CPT: Representing internal-wave driven mixing in global ocean models

http://www-pord.ucsd.edu/~jen/cpt/

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The Team:

simulations. parameterizations for use in representing internal wave driven mixing in climate Goal: Development of ocean mixing

mixing processes at internal wave generation sites Near-field parameterizations accounting for

2) A new parameterization for the mixing resulting from the breakdown of near inertial energy transported in the wave field

sources wave energy in the ocean interior far away from 3) Parameterization for the breakdown of internal

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Overview

- . Most diapycnal (vertical) mixing in the ocean interior is due to breaking internal gravity waves
- 2. Mixing is patchy in space and time, reflecting the complex geography of internal wave generation, propagation, and dissipation.
- 3. Patchy mixing matters for ocean circulation and fluxes. It's important to "get it right".
- 4. Our plan: use what we collectively know about internal wave physics to develop dynamic parameterizations of diapycnal mixing that can evolve in a changing climate.

Internal gravity wave primer



- Low-mode ~interfacial waves
- High-mode ~ plane waves
- **Rapid:** $f \leq \omega \leq N$
- Breaking waves are at small (I-I0 m) scales



Parameterizing mixing

Cant' explicitly resolve internal waves in climate models. 3 steps to parameterize their role:

Wave generation
 Wave propagation
 Wave dissipation

Near-Inertial Waves







GOLD SIMULATION Alford NIW fluxes (80 moorings)



Wednesday, July 20, 2011

Internal Tides



Internal Tides



N.B., strong dependence on the wind forcing employed!



Maps of Near-Inertial Wave Generation

2) Wave propagation 3) Wave dissipation producing elevated mixing up) Wave generation Internal tides propagating up into the main thermocline bathymetry steadily break, from the rough (eastern) Cant' explicitly resolve internal waves in climate models. Global pattern of mixing that mirrors wave generation 3 steps to parameterize their role: Brazil Basin Parameterizing mixing Water depth (m) -3500 -3000--2000--2500--1500--1000--5000 -5500 -4500 -4000 -500-Polzin et al. 97 0.1 ŝ 0.2 0.3 <u></u> ³ 4 ŝ 0.4 -30 0.5 Diffusivity (10-4 m²s-1) 0.6 Longitude -28 0.7 -26 0.8 -24 0.9 -22 2.0 5.0 20

8.0

22.0

-18

-16





Desirable Properties of Parameterizations

- physical inconsistencies. Energy-based prescriptions are more likely to avoid
- calculating the energy conversion to allow for negative Should use model's own properties (e.g., stratification) for feedbacks on mixing.
- Static maps can get scary when the control the ocean structure.
- to the changing model state. Energy fluxes (and hence buoyancy fluxes) should respond
- of wavenumber bins will be a challenge. Vertically elliptic or iterative equations are O.K.; Horizontally elliptic equations, iterations, or large numbers
- Minimize dimensional "parameters" wherever possible.
- Should work across a range of model resolutions.
- $(1^{\circ} 1/8^{\circ} \text{ climate models now.})$

Some more complex implementation in more complex models Low Latitudes t conveyo heat **High Latitudes**

Turbulent mixing helps make the ocean go round

Some more complex implementation in more complex models





Turbulent mixing helps make the ocean go round





0 ₀

10⁰ ε[m² s⁻³]





10⁻¹⁰

10[°] ε[m² s⁻³]





10⁻¹⁰

10[°] ε[m² s⁻³]





mixing in the upper ocean under storm tracks

10⁻¹⁰

10[°] ε[m² s⁻³]







Hallberg

Elevated mixing over rough topography CM2M (in AR5)

CM2.1 (in AR4)





Zonal mean Atlantic temperature bias





Simmons/Alford, this talk. Continued process-oriented GCM analysis & data comparisons:

GCM investigation of energetically constrained, spatially variable mixing: Jochum-NCAR - Implementation of NIW parameterization --Upside down version of St. Laurent et al (2002)

Hallberg, Legg & Griffies (GFDL) - tidal and other contributions to spatial variable & topographically controlled mixing --this talk

