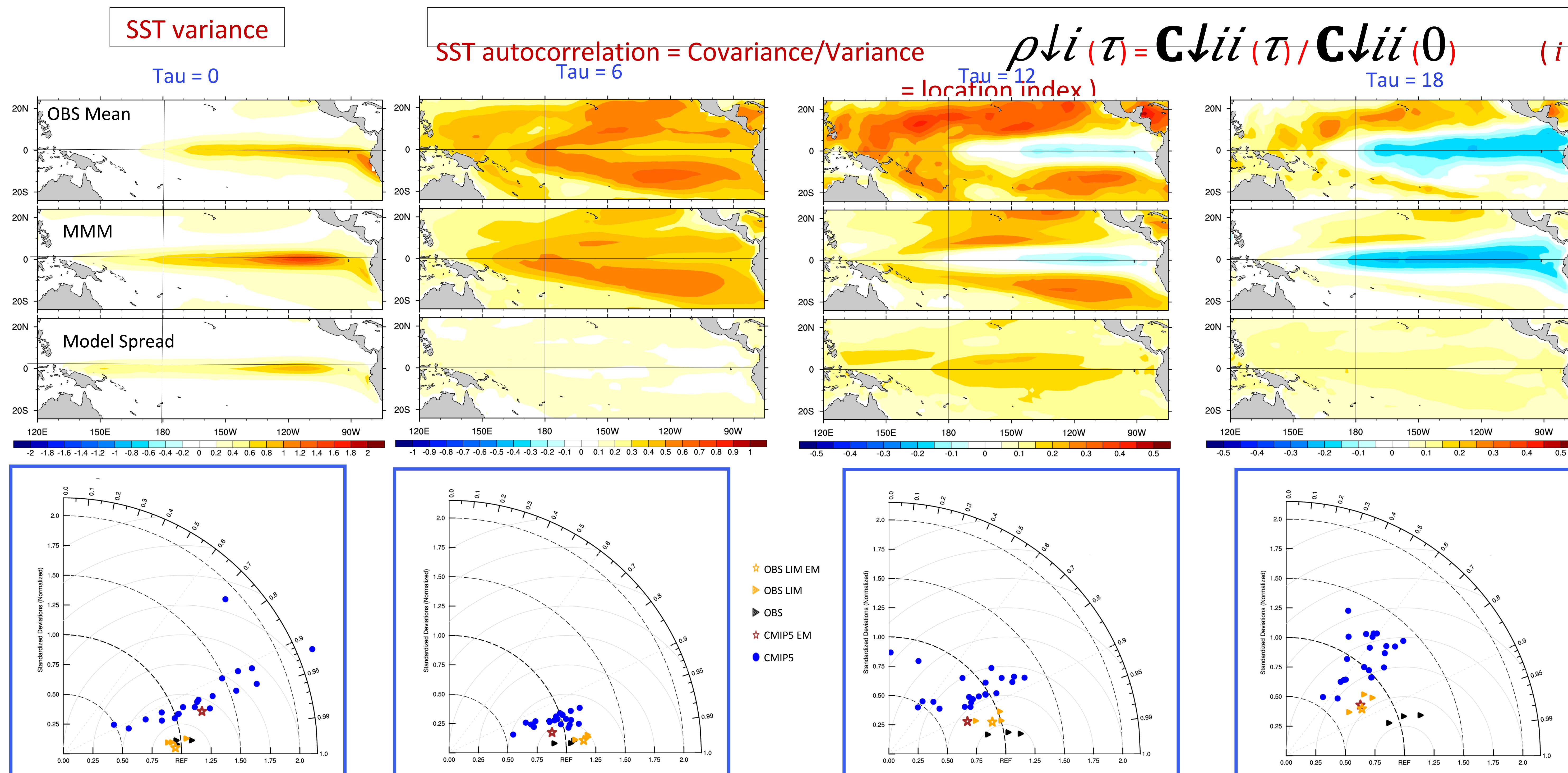


## Persistence of Tropical SSTs in Climate Models

**Antonietta Capotondi<sup>1</sup>, Matt Newman<sup>1</sup>, Prashant D. Sardeshmukh<sup>1</sup>, Yan Wang<sup>1</sup>, and Andrew T. Wittenberg<sup>2</sup>**

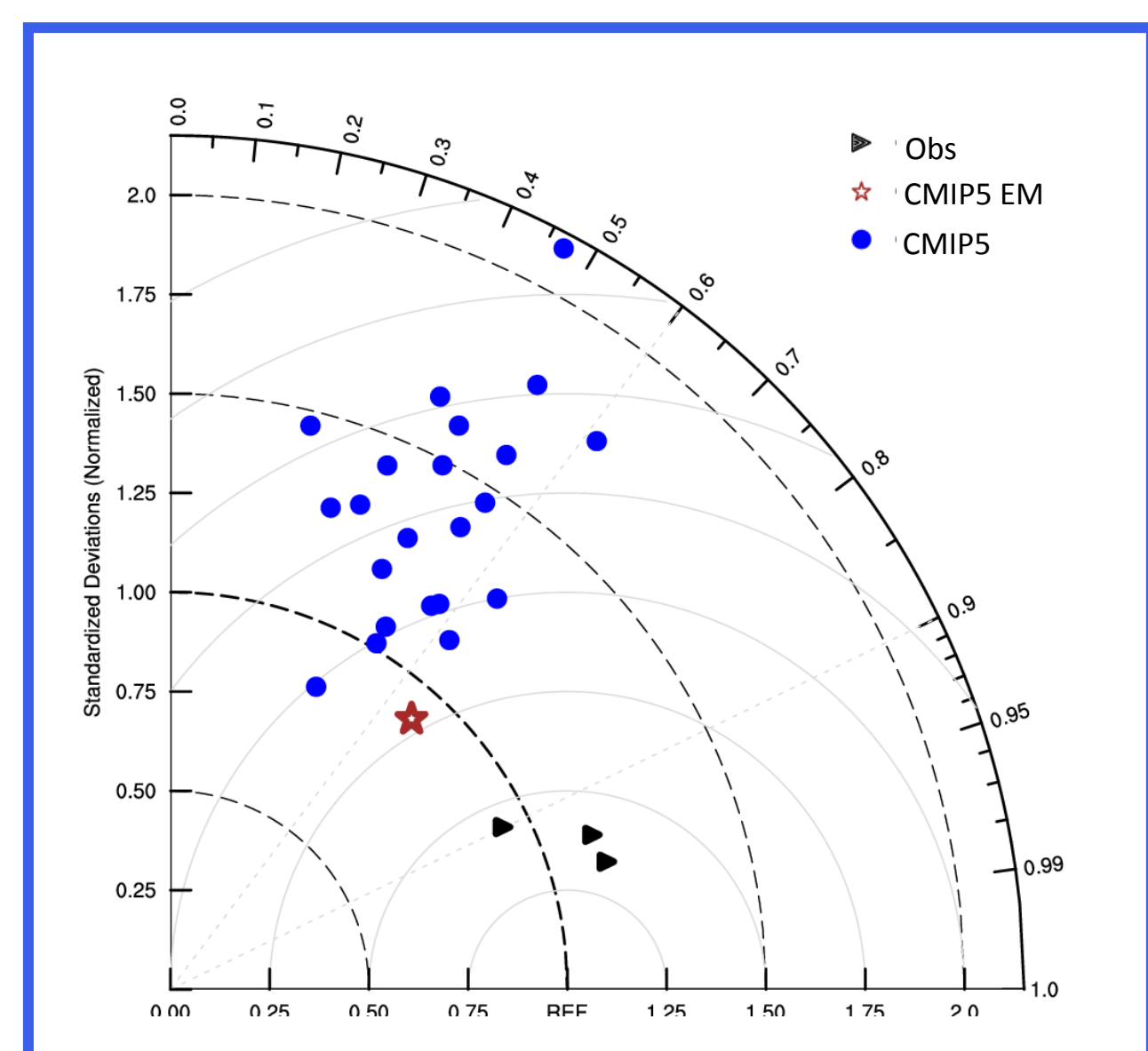
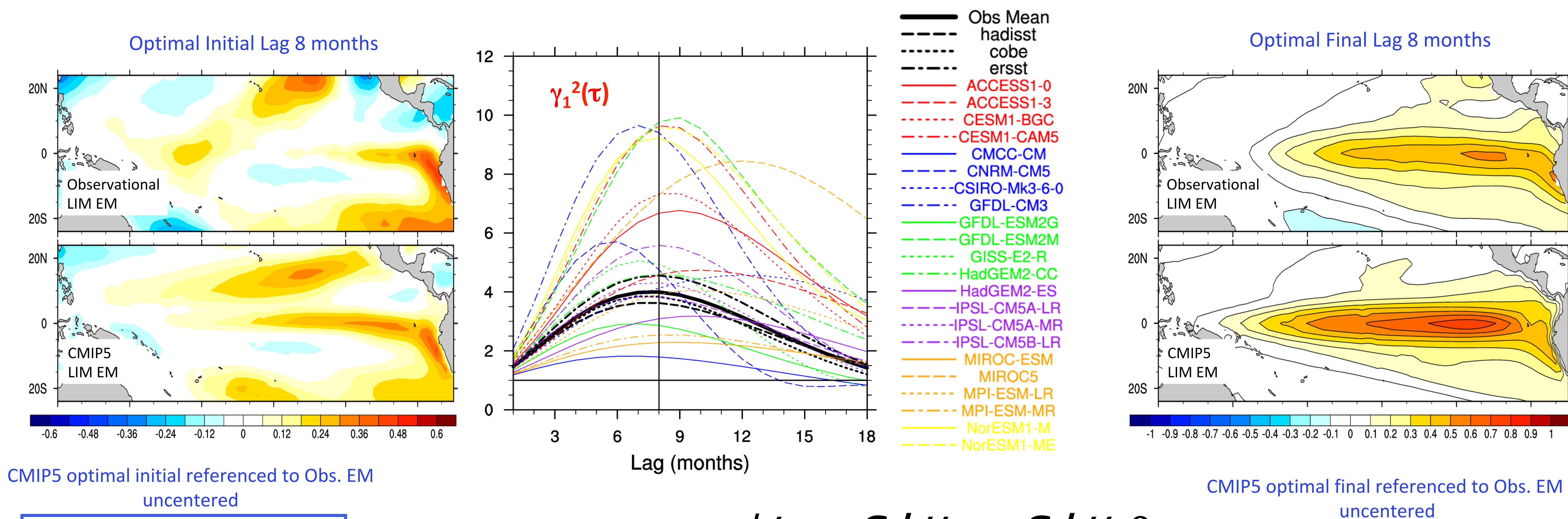
**University of Colorado/CIRES and NOAA Earth System Research Laboratory – Physical Sciences Division, Boulder CO**  
**NOAA Geophysical Fluid Dynamics Laboratory, Princeton, NJ**

Dynamically meaningful metrics should capture the *evolving* nature of ENSO, not just static variance. Lag-covariances provide both spatial pattern and temporal evolution, which are the expression of the system dynamics.



Observational and model monthly SST anomalies have been detrended, and a 3-month running average has been applied prior to the analysis. Taylor diagrams use observational ensemble mean (EM) as reference point and are uncentered (include spatial means).

Linear Inverse Modeling (LIM) captures the system dynamics



Local autocorrelation:  $\rho_{ii}(\tau) = \mathbf{C}_{ii}(\tau) / \mathbf{C}_{ii}(0)$

Define  $\mathbf{G}(\tau) = \mathbf{C}(\tau) \mathbf{C}(0)^{\dagger} - 1$       Note!  $\mathbf{G} \downarrow ii$   
 $(\tau) \neq \rho \downarrow i(\tau)$

$$\text{LIM: } \mathbf{G}(\tau) = [\mathbf{G}(1)]^\tau$$

Maximum anomaly growth over time interval  $\mathcal{T}$  is achieved

Models show a large spread in the optimal initial states, as well as in the  $\gamma_1^2$ , representing differences in their dynamics. Analyses of current and projected climate variability need to account for errors in model dynamics.

(using SVD of  $\mathbf{G}(\tau) = \mathbf{U}\boldsymbol{\gamma}\mathbf{V}^T$ ), so that domain variance increases by a factor of  $\gamma_1^2$

