

White Paper “Dynamics of Decadal Climate Variability and Implications for its Prediction”

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Abstract

The temperature record of the last 150 years is characterized by a long-term warming trend, with strong multidecadal variability superimposed. The multidecadal variability is also seen in other (societal important) parameters such as Sahel rainfall or Atlantic hurricane activity. The existence of the multidecadal variability makes climate change detection a challenge, since Global Warming evolves on a similar timescale. The ongoing discussion about a potential anthropogenic signal in the Atlantic hurricane activity is an example. A lot of work was devoted during the last years to understand the dynamics of the multidecadal variability, and external as well as internal mechanisms were proposed. This White Paper focuses on the internal mechanisms relevant to the Atlantic Multidecadal Oscillation/Variability (AMO/V) and the Pacific Decadal Oscillation/Variability (PDO/V). Specific attention is given to the role of the Meridional Overturning Circulation (MOC) in the Atlantic. The implications for decadal predictability and prediction are discussed.

1. Introduction

Climate variability can be either generated internally by interactions within or between the individual climate subcomponents (e.g., atmosphere, ocean, and sea ice) or externally by e.g. volcanic eruptions, variations in the solar insolation at the top of the atmosphere, or changed atmospheric greenhouse gas concentrations in response to anthropogenic emissions. Examples of internal variations are the North Atlantic Oscillation (NAO), the El Niño/Southern Oscillation (ENSO), the Pacific Decadal Variability (PDV), or the Atlantic Multidecadal Variability (AMV). We present in Fig.1 indices derived from observations to visualize the decadal variability in selected variables and regions. The North Atlantic Sector is one region of strong multidecadal variability, and multidecadal variability can be readily seen, for instance, in European surface air temperature (SAT), Sahel rainfall, and Atlantic Hurricane activity. All three indices are highly correlated with the fluctuations in North Atlantic sea surface temperature (SST) on decadal timescales. Similar behaviour is seen in the North Pacific; the correspondence between SAT in the Southwest United States, the region most strongly affected by the PDV, and North Pacific SST (as given by the PDO index), however, is obvious but less significant.

Uncertainty in climate change projections for the 21st century arises from three distinct sources: internal variability, model errors and scenario uncertainty. Using data from a suite of climate models Hawkins and Sutton (2009) separate and quantify these sources (Fig. 2). For lead times of the next few decades the dominant contributions are internal variability and model uncertainty. The importance of internal variability generally increases at shorter time and space scales. ENSO, for instance, is one of the major factors affecting inter-annual variability, even on a global scale. The analyses of Hawkins and Sutton (2009) suggest that for decadal timescales and regional spatial scales (~2000km), model uncertainty is of greater importance than internal variability. Both the contributions to prediction uncertainty from internal variability and, especially, model uncertainty are potentially reducible through

progress in climate research. Model biases are still rather large and model improvement will most likely lead to higher prediction skill, an experience made in numerical weather forecasting (NWP) and seasonal forecasting. Another lesson from NWP and seasonal forecasting is that predictions considerably benefit from better knowledge of the initial conditions and initialization methods.

In this White Paper, we focus on the internal decadal variability and its dynamics. A large body of literature exists on these dynamics and we try to summarize this knowledge in a concise way. Bjerknes (1964) concluded from his early analysis of the observations in the mid-latitude Atlantic region that the atmosphere drives the ocean at interannual time scales, while at the decadal to multidecadal time scales it is the ocean dynamics that matters. Many subsequent observational and modelling studies agree basically with this view (e.g., Delworth et al. 1993; Latif 1998; and references therein), so the predictability potential in the Atlantic Sector is probably large at the decadal to multidecadal timescales (Latif et al. 2006a).

Although the mechanisms behind the decadal to multidecadal variability in the Atlantic Sector are still controversial, there is some consensus that some of the longer-term multidecadal variability is driven by variations in the thermohaline circulation (THC). In the North Pacific, a strong overturning circulation does not exist, and variations in the wind-driven circulation are the most likely candidate for the generation of decadal to multidecadal variability (e.g., Latif 1998, Schneider and Cornuelle 2005, Latif 2006). Rossby wave propagation is important in this context.

Here we concentrate on the North Atlantic and the North Pacific Sectors, two regions of considerable potential decadal predictability (Fig. 3). We note, however, that other regions such as the Southern Ocean also exhibit relatively high predictability potentials. In Section 2, we discuss some aspects of potential decadal predictability. Section 3 provides a conceptual description of the different possible mechanisms that can lead to decadal variability. Again, we consider only variability that is driven internally and involves the interaction between the

ocean-sea ice system and the atmosphere. This type of variability must be treated within a stochastic framework. We do not consider chaotic variability originating in either only the atmosphere or the ocean-sea ice system. The main conclusions are drawn in Section 4.

2. Decadal predictability

Climate prediction has been mostly considered on two different time scales: seasonal and centennial. Seasonal prediction is primarily an initial value problem, i.e. the evolution of the system depends on the initial state. Whereas centennial-scale prediction is normally considered a boundary value problem (e.g., Palmer et al. 2004), i.e. the evolution of climate depends on external changes in radiative forcing, such as anthropogenic changes in atmospheric composition or solar forcing (IPCC 2007). What class of problem is decadal prediction: initial value or boundary value? As suggested by observations and models decadal climate variations - global and regional - may arise from internal modes of the climate system and be potentially predictable (i.e. an initial value problem). On the other hand, projections of future climate indicate a rise in global mean temperature of between 2 and 4°C by 2100, dependant on emission scenario and model. This translates to an average rise in global mean temperature of order 0.3°C per decade. This is large compared, for instance, with the observed increase of about 0.7°C during the last century, and argues that decadal prediction is also a boundary value problem. Thus the prediction of the climate over the next few decades poses a joint initial/boundary value problem.

While the predictability of internal fluctuations on seasonal timescales has been intensively studied for more than twenty years, decadal predictability has been systematically investigated for only a few years. One reason is the much longer timescale, which requires rather long model integrations and which is therefore closely related to the availability of large computer resources. One distinguishes between potential (diagnostic) and classical (prognostic) predictability studies. Potential predictability (e.g., Boer 2000 and 2004) attempts

to quantify the fraction of long-term variability that may be distinguished from the internally generated natural variability, which is not predictable on long time scales and considered as “noise”. The long-term variability “signal” that rises above this noise is deemed to arise from processes operating in the physical system that are assumed to be, at least potentially, predictable. Decadal potential predictability is simply defined as the ratio of the variance on the decadal time scales to the total variance. Classical predictability studies consist of performing ensemble experiments with a single coupled model perturbing the initial conditions (Griffies and Bryan 1997a, b; Grötzner et al. 1999; Boer 2000; Collins 2002; Collins and Sinha 2003; Pohlmann et al. 2004). The predictability of a variable is given by the ratio of the actual signal variance to the ensemble variance. This method provides in most cases an upper limit of predictability since they assume a perfect model and, very often, near-perfect initial conditions. A third method compares the variability simulated with and without active ocean dynamics. Those regions in which ocean dynamics are important in generating the decadal variability are believed to be regions of high decadal predictability potential (Park and Latif 2005).

All three types of studies yield similar results for decadal predictability (Latif et al. 2006a). In contrast to seasonal to interannual predictability potential decadal predictability is found predominately over the mid to high-latitude oceans (Boer 2004). The potential decadal predictability decreases with increasing timescale but appreciable values exist up to multidecadal timescales, especially for the North Atlantic and the Southern Ocean (Fig. 3). In the North Pacific, the decadal predictability potential is considerably smaller, but probably still useful. It should be mentioned that these results hold only for the internal variability. Results obtained by including externally driven variability such as that related to an increase in atmospheric greenhouse gas concentrations yield rather different results (Hawkins and Sutton 2009). This can be easily understood by considering, for instance, the North Atlantic. MOC- related decadal variations are strong in this region and appear to be predictable. In

contrast, the expected anthropogenic weakening of the MOC may not be well detectable for many decades due to the existence of the strong internal variability. So, predictability will critically depend on the lead time. On short lead times of a decade, the internal variability may dominate. On long lead times of a century, the weakening in response to changing boundary conditions may prevail.

3. Mechanisms for decadal variability

The climate system displays variability over a broad range of timescales, from monthly to millennial and to even longer timescales. It is impossible to describe the full range of climate variability deterministically with one model, since the governing equations are too complicated so that an analytical solution is not known. A numerical solution of the complete equations is not feasible, because the necessary computer resources are not available, and will not be available over the next years. The climate system is comprised of components with very different internal timescales. Weather phenomena, for instance, have typical lifetimes of hours or days, while the deep ocean needs many centuries to adjust to changes in surface boundary conditions. Hasselmann (1976) introduced an approach to modelling the effect of the fast variables on the slow in analogy to Brownian motion. He suggested treating the former not as deterministic variables, but as *stochastic* variables, so that the slow variables evolve following dynamical equations with stochastic forcing. The chaotic components of the system often have well-defined statistical properties and these can be built into approximate stochastic representations of the high-frequency variability. The resulting models for the slow variables are referred to collectively as *stochastic climate models*, although the precise timescale considered slow may vary greatly from model to model. We describe in the following the different models which were suggested for the generation of the internally driven decadal variability.

2.1 The zero-order stochastic climate model

We consider in the following discussion the atmosphere as the fast and the ocean-sea ice system as the slow component. In the simplest case, the atmosphere is treated as a *white noise* process, i.e. the spectrum of the atmospheric forcing agents such as the air-sea heat flux is white, i.e. does not have a variation with frequency. Furthermore, we assume first a local model in which the atmospheric forcing at one location drives only changes in the ocean-sea ice system at this very point; and neither the atmosphere nor the ocean-ice system exhibit spatial coherence. The resulting spectrum of an ocean variable say sea surface temperature (SST) is *red*, which means that the power increases with timescale. To avoid a singularity at zero frequency a damping was introduced by Hasselmann 1976. Frankignoul and Hasselmann (1977) and Hall and Manabe (1997) have shown that such a local model can indeed explain observed SST and salinity variability in parts of the Mid-Latitudes, away from coasts and fronts, and that climate models do reproduce this behaviour. In regions where meso-scale eddies or advection is important, the simple stochastic models fail too. Nevertheless, the simple local model demonstrates that decadal variability in the ocean-sea ice system can be easily be generated by the integration of atmospheric weather noise. This model can be considered as a null hypothesis for the generation of decadal variability.

2.2 Stochastic models with mean advection

Atmospheric variability on timescales of a month or longer is dominated by a small number of large-scale spatial patterns, whose time evolution has a significant stochastic component. One prominent example is the North Atlantic Oscillation (NAO, Hurrell 1995), and we shall discuss the role of the NAO in driving variations in the thermohaline circulation (THC) below. One may expect the atmospheric patterns to play an important role in ocean-atmosphere interaction. Advection in the ocean-sea ice system plays an important role in this coupling. Lemke et al. (1980) have fitted a stochastic model incorporating horizontal transport to

observed polar sea ice variability. Frankignoul and Reynolds (1983) describe the use of a local stochastic model, including the effects of advection by the observed mean current, to predict the statistical characteristics of observed SST anomalies in the North Pacific on timescales of several months. They find that mean advection has only a small effect in general, although in regions of large currents, the advection effects were important at lower frequencies. Herterich and Hasselmann (1987) have fitted a more general nonlocal stochastic model, incorporating advection and diffusion, to observed SST anomalies over the same region.

A simple one-dimensional stochastic model of the interaction between spatially coherent atmospheric forcing patterns and an advective ocean was developed by Saravanan and McWilliams (1998). Their model equations are simple enough and allow analytical treatment. A *slow–shallow* regime where local damping effects dominate and a *fast–deep* regime where nonlocal advective effects dominate were found. An interesting feature of the fast–deep regime is that the ocean–atmosphere system shows preferred timescales, although there is no underlying oscillatory mechanism, neither in the ocean nor in the atmosphere. The existence of the preferred timescale in the ocean does not depend on the existence of an atmospheric response to SST anomalies. It is determined by the advective velocity scale associated with the upper ocean and the length scale associated with low-frequency atmospheric variability. This mechanism is often referred to as “spatial resonance” or “optimal forcing”. For the extra-tropical North Atlantic basin, this timescale would be of the order of a decade. Interestingly, Deser and Blackmon (1993), Sutton and Allen (1997), and Alvarez et al. (2008) find such a decadal timescale in surface observations of the North Atlantic. However, the studies differ in the derived propagation characteristics. The stochastic-advective mechanism may also underlie the Antarctic Circumpolar Wave (ACW, e. g. White and Peterson 1996), as shown in the model study of Weisse et al. (1999) who drove an ocean-sea ice general circulation model

by spatially coherent but temporally white forcing. The analyzed model experiment is also described below, when we discuss the variability of the THC.

2.3 Stochastic wind stress forcing of a dynamical ocean

We have considered so far no varying ocean dynamics and only thermohaline forcing, i.e. heat and freshwater forcing. Frankignoul et al. (1997) used a simple linear model to estimate the dynamical response of the extra-tropical ocean to stochastic wind stress forcing with a white frequency spectrum. The barotropic fields are governed by a time-dependent Sverdrup balance, the baroclinic ones by the long Rossby wave equation. At each frequency, the baroclinic response consists of a forced response plus a Rossby wave generated at the eastern boundary. For zonally independent forcing, the response propagates westward at twice the Rossby phase speed. The model predicts the shape and level of the frequency spectra of the oceanic pressure field and their variation with longitude and latitude. The baroclinic response is spread over a continuum of frequencies, with a dominant timescale determined by the time it takes a long baroclinic Rossby wave to propagate across the basin and thus increases with the basin width. The baroclinic predictions for a white wind stress curl spectrum are broadly consistent with the frequency spectrum of sea level changes and temperature fluctuations in the thermocline observed near Bermuda. Schneider et al. (2002) found some evidence for the accumulation of stochastic atmospheric forcing along Rossby wave trajectories in the North Pacific. Stochastic wind stress forcing may thus explain a substantial part of the decadal variability of the oceanic gyres, especially in the North Pacific. The importance of stochastically driven Rossby waves was also described in Latif (2006) who studied the multidecadal variability in the North Pacific in a coupled ocean-atmosphere general circulation model. His results suggest that the multidecadal variability can be partly explained by the dynamical ocean response to stochastic wind stress forcing. Superimposed on the red background variability, a multidecadal mode with a period of about 40 yr is simulated which

can be understood through the concept of spatial resonance between the ocean and the atmosphere, as described above. In summary, the bulk of the potential predictability found in the North Pacific (e.g., Boer 2004) is probably related to the propagation of long baroclinic Rossby waves (e.g., Schneider and Cornuelle 2005). The degree of air-sea coupling, however, needs to be considered in this context (e.g., Latif and Barnett 1994), as well as the role of remote forcing by the Tropics (e.g., Gu and Philander 1997, Jacobs et al. 1994).

2.4 Hyper modes

We describe now a case in which the atmosphere is no longer represented by a simple stochastic model but deterministically by an atmospheric general circulation model. The ocean is represented by a vertical column model that only includes vertical diffusion. Such a coupled model (AGCM-ML) was studied, for instance, by Dommenges and Latif (2008) and displays many features described above. Since varying ocean dynamics are not considered, air-sea interactions are purely local. It is shown that some important aspects of the space-time structure of multidecadal sea surface temperature (SST) variability can be explained by such a model. Dommenges and Latif (2008) formulate the concept of “Global Hyper Climate Modes”, in which surface heat flux variability associated with regional atmospheric variability patterns is integrated by the large heat capacity of the extra-tropical oceans, leading to a continuous increase of SST variance towards longer timescales. Atmospheric teleconnections spread the extra-tropical signal to the tropical regions. Once SST anomalies have developed in the Tropics, global atmospheric teleconnections spread the signal around the world creating a global “hyper” climate mode. Estimates with a further reduced model suggest that hyper climate modes can vary on timescales longer than 1,000 years. The SST anomaly patterns simulated at multidecadal timescales in the model of Dommenges and Latif (2008) are remarkably similar to those derived from observations and from long control integrations with sophisticated coupled ocean-atmosphere general circulation models (Fig. 4). The hyper mode

mechanism could, for instance, underlie the Pacific Decadal Oscillation, whose structure is reasonably well reproduced in the Dommenget and Latif (2008) model at multidecadal timescales. Ocean dynamics and large-scale ocean-atmosphere coupling may amplify the hyper modes, especially in the Tropics, and may influence the regional expression of the associated variability. Equatorial ocean dynamics such as those operating in ENSO, for instance, would enhance the variability in the eastern and central Equatorial Pacific. Such feedbacks would make the model certainly more realistic, but are not at the heart of the mechanism which produces the hyper modes.

2.5 Stochastically driven variability of the thermohaline circulation

We now discuss the decadal variability of the thermohaline circulation (THC). Competing mechanisms were proposed for the North Atlantic Meridional Overturning Circulation (MOC) variability. One idea is that multidecadal MOC variability is driven by the low-frequency portion of the spectrum of atmospheric flux forcing. Mikolajewicz and Maier-Reimer (1990) describe results from a multi-millennial integration with the Hamburg Large-Scale Geostrophic (LSG) Ocean General Circulation Model which was driven by spatially correlated white-noise freshwater flux anomalies. In addition to the expected red-noise character of the oceanic response, the model simulated pronounced variability in a frequency band around 320 years in the Atlantic basin. This is due to the excitation of a damped oceanic eigenmode by the stochastic freshwater flux forcing. The physics behind the variability involve a dipole-like salinity anomaly advected by and interacting with the mean THC. Weisse et al. (1994) describe decadal variability with a timescale of the order of 10 to 40 years in the North Atlantic in the same experiment. It describes the generation of salinity anomalies in the Labrador Sea and the following discharge into the North Atlantic. The generation of the salinity anomalies is mainly due to an almost undisturbed local integration of the white noise freshwater fluxes. The timescale and damping term of the integration

process are determined by the flushing time of the well-mixed upper layer. The decadal mode affects the MOC and represents a “discharge process” that depends nonlinearly on the modulated circulation structure rather than a regular linear oscillator.

Delworth and Greatbatch (2000), investigating the multidecadal variability in the coupled model simulation of Delworth et al. (1993) also describe an internal ocean mode in their analysis of a series of coupled and uncoupled ocean model integrations. The multidecadal variability simulated in the model discussed in Delworth et al. (1993) involves interactions of the gyre and thermohaline circulations, in which the anomalous salt advection into the sinking region plays a crucial role in determining deep convection. Delworth and Greatbatch (2000) show that the multidecadal MOC fluctuations are driven by a spatial pattern of surface heat flux variations that bear a strong resemblance to the NAO. No conclusive evidence is found that the MOC variability is part of a dynamically coupled atmosphere-ocean mode in this particular model. Griffies and Tziperman (1995) interpreted the same coupled model simulation in terms of a stochastically forced four-box model of the MOC. The box model was placed in a linearly stable thermally dominant mean state under mixed boundary conditions (Stommel 1961). A linear stability analysis of this state reveals one damped oscillatory THC mode in addition to purely damped modes. Direct comparison of the variability in the box model and coupled ocean-atmosphere model reveals common qualitative aspects. Such a comparison supports the hypothesis that the coupled model’s MOC variability can be described by the stochastic excitation of a linear damped oscillatory THC mode.

Coupled air-sea modes were also proposed as mechanisms for decadal variability.

Timmermann et al. (1998) describe coupled variability with a 35-yr period in a multicentury integration of the ECHAM3/LSG climate model. An anomalously strong MOC drives positive SST anomalies in the North Atlantic. The atmospheric response to these involves a strengthened NAO, which leads to anomalously weak evaporation and Ekman transport off Newfoundland and in the Greenland Sea and the generation of negative sea surface salinity

(SSS) anomalies. Reduced deep convection in the oceanic sinking regions and a subsequently weakened MOC eventually lead to the formation of negative SST anomalies.

Eden and Greatbatch (2003) describe results from a simple stochastic atmosphere model coupled to a realistic model of the North Atlantic Ocean. A north-south SST dipole, with its zero line centred along the sub-polar front, drives the atmosphere model, which in turn forces the ocean model by NAO-related surface fluxes. The model simulates a damped decadal oscillation. It consists of a fast wind-driven, positive feedback of the ocean and a delayed negative feedback orchestrated by shallow overturning circulation anomalies. The positive feedback turns out to be necessary to distinguish the coupled oscillation from that in a model without any feedback from the ocean to the atmosphere. Anomalous geostrophic advection appears to be important in sustaining the oscillation.

The uncertainties in the atmospheric response to mid-latitude SST anomalies and thus the strength of air-sea coupling is still an open scientific question and explain some of the major differences between models. Recent studies suggest that the resolution (horizontal and vertical) of atmosphere models may not be sufficient to resolve phenomena such as variations in the position of the polar front (Minobe et al. 2008). Furthermore, stratospheric and tropospheric variability are linked on seasonal time scales, as shown by observational and modelling studies (see e.g., Keenlyside et al. 2008 and references therein). It follows that low-frequency stratospheric change, of either natural or anthropogenic origin, can influence tropospheric circulation. This was recently highlighted in experiments that showed that the observed strengthening of the stratospheric jet from 1965-1995 could reproduce the observed changes in the NAO and North Atlantic Sector climate (Scaife et al. 2005).

4. Conclusions

Decadal climate prediction is of socio-economic importance and has a potentially important role to play in policy making. In contrast to seasonal prediction and centennial climate

projections, it is a joint initial value/boundary value problem. Thus, both accurate projections of changes in radiative forcing and initialisation of the climate state, particularly the ocean, are required. Although the first promising steps towards decadal prediction have been made, much more work is required. Most importantly, understanding of the mechanisms of decadal-to-multidecadal variability is lacking. The atmospheric response to mid-latitude SST anomalies, for instance, is still highly controversial and future research should treat this as a key topic.

The decadal predictability potential is rather large in the North Atlantic Sector. Although the mechanisms behind the decadal to multidecadal variability in the North Atlantic Sector are still controversial, there is some consensus that some of the longer-term multidecadal variability is driven by variations in the thermohaline circulation. Analyses of ocean observations and model simulations by Latif et al. (1996b) suggest that there have been considerable changes in the thermohaline circulation during the last century. These changes are likely to be the result of natural multidecadal climate variability and are driven by low-frequency variations of the North Atlantic Oscillation (NAO) through changes in Labrador Sea convection (Fig. 5a). The variations in the North Atlantic thermohaline circulation appear to be predictable one to two decades ahead, as shown by a number of perfect model predictability experiments. North Atlantic SST is strongly influenced by the MOC changes and the two quantities exhibit a clear lead-lag relationship in some models, as visualized in Fig. 5b. We expect that the next few decades will be strongly influenced by such multidecadal variations, although the effects of anthropogenic climate change are likely to introduce trends. Some impact of the variations of the thermohaline circulation on the atmosphere has been demonstrated in some studies, so that useful decadal predictions with economic benefit may be possible. However, unpredictable external forcing through strong volcanic eruptions and/or anomalous solar radiation may offset the internal variations and introduce an additional source of uncertainty.

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Figure Captions

Fig. 1: From top to bottom: Annual mean European SAT (5°W-10°E, 35-60°N), linearly de-trended annual mean European SAT, linearly de-trended North Atlantic SST (0-60°N), summer Sahel rainfall, Atlantic hurricane activity (ACE index), and North Atlantic SST repeated from above. The bottom three panels show northwestern United States SAT, the linearly de-trended version, and the PDO index. All timeseries are deviations from the long-term mean. All temperature timeseries are in units of [°C]. The Sahel rainfall and the ACE index were normalized with the long-term standard deviation.

Fig.2.: The relative importance of each source of uncertainty in decadal mean surface temperature projections is shown by the fractional uncertainty (the 90% confidence level divided by the mean prediction), for **(A)** global mean, relative to the warming from the 1971-2000 mean, and **(B)** British Isles mean, relative to the warming from the 1971-2000 mean. Internal variability grows in importance for the smaller region. Scenario uncertainty only becomes important at multi-decadal lead times. The dashed lines in **(A)** indicate reductions in internal variability, and hence total uncertainty, that may be possible through proper initialisation of the predictions through assimilation of ocean observations (Smith et al., 2007). The fraction of total variance in decadal mean surface air temperature predictions explained by the three components of total uncertainty is shown for, **(C)** a global mean, **(D)** a British Isles mean. Green regions represent scenario uncertainty, blue regions represent model uncertainty and orange regions represent the internal variability component. As the size of the region is reduced, the relative importance of internal variability increases. From Hawkins and Sutton (2009).

Fig. 3: Potential decadal predictability considering 10-year means obtained from the ensemble of coupled ocean-atmosphere models. See text for details. From Boer (2004).

Fig. 4: Correlation maps with the EOF-1 time series of the simple global climate model (AGCM-ML) at different timescales: a) timescales of 1-5 years, b) timescales of 5-41 years,

c) timescales longer than 41 years. Filtering was performed by applying and/or subtracting running means. The hyper mode is fully developed at multidecadal timescales. From Dommenget and Latif (2008).

Fig. 5: a) Time series of the winter [December–March (DJFM)] NAO index (shaded curve), a measure of the strength of the westerlies and heat fluxes over the North Atlantic and the Atlantic dipole SST anomaly index ($^{\circ}\text{C}$, black curve), and a measure of the strength of the MOC. The NAO index is smoothed with an 11-yr running mean; the dipole index is unsmoothed (thin line) and smoothed with a 11-yr running mean filter (thick line).

Multidecadal changes of the MOC as indicated by the dipole index lag those of the NAO by about a decade, supporting the notion that a significant fraction of the low-frequency variability of the MOC is driven by that of the NAO. (top) Annual data of LSW thickness (m), a measure of convection in the Labrador Sea, at ocean weather ship Bravo, defined between isopycnals $\sigma_{1.5} = 34.72\text{--}34.62$, following (Curry et al. 1998). b) Correlation of an MOC index with North Atlantic SST as a function of the time lag computed from a multimillennial control integration with the Kiel Climate Model (KCM). From Latif et al. (2009).

Figures

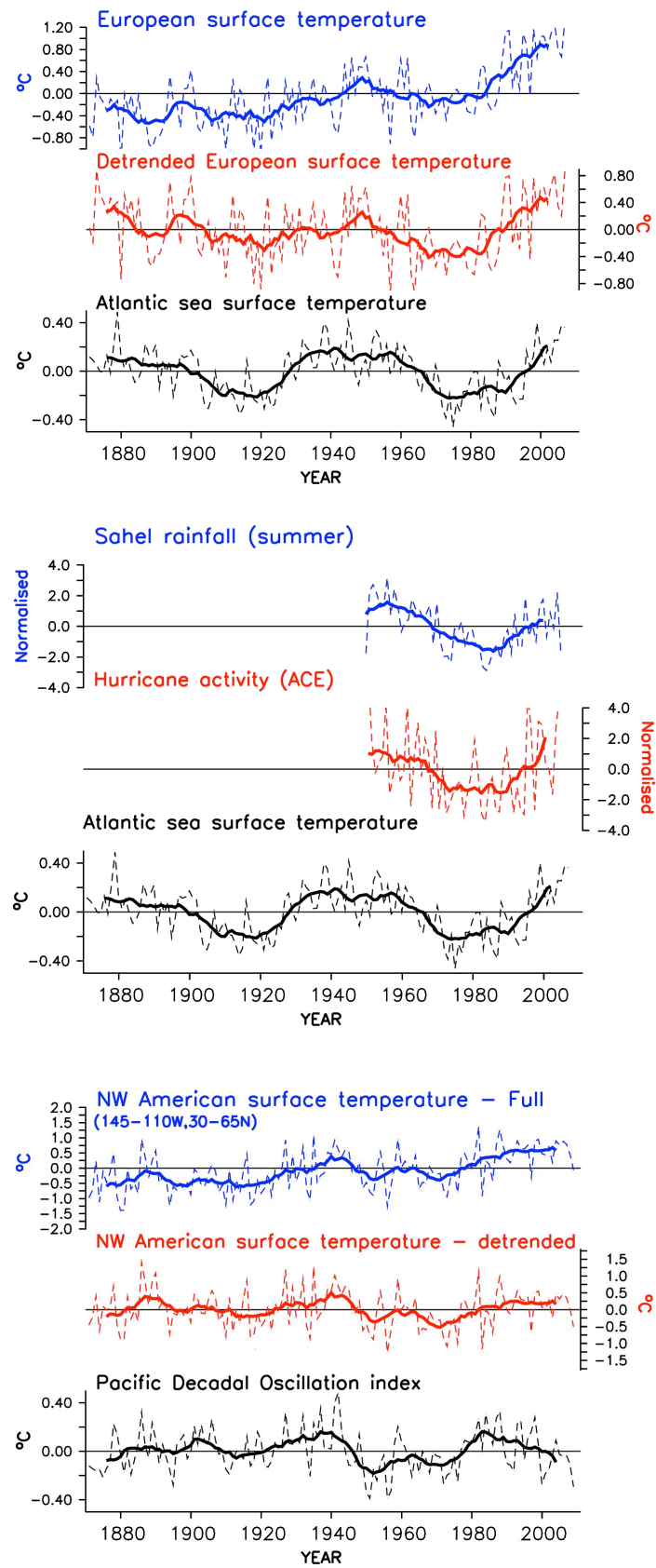


Fig. 1

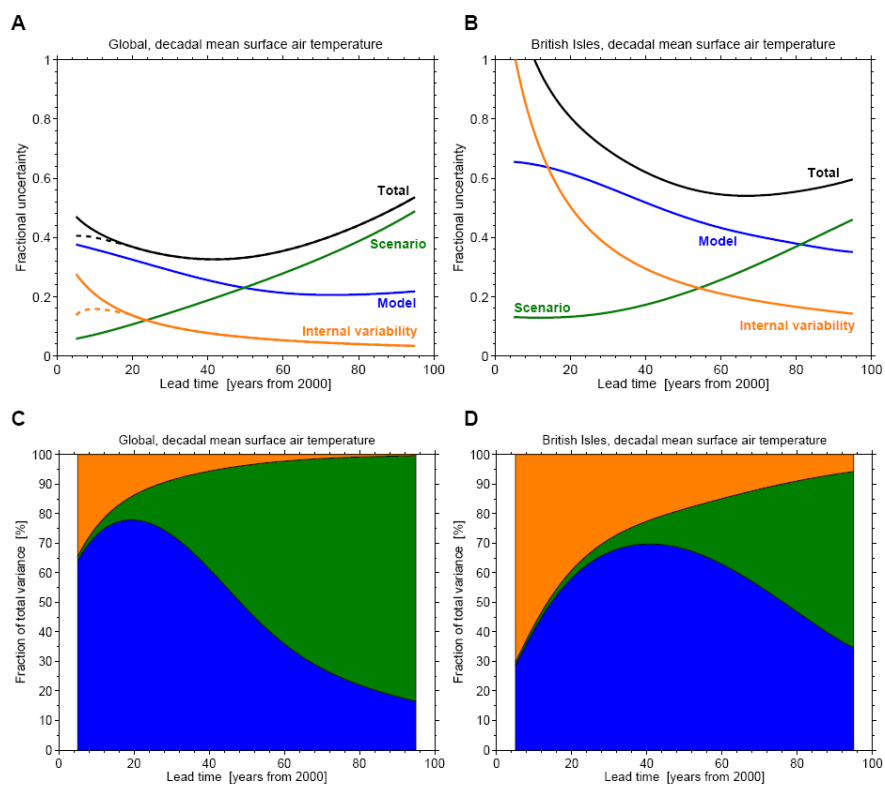


Fig. 2

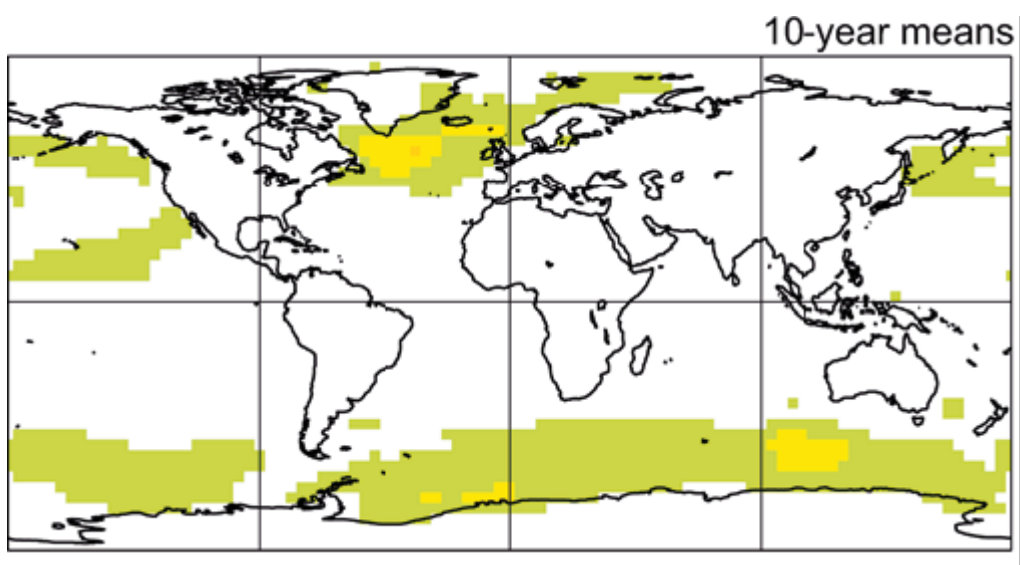


Fig. 3

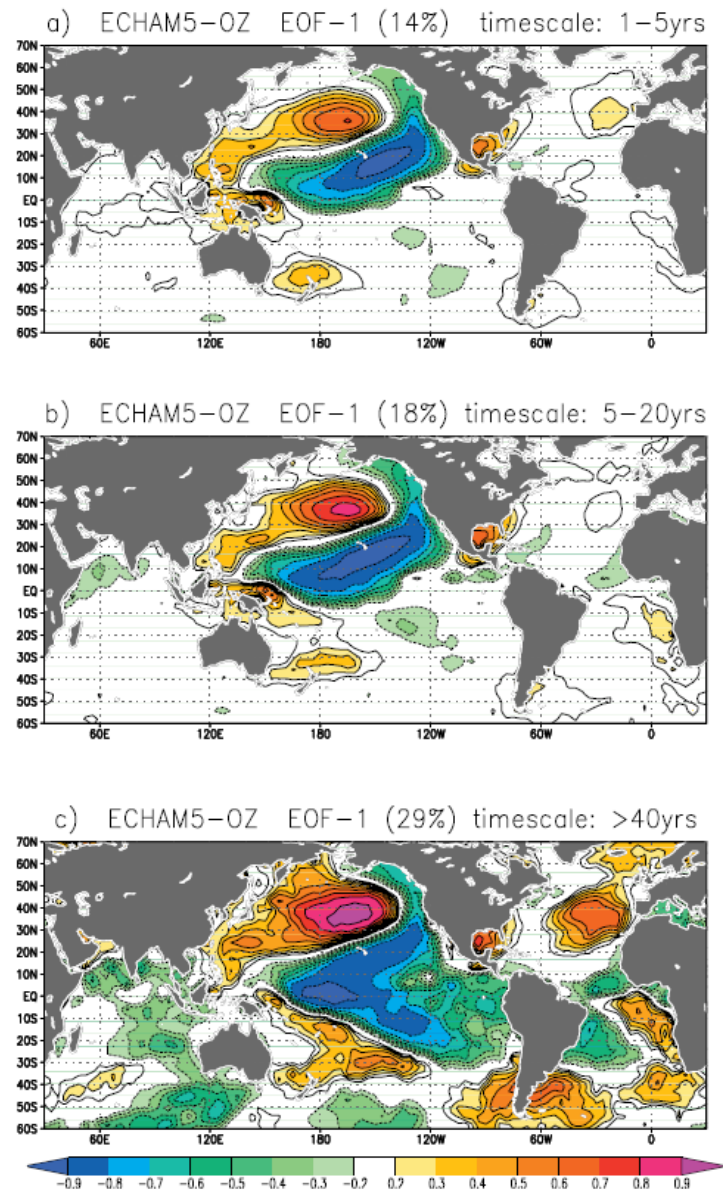


Fig. 4

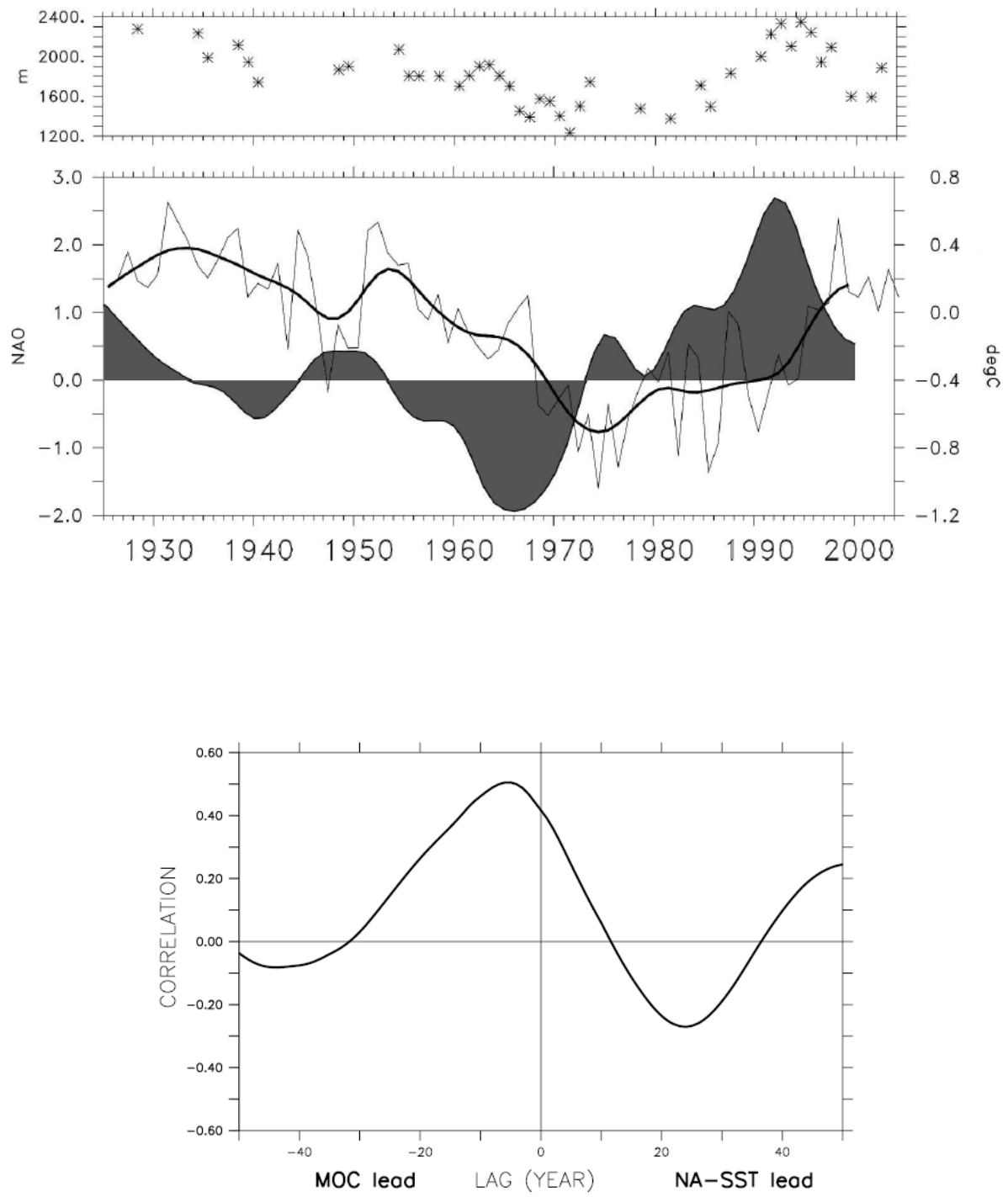


Fig. 5