

# Observations and models of oceanic macroturbulence:

meet the new bias same as the old bias

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with

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Valentin Resseguier (L@b SCALIAN)

Sources and Sinks of Eddy Mesoscale Energy,  
3/14/19, 10:40-11:00AM

Sponsors: NSF (OCE 1350795), ONR (N00014-17-1-2963), CARTHE

## Old Bias—Reynolds Avg.—LowRes

Questionable eddy parameterizations  
Diversity of approaches, no cross-evaluation  
Not much data to assess them  
Unclear relationship of key impacts to params.

## New Bias—Large Eddy Sims. (LES)

Questionable subgrid parameterizations  
Diversity of approaches, no cross-evaluation  
Not much data to assess them  
Unclear relationship of key impacts to schemes.

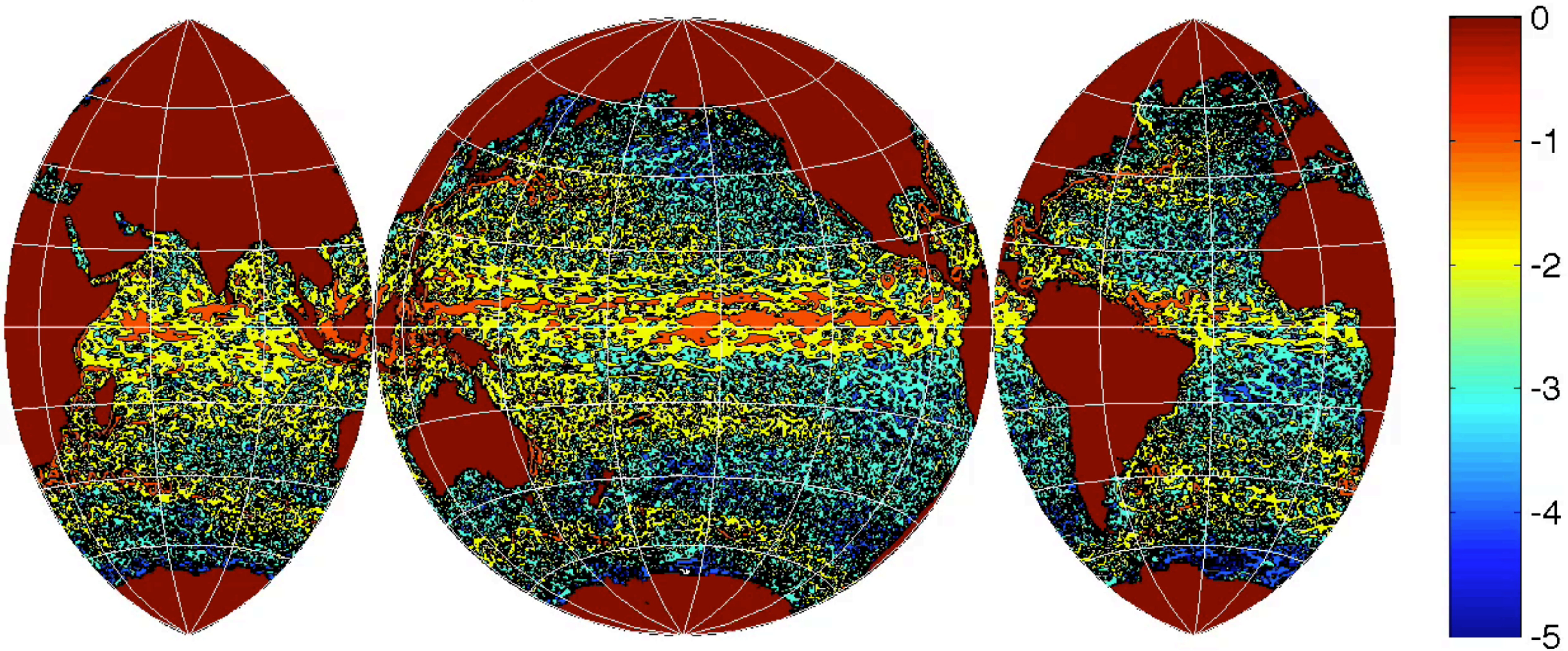




# Largest (Mesoscale) Ocean Eddies are Small vs. Earth



AVISO:  $\log_{10}(0.5 (u^2 + v^2))$  on 19940101





# Global Models have not resolved many, until now...

MITgcm LLC4320 (2km global)

Courtesy:

Menemenlis & Fenty

1.000  
0.750  
0.500  
0.250  
0.000

Surface speed (m/s)

0h 13 Sep 2011





# Absent realistic global models, We studied "Cascade" Scalings



3D: Richardson/Kolmogorov/Smagorinsky/Corrsin

$$E \propto \epsilon^{2/3} \ell^{5/3}, \quad S_2 \propto \epsilon^{2/3} r^{2/3}, \quad \epsilon \propto \nu \alpha^2, \quad \nu = \text{Pr} \kappa \propto \Delta x^2 |\alpha| \propto \epsilon^{1/3} \ell^{4/3}$$

$\epsilon$  = conserved energy flux

2D: Barnier/Kraichnan/Leith

$$E \propto \eta^{2/3} \ell^3, \quad S_2 \propto \eta^{2/3} r^2, \quad \eta \propto \nu (\nabla \omega)^2, \quad \nu \propto \Delta x^3 |\nabla \omega| \propto \eta^{1/3} \ell^2, \quad \kappa \propto ?$$

$\eta$  = conserved enstrophy flux

Quasigeostrophy: Barnier/Charney/QGLEith

$$E \propto \eta^{2/3} \ell^3, \quad S_2 \propto \eta^{2/3} r^2, \quad \eta \propto \nu (\nabla q)^2, \quad \nu = \kappa_{Redi} = k_{GM} \propto \Delta x^3 |\nabla q| \propto \eta^{1/3} \ell^2$$

$\eta$  = conserved potential enstrophy flux

SQG?

Submesoscale: McWilliams/?/?F-K?

$$E \propto \ell^2, \quad S_2 \propto r^1, \quad d/dt(PE + KE) = ??, \quad \nu = ?, \quad \kappa = ?$$



# QG & 2D Leith

$$v^* = \left( \frac{\Lambda \Delta x}{\pi} \right)^3 |\nabla q^*|$$

Different (Pot'l) Vorticity Gradients:

$$q_{2d}^* = f + \hat{k} \cdot \nabla \times u^*$$

$$q_{qg}^* = f + \hat{k} \cdot \nabla \times u^* + \frac{\partial}{\partial z} \frac{f^2}{N^2} b^*$$

$$\Lambda = 1$$

stretching—needs “taming” where QG is a bad approx (equator, boundary layers, etc.)

Use grid-scale nondims to determine when on the fly

$$Ro^* = \frac{U^*}{f \Delta x} \quad Bu^* = \frac{N^{*2} \Delta z^2}{f^2 \Delta x^2} \sim Ro^{*2} Ri^*$$

B. Pearson, BFK, S. D. Bachman, and F. O. Bryan, 2017: Evaluation of scale-aware subgrid mesoscale eddy models in a global eddy-rich model. *Ocean Modelling*, 115:42–58.

S. D. Bachman, BFK, and B. Pearson. A scale-aware subgrid model for quasigeostrophic turbulence. *Journal of Geophysical Research-Oceans*, 122:1529-1554, March 2017.



\* = as resolved

# QG & 2D Leith

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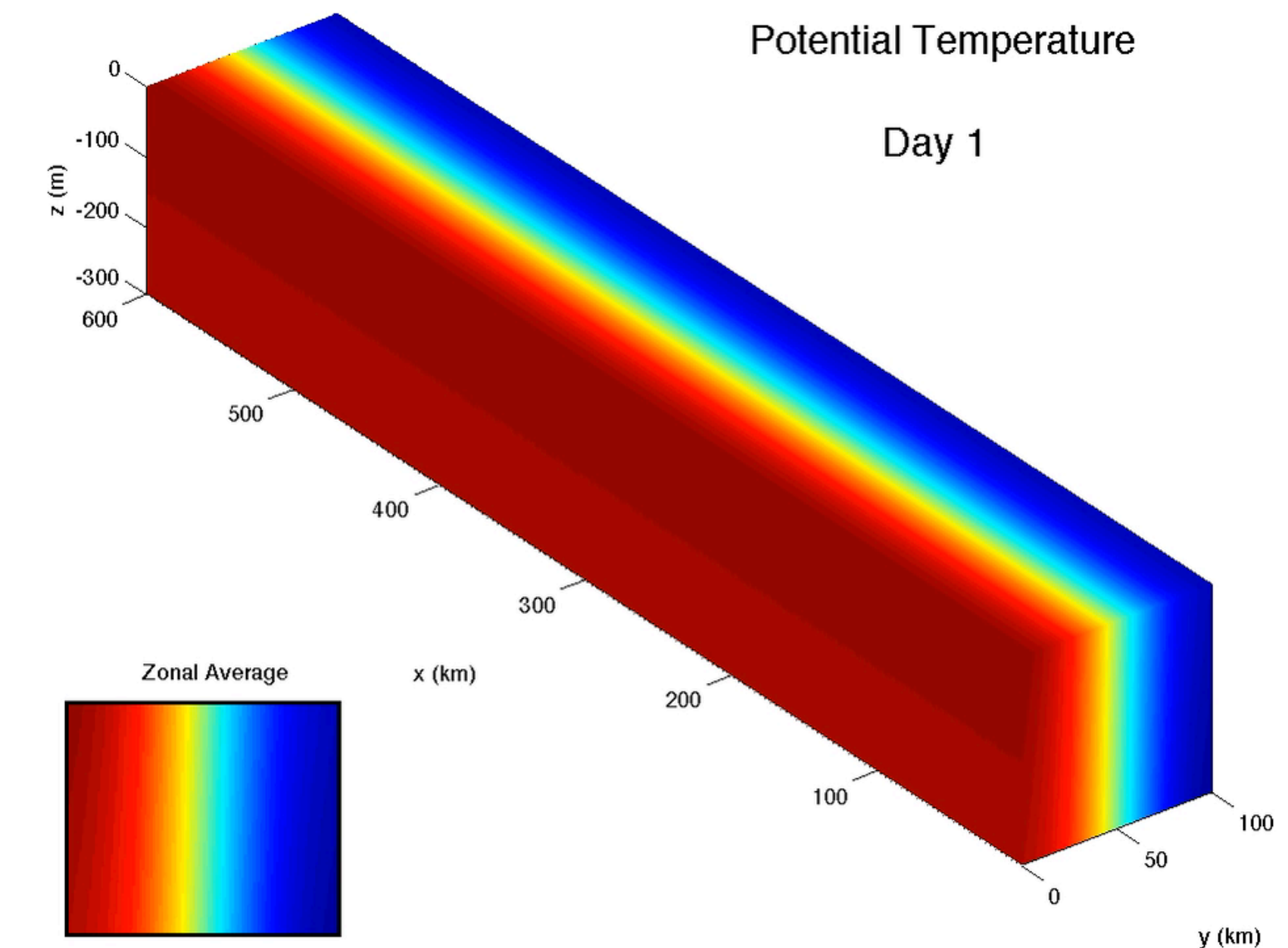
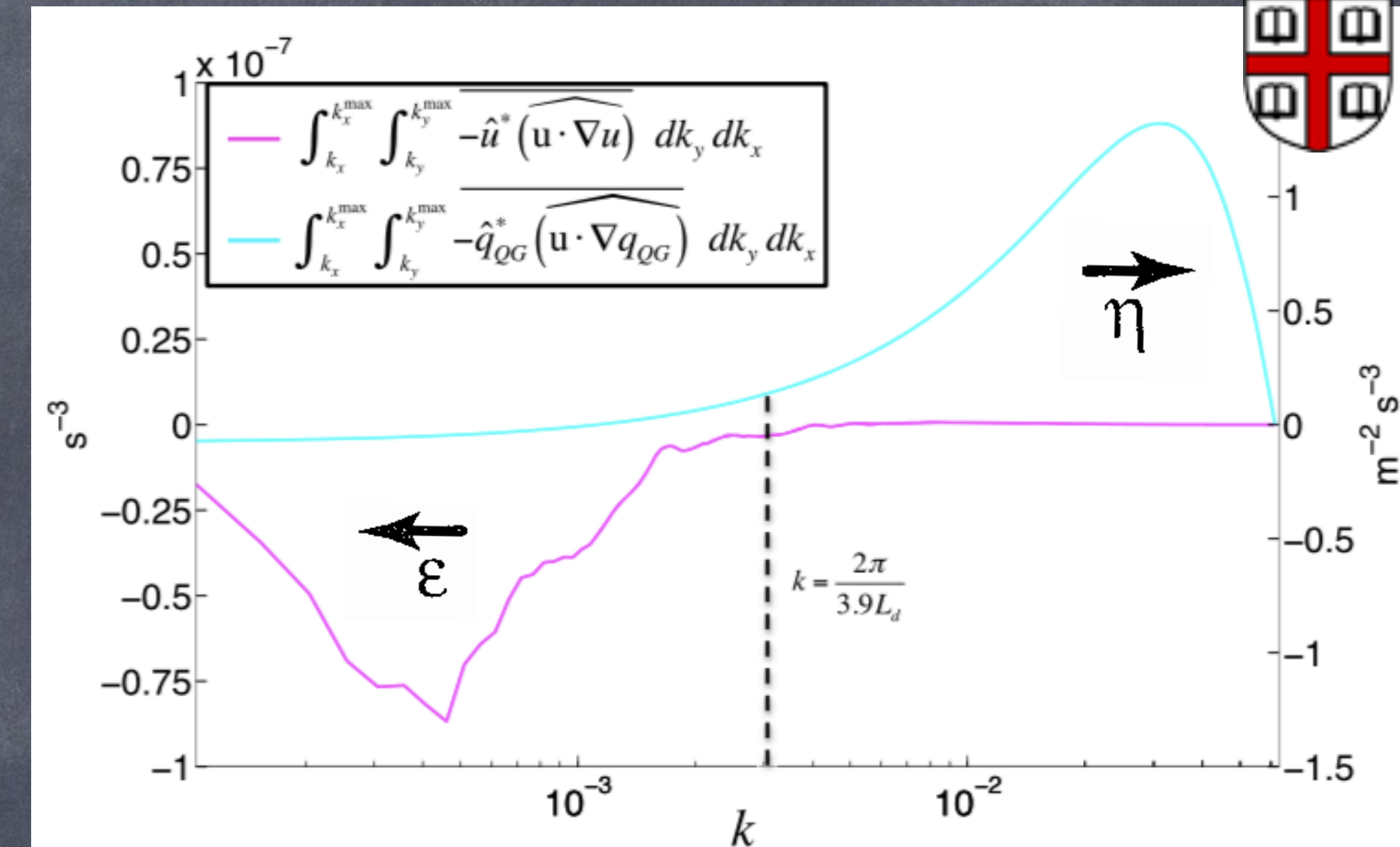
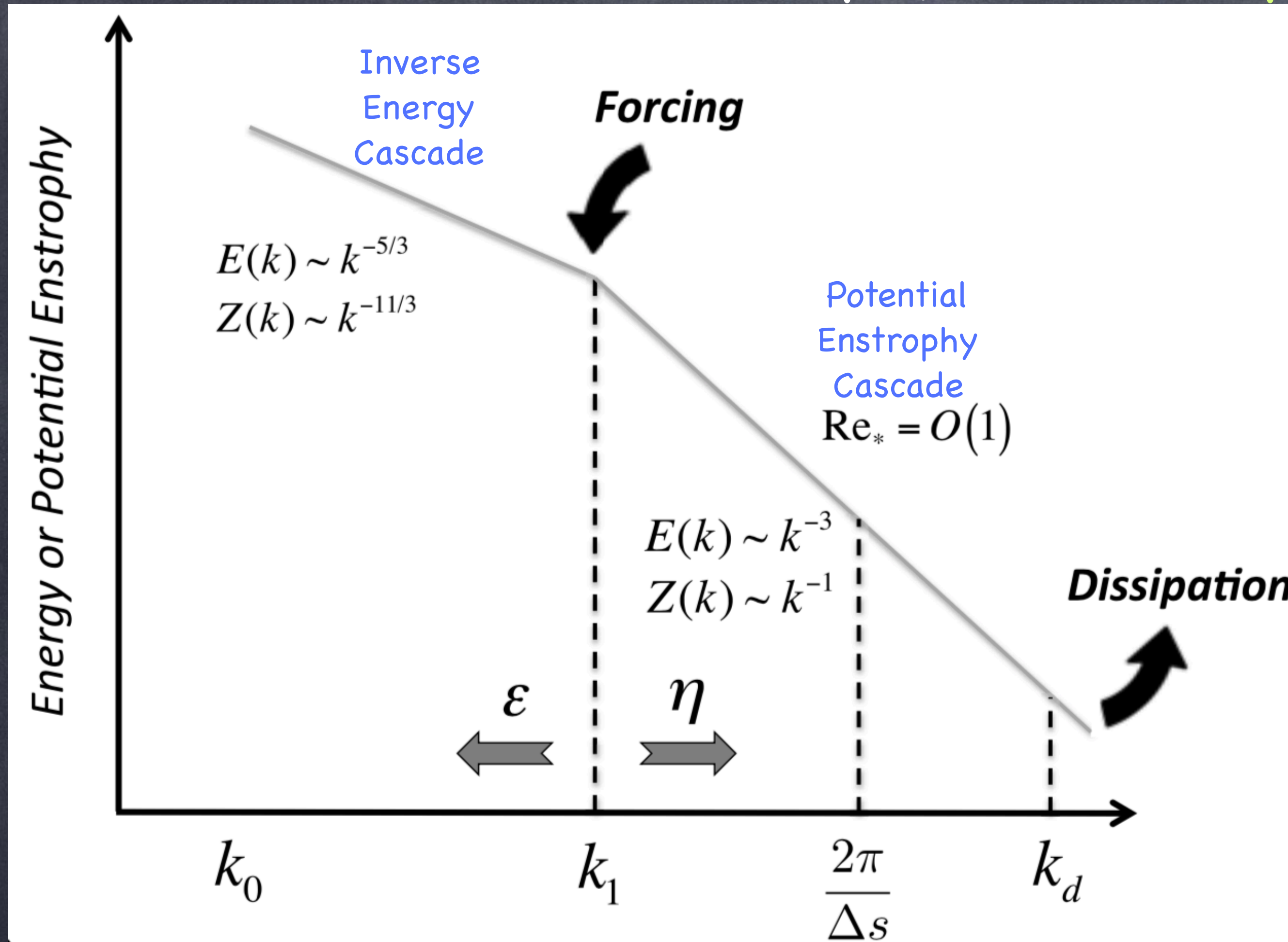
$$\Lambda = 1$$

Also, different implications, because **relative vorticity**, **buoyancy**, **T**, **S** dissipation now must be consistent with PV:

$$\frac{Dq_{qg}^*}{Dt} = -\nabla \cdot \overline{u' q'_{qg}} \approx \nabla \cdot [\nu^* \nabla q_{2d} + \kappa_{gm}^* \nabla (q_{qg} - q_{2d})] \rightarrow \kappa_{gm}^* = \nu^* = \kappa_i^*$$



# QG Turbulence: Pot'l Enstrophy cascade (potential vorticity<sup>2</sup>)



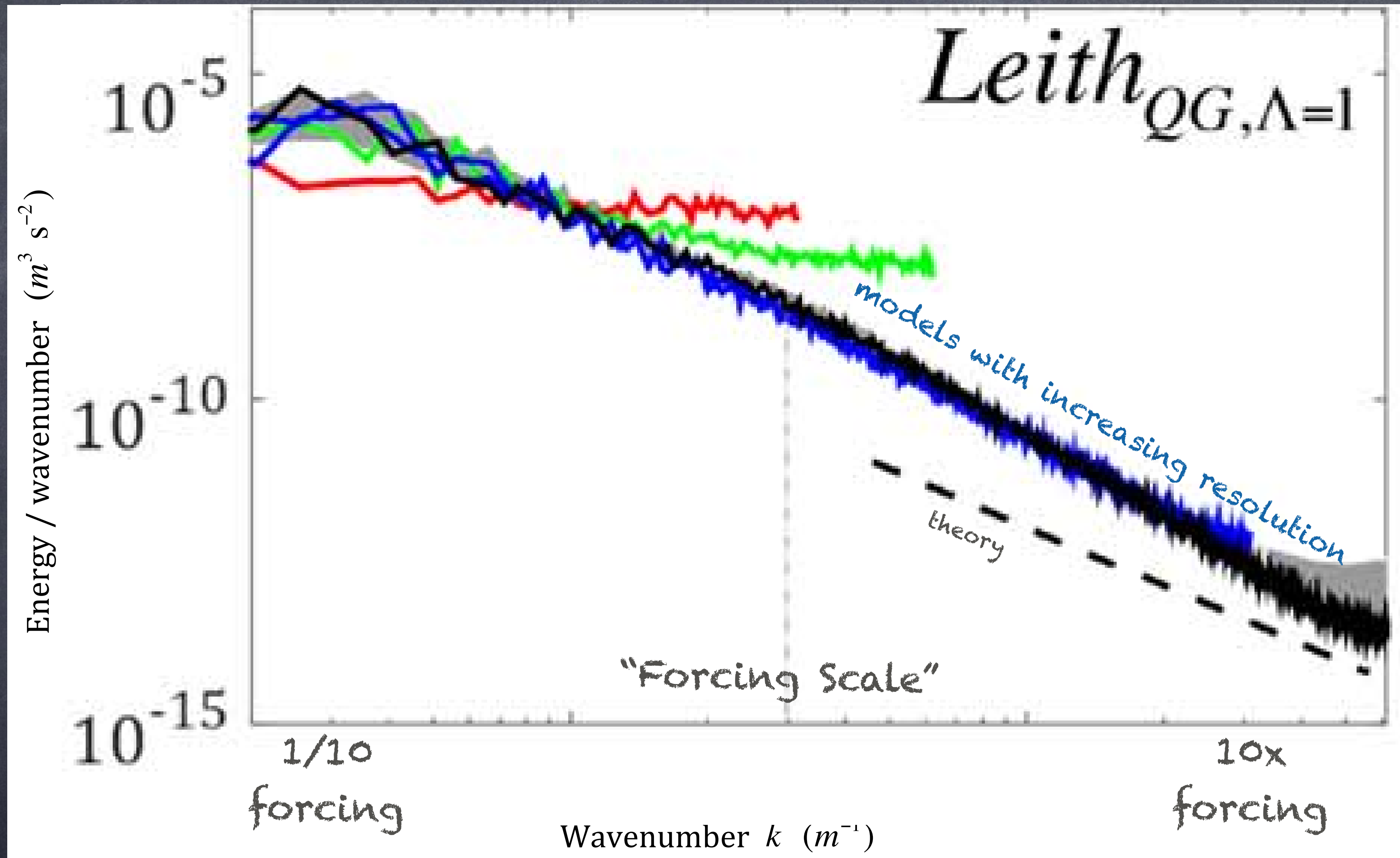
S. D. Bachman, B. Fox-Kemper, and B. Pearson. A scale-aware subgrid model for quasigeostrophic turbulence. *Journal of Geophysical Research-Oceans*, 122:1529-1554, March 2017.

BFK, W. Pan and V. Resseguier. Data-driven versus self-similar parameterizations for Stochastic Lie Transport and Location Uncertainty. In preparation.



# Where does ocean energy go?

Spectrally speaking



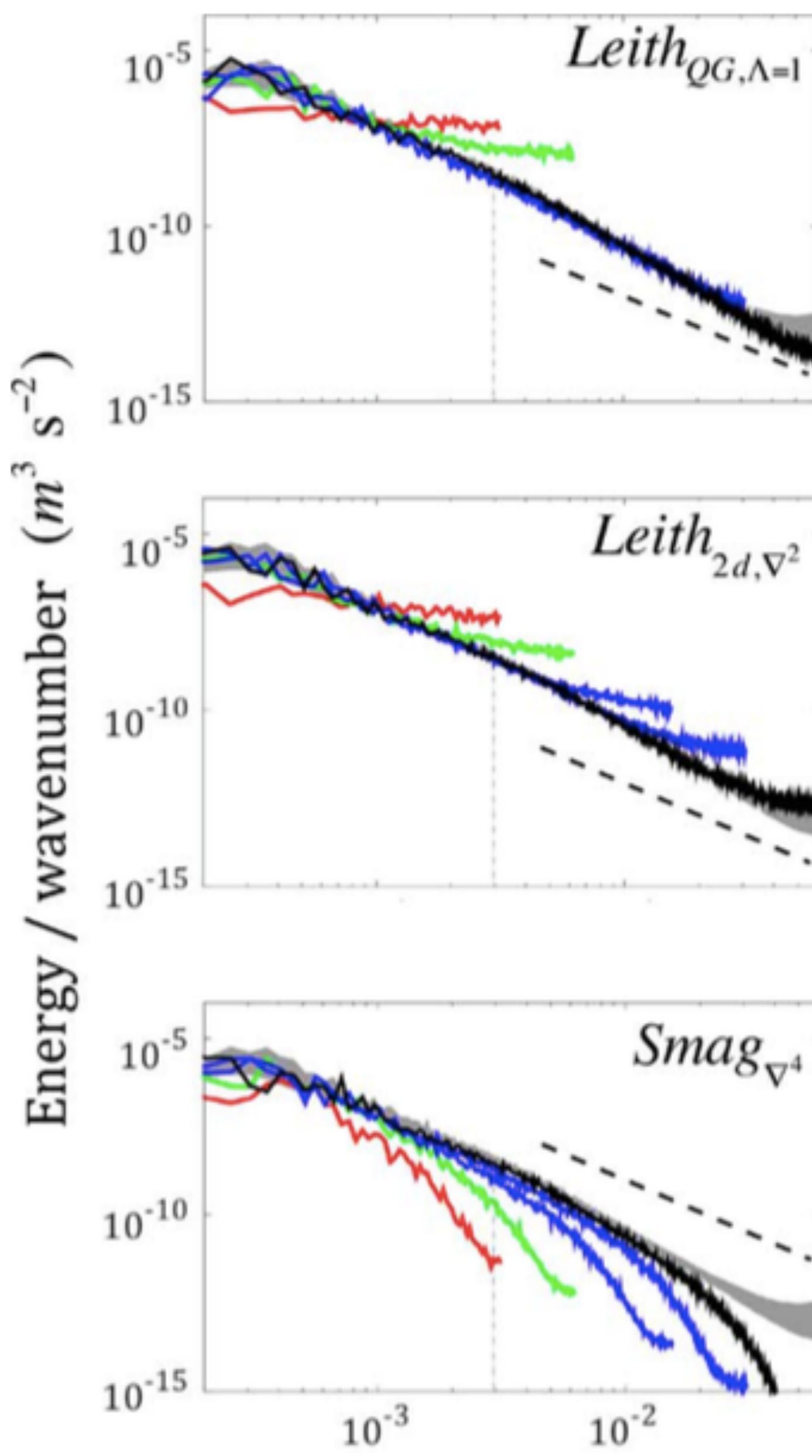
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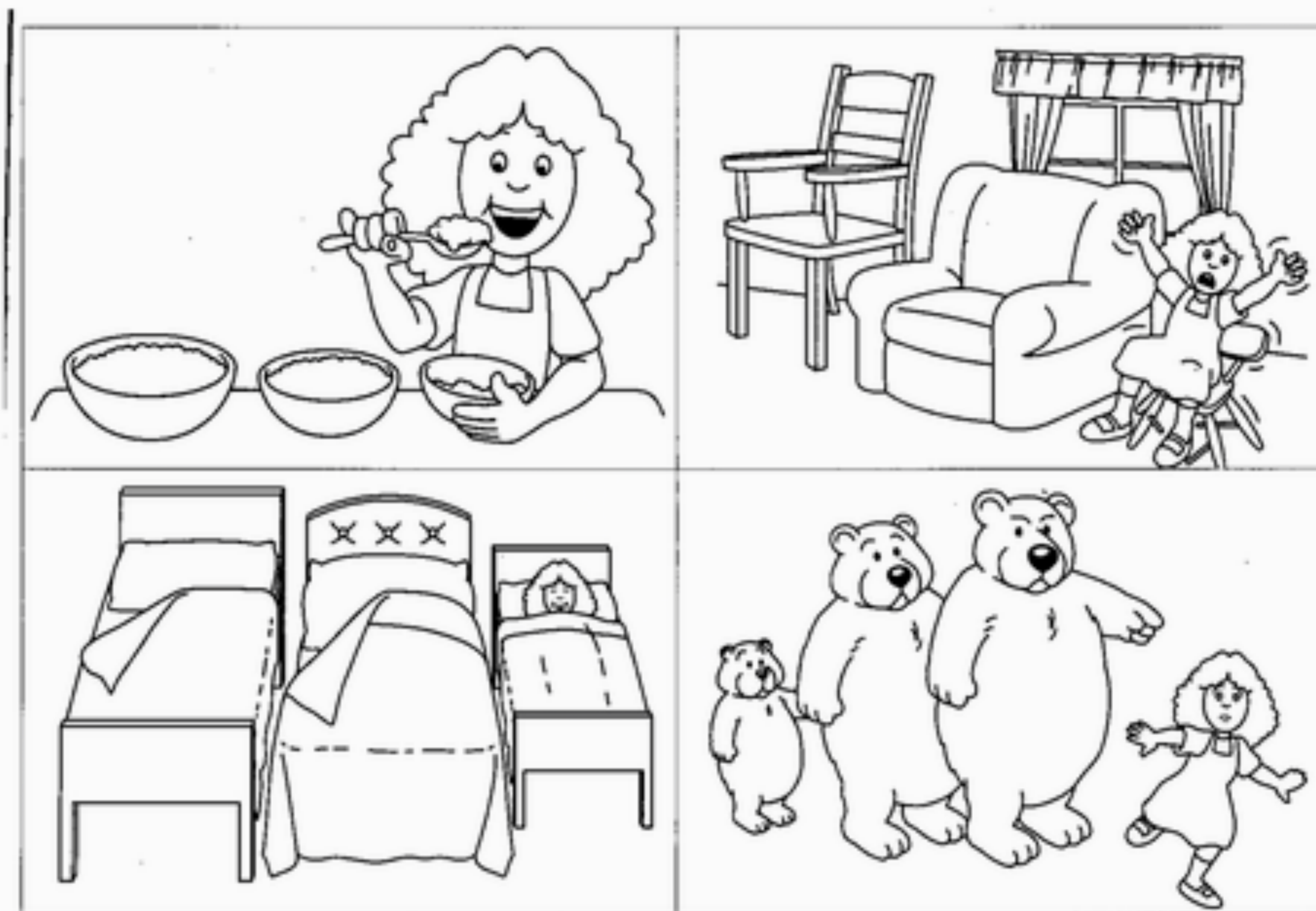
Spectrally speaking



QG Leith:  
Just Right!

2D Leith:  
Too Noisy

3D Smagorinsky:  
Too Smooth



S. D. Bachman, B. Fox-Kemper, and B. Pearson, 2017: A scale-aware subgrid model for quasi-geostrophic turbulence. *Journal of Geophysical Research—Oceans*, 122:1529–1554. URL <http://dx.doi.org/10.1002/2016JC012265>.





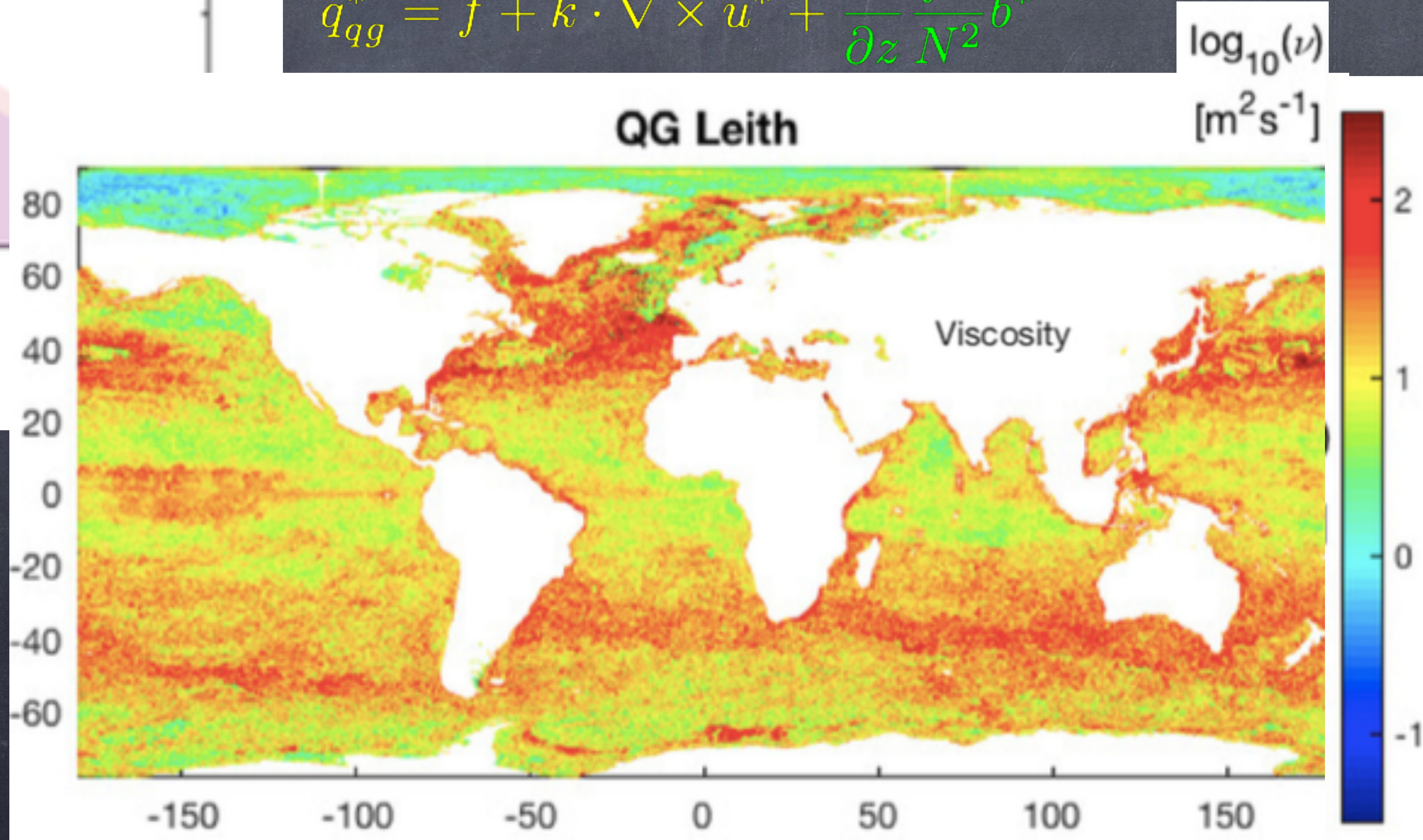
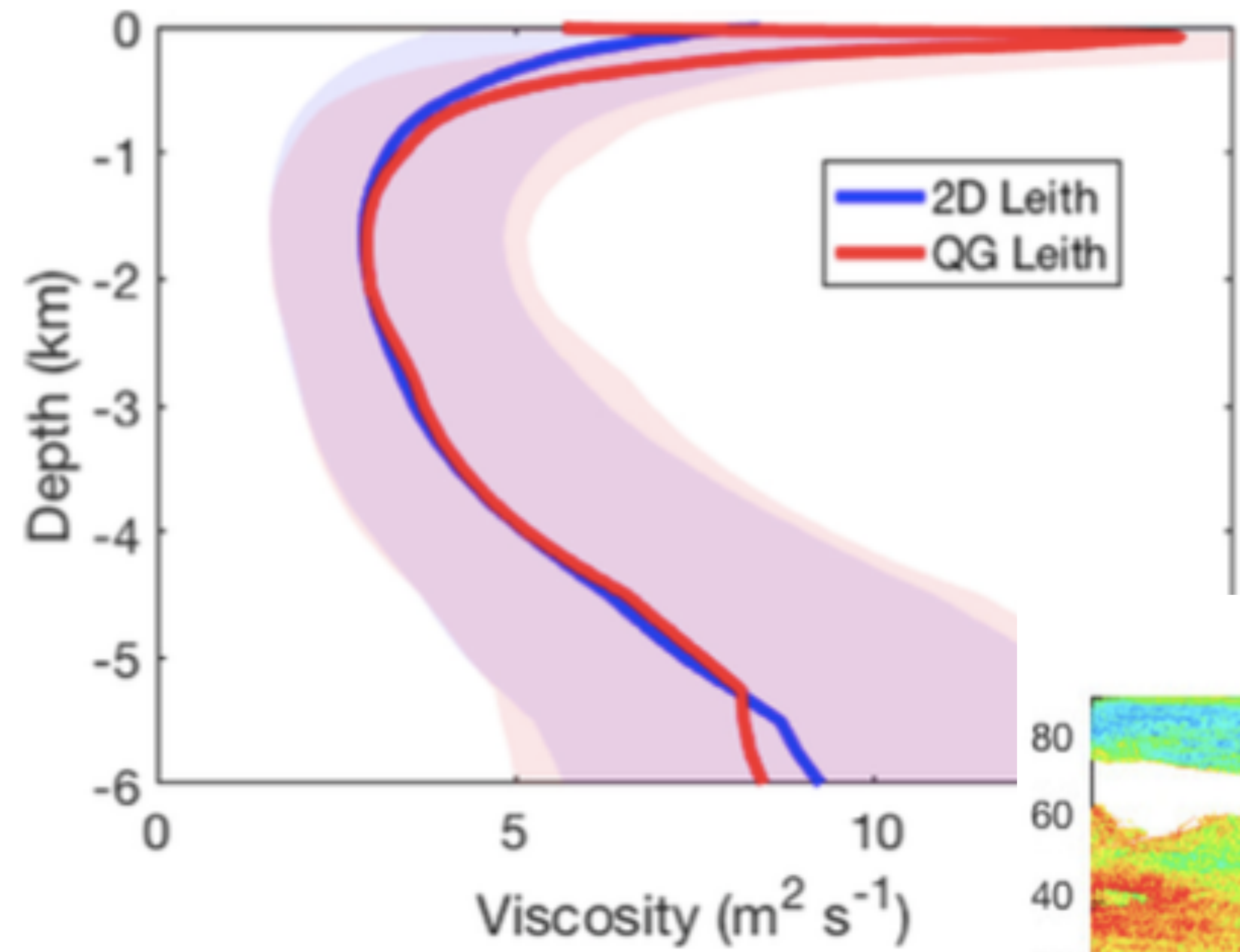
# QG Leith:

Works OK in an idealized flow:  
Let's try it in a realistic, 10km  
CORE-forced global model!

$$\nu^* = \kappa_i^* = k_{GM}^* \propto \Delta x^3 |\nabla q|$$

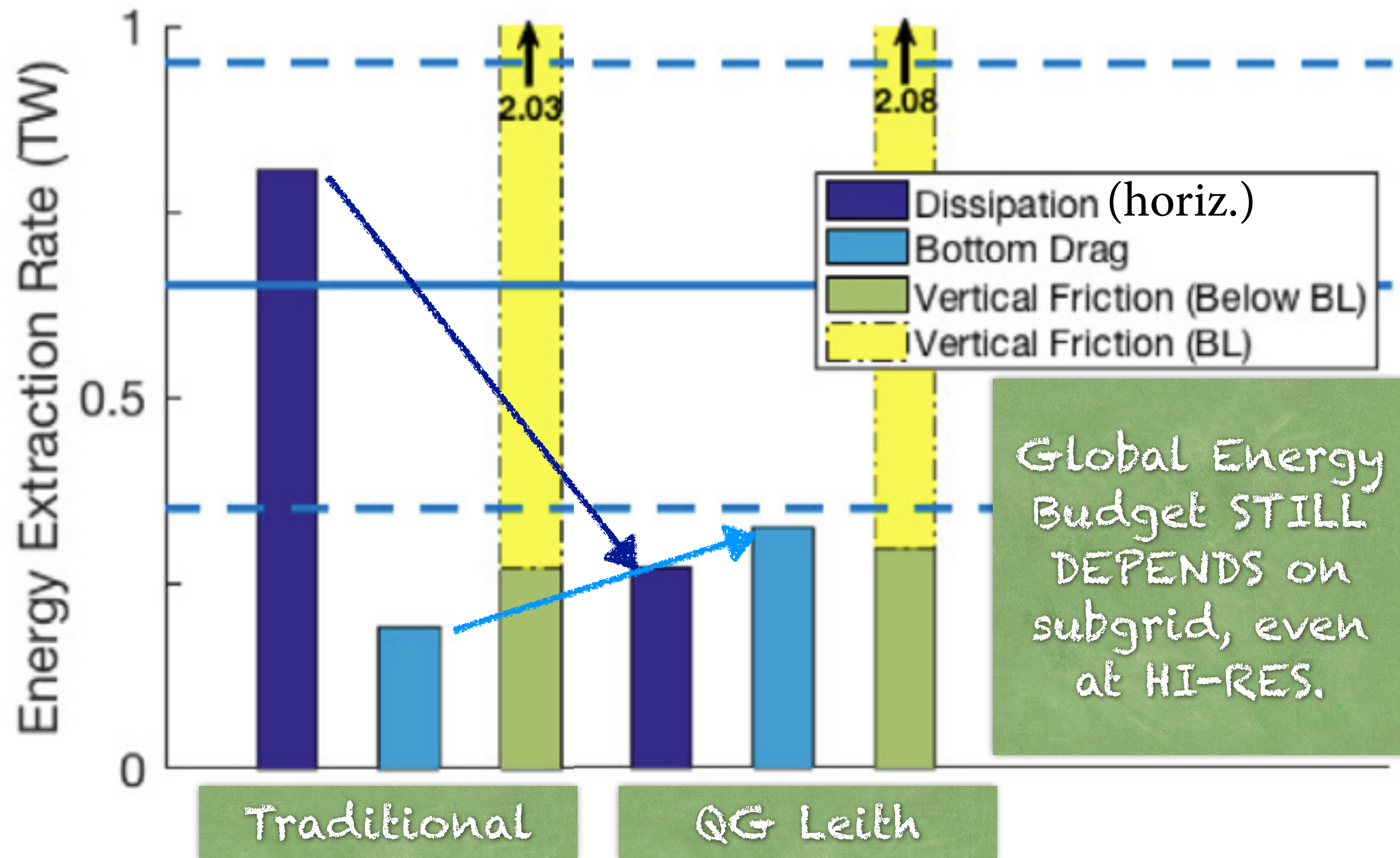
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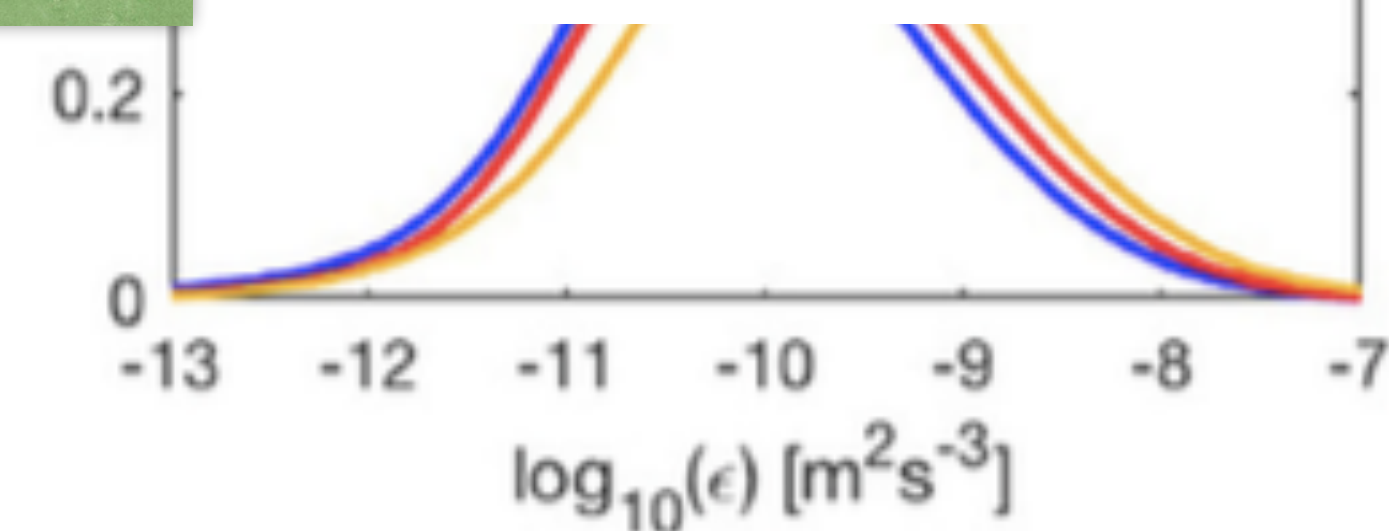
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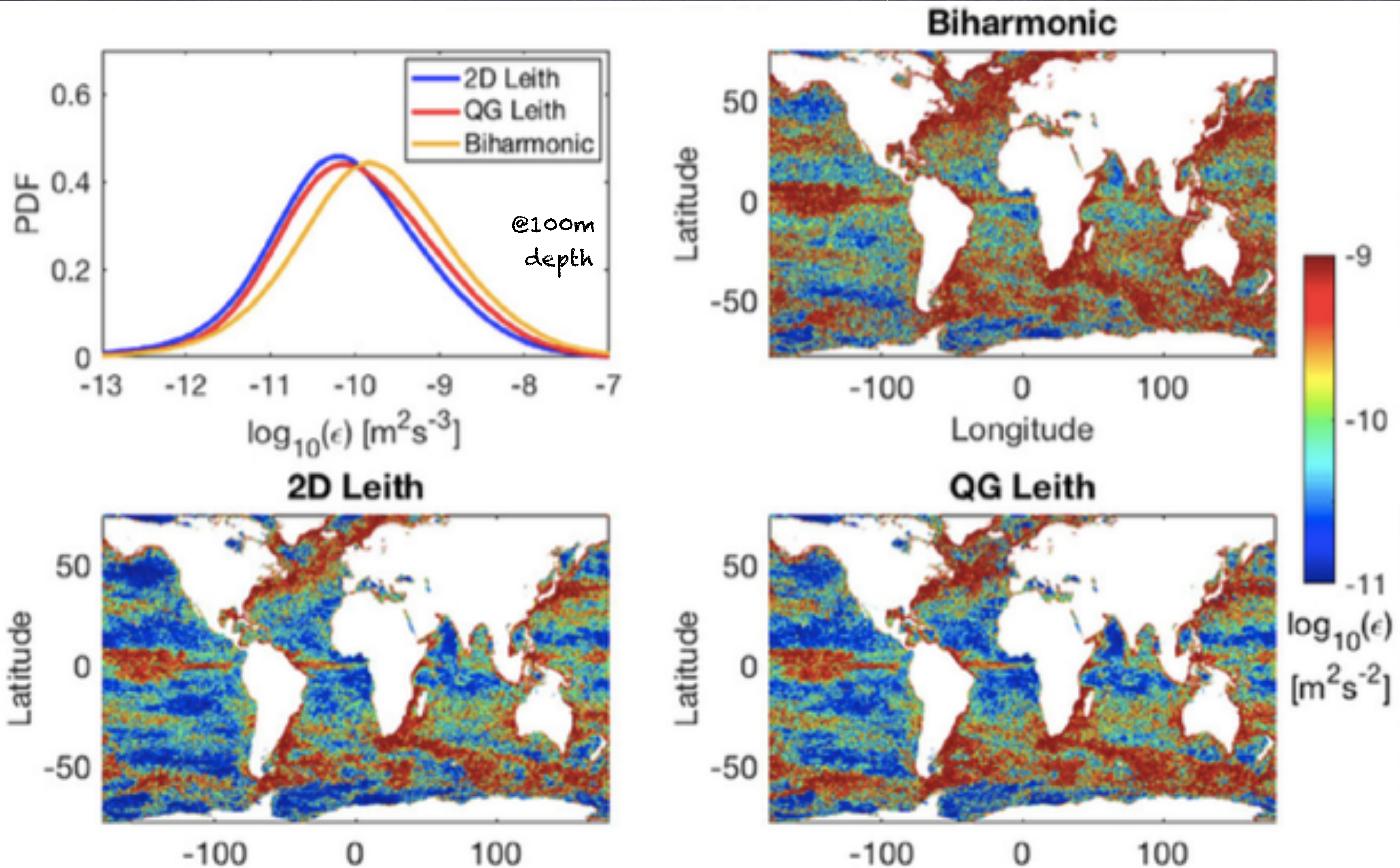
(most in upper 200m)

B. Pearson, BFK, S. D. Bachman, and F. O. Bryan, 2017:  
Evaluation of scale-aware subgrid mesoscale eddy models in a  
global eddy-rich model. *Ocean Modelling*, 115:42–58.





Energy Dissipation is Approx. Lognormally distributed—  
AND knows where the Gulf Stream is!





# Energy Dissipation Stats are Self-Similar

A (weak)  
dissipation of  
energy  
with pot'l  
enstrophy  
cascade

...

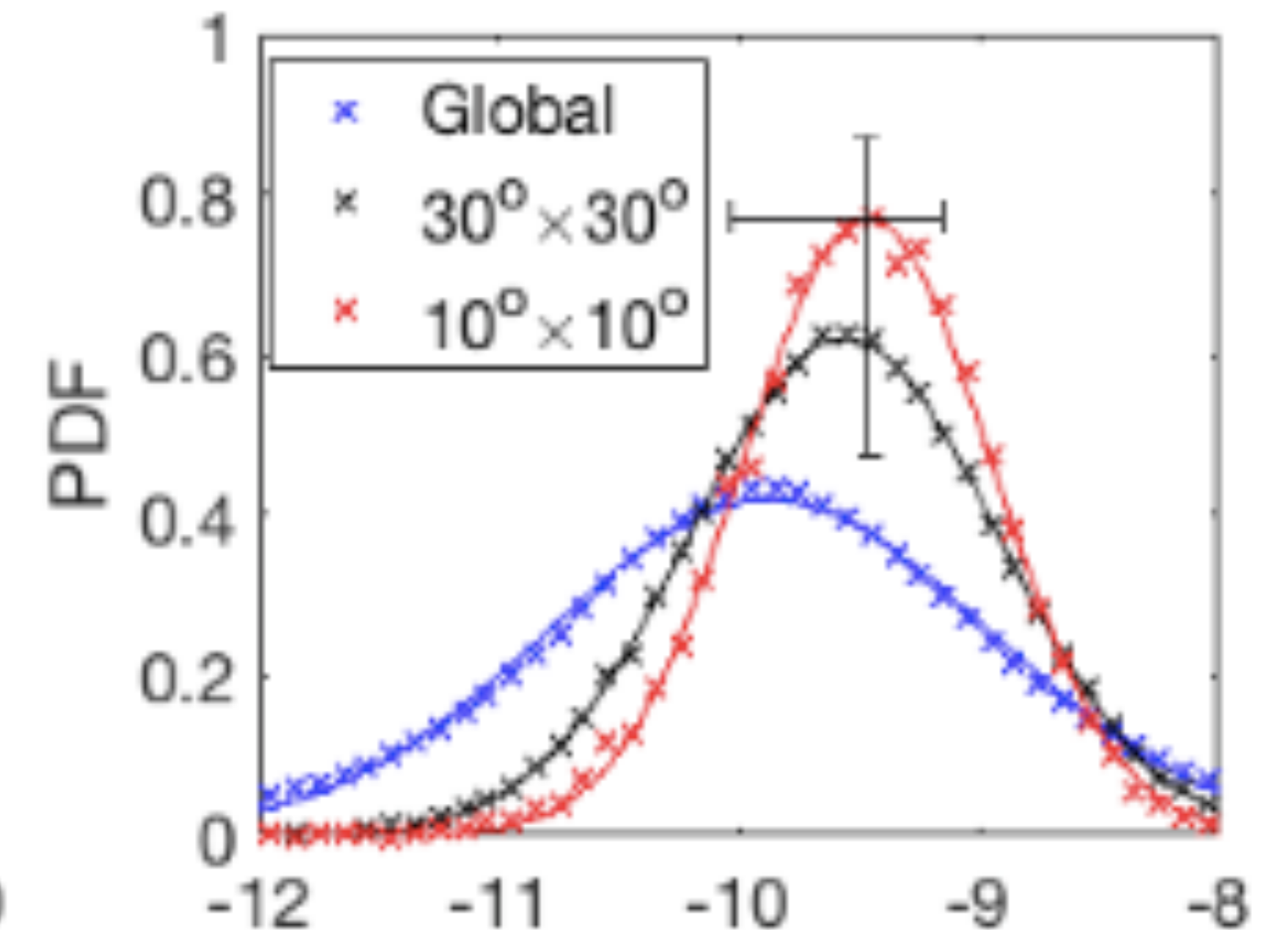
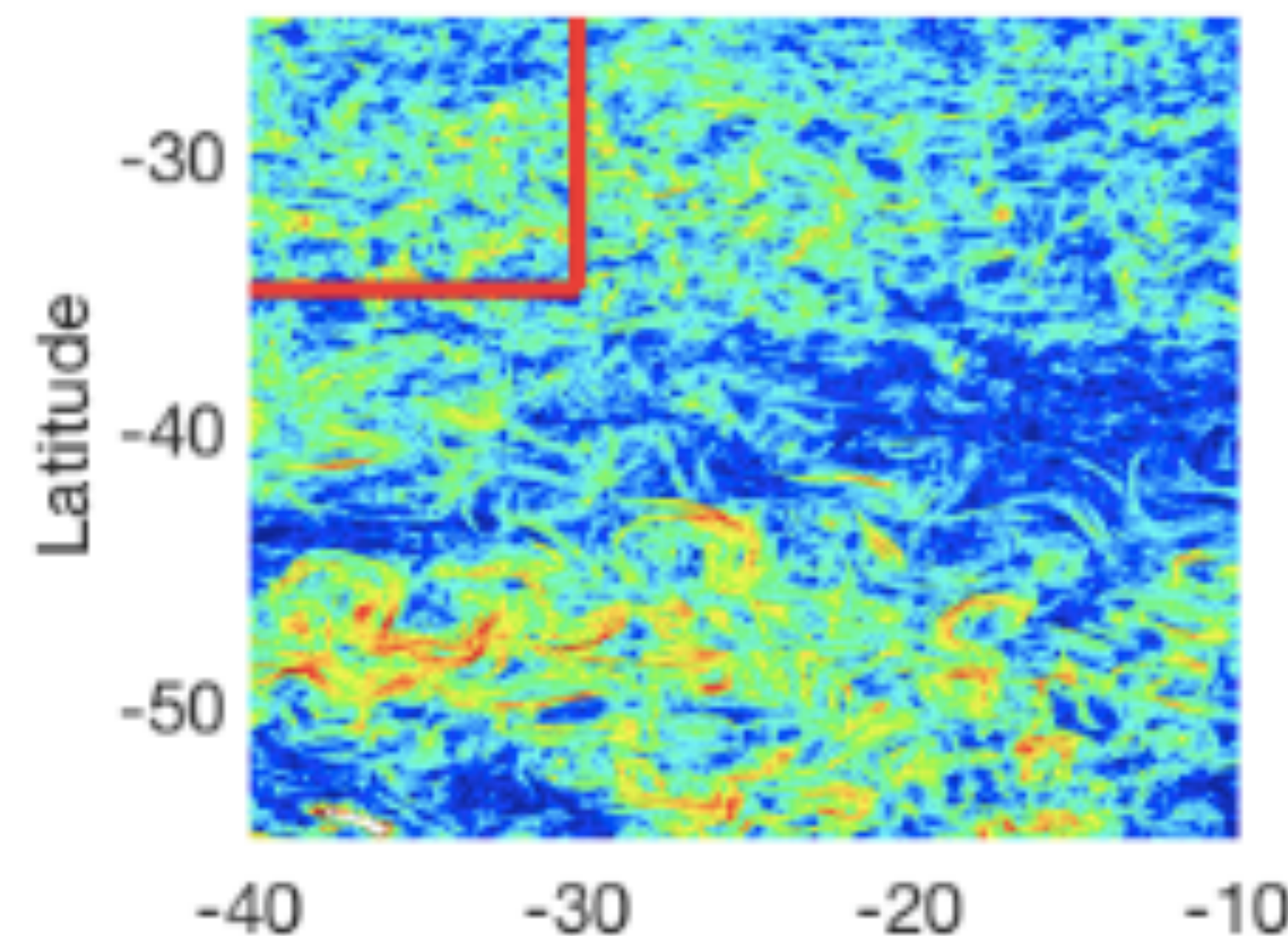
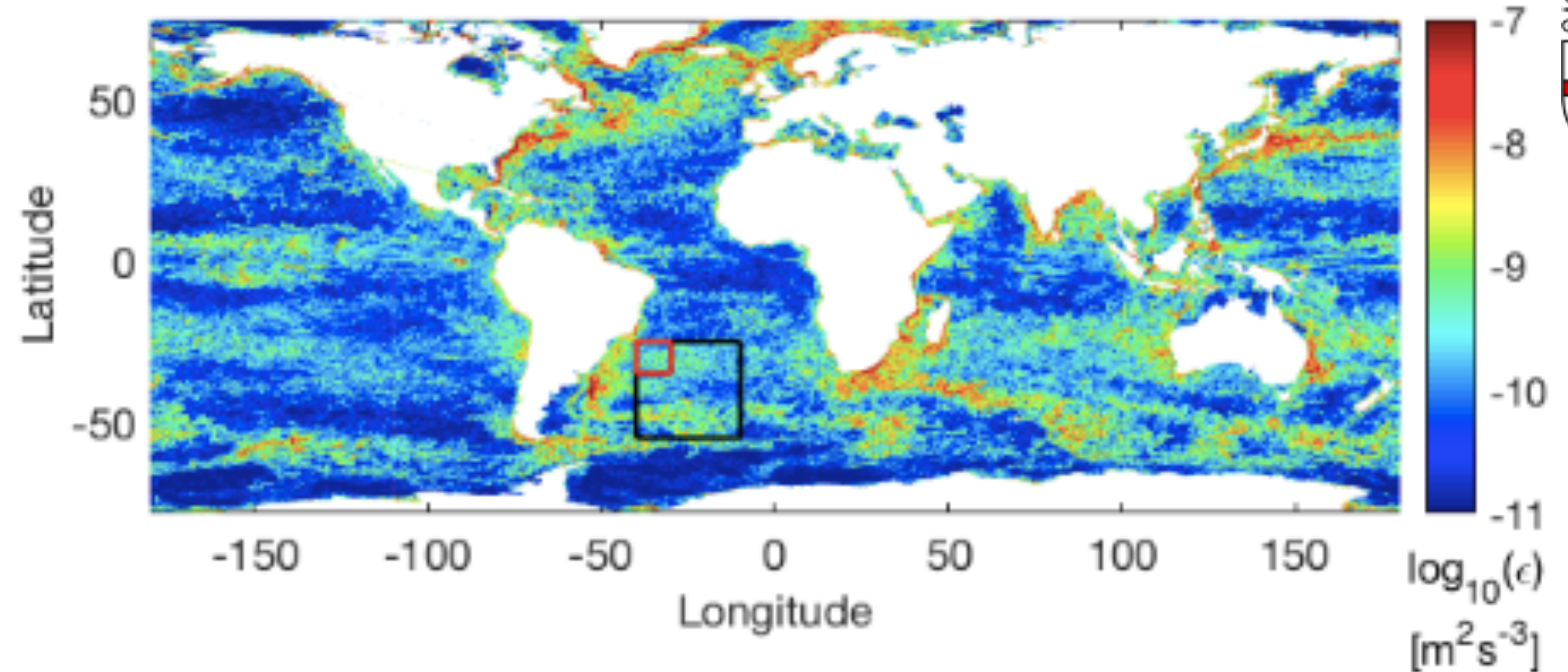
approx.  
lognormally  
distributed

(super-Yaglom '66)

...

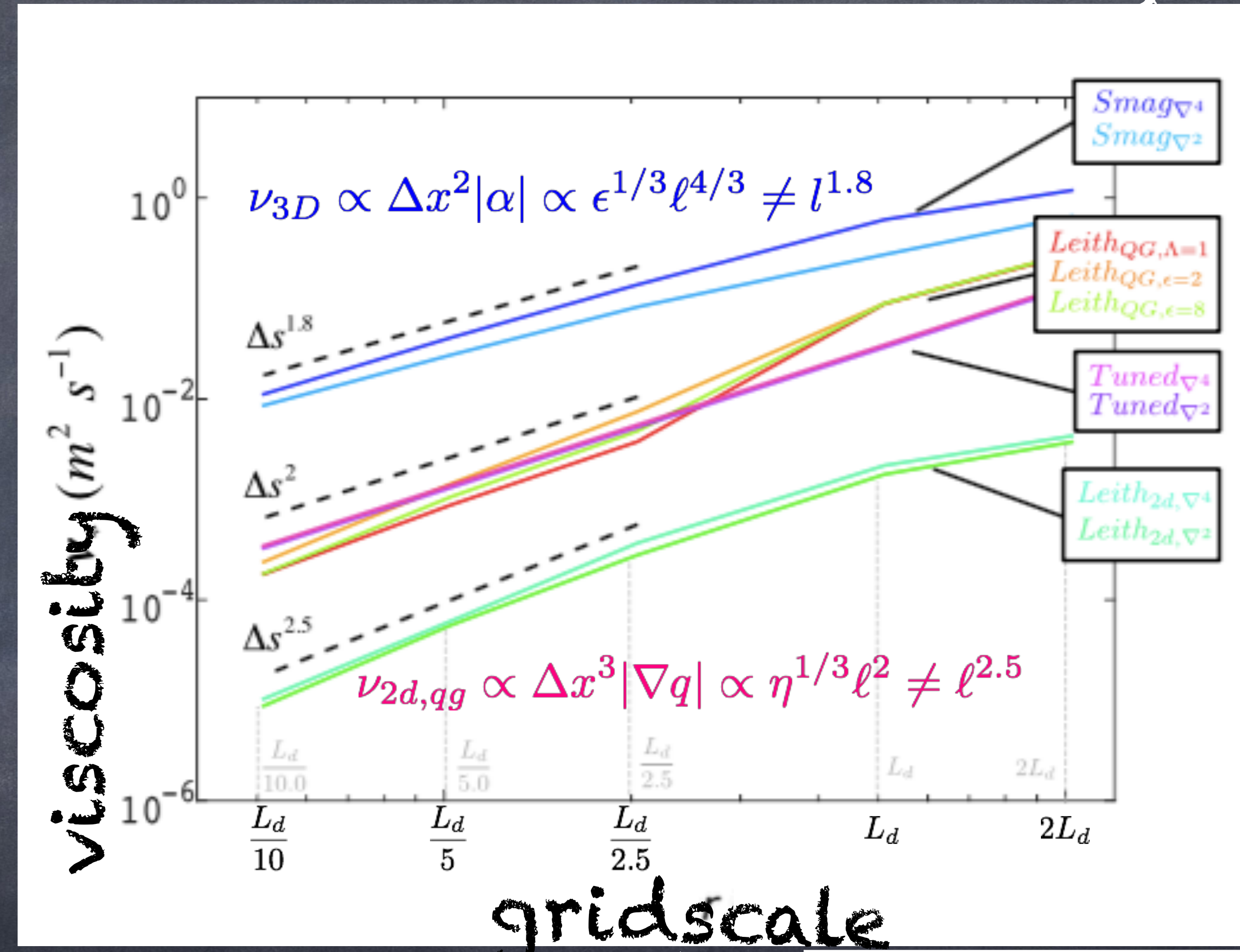
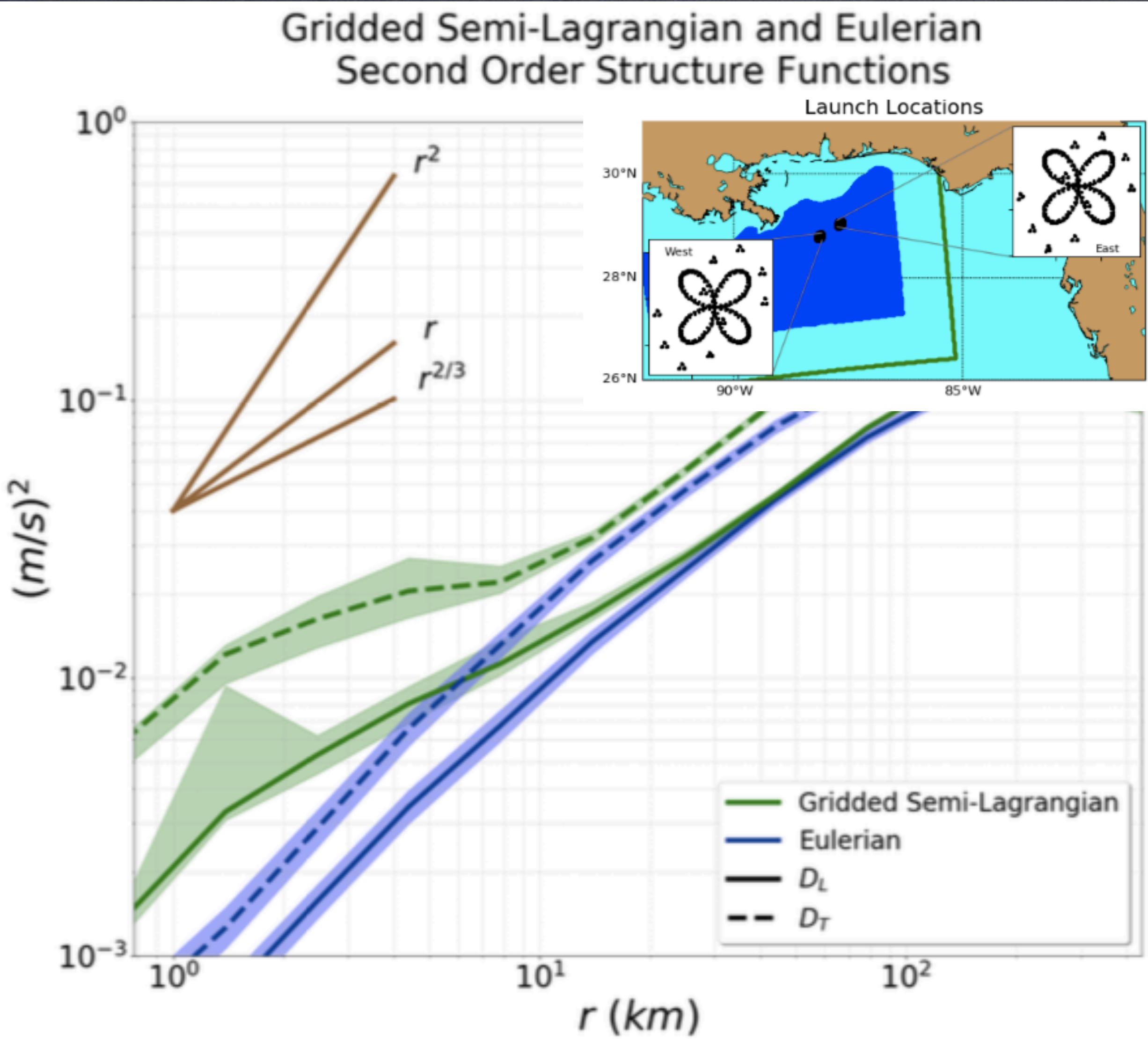
90% of KE  
dissipation in  
10% of ocean

B. Pearson and BFK. Log-normal  
turbulence dissipation in global  
ocean models. Physical Review  
Letters, 120(9):094501, March 2018.





# 'Observed' scale-sensitivity?



Some Theory/Model combos  
are inconsistent  
(e.g., Smagorinsky in a QG regime)

S. D. Bachman, B. Fox-Kemper, and B. Pearson, 2017: A scale-aware subgrid model for quasi-geostrophic turbulence. Journal of Geophysical Research—Oceans, 122:1529–1554.



J. Pearson, B. Fox-Kemper, R. Barkan, J. Choi, A. Bracco, and J. C. McWilliams. Impacts of convergence on Lagrangian statistics in the Gulf of Mexico. Journal of Physical Oceanography, 49(3):675-690, March 2019.





# Under "Cascade" Scalings, new bias is a little different

Grounded Param: Following Smagorinsky's 3D approach, we built schemes suitable for mesoscale-permitting ocean models, where 2D or QG cascades rule.

Intercomparison: By comparing across schemes, the Goldilocks test, the self-consistent scaling test, and numerical robustness select QGLEITH

Eval. Data (challenges): Lognormal dissipation, together with limited observing platforms (e.g., drifters), makes observing dissipation & scalings challenging.

Key Impacts: Global KE budget, major currents, adiabatic scheme. Watermasses? Heat uptake?



# On Redi-ness

$$\nu = \frac{\partial_z \left( k_{gm} \frac{\nabla_h b}{\partial_z b} \right)}{\partial_z \left( \frac{\nabla_h b}{\partial_z b} \right) - \frac{\partial_y f}{f}} \approx k_{gm} + \frac{\partial_z k_{gm}}{\partial_z \left( \ln \frac{|\nabla_h b|}{\partial_z b} \right)}$$

$$k_{gm} = \nu - \frac{\int_{-H}^z \frac{\nabla_h b}{\partial_z b} \partial_z \nu - \nu \frac{\partial_y f}{f} dz}{\frac{\nabla_h b}{\partial_z b}} \approx - \frac{\int_{-H}^z \frac{|\nabla_h| b}{\partial_z b} \partial_z \nu dz}{\frac{|\nabla_h| b}{\partial_z b}}$$