## Relationship between Air-Sea Density Flux and Isopycnal Meridional Overturning Circulation Igor Kamenkovich (RSMAS, Miami, FL), Myeonghee Han (current affiliation unknown), Timour Radko (Naval Postgraduate School, Monterrey, CA), and William E. Johns (RSMAS, Miami, FL)

# AMOC, using the IPCC AR4 model projections of the 21<sup>st</sup> century climate.

Hemisphere and pull out of the deep ocean in the Southern Hemisphere:



$$PP(\sigma) = \frac{1}{2} [B_s - B_n]_{\sigma_{max}}^{\sigma} + \frac{1}{2} (\Psi_a^{30S} + \Psi_a^{65})$$

$$B_{s,n}(\sigma) = \lim_{\Delta \sigma \to 0} \frac{1}{\Delta \sigma} \iint_{\Delta \sigma} D(x, y) dx dy$$

**PP** =  $\Psi_{a}$  if the diapycnal mixing is negligible and stratification is steady.

change simulation with the (GFDL) Climate Model version 2.1 (CM2.1).

## References

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- PP and AMOC at the equator  $(\Psi_a^{0})$  are similar
- the linear trends in the maximum PP and  $\Psi_{a}^{0}$ are similar as well (0.04-0.05 Sv  $yr^{-1}$ );
- contribution of the surface heat flux to PP is greater than the contribution of the
- exchanges with the Southern Ocean play a large role in variability of the PP mode

The overall agreement between the push-pull mode and actual isopycnal AMOC suggests a largely adiabatic, pole-to-pole response of the