

# **CMIP5: Uncertainties in tropical climate and carbon uptake**

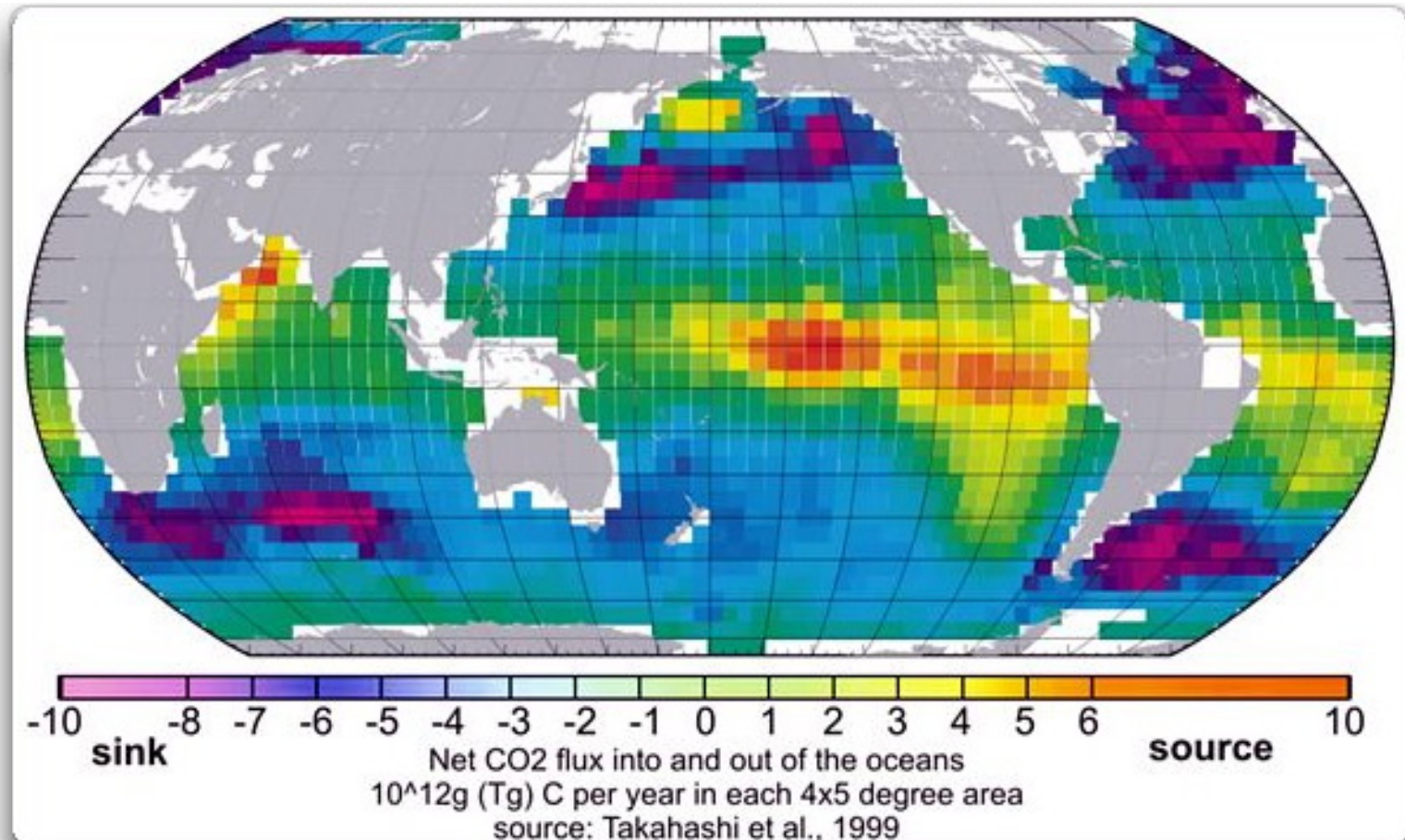
Pedro DiNezio

University of Hawaii

# Outline

- Biases in CMIP5 vs. CMIP3
  - Mean tropical climate
  - ENSO
- Role of resolution
  - Results from CESM1 high-resolution simulation
- Uncertainties in the 21<sup>st</sup> Century evolution of carbon uptake
  - CMIP5 models

# Importance of the tropical Pacific for carbon cycle

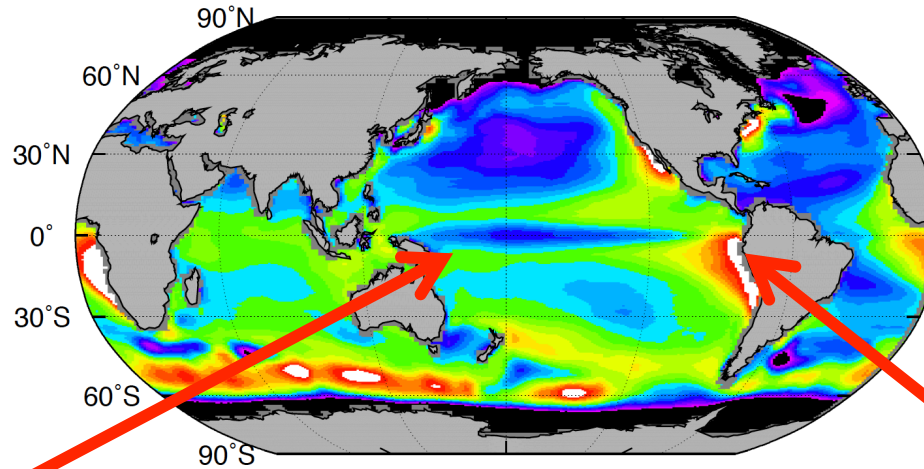


The tropical pacific is a massive source of CO<sub>2</sub>  
obvious to everybody in this room ;)

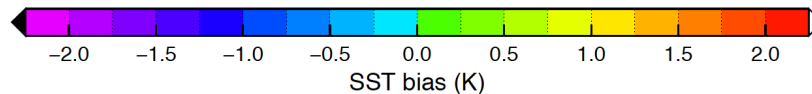
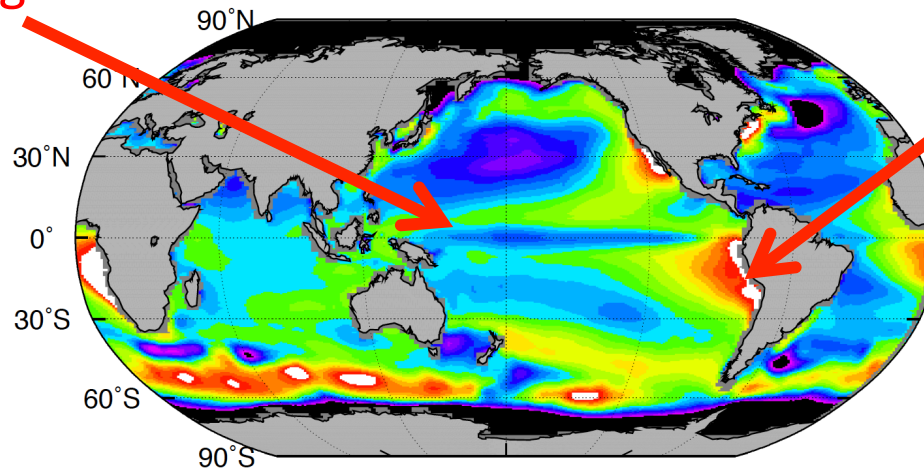
# SST biases have not improved in CMIP5

Model climatology minus ERSSTv3

CMIP3



CMIP5

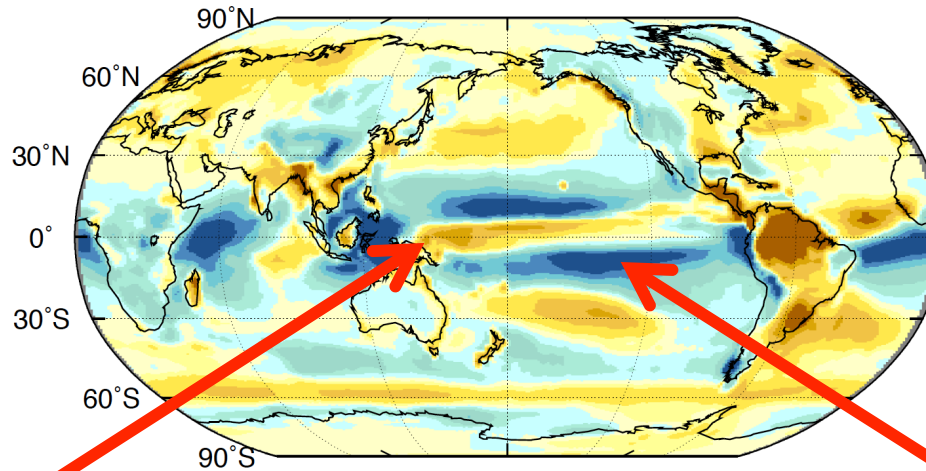




# Rainfall biases have not improved in CMIP5

Model climatology minus GPCPv2.2

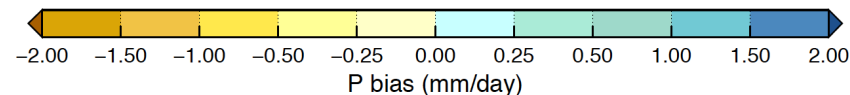
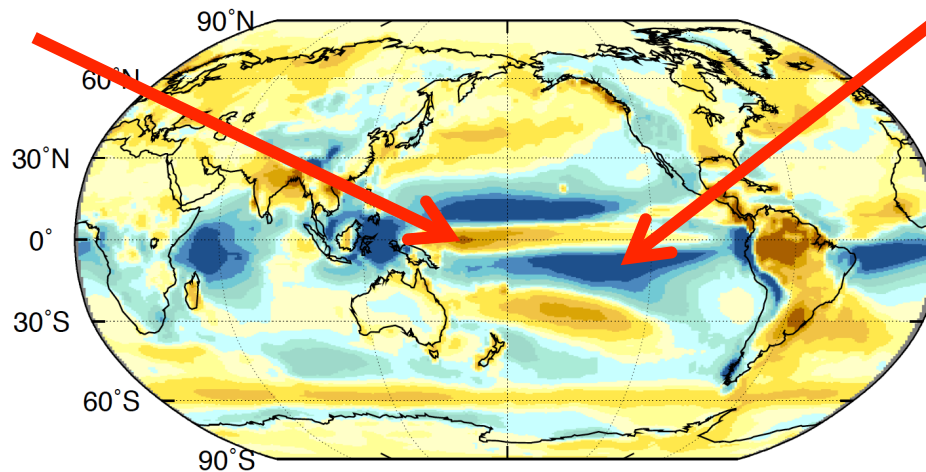
CMIP3



dry western  
equatorial Pacific

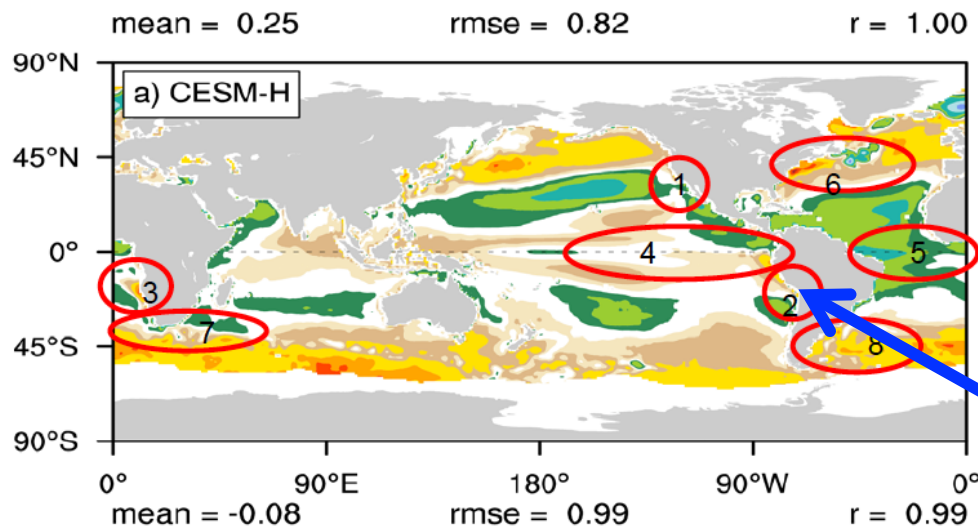
Double ITCZ bias

CMIP5



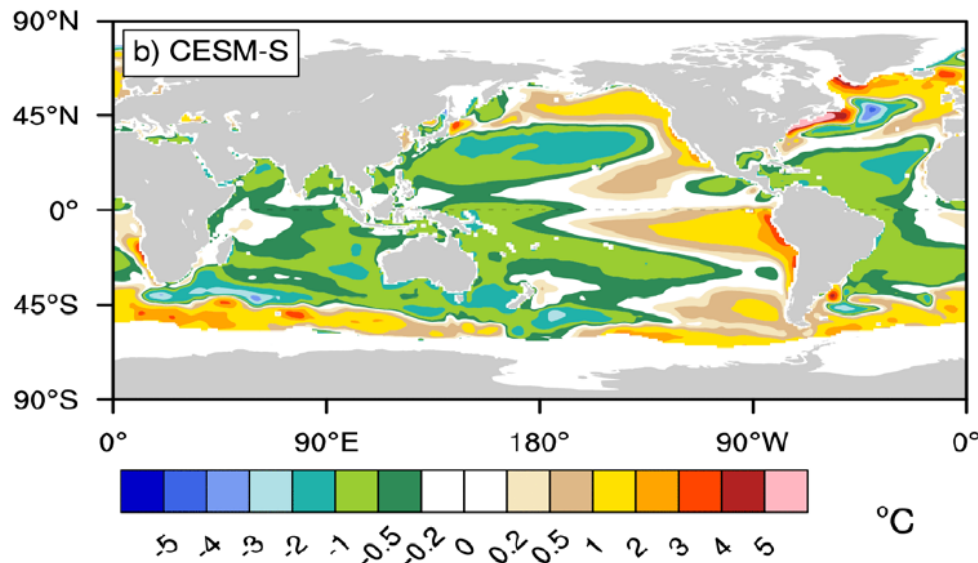
# Some of these biases maybe resolution dependent

high  
resolution  
CESM1



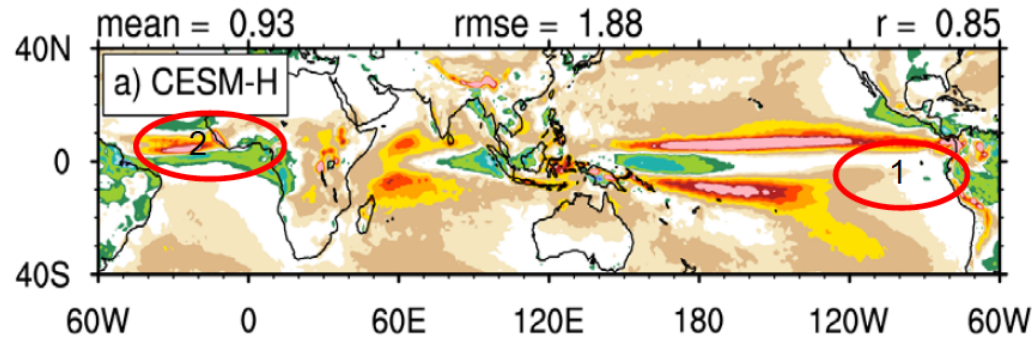
alleviated biases in  
the Pacific  
upwelling zones

standard  
resolution  
CESM1

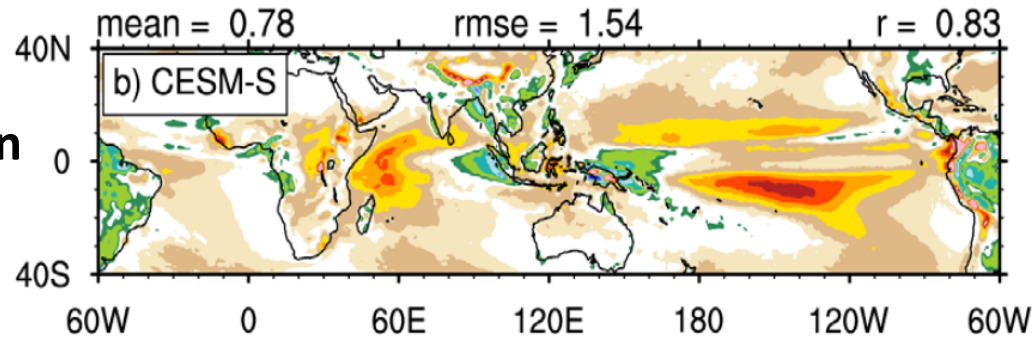


# High ocean resolution improves double ITCZ problem

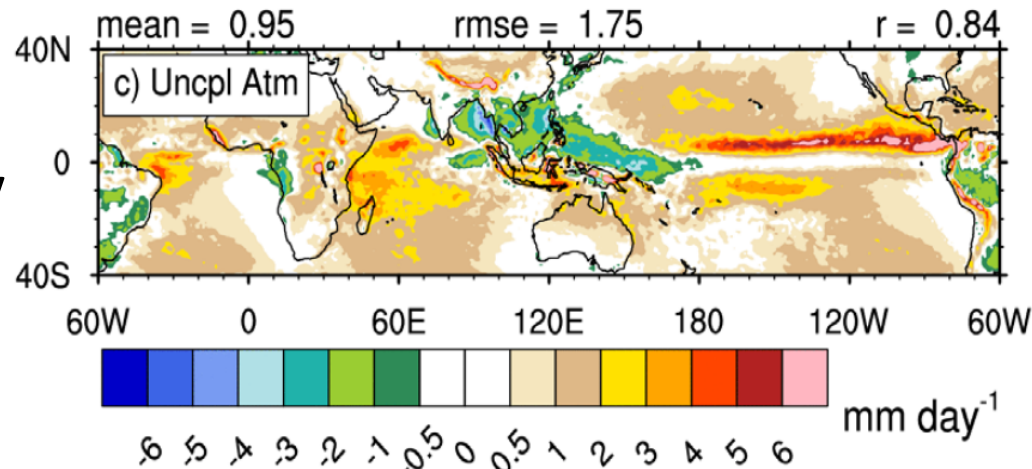
high resolution  
CESM1



standard resolution  
CESM1

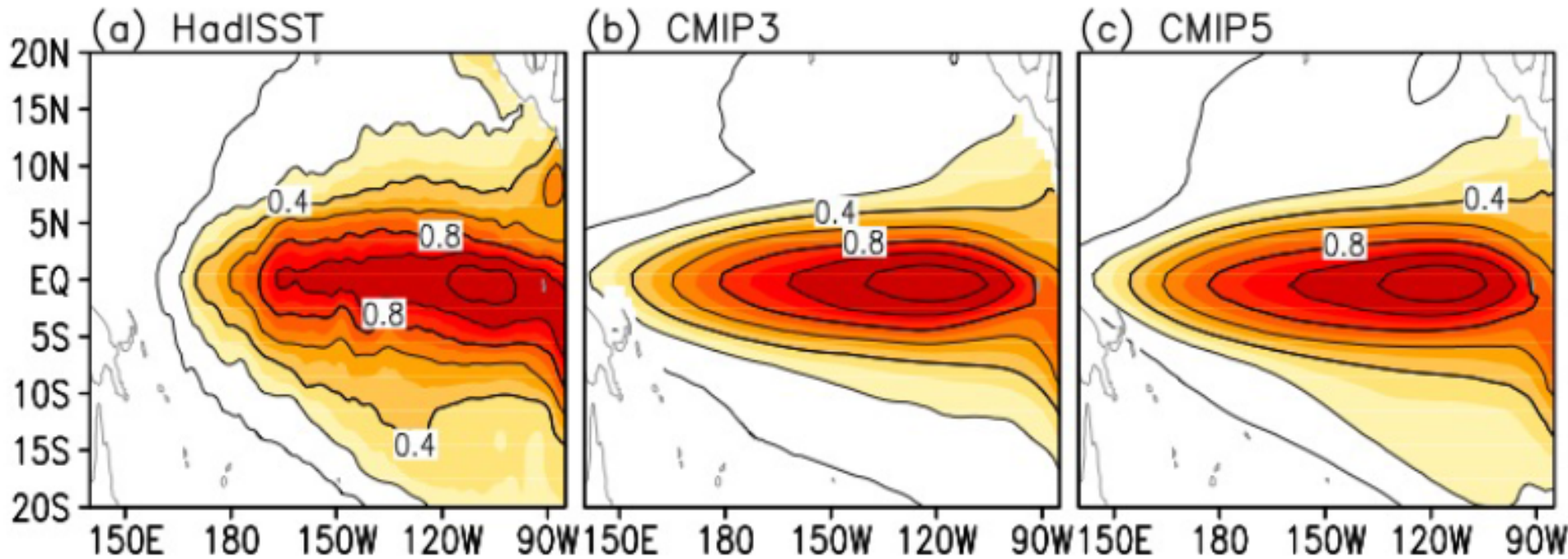


High resolution  
atmosphere-only



# ENSO simulation has not improved either

## Spatial pattern of SST anomalies

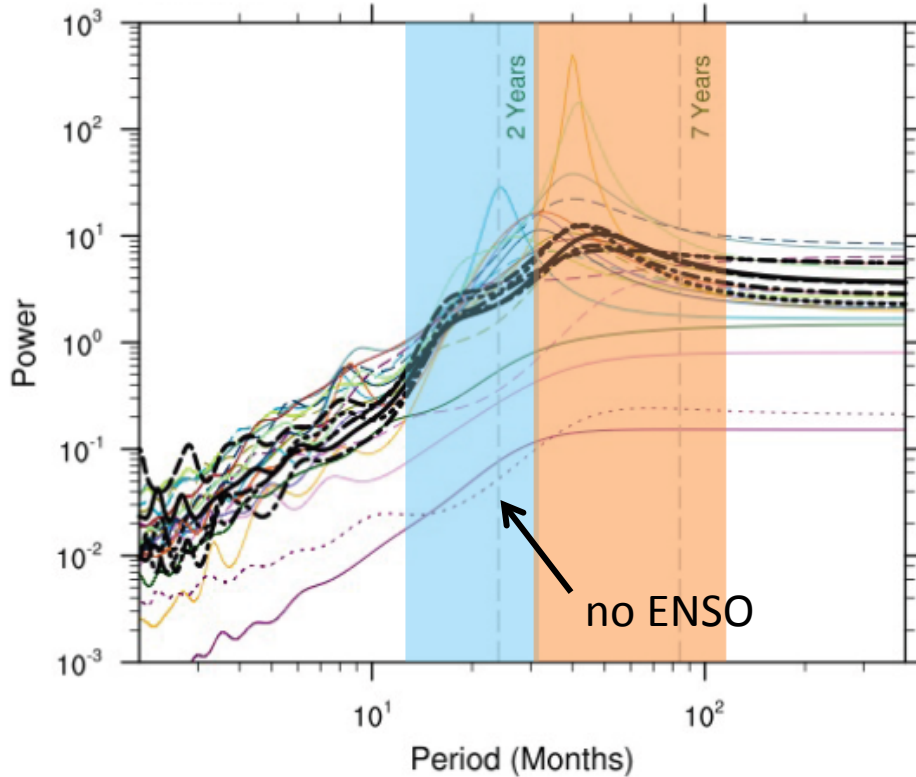


Excessive westward extension of SST anomalies  
still present in CMIP5



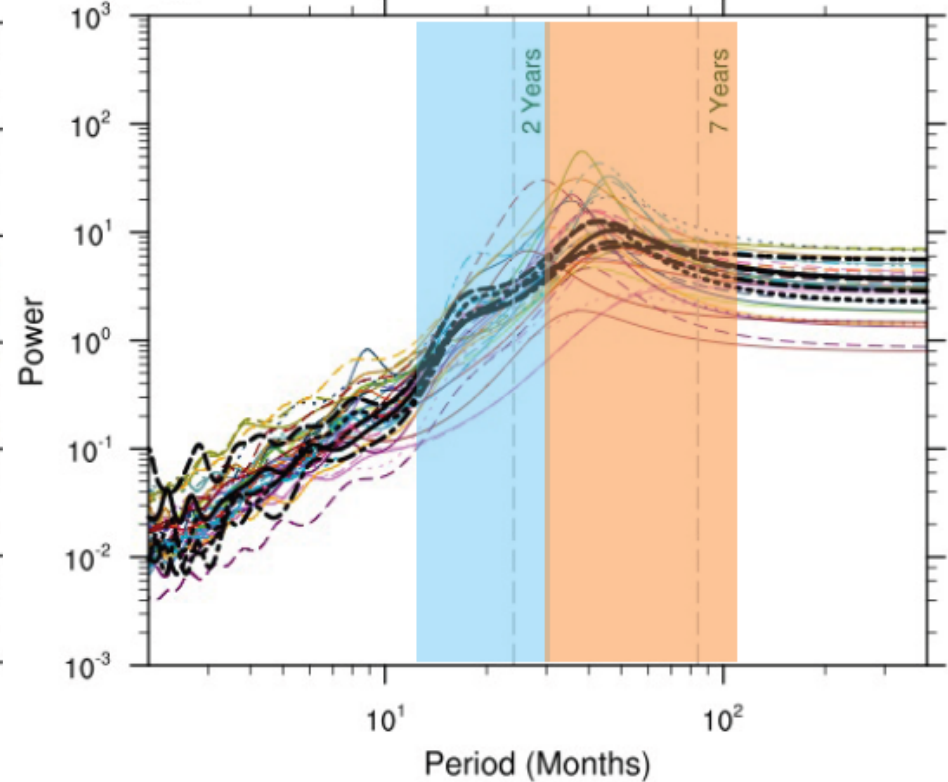
# Improved Nino-3.4 spectrum

CMIP3



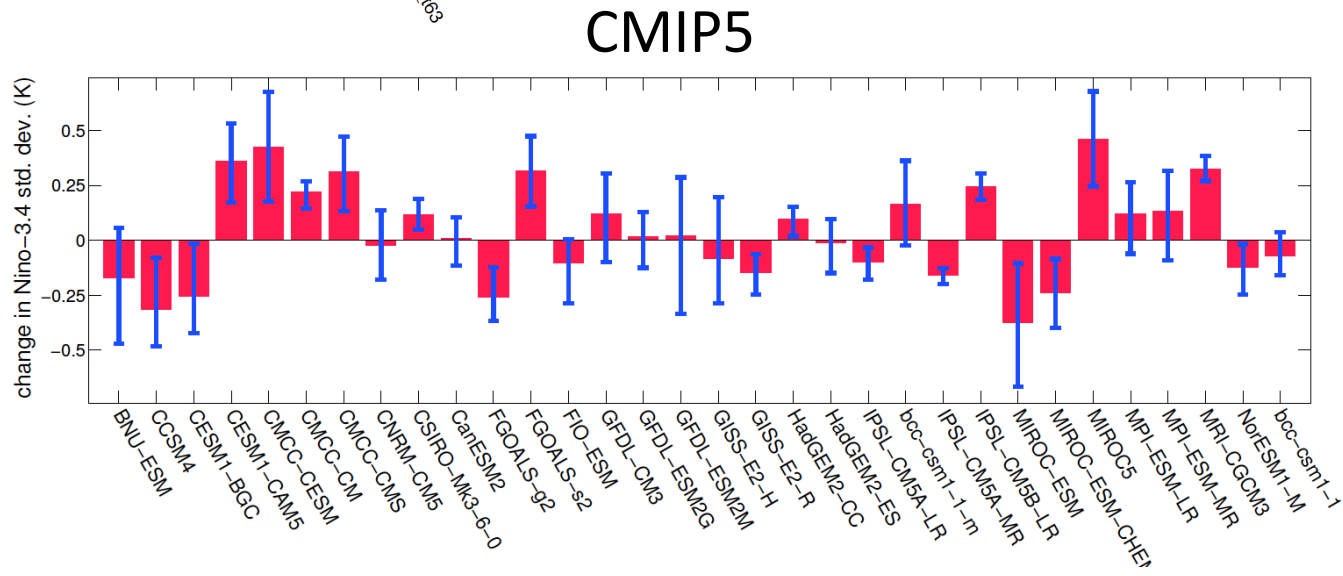
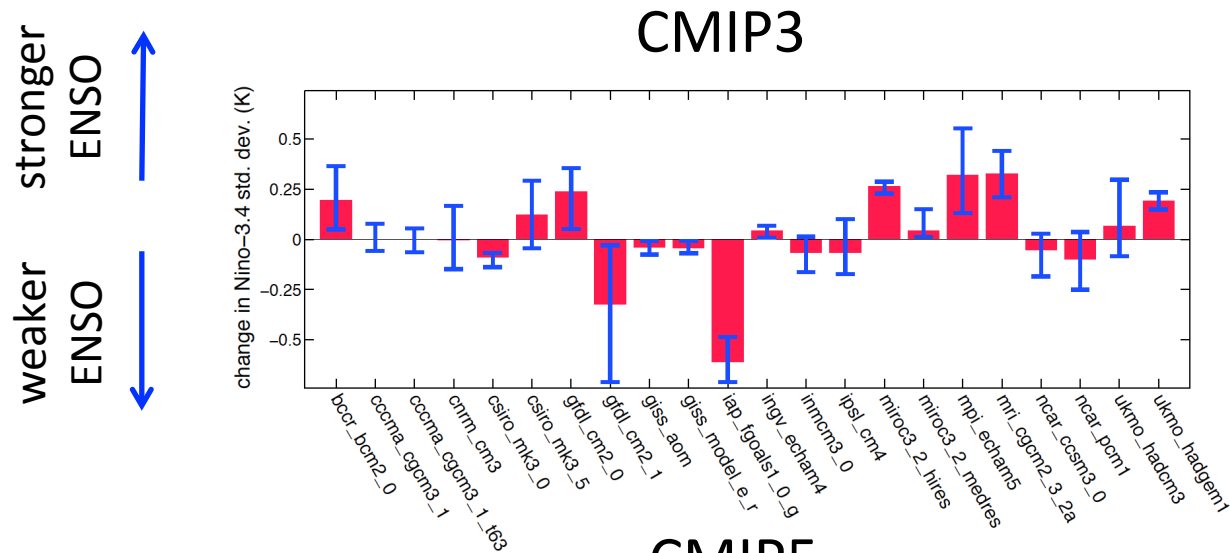
- Unrealistic sharp spectral peaks
- Some modes without ENSO

CMIP5



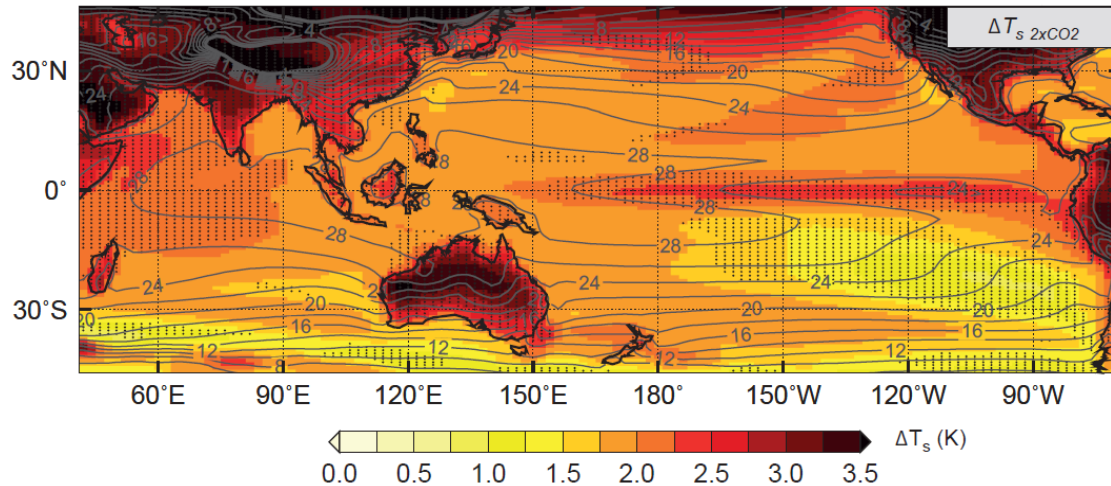
- Improved (broader) spectra in CMIP5
- No more models with no ENSO

# Models still do not agree whether ENSO will become stronger or weaker

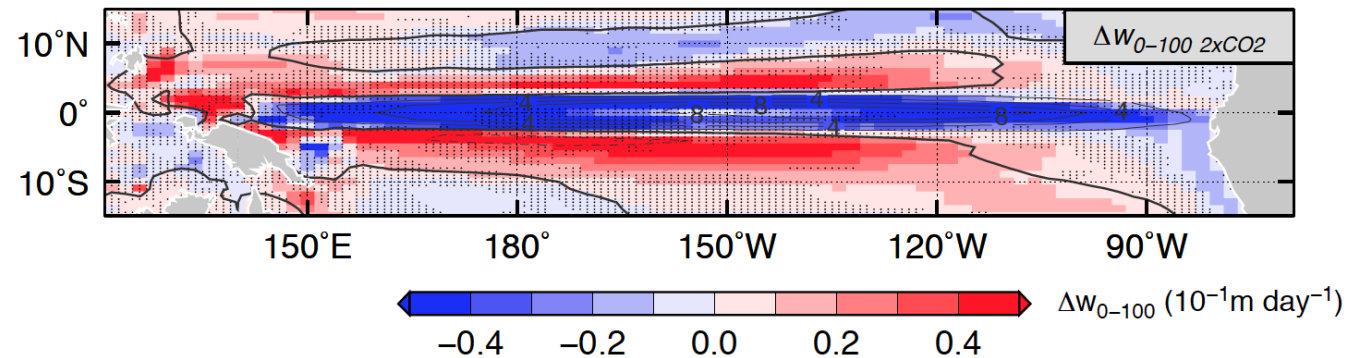


# But models agree about the mean climate response of the equatorial Pacific

Enhanced equatorial warming

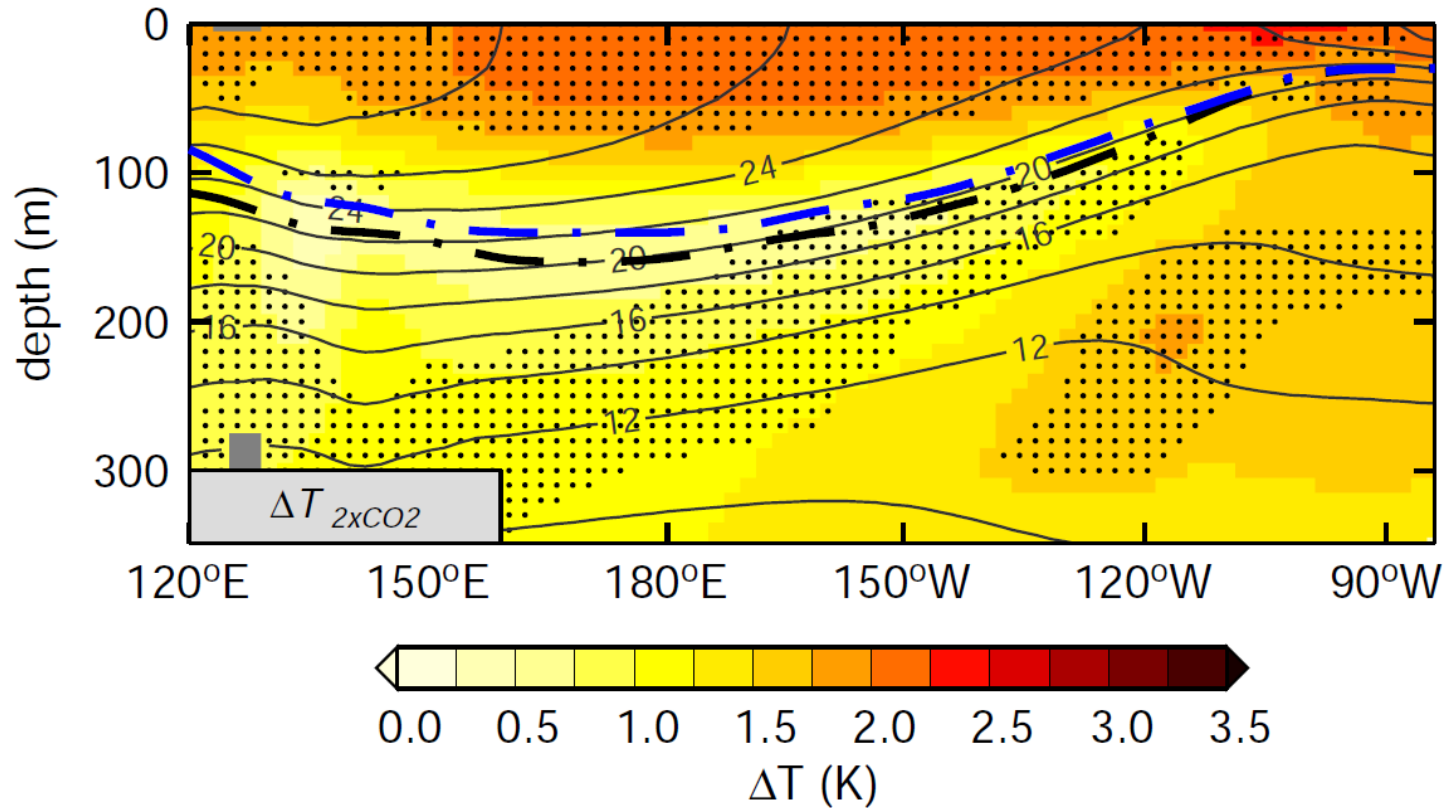


weaker upwelling





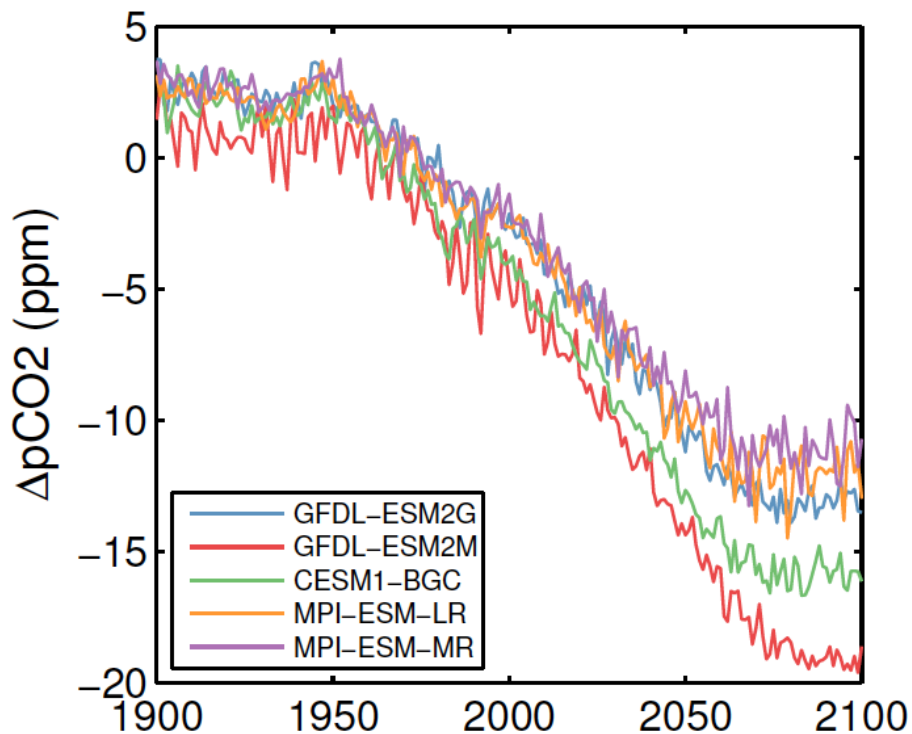
# And a sharper equatorial thermocline



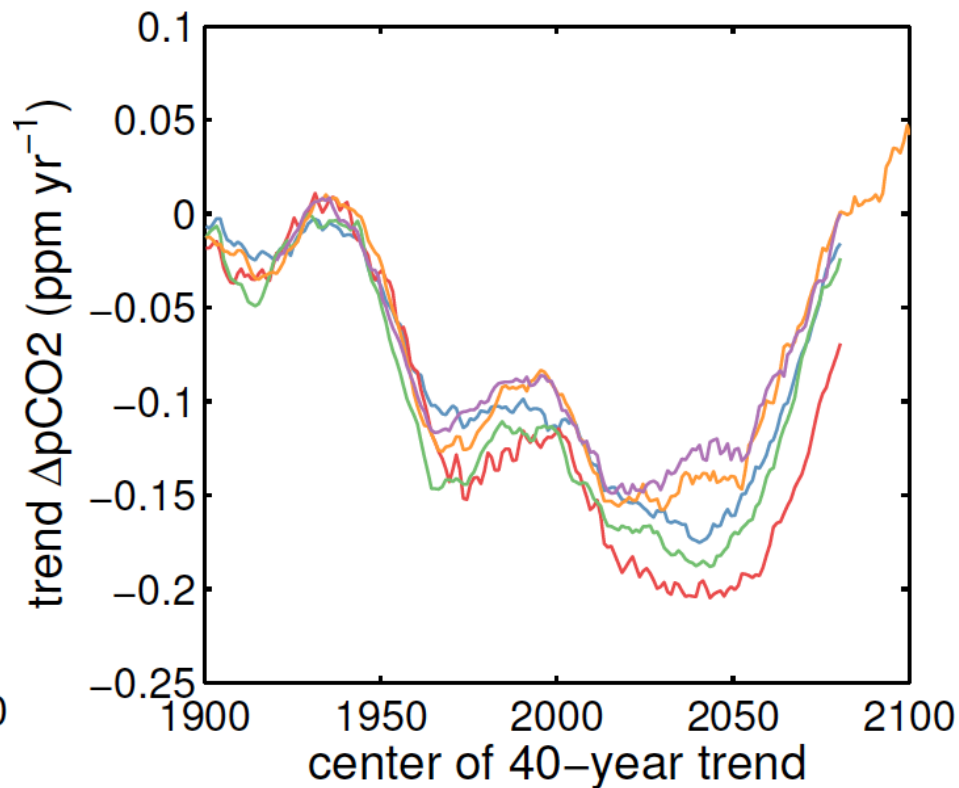
Dash-dotted black (blue) line is the thermocline in the control (2xCO<sub>2</sub>) climate.

DiNezio *et al.* 2009

# Do any of these processes play a role in the evolution of CO<sub>2</sub> uptake during the 21<sup>st</sup> Century?



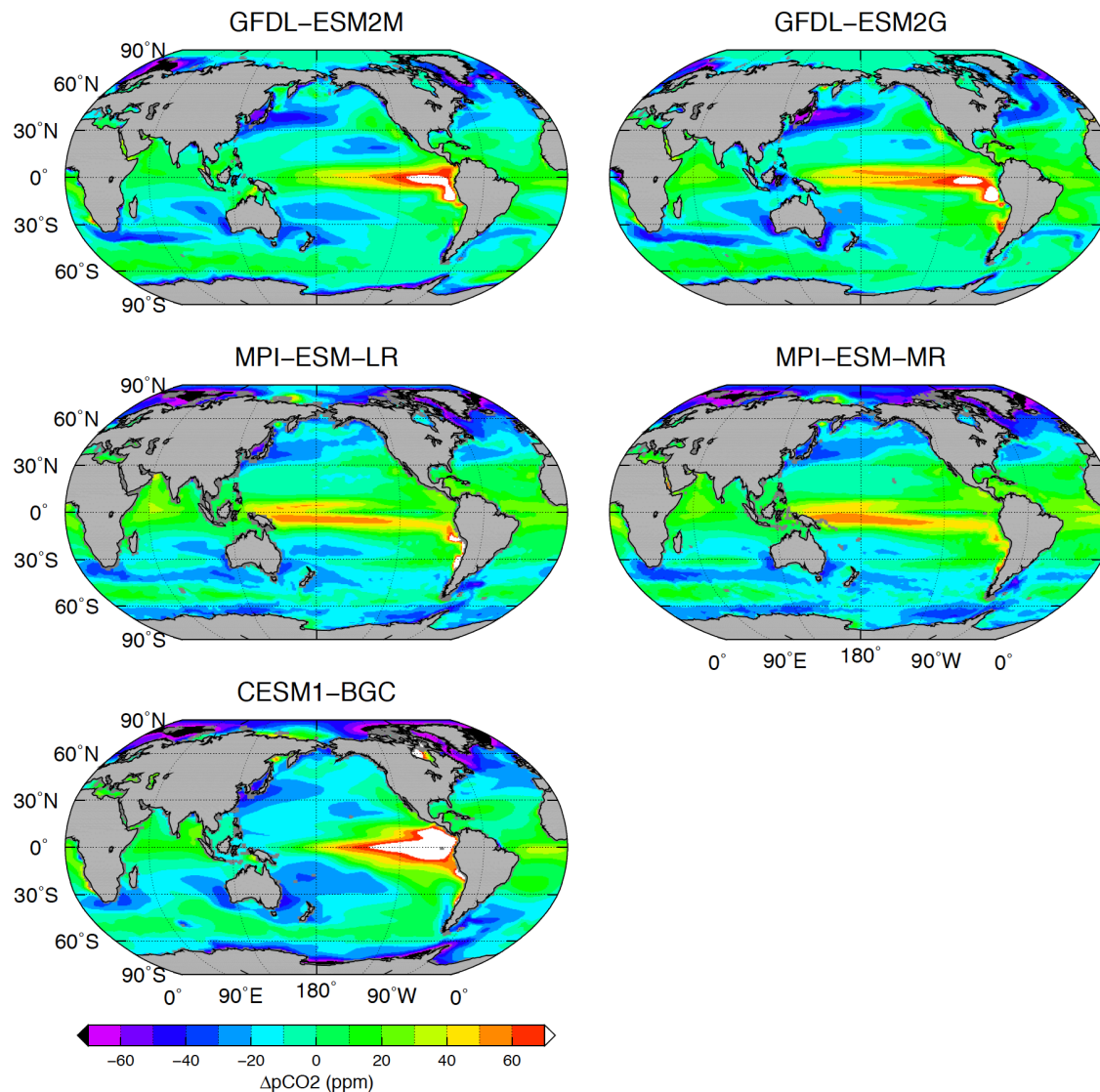
Large spread among the models  
after year 2000



All models simulate  
accelerations and decelerations  
in carbon uptake, hummm....

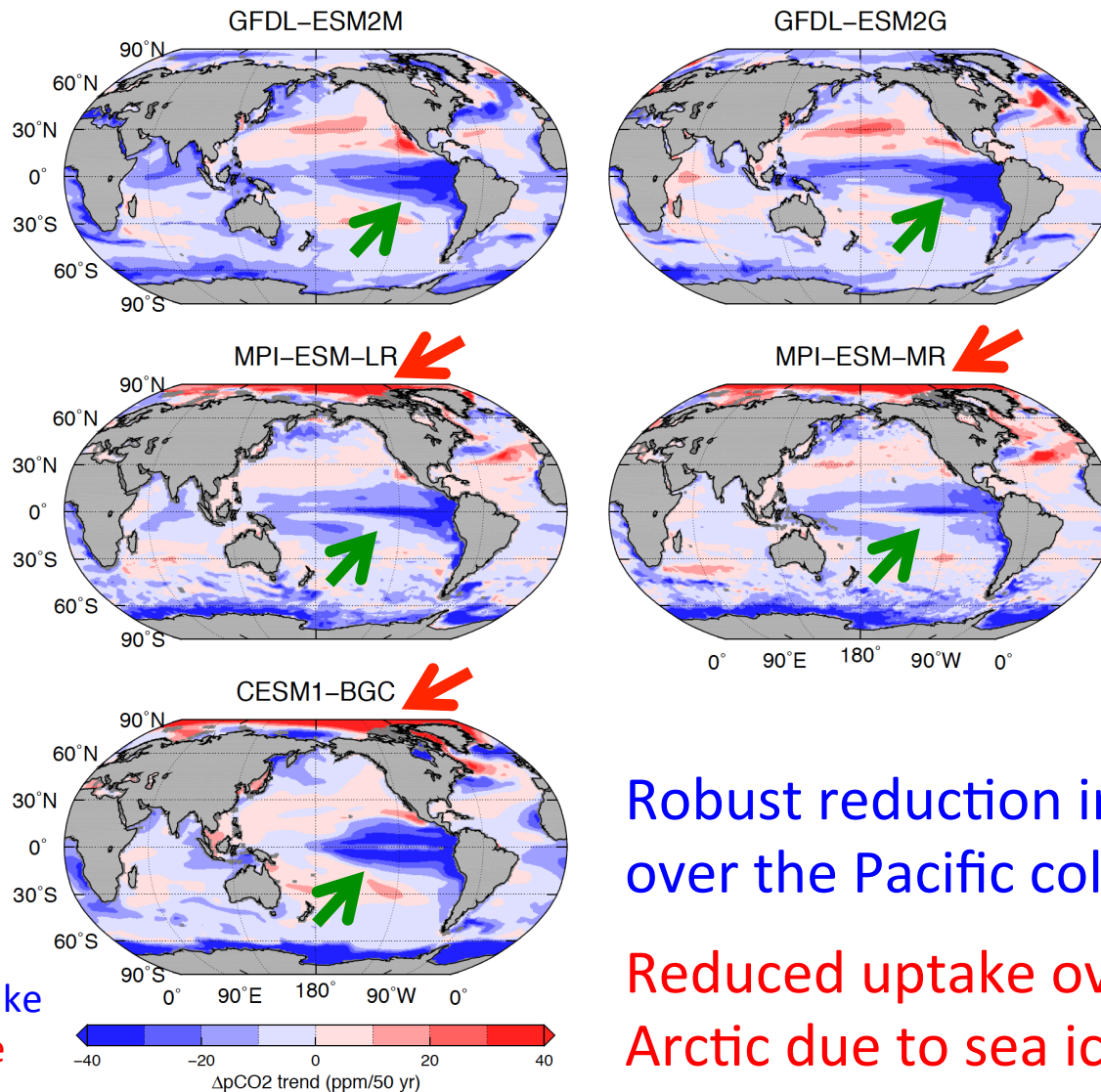
# Some models show pronounced biases in $\Delta p\text{CO}_2$ over the tropical Pacific

1991–2010  $\Delta p\text{CO}_2$  climatology



# But all of them project reduced outgassing over the tropical Pacific

2030–2080  $\Delta p\text{CO}_2$  change

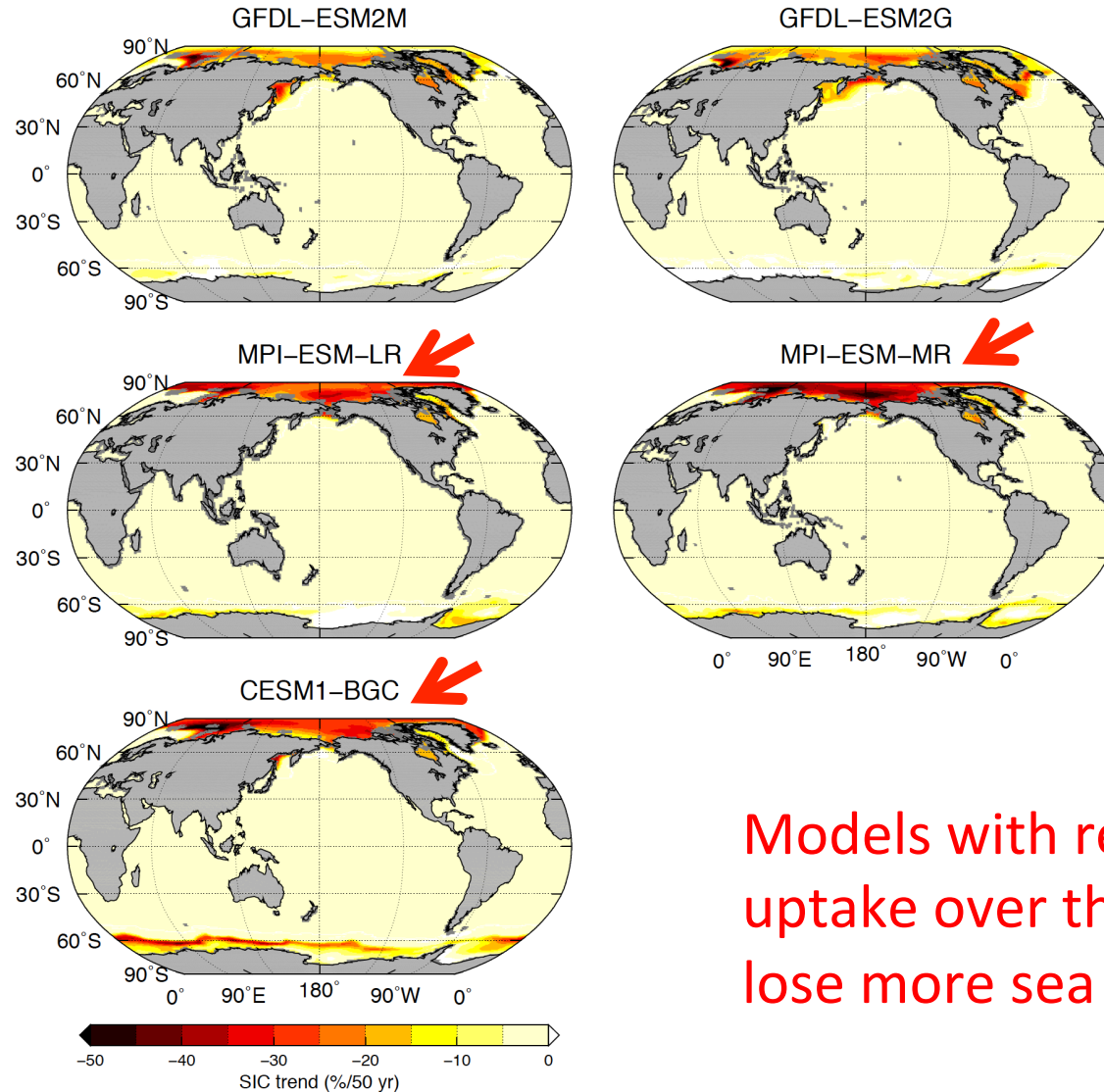


Robust reduction in outgassing over the Pacific cold tongue

Reduced uptake over the Arctic due to sea ice loss?

# Arctic sea ice loss could be a source of model spread:

2030–2080 sea ice concentration change



Models with reduced CO<sub>2</sub> uptake over the Arctic also lose more sea ice

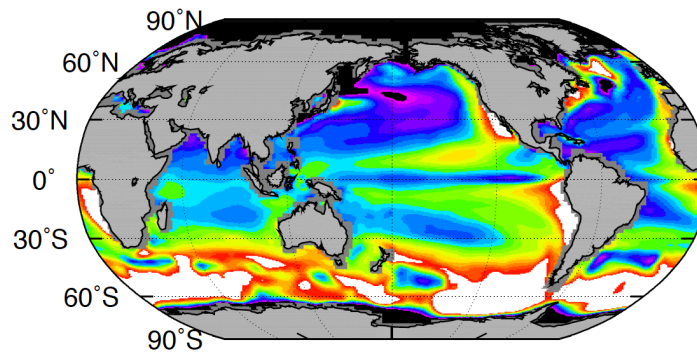


# Conclusions

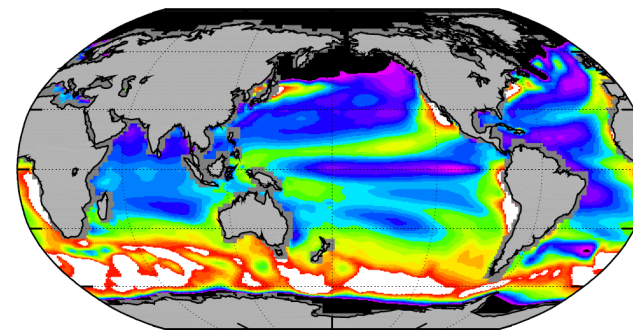
- Not much improvement in the simulation of tropical climate from CMIP3 to CMIP5
- Ocean resolution may play a role improving warm biases in upwelling zones
- Uncertainties in the 21<sup>st</sup> Century evolution of carbon uptake:
  - Interplay of aerosols and GHGs?
  - Robust reduction in outgassing over the Pacific cold tongue
  - Arctic sea ice loss could be a source of model spread

# SST bias

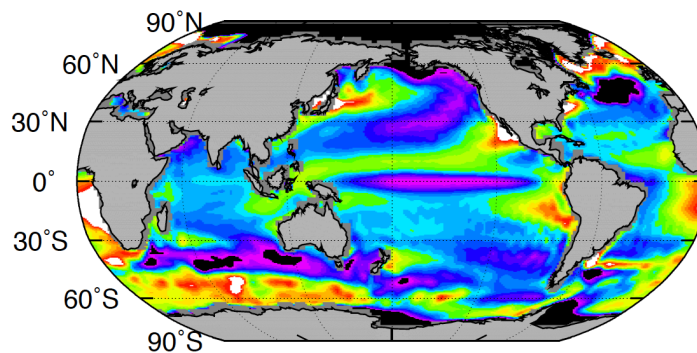
GFDL-ESM2M



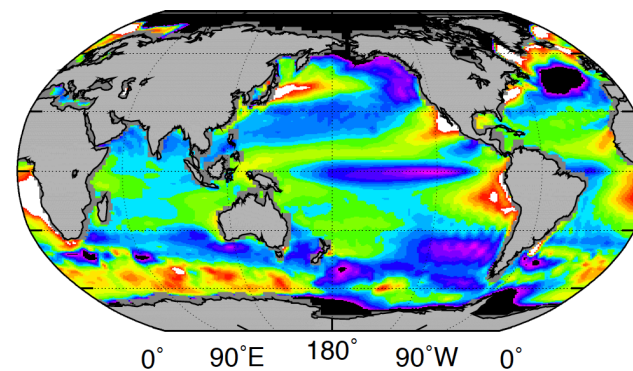
GFDL-ESM2G



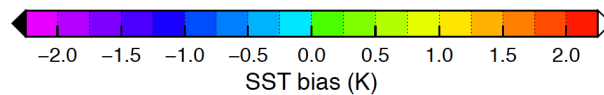
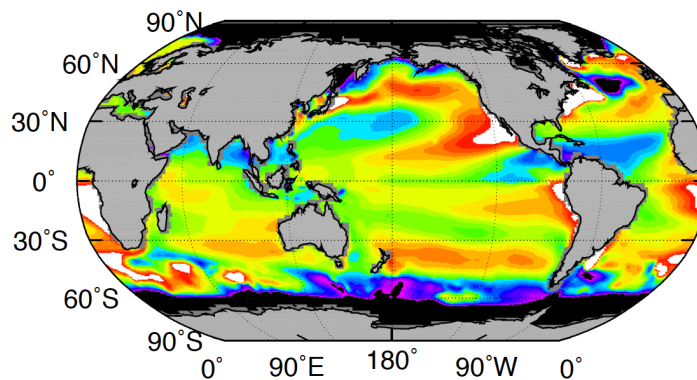
MPI-ESM-LR



MPI-ESM-MR

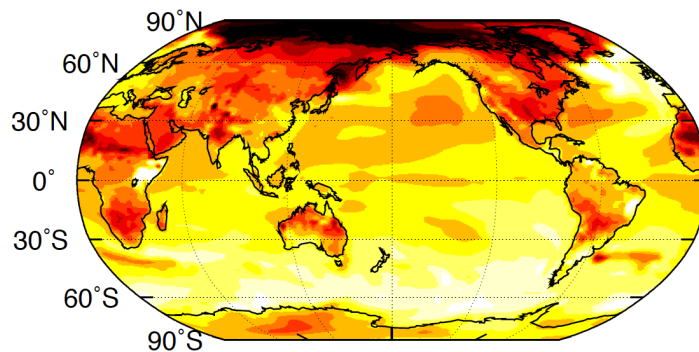


CESM1-BGC

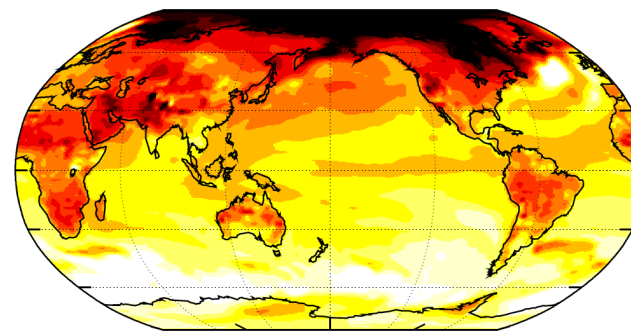


# 2030–2080 surface temperature change

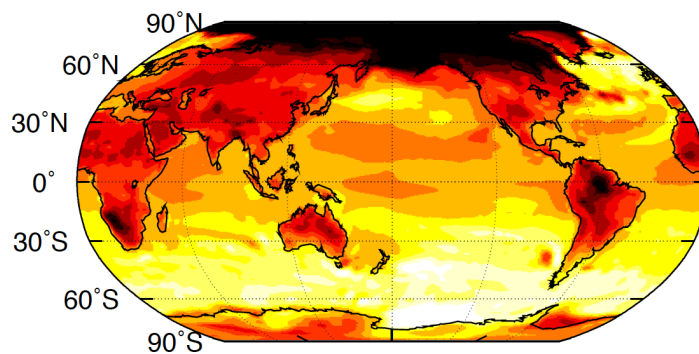
GFDL-ESM2M



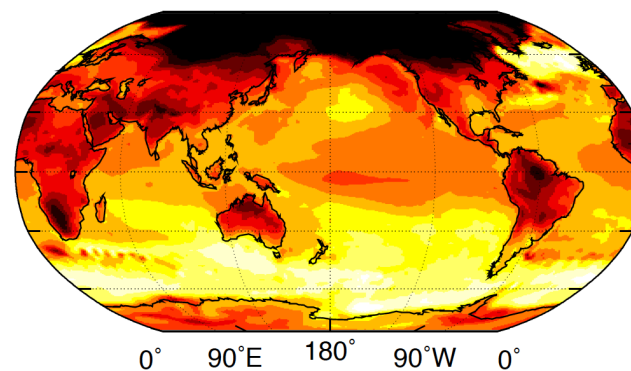
GFDL-ESM2G



MPI-ESM-LR



MPI-ESM-MR



CESM1-BGC

