

The wind stress drives the coherence of the North Atlantic Meridional OVERTURNING CIRCULATION on seasonal and longer time scales

Shane Eliot (RSMAS)

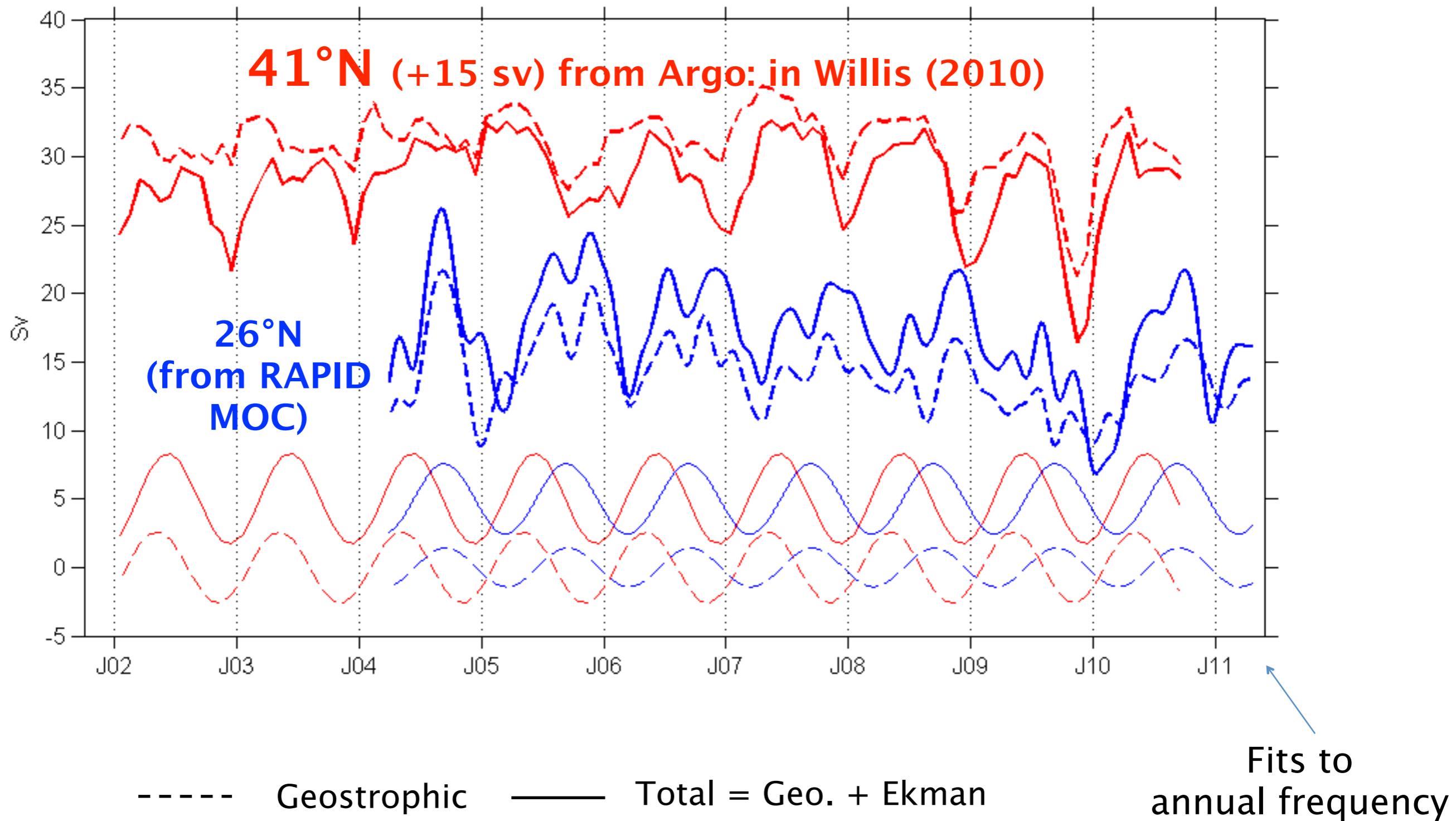
with Eleanor Frajka-Williams (Southampton-NOC), Chris Hughes (NOC Liverpool), Sofia Olhede (UCL), Matthias Lankhorst (SIO)

With thanks to:

[Ric Williams](#) (University of Liverpool), [Miguel Morales Maqueda](#), (NOC), [Josh Willis](#), (NASA/JPL)
NERC Rapid WAVE project in collaboration with :

[Woods Hole Oceanographic Institution – Line W program \(John Toole\)](#)
[MOVE Project at Scripps Institution of Oceanography \(Uwe Send\)](#)

MOC strength at two latitudes:



See e.g Kanzow et al. 2007, Kanzow et al. 2010, Johns et al. 2011, McCarthy et al. 2012, Mielke et al. 2012

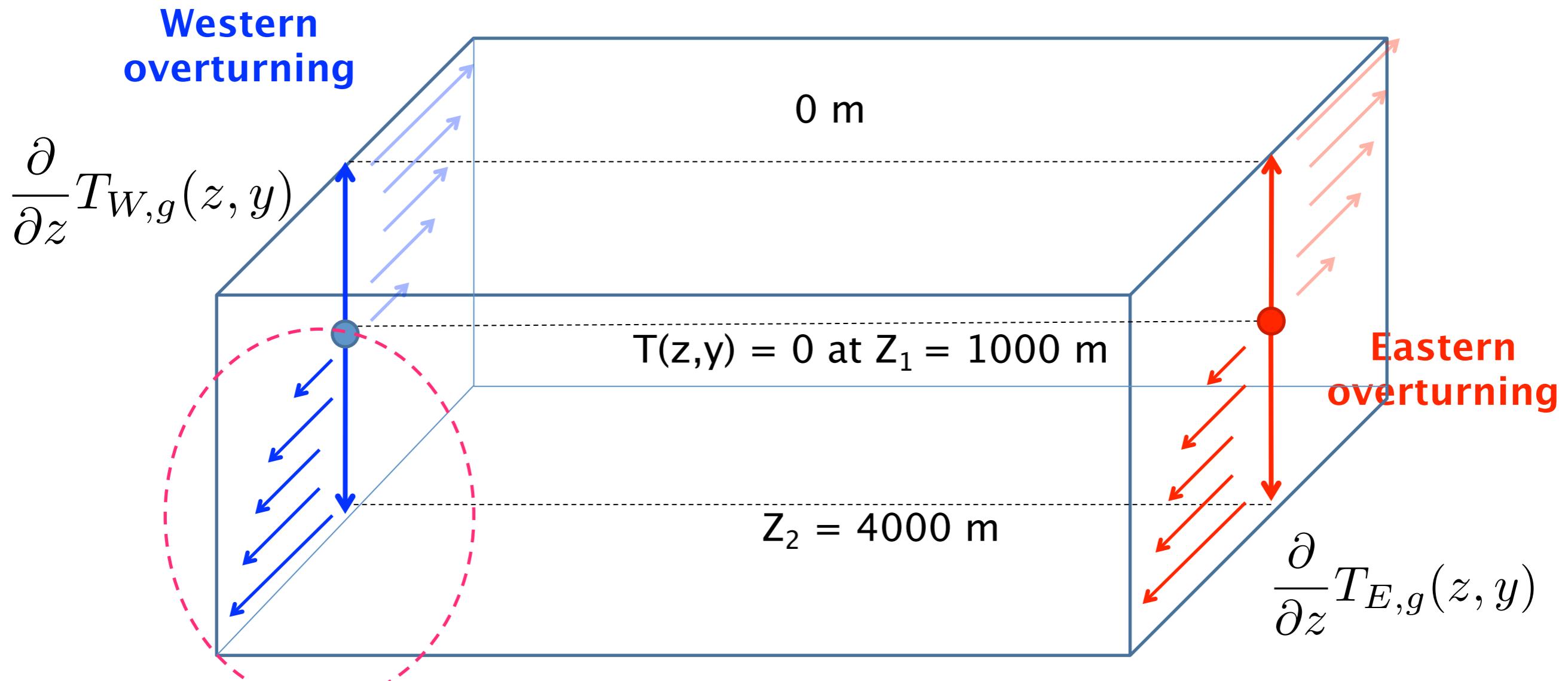
Goals of study:

1. Derive a quantity from **observations** which is representative of the MOC in the North Atlantic ocean
2. Study the **meridional coherence of overturning processes** in the North Atlantic: role of wind forcing?

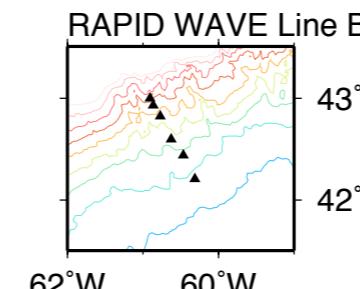
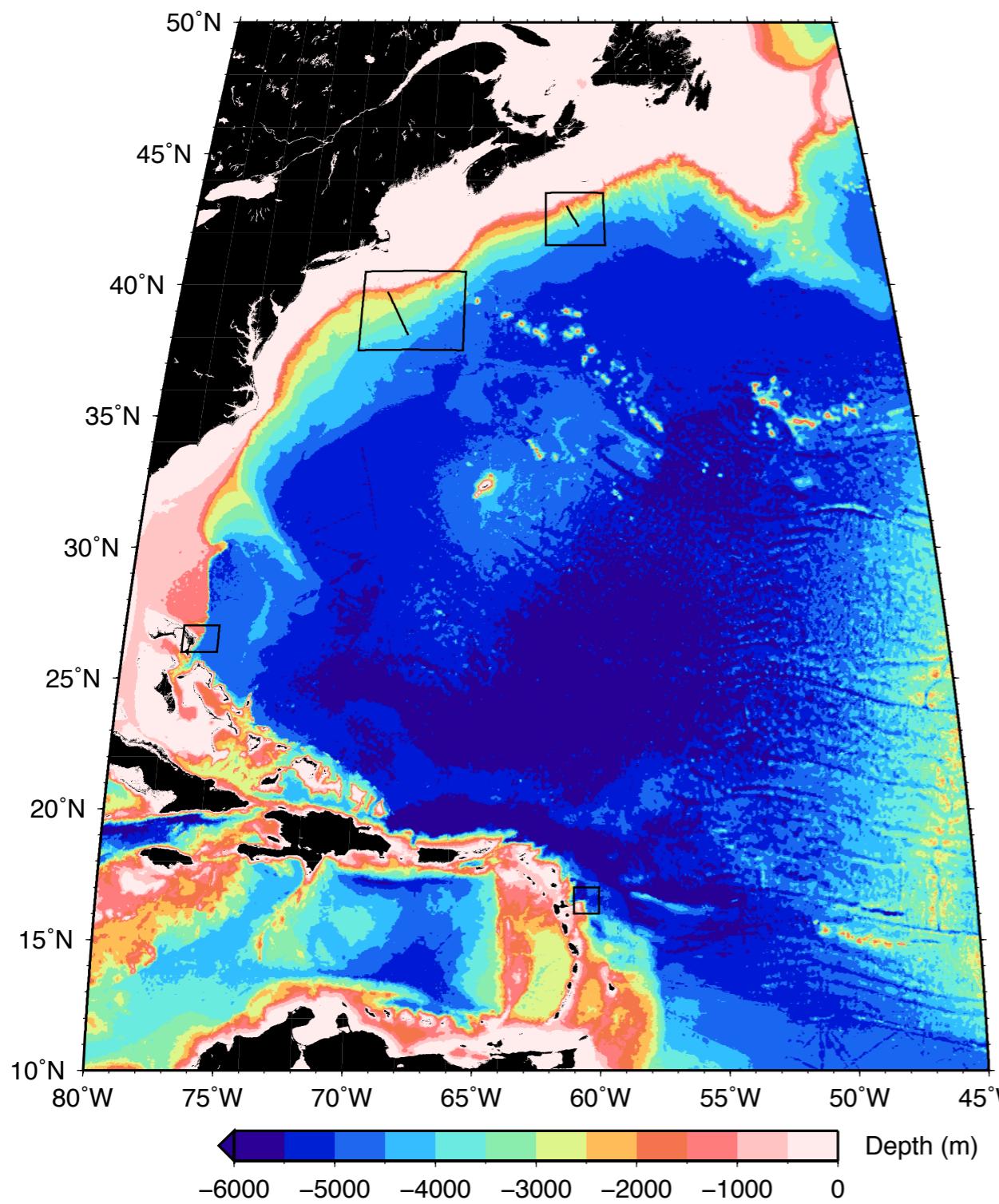
Eastern and western overturning transports below 1000 m

$$T_g(z, y) = \int_W^E \rho v_g dx \quad : \text{zonally integrated meridional geostrophic transport}$$

$$\frac{\partial}{\partial z} T_g(z, y) = \frac{1}{f} \frac{\partial}{\partial z} [p_E(z, y) - p_W(z, y)] = \frac{\partial}{\partial z} T_{E,g}(z, y) + \frac{\partial}{\partial z} T_{W,g}(z, y)$$

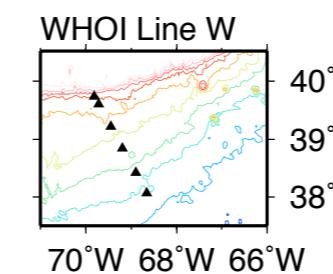


4 arrays to obtain western overturning transports below and relative to 1000 m:



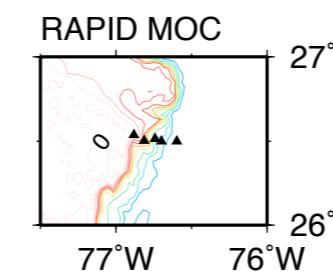
RAPID WAVE array
Line B/Line RS
(NO/C/BIO)

Hughes et al. (2013)
Elipot et al. (2013)



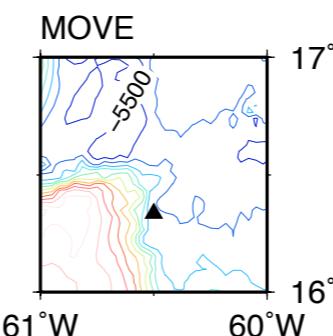
WHOI Line W array

Toole et al. (2011)



RAPID MOC/MOCHA array
RSMAS/AOML/NOC

Rayner et al. (2011)

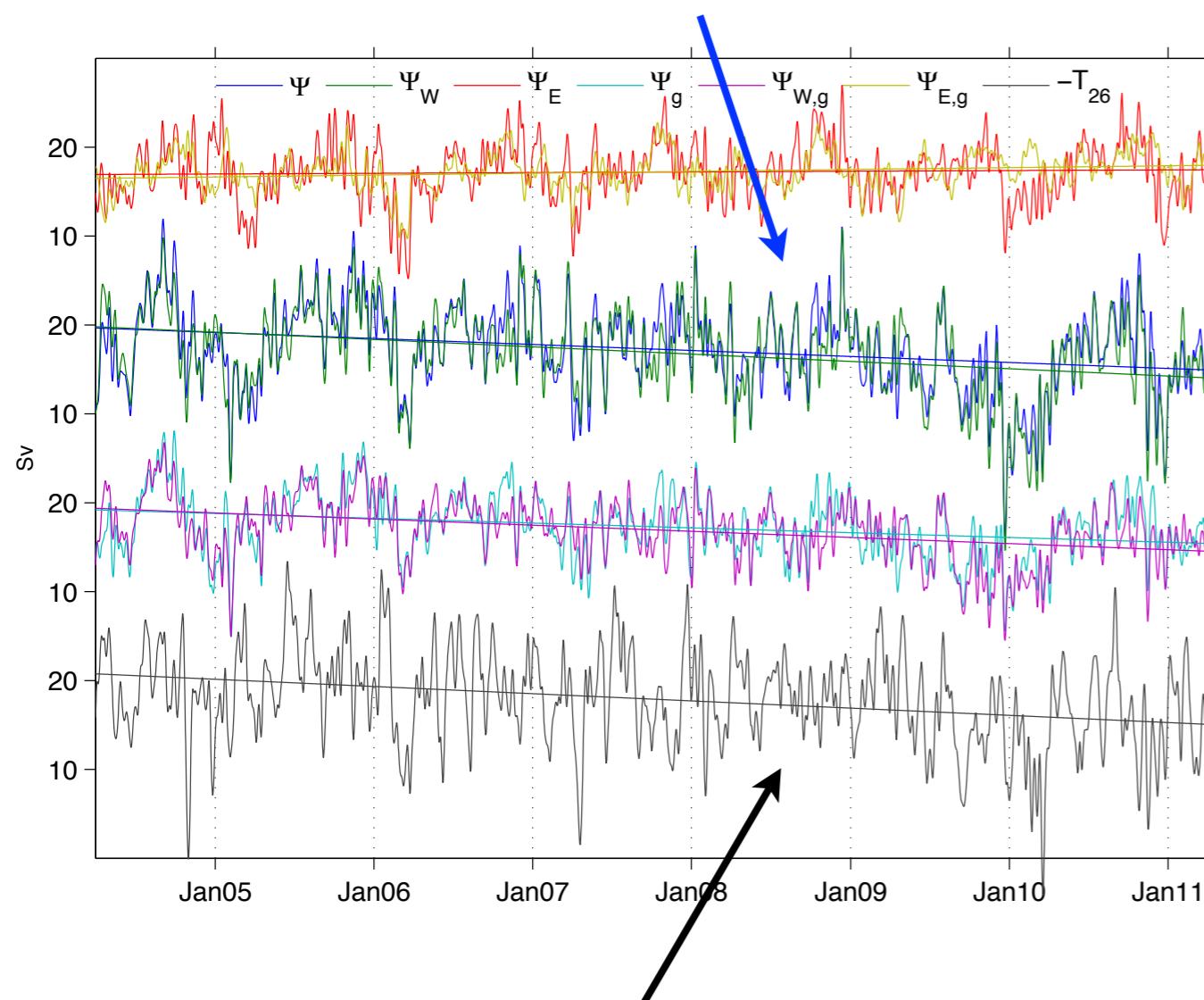


MOVE array
(SIO)

Send et al. (2011)

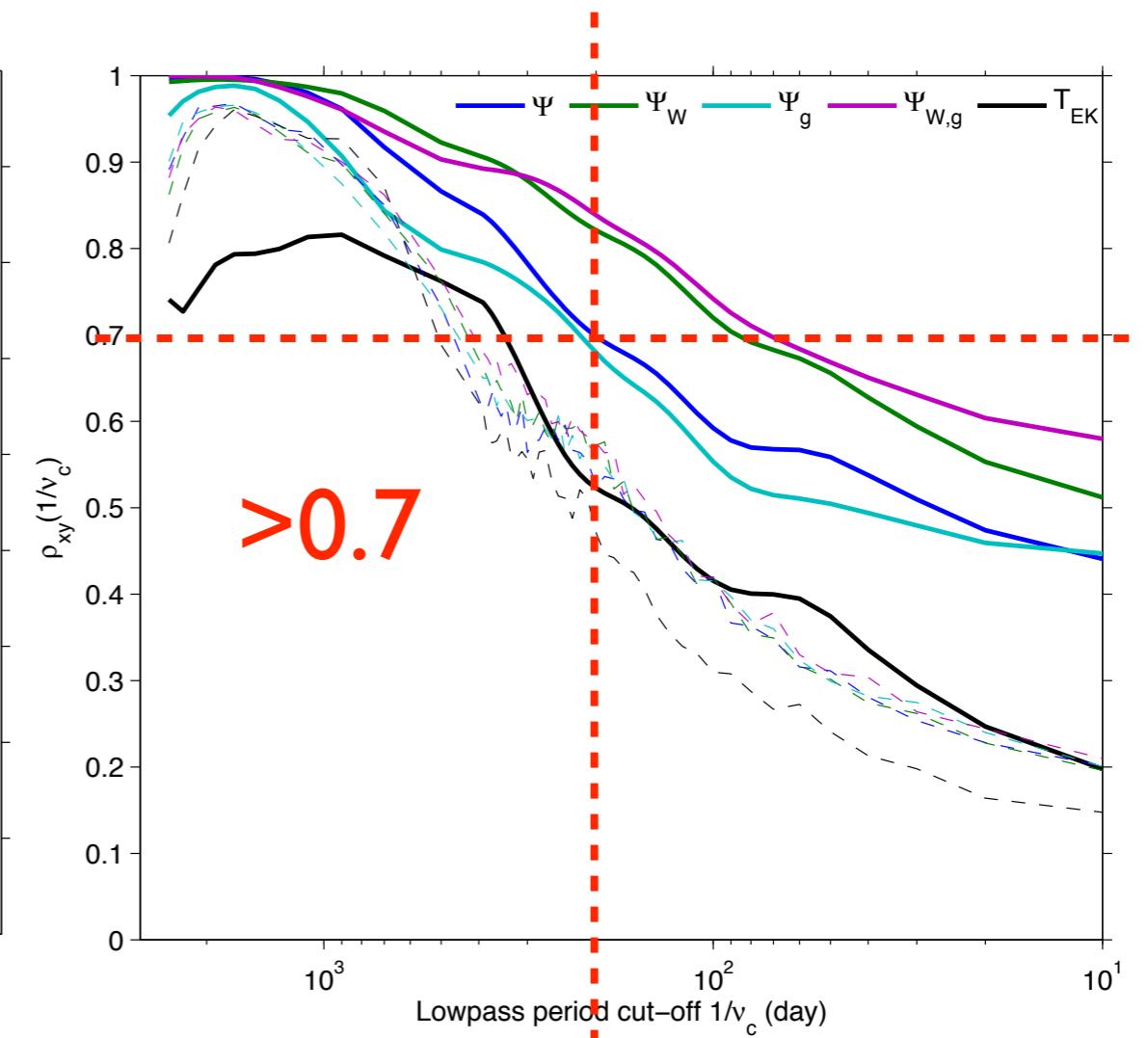
OVERTURNING TRANSPORT BELOW 1000 m AND MOC AT 26°N?

“MOC” transport



Negative of overturning transport
below and relative to 1000 m (\mathcal{T}_{26})

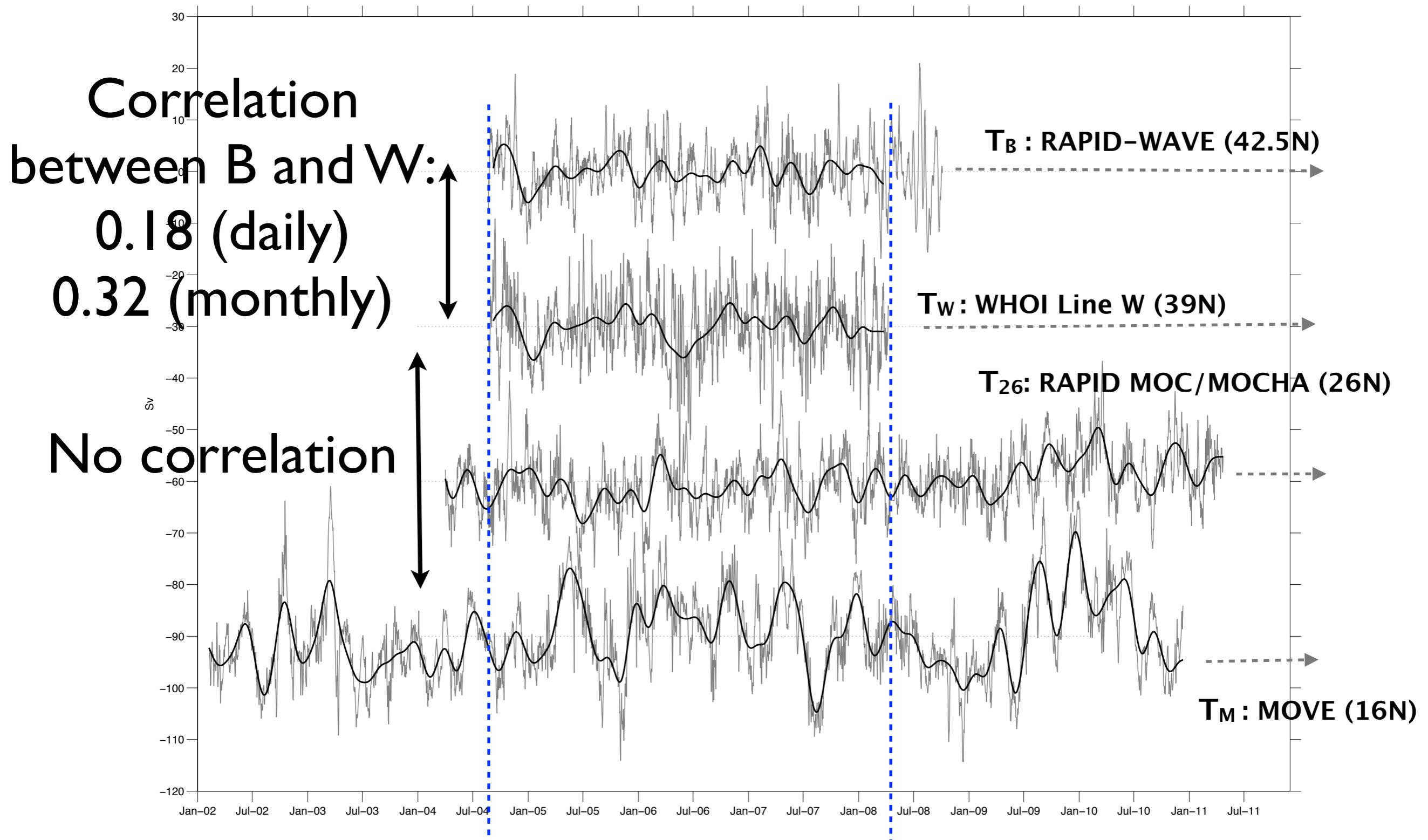
Correlation



~ 6 months

In “Ocean bottom pressure data capture the North Atlantic Meridional
Overturning Circulation and its meridional coherence” Elipot et al. 2013 (in revision for JPO)

So, 4 time series of overturning transport below 1000 m ...



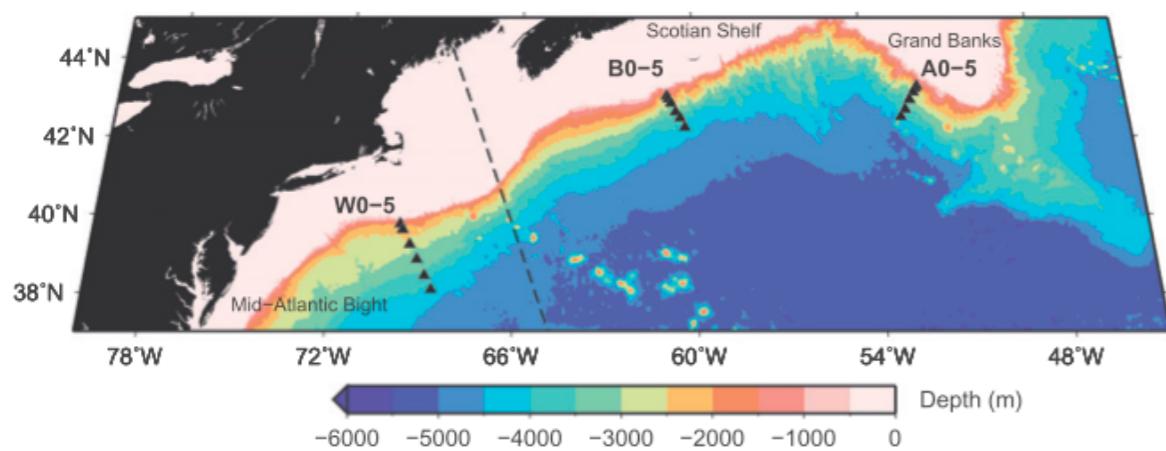
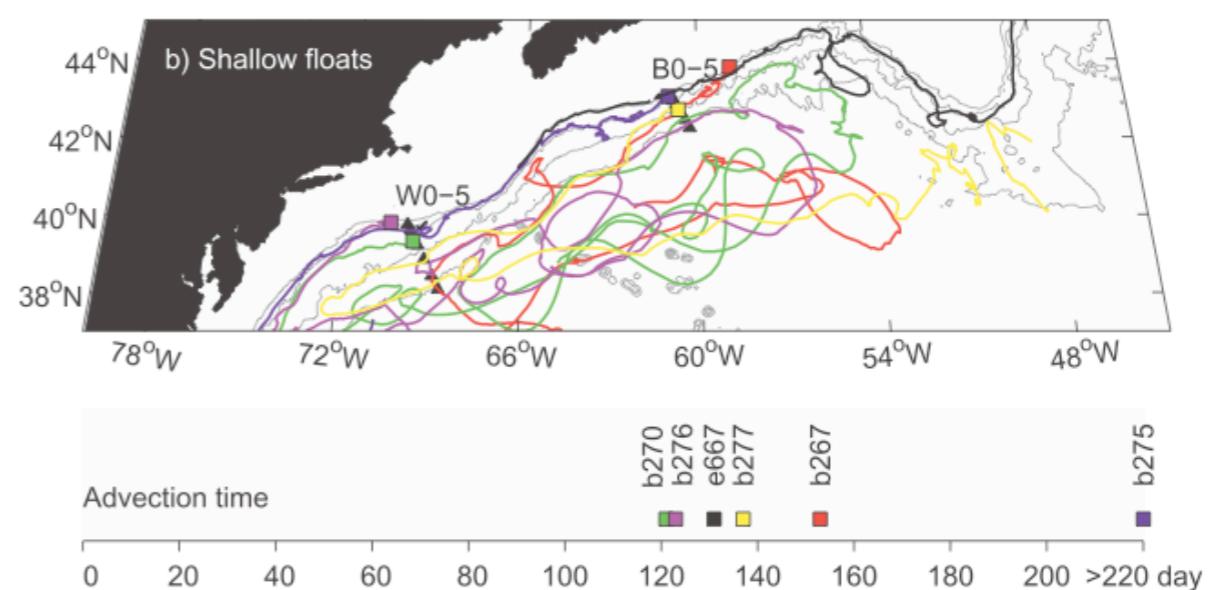
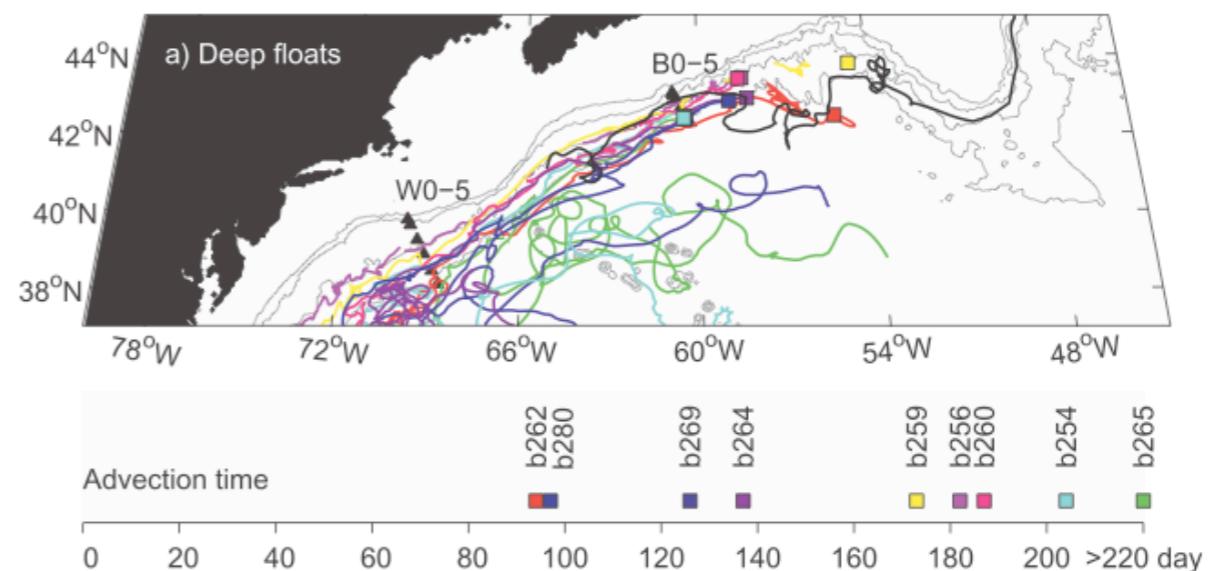
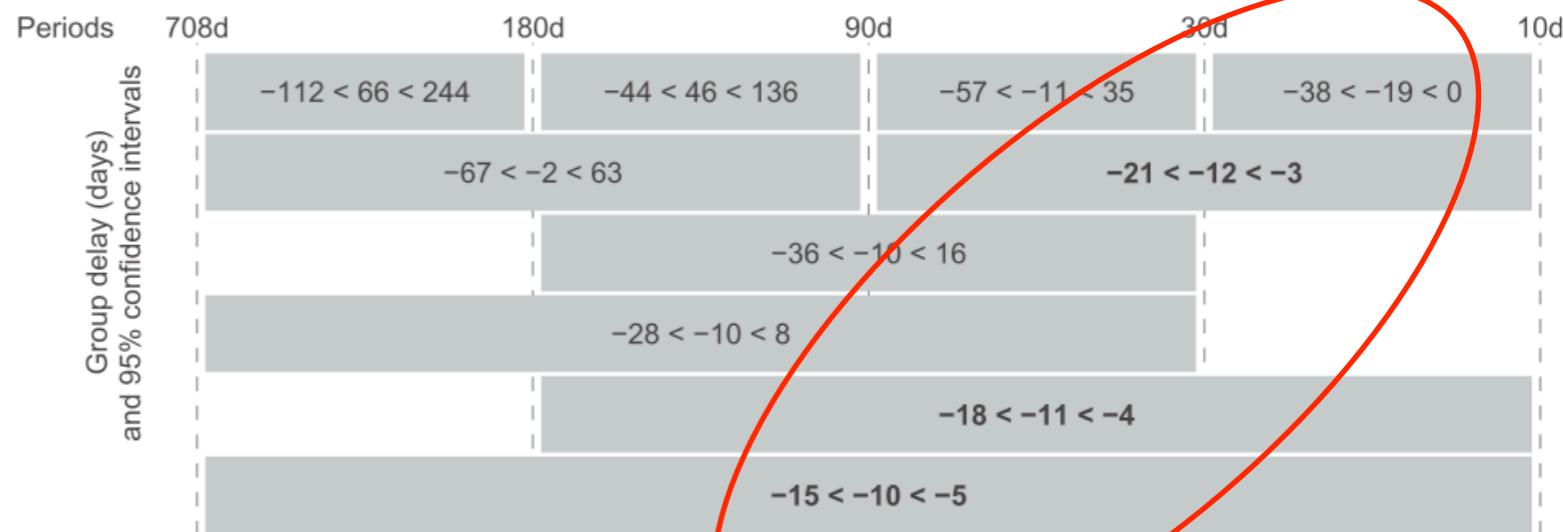
Study from 22 August 2004 to 8 April 2008 : 3.7 years

Evidence for boundary waves
propagation at time scales < 3 months
between Lines B and W at $\sim 1 \text{ m s}^{-1}$

Outstanding question:

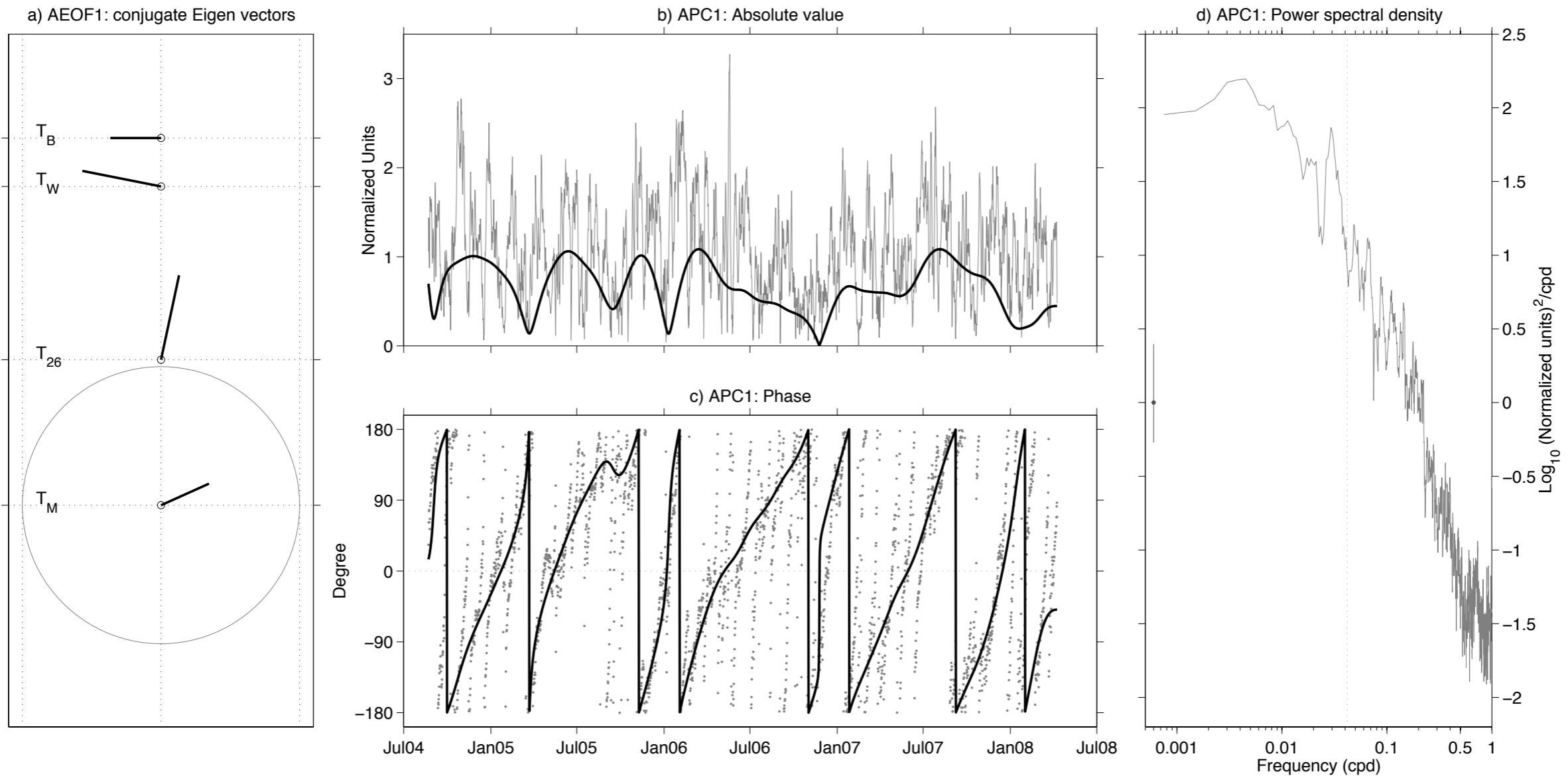
How does one explain overturning
coherence at time scales longer than 3
months?

From RAFOS floats: NOT advection by
DWBC



Analytic (complex) EOF analysis

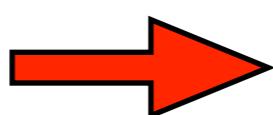
First mode: AEOF1: 36% of the covariance



Eigen
vectors

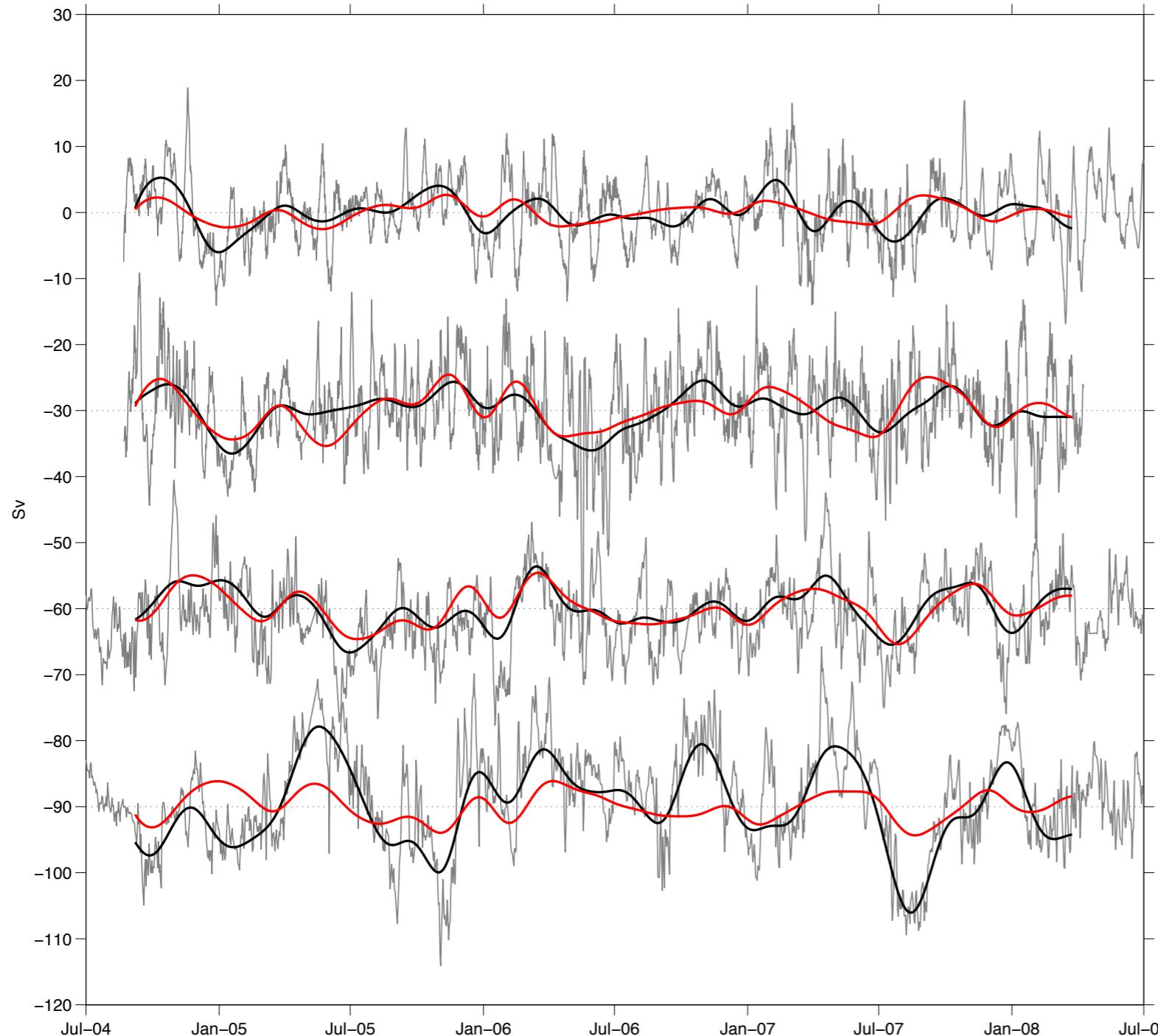
Principal Component
time series

“Red” spectrum
for this mode



dubious wave interpretation

Analytic EOF1: projection on time series



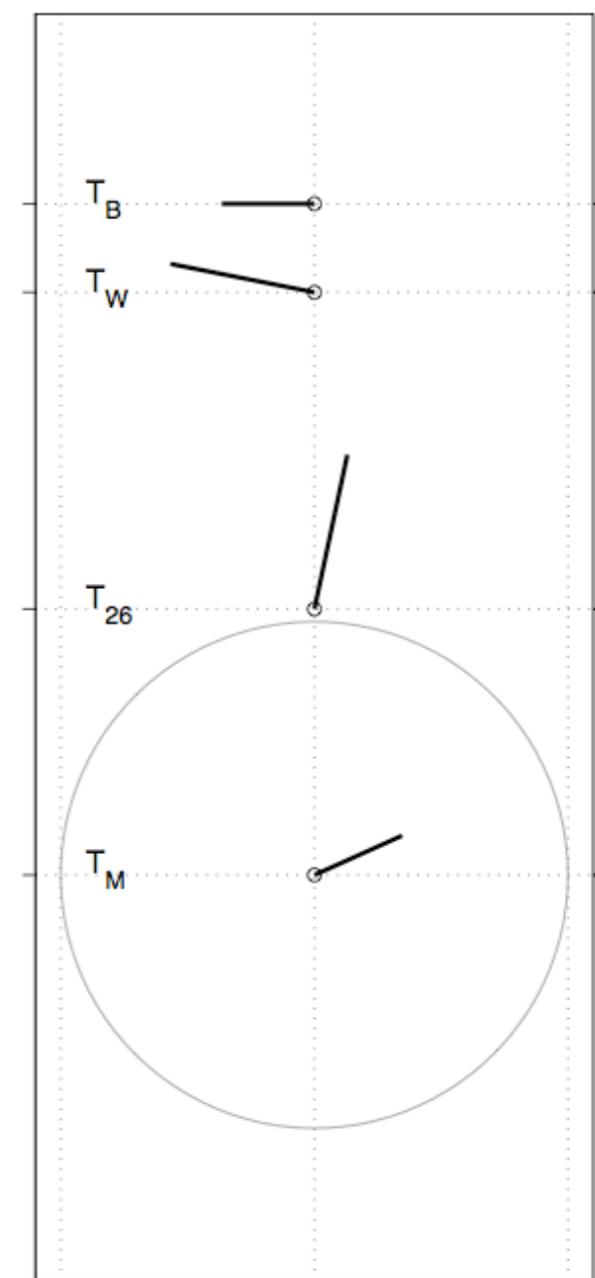
— 3-month lowpass

— 3-month lowpass AEOF1

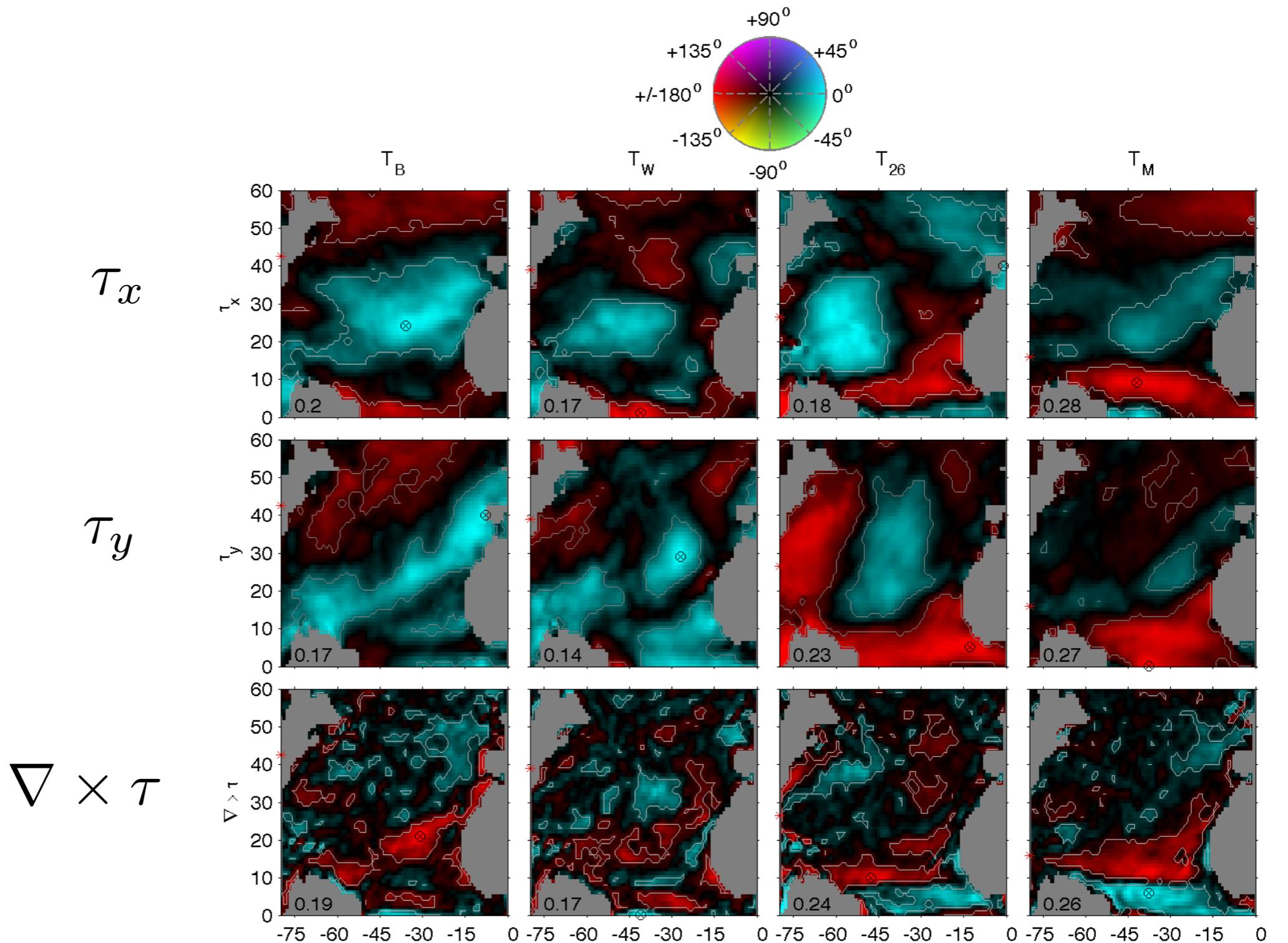
Variance explained (12h-step):

$AEOF1(\%)$	T_B	T_W	T_{26}	T_M
	19	48	56	21

a) AEOF1: conjugate Eigen vectors

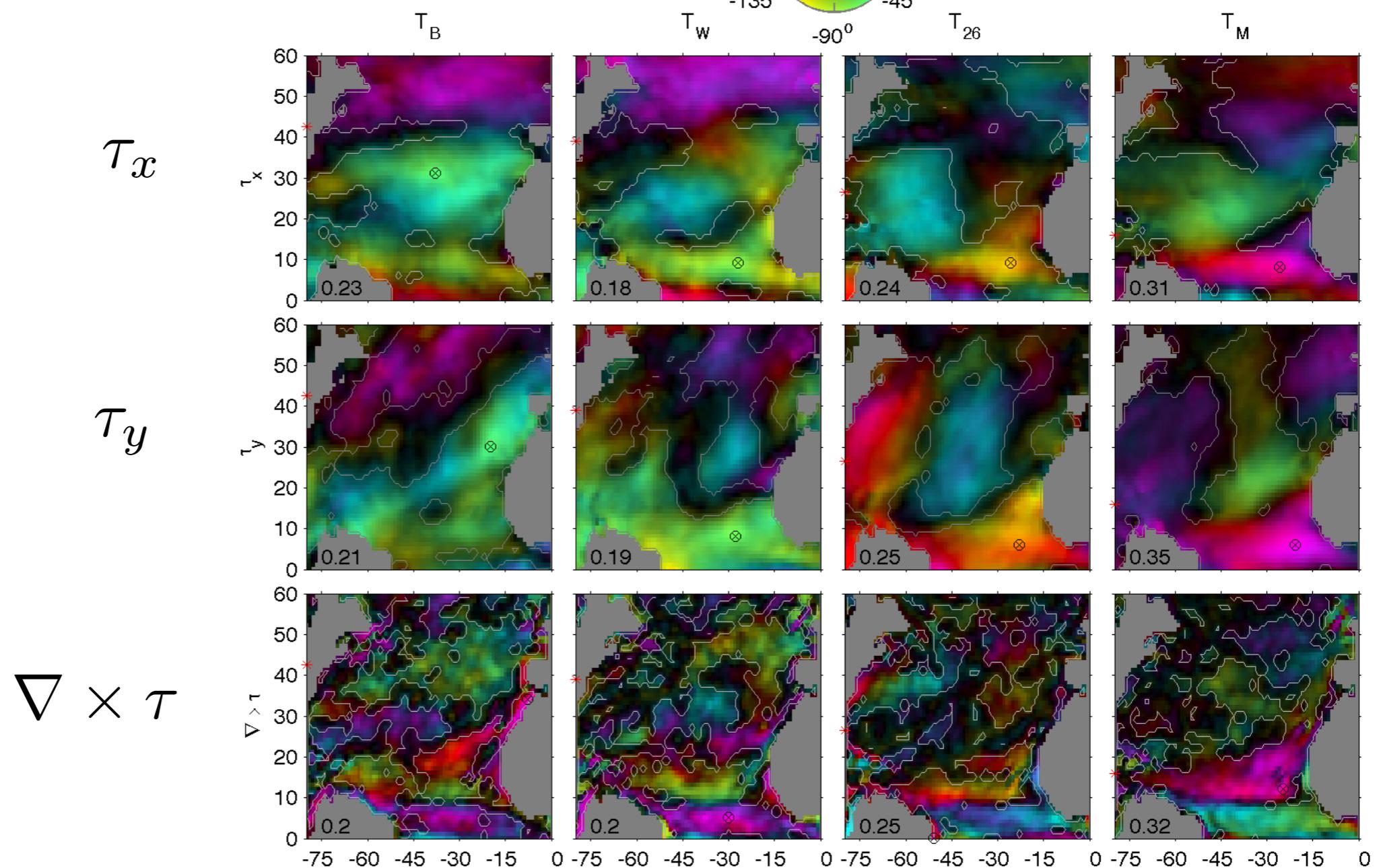
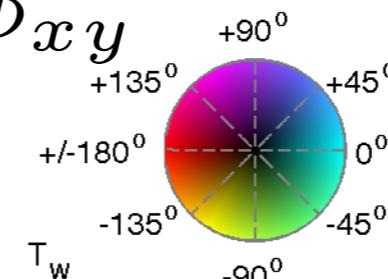


Correlation of transports with wind stress Cartesian components and wind stress curl:



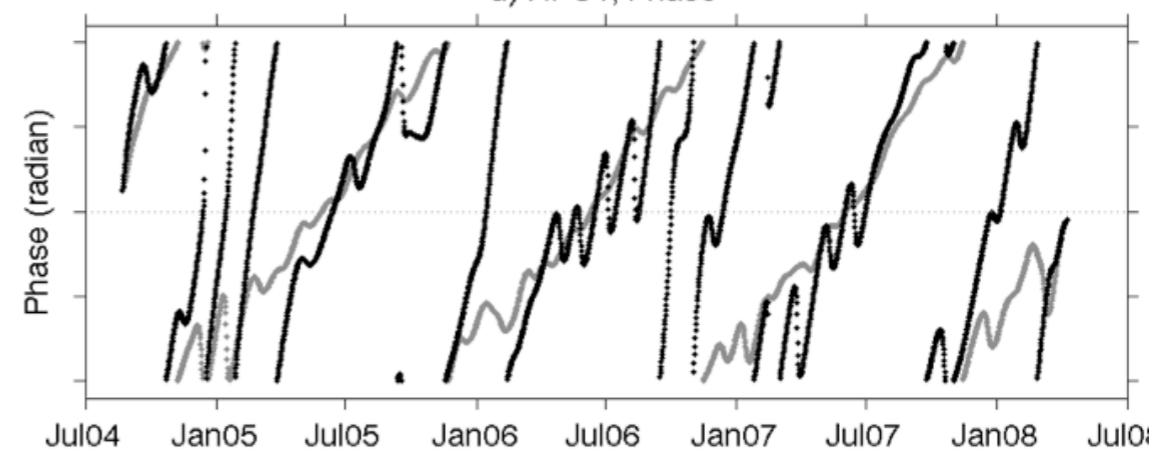
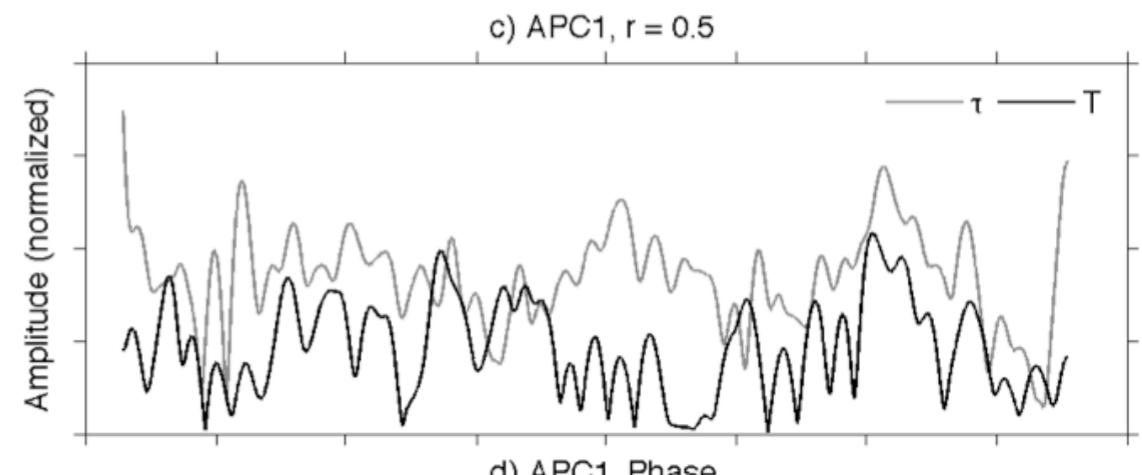
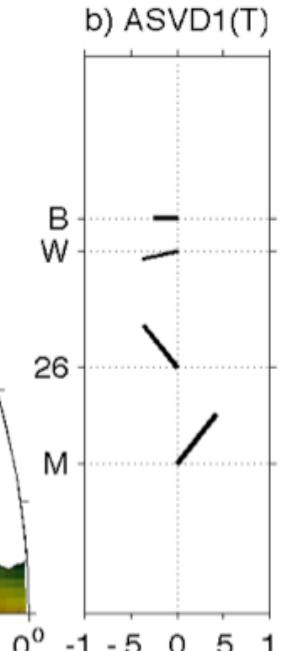
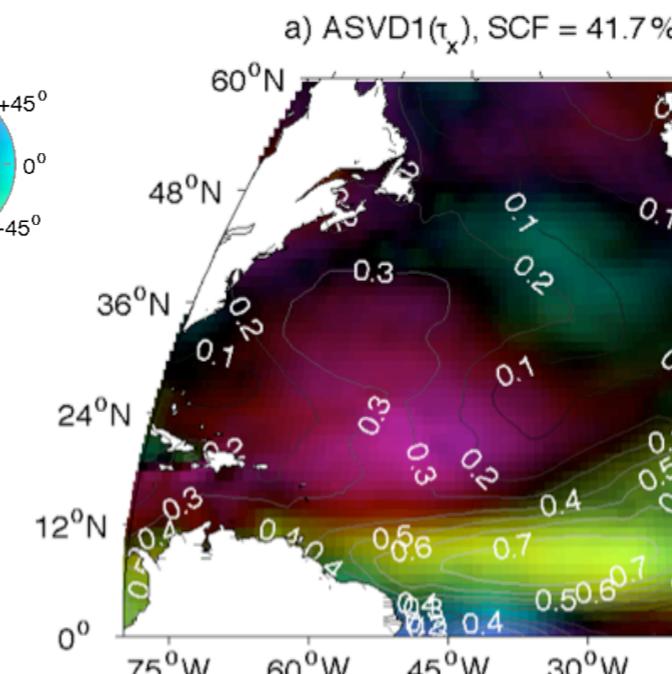
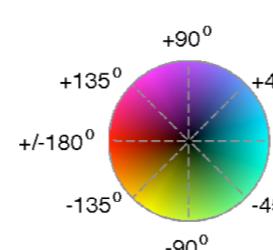
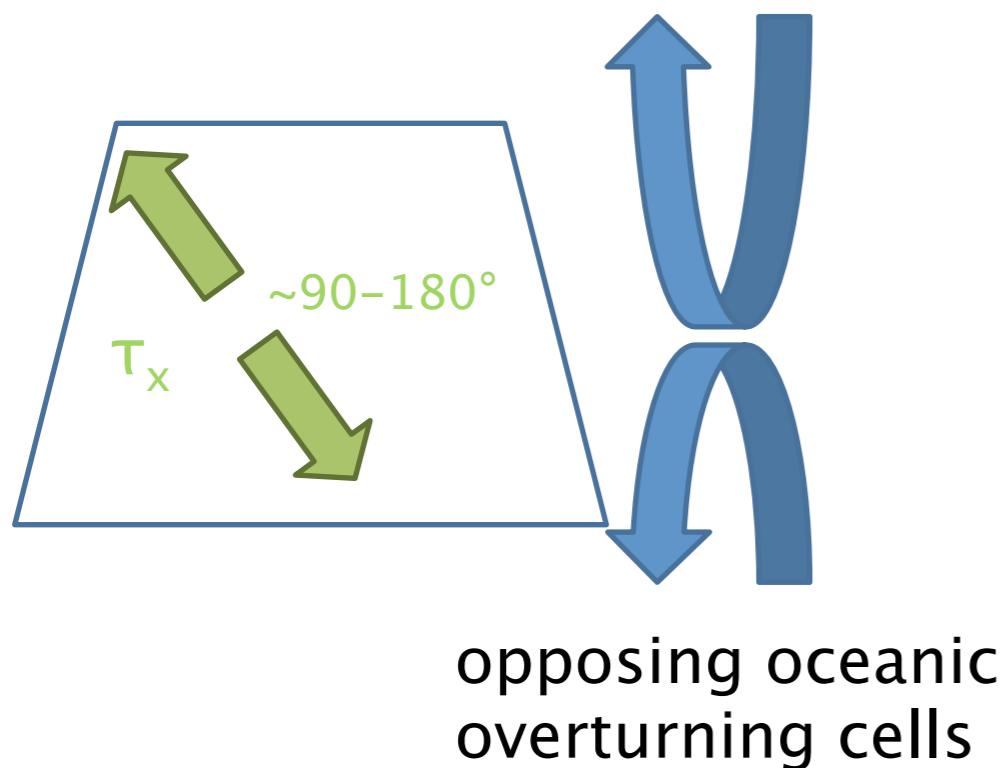
Analytic correlation of transports with wind stress Cartesian components and wind stress curl:

$$r_{x+y+}(0) = |r_{x+y+}(0)| e^{i\phi_{xy}}$$



First mode of correlation between transports and wind stress: ASVD1

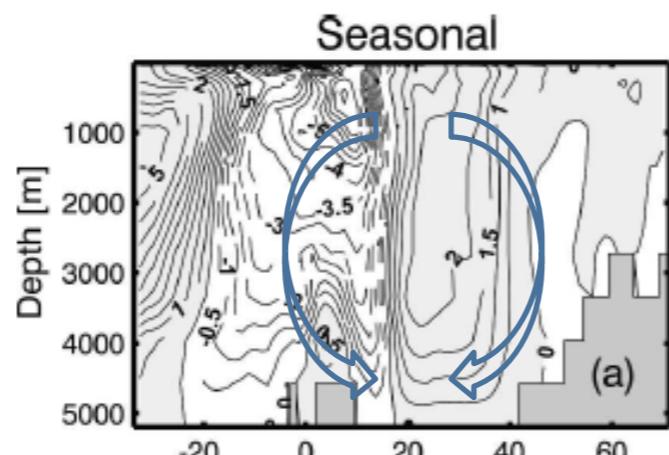
Singular Vectors



Principal Component time series: →
APC1 (τ) APC1 (T)

Coupling coefficient: $r = 0.50$

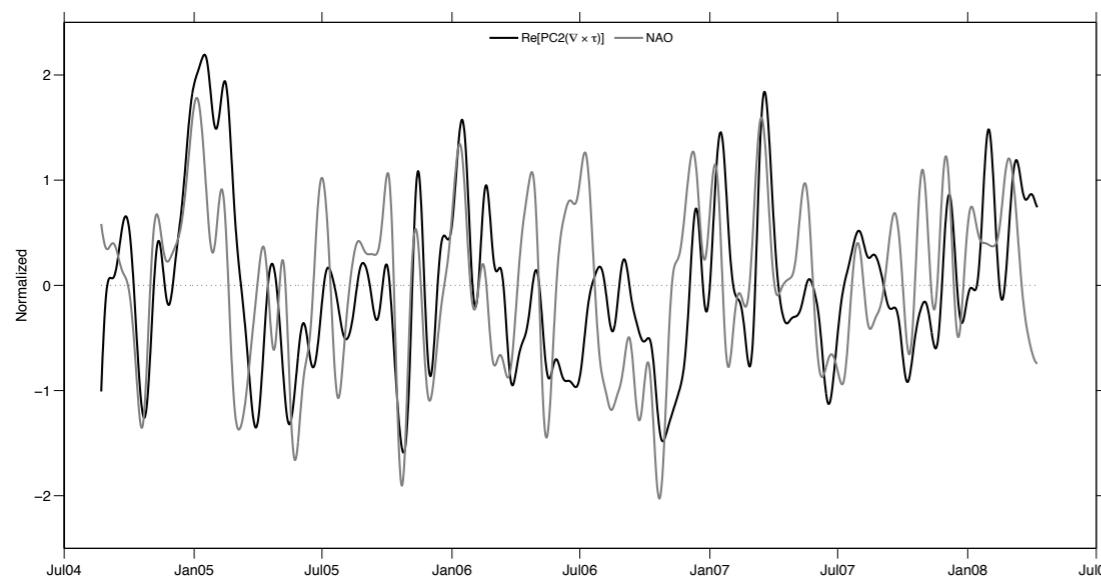
Sime et al. (2006):
Mean seasonal
anomaly
in MOC diagnosed
in HadCM3



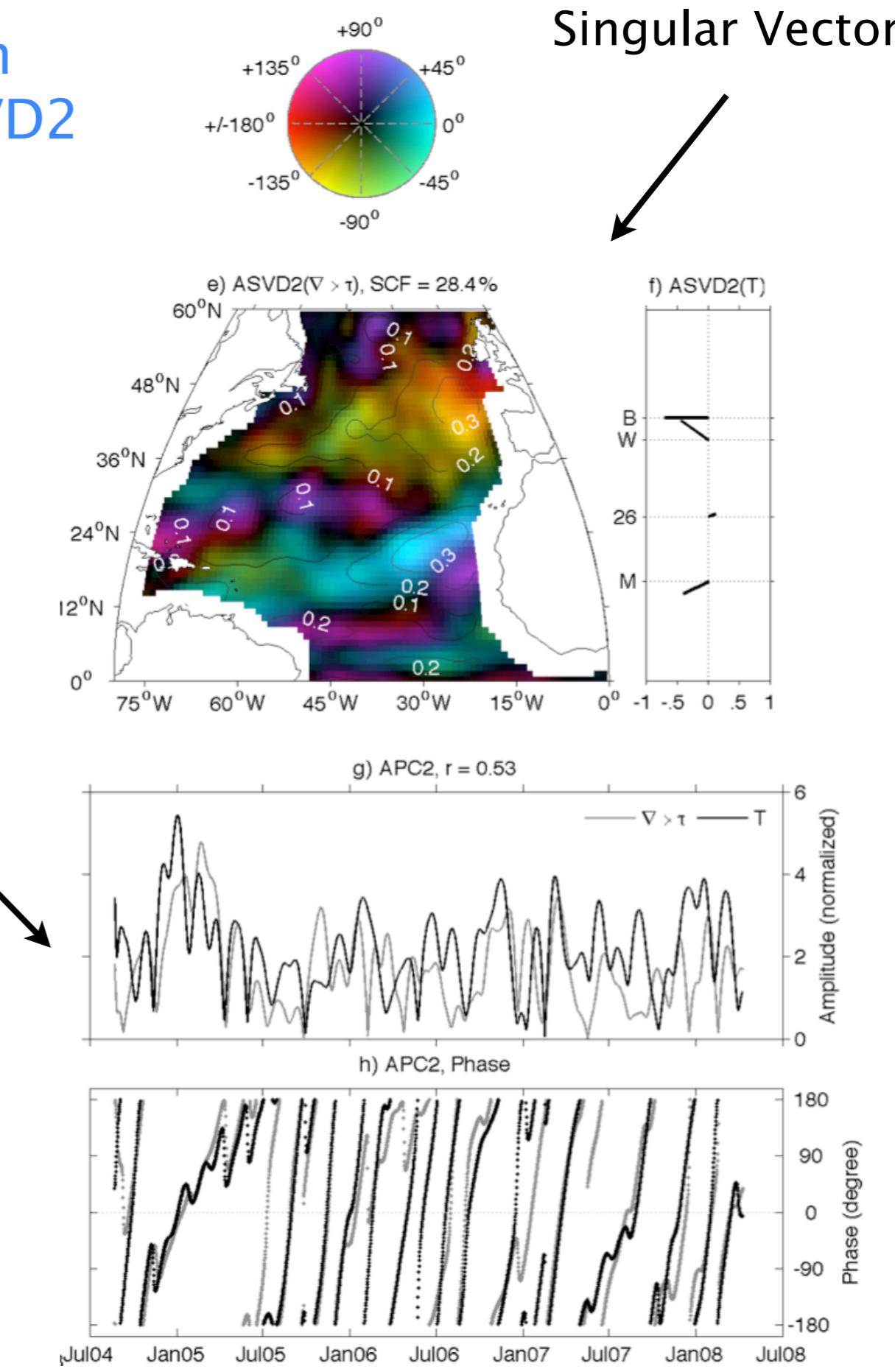
Second mode of correlation between transports and wind stress curl: ASVD2

Principal Component time series:
 APC1 ($\text{curl } \tau$) APC1 (T)

Coupling coefficient: $r = 0.53$

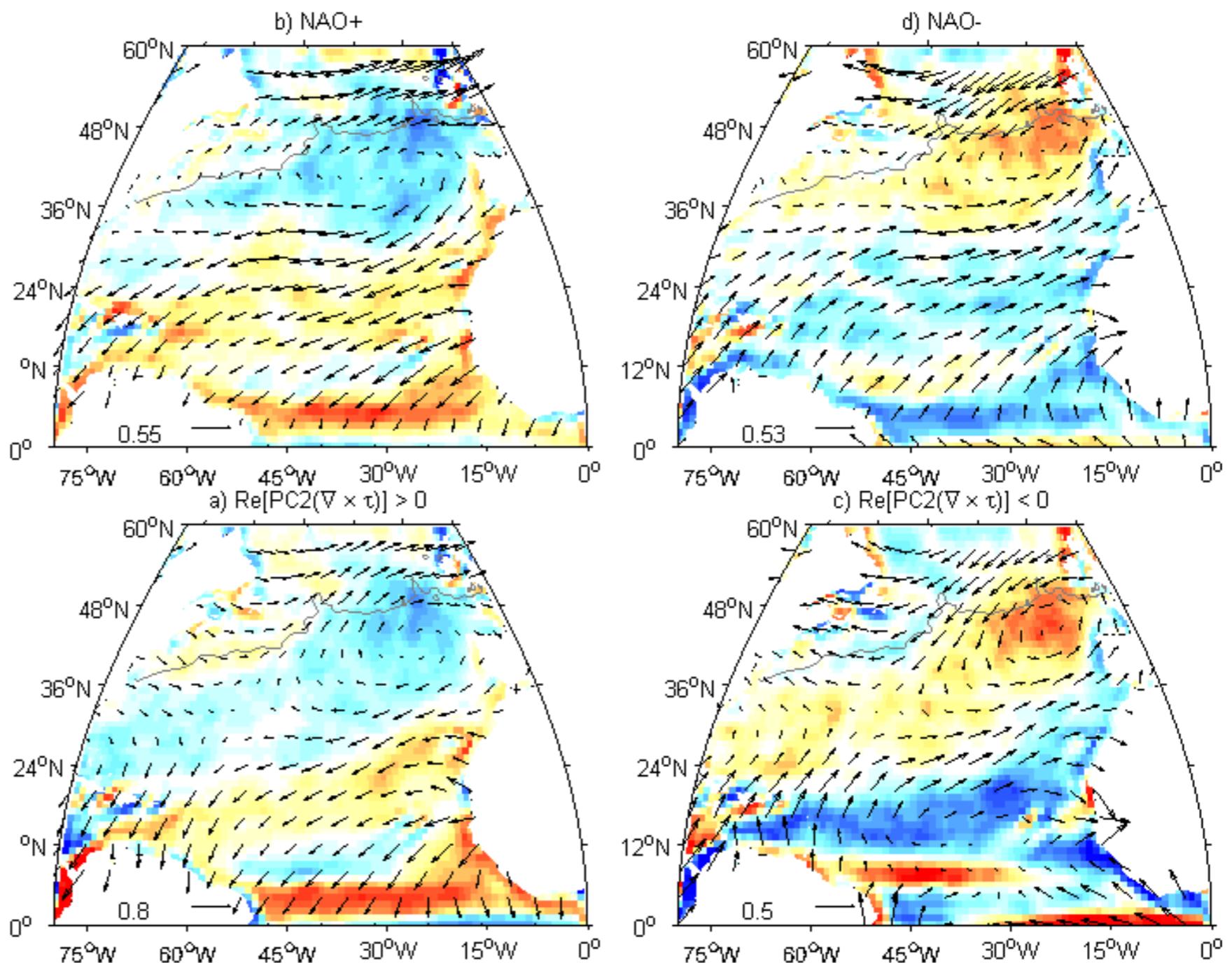


Correlation of mode
 with NAO index: **0.38**



Wind stress and curl changes associated with ASVD2:

NAO
phases

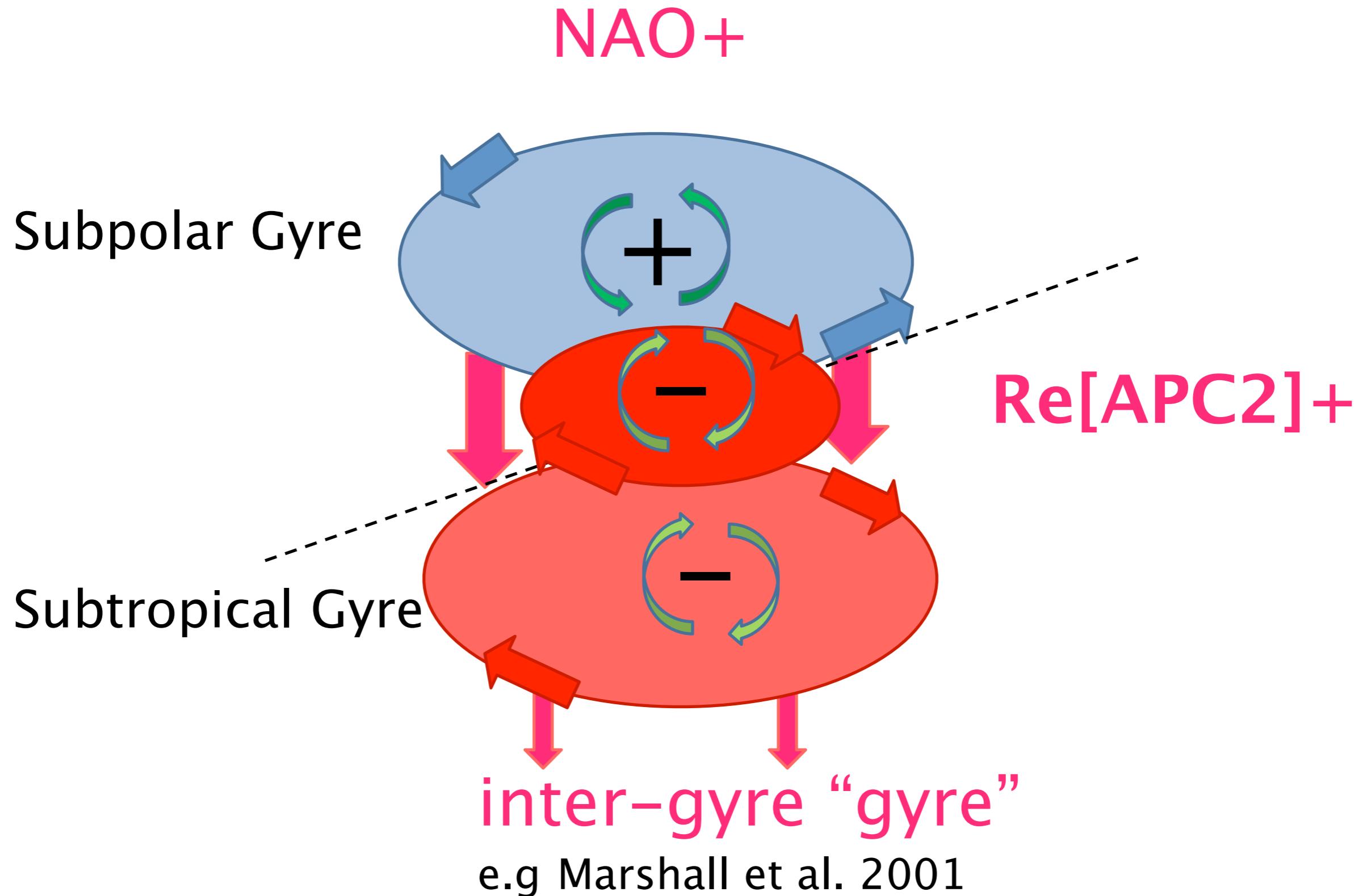


Positive or negative
 $\text{Re}[\text{APC}(\text{curl stress})]$

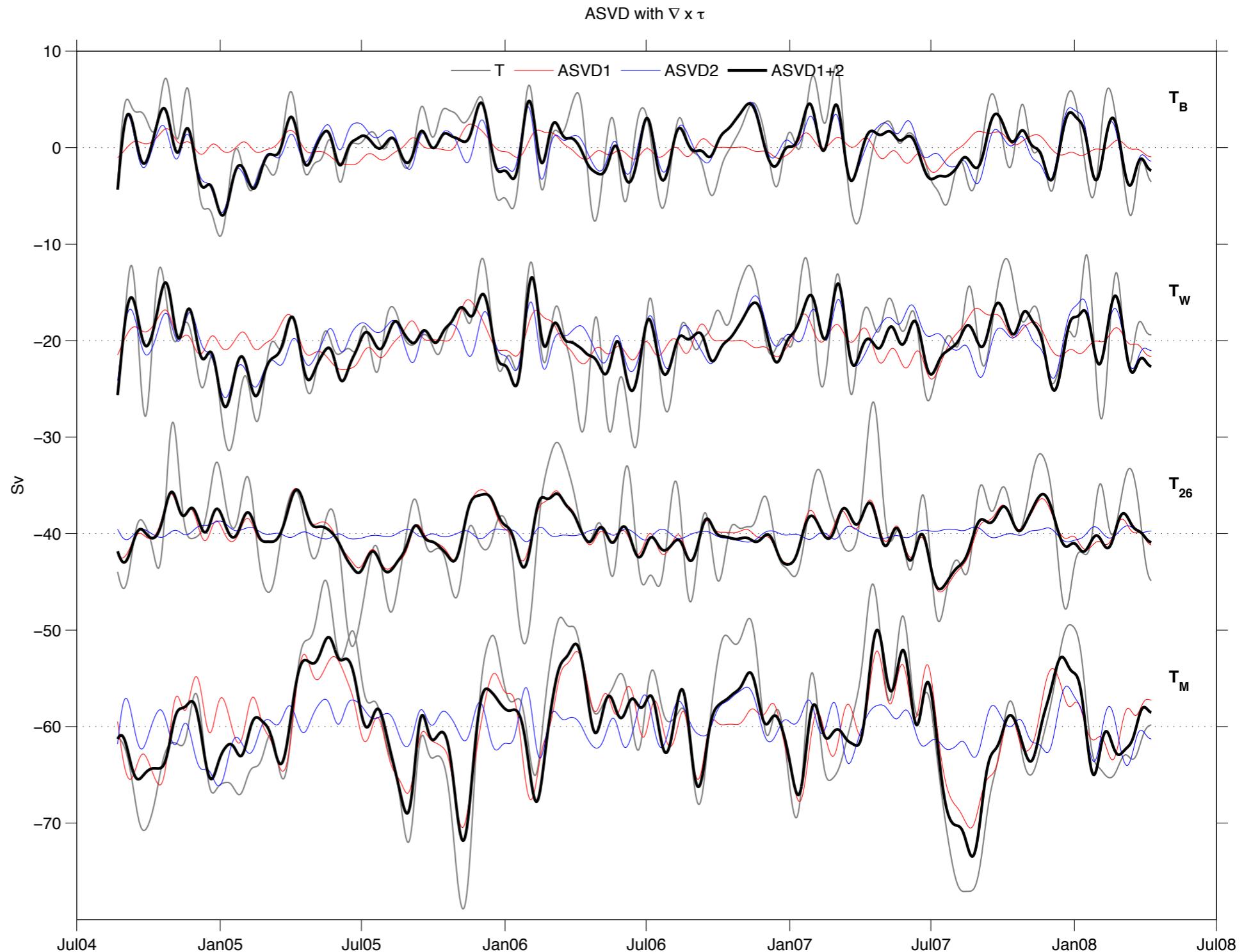


Composite anomalies of normalized wind stress curl (shading) and normalized wind stress vector (arrows)

Mechanism for deep overturning transport fast response to ASVD2?



Summary of wind covariance analysis: ASVD1 & ASVD2 for overturning transport time series



Total amount of variance explained
by AEOF and wind stress curl modes:

	T_B	T_W	T_{26}	T_M
AEOF1 (%)	19.1	47.9	55.7	20.6
$\nabla \times \tau$				
SVD1	9.4	20.7	45.7	59.4
SVD2	56.6	39.2	6.1	14.2
SVD1+SVD2	63.4	54.5	49.6	80.0

Summary

- 4 time series of western overturning transport below and relative to 1000 m at 42°N, 39°N, 26°N, and 16°N
- Time series are representative of geostrophic overturning processes on semi-annual and longer time scales
- Evidenced semi-annual and longer periods covariability between transports
- 1st mode of variability with near annual & semi-annual phase cycle associated with basin-wide Ekman forcing
- 2nd mode of variability related to large-scale NAO-like wind pattern

Thank you

selipot@rsmas.miami.edu