Observations and dynamics of seasonal to interannual timescales: perspectives and future challenges

David Marshall (University of Oxford)

- 1. RAPID/MOCHA array: examples of new science the observations have stimulated.
- 2. Opportunities arising from new and continuing observations over the next 5 years.
- 3. Some (fairly random) thoughts on state estimation, numerical models, conceptual models, AMOC prediction.

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The contribution of eastern-boundary density variations to the Atlantic meridional overturning circulation at 26.5° N

M. P. Chidichimo^{1,2}, T. Kanzow^{3,4}, S. A. Cunningham³, W. E. Johns⁵, and J. Marotzke¹

Monthly-mean contributions of eastern boundary density to seasonal cycle of the AMOC:



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Wind-Stress Forcing of Seasonal AMOC Variability in a Two-Layer Model

Jiayan Yang

Department of Physical Oceanography, Woods Hole Oceanographic Institution



-6000 -5400 -4800 -4200 -3600 -3000 -2400 -1800 -1200 Model Bathymetry (meter)



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LETTERS

Mass and volume transport variability in an eddy-filled ocean

CARL WUNSCH

claimed AMOC signal should be swamped by eddies



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Kansow et al (2009)



Kansow et al (2009)



idealised numerical model experiments:



Kansow et al (2009)



idealised numerical model experiments:



linear wave theory:





Double Dip: Winter 2010/11

- Reemerging SSTs are observed in 1969/70 as well as in 2010/11
- These were conducive to the development of the negative NAO in winter 2010

Buchan et al. submitted, North Atlantic SST anomalies and the cold north European weather events of winter 2009/10 and December 2010. *submitted to Monthly Weather Review*

• Evidence that this second negative is predictable due to the ocean

Maidens et al. in prep, The Influence of Surface Forcings on Prediction of the North Atlantic Oscillation Regime of Winter 2010-11. *submitted to Monthly Weather Review*



Winters of 2010/11 (green), 1969/70 (blue) and 1978/79 (red). Black shows mean (1960-2011) with 1, 2 std envelopes

courtesy: Gerard McCarthy



Pillar, Johnson and Marshall (in prep.): Attribution of AMOC variability at 25N (in 1° MITgcm)

Maps of linear sensitivity of AMOC (25N) to surface forcing at different time lags: **zonal wind stress** (Sv/Nm⁻²) meridional wind stress (Sv/Nm⁻²) heat flux (Sv/Wm⁻²) 30 0° 0° 0° 30°S 30°S 30°S -2 months $_{60^{\circ}\text{S}}$ -1 month -6 years -9 years -2 months 1 month 60^c 60 60[°] 30° 30° 30 0° 0° 0° 30°S 30°S $30^{\circ}S$



-3 months

10 months

-0.01

0

-0.1

60°S

30

00

 $30^{\circ}S$

-5 months

-1.1 years

0.1

0.01



-0.0001 -1e-05 -1e-06 0 1e-06 1e-05 0.0001

30°

 0°

30°S

AMOC response = \sum (monthly sensitivities) x (forcings) ... over all grid points, preceding 15 years:



AMOC response = \sum (monthly sensitivities) x (forcings) ... over all grid points, preceding 15 years:



Dependence of AMOC response on integration period: (= **memory** of past forcing)





Decomposition of the AMOC in HadCM3 (Sime et al., 2006)



variability in the external mode?

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OSNAP

U.S. : Susan Lozier (Duke); Bill Johns (U. Miami); Amy Bower, Bob Pickart and Fiamma Straneo (WHOI)

UK: Sheldon Bacon, Penny Holliday and Chris Wilson (NOC); Stuart Cunningham and Mark Inall (SAMS); David Marshall and Helen Johnson (Oxford) and Ric Williams (Liverpool)

Netherlands: Laura de Steur (NIOZ)

Germany: Jürgen Fischer, Johannes Karstensen and Martin Visbeck (GEOMAR) and Torsten Kanzow (AWI)

Canada: Blair Greenan (BIO); Brad de Young (Memorial U.)

France: Herlé Mercier and the OVIDE group (IFREMER)

courtesy: Susan Lozier

OSNAP overall goal: To quantify the large-scale, low-frequency, full watercolumn net fluxes of mass, heat and fresh water associated with the meridional overturning circulation in the subpolar North Atlantic.



- (A) German 53°N western boundary array and Canadian shelfbreak array;
- (B) US West Greenland boundary array;
- (C) US/UK East Greenland boundary array;
- (D) Netherlands western Mid-Atlantic Ridge array;
- (E) US eastern Mid-Atlantic Ridge array;
- (F) UK survey over the Hatton-Rockall Bank and Rockall Trough;
- (G) UK Scottish Slope current array.
- Red dots: US float launch sites.

Blue star: US OOI Irminger Sea global node. Black concentric circles: US sound sources.

courtesy: Susan Lozier

In the water: Summer 2014

Specific OSNAP objectives:

 Quantify the subpolar AMOC and its intra-seasonal to interannual variability via overturning metrics, including associated fluxes of heat and freshwater.



- 2. Determine the pathways of overflow waters in the NASPG to investigate the connectivity of the deep boundary current system.
- 3. Relate AMOC variability to deep water mass variability and basin-scale wind forcing.
- 4. Determine the nature and degree of the subpolar-subtropical AMOC connectivity.
- 5. Determine from new OSNAP measurements the configuration of an optimally efficient long-term AMOC monitoring system in the NASPG.
- courtesy: Susan Lozier

South Atlantic MOC array (SAMOC):



courtesy: Chris Meinen

Equatorial buffer

=> limited coherence between North and South Atlantic on seasonal to interannual time scales:

(Johnson and Marshall, 2002)



but



Intrinsic AMOC_{σ_2} : meridional coherence



min -0.3Sv max +0.3Sv





courtesy: Thierry Penduff

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Monthly AMOC and Atlantic OHT variability



Note: Neither Florida current transports (cable observations) or RAPID AMOC transport estimates at 26°N are used as constraints in ECCO

courtesy: Martha Buckley

How many critical processes for AMOC variability are adequately-resolved by current models?



Gelderloos et al. (2011) - "GFD in a GCM"



 $(\Delta x = 3.75 \text{ km})$



SST

Gelderloos et al. (2011) - "GFD in a GCM"



SST

Gelderloos et al. (2011) - "GFD in a GCM"



SST

f(x) = f(x) =e Warshall and Johnson (2013) ced gravity (and rections, coefficient of linear friction. Wetadoptiacon plane to proving the such that the formation of $f = \beta y$. roximation such that the Coriolis parameter is $f = \beta y$. For similar we will estrict out attenduced gravity source in the strict out attenduced gravity tal stown boundarys constitute: place at Prentich ill Fres u = 0western boundary
boundary u = 0western boundary (4)boundary (4)western boundary (4)boundary (4)boundary
boundary (4)boundary (4)io-nd now boundary ion: **– 2** – 0 stern boundary so ane $x \ge 0$ and eastern boundary asplaendh ath p oltions, we can derive the vorticity equation: he vorticity eq ation: $\mathbf{s},$ • adjustment is not by Kelvin waves but Rossby $\frac{\partial^2 h}{\partial x^2} = \frac{h}{d_{2D}^2} \xrightarrow{h}{\partial x} \xrightarrow{h}{\partial x} \xrightarrow{h}{\partial x^2} \xrightarrow{h}{\partial x^2}$ $\left(\frac{1}{\partial x^2} - \frac{h}{L_D^2}\right) + \beta \frac{\partial h}{\partial r}$ $\overline{\partial t}$ $L_D(y) = \frac{\sqrt{g'H}}{2}$ where $L_D(y) = \frac{\sqrt{g'H}}{\beta y}$ is the Rossby deformation radius. • in reality, bottom topography likely to dominate.

bothdary comentary be model dependent

$$\frac{\partial^2 h}{\partial t \partial x} + \beta y \frac{\partial h}{\partial y} + r \frac{\partial h}{\partial x} = \frac{\partial^2 h}{\partial t \partial x} + \begin{pmatrix} \beta y \frac{\partial h}{\partial y} + r \frac{\partial h}{\partial x} = 0 \quad (x = 0) \end{pmatrix}$$
(6)

Comment on equilibration time scale for models: O (100s - 1000s years)

Journal of Marine Research, 69, 167-189, 2011

Spin-up and adjustment of the Antarctic Circumpolar Current and global pycnocline

by Lesley C. Allison¹, Helen L. Johnson² and David P. Marshall³



Gnanadesikan (1999) with time dependence

time scale set by

Southern Ocean eddy diffusivity and d(AMOC)/d(pycnocline depth)

adjustment likely to be faster with explicit eddies

(also see Jones et al., 2011; Samelson, 2011)



Msadek et al. submitted

courtesy: Rym Msadek

Concluding remarks

- The RAPID/MOCHA observations have led to many new discoveries about the AMOC, many of which were not anticipated at the outset.
- The new and continuing observations (OSNAP, RAPID/MOCHA, SAMOC, ...) will lead to further new discoveries, the most exciting of which have not yet been anticipated.
- Need to think hard about role of new technologies, cheaper array designs, for the longer term.
- Much progress has been made with state estimation, but direct AMOC observations remain essential. Still unclear how to best assimilate boundary hydrographic measurements.
- Some encouraging signs for predictability of the AMOC and its impact on SSTs, etc.
- Attribution of observed changes is also important (e.g., recent hiatus in global warming).
- Don't forget conceptual models "all models are wrong, some models are useful".