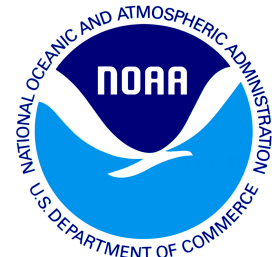


Modulation of Arctic Sea-Ice Loss by Atmospheric Teleconnections from Atlantic Multi-Decadal Variability (AMV)

Gokhan Danabasoglu

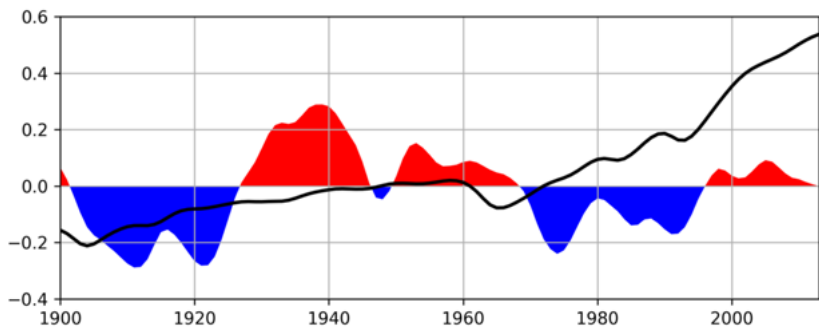
Frederic Castruccio, Yohan Ruprich-Robert,
Steve Yeager, Rym Msadek, and Thomas Delworth



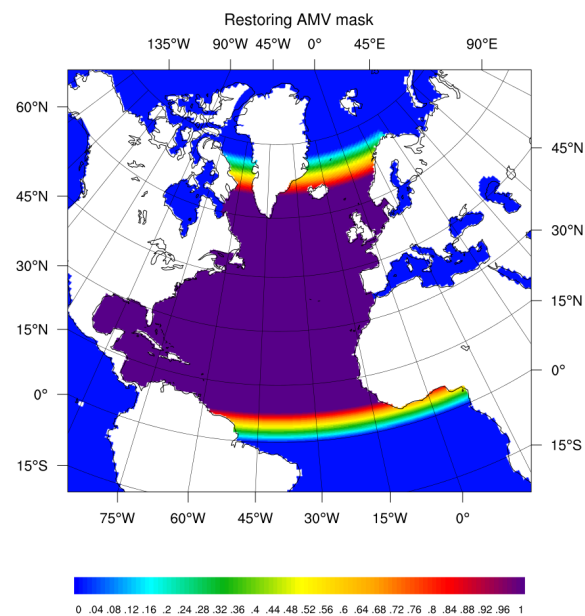
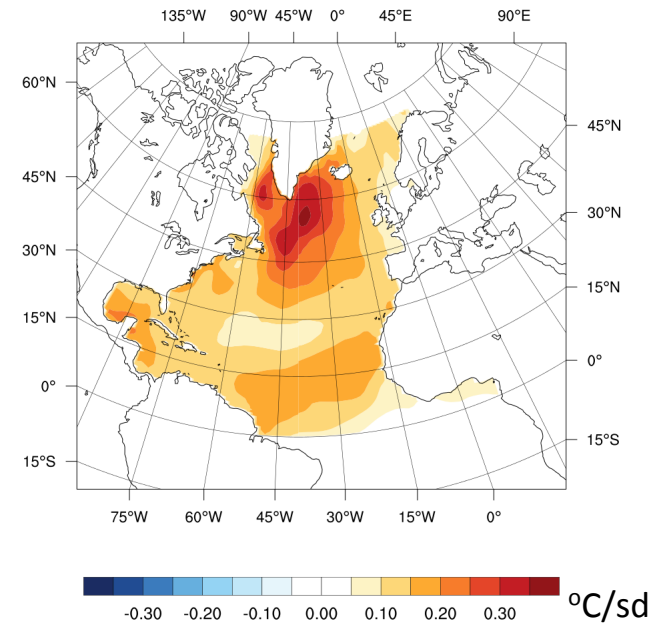
Background and Motivation

- The present study represents a component of our overarching goal of documenting climate impacts of sea surface temperature (SST) variability associated with AMV.
- We follow an experimental protocol designed to isolate impacts from atmospheric teleconnections that result from imposed SST anomalies, i.e., the dynamical adjustments of the ocean are minimized.
 - Global impacts (Ruprich-Robert et al. 2017, *J. Climate*)
 - Impacts on North American summer climate and heat waves (Ruprich-Robert et al. 2018, *J. Climate*)

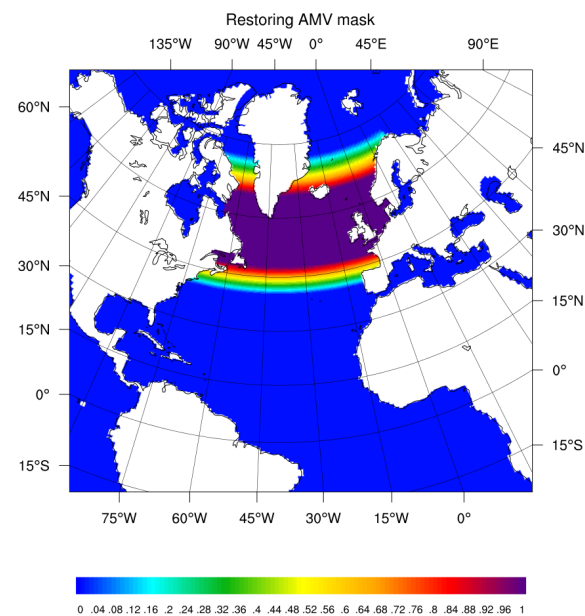
Experimental Setup



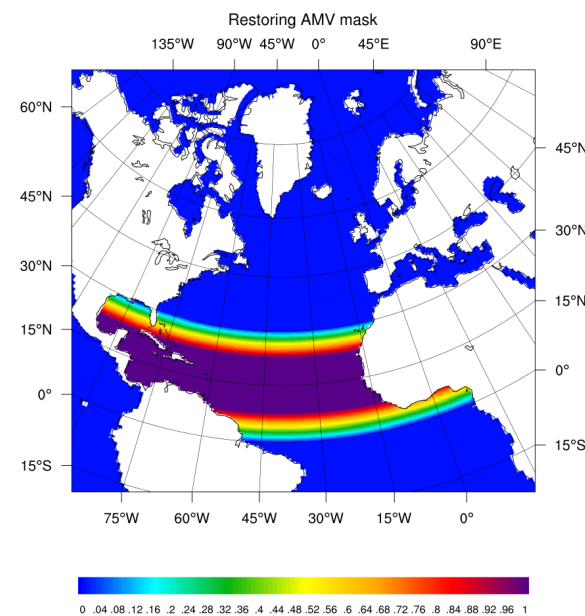
AMV pattern (based on Ting et al. 2009)



Full NA restoring



SPNA restoring



TA restoring

Time-independent SST anomalies corresponding to 1 SD of the AMV index are added to (subtracted from) the model daily climatological SSTs for the AMV+ (AMV-) experiments.

The restoring time scale is usually 5 days over 10 m.

10-year simulations under pre-industrial conditions: long enough for atmospheric teleconnections to arise, yet short enough to limit oceanic drift issues in the North Atlantic.

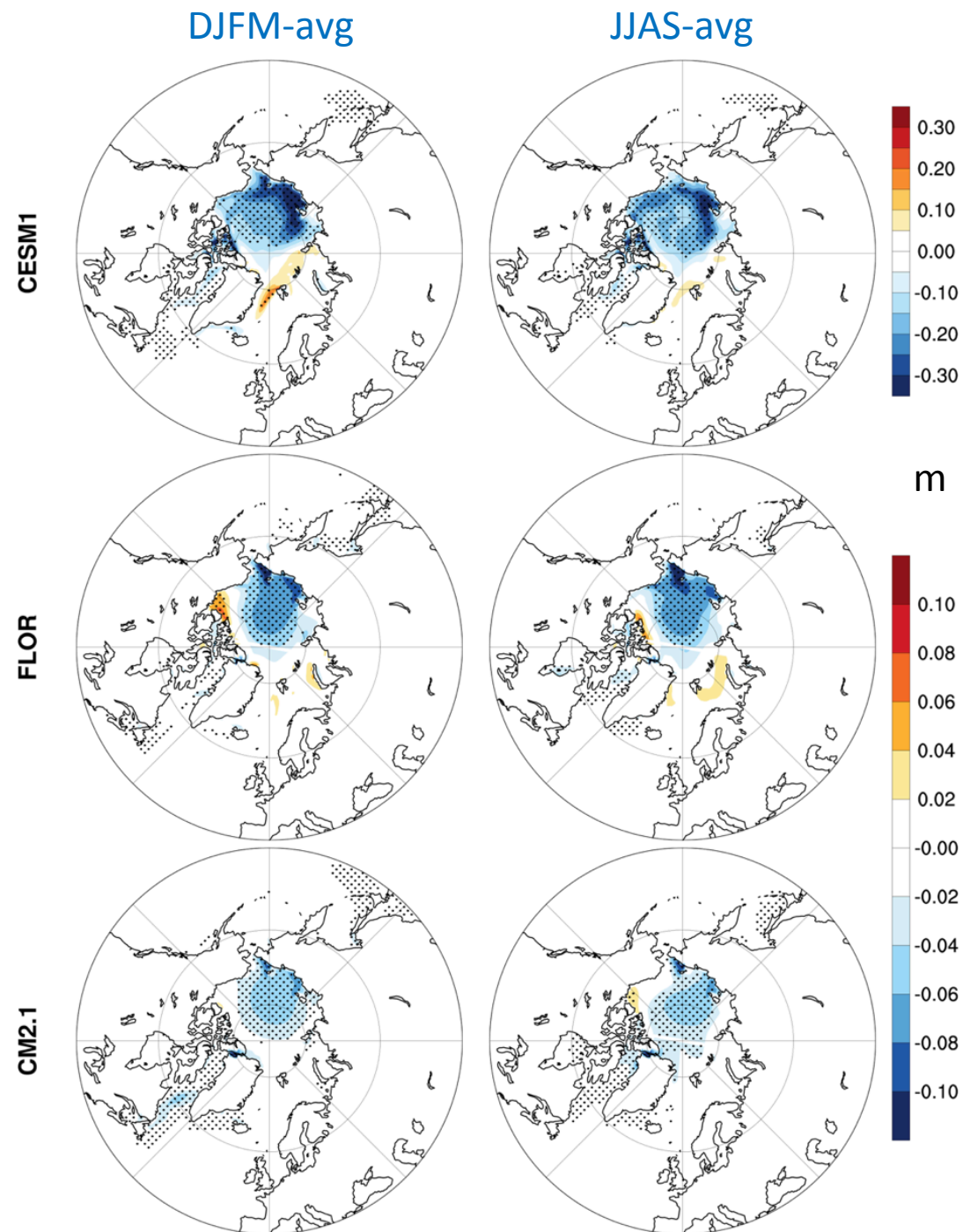
Models

- Community Earth System Model version 1 (CESM1): 30 members
- GFDL Forecast-oriented Low Ocean Resolution (FLOR): 50 members
- GFDL Climate Model version 2.1 (CM2.1): 100 members

All three models use nominal 1° horizontal resolution in their ocean components, but employ different atmospheric resolutions: 2° CM2.1; 1° in CESM1; and 0.5° in FLOR.

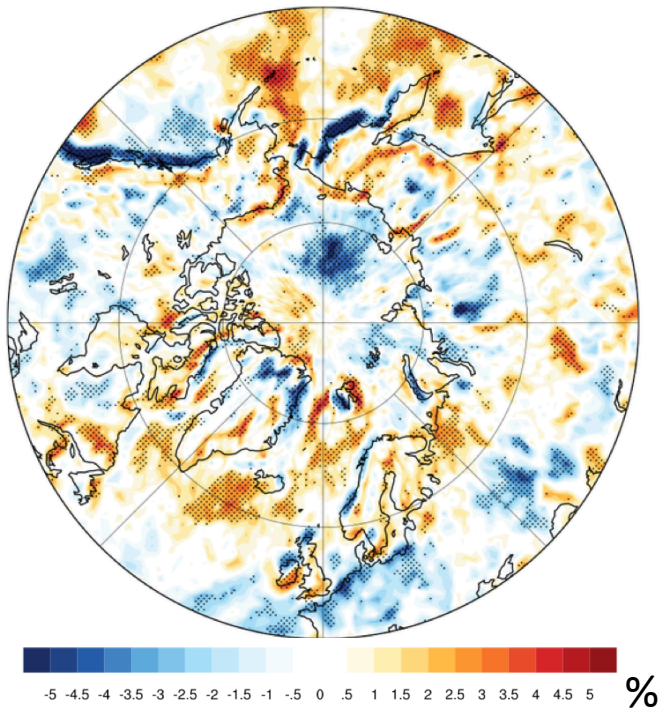
10-year average, AMV+ minus AMV- ensemble-mean differences are shown.

Differences in Sea-Ice Thickness



Differences in Sea Level Pressure and Winds

Differences in DJFM Surface
Anticyclonic Winds



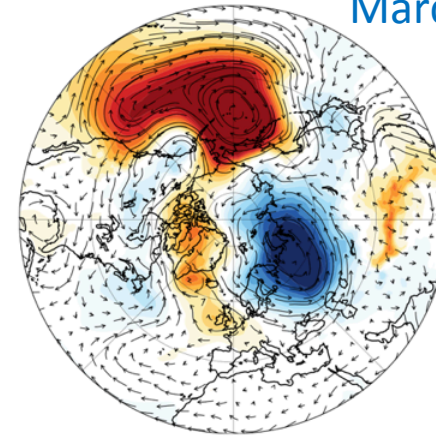
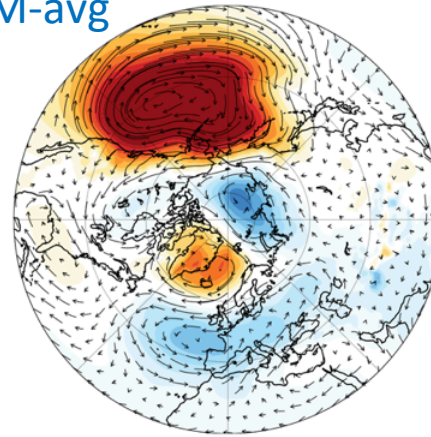
From CESM1

Consistent w/ Wernli and Papritz (2018)

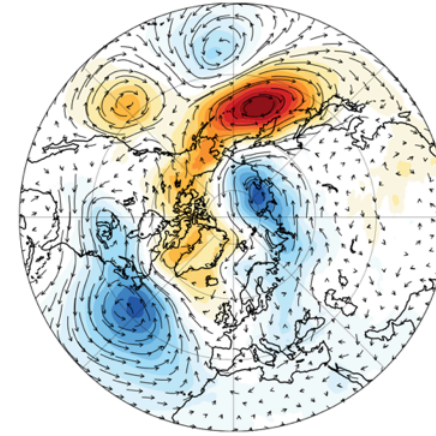
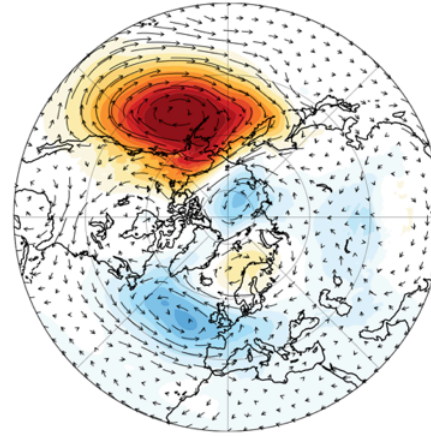
DJFM-avg

March/April

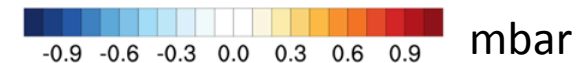
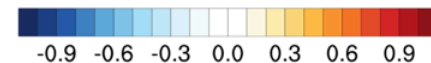
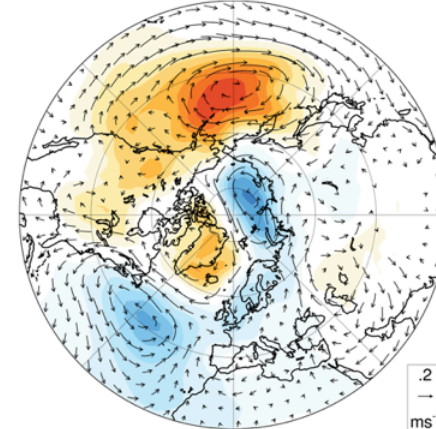
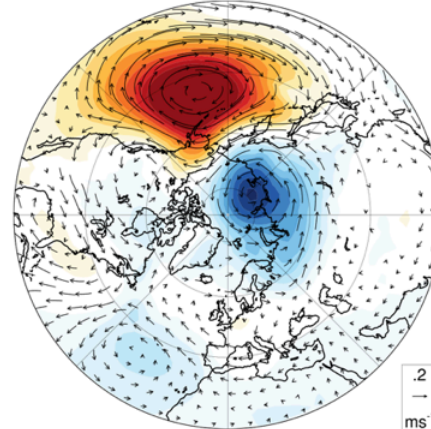
CESM1



FLOR

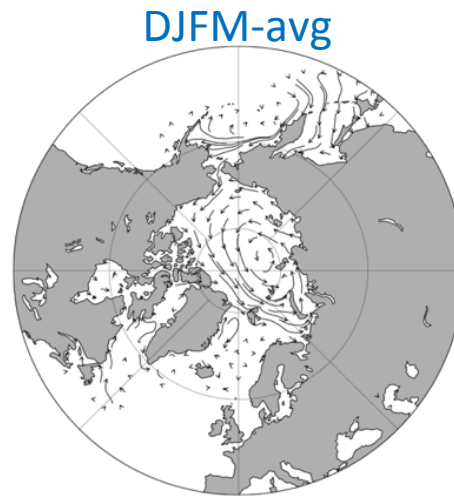


CM2.1

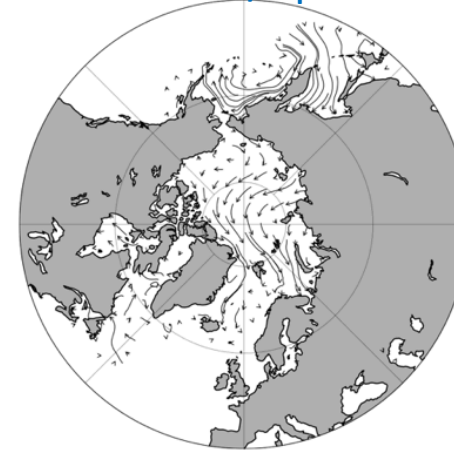


Differences in Sea-Ice Motion

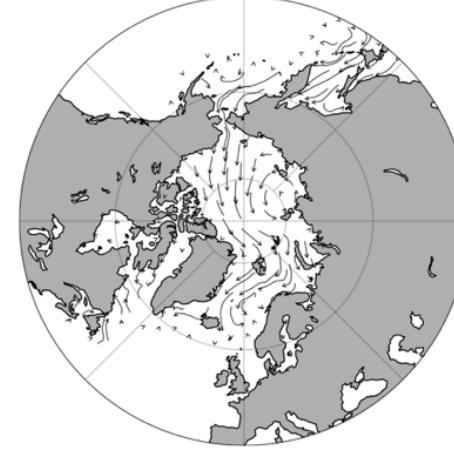
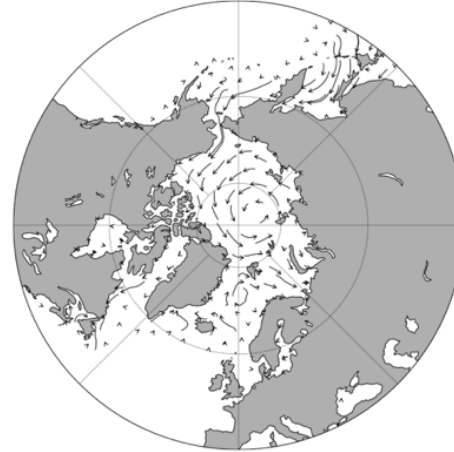
CESM1



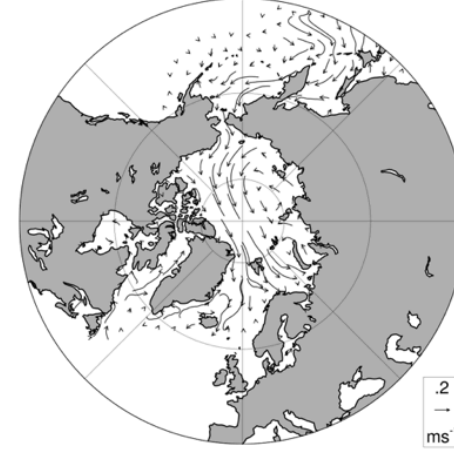
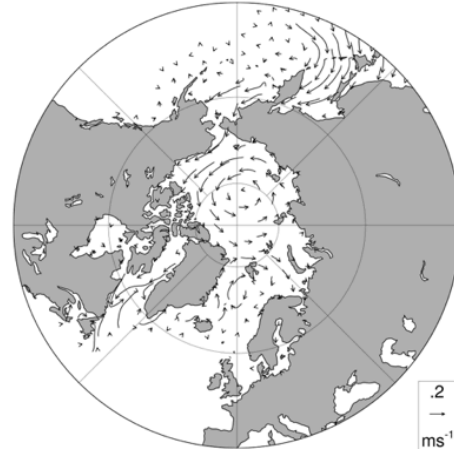
March/April



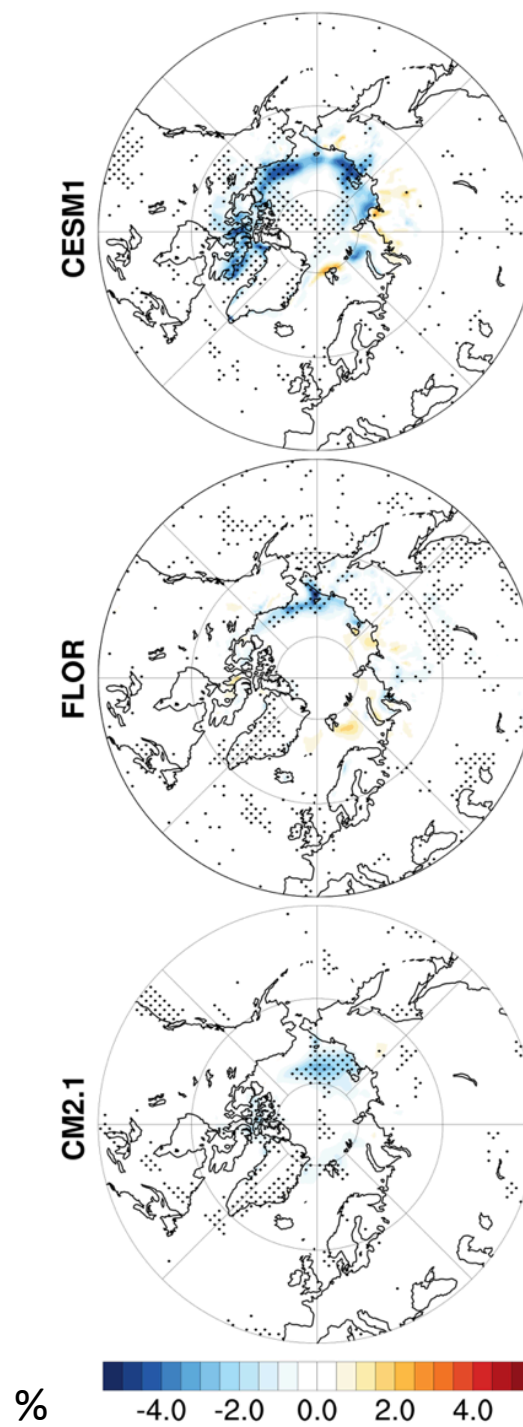
FLOR



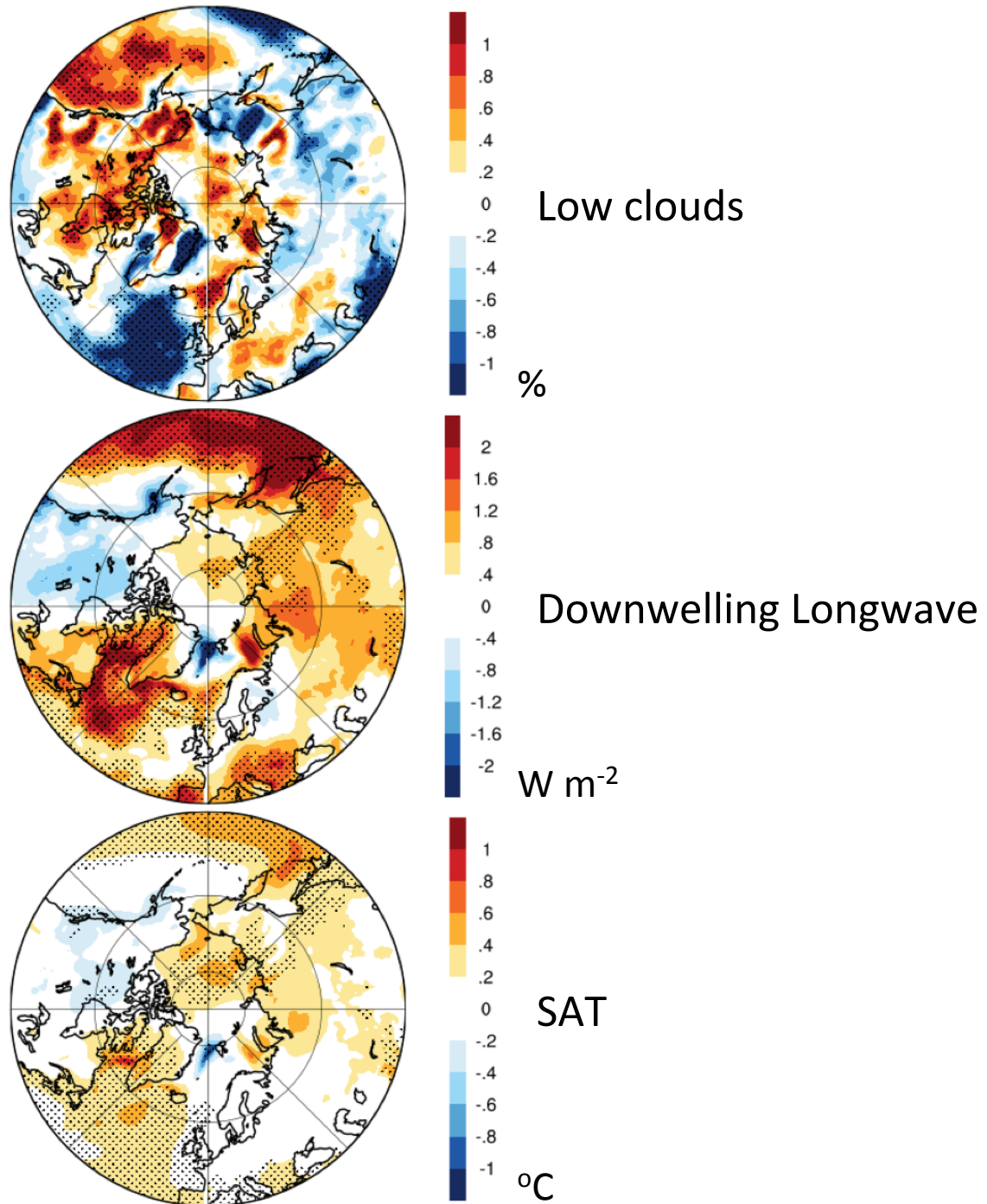
CM2.1



Differences in September Albedoes



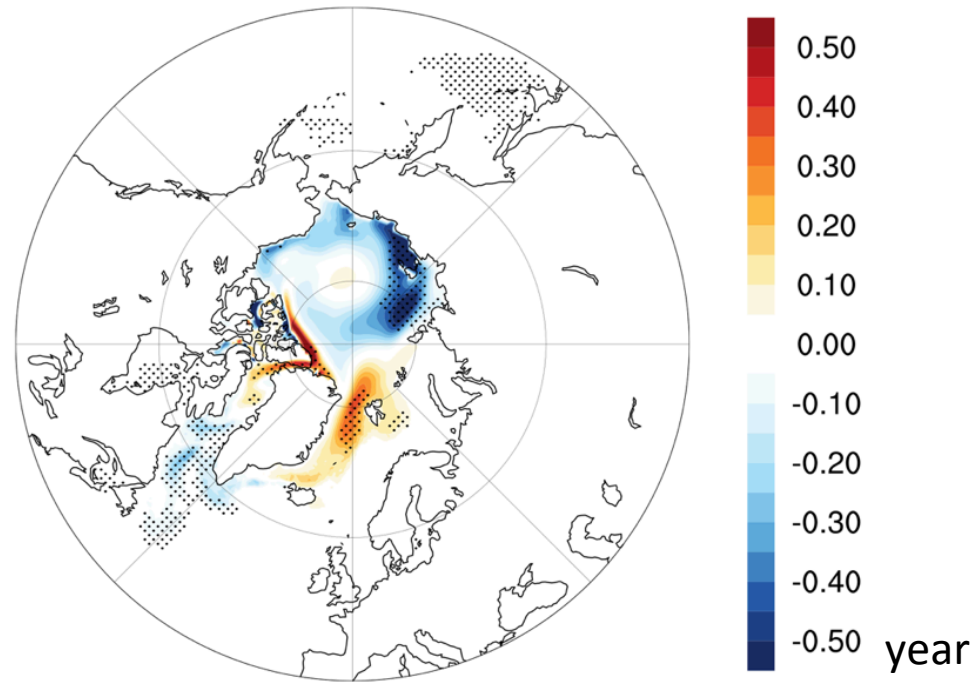
Differences in DJFM- average



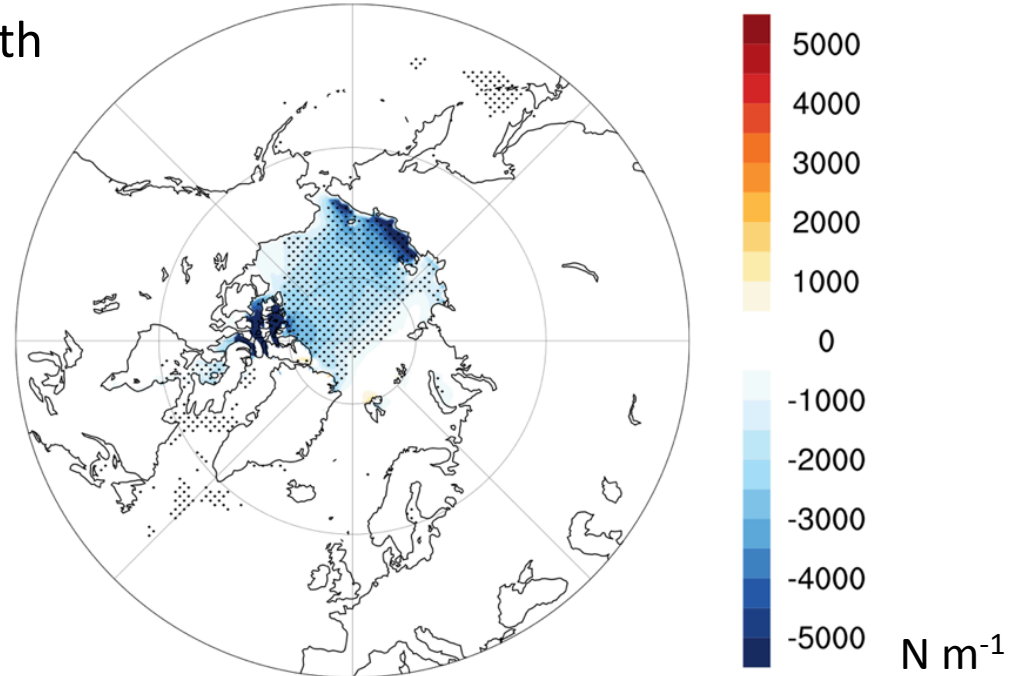
From CESM1

Differences in 10-year Average

Ice Age



Ice Strength

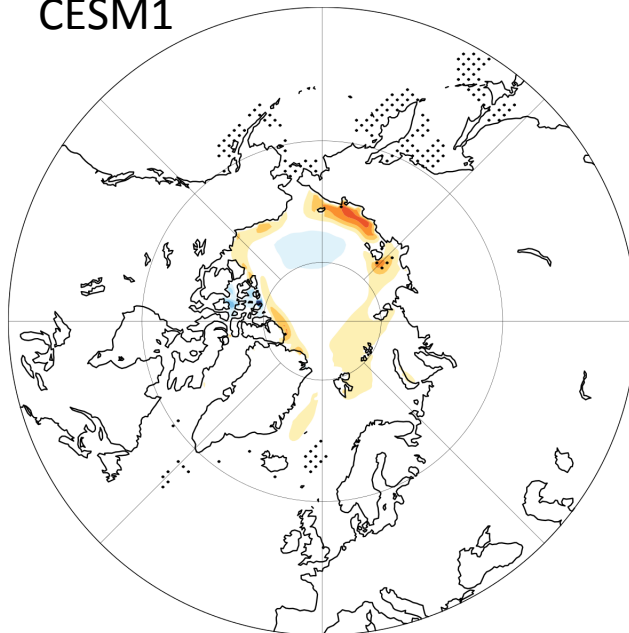


From CESM1

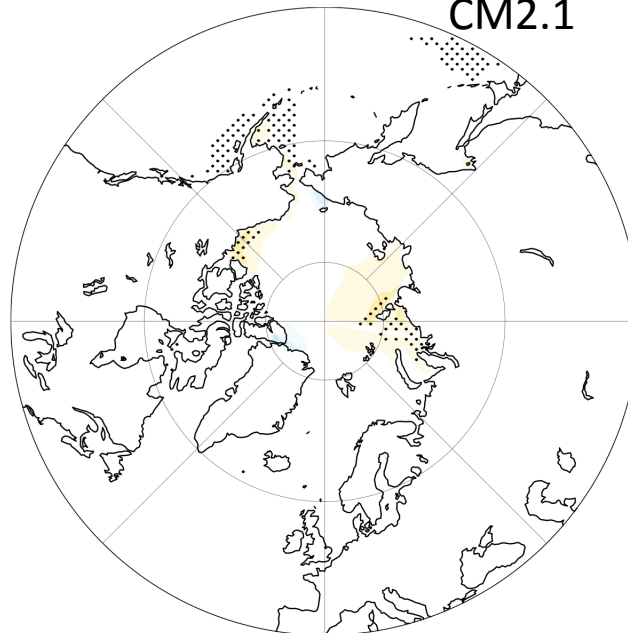
DJFM-Average Sea-Ice Thickness Differences

TA-only

CESM1

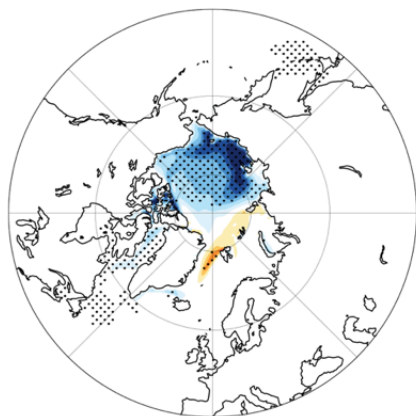


CM2.1

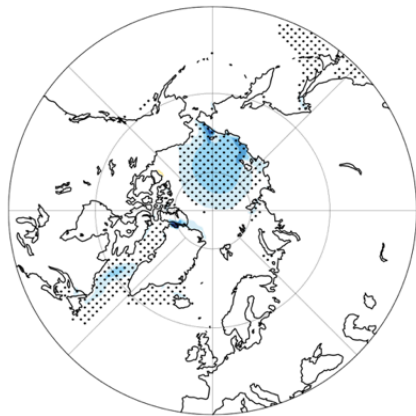


Full NA

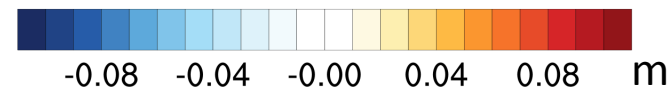
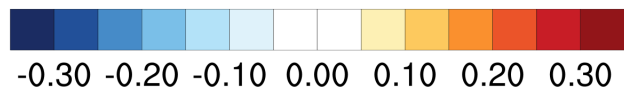
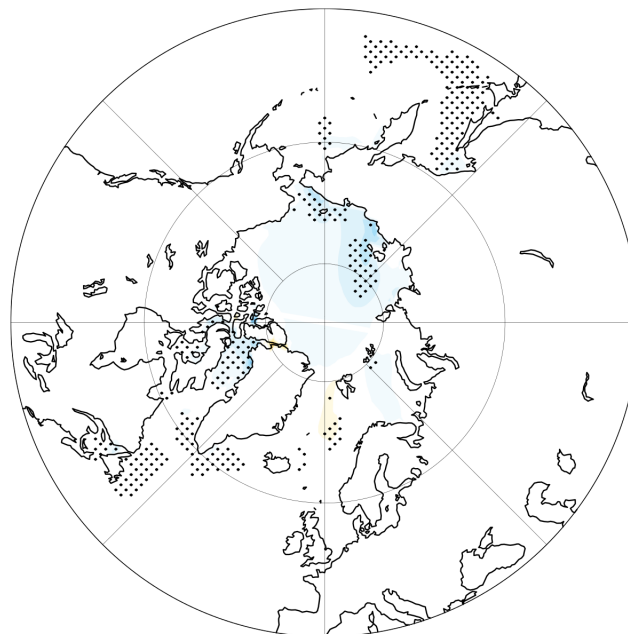
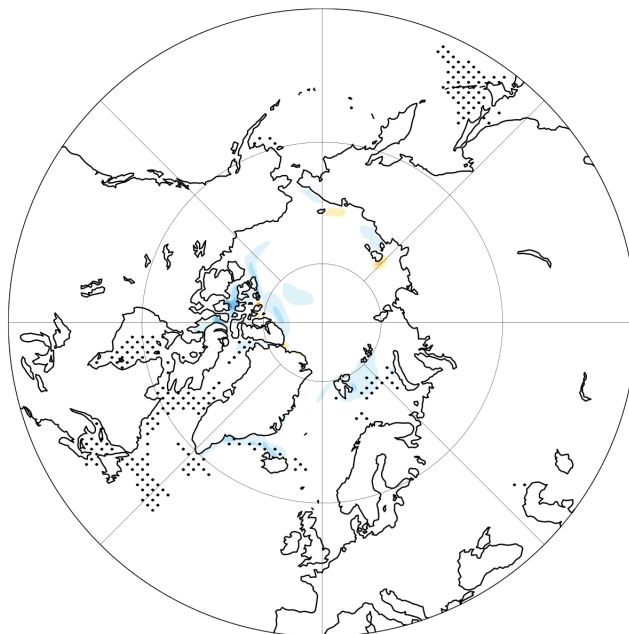
CESM1



CM2.1



SPNA-only



Summary and Conclusions

Idealized AMV simulations with three global coupled climate models robustly show thinner sea-ice pack in the Arctic in response to AMV+.

Such thinner ice results from atmospheric teleconnections / circulation changes (dynamical effect) that lead to:

- Weakening of the Beaufort Sea High, driving an anomalous cyclonic ice motion;
- A dipole pattern in SLP in late winter / early spring that acts to enhance the Transpolar Drift Stream, enhancing sea-ice transport out of the Arctic Basin.

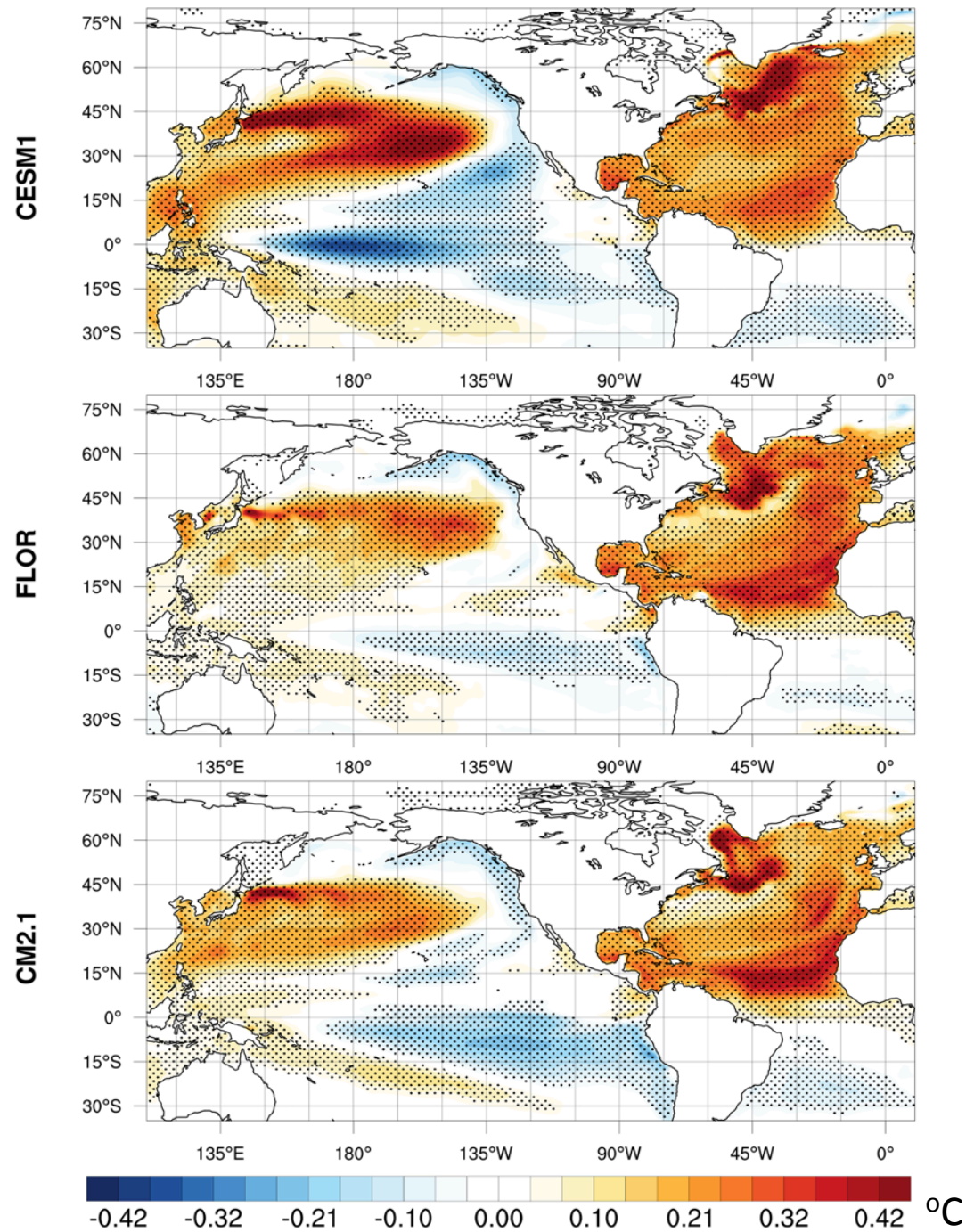
The AMV+ is also associated with (thermodynamic effect):

- Lower ice albedoes; increased low clouds; increased downwelling radiation;
- Higher surface temperatures

Thinner and younger ice is more prone to melt, resulting in less extensive sea-ice cover at the end of the melting season.

Weaker ice is more responsive to winds, i.e., positive feedback.

A Global Impact Example: Differences in DJFM Sea Surface Temperature



Differences in Sea-Ice Concentration

