Coastal ocean response to ENSO from models and observations

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CLIVAR ENSO Meeting Scripps Institution of Oceanography

August 10-11, 2016



Mechanisms

$$\frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = K_v \frac{\partial^2 T}{\partial z^2} + K_H \left(\frac{\partial^2 T}{\partial x^2} + \frac{\partial^2 T}{\partial y^2}\right) + Q_T$$

- Local rate of change
- Advection
- Mixing
- Surface heat flux

- Atmospheric teleconnections
- Coastally trapped waves
- Advection



California Underwater Glider Network

- Climate effects on region
- Lines 90, 80, since <u>2006</u>
- Line 66.7, since <u>2007</u>
- Data distributed in real time, assimilated in several models
- 28 glider-years, 200,000 km, 95,000 dives



• Gridded climatology





Mean on line 90





- Cores of poleward undercurrent
- Equatorward California current
- Maximum poleward flow at surface

Mean on line 66.7





- Poleward undercurrent
- Equatorward California current
- Undercurrent maximum near 100 m

Annual cycle of temperature on line 90



- Surface driven by heat flux from atmosphere
- Thermocline driven by wind stress (curl)





Data Assimilative CCS Reanalyses 1980-2010 (Andy Moore)

Table 1

A summary of the observation types, observing platforms, data sources, the nominal measurement errors, and the period covered.

Observation Type	Observing platform	Source	Instrument error	Period covered
SSH	Altimeter	Aviso, 1 day average	0.04 m	1993–2012
SST	AVHRR/ Pathfinder	NOAA Coast watch	0.6 °C	1981–2012
SST	AMSR-E	NOAA Coast watch	0.7 °C	2002–2010
SST	MODIS-Terra	NASA JPL	0.3 °C	2000–2012
Hydrographic data	Various	UK Met Office	0.5 °C for T 0.1 for S	1950–2012



- ROMS 4D-Var (PSAS)
- SODA BCs
- Controls: IC, BCs, Forcing
- 8-day assimilation cycles
- 4-day overlap

 $t_{0}^{j} + 8$

 t_0^{j+2}

1 outer, 15 inner loops

 $t_0^{j+1} + 8$

 $t_0^{j+2} + 4$

 $t_0^{j+2} + 8$

Model vs. wind-estimated upwelling



What does the CCS look like during El Niño?

- Coastal averages (2 degree meridionally and 50 km from shore)
- W: Vertical transport across 40 m depth
- d_{26.0}: depth of the 26.0 kg m⁻³ isopycnal
- d_s: a measure of depth of upwelled source water 30 days prior
- σ_s : a measure of density of source water 30 days prior



Interannual variability

Significant lagged correlations of anomalies with ONI

- W: 1 months (suggests atmospheric teleconnections)
- d_s: 2 months (suggests indirect atmospheric teleconnections)
- d_{26.0}: 2 months (suggests indirect atmospheric teleconnections)



The 2015-16 El Niño

-125

- Among strongest events by tropical Pacific surface indices
- Subsurface signature much weaker
- Local CCS surface signature much weaker

CCS SST anomalies (NOAA OISST)





Tropical Pacific Niño 3.4 and d_{20.0}



2015-2016 El Niño showed weak CCS response (from combined glider data/model reanalysis)



Temperature anomalies



Temperature anomalies vs. time, depth



- Inshore 200 km
- Line 90: 2009-10 El Niño
- 2014-15 anomalies near surface
- 2015-16 anomalies penetrate deeply, especially on 90

Isopycnal salinity anomalies on line 90 (evidence of advection)



- Fresh on 25.3 during 2014-2015
- Salty on 26.0 during 2015-2016





Summary

- CCS has a robust annual hydrographic cycle at both surface and thermocline depths.
 - o Driven by local surface fluxes.
- CCS interannual variability is strongly linked to tropical Pacific anomalies.
 - o Atmospheric teleconnections (weeks).
 - o Oceanic waveguide influences (months).
- Predictive skill for local impact of ENSO events based on trpical Pacific indices is significant, but individual responses do vary.
- 2015-2016 event locally much weaker than expected, in part due to pre-warming by 2014-2015 warm blob and in part due to different character of this years El Niño.



