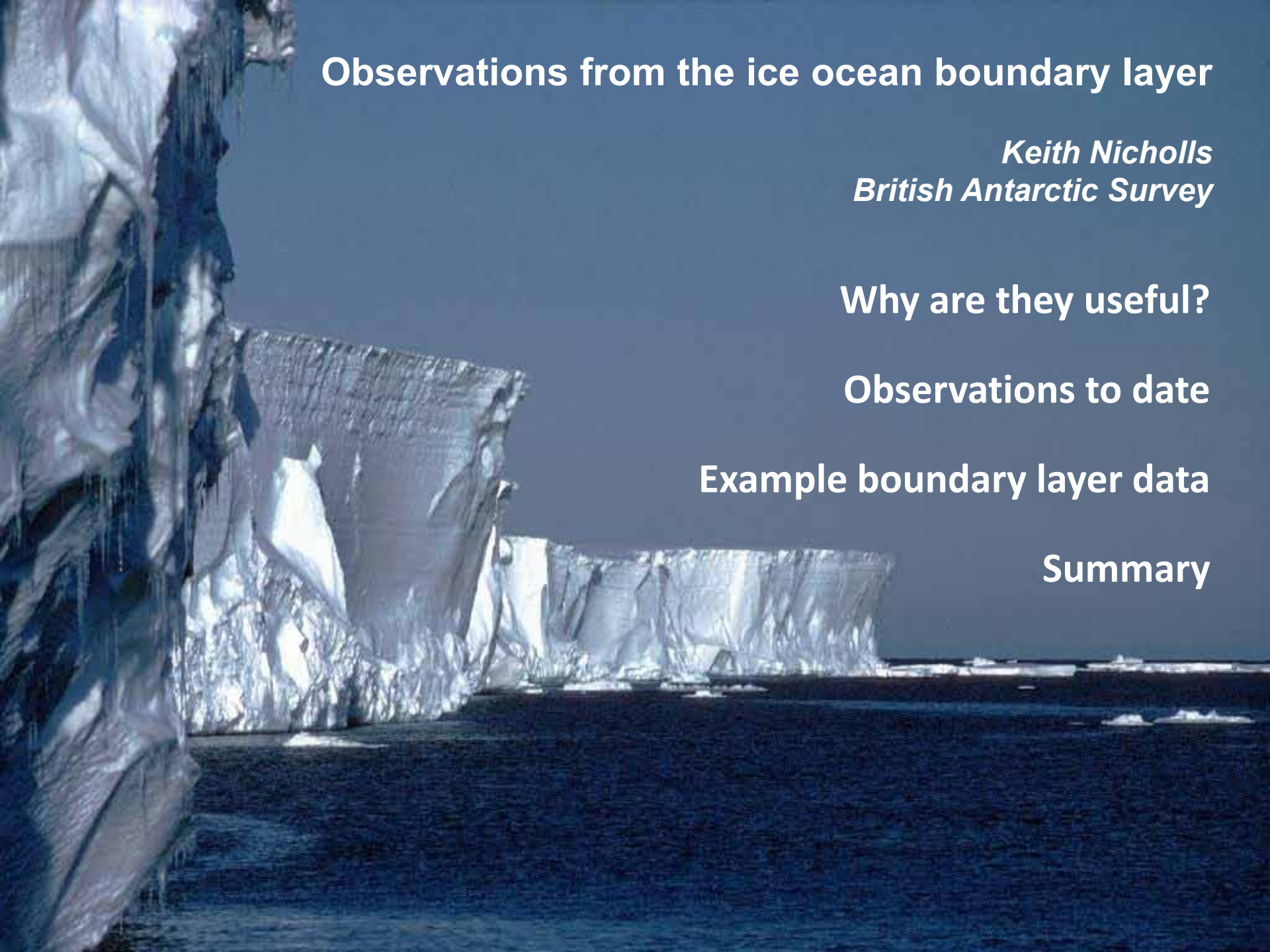




Observations from the ice ocean boundary layer

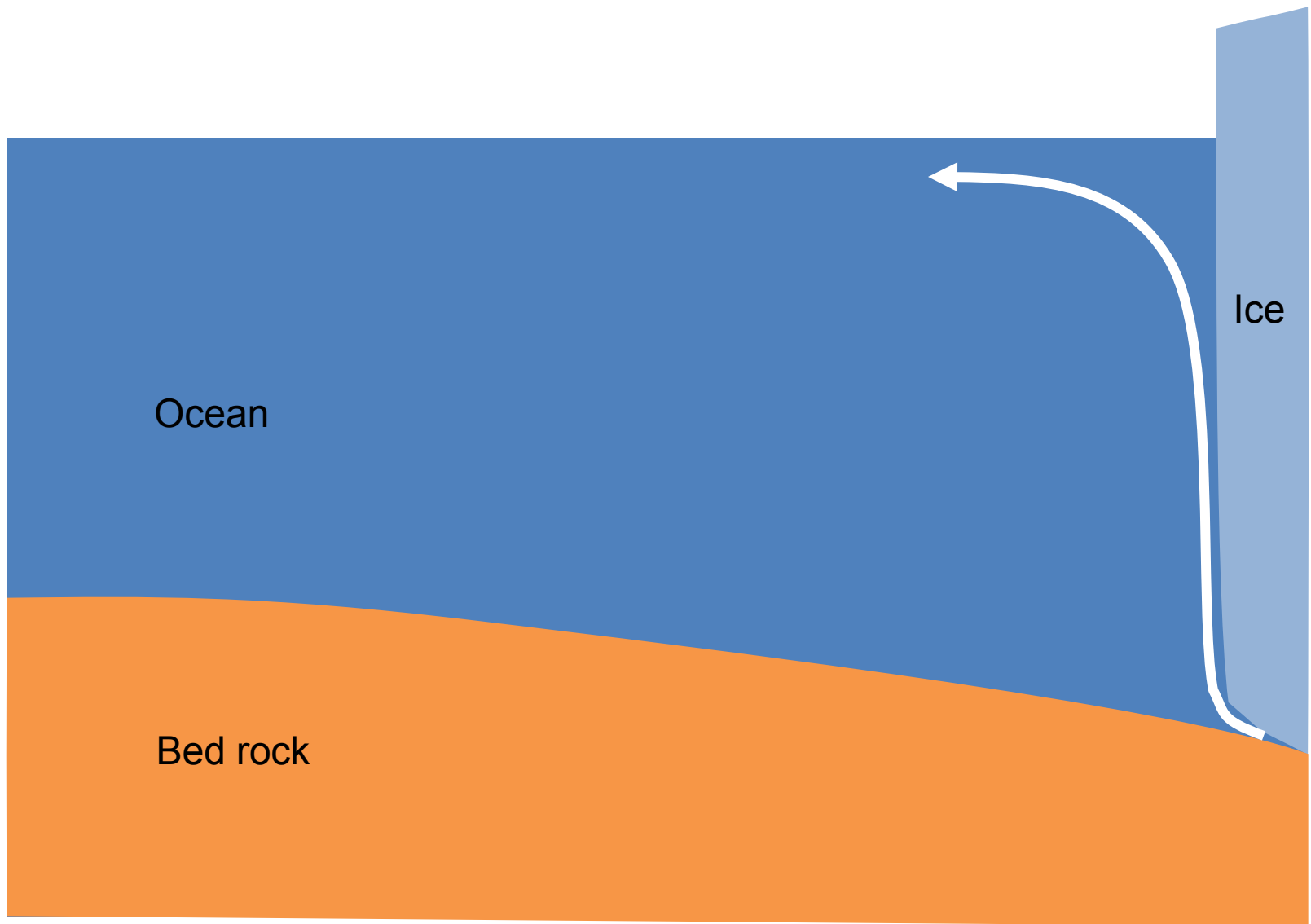
Keith Nicholls
British Antarctic Survey

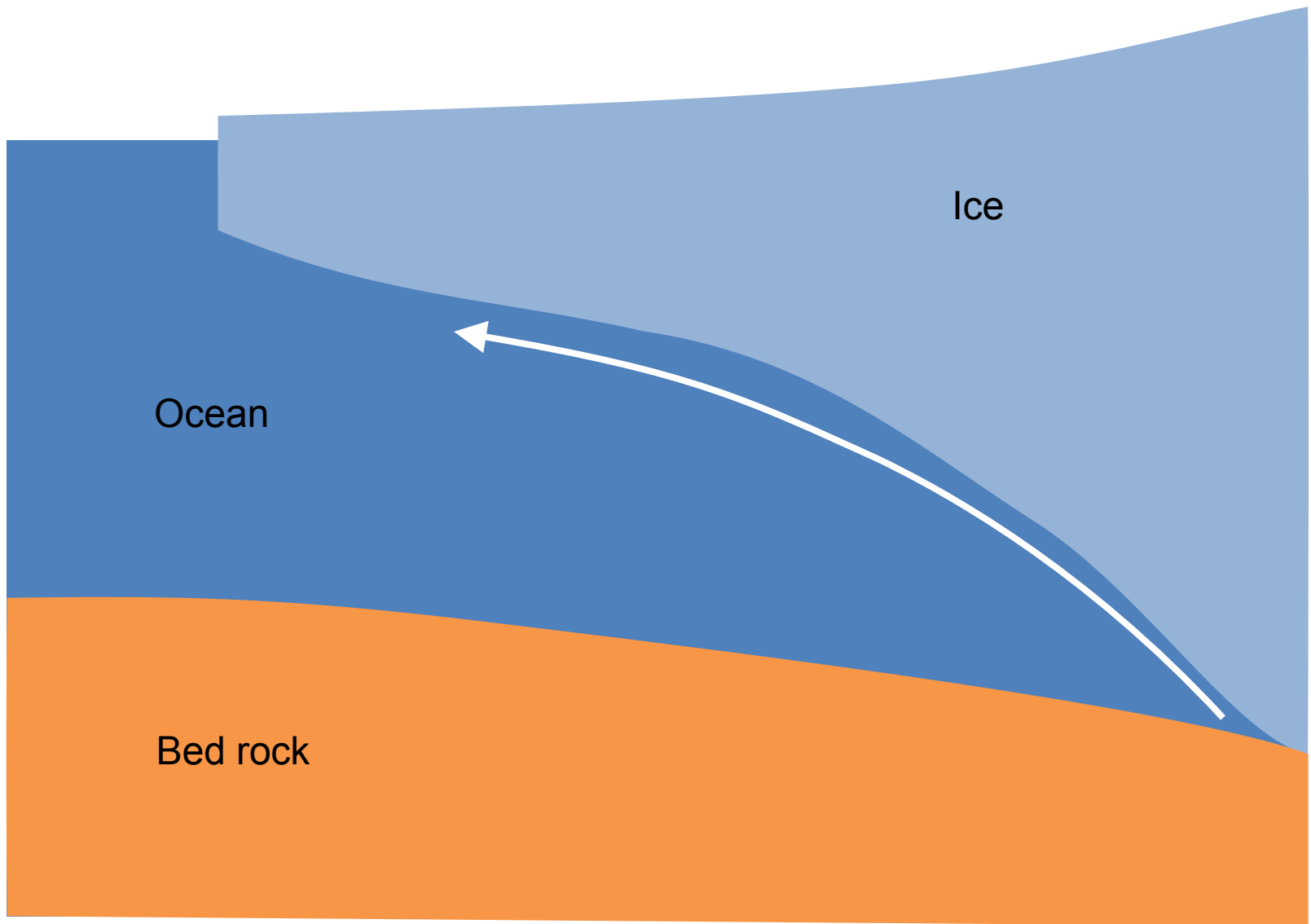
Why are they useful?

Observations to date

Example boundary layer data

Summary



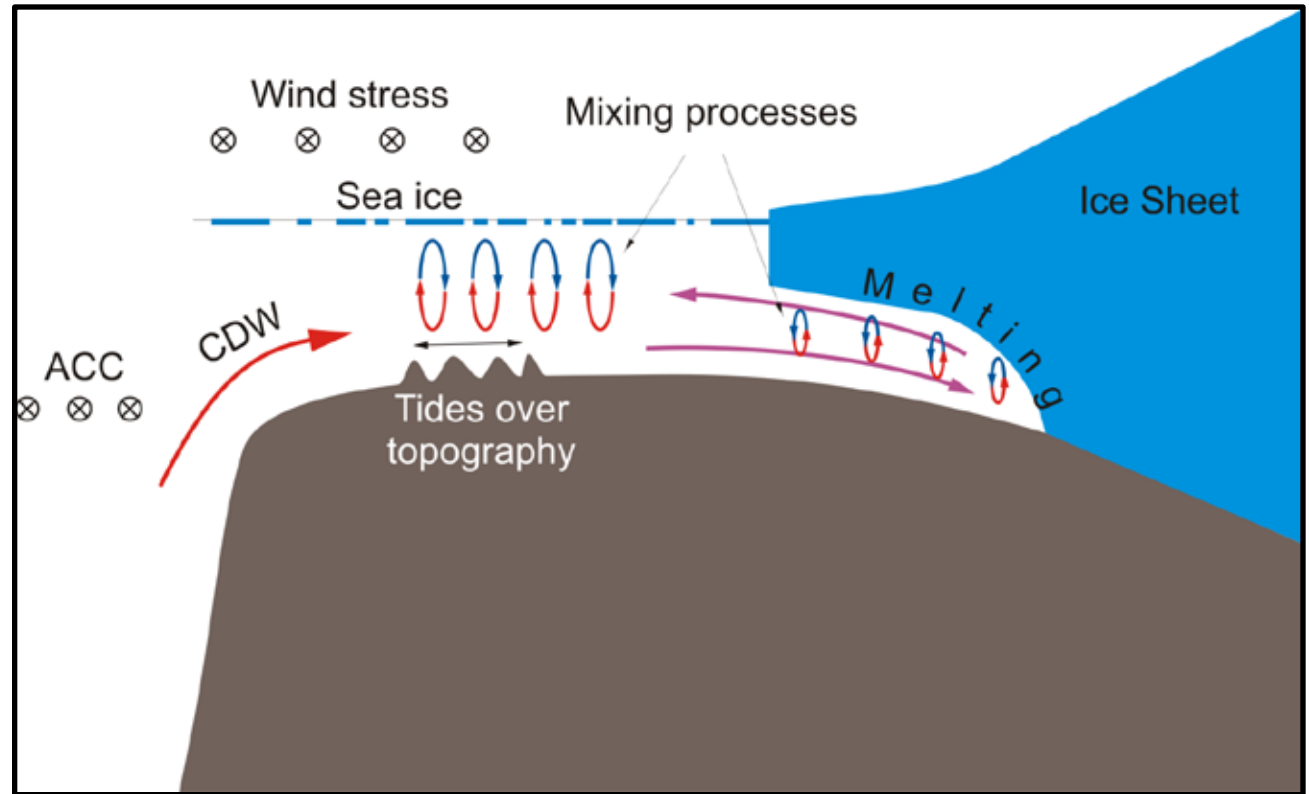


Need to predict:

Oceanographic forcing
on the ice shelves

Response of the ice
shelves to that
changing forcing

Response of the inland
ice sheet



Lynch-pin is the ability to calculate how fast the ice shelf base will melt given the oceanographic properties in the sub-ice shelf cavity

How do we calculate the rate of melting at the ice base?

Conservation of heat:

Melt rate $\sim$ Heat diffused across boundary layer - heat diffused into ice shelf

Conservation of salt:

Freshening at base balanced by salt diffusion across boundary layer

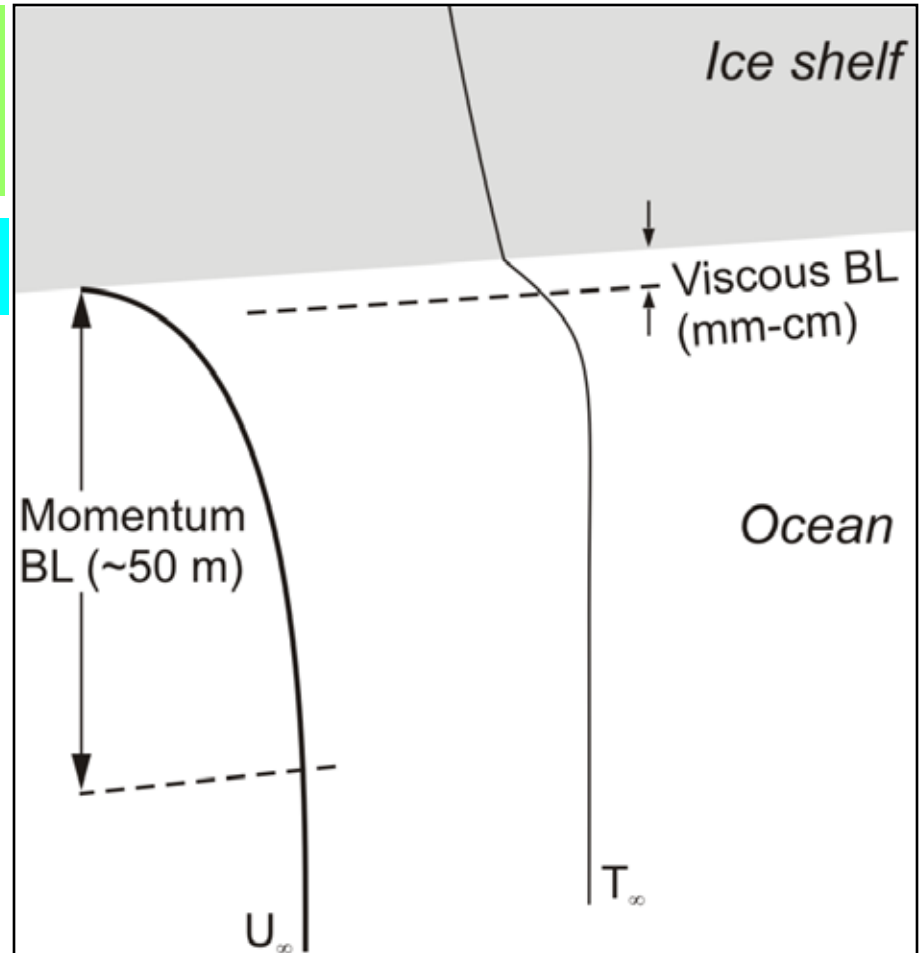
Ice base at the *in situ* freezing point

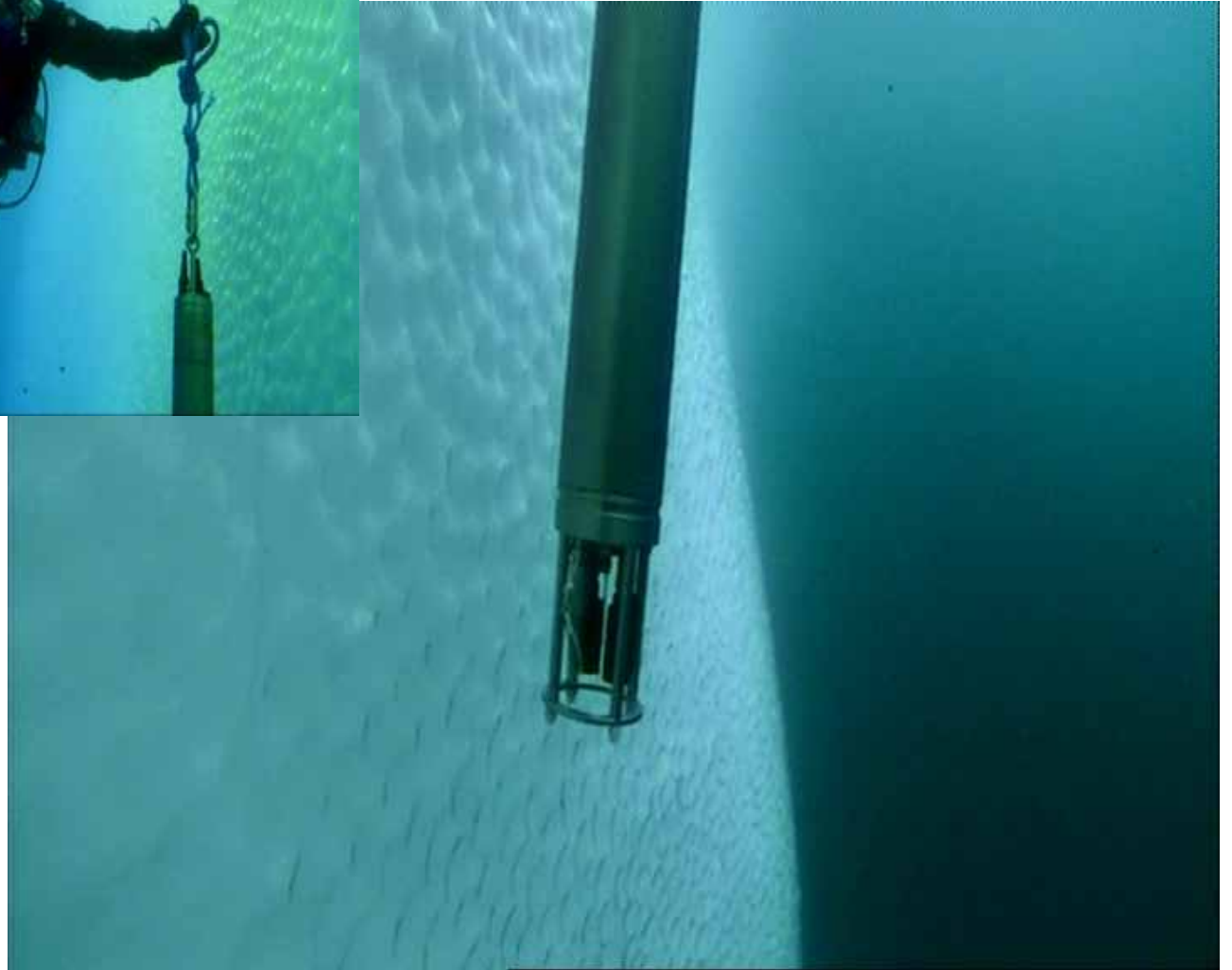
Need parameterisation for bulk transport coefficients for heat and salt

Traditionally used *Kader & Yaglom* (1972)

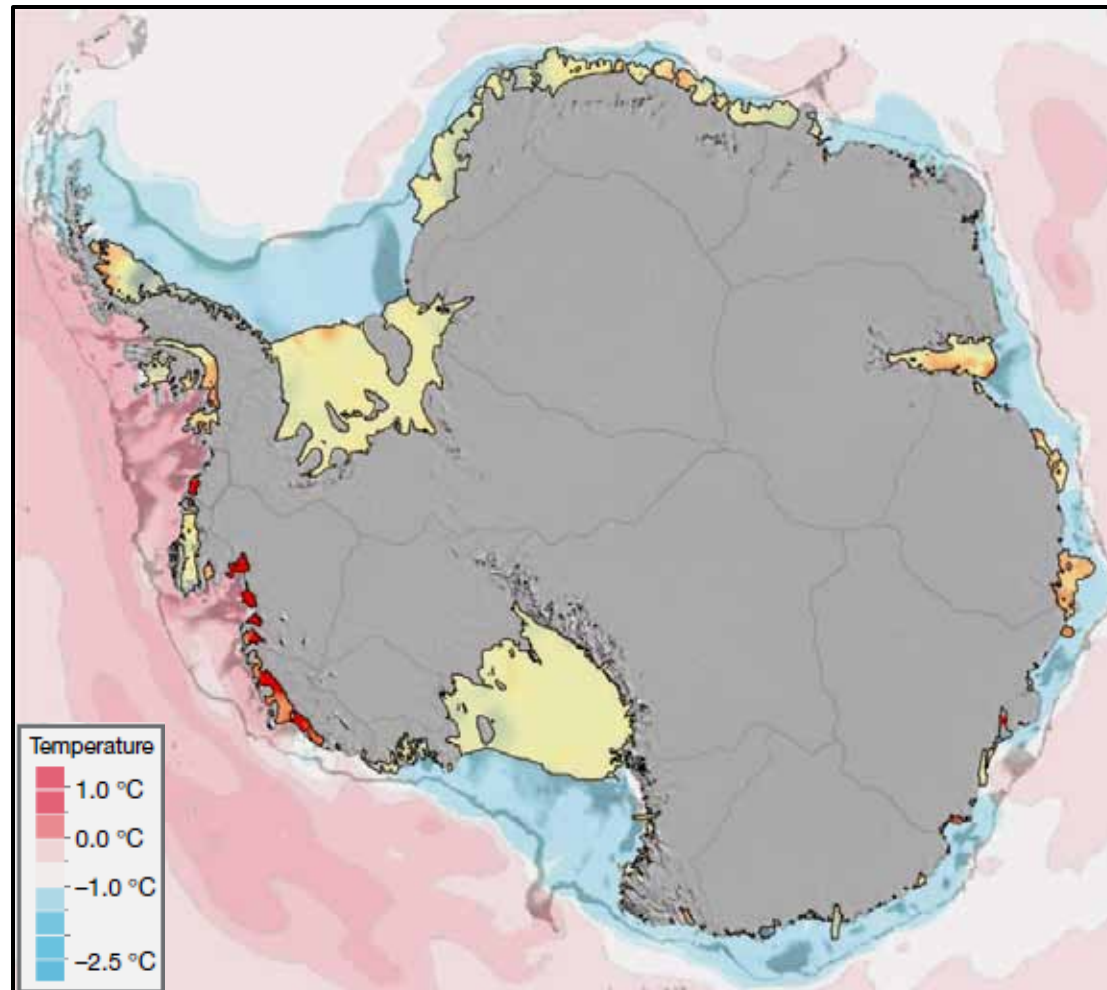
Parameterisation assumes

- fully turbulent flow
- neutrally stable
- “smooth, flat” ice base



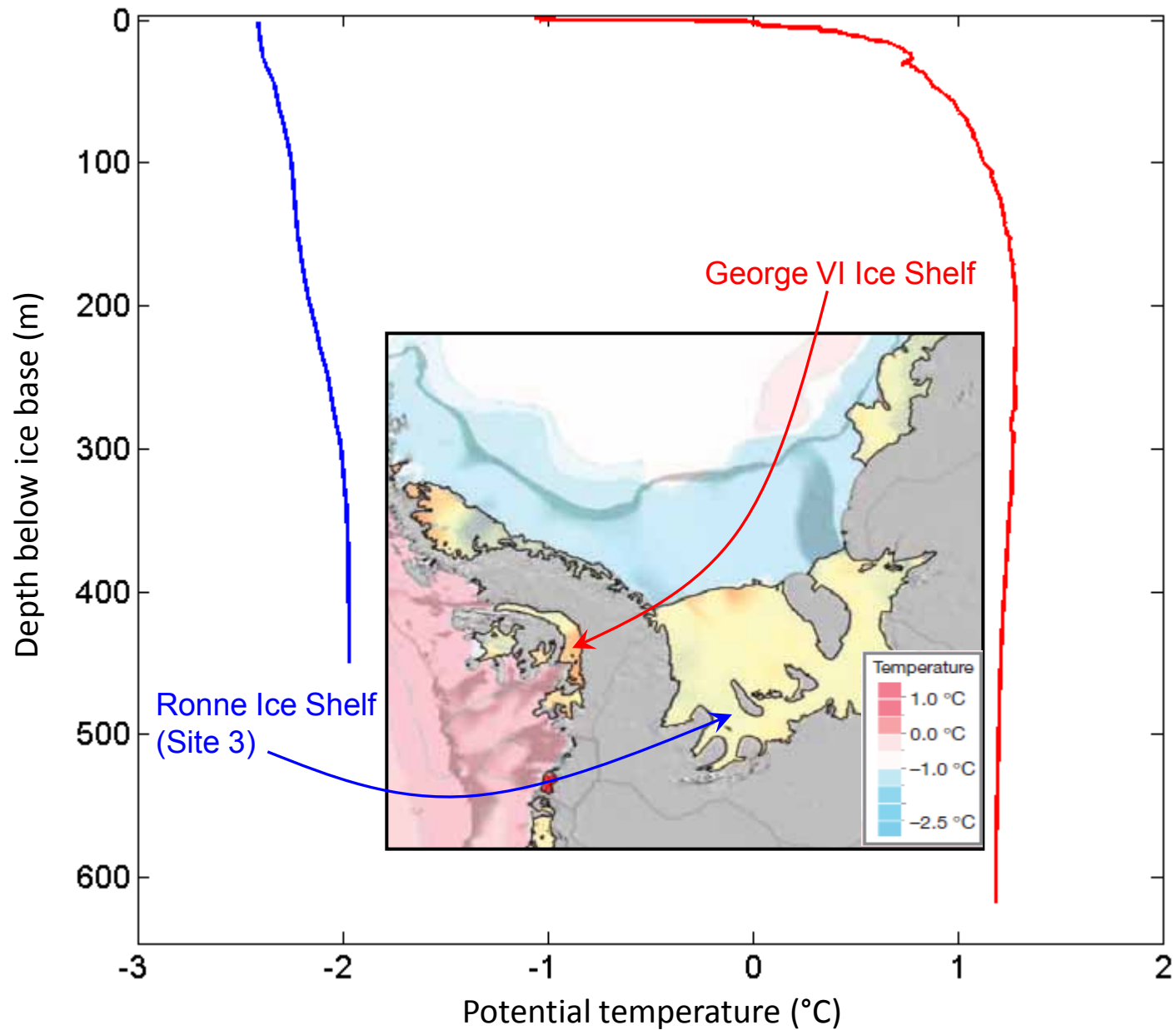


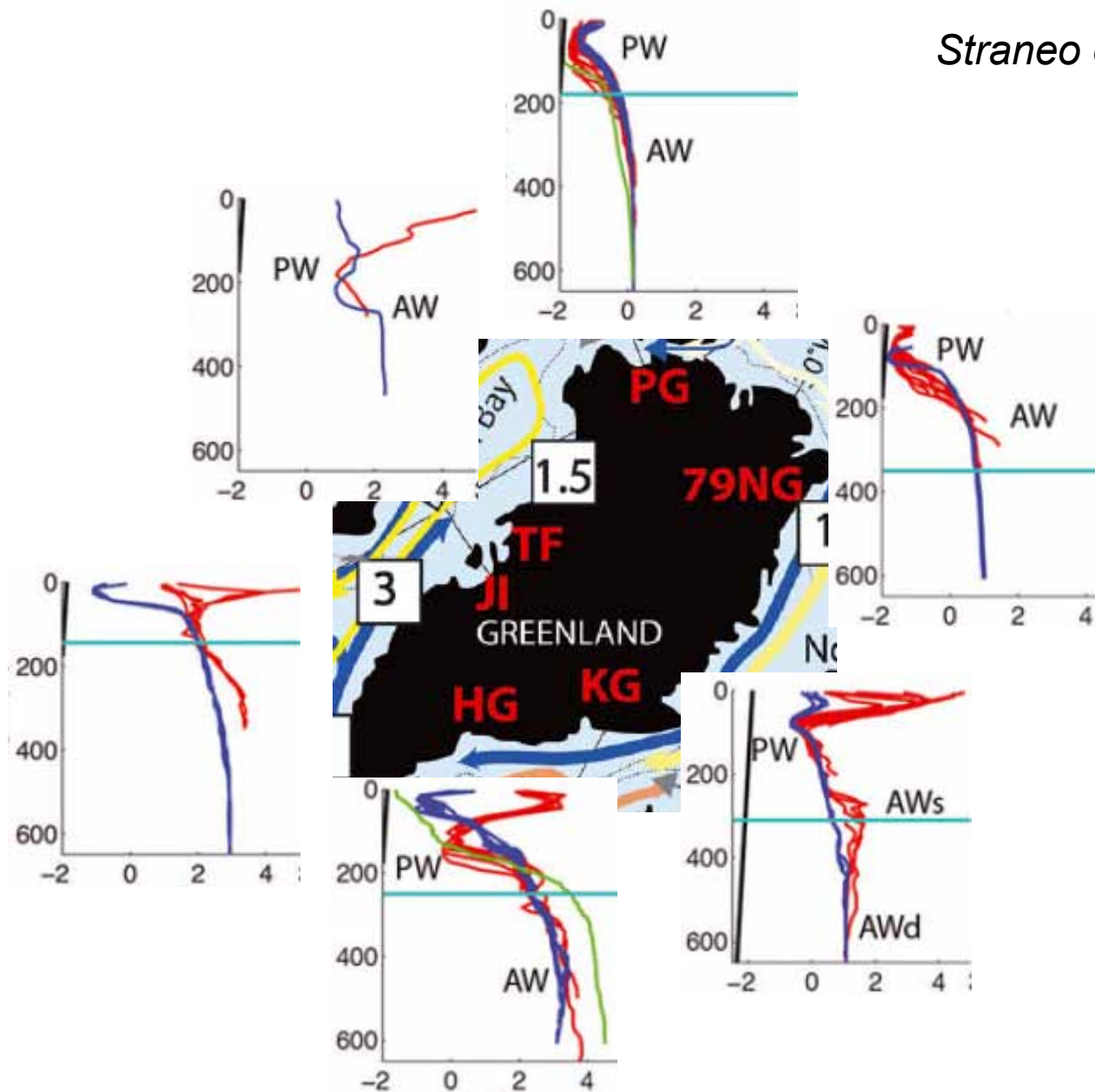
Sea floor temperatures



Pritchard et al

Warm and cold sectors of the Antarctic continental shelf



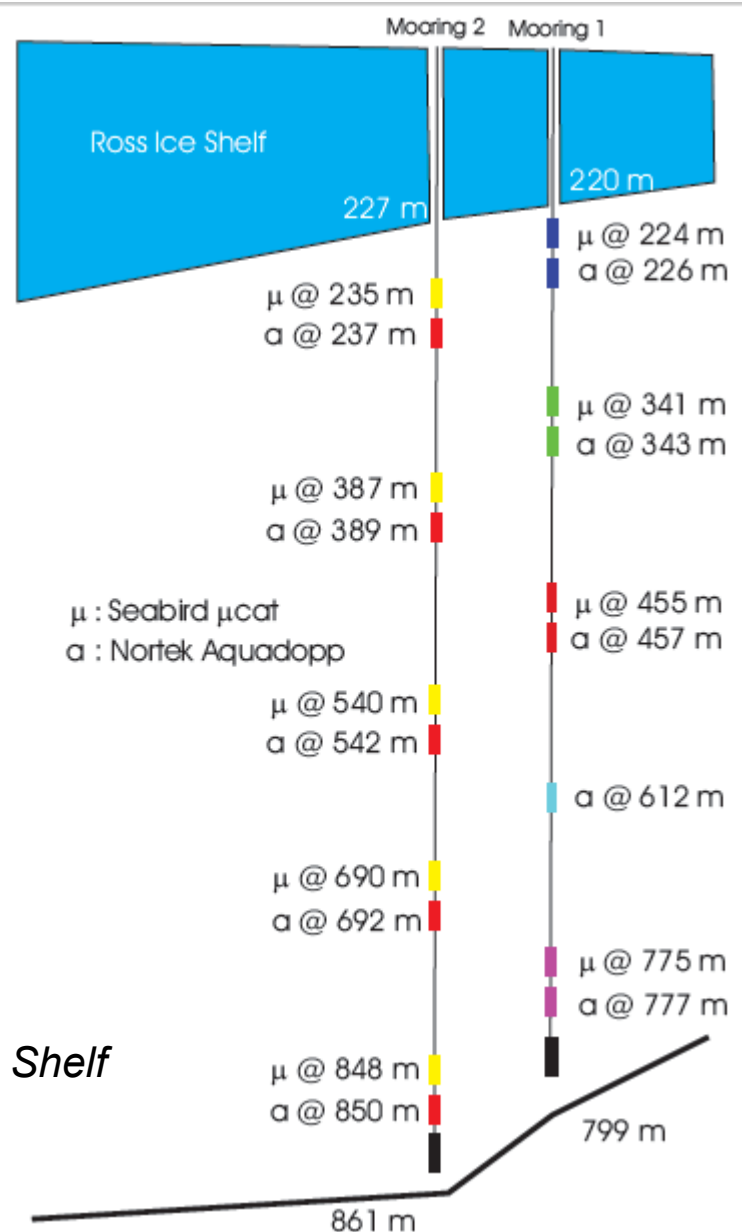


Moored instruments

Typically point measurements
Some short-term ADCP deployments
Generally focussed on big picture

Must be cabled (rarely recovered)
Major benefits in inductive modems

*WHOI and NIWA sites, Ross Ice Shelf
Courtesy Craig Stewart*

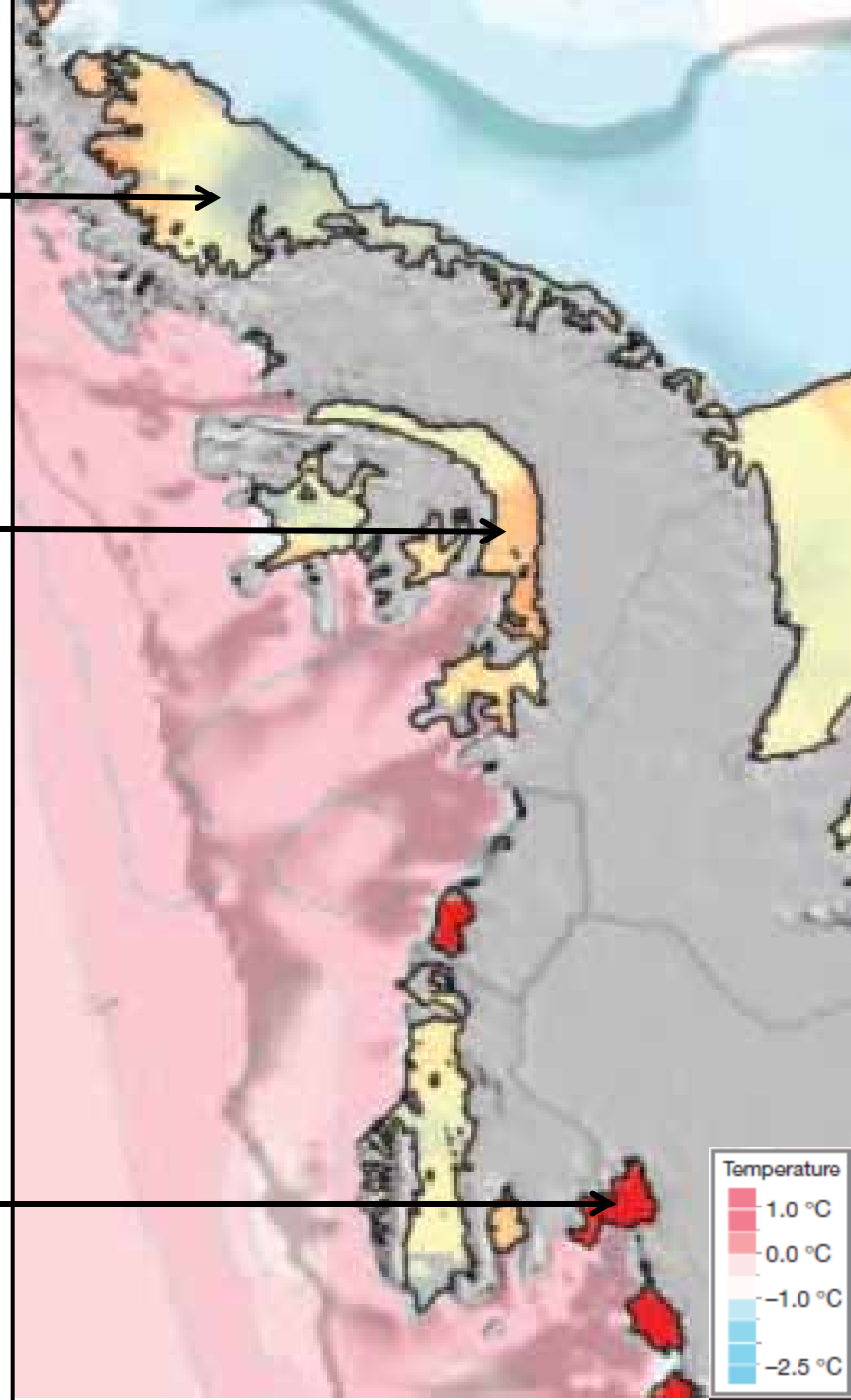


Larsen C Ice Shelf
(2011-12)

George VI Ice Shelf
(2011-12)

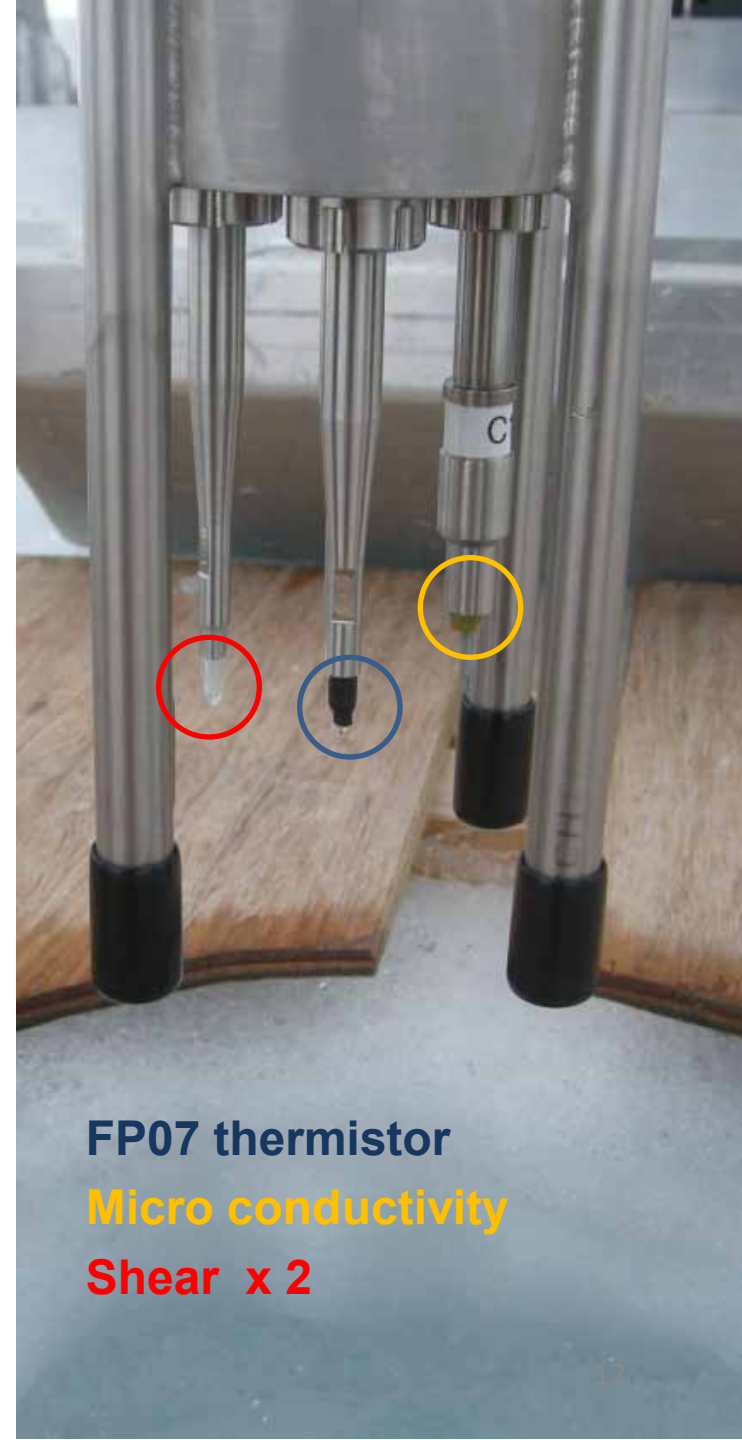
Pine Island Glacier
(2012-13)

GRISO 2013



Microstructure profiling

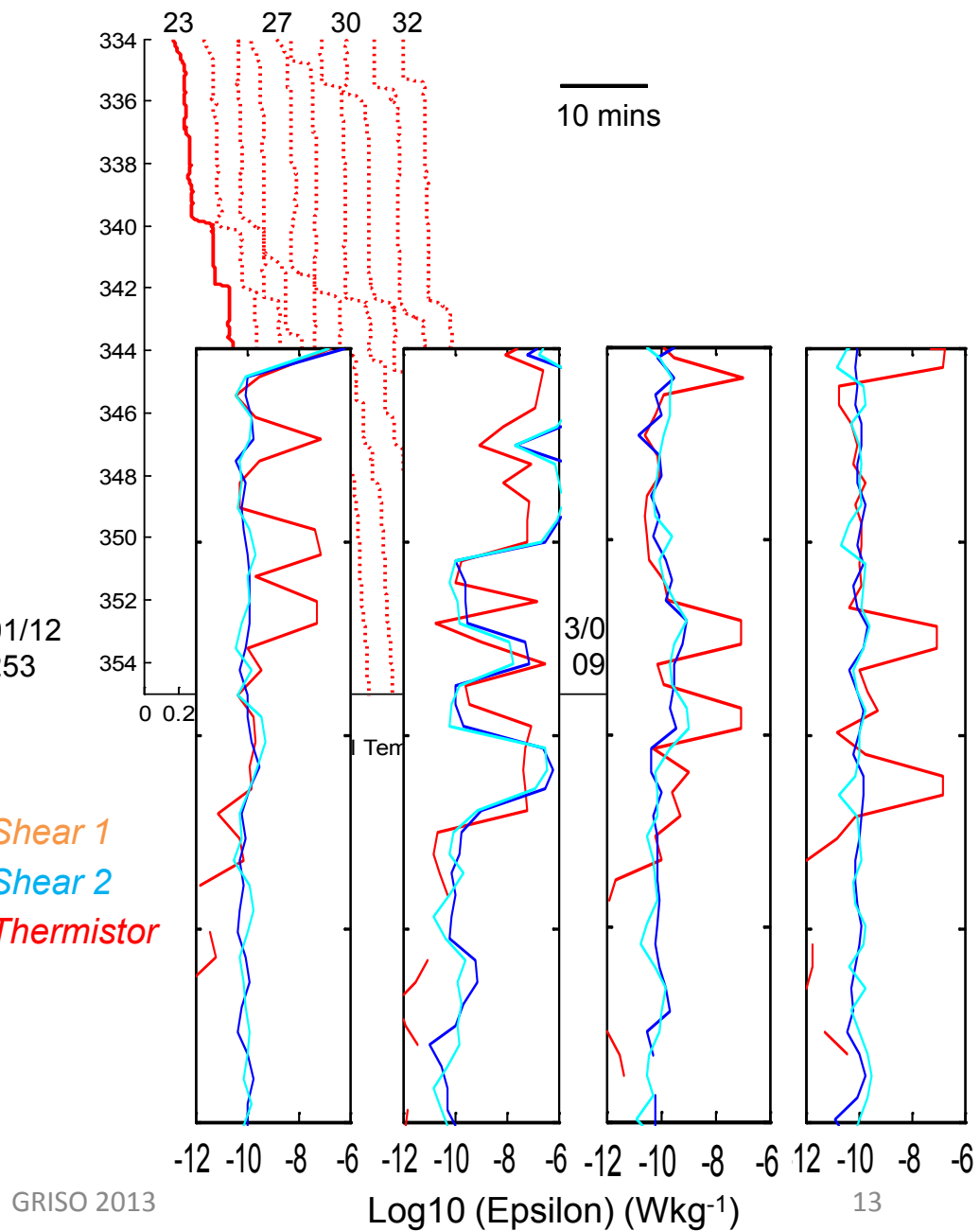
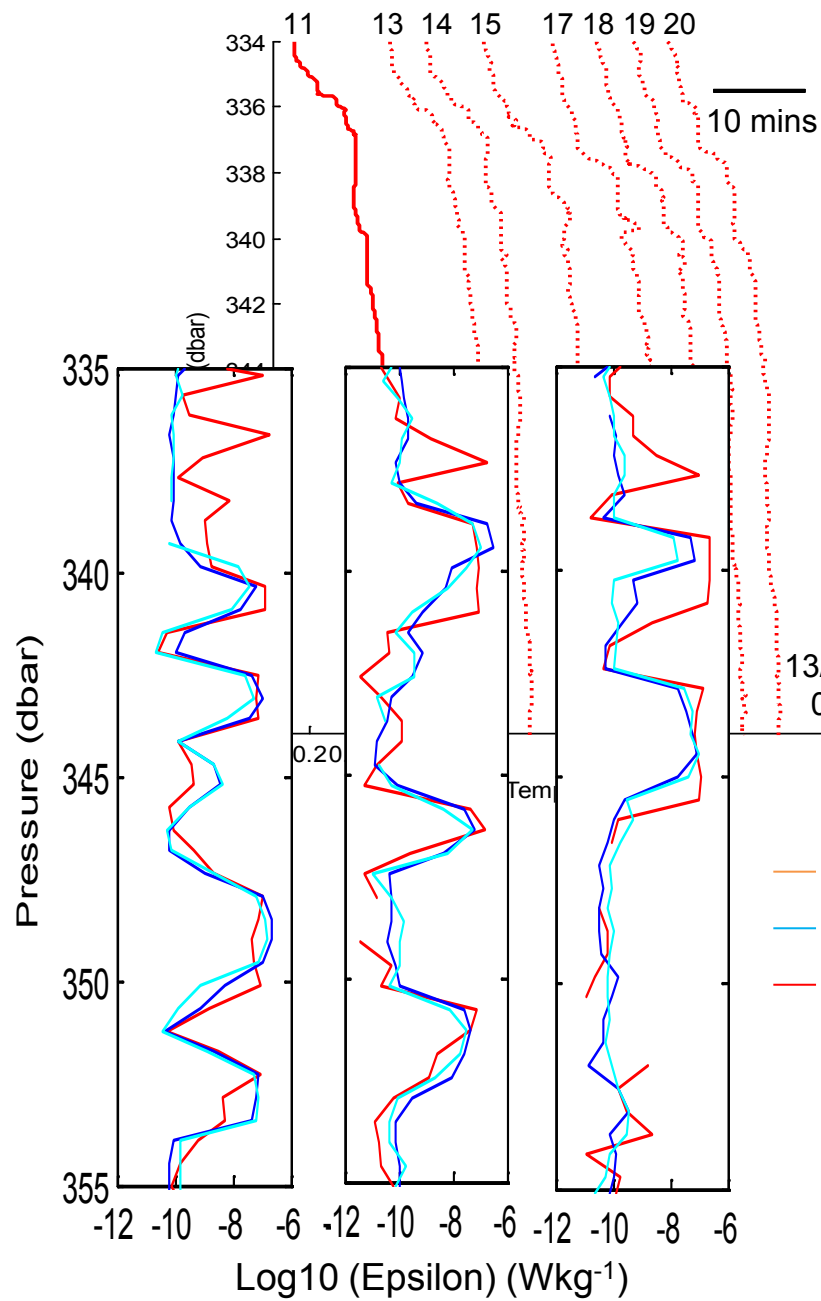
- Rockland VMP 200
- Smoothly moving platform
- Appropriate fall speed
- Sensors at tip sampled at 512 Hz
- Further sensors inside body

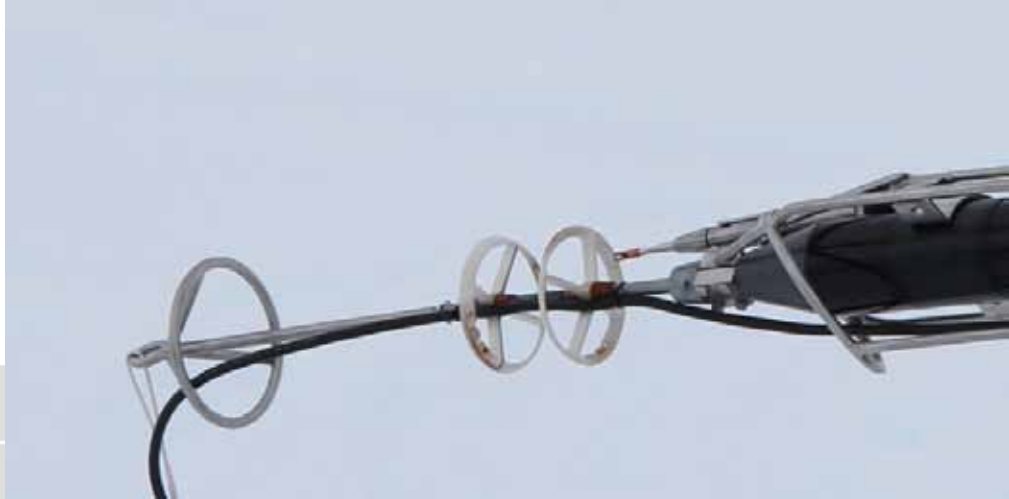
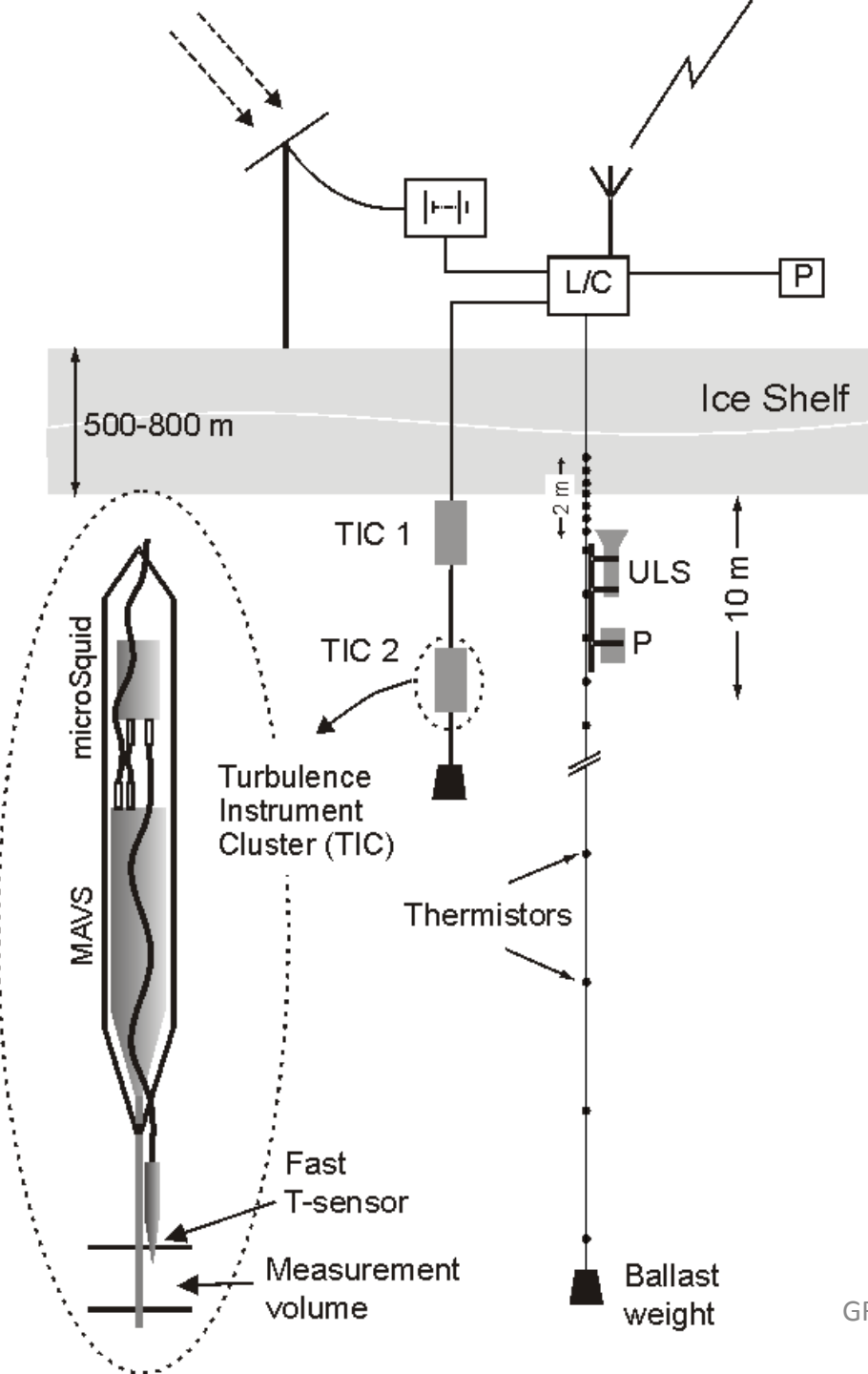


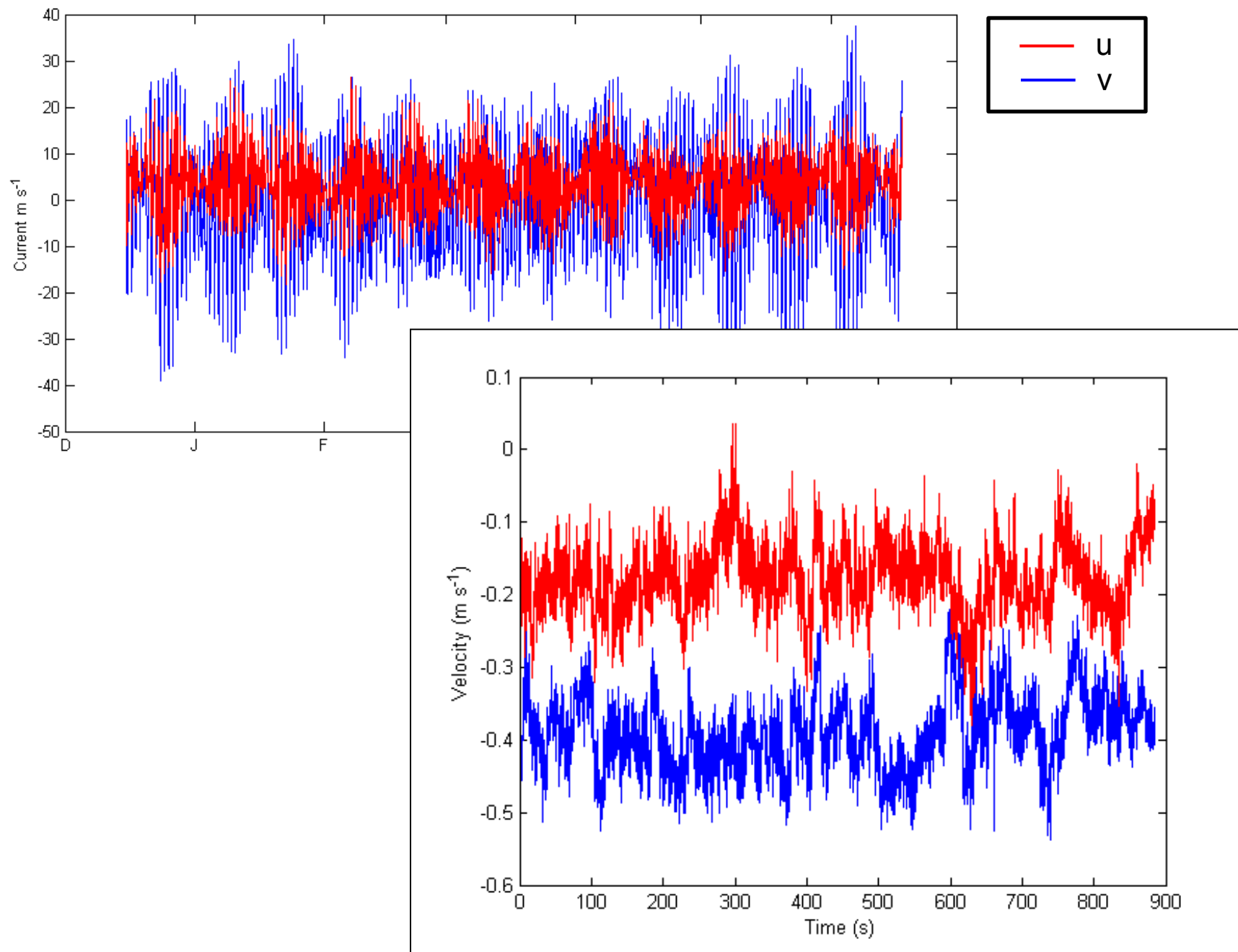
FP07 thermistor

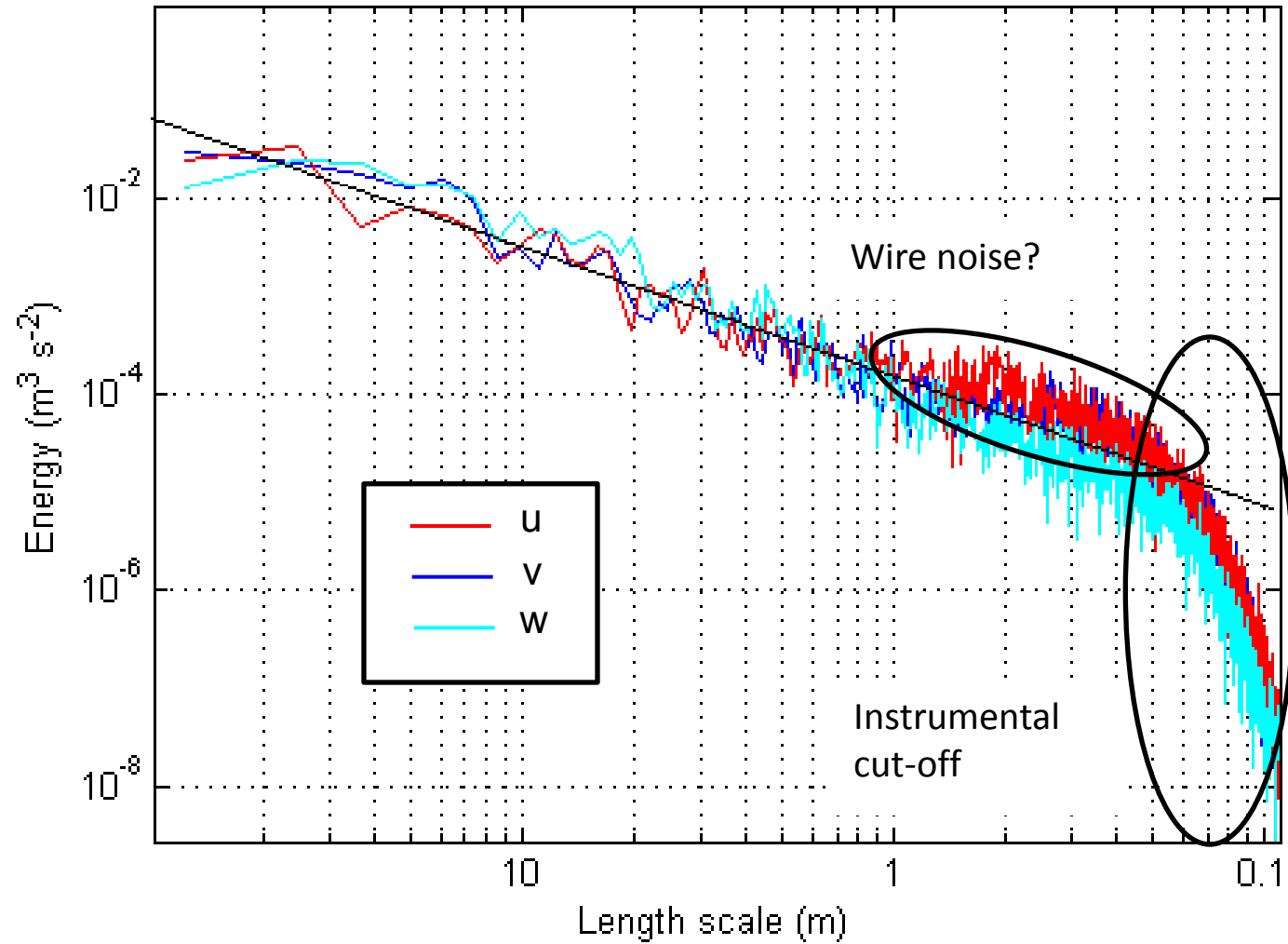
Micro conductivity

Shear x 2

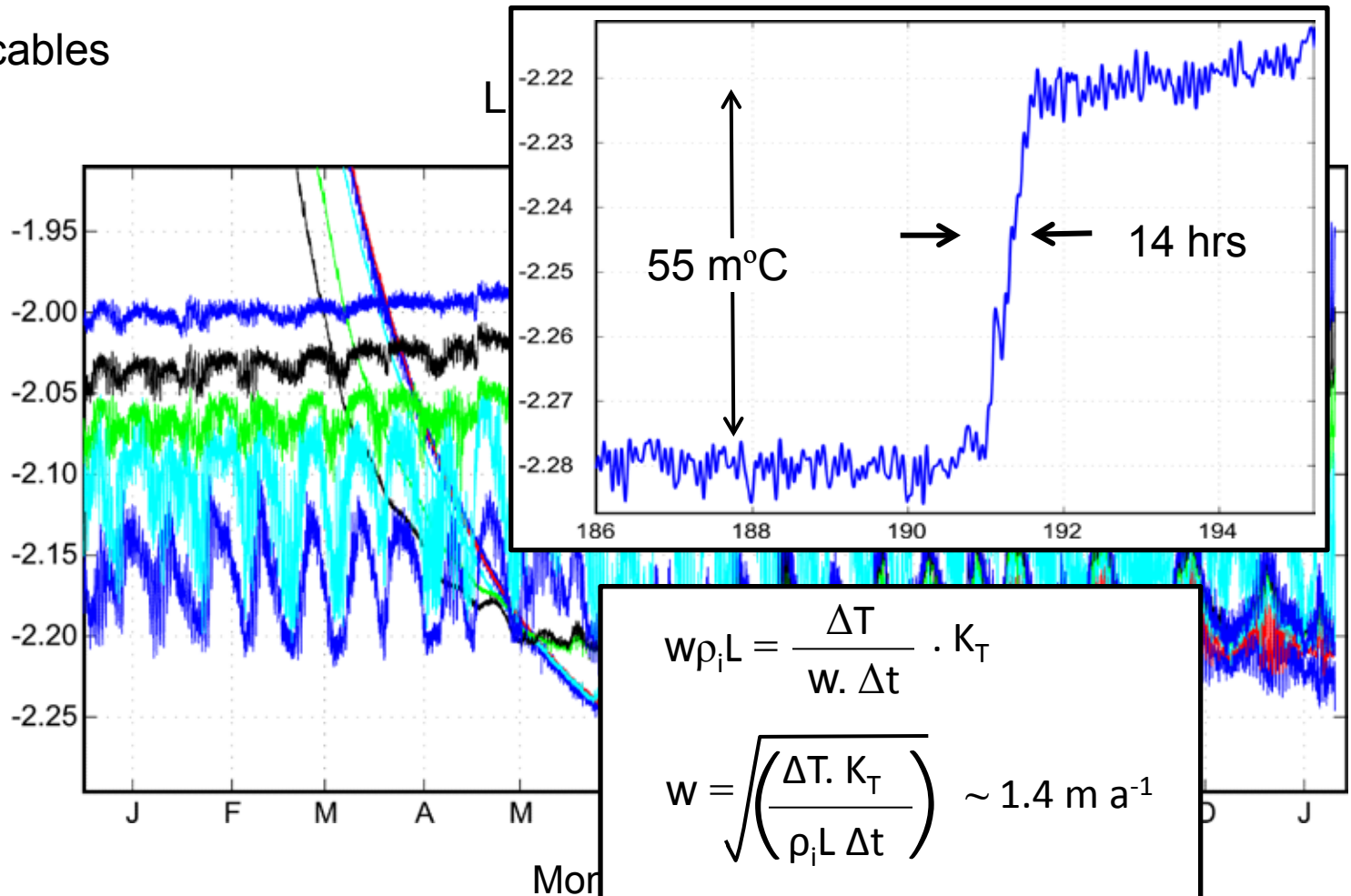
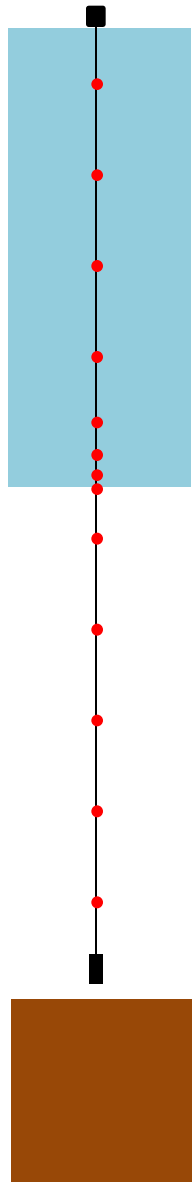








Thermistor cables



$$w\rho_i L = \frac{\Delta T}{w \cdot \Delta t} \cdot K_T$$

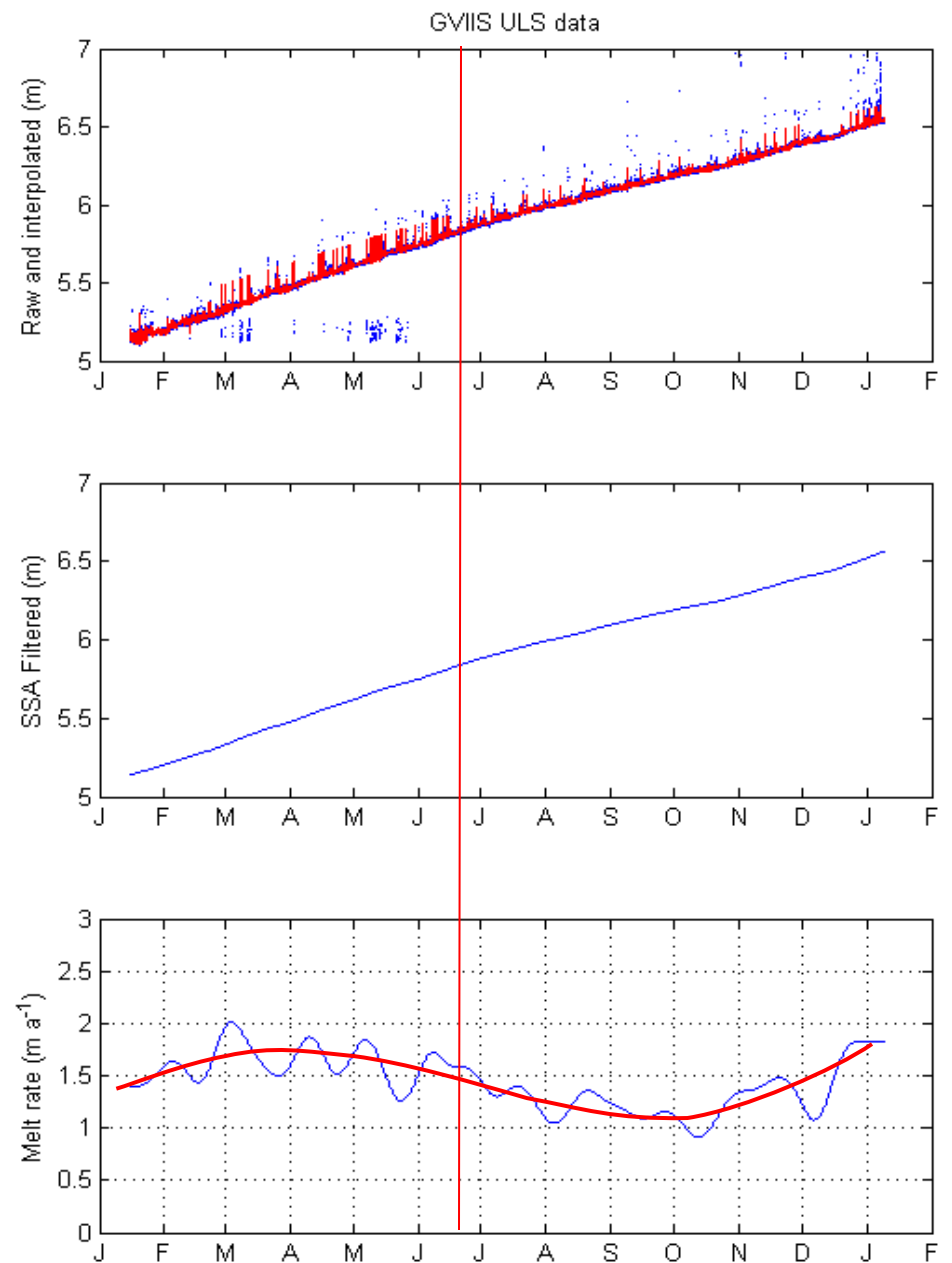
$$w = \sqrt{\left(\frac{\Delta T \cdot K_T}{\rho_i L \Delta t} \right)} \sim 1.4 \text{ m a}^{-1}$$

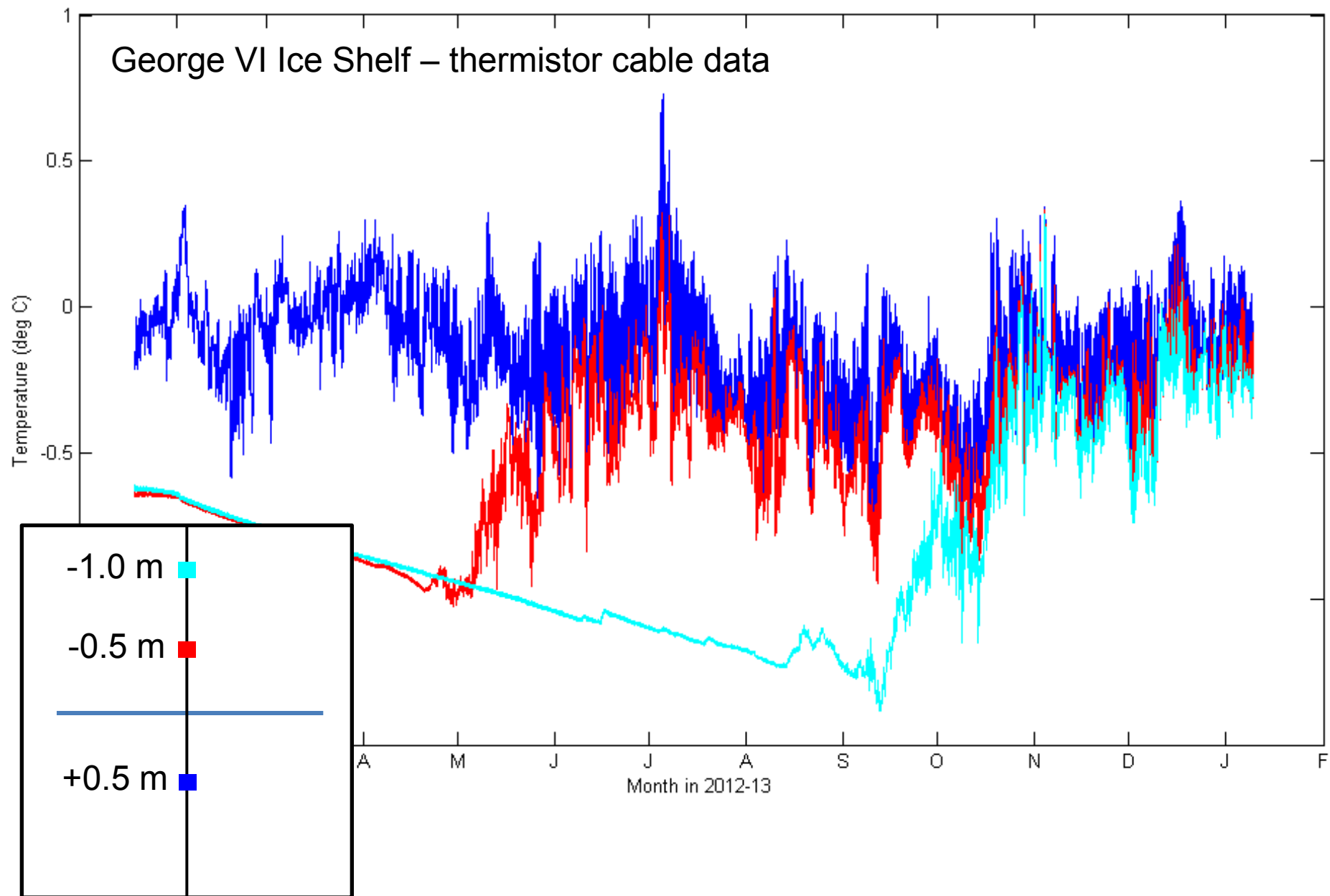
W = melt rate

L = latent heat of fusion

K_T = thermal conductivity

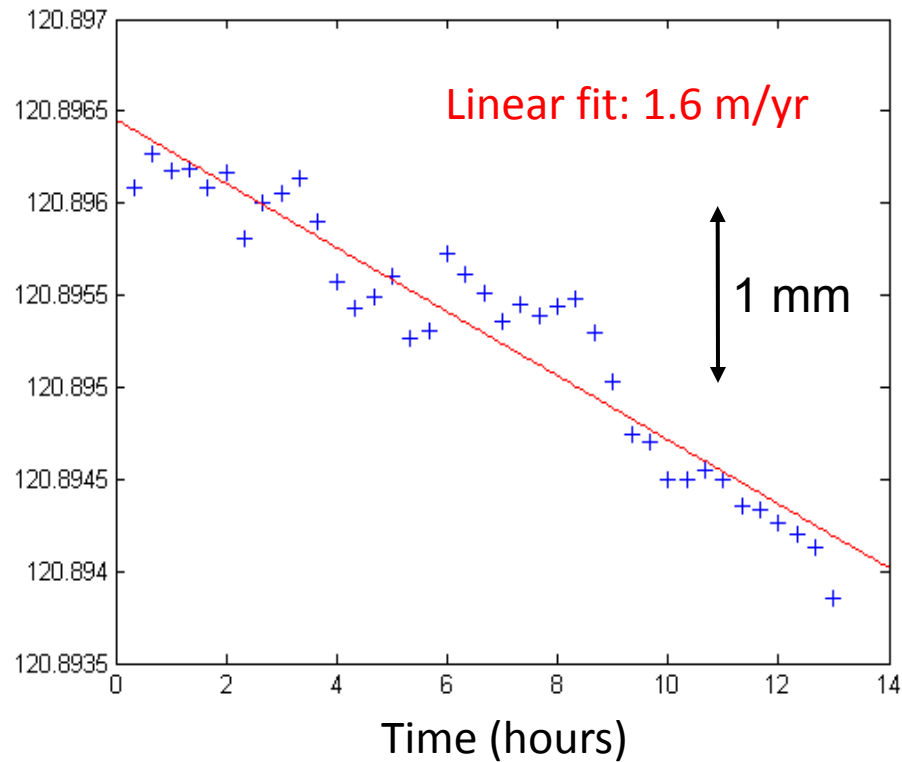
ρ_i = density of ice





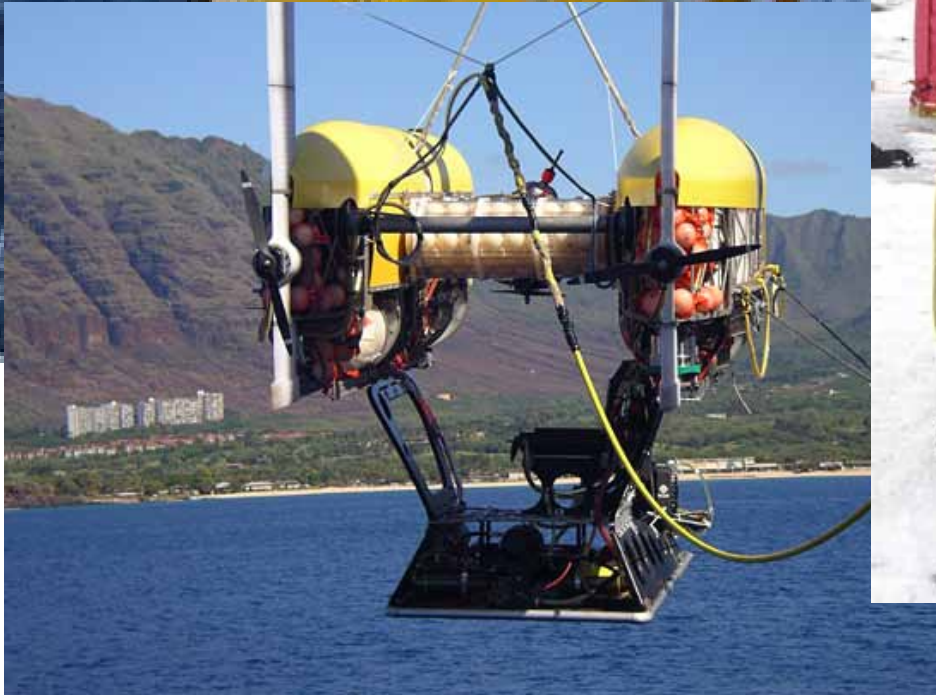
Difference between range to internal and range to ice base (m) (Ross Ice Shelf site)

Precision radar



Measures relative vertical motion of internal and basal reflectors to yield thinning rate of ice shelf, vertical strain rate of ice column and therefore basal melt rate.

AUVs and ROVs



Nereus derivative

SCINI



Summary

- Greenland sub-ice shelf conditions are similar to the warm shelf sector of Antarctica;
- Very few sub-ice shelf boundary layer experiments to date;
- Appropriate instrumentation is now available;
- Free-fall microstructure profiling can work through ice boreholes;
- ROVs hold promise for the future;
- Thermistor cables are highly cost effective;
- Melt-rate time series from precision radar.