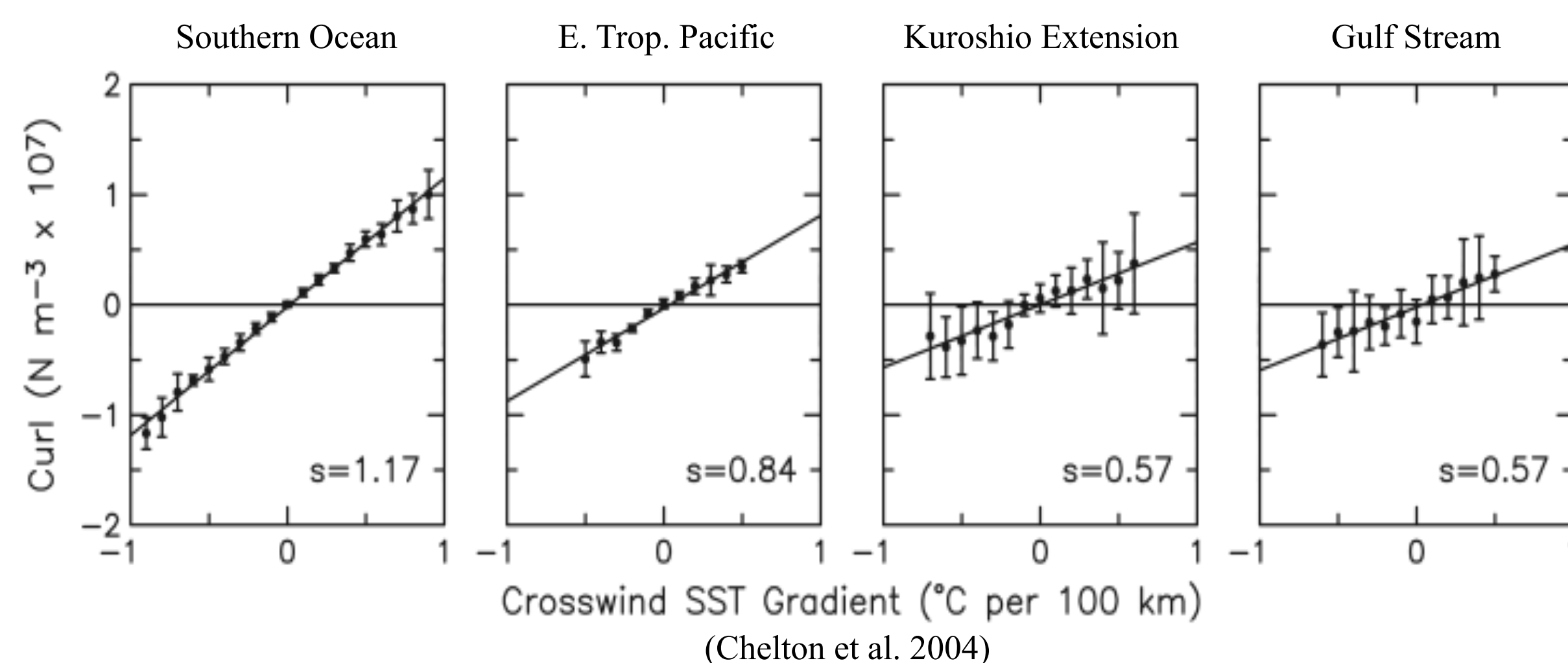
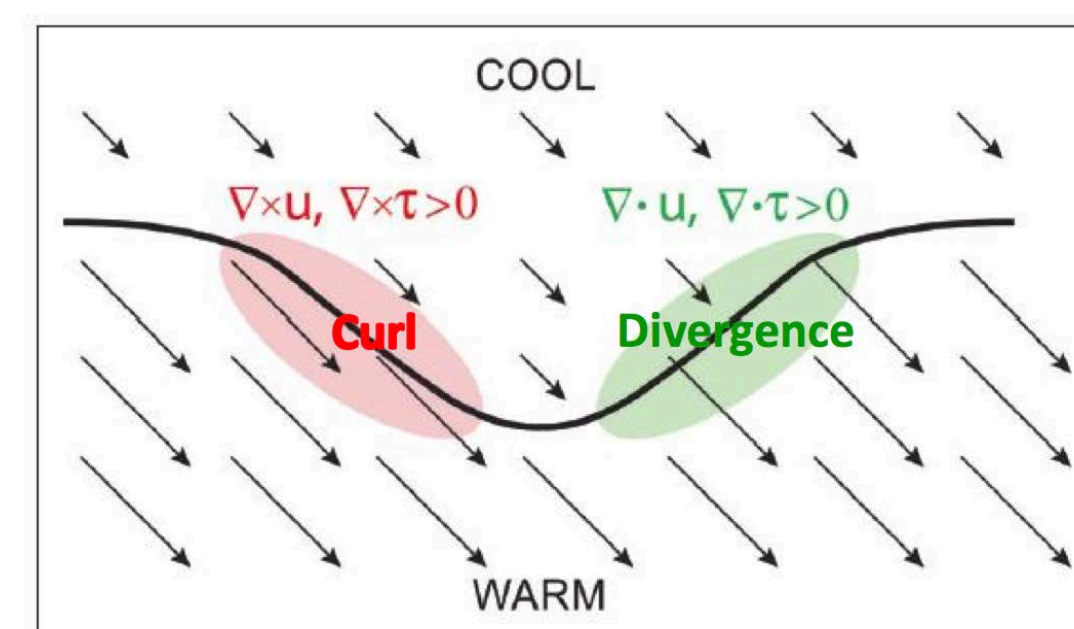


Introduction

- Positive linear relationship between wind stress curl and crosswind sea surface temperature (SST) gradient has been found by Chelton et al. (2004) on satellite observations (25-km resolution).



- Stronger wind stress over warm water is argued to be the mechanism.



- In the submesoscale regime, SST gradients are much stronger than those in the mesoscale and larger scale regimes. And surface current vorticity is significantly robust in submesoscale regime.

$$W_{total} = \frac{1}{\rho} \nabla \times \left(\frac{\vec{\tau}}{f + \zeta} \right) \quad (\text{Stem, 1965})$$

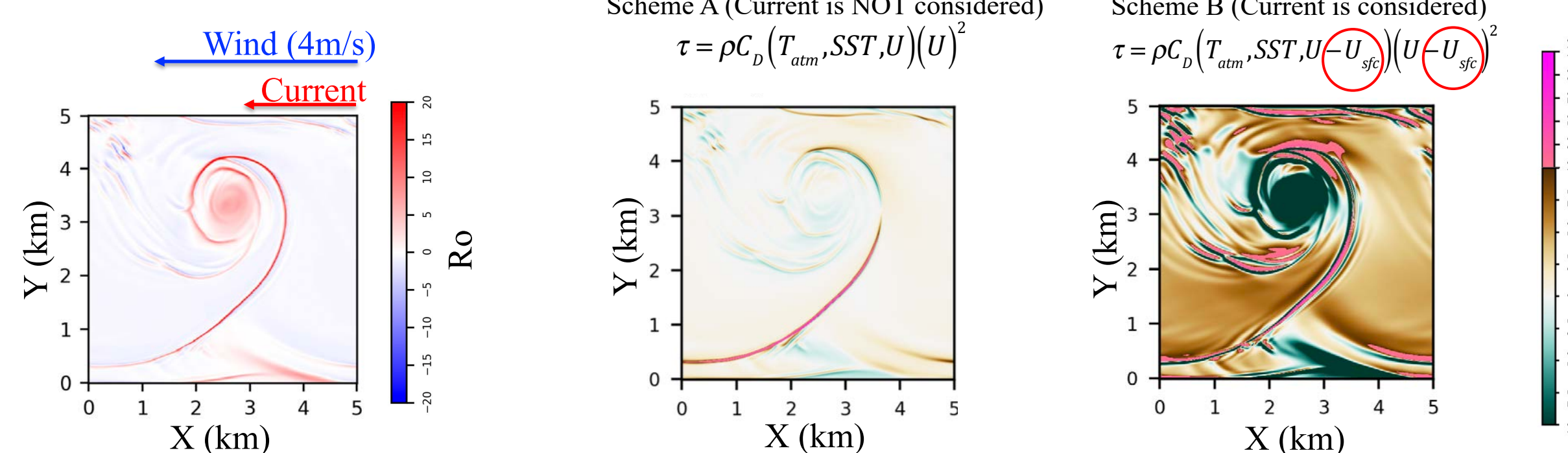
- Negative potential vorticity (PV) injection from the ocean surface is crucial for triggering instability in the upper ocean layer.

Goals

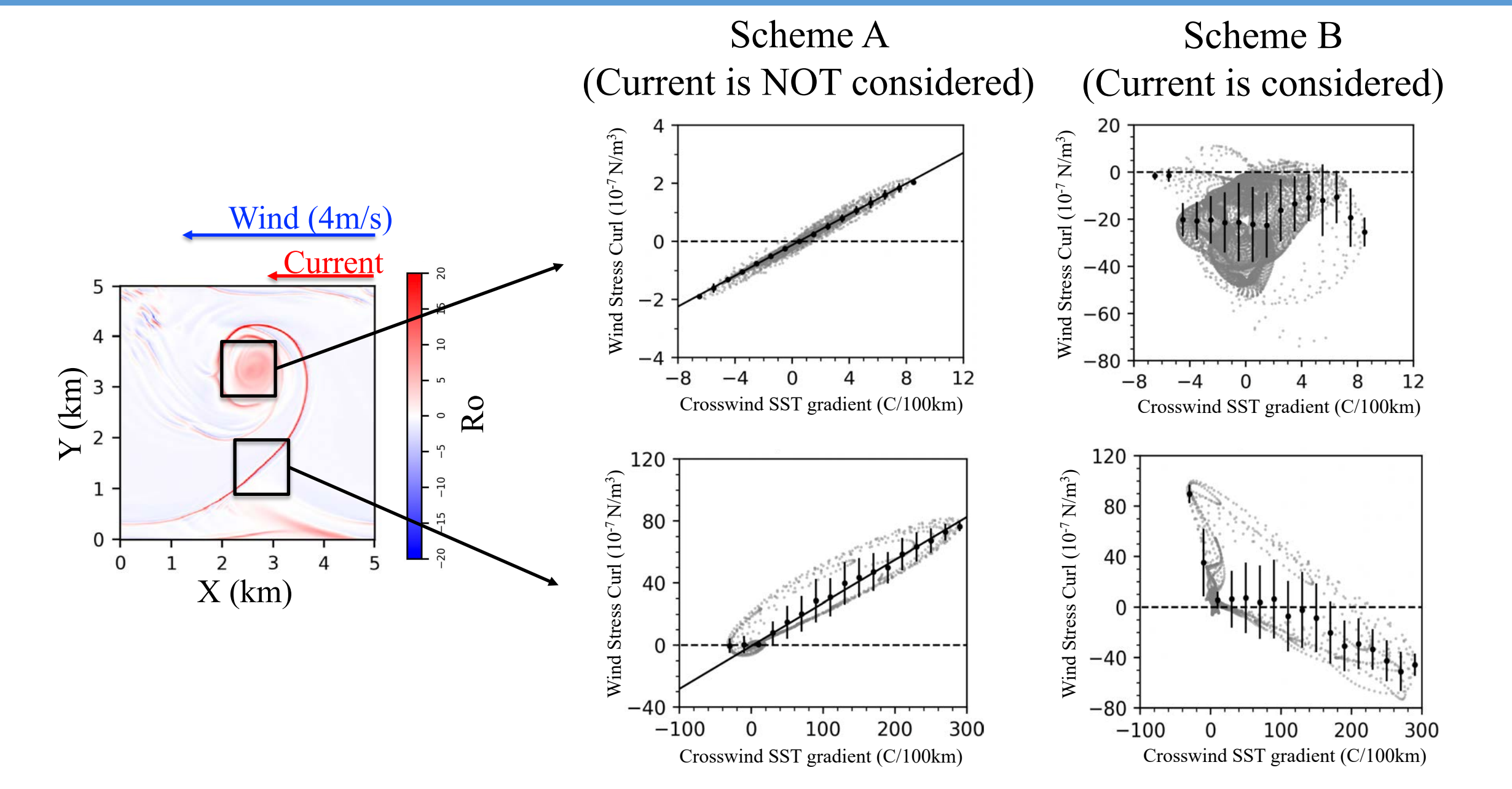
- Determine (quantitatively) if wind stress field is significantly influenced by submesoscale surface features (SST gradient and current vorticity)?
- Assess (quantitatively) the effect of including surface current in air-sea turbulent flux on the PV surface flux and vertical transports.

Model Setup

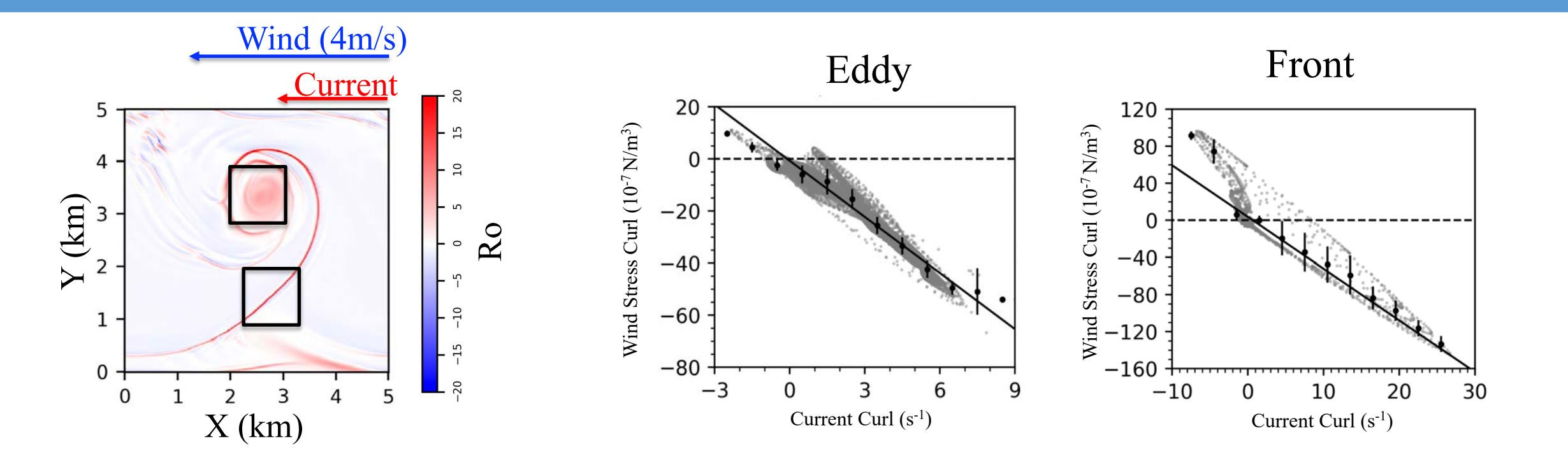
- Ocean: MITgcm; Atmospheric Boundary Layer: CheapAML
- Resolution: 10 m horizontal; 2 m vertical.



Wind Stress Curl & Crosswind SST Gradient (Scheme A & B)

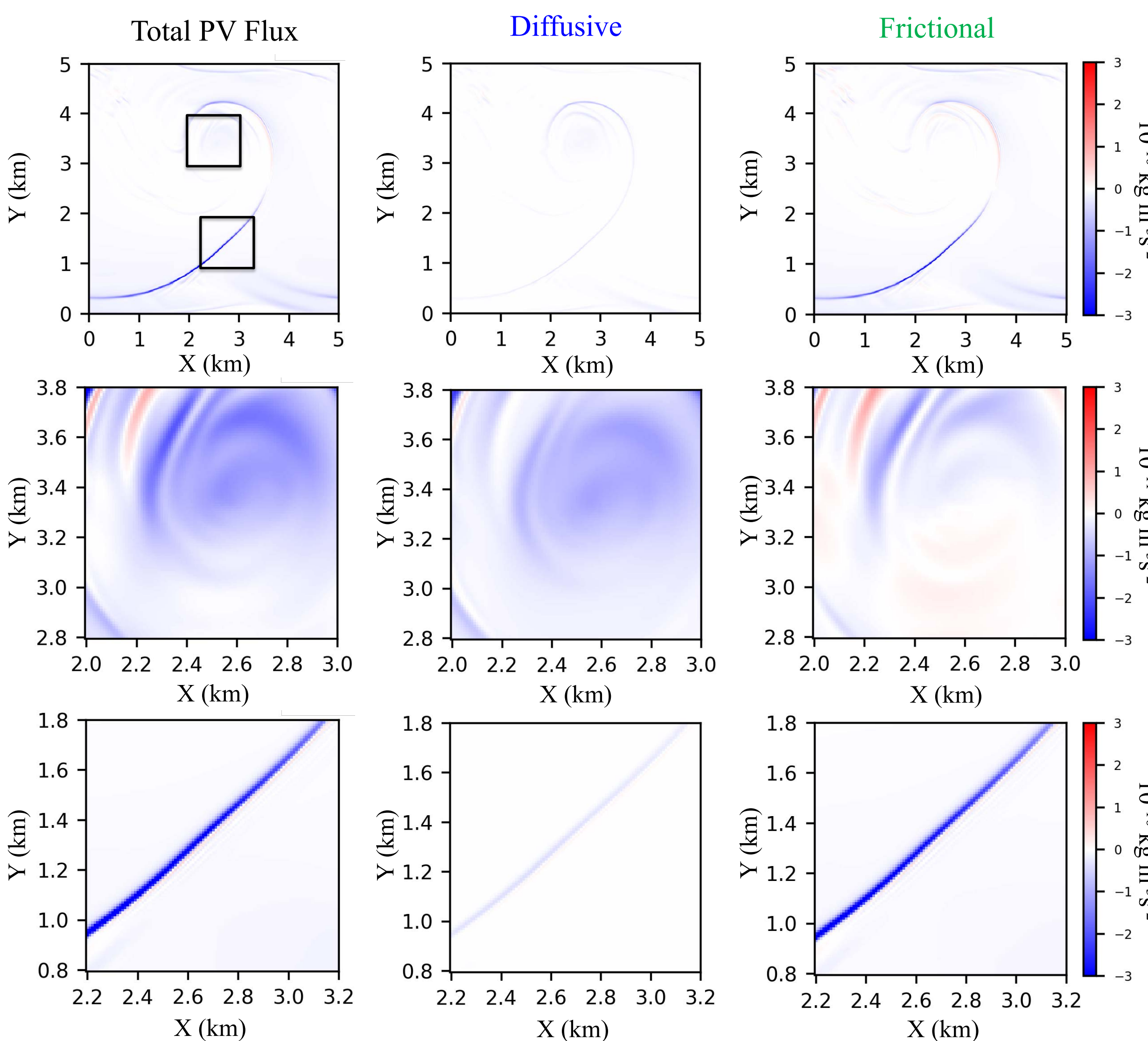


Wind Stress Curl & Surface Current Vorticity (Scheme B)



PV Flux at Surface (Scheme B – Scheme A)

$$\vec{j} = \vec{\omega} \frac{D\sigma}{Dt} + \vec{F} \times \nabla \sigma \quad (\text{Marshall et al. 2001})$$

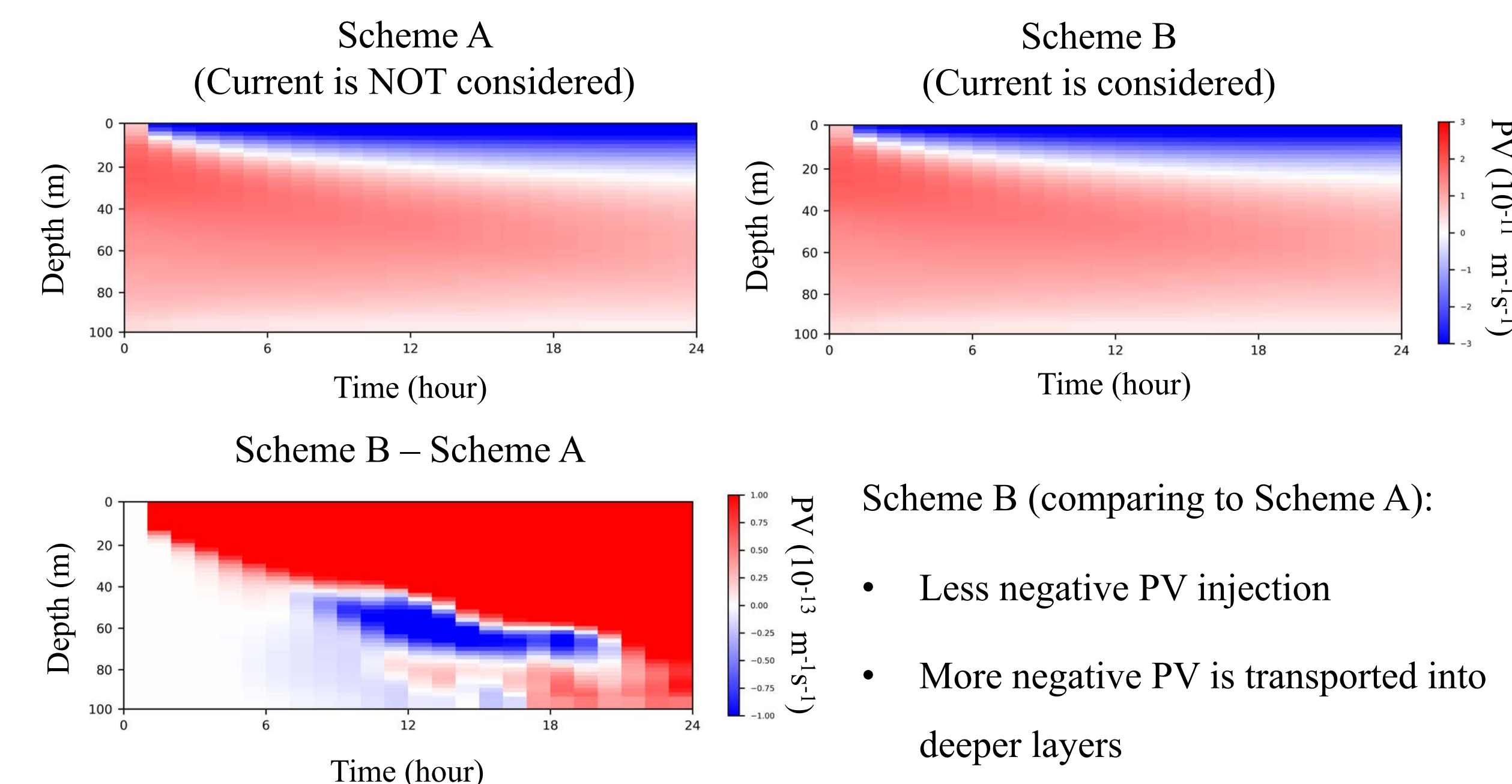


References

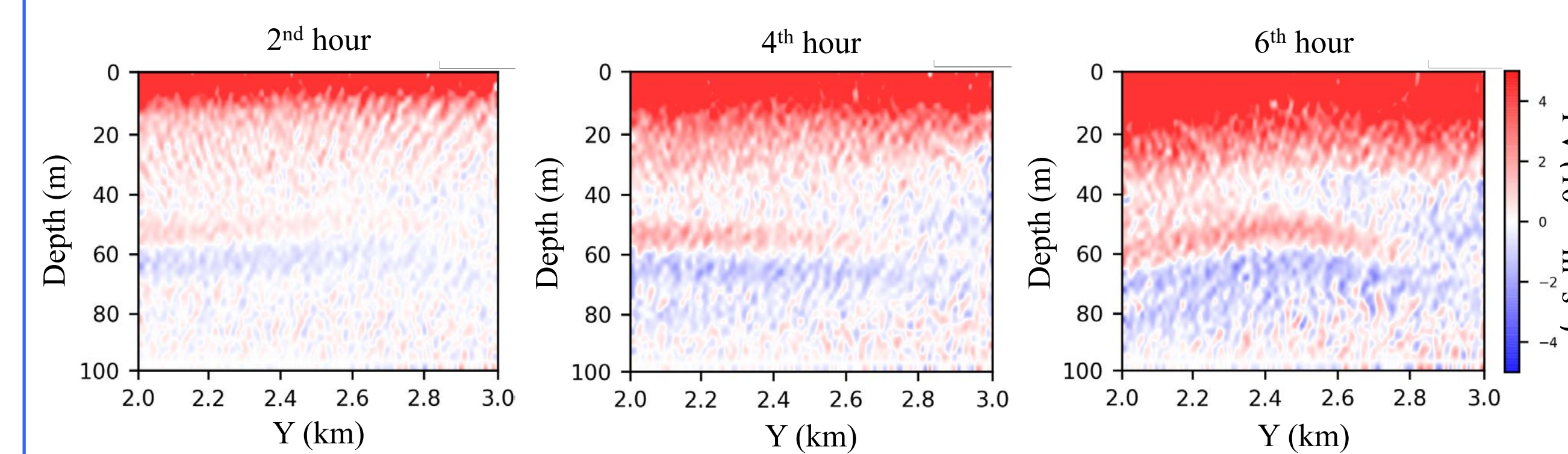
- Chelton, Dudley B., Schlax, M. G., Freilich, M. H., & Milliff, R. F. (2004). Satellite measurements reveal persistent small-scale features in ocean winds. *Science*, 303, 978–983, <https://doi.org/10.1126/science.1091901>.
- Chelton, Dudley B., Schlax, M. G., and R. M. Samelson, (2007). Summertime coupling between sea surface temperature and wind stress in the California Current System. *J. Phys. Oceanogr.*, 37, 495–517, <https://doi.org/10.1175/JPO3025.1>.
- Marshall, J. C., D. Jamous, and J. Nilsson, (2001) Entry, flux, and exit of potential vorticity in ocean circulation. *J. Phys. Oceanogr.*, 31, 777–789, [https://doi.org/10.1175/1520-0485\(2001\)031<0777:EFAEOP.2.CO;2](https://doi.org/10.1175/1520-0485(2001)031<0777:EFAEOP.2.CO;2)
- Stern, W., (1965). Interaction of a uniform wind stress with a geostrophic vortex. *Deep-Sea Res.*, 12, 355-367.

Hovmöller Diagrams of Horizontally Averaged PV

- Comparison between 24-hour experiments using Scheme A and Scheme B

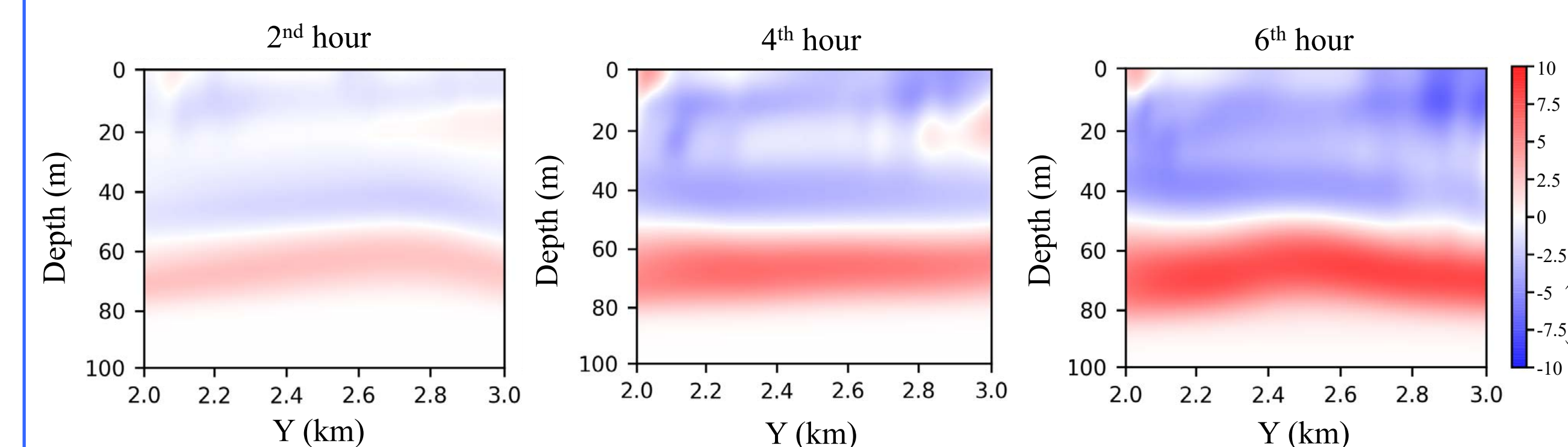


Difference of PV Averaged along X direction (Scheme B – A)

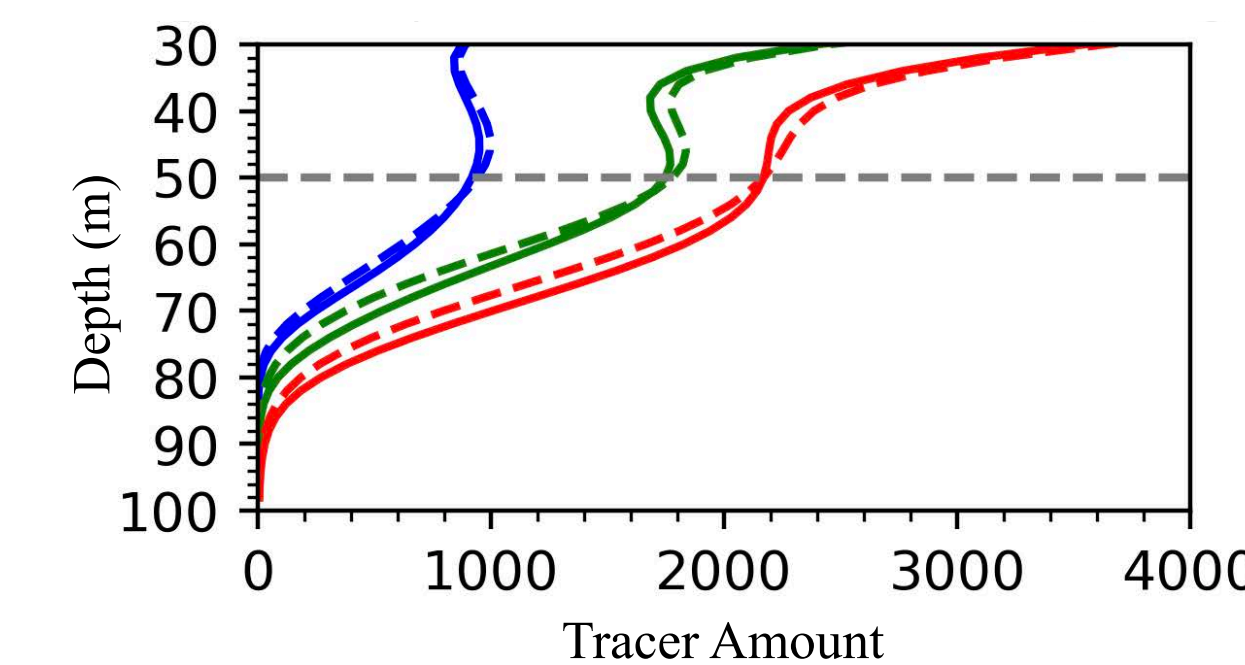


Vertical Transport of Passive Tracer from Surface

- Difference of Tracer Averaged along X direction (Scheme B – Scheme A)



- Comparison of horizontally averaged tracer profiles at 2nd, 4th, 6th hour (dashed: A; solid: B).



In Scheme B (compare to Scheme A):

- More passive tracer is transported into layers beneath 50 m depth.

- Table: Quantitative assessment of passive tracer transported into layers beneath 50 m depth.

Time (hour)	Current (B)	No Current (A)	Current – No Current (B-A)	Percentage Increase (%)
1	4064	3966	98	2.5
2	8294	7790	504	6.5
3	11856	10808	1048	9.7
4	16287	14788	1499	10.1
5	20477	18529	1948	10.5
6	24339	21963	2376	10.8