TROPICAL PACIFIC OCEAN MODEL PARAMETERIZATIONS

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- GOAL ? : Exploit global predictability arising from Tropical Pacific variability (e.g. ENSO, MJO)
- OBJECTIVE ? : Improve sub-seasonal to annual predictions over the U.S.
- PREMISE ? : Forecasts are an *initial value problem* that depend on
- Observations TPOS Backbone and Global Observing Systems (e.g. satellites)
- Data Assimilation Tropical Pacific has potential for <u>coupled data</u> <u>assimilation</u>

3TPG§n2G2CoupleUM6de2018,<u>MBGUUGE</u>feads to systematic <mark>foreca</mark>st & AB



Sustained (Backbone) observations to:

- 1) the *initial condition* via *data assimilation* (DA)
- 2) the drift metric (DA increments)
- 3) the context for process and regional models

Breakouts 4-5 ?

Embedded "process" studies to :

- 1) improve (reduce drift) the prognostic models (e.g. parameterizations)
- 2) develop coupled data assimilation
- 3) inform "Backbone" decisions

Breakouts 2-5-6?



- 1) The Grand Challenge
- 2) "Parameterization" basics (e.g. non-linear advection)
- 3) Some common ocean model parameterizations & schemes
- 4) A framework
- 5) Transition to Operations (Not Today)





Outline

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1) The Grand Challenge



For a quantity
$$Q = \{ U, V, W, T, S, \text{ etc. } \}$$
:
 $\partial_t Q \sim -U \bullet \nabla Q = -\partial_x (UQ) - \partial_y (VQ) - \partial_z (WQ)$

Reynolds' decomposition into mean + fluctuation $\rightarrow Q = Q + q : \langle q \rangle = 0$ Model discretization into resolved + unresolved $\rightarrow Q = Q + q : \langle q \rangle \neq 0$



Correlations

e.g.
$$\langle wq \rangle = \int \phi_{wq}(f,k) df dk = \Sigma processes$$

Different processes occupy different regions of frequency–wavenumber space, so "Process Parameterizations" are additive.

"Parameterization schemes" (e.g. mixed layer, diurnal cycling) can encompass a variety of regions of frequency–wavenumber space.

Discussions & Breakouts: TPOS Priorities ? Double Counting ?



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MEET TPOS NEEDS ? HOW TO QUANTIFY ?

Processes :

- Gent McWilliams : **u**T, **u**S (mesoscale) ; for an unresolved "bolus" velocity **u***
- Vertical "Mixing" : wq (high f,k) ; entrainment velocity = $\partial_{t}h + w(-h)$
- Lateral Viscosity : [uQ + Uq + uq + vQ + Vq + vq] (Numerics)
- "Phenomenon" (schemes)
- "Pipes" over or through unresolved topography (slopes) (ITF)
- Boundary (Mixed) Layer
- Diurnal Cycling (SST, S, U) ;
- Lateral boundary conditions



A <u>parameterization</u> of the deviatoric component of the stress tensor σ_{ij} ; in terms of the resolved shear [$\nabla_k U_l + \nabla_l U_k$]: Symmetries $T_{ijkl} = T_{jikl} = T_{ijlk} = T_{klij}$ leave only 21 independent elements

Models typically specify a vertical K_v plus one (K_u) or two (e.g. A_H downstream, B_H , cross-stream) grid-dependent lateral viscosities

*K*_µ, *B*_µ control of equatorial zonal currents ?

Sufficient for equatorial dynamics ? ($\sigma_{23} \neq 0$, where [$\nabla_3 U_2 + \nabla_3 U_2$] = 0

What do observations say ?

Are there numerical constraints ?



OCEAN MODEL PARAMETERIZATIONS 3) Diurnal Cycling (SST & SSS)





INGREDIENTS?

- 1) What needs to be improved (e.g. sources of drift over a forecast interval)?
- 2) What are the observational/theoretical/process model needs (CLIVAR CPTs)?
- 3) A protocol, including data assimilation (DA)
- 4) A developmental "testbed"
- 5) **<u>Objective</u>** Evaluation (process evaluation, analysis of increments)



4) Framework



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4) Regional Models

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Modulation of Equatorial Turbulence by Tropical Instability Waves in a Regional Ocean Model



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4) Large Eddy Simulation





QUESTIONS ?



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A <u>parameterization</u> of the deviatoric component of the stress tensor σ_{ij} ; in terms of the resolved shear [$\nabla_k U_l + \nabla_l U_k$]: Symmetries $T_{ijkl} = T_{jikl} = T_{ijlk} = T_{klij}$ leave only 21 independent elements

Models typically specify a vertical K_v plus one (K_u) or two (e.g. A_H downstream, B_H , cross-stream) grid-dependent lateral <u>viscosities</u>

$$T_{1111} = T_{2222} = 2 K_{u} = A_{H} + B_{H}$$

$$T_{3333} = K_{v} + K_{u} = K_{v} + A_{H}$$

$$T_{1212} = K_{u} = B_{H}$$

$$T_{1313} = T_{2323} = K_{v}$$

$$T_{1133} = T_{2233} = K_{u} - K_{v} = B_{H} - K_{v}$$

$$K_{u}, B_{H} \text{ control equatorial zonal currents}$$

$$Sufficient for equatorial dynamics?$$

$$\sigma_{23} \neq 0, \text{ where } [\nabla_{3}U_{2} + \nabla_{3}U_{2}] = 0$$
What do observations say?
Are there numerical constraints?