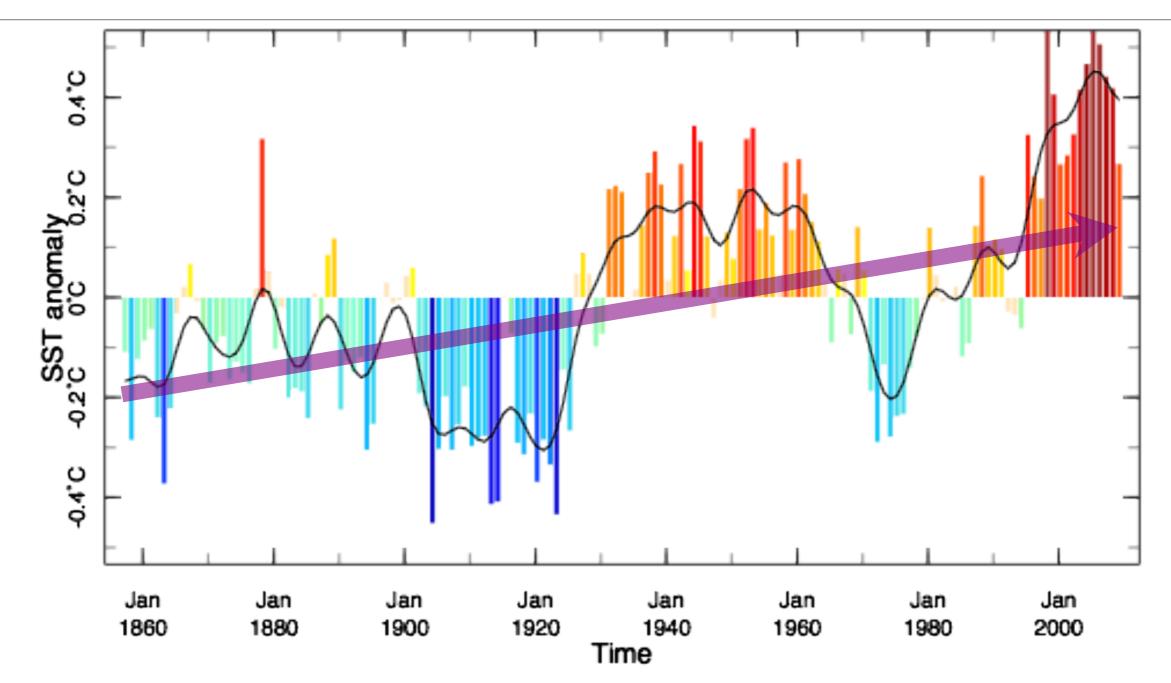
# Climatic impacts of long-term North Atlantic SST Variability: A brief review

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2013 U.S. AMOC/U.K. Rapid International Science Meeting - AMOC Variability: Dynamics and Impacts Baltimore Maryland, July 2013 •Further study is required to understand the connections between AMOC/North Atlantic SST and climate variability elsewhere, the physical mechanisms of these teleconnections, and the related impacts on humans and ecosystems.

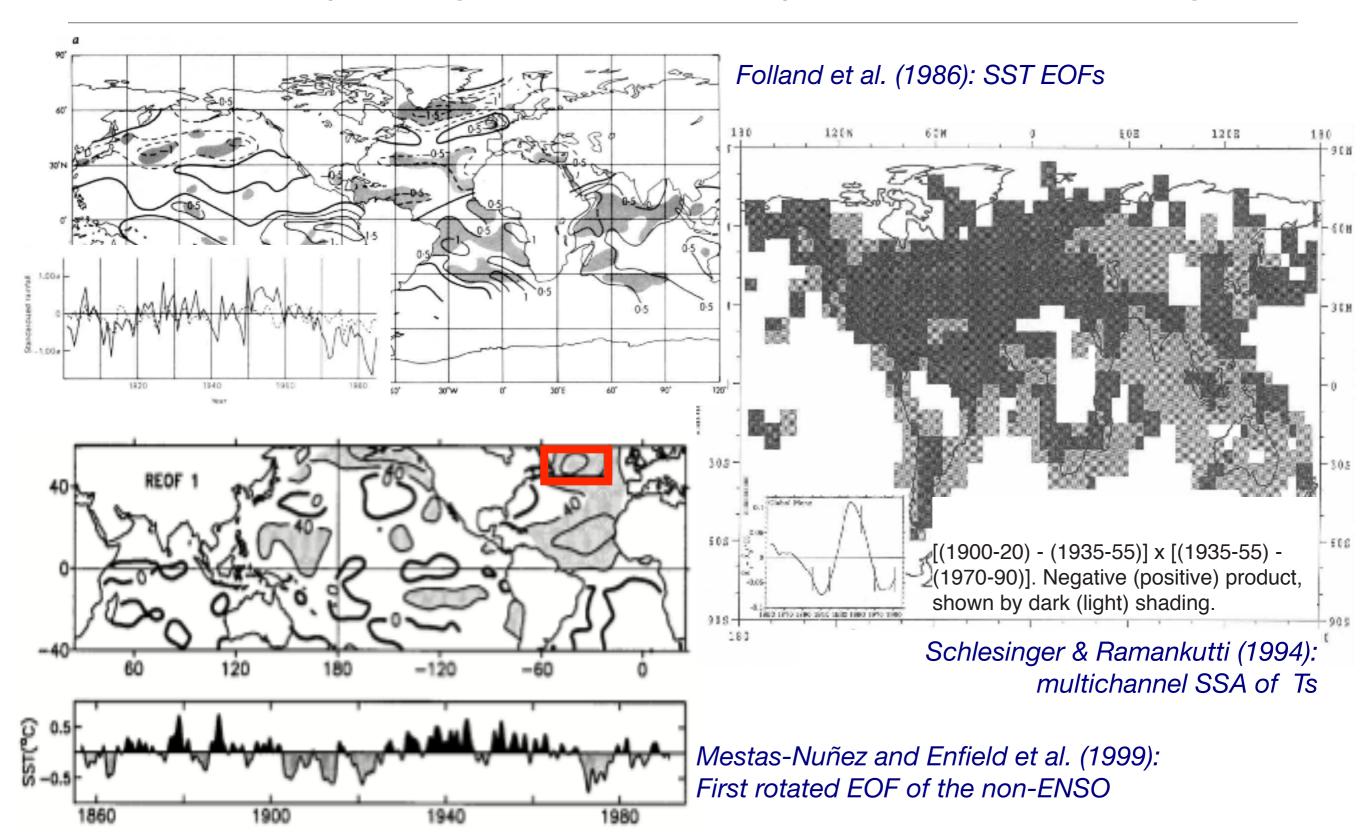
### Atlantic multidecadal variability

### North Atlantic observed SST variability

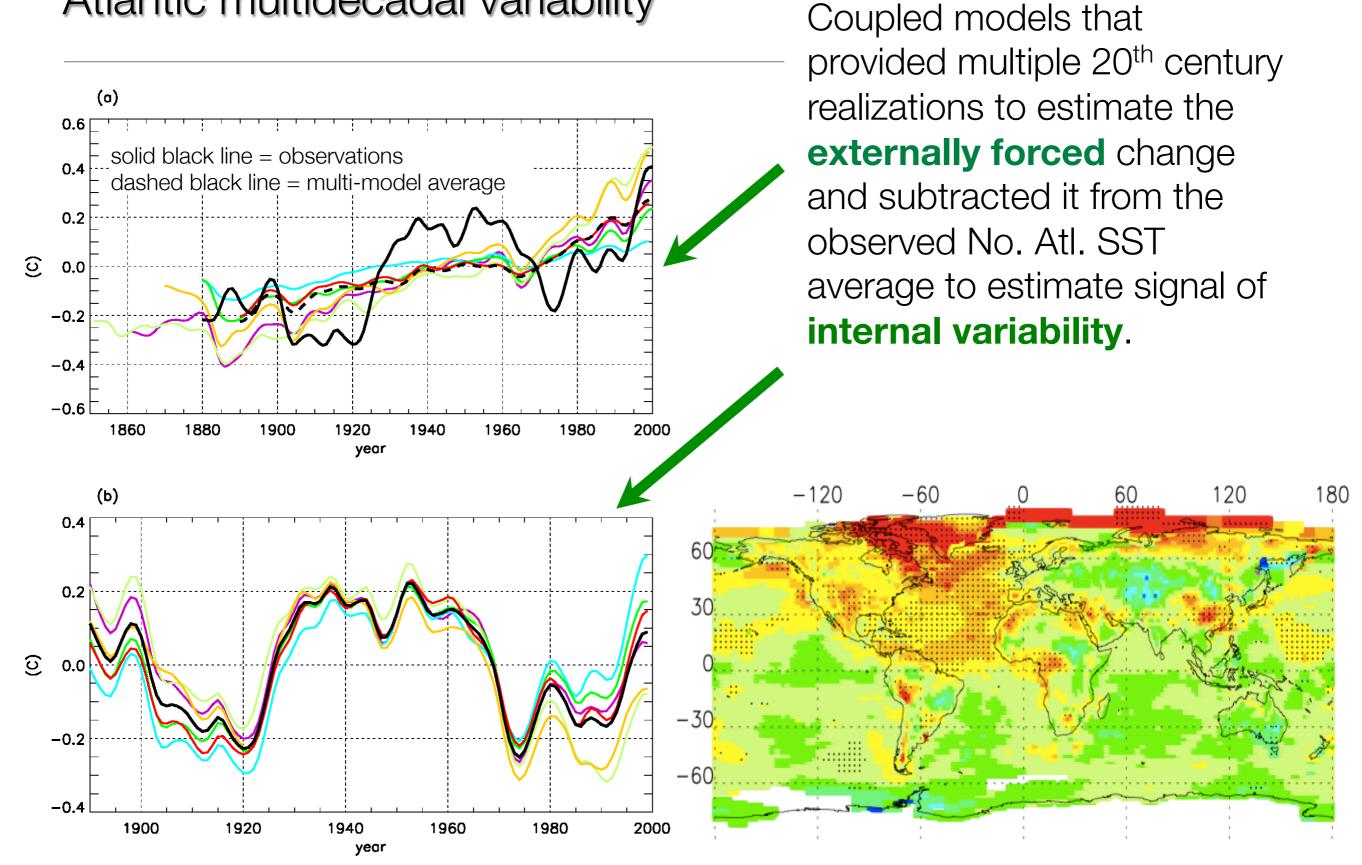


North Atl. (eq. to 60°N) averaged SST anomaly, 1856-2009: Color bars show the annually averaged SST anomaly, the solid line is the 10 year lowpass anomaly.

# Atlantic Multidecadal Oscillation: a distinct pattern of low-frequency surface temperature variability



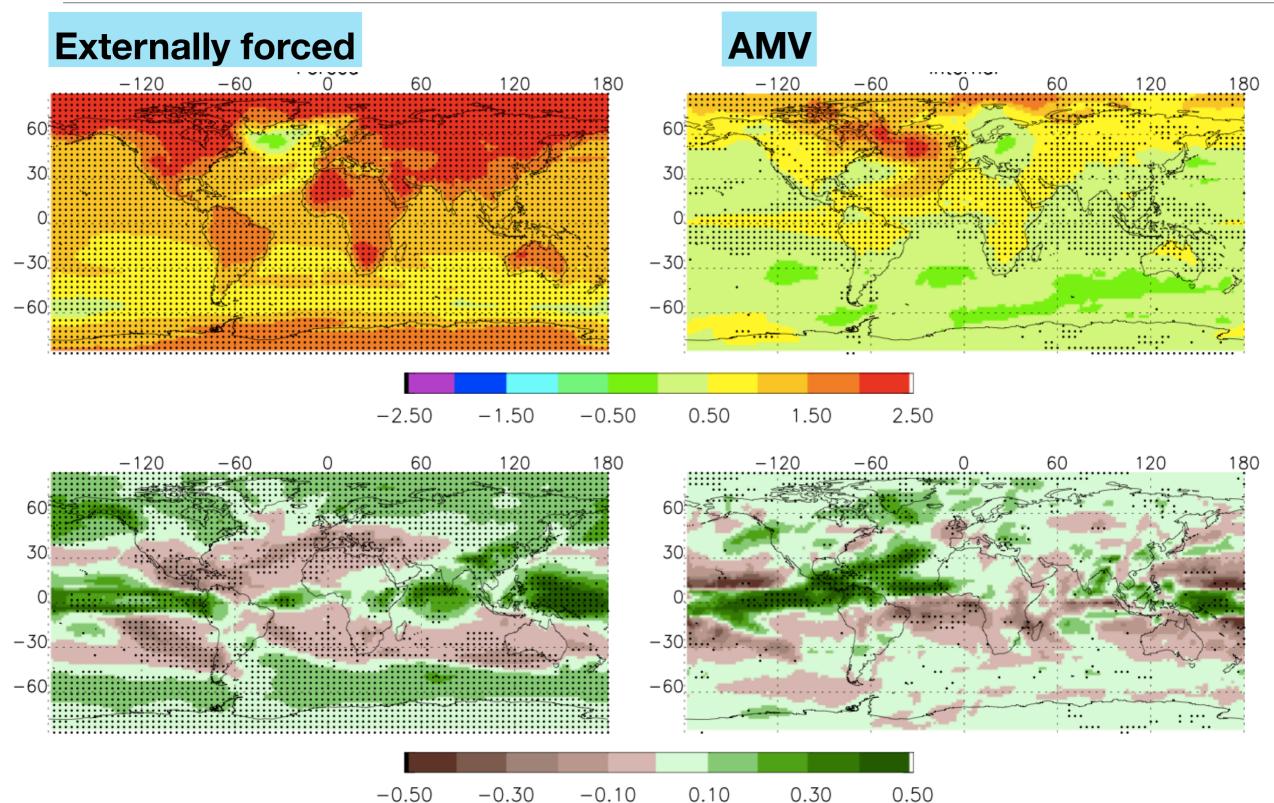
# Separating forced and internal Atlantic multidecadal variability



*Ting et al. (2009)* applied S/N

maximizing PCA to 9 CMIP3

#### CMIP3 20th century variability Ting et al. [2009; 2011]



-0.50-0.30-0.100.10 0.30

### Ocean role in AMV - link to AMOC

### Bjerknes (1964) in Atlantic airsea interactions ...

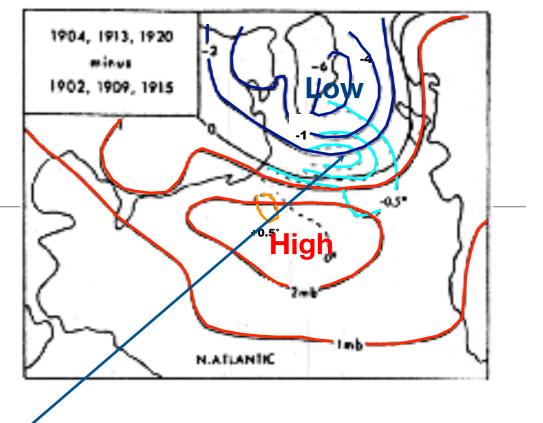
... pointed out that there are two patterns of Atmo/Ocean variability in the North Atlantic fast and slow. The slow (multi-year) pattern displays spatially uniform SST variations and a different ocean-atmosphere relationship than the fast one.

Interannual variability is driven by atmospheric wind fluctuations: strong westerlies ⇒ cold subpolar gyre weak westerlies ⇒ warm subpolar gyre

Multi-yeat variations are driven by changes in ocean heat transport:

Uniform warming of No. Atlantic  $\Rightarrow$  weak subtropical anticyclone (& weak westerlies in the north)

#### Figure from Bryan and Stouffer (1991) after Bjerknes (1964)



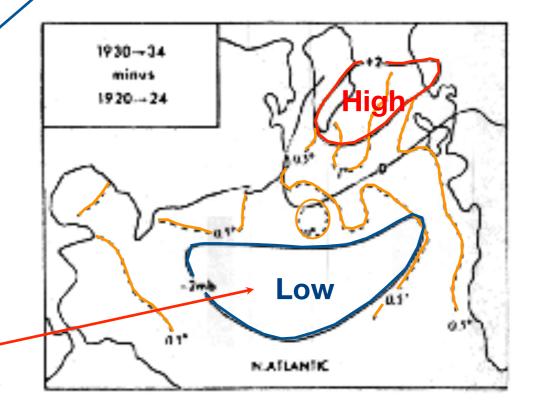
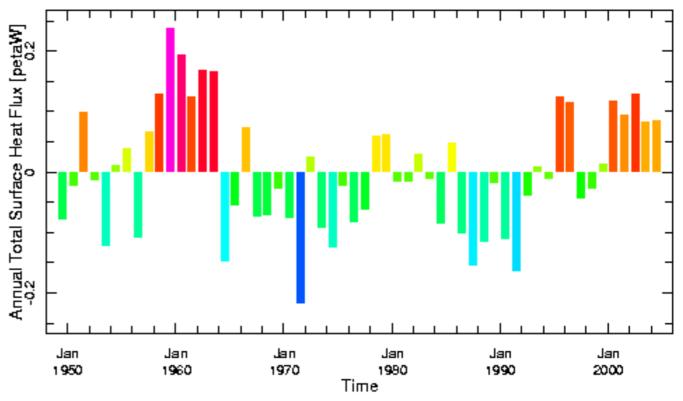


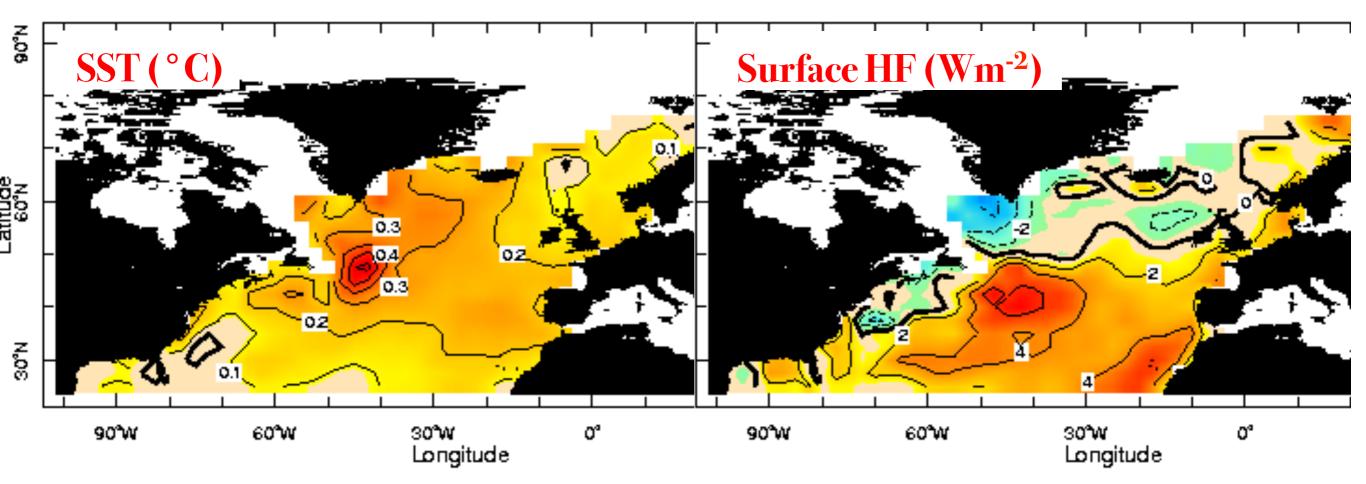
Fig. 4. SST and surface pressure anomalies from Bjerknes (1964). (a) The difference between isolated years of high and years of low North Atlantic Index, where the North Atlantic Index is defined as the difference surface pressure between the Azores and Iceland. (b) The difference between an average over a period of persistent high index and a period of persistent low index (1920-1924 minus 1930-1934).

# O/A surface heat flux and the AMO

NCEP reanalysis O/A fluxes support Bjerknes hypothesis: On multidecadal timescales the atmosphere is heated by a warm ocean surface possibly forced by the convergence of ocean heat transport (AMOC?)

#### Surface upward O/A HF from NCEP-NCAR Reanalysis

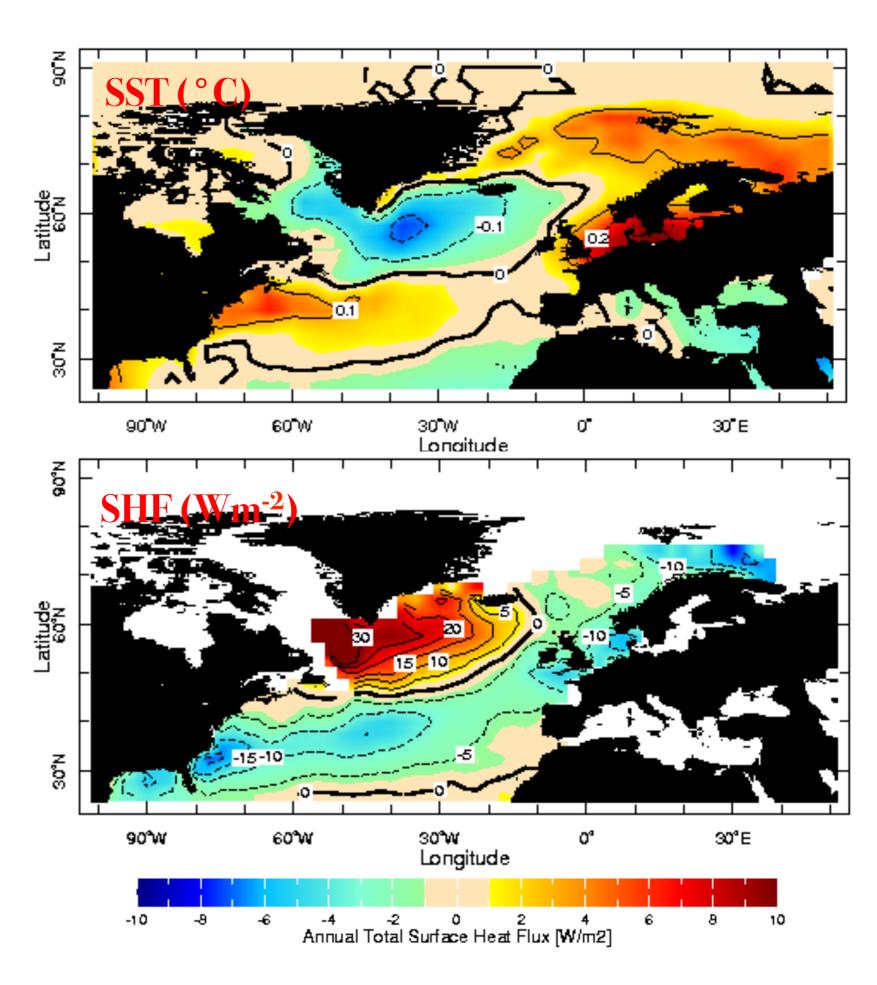




# Inter-annual variability

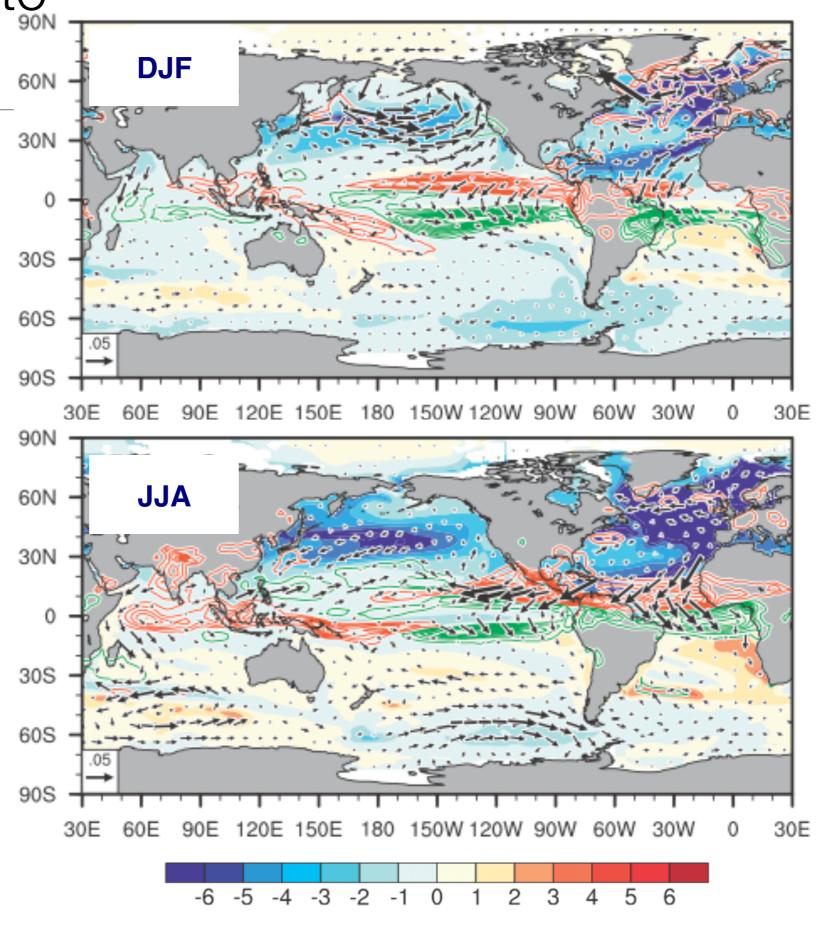
THE OCEAN IS COOLED & WARMED BY ATMOSPHERIC CONTROLLED SURFACE HEAT FLUXES & OHT BY EKMAN TRANSPORT

Regression on the NAO index show that surface heat fluxes forced by the atmospheric wind variations are linked with SST warming where heat loss to the atmosphere is reduced and cooling where heat loss is increased.

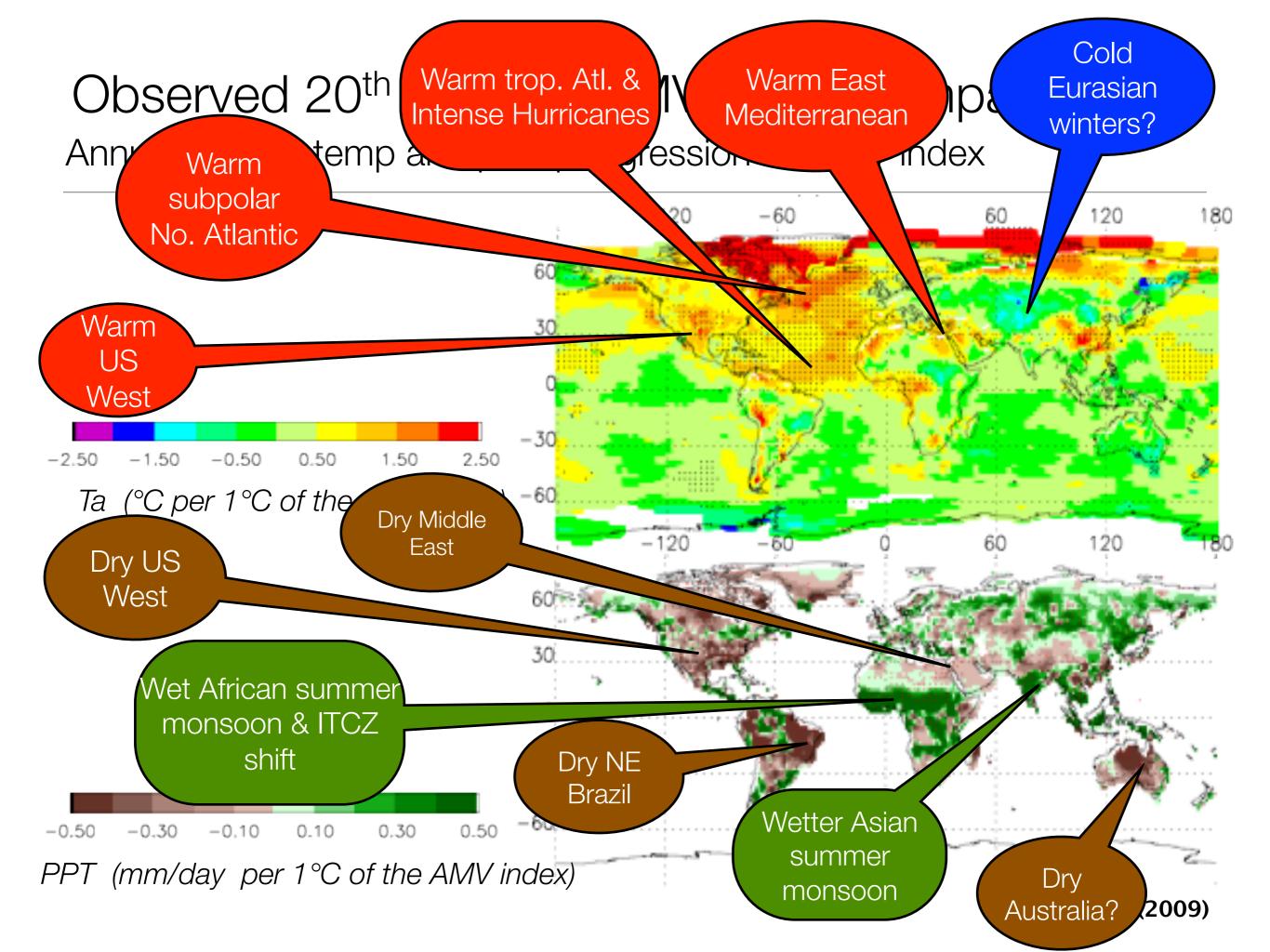


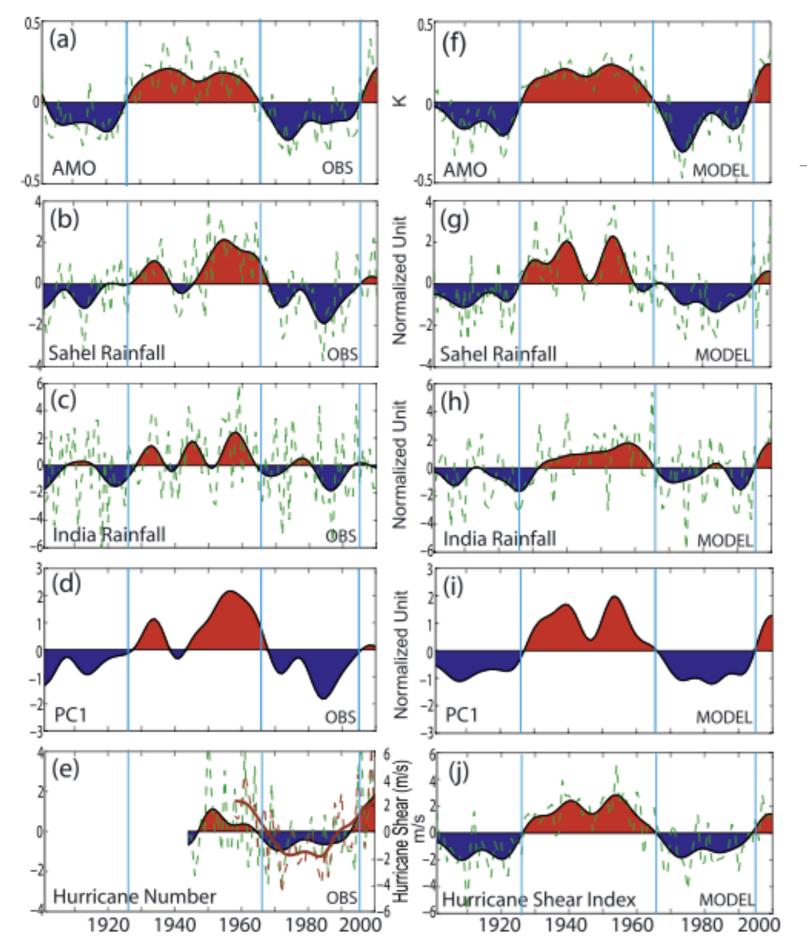
# Surface response to AMOC collapse

- Typical Coupled GCM (here GFDL CM2.1) response to a 1-Sv AMOC water-hosing experiment: SST (color), surface wind stress (vectors, N/m2), and precipitation (green contours > 1.0 mm/day and orange contours < -1.0 mm/day intv. 1.0 mm/day).
- Okumura et. al. (2009)



### Climatic impacts of AMV





# Modeling AMV impacts

Time series of Sahel and West +Central India rainfall, and of number of Atlantic hurricanes, exhibit AMV time scales.

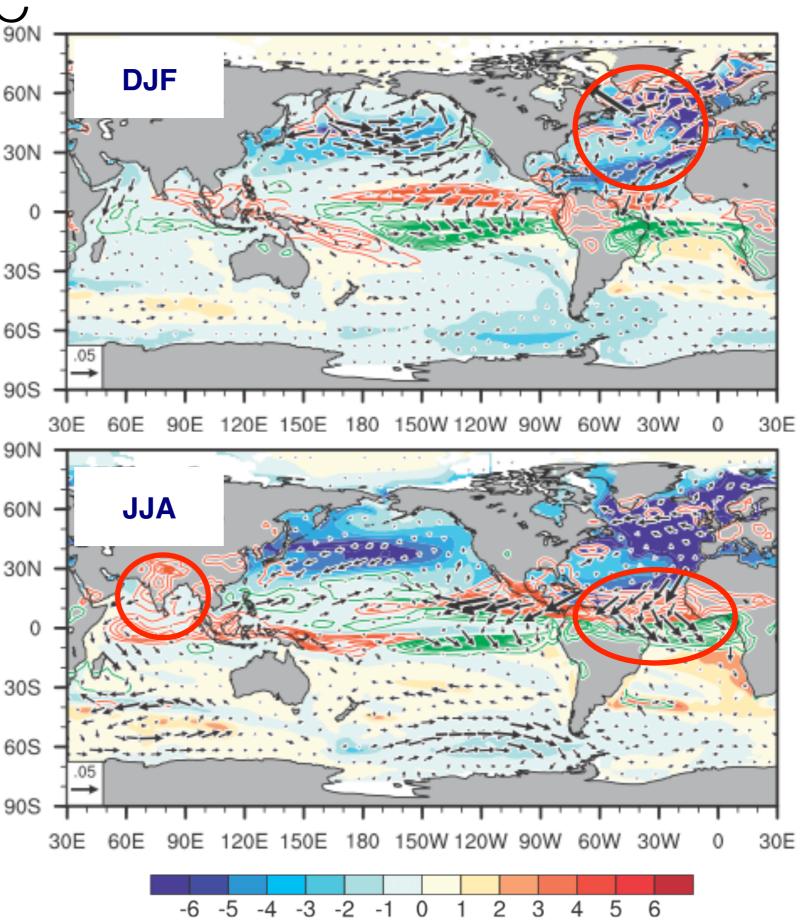
A coupled GCM with imposed AMV SST change indicates these changes are caused by AMV.

Zhang and Delworth (2006)

## Response to AMOC collapse

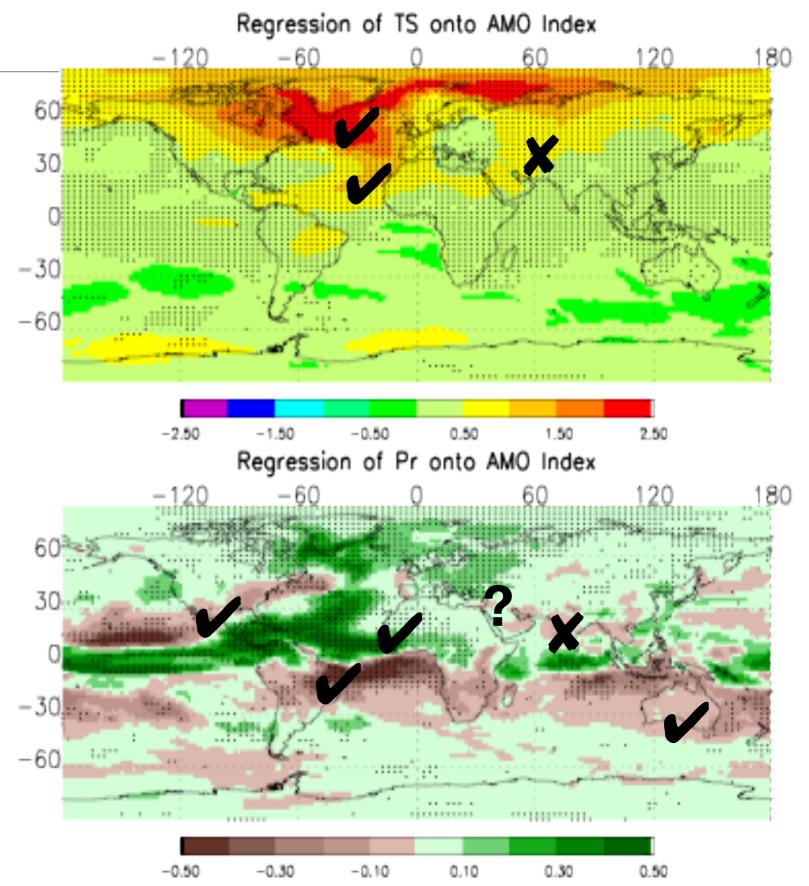
- Typical Coupled GCM (here GFDL CM2.1) response to

   a 1-Sv AMOC water-hosing 30S
   experiment confirms 60S
   AMOC-AMV link and the climatic impacts.
- SST (°C, color), surface wind stress (vectors, N/ m2), and precipitation (green contours > 1.0 mm/ day and orange contours < -1.0 mm/day intv. 1.0 mm/ day).
- Okumura et. al. (2009)

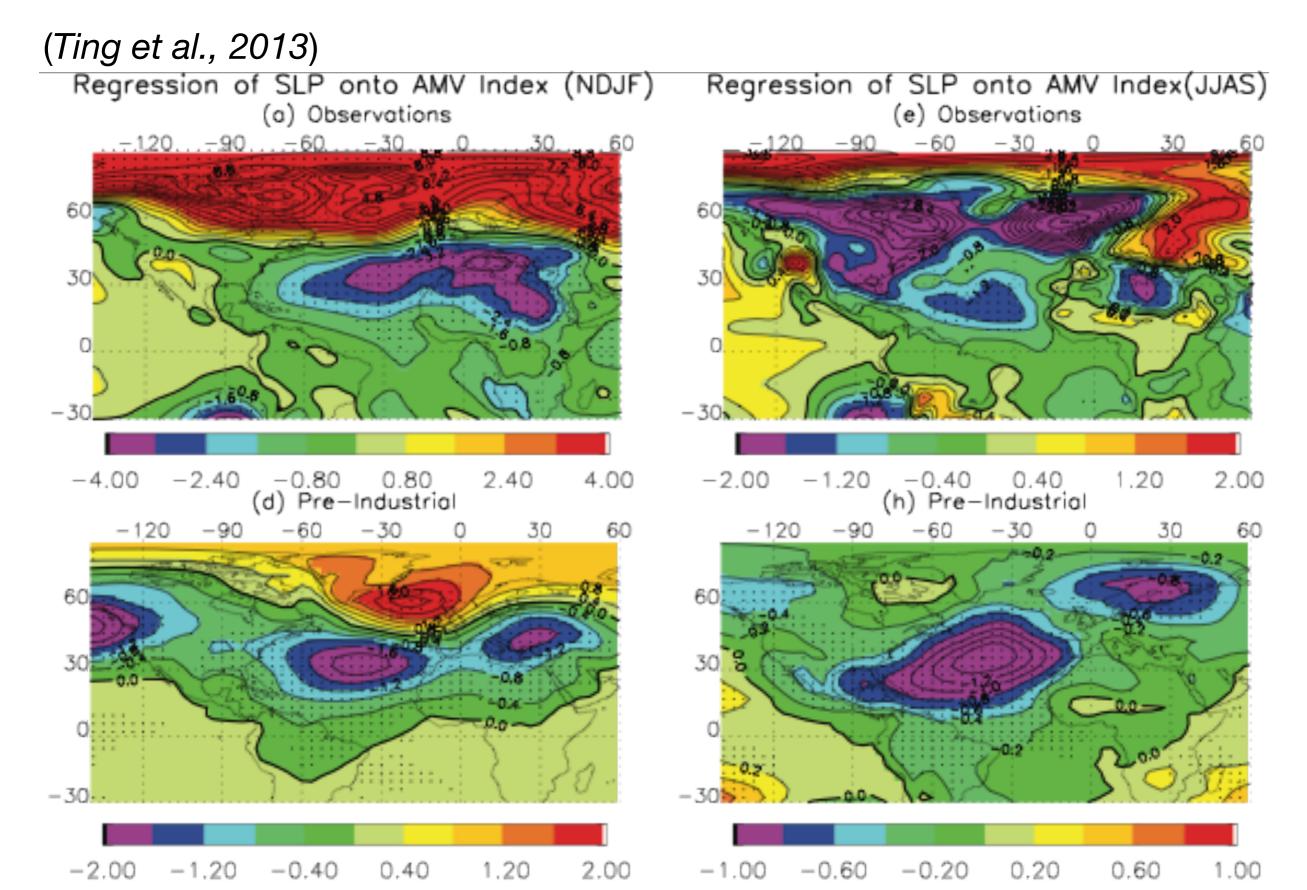


# CMIP5 multi-model AMV simulation

- There is a broad model agreement regarding the global surface temperature pattern associated with AMV.
- There is no model agreement on timescale (for most model timescale << observed).</li>
- The similarity in the surface temperature pattern in the North Atlantic indicates a common underlying atmospheric or oceanic or coupled mechanism (AMOC related?)



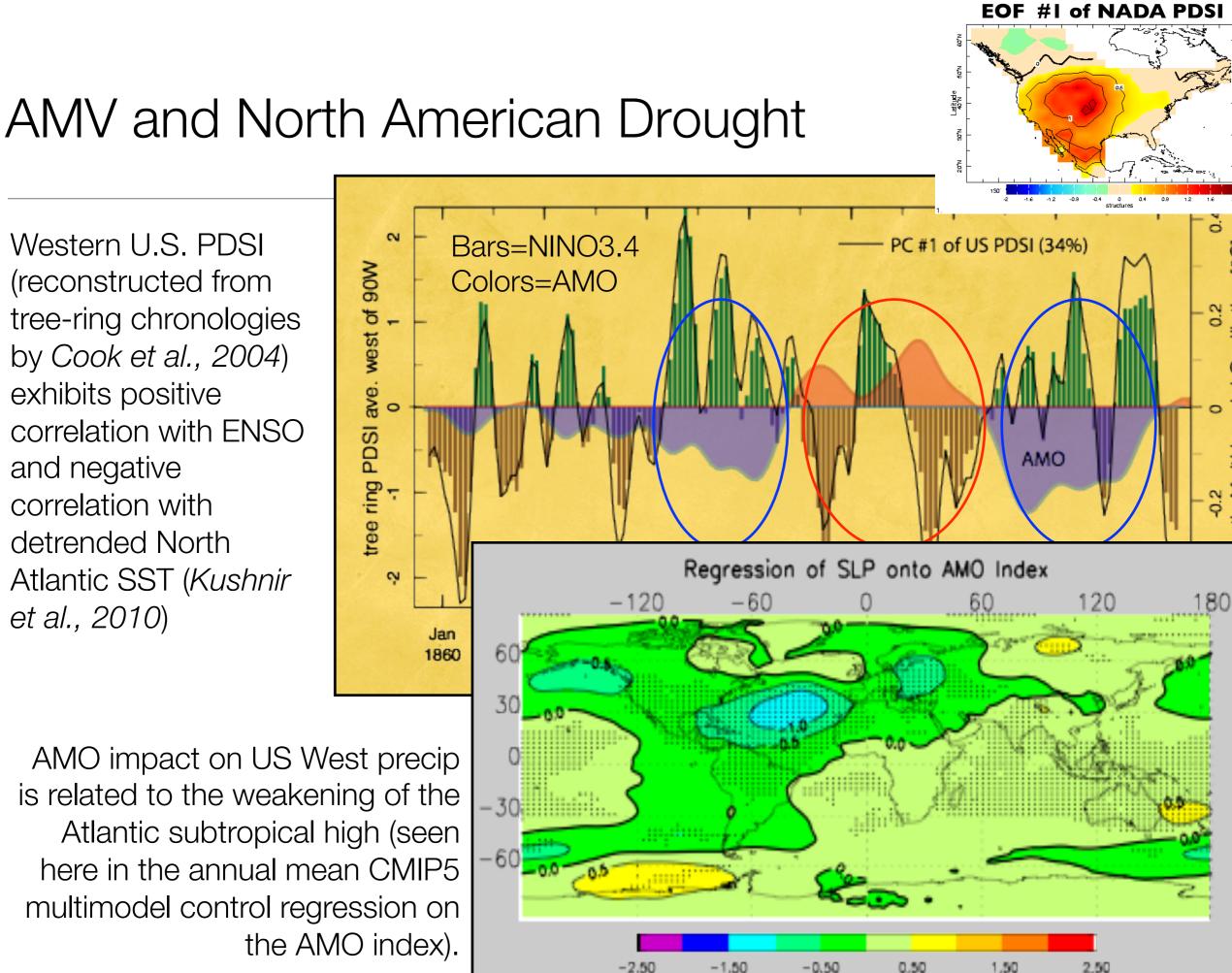
### Atmospheric circulation changes



### Mechanisms

Western U.S. PDSI (reconstructed from tree-ring chronologies by Cook et al., 2004) exhibits positive correlation with ENSO and negative correlation with detrended North Atlantic SST (Kushnir et al., 2010)

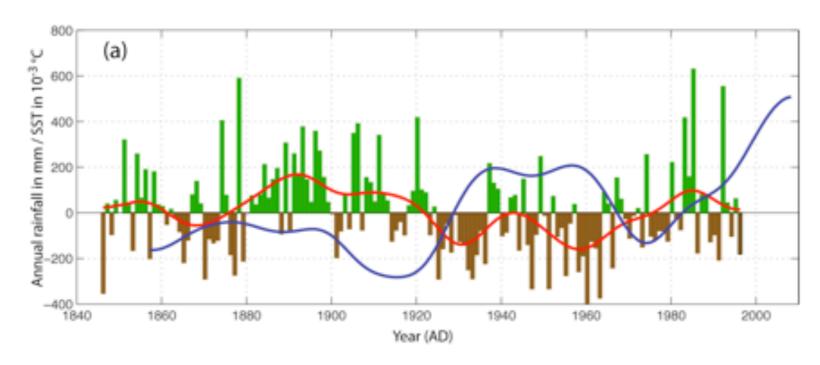
AMO impact on US West precip is related to the weakening of the Atlantic subtropical high (seen here in the annual mean CMIP5 multimodel control regression on the AMO index).



ntic Multidecadal Oscillation

# Levant rainfall responds to AMV

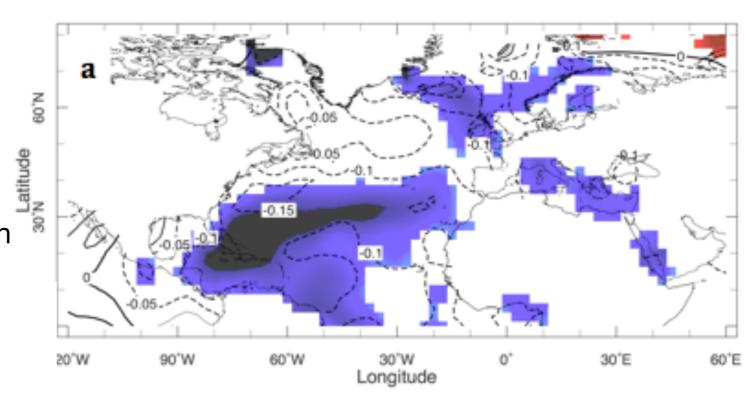
#### Kushnir and Stein (2010)



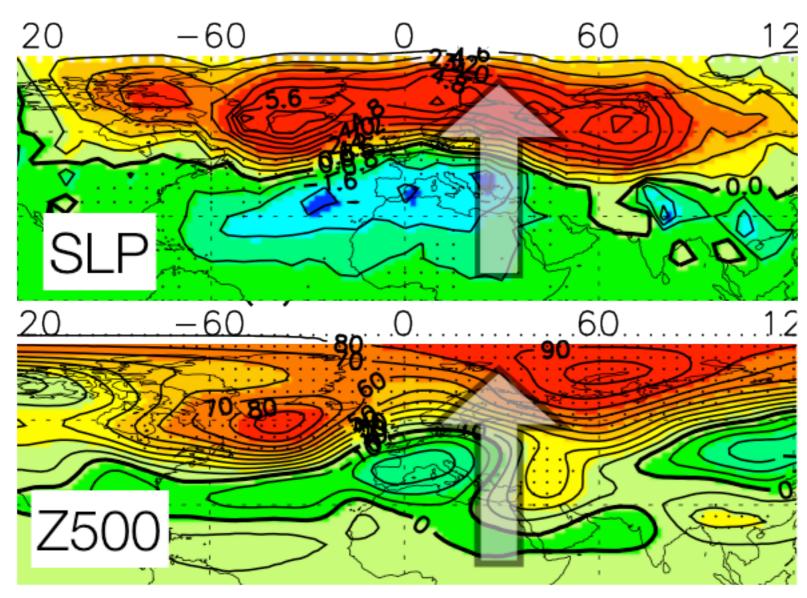
Jerusalem annual (October-September) prcp anomaly: annual (color bars in mm) and low-pass filtered (red line); low pass filtered SST anomalies averaged over the extratropical North Atlantic (30°N to 70°N) in units of 10<sup>-3</sup> °C (in blue). Anomalies are wrt 1961-1990.

The regressions of SST on Jerusalem precipitation. Colors indicate regions where the values are significant at the 10% level (non-directional). Data filtered with a 10-yr low-pass. Units are °C per one standard deviation of the filtered seasonal precipitation time series (~110 mm/yr).

#### Cold No. Atl. ⇒ Wetter Levant



## AMV influences the west-east Mediterranean seesaw

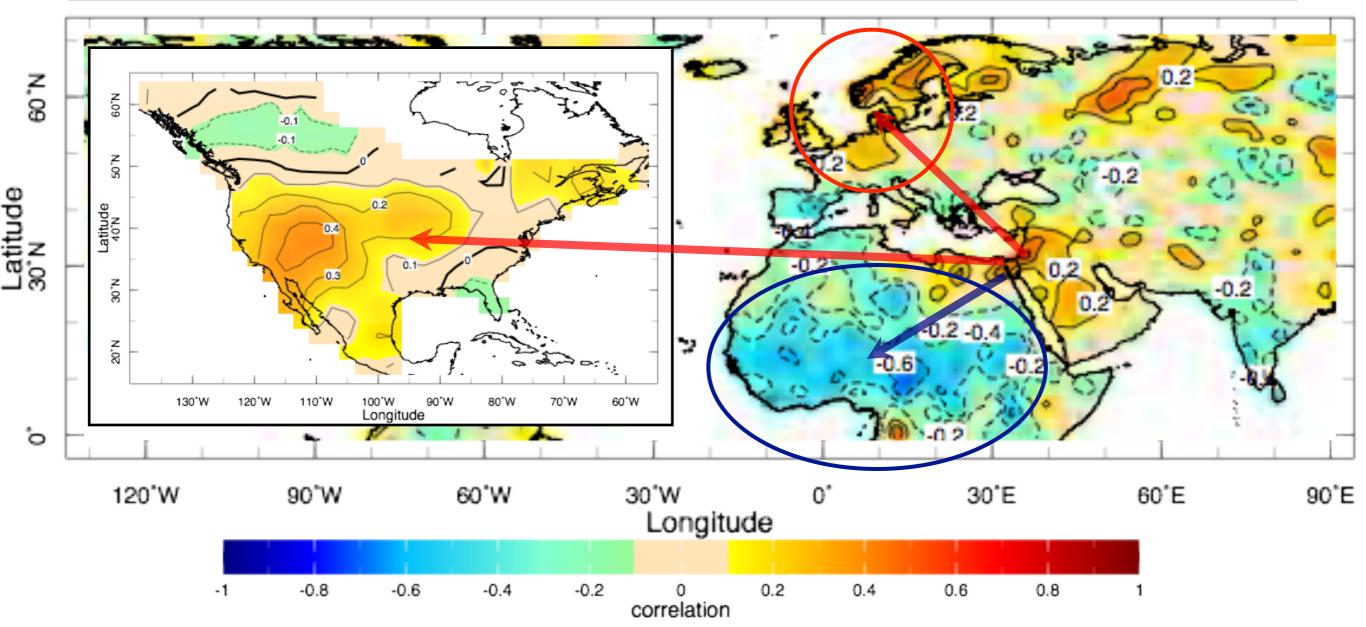


When the AMV is in its warm phase both surface and upper level flows over the Eastern Mediterranean are southerly and cold air incursions from Europe are blocked, leading to reduced cyclogenesis in the and relatively dry weather in the Levant.

The situation is reversed when the Atlantic is cold.

#### Jerusalem hemispheric precip teleconnections point at an "Atlantic governor"

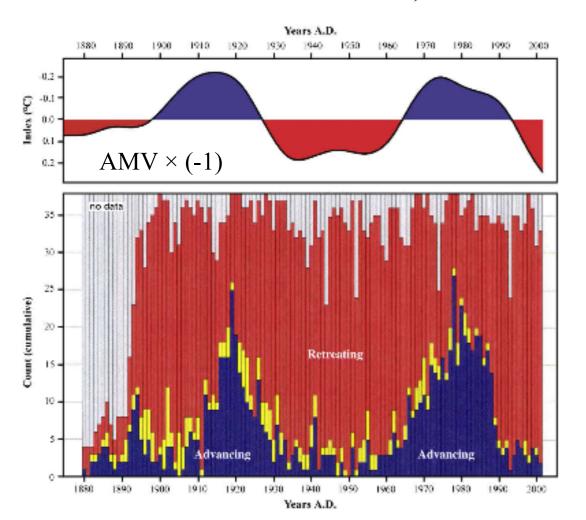
Kushnir and Stein (2010)



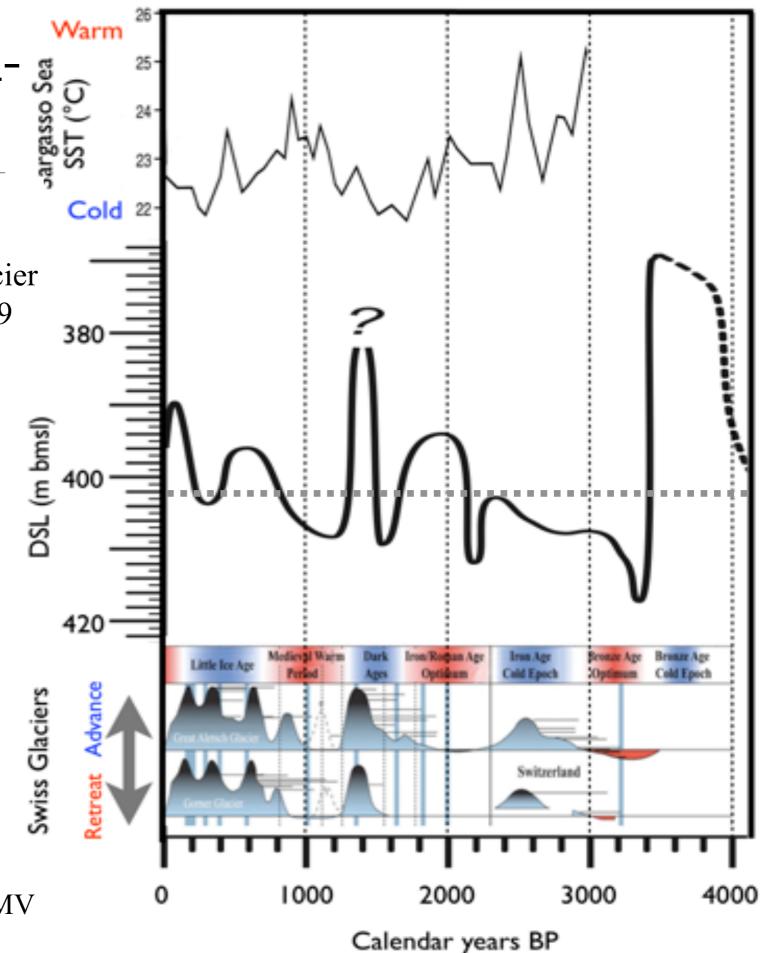
Annual (Oct-Sep) Jerusalem precipitation correlated with precipitation in surrounding land areas. Time series were smoothed by 1 pass of a binomial filter. Precipitation from GPCC 1930-1995. A correlation of 0.38 is significant at the 5% level (non-directional) assuming every fourth sample in the record is independent of the other (*Kushnir and Stein, 2010*).

# 4000 yr record of DSL-AMV teleconnections

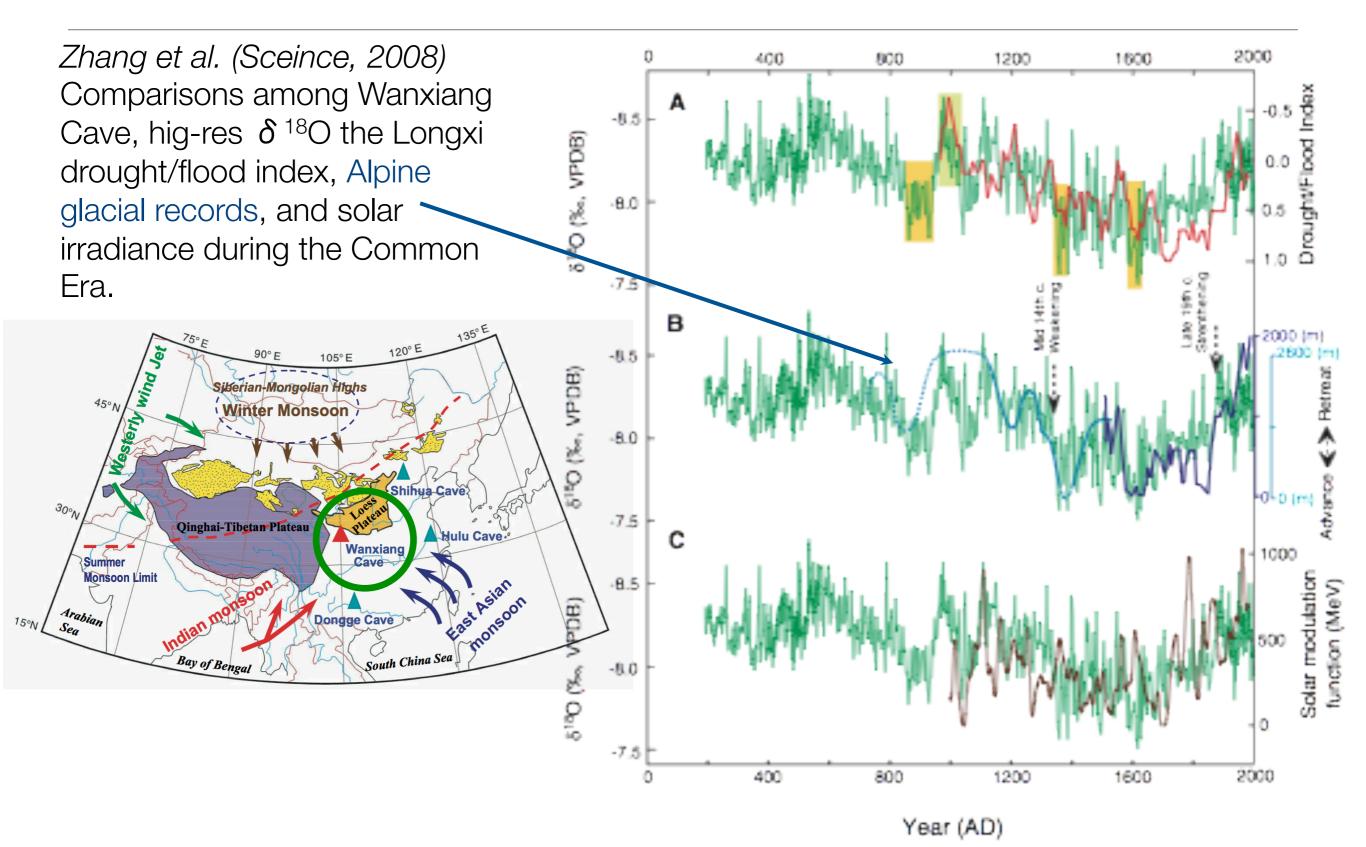
Dead Sea level (DSL) compared with the record of reconstructed Sargasso Sea SST(Keigwin et al., 1995) and Alpine glacier advances and retreats (Schaefer et al., 2009 after Holzhauser et al., 2005).



Alpine glaciers advances and retreats track AMV fluctuations (Denton and Brocker, 2008)

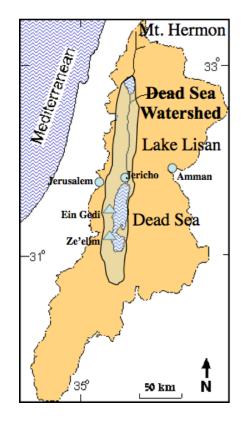


### AMV-East Asian monsoon connection

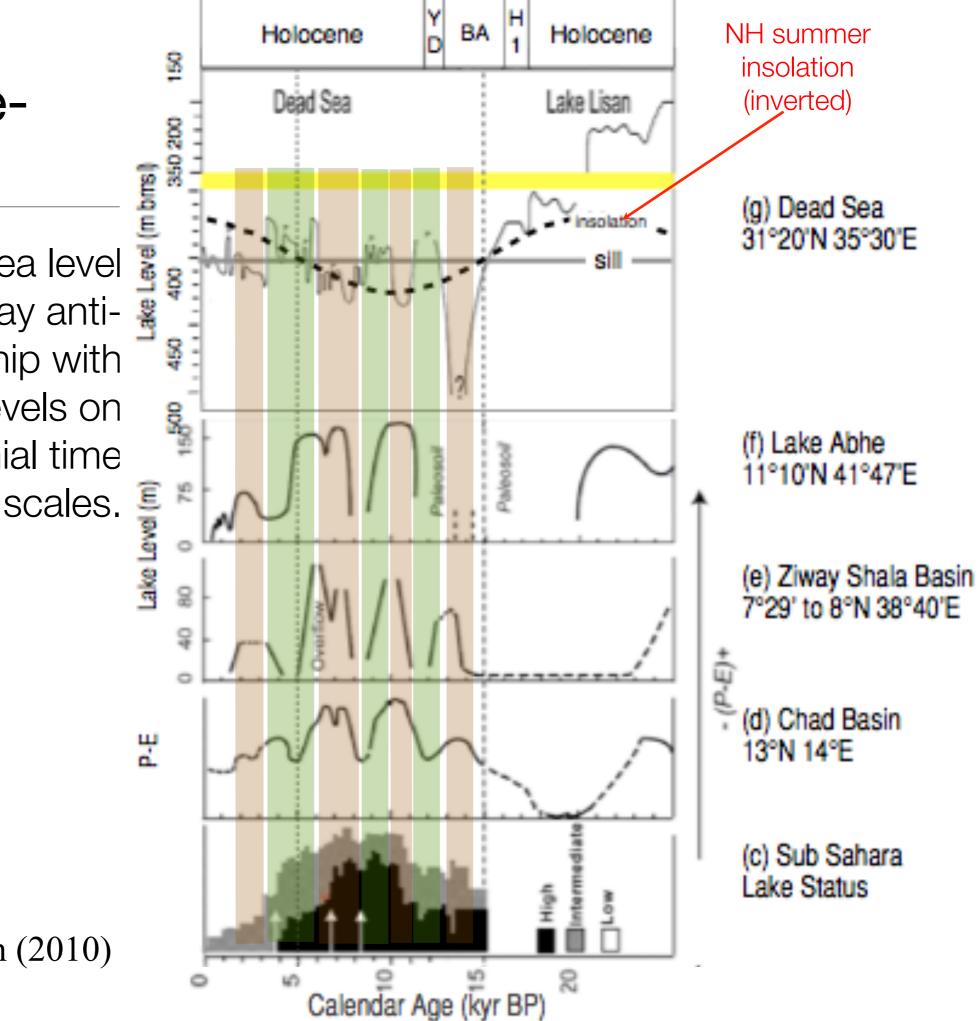


## Holocene Millennial Teleconnections

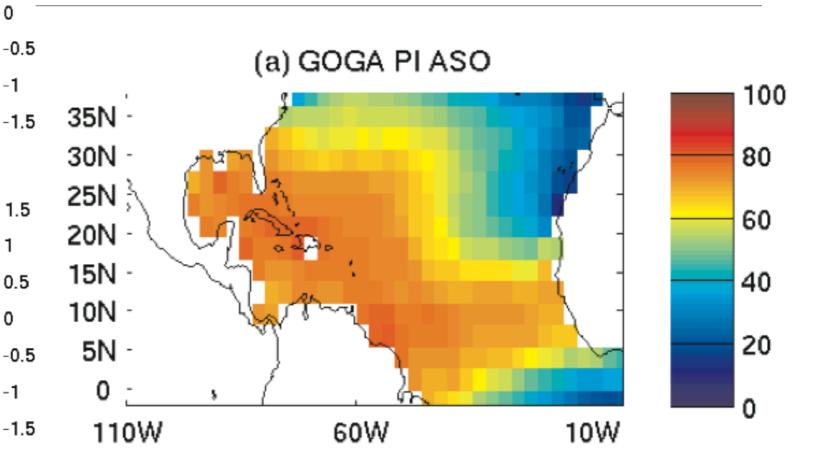
Holocene Dead Sea level variations display antiphase relationship with sub-Saharan lake levels on epochal and millennial time

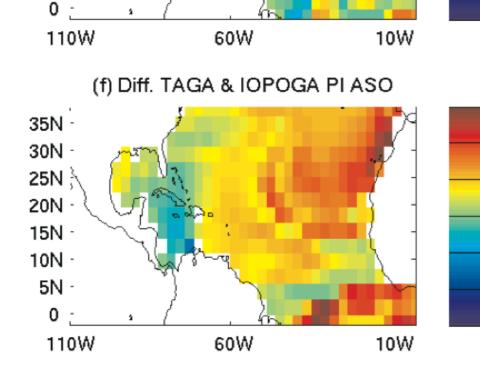


Kushnir and Stein (2010)



# Tropical Atl. hurricane Pl response to remote forcing





(b) Diff. GOGA & TAGA PI ASO

60W

(d) Diff. GOGA & IOPOGA PI ASO

35N -

30N ·

25N 3

20N

15N -

10N -

5N -

0 -

35N ·

30N ·

25N 3

20N

15N -

10N -

5N -

110W

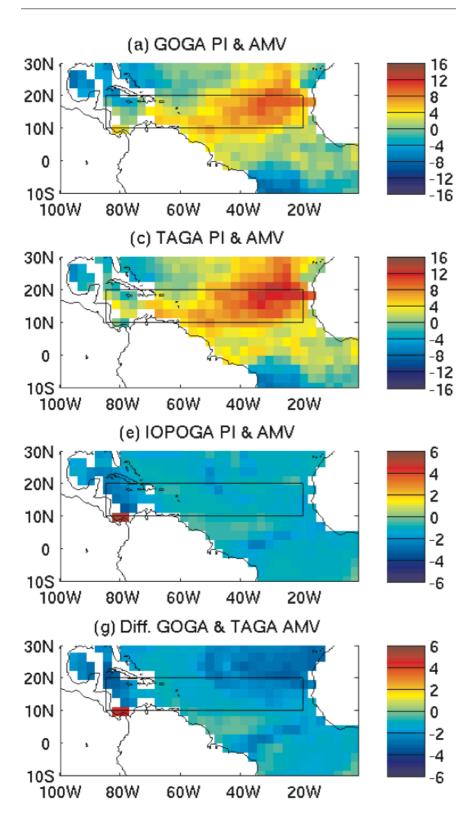
1.5

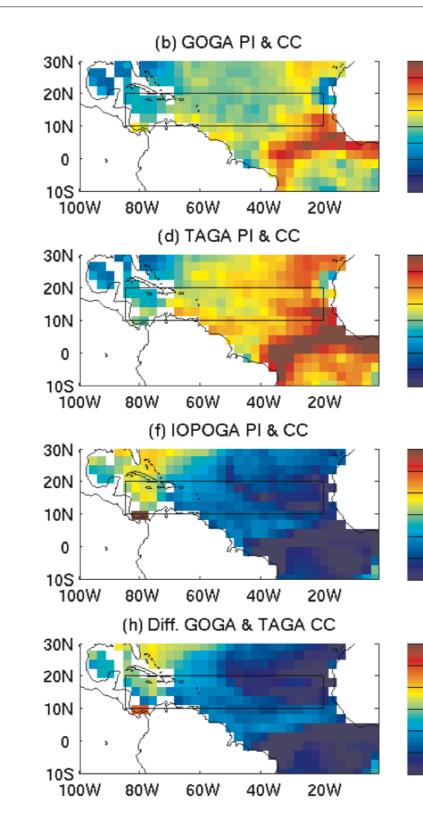
0.5

10W

 Ensemble mean, 1856 - 2006 climatological PI (m/s) for the peak hurricane season ASO in the tropical Atlantic for (a) GOGA simulation.
 Differences between climatological PI in ASO for (b) GOGA and TAGA, (d) GOGA and
 IOPOGA, (e) TAGA and IOPOGA.

### Impact of local and remote forcing on internal and externally forced Atlantic PI variability





Regression of anomalous 16 ASO PI (m/s) on AMV (left 12 panels) and climate -4 -8 change (CC) indices (right panels). Regression pattern of AMV and GOGA (a), TAGA (c), and 16 12 IOPOGA (e) . Regression pattern of CC index and GOGA (b), TAGA (d), and -8 -12 IOPOGA (f). The -16 difference between the PI regression patterns in GOGA and TAGA are shown in (g) and (h) for -2 the AMV and CC patterns, -4 respectively. The region of the main development region (MDR) is indicated by the black box.

8

0

4

2

0

4

Λ

-2

-4

-6

## Summary

- AMV is a prominent low-frequency phenomenon acting on a broad range of time scales and is likely the surface expression of AMOC variability.
- AMV impact climate in and around the Atlantic Basin as far as the South Atlantic and Asia and affests decadal variations in global surface temperature and precipitation.
- AMV played an important role in orchestrating significant multidecadal to millennial variability across the NH and beyond, throughout the past.
- Questions:
  - What are the dynamical mechanisms involved in shaping the AMV, particularly its tropical arm?
  - What is the response of AMV to different sources of external forcing: solar, volcanoes, CO<sub>2</sub>, aerosols?
  - How do we build on the emerging predictive skill?