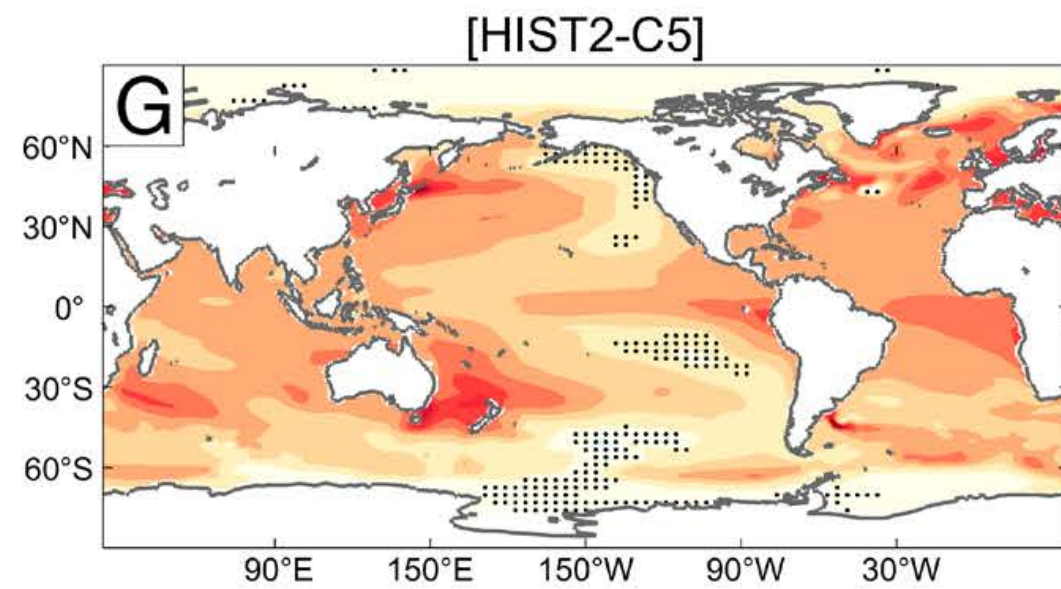
CESM1 + CMIP5 forcing



CESM2 + CMIP6 forcing

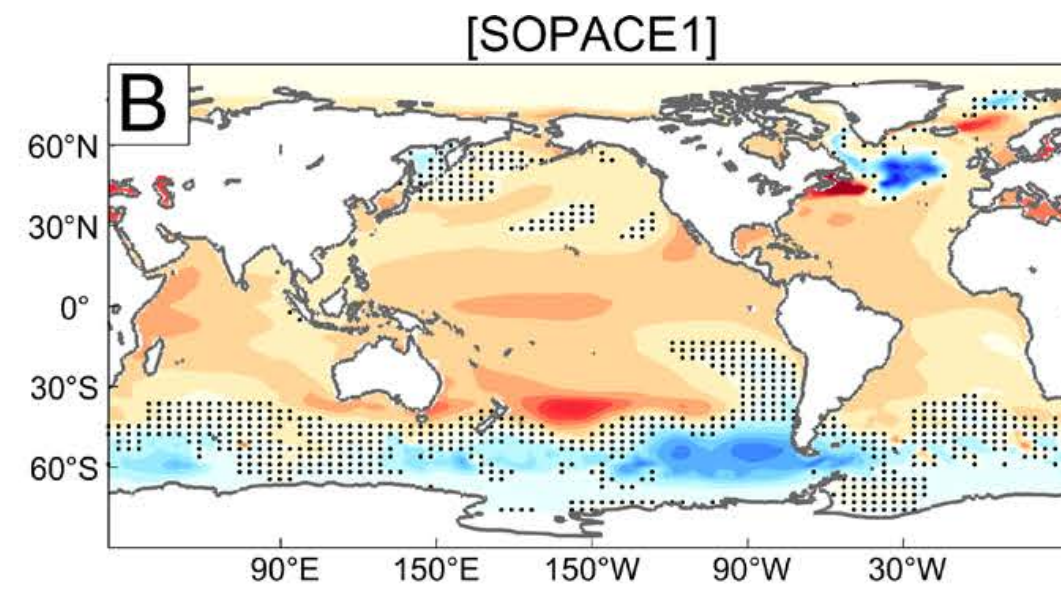
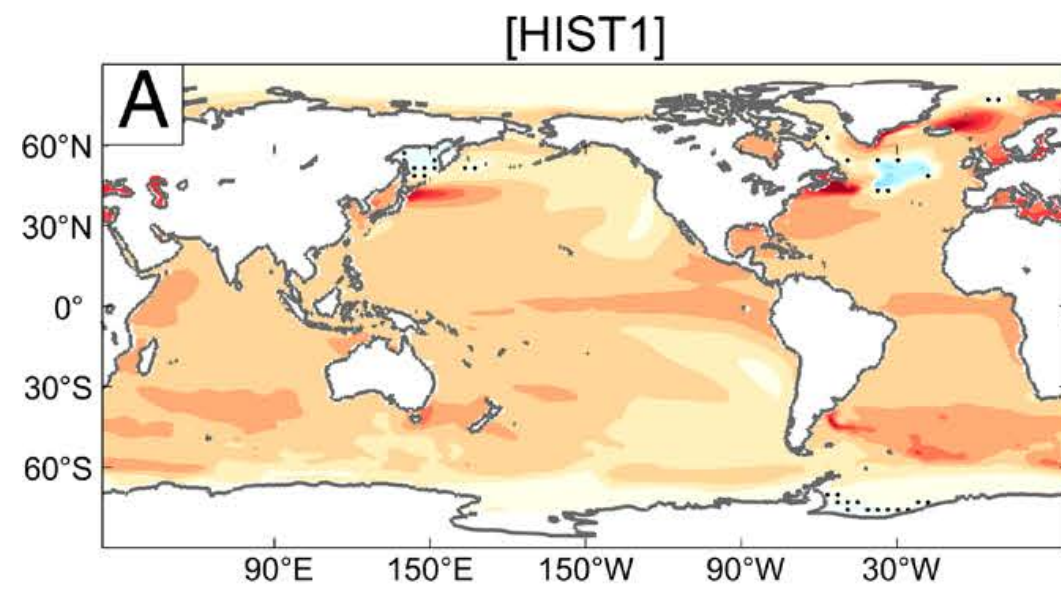


CESM2 + CMIP5 forcing

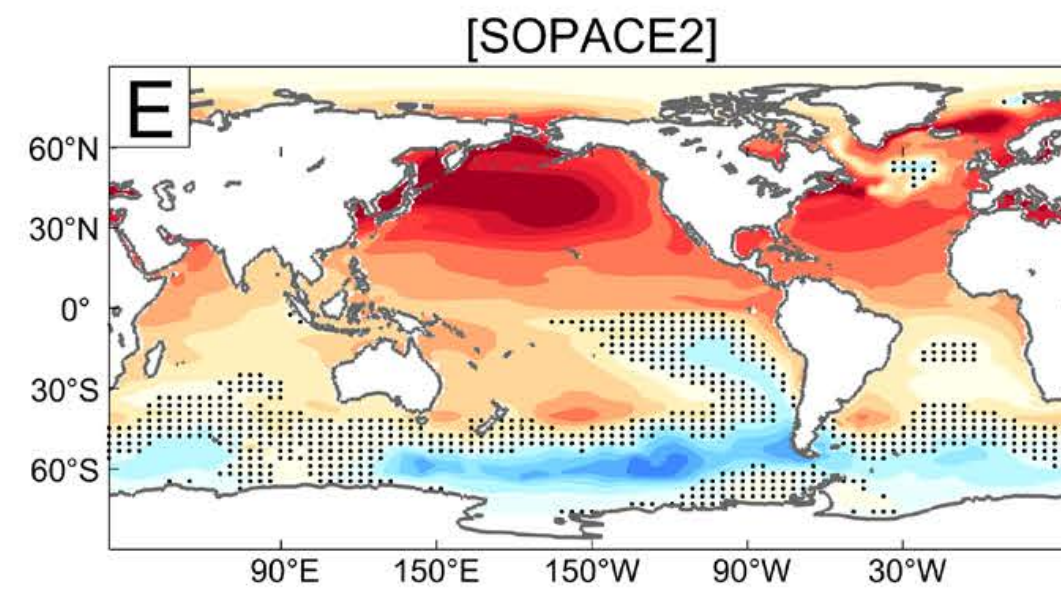
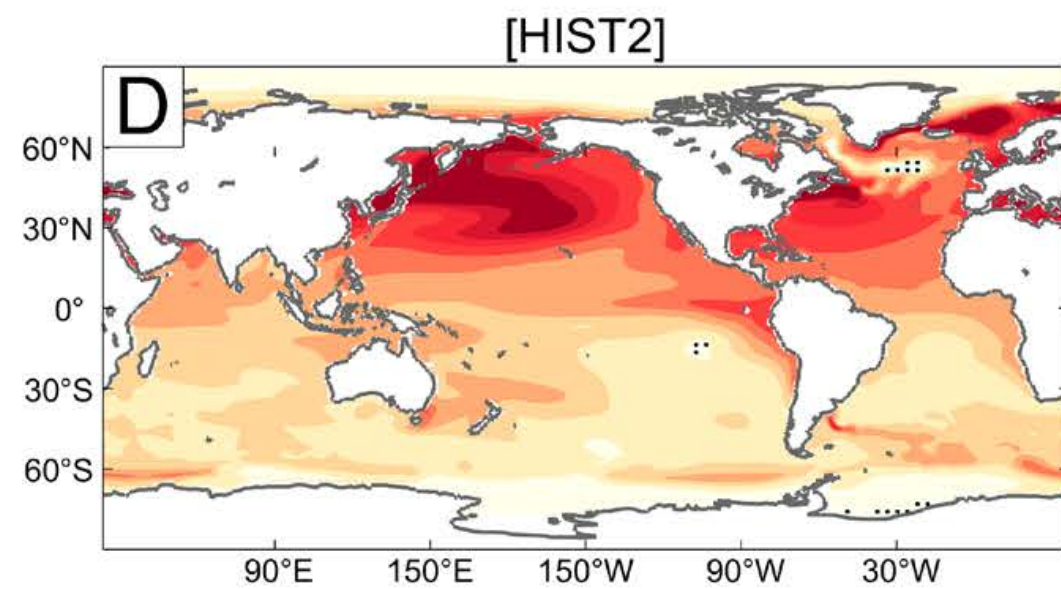


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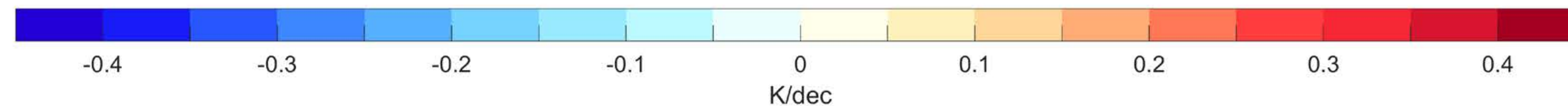
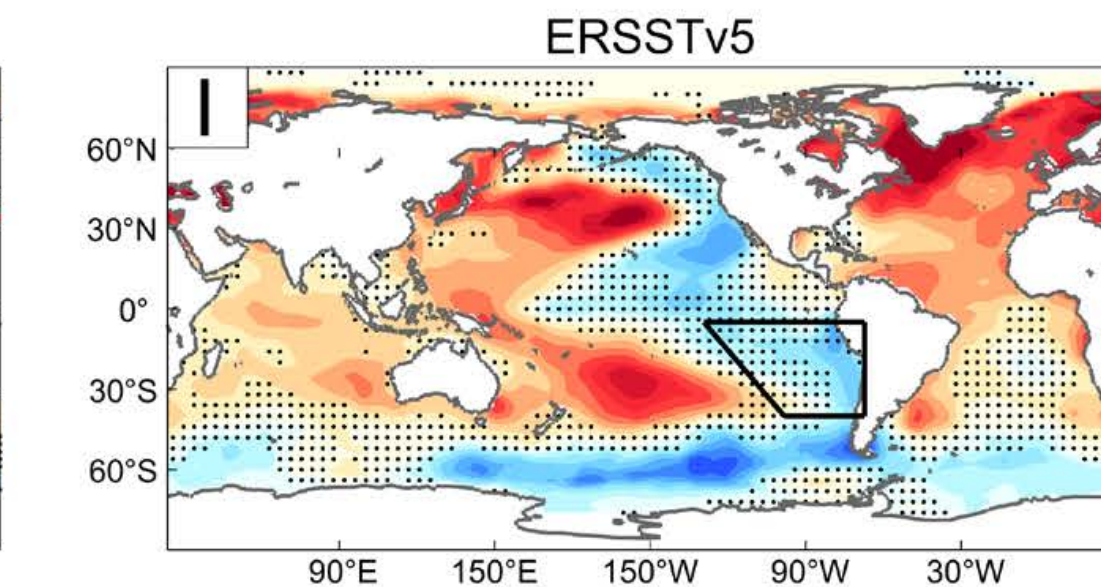
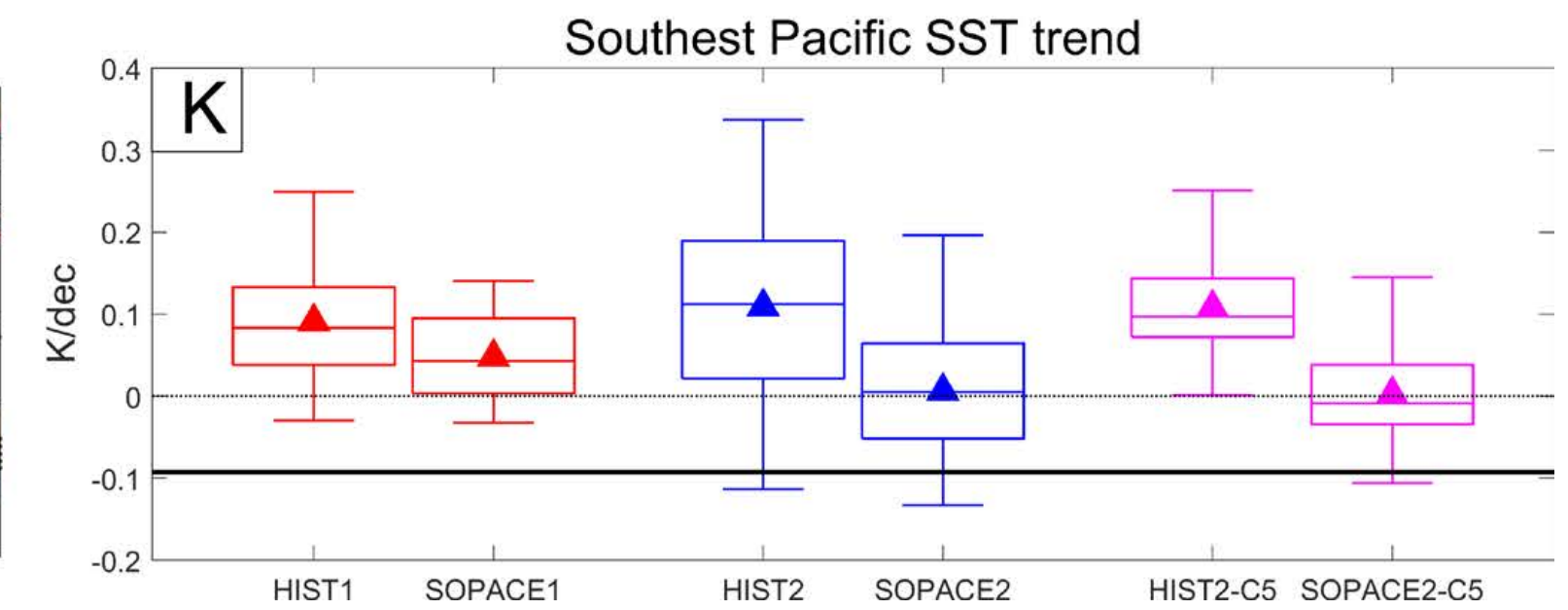
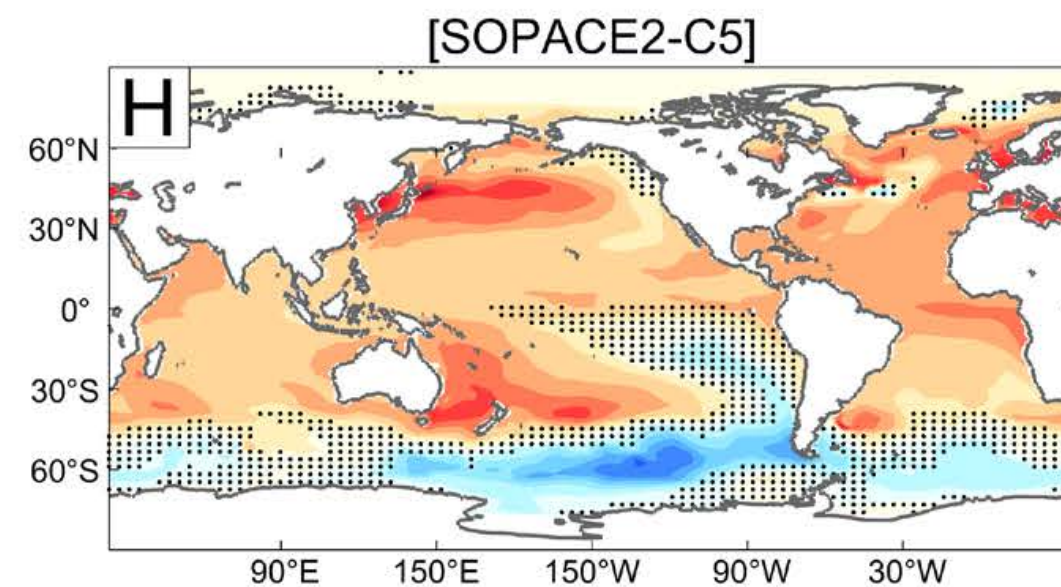
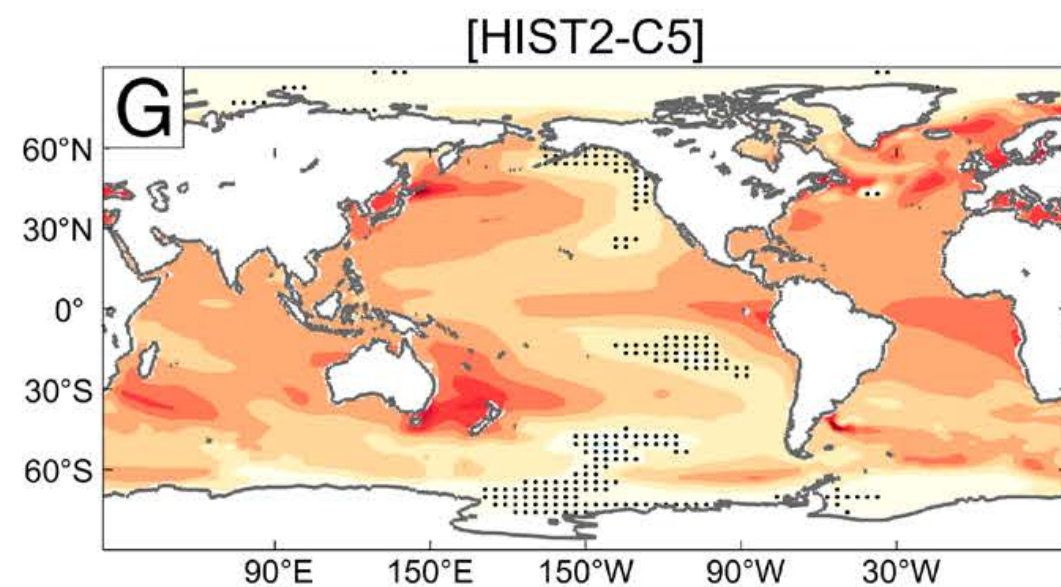
CESM1 + CMIP5 forcing



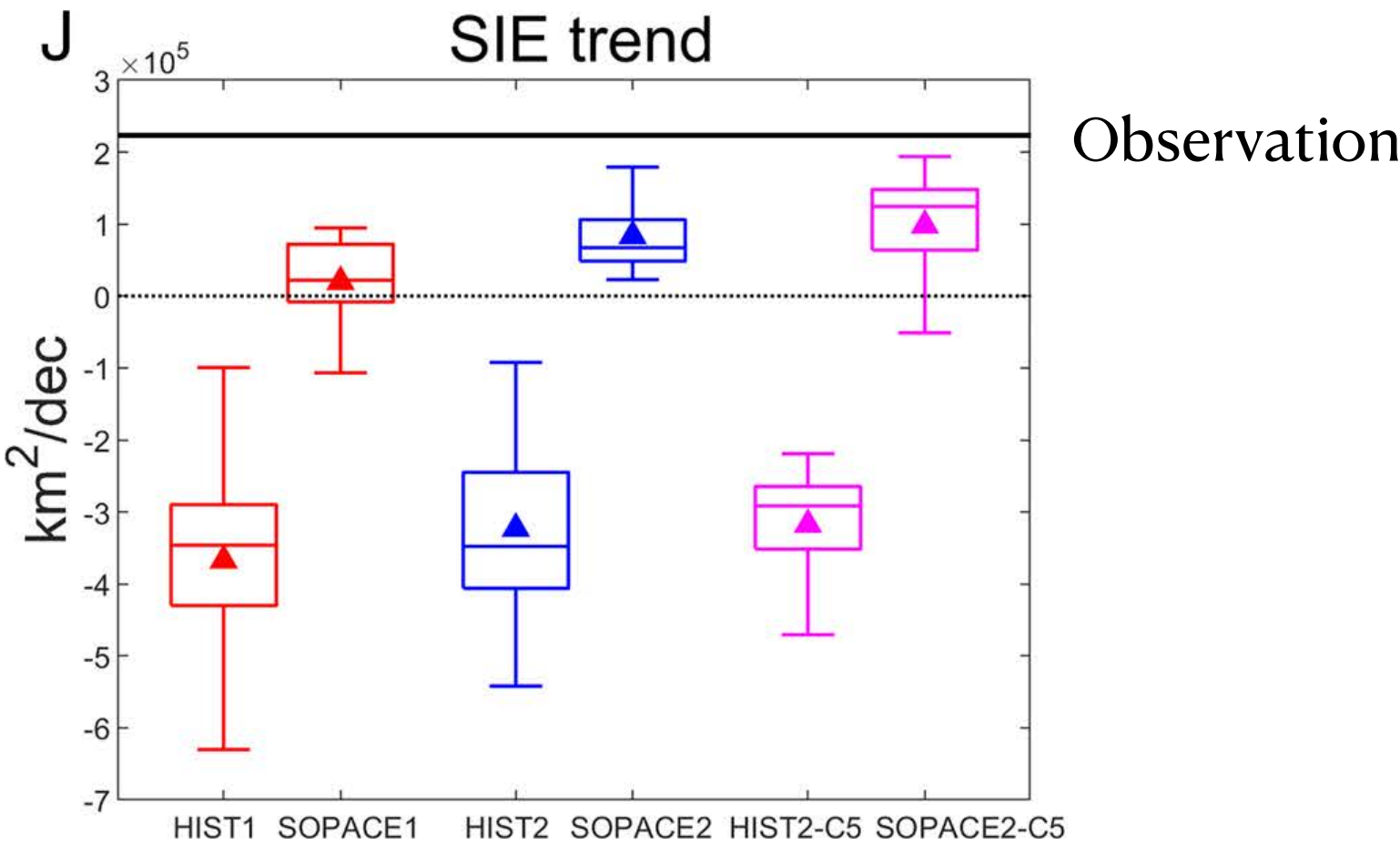
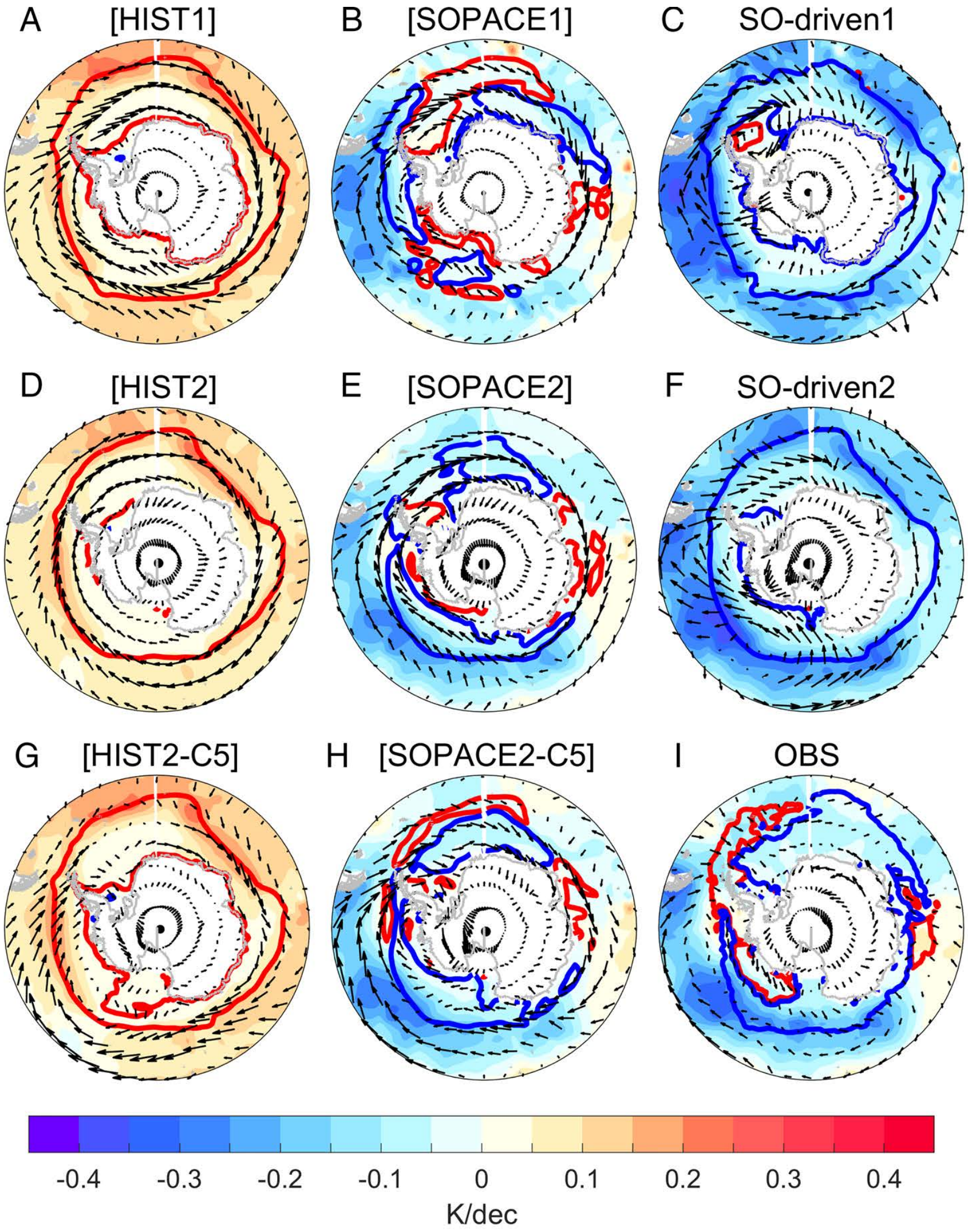
CESM2 + CMIP6 forcing



CESM2 + CMIP5 forcing



Observed Antarctic sea ice trends are also better represented in SOPACE2



Summary

- The global response of observed SO surface cooling includes cooling of the southeastern tropical Pacific and Atlantic, as well as Antarctic sea ice expansion
- Observed SO surface cooling from 1979 to 2013 is partly responsible for driving cooling of the southeastern tropical Pacific SST
- The SO-tropical teleconnection is highly sensitive to the strength of the subtropical low cloud feedback
- There are implications for future warming patterns as the SO transitions from cooling to warming under increasing greenhouse gases

Geophysical Research Letters

RESEARCH LETTER
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Key Points:

- The global impact of recent

Is There a Tropical Response to Recent Observed Southern Ocean Cooling?

Xiyue Zhang^{1,2} , Clara Deser¹ , and Lantao Sun³ 

PNAS

RESEARCH ARTICLE

EARTH, ATMOSPHERIC, AND PLANETARY SCIENCES

 OPEN ACCESS

Global impacts of recent Southern Ocean cooling

Sarah M. Kang^{a,b,1} , Yue Yu^c , Clara Deser^{d,1} , Xiyue Zhang^e, In-Sik Kang^c, Sun-Seon Lee^{f,g}, Keith B. Rodgers^{f,g} , and Paulo Ceppi^h 

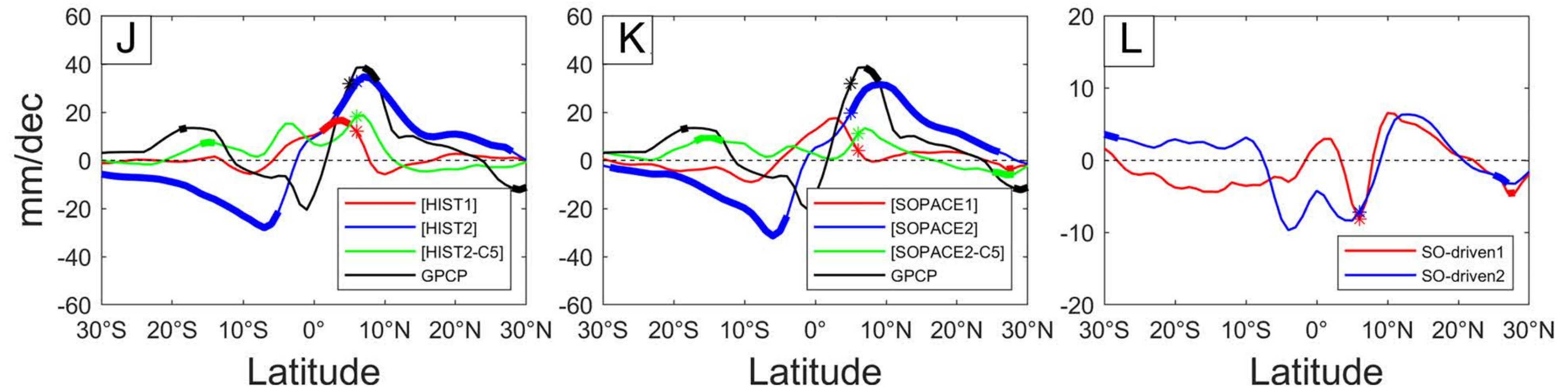
Challenges and open questions

- Uncertainties in historical radiative forcing hinder our understanding of the SO-tropical teleconnection (especially in the northern extratropics)
- Can we quantify the causes of SO SST multi-decadal variability (e.g., internal variability, CO₂ or ozone, ice melt...)?
- What other model/resolution-dependent feedbacks can influence the SO-driven teleconnection?

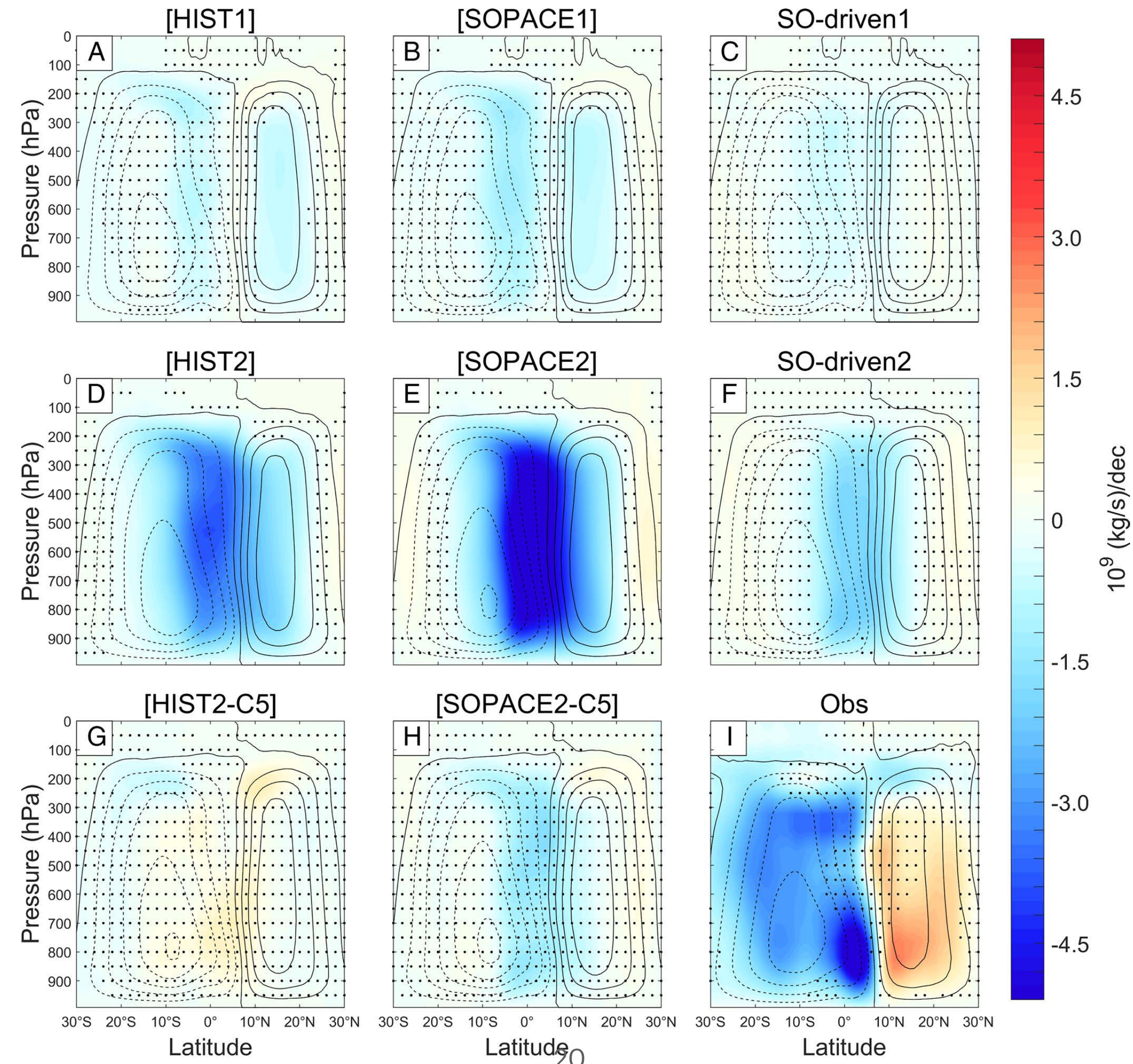
Questions and feedbacks?  sallyz@jhu.edu

Extra Slides

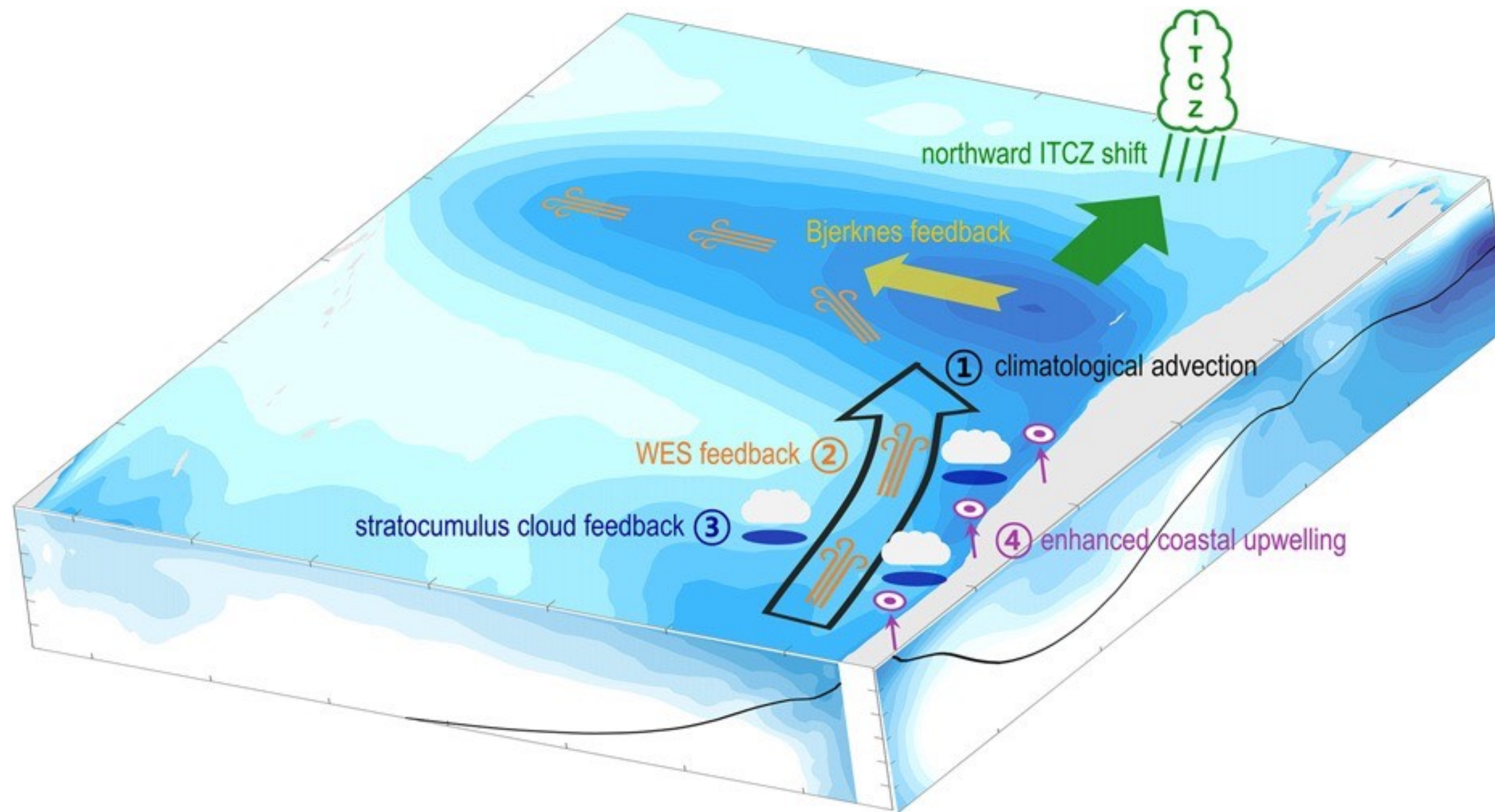
SO cooling's impact on tropical precipitation



SO cooling's impact on tropical circulation

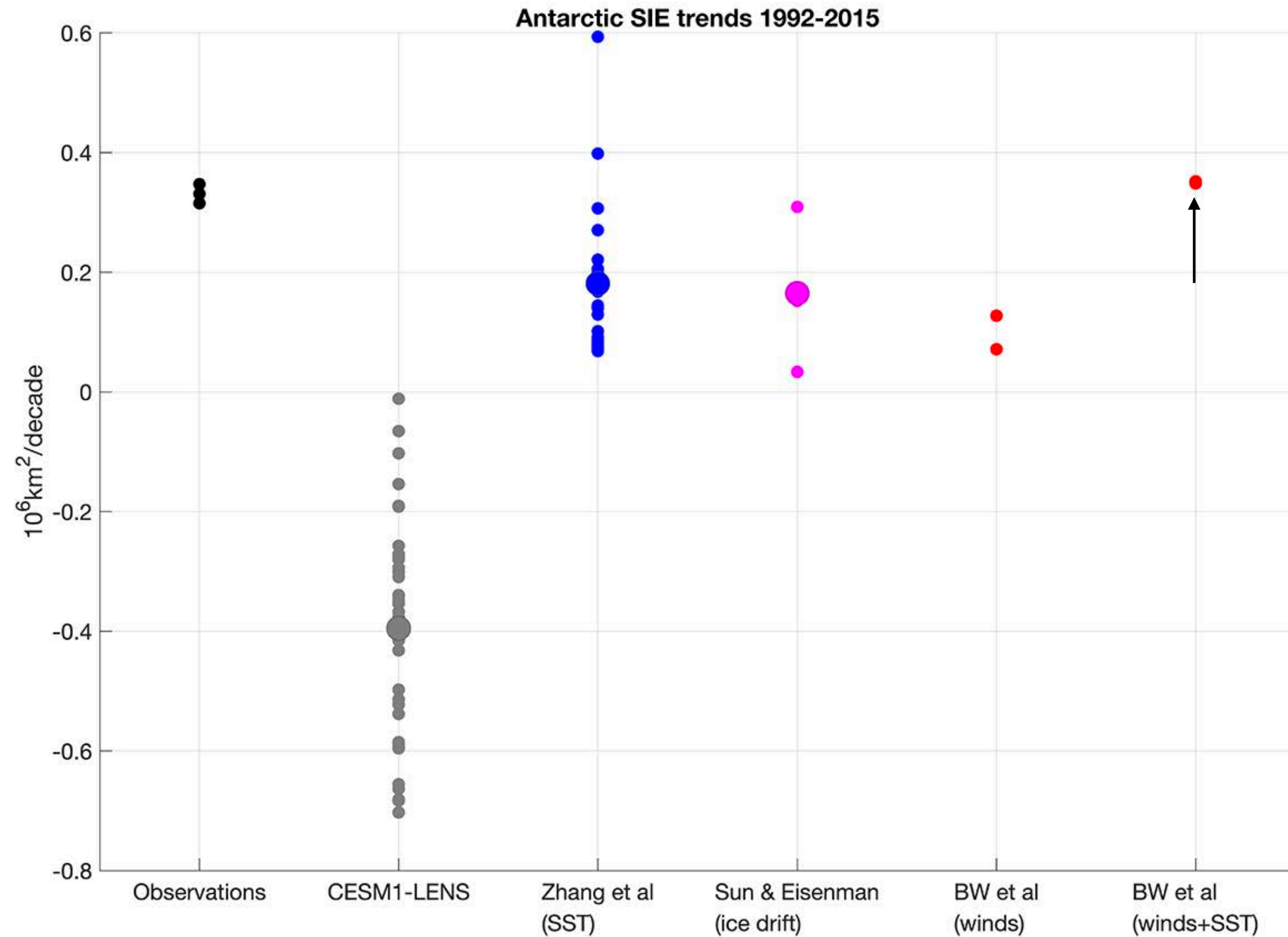


Schematic of SO teleconnection



Antarctic sea ice trends are better captured

when observed SO SST trends are included



Origin of SO SST variability is under debate

Forced response

- Ocean heat uptake (Marshall et al., 2015; Armour et al., 2016)
- Ozone (e.g., Ferreira et al., 2015; Hartmann 2022)
- Antarctic meltwater (e.g., Bronselaer et al., 2018)

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Internal variability

- Linked to tropical variability (e.g., Schneider and Deser 2018; Chung et al., 2022)
- Ocean deep convection (Latif et al., 2013; Zhang et al., 2017; Cabré et al., 2017)

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Observations: limited coverage; some evidence from paleo records (Latif et al., 2013)

Models: sensitive to parameterization; captured in high-resolution models (Chang et al., 2020)

Sea level pressure

