

An interdecadal oscillatory mode of the AMOC related to westward-propagating temperature anomalies in ocean and climate models (CMIP5)

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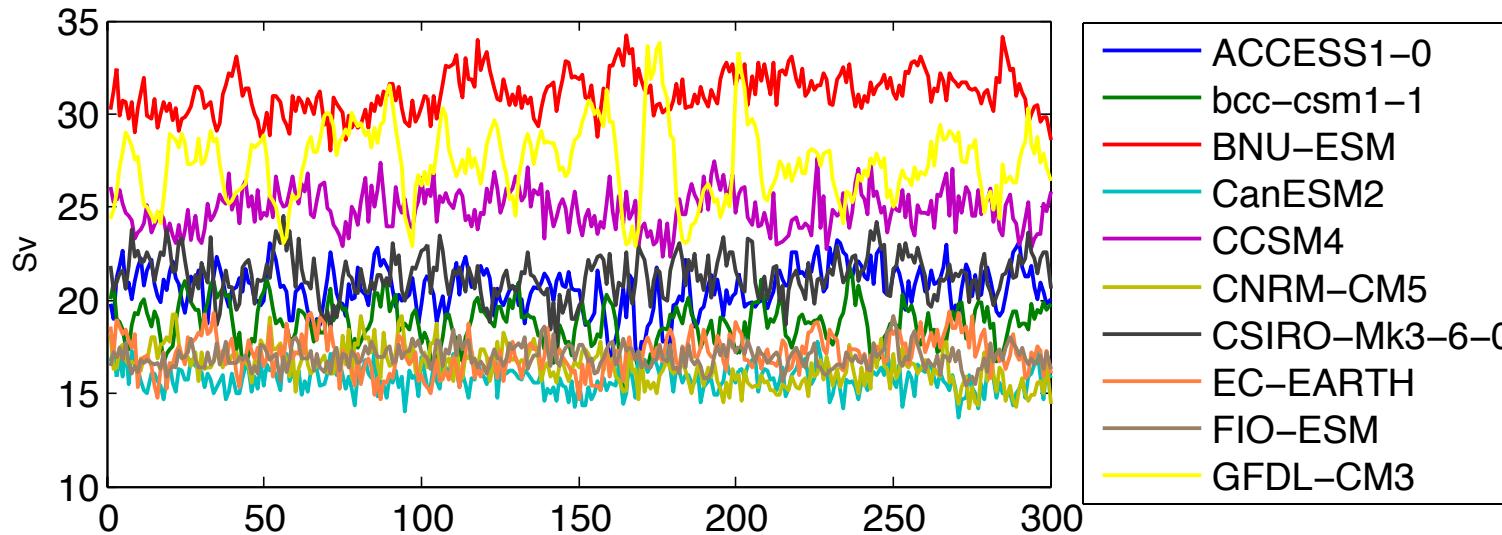


August 2012

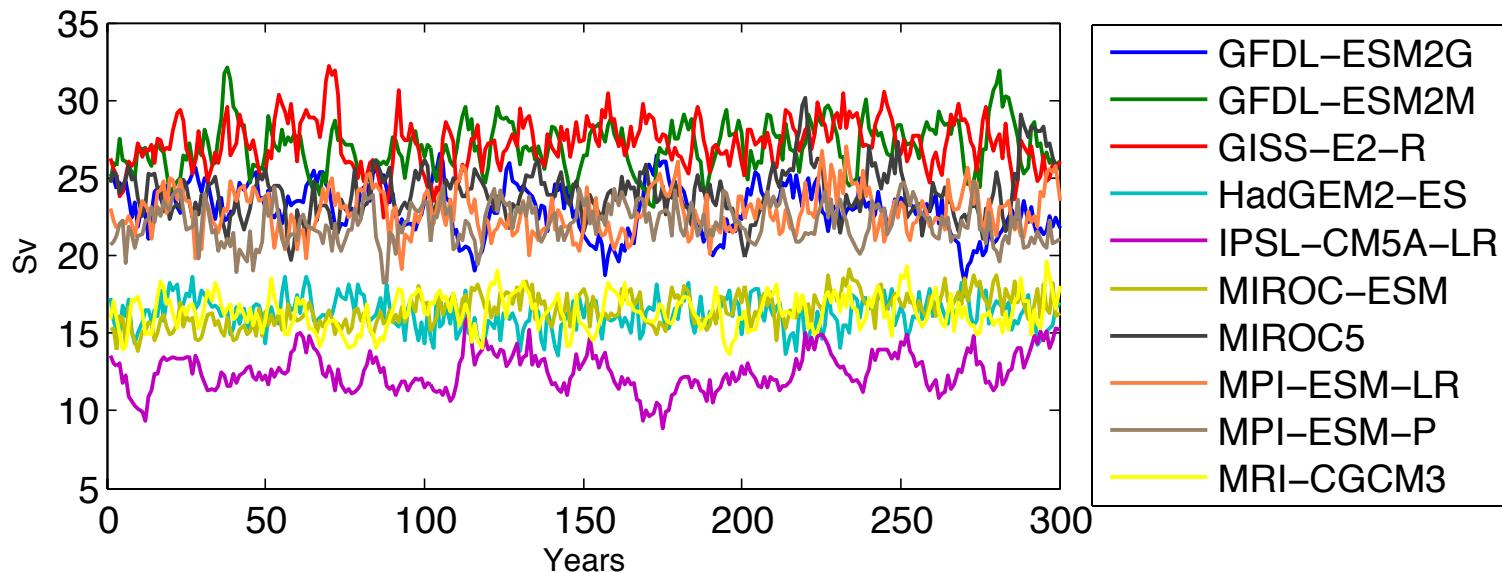
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Atlantic Meridional Overturning Circulation (AMOC)

**AMOC
decadal
and multi-
decadal
variability**



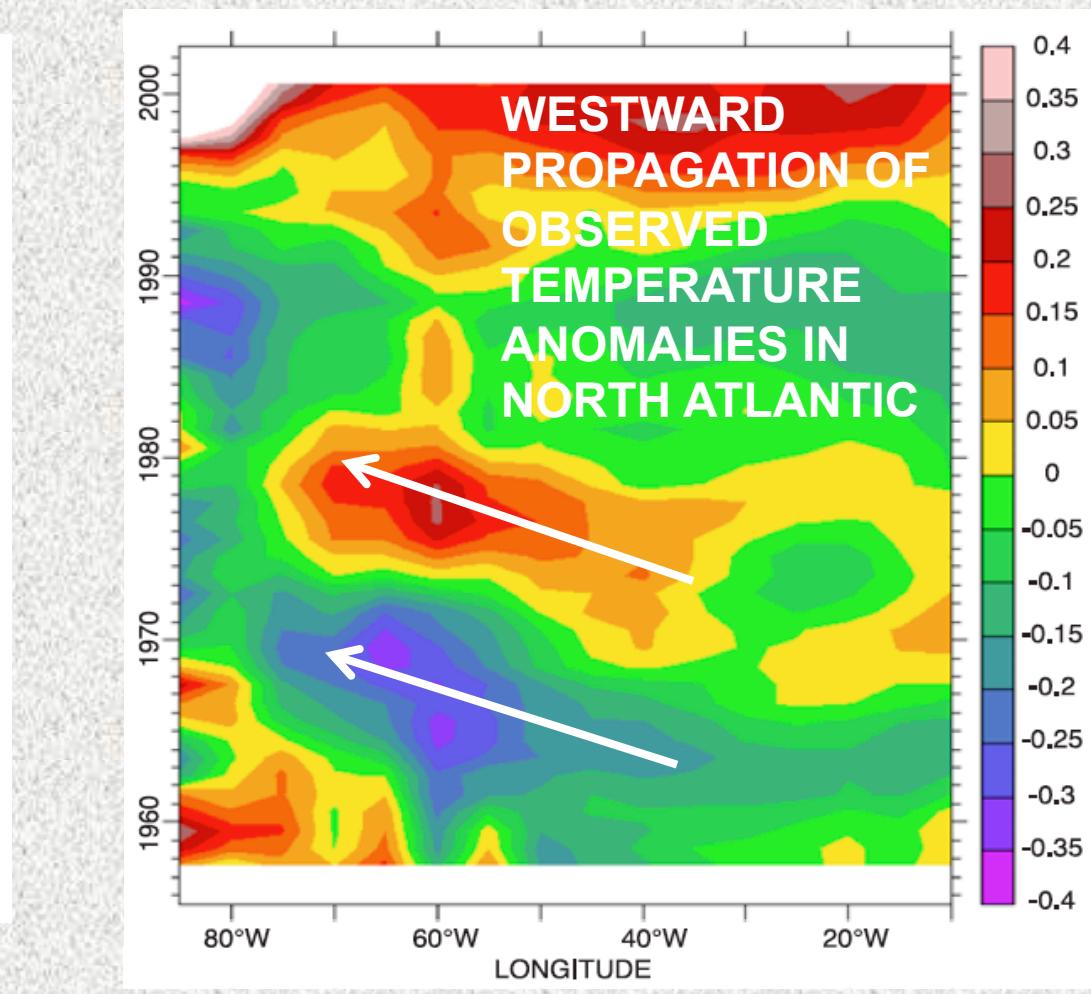
CMIP5 (Muir and Fedorov 2012, in prep)



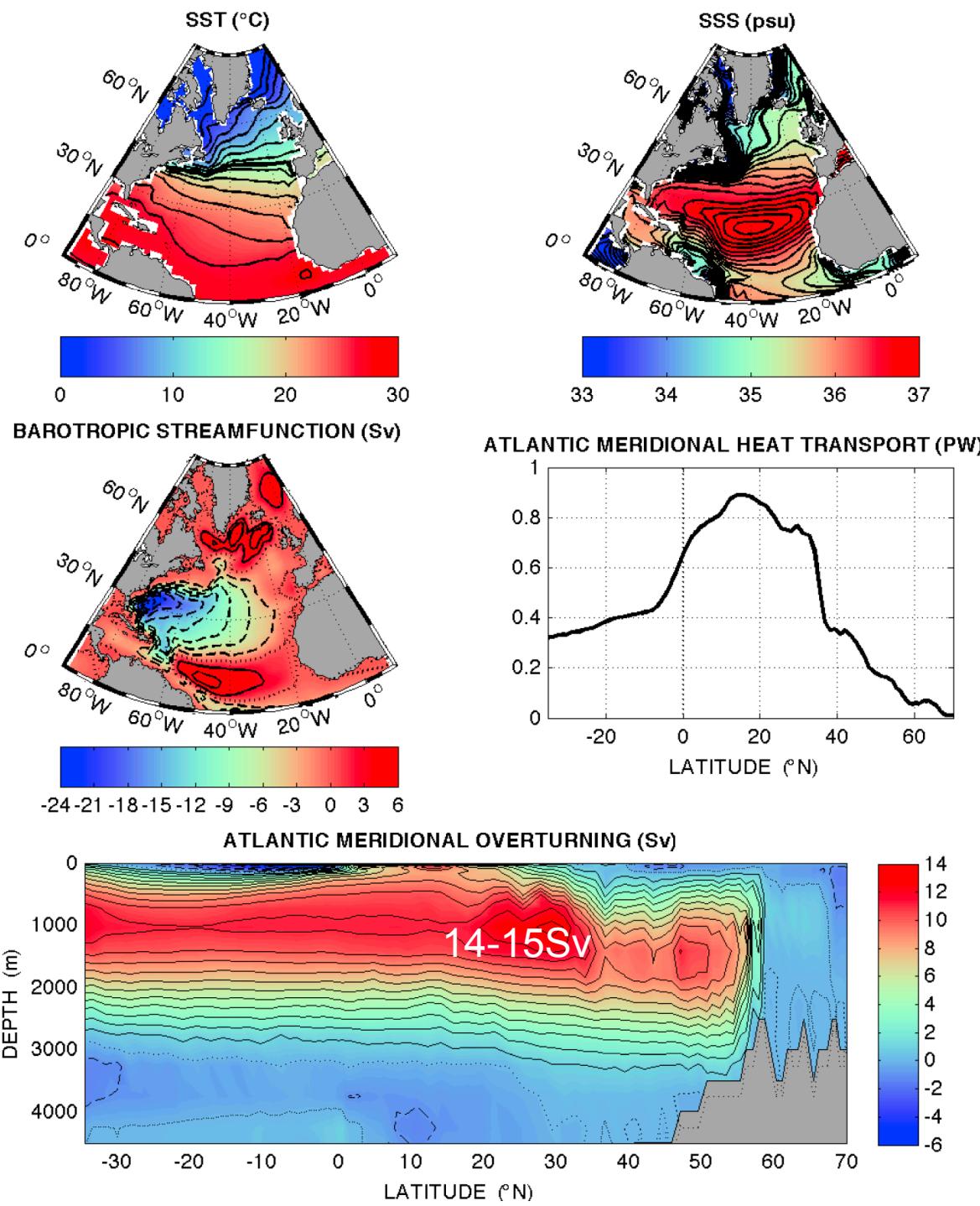
Working hypothesis: AMOC variability in GCMs is related to an oceanic internal mode sustained by noise or weak ocean-atmosphere coupling

➤ **Emphasis:** Westward propagation of temperature (density) anomalies in the northern Atlantic interacting with the AMOC

Huck et al. 1999, Colin de Verdière and Huck 1999, Marshall et al 2000, te Raa and Dijkstra 2002, Dijkstra et al. 2006, Frankcombe et al. 2009, Sévellec et al. 2009, Sévellec and Fedorov 2012, Muir and Fedorov 2012, in prep.



Hovmoller diagram of observed temperature anomalies averaged between 300-400m and 10–60°N in the North Atlantic (XBT data), Frankcombe et al 2008



Ocean GCM:
OPA 8.2
2 $^{\circ}$ -global
configuration
31 levels (ORCA2)

We use tangent linear and adjoint versions of the model

**Could we extract
the leading AMOC
eigenmode in this
ocean GCM?**

**Sevellec and
Fedorov 2012**

1. Ocean GCM :

Method:

$$\frac{d\mathbf{X}}{dt} = \mathbf{F}(\mathbf{X}, t)$$

Non-autonomous

X - the state vector

2. Linearize

$$\frac{dx'}{dt} = \left. \frac{\partial F}{\partial X} \right|_{X_o} x'$$

$$X = X_o + x'$$

X_o - seasonally varying mean state of the ocean

x' - anomalies

3. Integrate between t_1 and t_2

$$\mathbf{x}(t_2) = \mathbf{M}(t_1, t_2) \mathbf{x}(t_1)$$

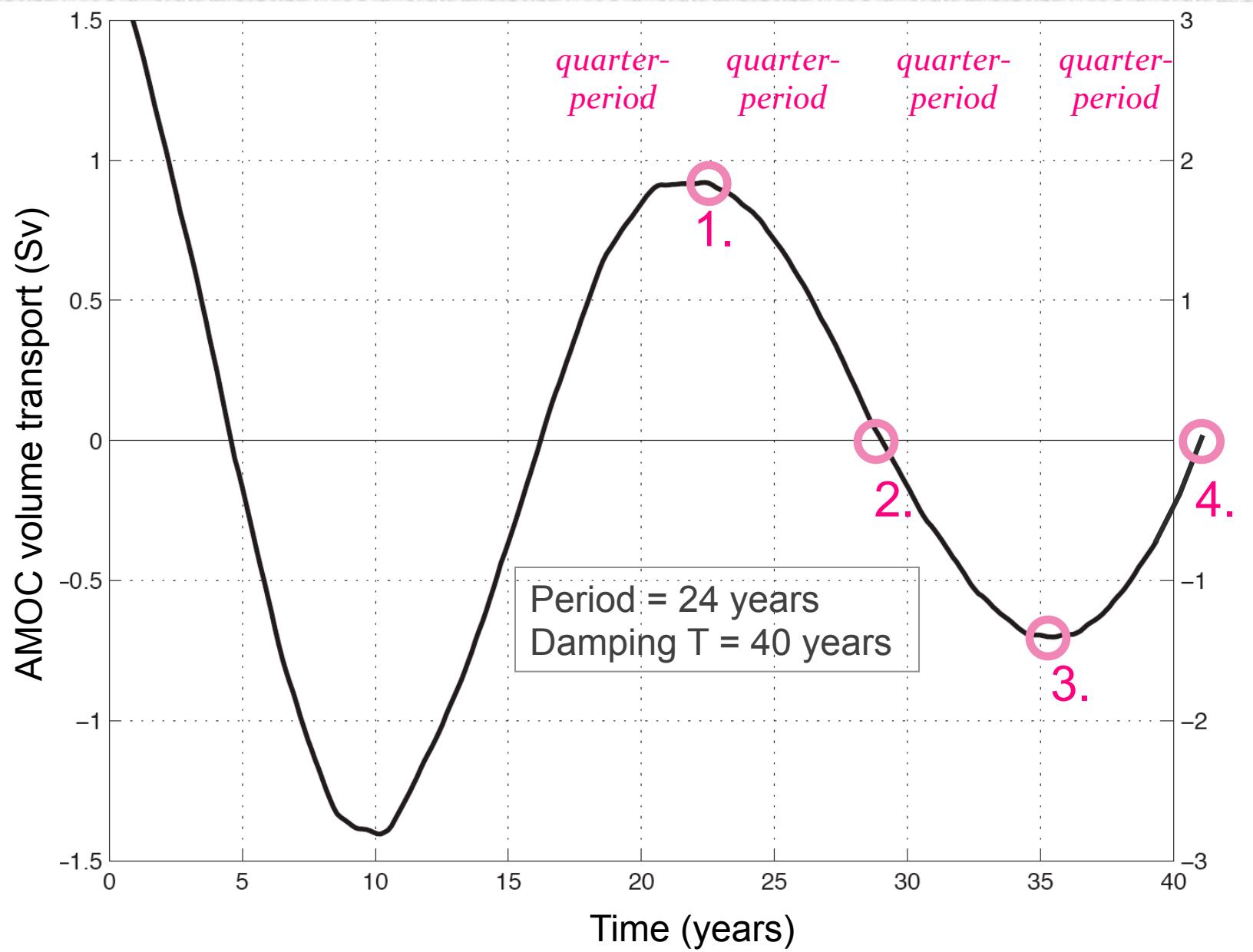
M - the linear propagator of the system

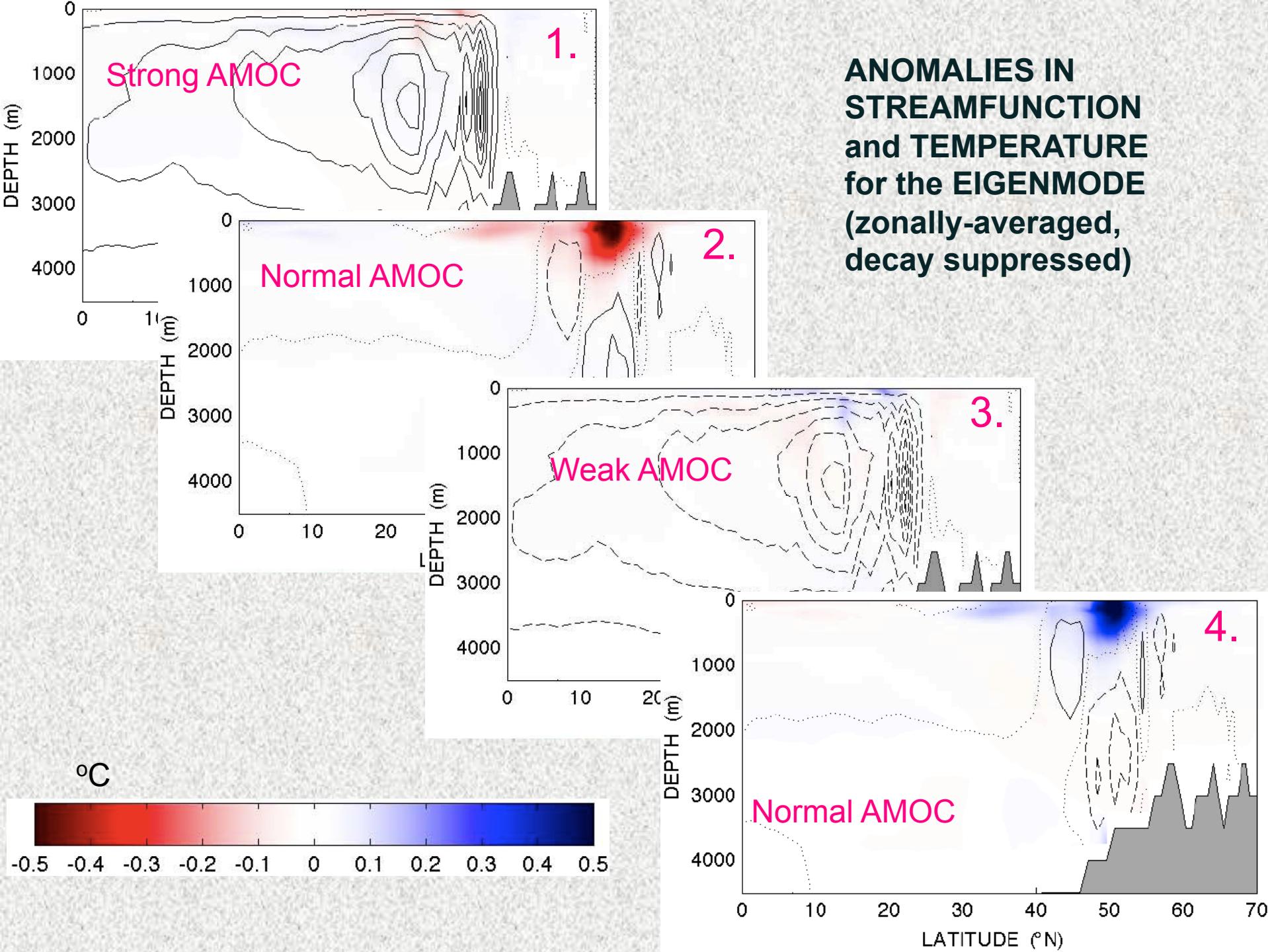
4. Eliminate the seasonal cycle from \mathbf{M}

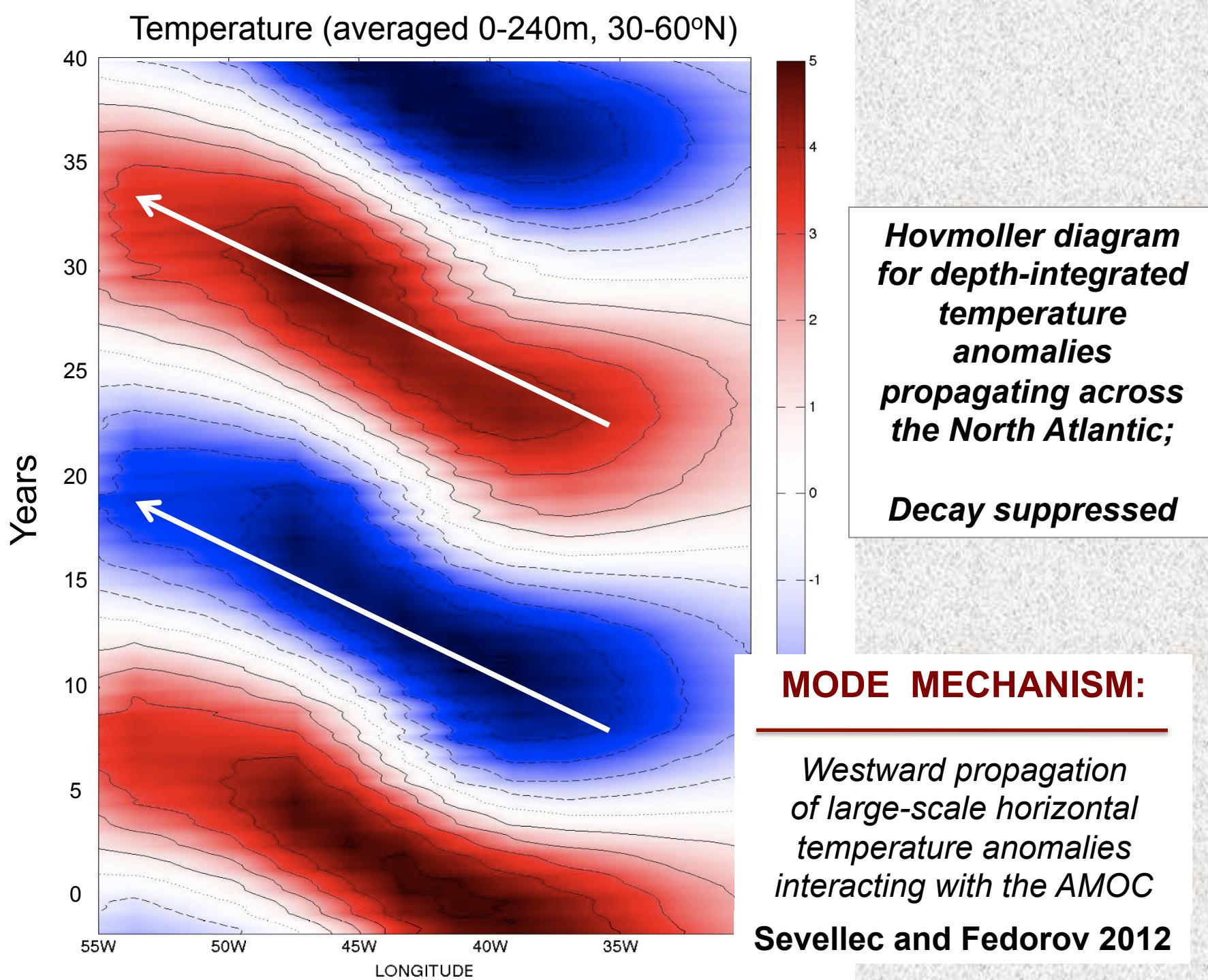
$\tilde{\mathbf{M}} = \mathbf{M}(t, t + n \cdot \text{year}) \rightarrow$ a Poincare section,
e.g. consider \mathbf{M} on every Jan 1

5. Obtain the least damped eigen-mode of $\tilde{\mathbf{M}}$

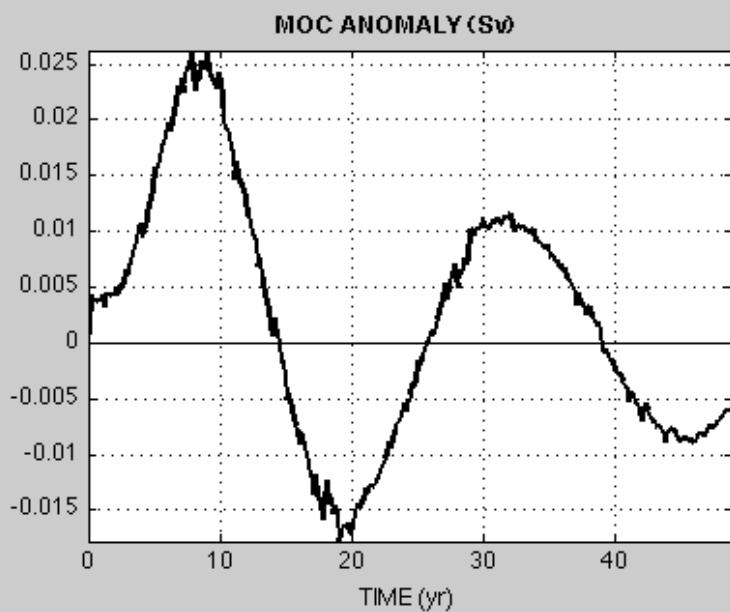
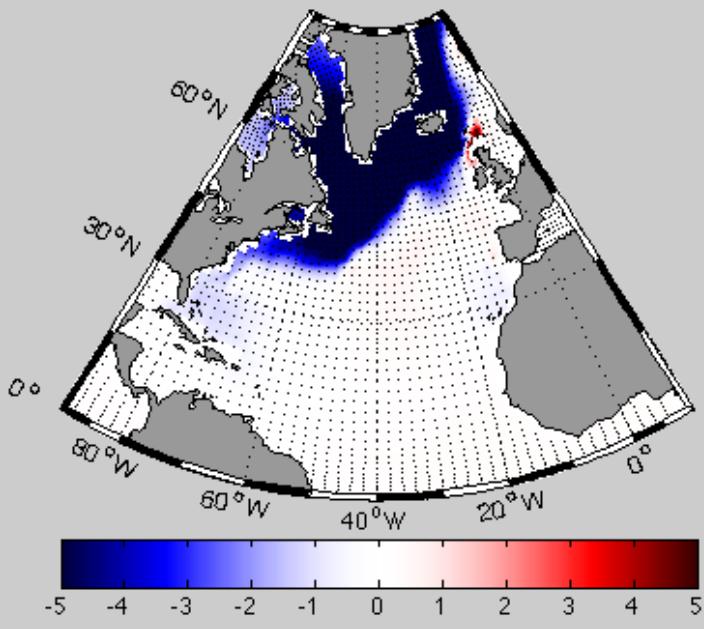
The least-damped mode: AMOC variations







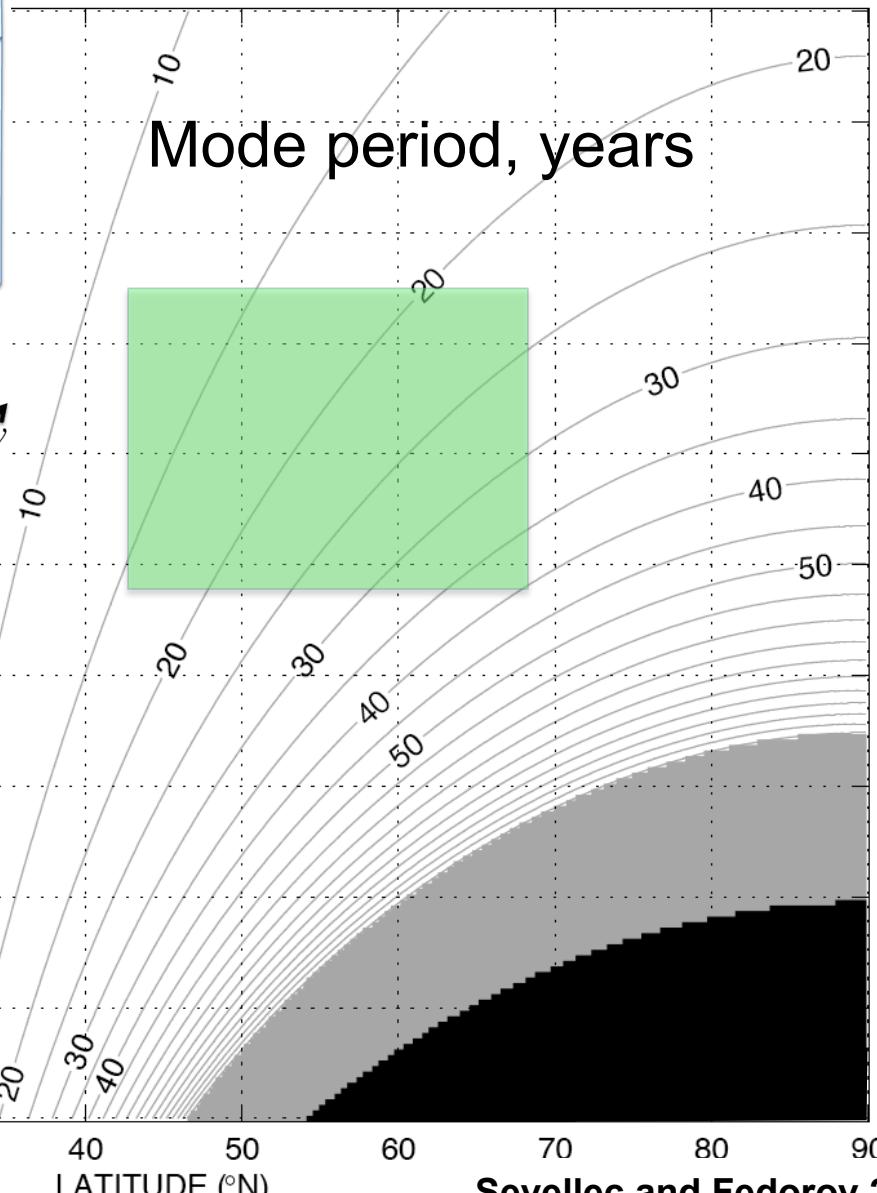
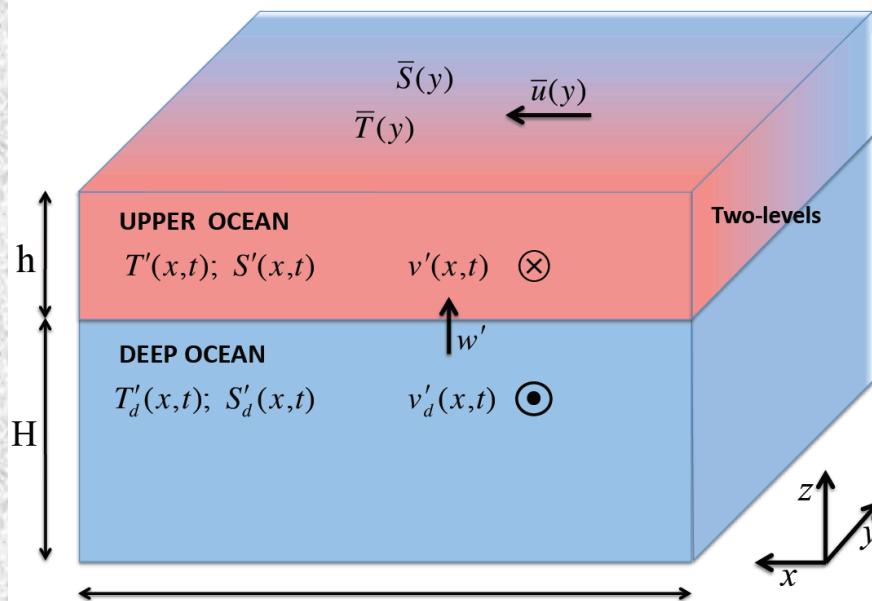
TEMPERATURE ($\times 10^{-3}$ K), Z-MEAN = 0 - 240 m



Excitation of the eigenmode by
optimal initial perturbations
(temperature anomalies
around
Greenland)

Sevellec and
Fedorov 2012b

OSCILLATION PERIOD: idealized 2-layer model

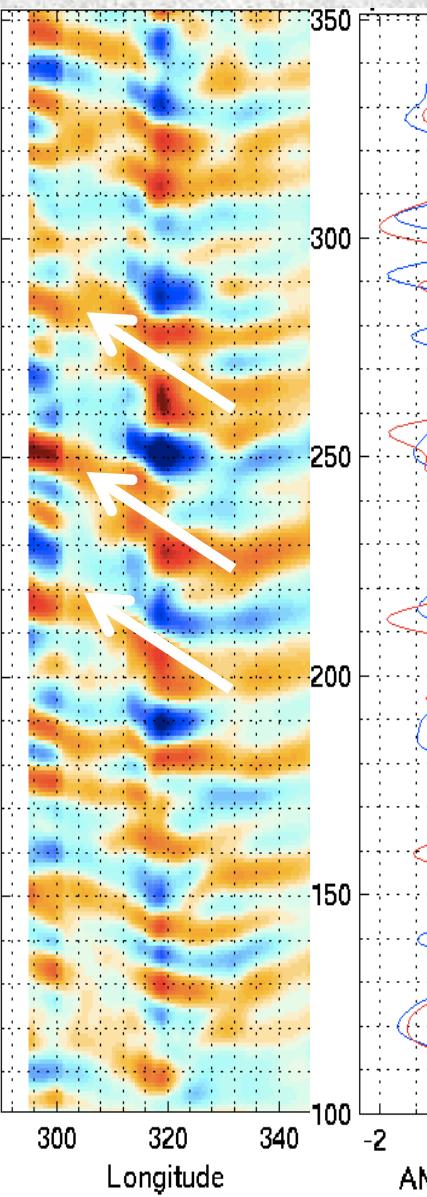
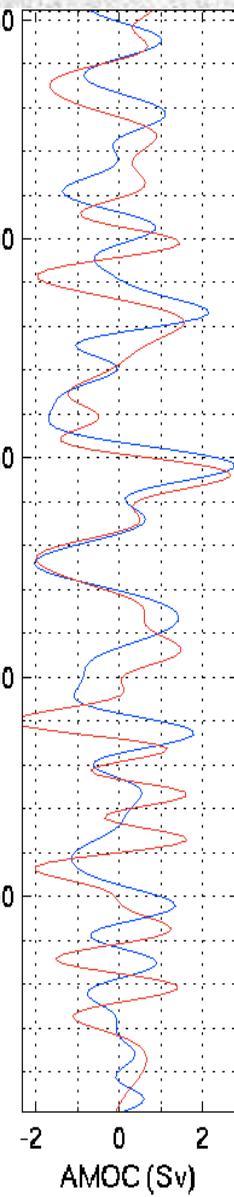
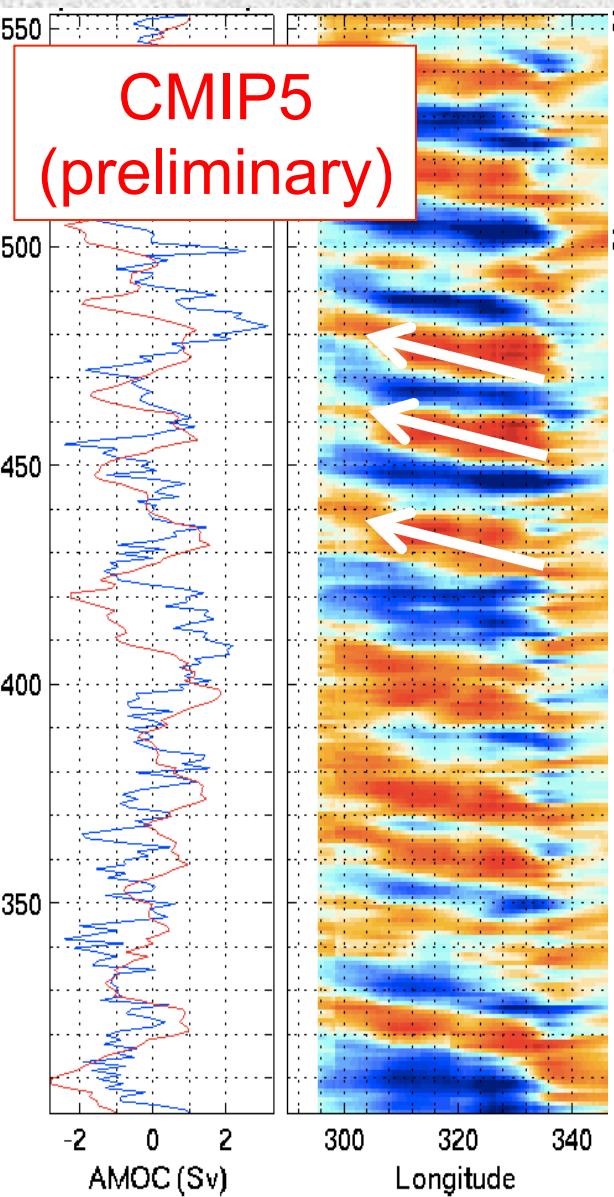


IPSL-CM5A-LR

GFDL-ESM2G

MIROC5

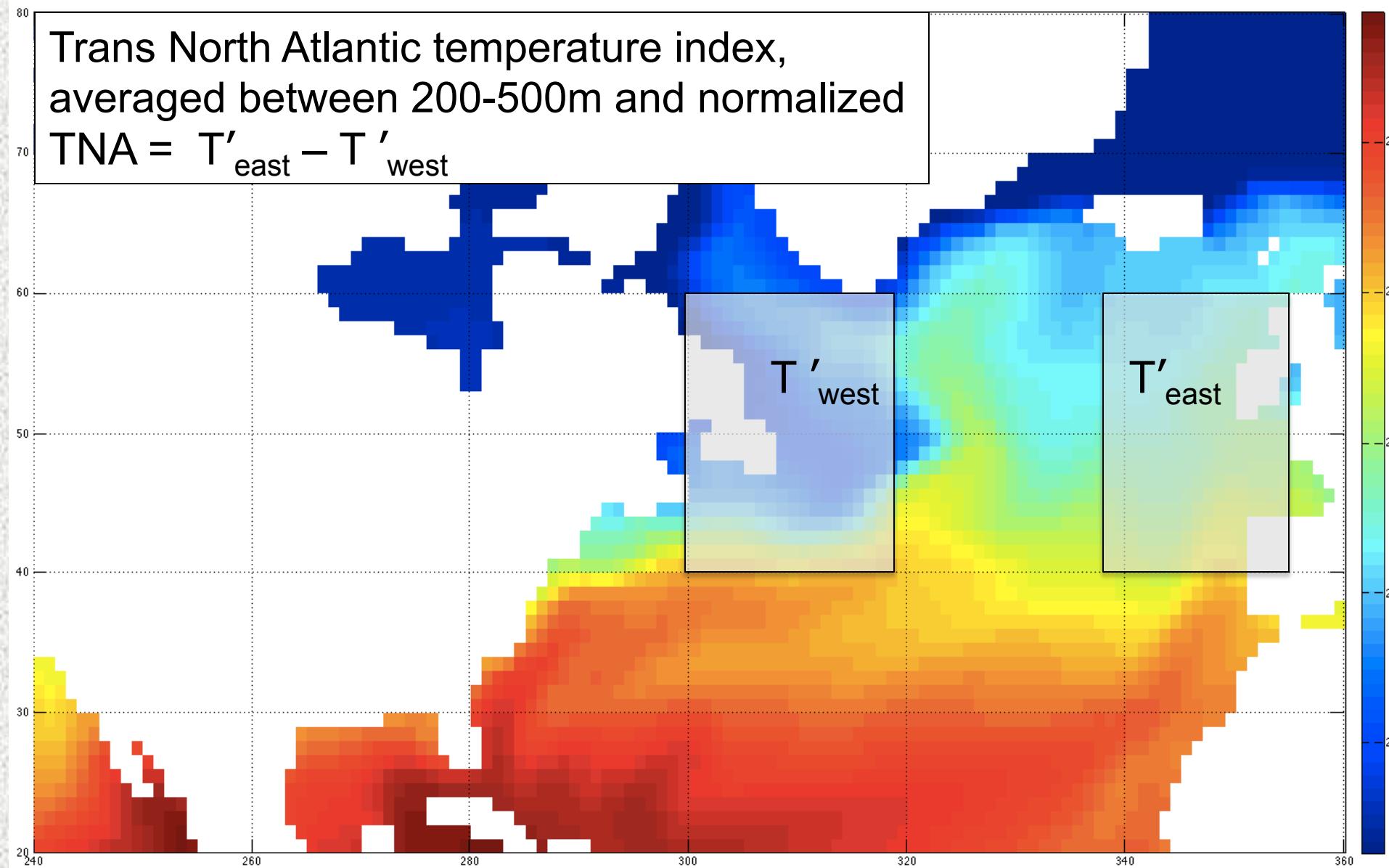
CMIP5
(preliminary)

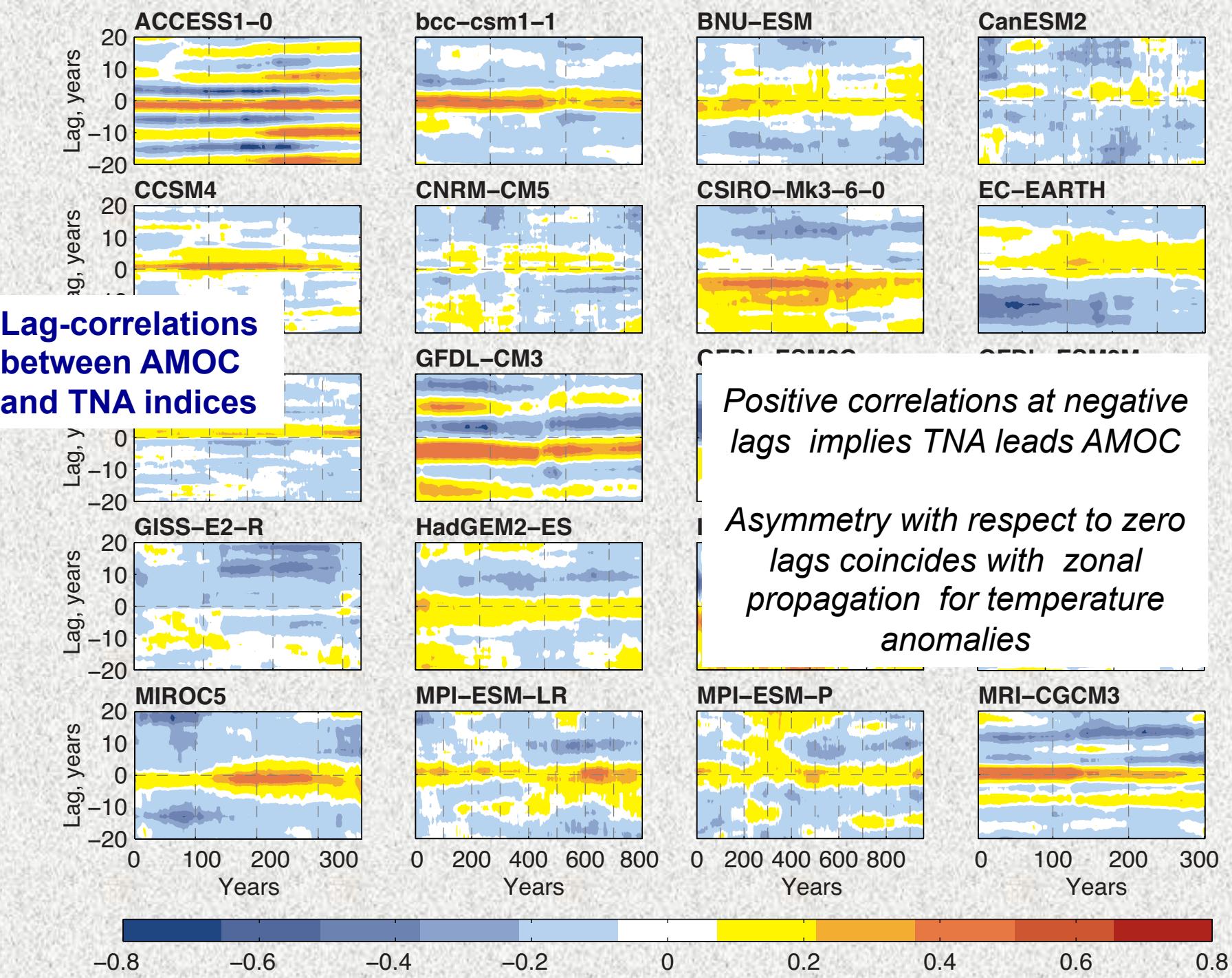


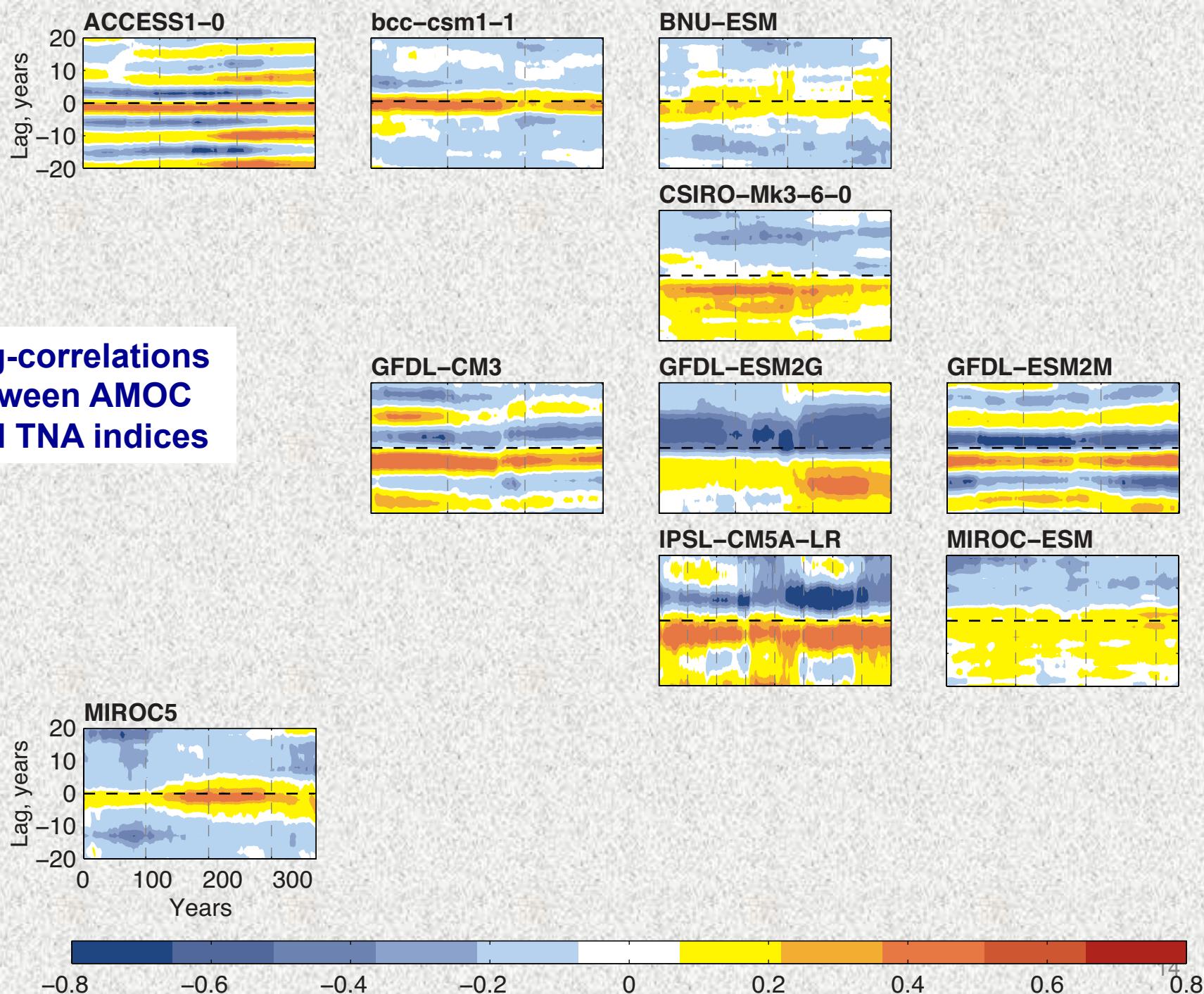
Blue line = AMOC (Sv), Red line = TNA temperature index (normalized)
Hovmoller maps = temperature anomalies correlated with the AMOC

TNA temperature index

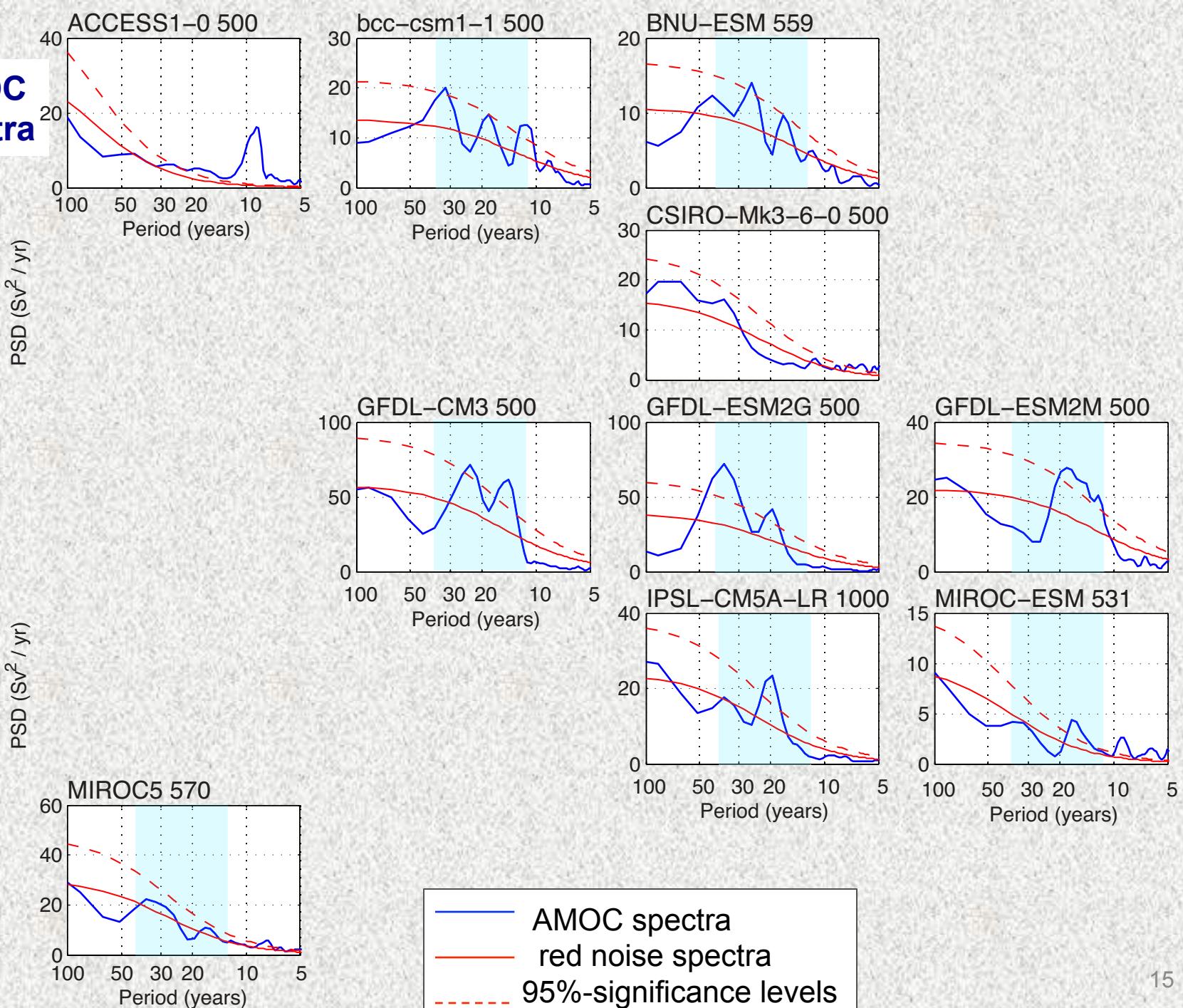
Trans North Atlantic temperature index,
averaged between 200-500m and normalized
 $TNA = T'_{east} - T'_{west}$



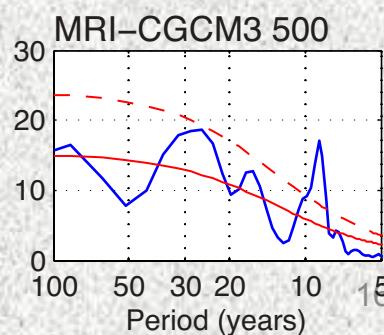
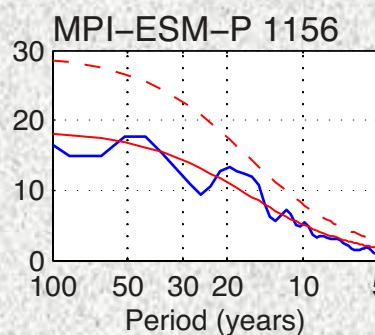
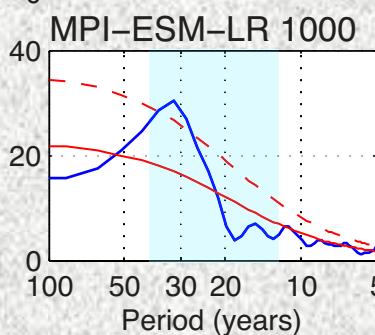
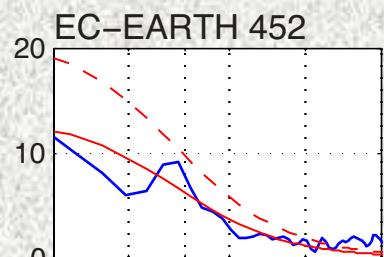
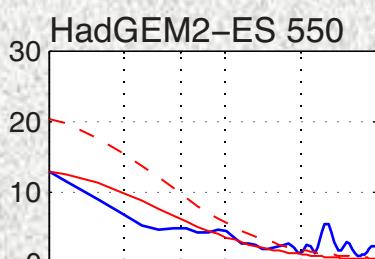
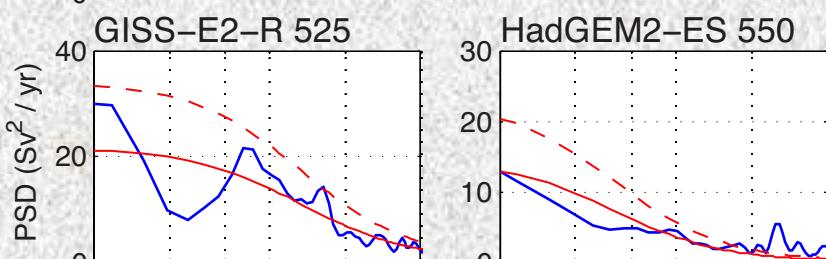
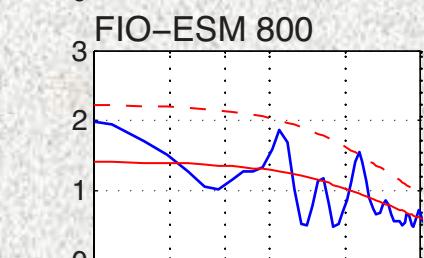
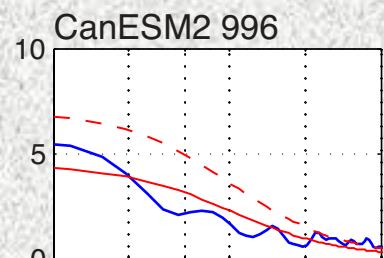
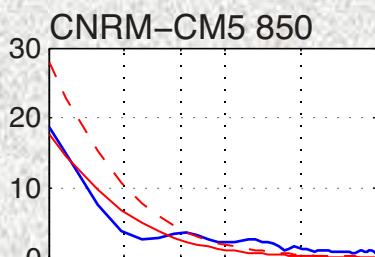
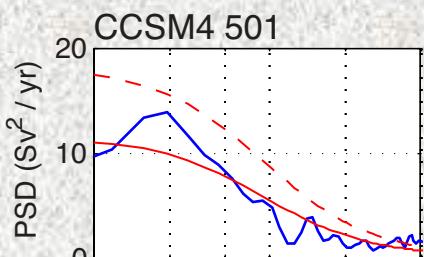




AMOC spectra



AMOC spectra



Summary

- We have identified the leading, interdecadal oscillatory eigenmode of the AMOC in a realistic ocean model ($T \sim 20$ years), also present in simple 2-layer models
- Its mechanism is related to westward propagation of temperature anomalies in the upper northern Atlantic ocean ($\sim 500m$) interacting with the AMOC
- This eigenmode is possibly present in CMIP5 models (IPSL-CM5, GFDL-ESM2G, MIROC5 ...), as evidenced by spectral peaks between 15-40 years and westward-propagating temperature anomalies correlated with the AMOC
- Lag-correlations between AMOC and TNA indices reach 0.8
- Most models without the westward propagation do not develop statistically significant interdecadal spectral peaks
- Also, a poster “AMOC variability and its effect on SST in CMIP5”

Lag-correlations between AMOC and TNA indices

