

Role of the Southern Ocean in climate: heat and carbon uptake

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Current state of knowledge of model biases and uncertainties in CMIP5 models

US CLIVAR & OCB workshop on Ocean's carbon and heat uptake: uncertainties and metrics



Nutrient cycle

Sarmiento et al. (2004) Marinov et al. (2006)

The Problem

- Sediment traps suggest that ~one-third of the particulate organic matter flux at 200 m continues past the base of the main thermocline (defined as $\sigma_{\theta} = 26.8$)
- If nitrate lost by the above particle sinking were not replaced, the thermocline nitrate would be depleted within ~50 years!
- <u>QUERY</u>: How do nutrients return from the deep ocean to the thermocline and thence to the surface?



Sarmiento et al. (2004)

Deep ocean overturning circulation simplified



Based on Talley (2013)

Upwelling at 1000 m in GFDL CM2.6



Morrison (pers. comm.)





Air-sea CO₂ flux

An analysis based on CMIP5 models. Froelicher et al. (in press, J. Climate)

Pre-industrial CO₂ flux in CMIP5

Multi-model mean (11 models)





Froelicher et al. (submitted)

Cumulative anthropogenic CO₂ uptake in CMIP5

Multi-model mean (11 models)





uptake



Froelicher et al. (submitted)

Cumulative CO₂ uptake in CMIP5 models & ESM2M ensemble in PgC (1950-2005)



- Southern Ocean accounts for 44% of global in models
- Observations (inverse model) give ~50%
- Most of the uncertainty comes from S. Ocean
- Internal variability accounts for \sim 24% of range.

Frölicher et al. (2014, in press)



Excess heat uptake by ocean

An analysis based on CMIP5 models. Froelicher et al. (in press, J. Climate)

Cumulative perturbation heat uptake in CMIP5

Multi-model mean (19 models)



Cumulative ocean heat flux (1986-2005) - (1861-1880) (10⁹J/m²) -15 -10 -5 0 5 10 15

uptake

Froelicher et al. (submitted)

Cumulative heat uptake in CMIP5 models & ESM2M ensemble (1950-2005)



- Southern Ocean accounts for \sim 69% of global.
- Internal variability accounts for $\sim 19\%$ of range.

Frölicher et al. (2014, in press)

Role of WBCs

An analysis by Dufour et al., (in preparation) based on GFDL CM2.6 climate model with 0.1° ocean & simple biogeochemistry

With acknowledgement to ludicone et al. (2011), Talley (2013) and others





Volume and phosphate transport versus sigma-2





- Goal: to separate surface ocean into regions dominated by
 - upwelling of ancient waters vs.
 - WBC cooling regions
- Strategy is based on the hypotheses that
 - Region to south of Subantarctic front is dominated by upwelling
 - Region to north of Subtropical front is dominated by WBC influence
 - Region between these fronts is influenced by both upwelling and WBC

"Operational" definition for the Subantarctic front

• The SSH contour corresponding to the region just south of the deepest winter mixed layers.

Example with GFDL-ESM2M



"Operational" definition for the Subtropical front

• The SSS contour that corresponds to the maximum lateral SSS gradient in the Western Boundary current regions





Ekman divergence

Note: Upwelling region is confined to south of Subantarctic Front



Air-sea heat fluxes

Note: *Cooling regions* due to Western Boundary Current influence are confined largely to north of Subtropical Front



Cumulative anthropogenic CO₂ uptake (70 years) by region in CM2.6



	Upwelling		Cooling
	region		region
Anthropogenic CO ₂ uptake (%)	55	28	17



Application to some CMIP5 models and CM2.6



- Separation into 3 regions of the 70 year cumulative sum of anthropogenic CO₂ flux
- Note: in red regions the method does not work well (e.g. polynyas)

Conclusions

Process	Southern Ocean (south of 30°S) contribution to world	Upwelling + boundary region	Cooling region
Nutrient resupply	~ 3/4 biological production north of 30°S	~100%	~0%
Anthropogenic CO ₂ uptake	Observations ~50% Models 40-45%	~70%	~30%
''Anthropogenic '' heat uptake	~70 ± 20%	?	?