

CLIMATE RESEARCH FACILITY

The use of CPOL data for validating GCM simulations of tropical convection



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1. Motivation Evaluate the DOE Energy Exascale Earth System Model (E3SM) convective parameterizations over	2. Darwin rainfall Nov. to May, Northern Australian Monsoon [1] +	3. Instrumentation CPOL: C-band POLarization radar, PPI scans @ 18 elevations every 10 min. from 1998-2017	4. Rainfall uncertainty analysis		
			2D histogram of radar estimated vs. disdrometer observed rain rates		Statistical uncertainty [6]
ne Maritime Continent with 10-25 km resolution	Madden-Julian Oscillation (MJO) [2] important + not well resolved in GCMs	Python ARM Radar Toolkit (Py-ART) [3] used to process/grid data	Using specific attenuation [4]	"CSU-blended" technique [5]	Darwin 10^2 — Fit for $R(A_h)$: U = 1.16 $R^{0.32}$ — Fit for $R(K_{dp})$: U = 0.93 $R^{0.52}$ — Fit for CSU-blended: U = 0.85 $R^{0.24}$
Convection in E3SM parameterized \rightarrow convection triggers based on available CAPE. Limits interactions	 19 years of continuous data from CPOL in Darwin: 	4/day rawinsondes: Monsoon/Break [1] Multidecadal dataset provides observational targets for GCM	Intate from CPOL [mm hr ⁻¹]	LO - 0.4 - 0.3 sales - 0.2 %	10 ⁴ 10 ⁰ 10 ⁰ 10 ⁻¹

7. Rainfall rates

with seabreezes and cold pools generated by surrounding convection.

The link between large scale forcing (i.e. MJO phase) + rainfall rates, cloud top heights convection in Darwin not well understood.





Using the tropical ocean "CSU-blended" technique [5] minimizes the difference in derived rainfall rates from disdrometer data + statistical uncertainty of retrieval.

5. Echo top heights



evaluation



Global run of E3SM using:

- November 2009 to April 2011
- IPCC AR5 scenario
- Temperature, humidity, winds nudged to ERA-Interim forcing.
- 1 degree spatial resolution

Compute: 3 hour rain accumulation in 4 grid cells below, same metric compared in both CPOL and E3SM data

8. Diurnal cycle evaluation of E3SM

Area coverage of grid boxes used in comparison:





[4] Ryzhkov, A., Diederich, M., Zhang, P., and Simmer, C., 2014: Potential Utilization of Specific Attenuation for Rainfall Estimation, Mitigation of Partial Beam Blockage, and Radar Networking, Journal of Atmospheric and Oceanic Technology, 31, 599-619, https://doi.org/10.1175/JTECH-D-13-00038.1, https://doi.org/10.1175/JTECH-D-13-00038.1, 2014

[5] Thompson, E. J., Rutledge, S. A., Dolan, B., Thurai, M., and Chandrasekar, V., 2018: Dual-Polarization Radar Rainfall Estimation over Tropical Oceans, Journal of Applied Meteorology and Climatology, 57,755–775, https://doi.org/10.1175/JAMC-D-17-0160.1, https://doi.org/10.1175/JAMC-D-17-0160.1

[6] Kirstetter, P.-E., Gourley, J. J., Hong, Y., Zhang, J., Moazamigoodarzi, S., Langston, C., and Arthur, A. (2015), Probabilistic precipitation rate estimates with ground-based radar networks, Water Resour. Res., 51, 1422–1442, doi:10.1002/2014WR015672

[7] Sakaeda, N., S.W. Powell, J. Dias, and G.N. Kiladis, 2018: The Diurnal Variability of Precipitating Cloud Populations during DYNAMO. J. Atmos. Sci., 75, 1307–1326, https://doi.org/10.1175/JAS-D-17-0312.1

- Strong afternoon diurnal cycle over land in break \rightarrow diurnal heating.
- Early morning peak over ocean stronger w/ active MJO, also observed in Indonesia [7], potentially due to convergence from differential radiative cooling.



- E3SM run fails to resolve phase and magnitude of diurnal cycle of rain. Likely due to inability of E3SM to resolve seabreezes.
- Future runs with improved resolution/machine learning based convective trigger to see if improvements result

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