



The role of sea ice and buoyancy fluxes in shaping glacial and modern overturning circulation

Alice Marzocchi and Malte Jansen

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MOC: a Southern Ocean perspective



Last Glacial Maximum

~21k yrs ago lower CO₂ (~190 ppm) lower temperatures (~3-6 °C cooler) lower sea level (~125 m less) higher ocean salinity (~1 psu) extensive <u>ice sheets</u> growth





→ expanded Artic and Antarctic sea ice cover

compilation from Roche et al. (2012)

Modern vs LGM overturning circulation (Atlantic)

δC¹³ (Western Atlantic) Ω -1000 upper cell ΆΔΙΫ -2000 Depth (m) IADW -3000 abyssal -4000 cell -5000 PI AABN PI 0 1.6 shallower upper cell -1000 AAW upper cell -2000 Depth (m) -3000 abyssal expanded cell -4000 abyssal cell -5000 LGM LGM credit: Jess Adkins -60 -50 -40 -30 -20 10 20 30 40 50 60 70 -10 Ν S Latitude Ν S

Curry and Oppo (2005)

What caused the inferred changes in ocean circulation?

Mechanistic "chain of events" from idealized simulations:

LGM atmospheric cooling → increased Antarctic sea ice formation → increased buoyancy loss rates → increased abyssal stratification → upward shift of upper overturning cell (NADW)



Key: Antarctic sea ice/buoyancy loss. See also: Shin et al. (2003); Ferrari et al. (2014); Watson et al. (2015); Jansen & Nadeau (2016)

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Can this mechanism be verified in fully-coupled simulations? Paleoclimate Modelling Intercomparison Project (PMIP)

Marzocchi and Jansen (GRL – pending minor revisions)



NEGATIVE B: atmospheric cooling and brine rejection

POSITIVE B: atmospheric heating and precipitation

B: buoyancy fluxes = heat fluxes + freshwater fluxes





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AMOC



CCSM3 (PMIP

AMOC

0

OSS

density

What about other models? (PMIP3/CMIP5)

Principal Component Analysis

"chain of events": surf. air temp. \rightarrow sea ice extent \rightarrow buoyancy loss rate \rightarrow stratification \rightarrow AMOC



Method #1: no missing data (CCSM3, CCSM4 and MRI) **Method #2:** missing data for at most one of the five variables (CCSM3, CCSM4, MRI, FGOALS, MPI, MIROC, CNRM) Method #3: all models (CCSM3, CCSM4, MRI, FGOALS, MPI, MIROC, CNRM, IPSL, GISS)

includes all PI and LGM simulations







- CCSM3
 CCSM4
 MRI
 FGOALS
- MIROC
- MPI
- ◀ IPSL
- ▼ CNRM
 - GISS







bias #1: lack of deep-ocean equilibration



CCSM4 model, LGM simulation

bias #2: insufficient Antarctic sea ice formation/export



--- observations from Roche et al. (2012)



Observational compilation (Roche et al., 2012)

Take-home messages

- Importance of Antarctic sea ice (buoyancy loss) in setting deep-ocean stratification and AMOC strength/depth.
- Discrepancies between paleo-proxies and simulated LGM ocean circulation and between different PMIP models can largely be reconciled:

#1: biases due to different simulation of Antarctic sea ice formation/export
#2: discrepancies amplified by short integration times (transient response)
#3: effect of other compensating forcings (often model-specific)

- Same models used for future projections → same issues! (PMIP4/CMIP6)
- What about a warmer world? Disappearance of Antarctic sea ice?

 Proposed mechanism would suggest AMOC strengthening/deepening BUT that is equilibrium response! Transient is likely AMOC weakening