

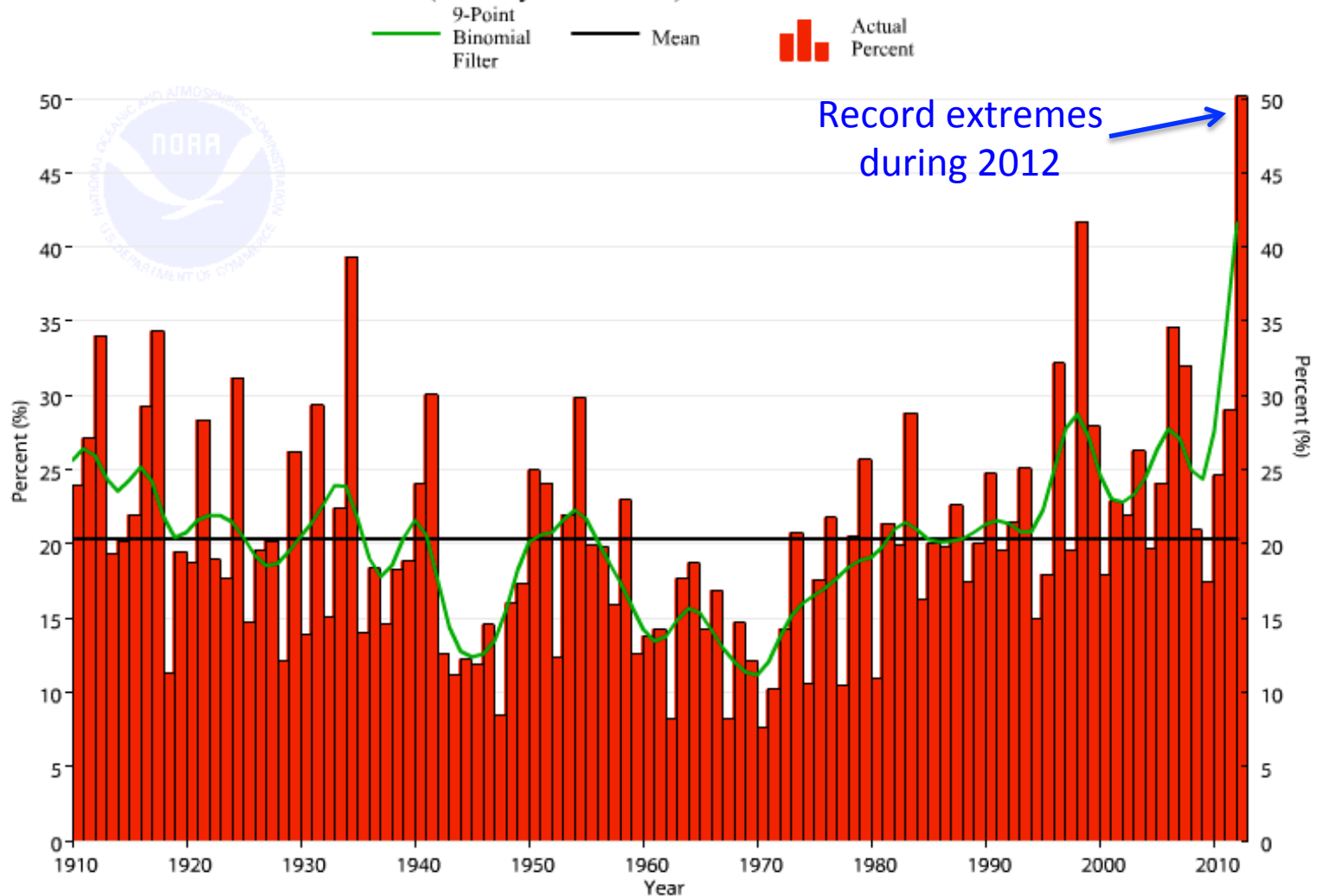
Relating Extreme Weather Events to Large-Scale Meteorological Patterns: Is the Glass half full or half empty?

Steve Vavrus
Center for Climatic Research
University of Wisconsin



U. S. Climate Extremes Index

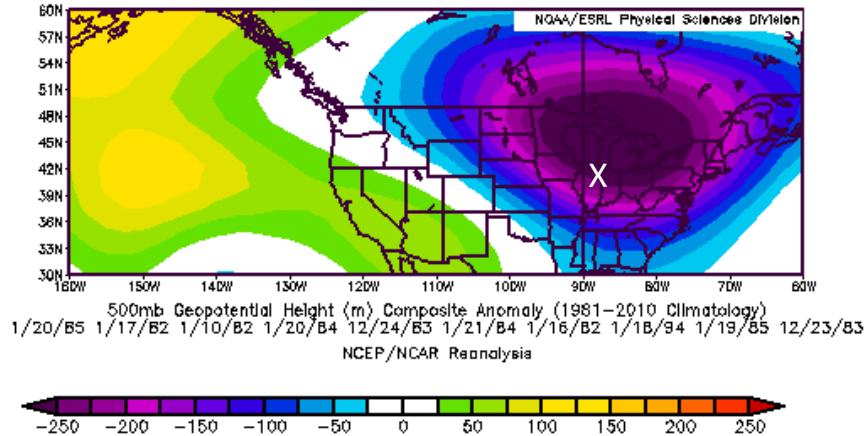
Contiguous U.S. Without Tropical Cyclone Indicator
Annual (January-December) 1910-2012



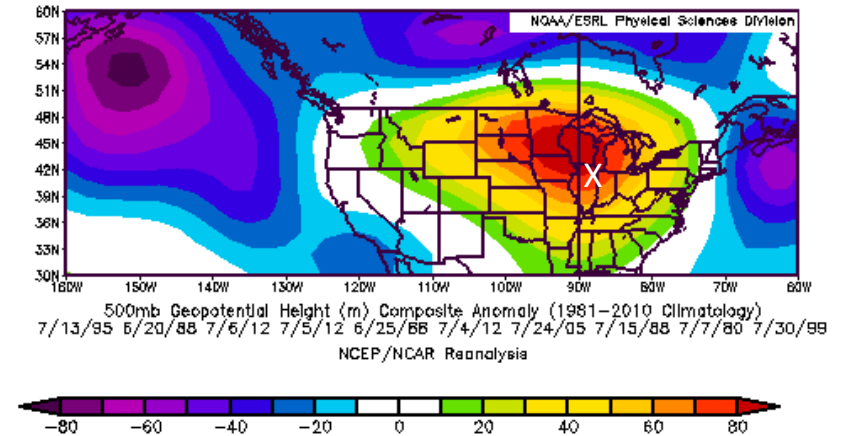
Upper-air Circulation Anomalies during Extreme Weather (Chicago)

High-Amplitude Wave Patterns

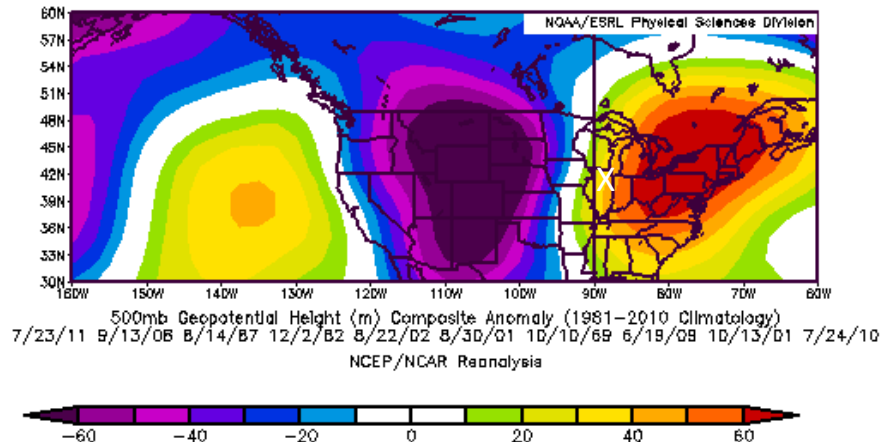
10 Coldest Days



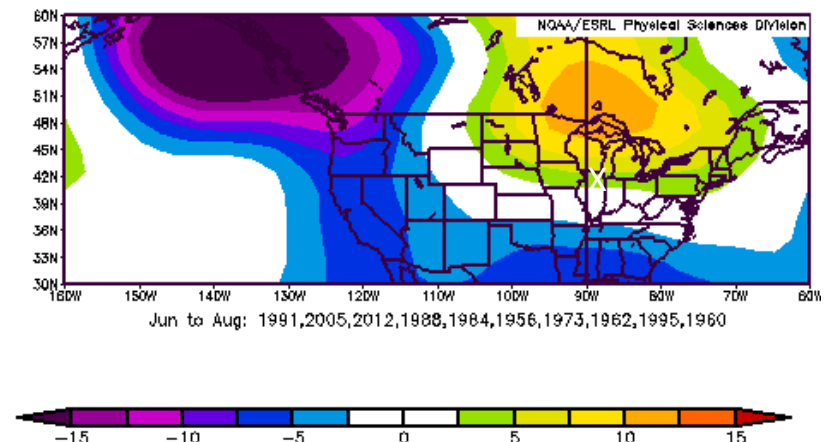
10 Hottest Days



10 Wettest Days

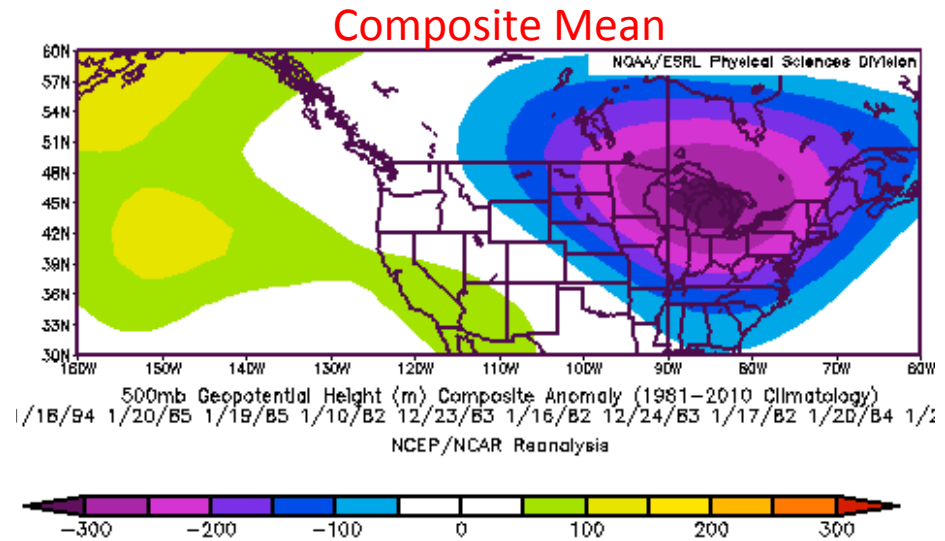


10 Driest Summers



Chicago's 10 Coldest Days

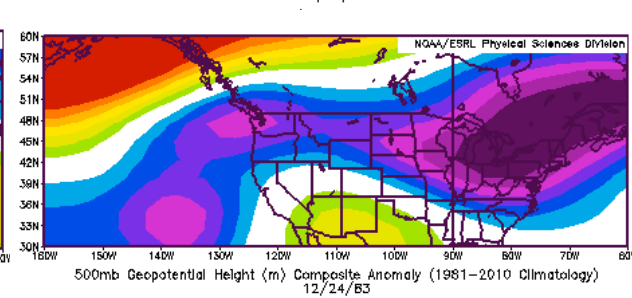
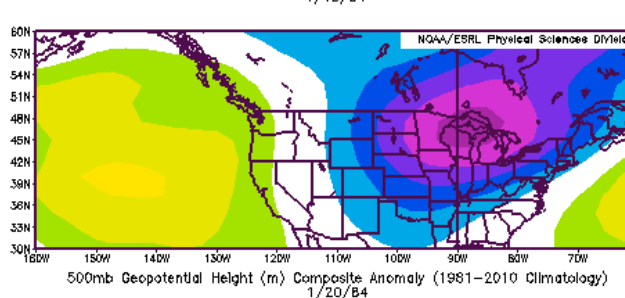
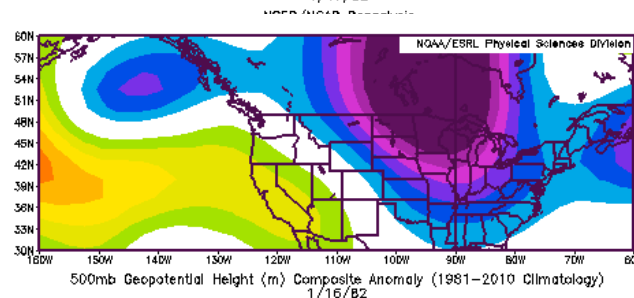
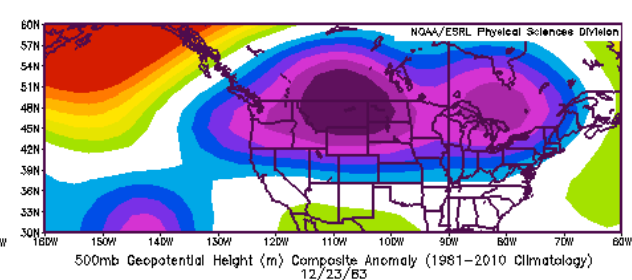
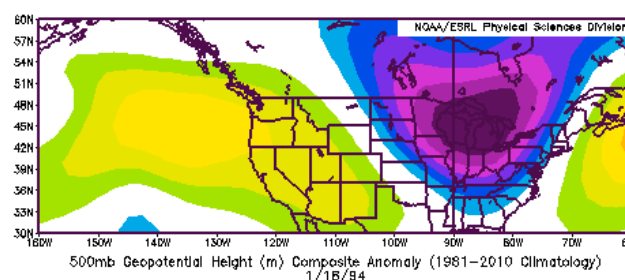
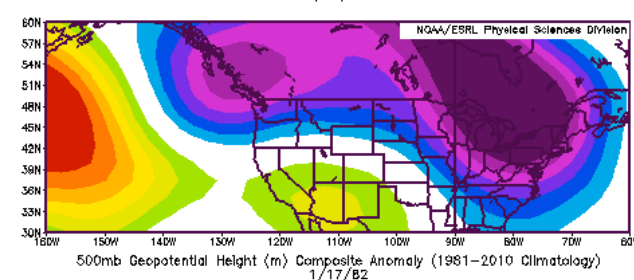
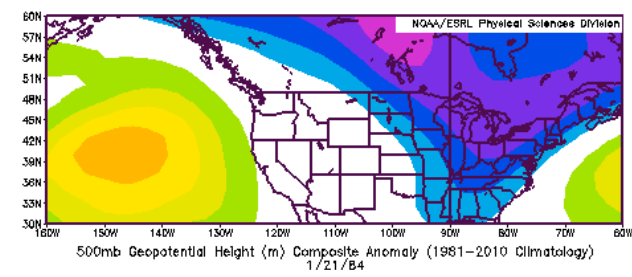
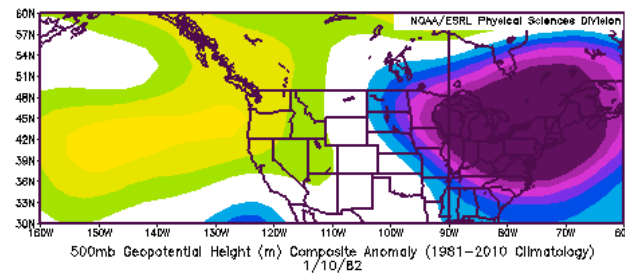
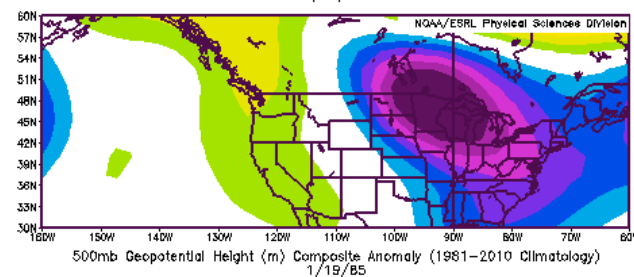
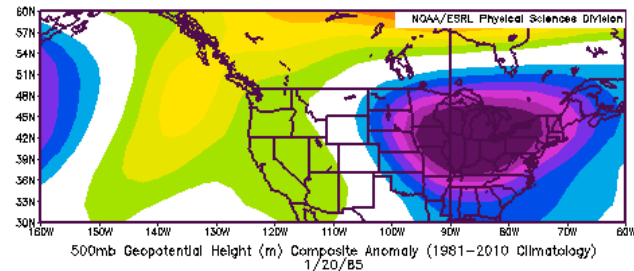
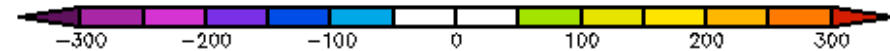
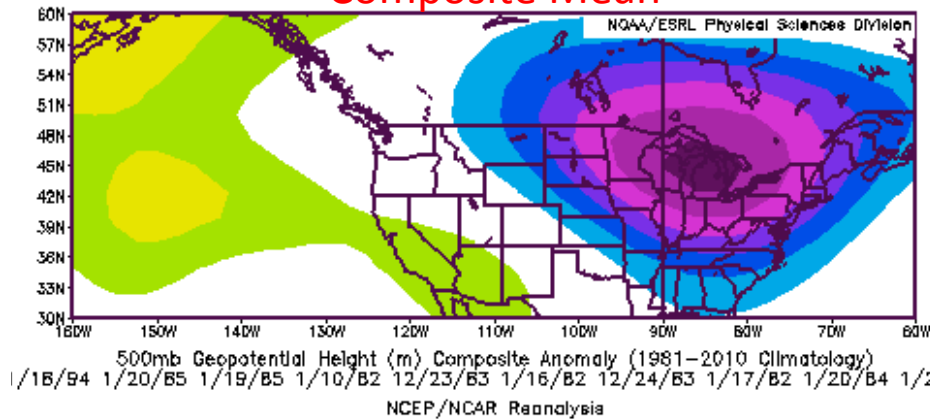
500 hPa Anomalies (m)



Chicago's 10 Coldest Days

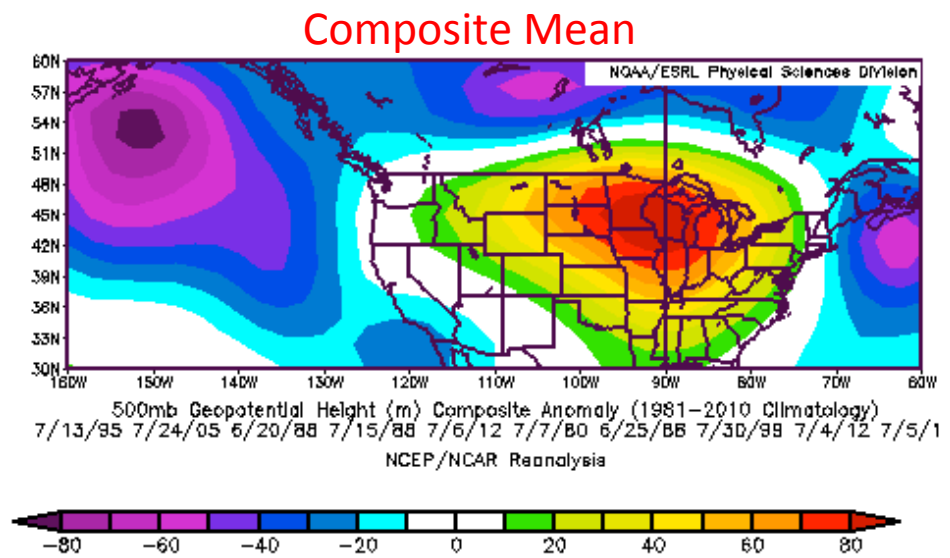
500 hPa Anomalies (m)

Composite Mean



Chicago's 10 Hottest Days

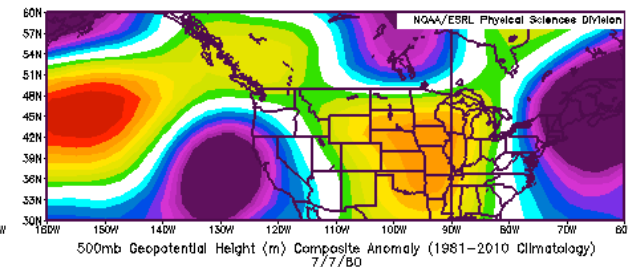
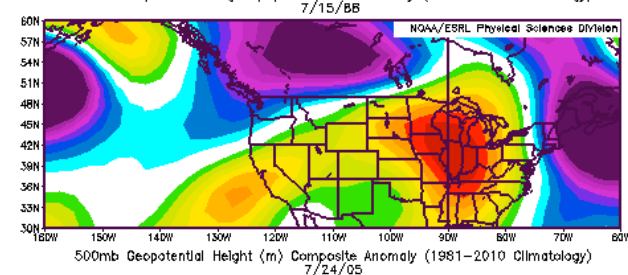
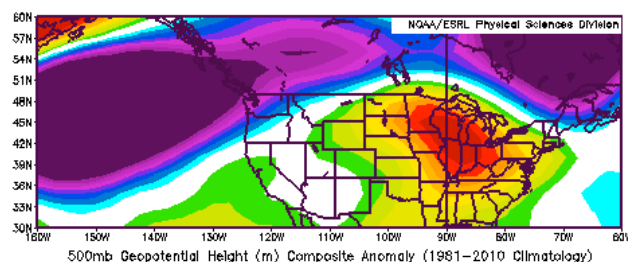
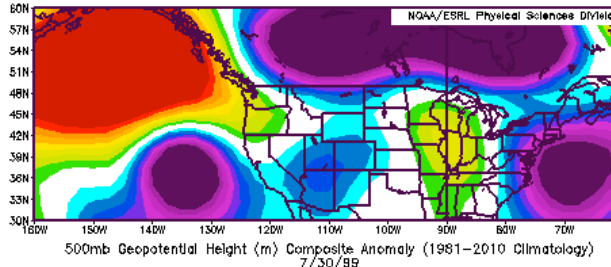
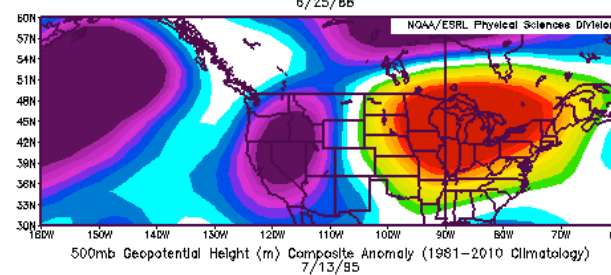
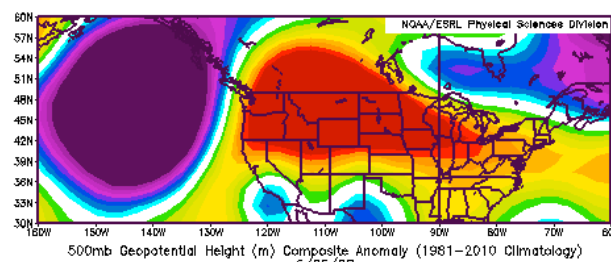
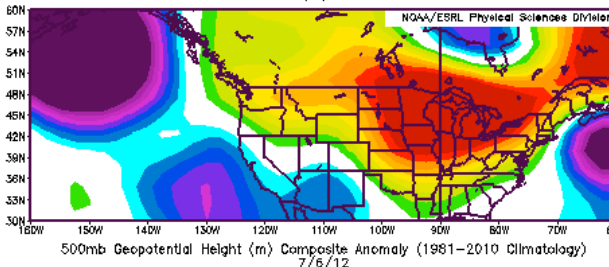
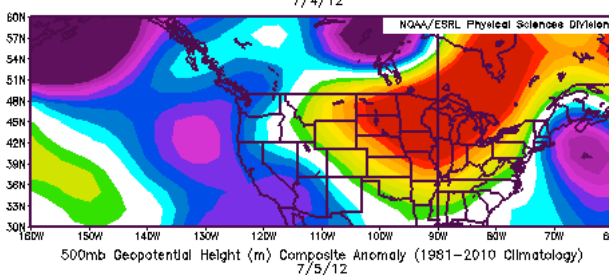
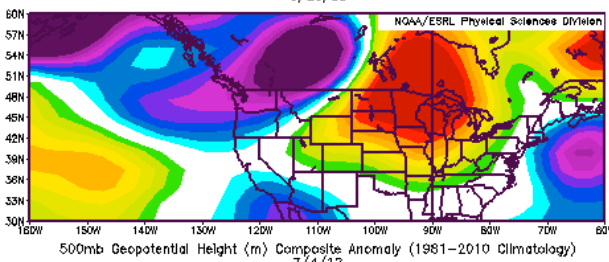
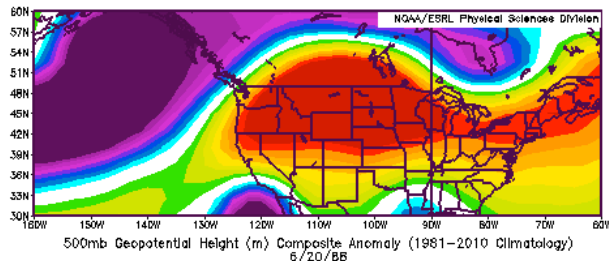
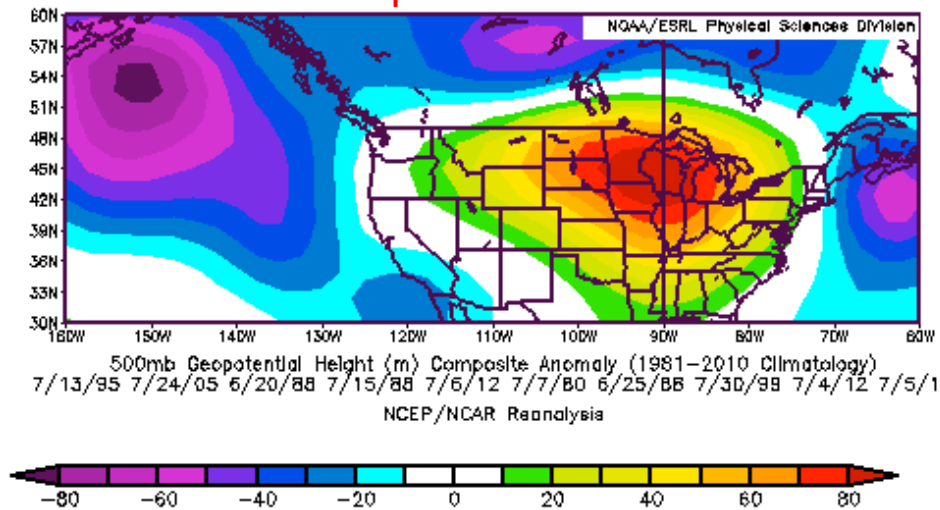
500 hPa Anomalies (m)



Chicago's 10 Hottest Days

500 hPa Anomalies (m)

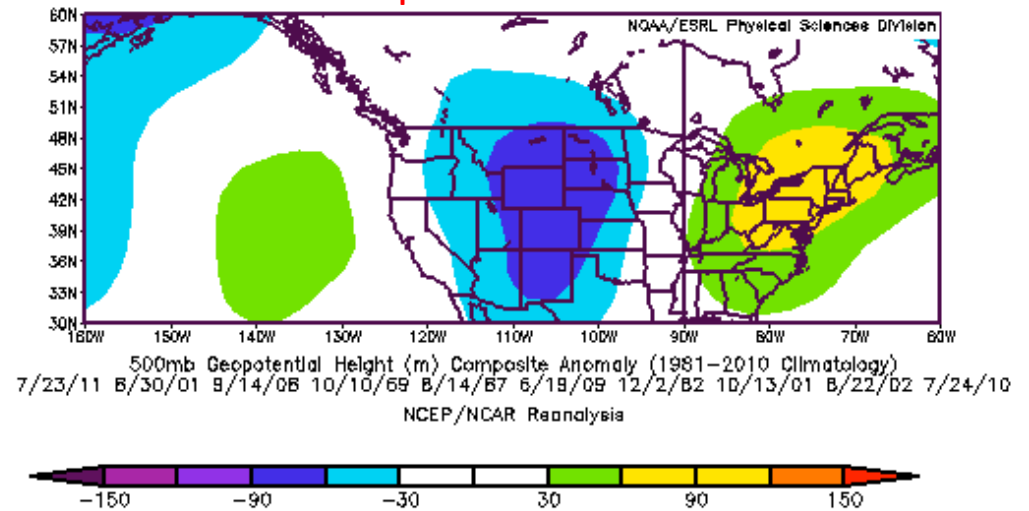
Composite Mean



Chicago's 10 Wettest Days

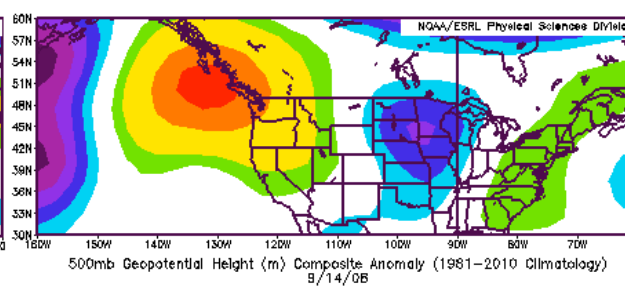
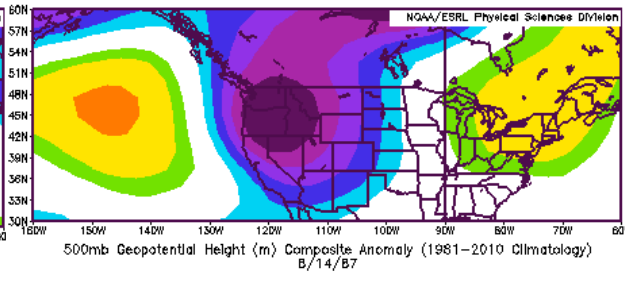
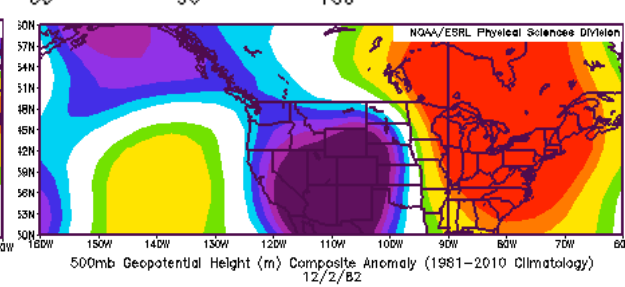
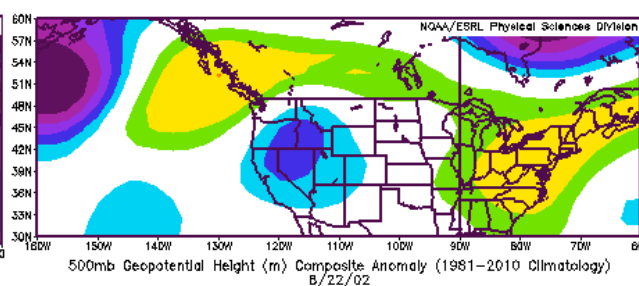
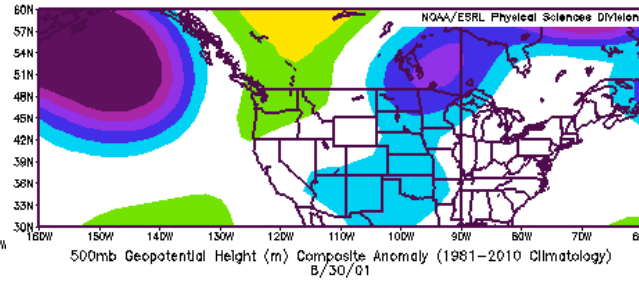
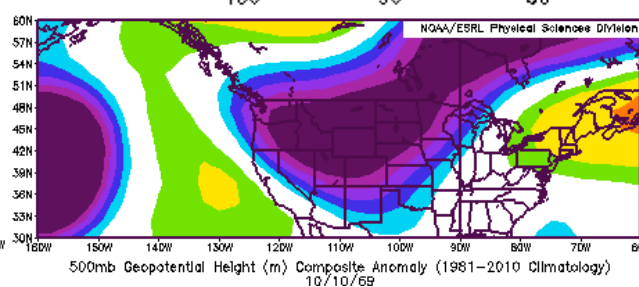
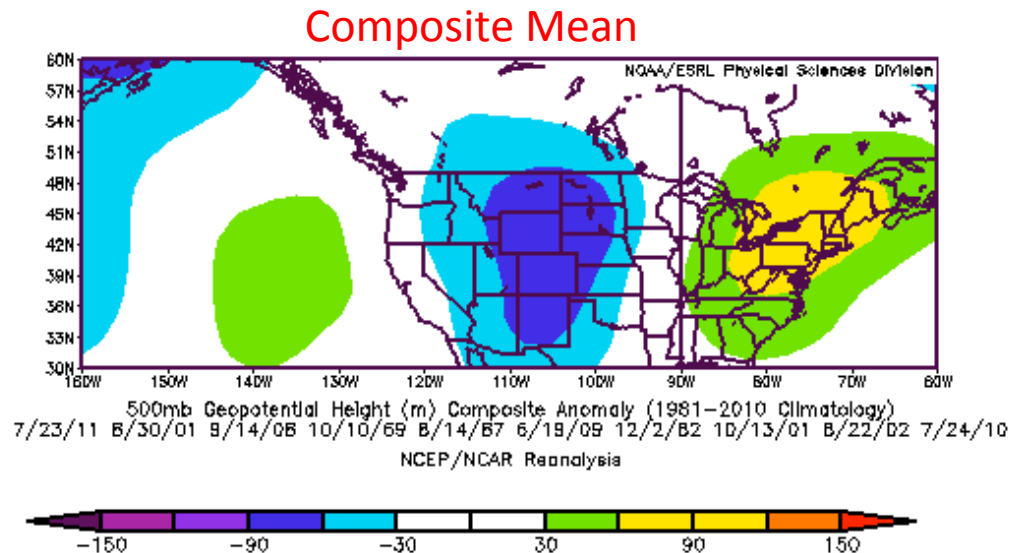
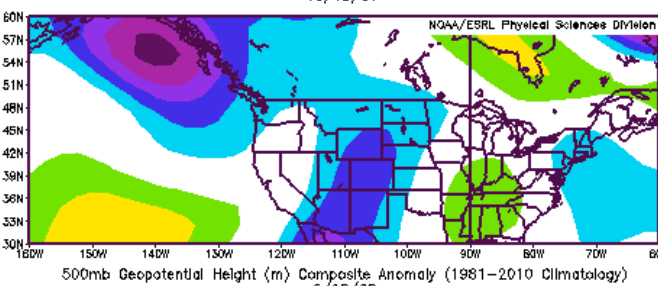
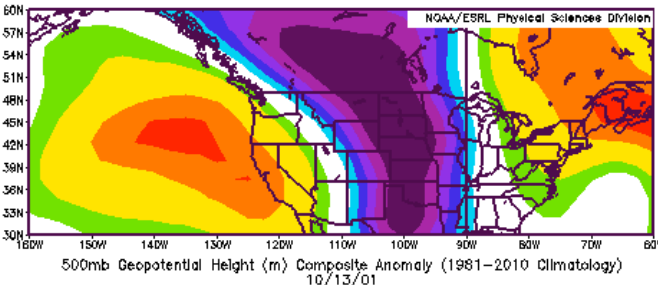
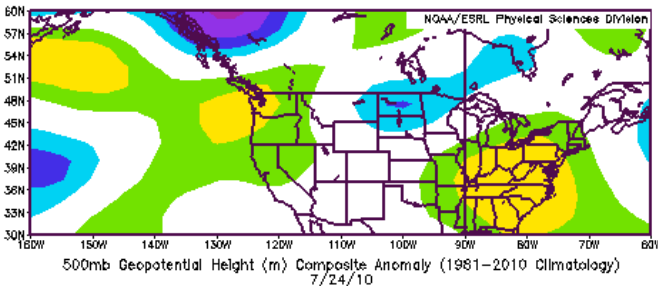
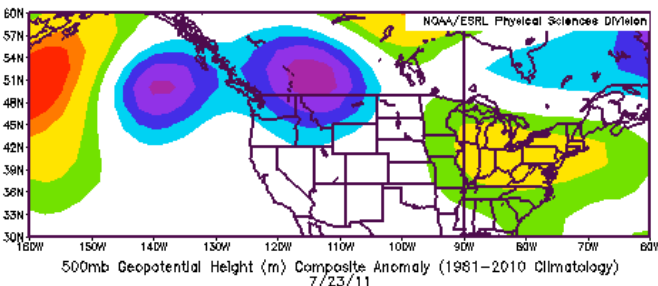
500 hPa Anomalies (m)

Composite Mean



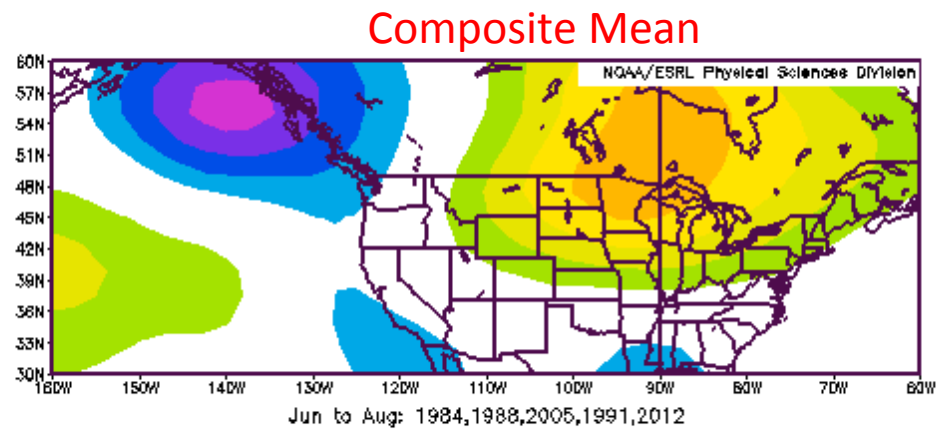
Chicago's 10 Wettest Days

500 hPa Anomalies (m)



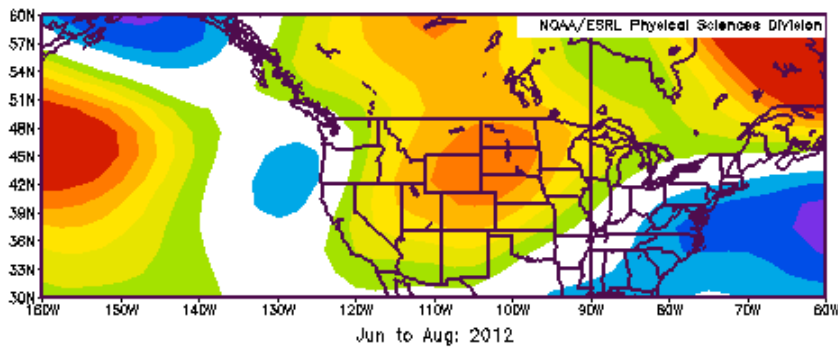
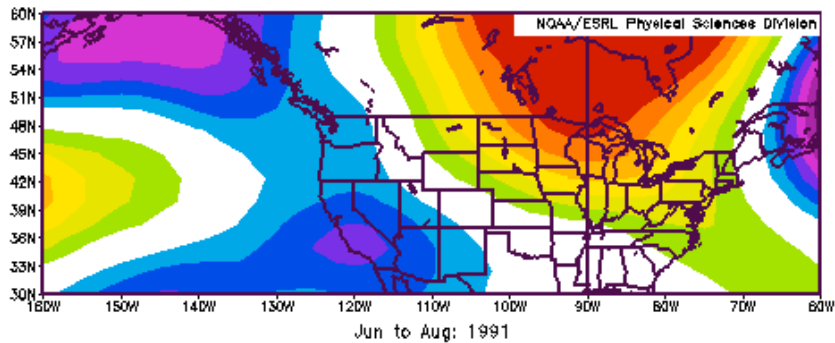
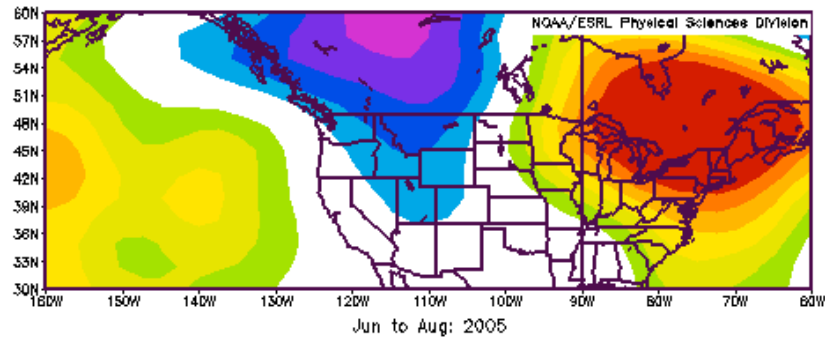
Chicago's 5 Driest Summers

500 hPa Anomalies (m)

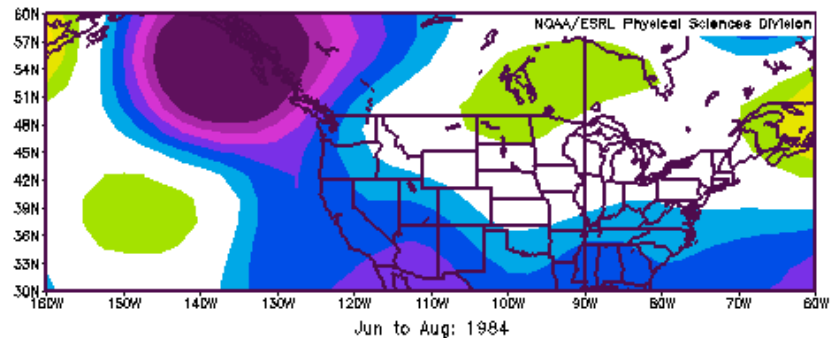
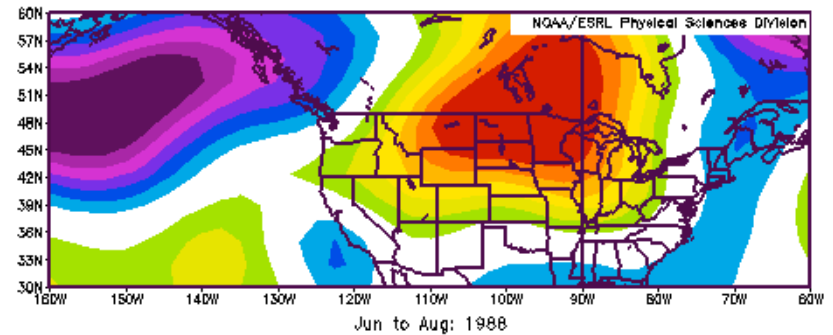
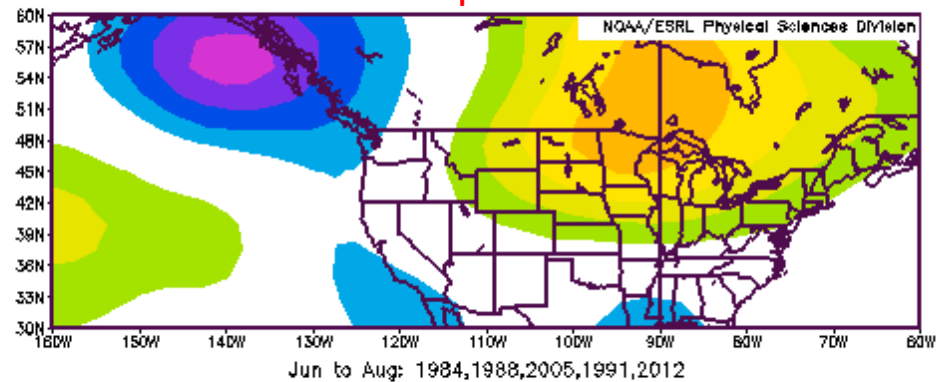


Chicago's 5 Driest Summers

500 hPa Anomalies (m)

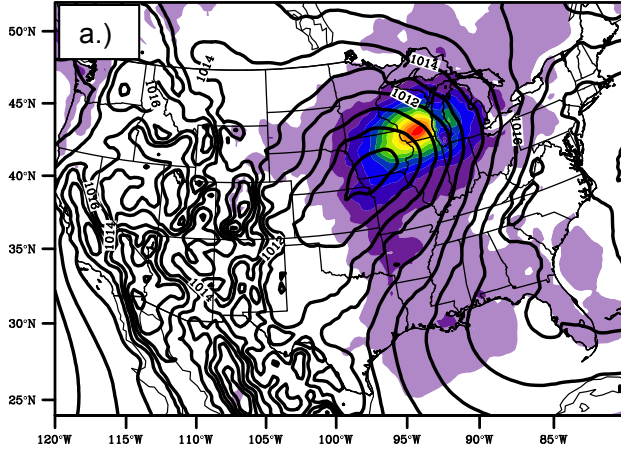


Composite Mean

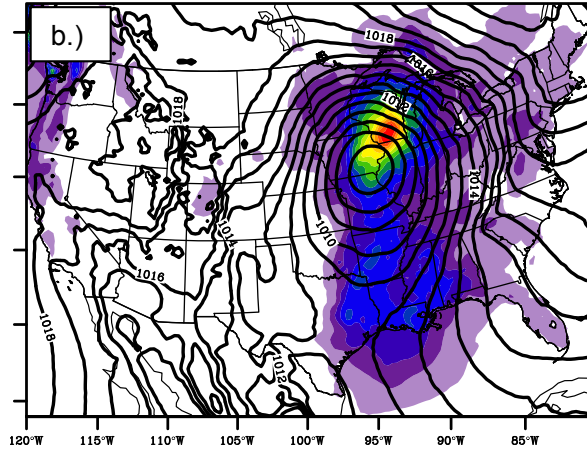


SLP and Precipitation on wettest 1% of days in Madison, WI North American Regional Reanalysis 1981-2000

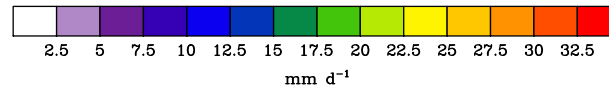
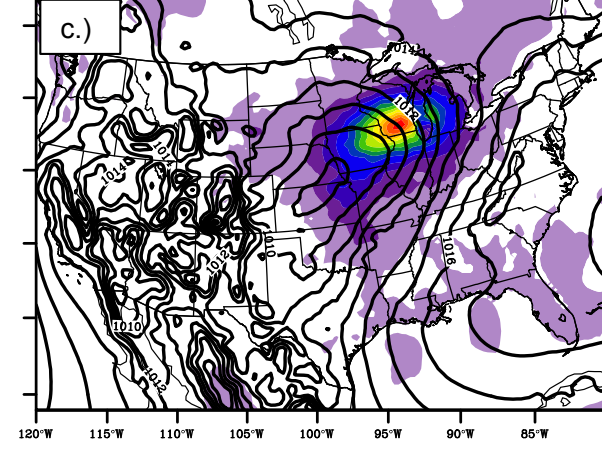
All days



Cold Season



Warm Season



Simulated SLP and Precipitation on wettest 1% of days in Madison, WI: 20th Century

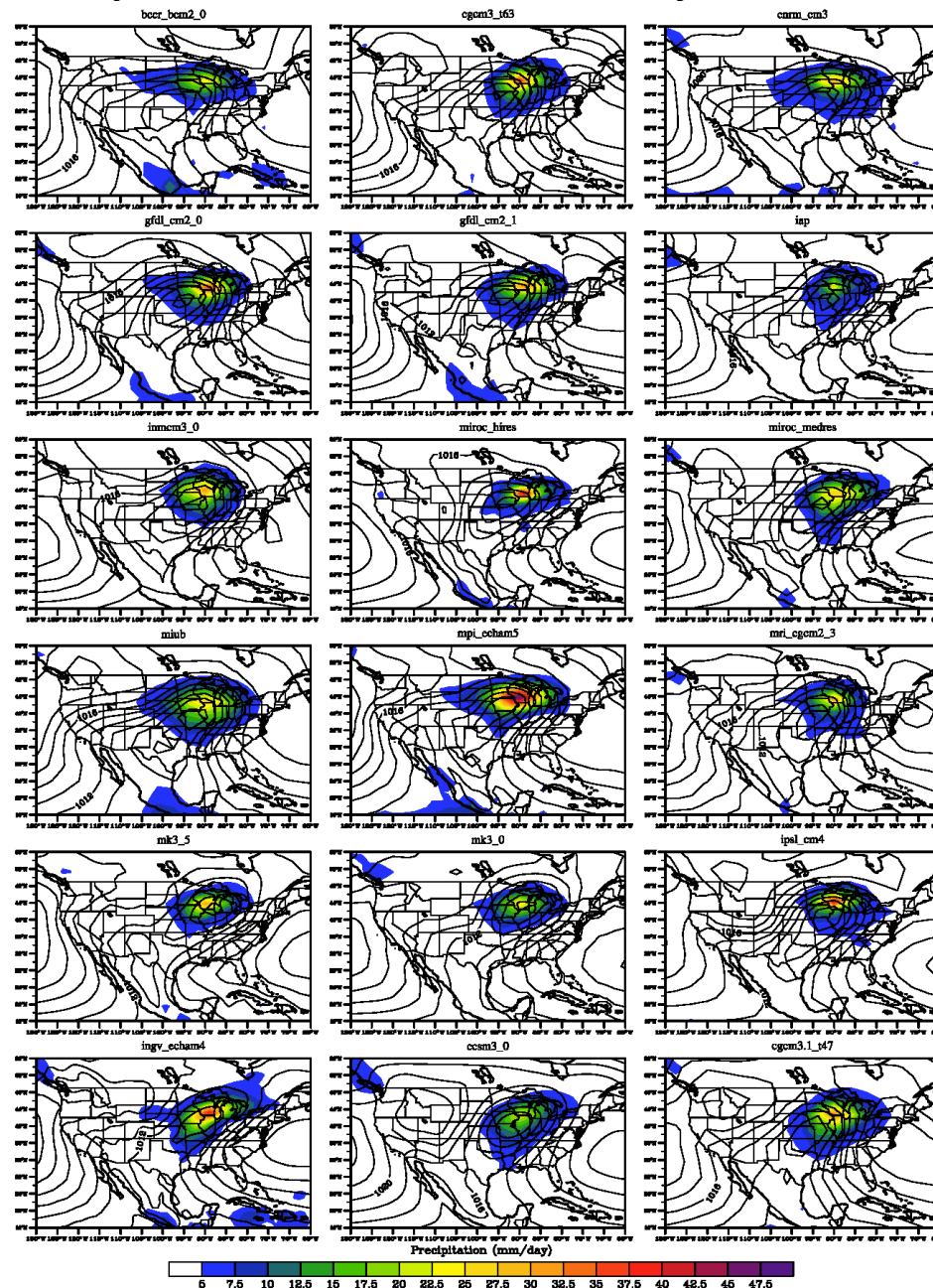


FIGURE 3.13: Late twentieth century (1981-2000) spatial composites of daily precipitation (mm/day) and sea-level pressure (hPa, contoured every 2hPa) during the wettest 1% of days for

Simulated SLP and Precipitation on wettest 1% of days in Madison, WI: 21st Century

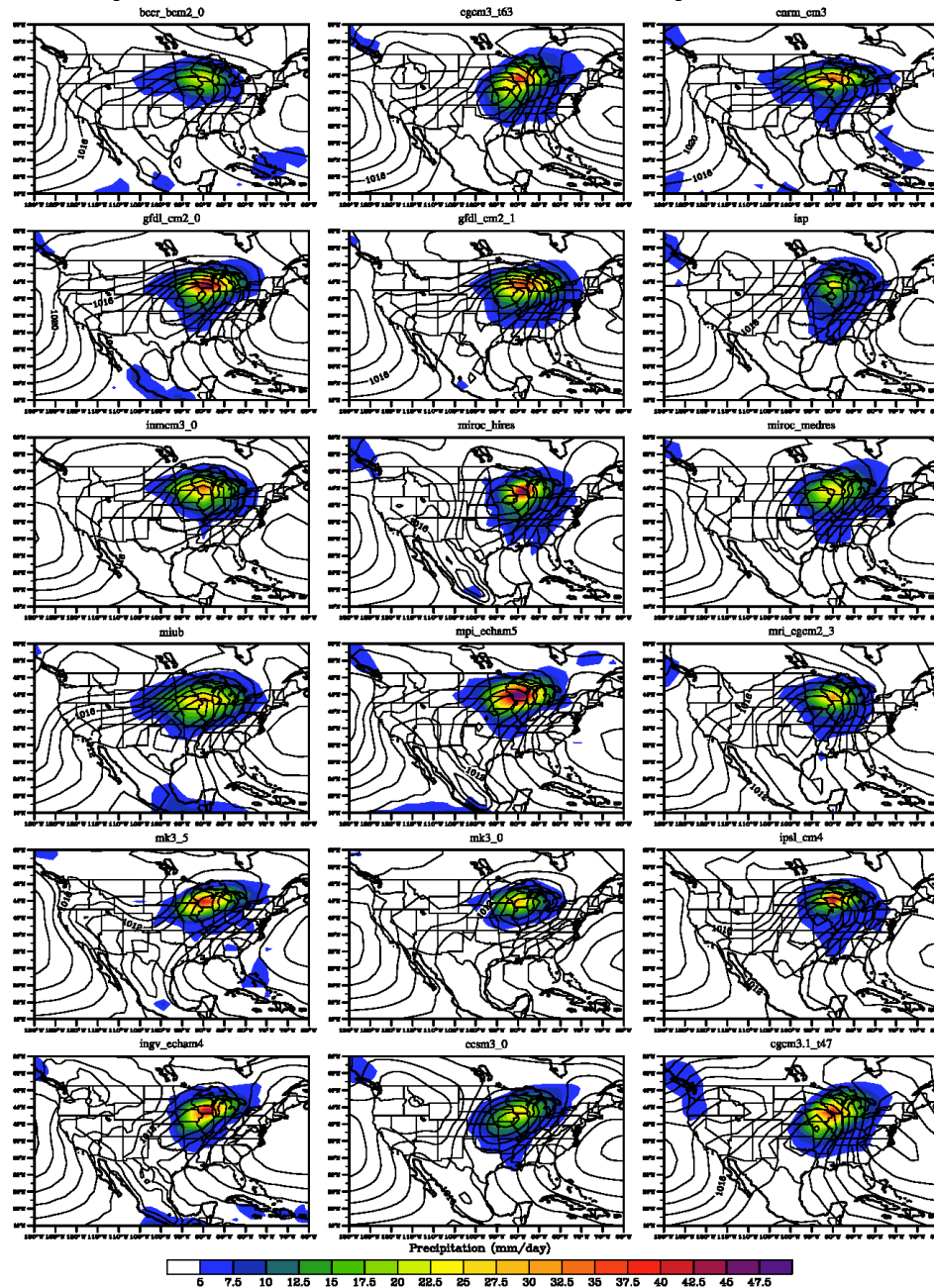
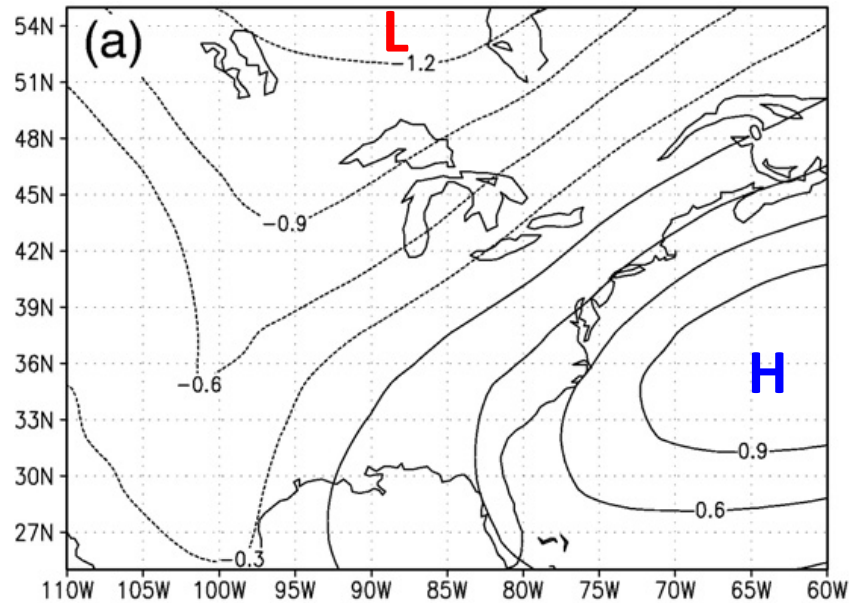


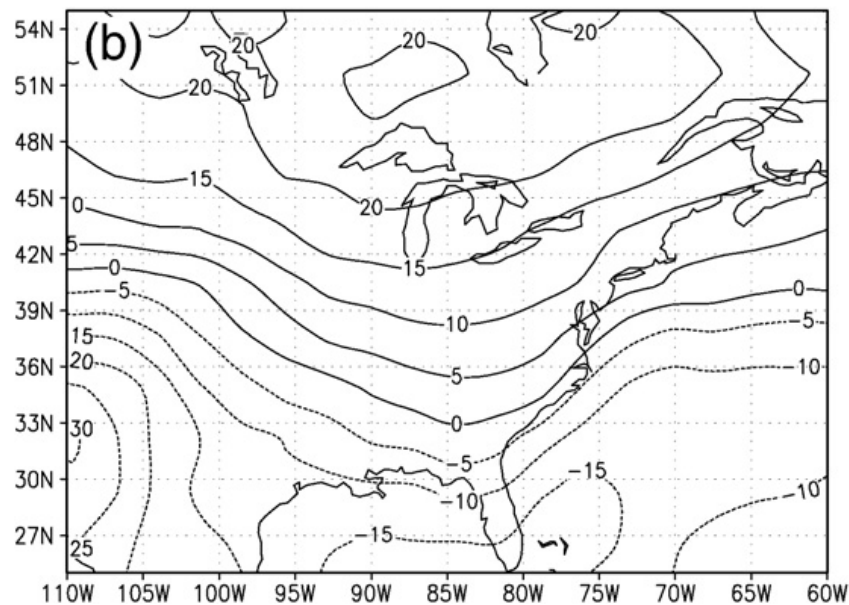
FIGURE 3.14: Same as Fig 3.13 except for the late twenty-first century (2081-2100).

Synoptic Pattern related to Projected Climate Change

Springtime Precipitation (mid-21st Century, A1B)



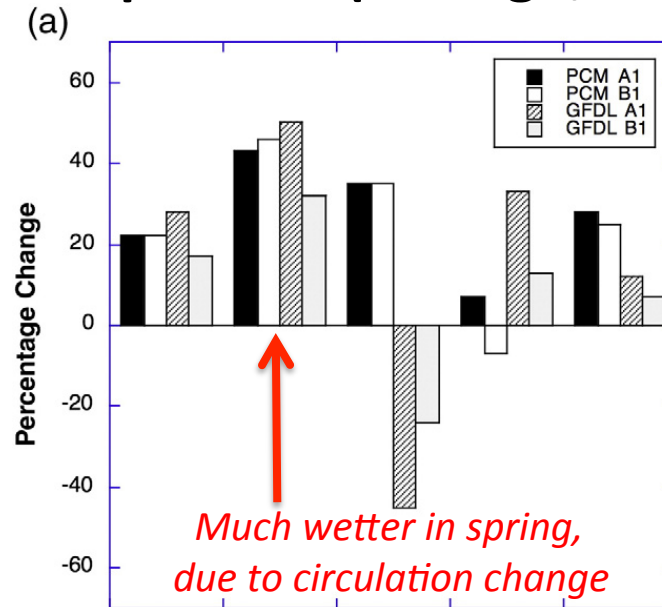
SLP Response (hPa)



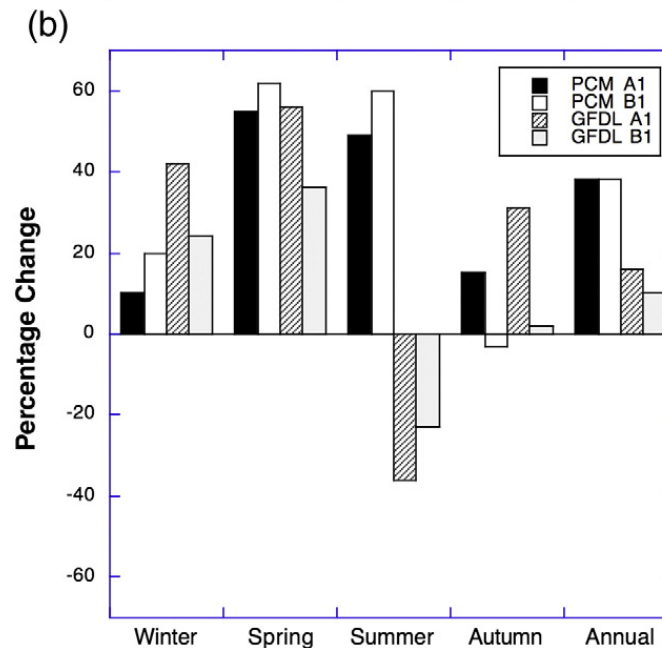
Precipitation Change (%)

Relationship between Changes in Mean and Extreme Precipitation (Chicago, late 21st Century)

Mean Precipitation



Extreme Precipitation
(wettest 5% of days)

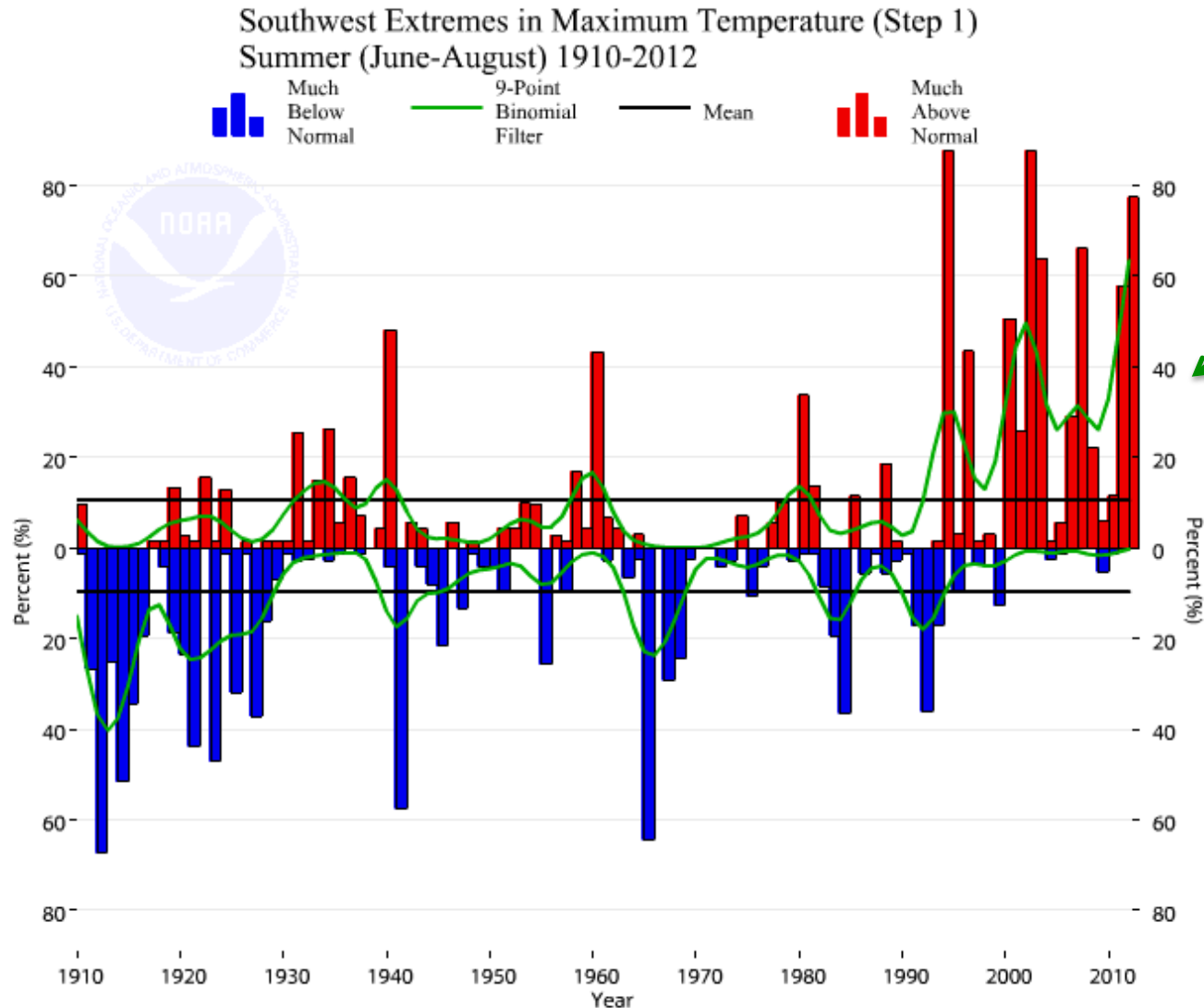


Increased heavy precipitation mirrors mean changes in every season

Back to the present day climate. . . .

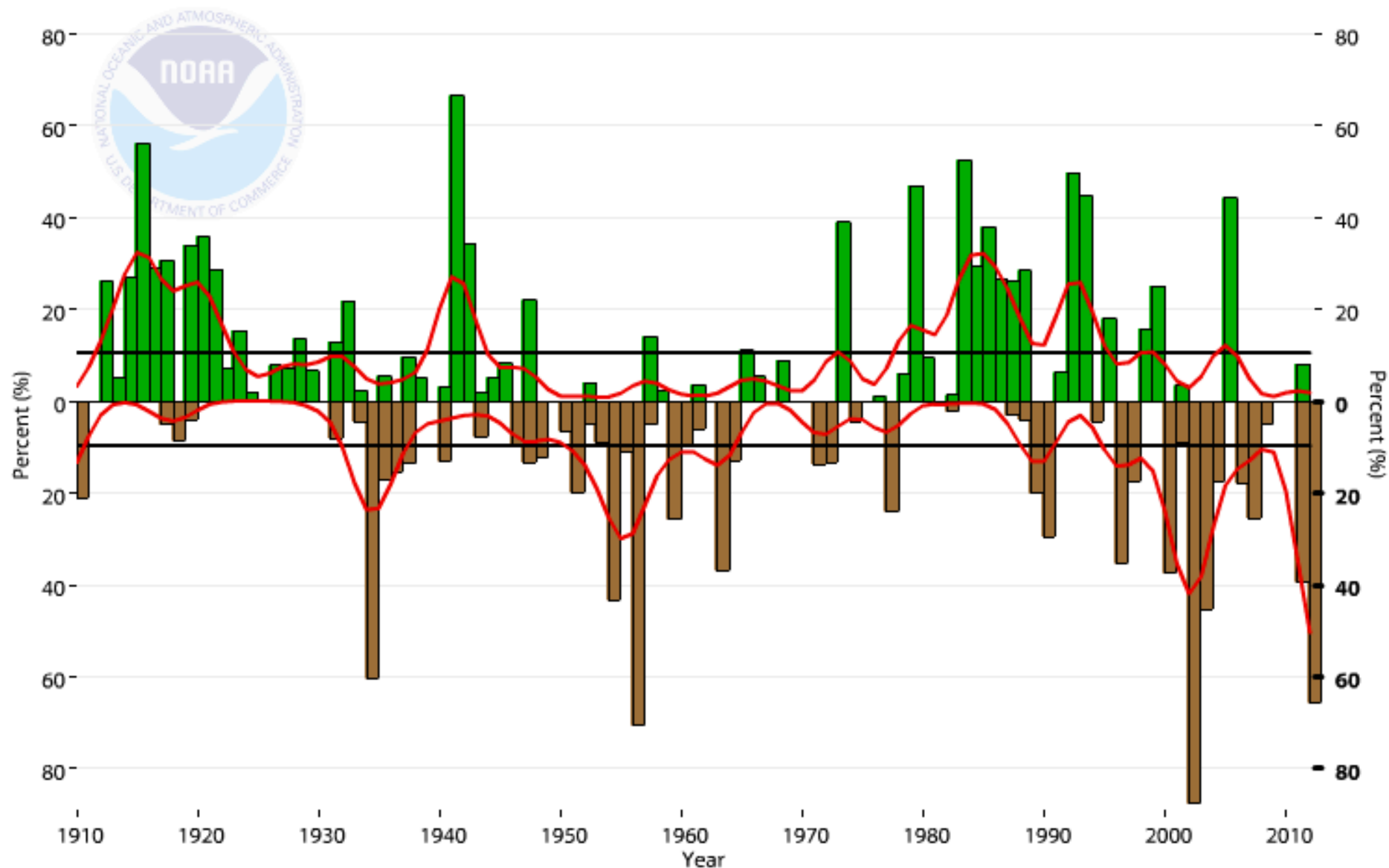
U. S. Climate Extremes Index

Summer Daily Maximum Temperature (Southwestern U. S.)



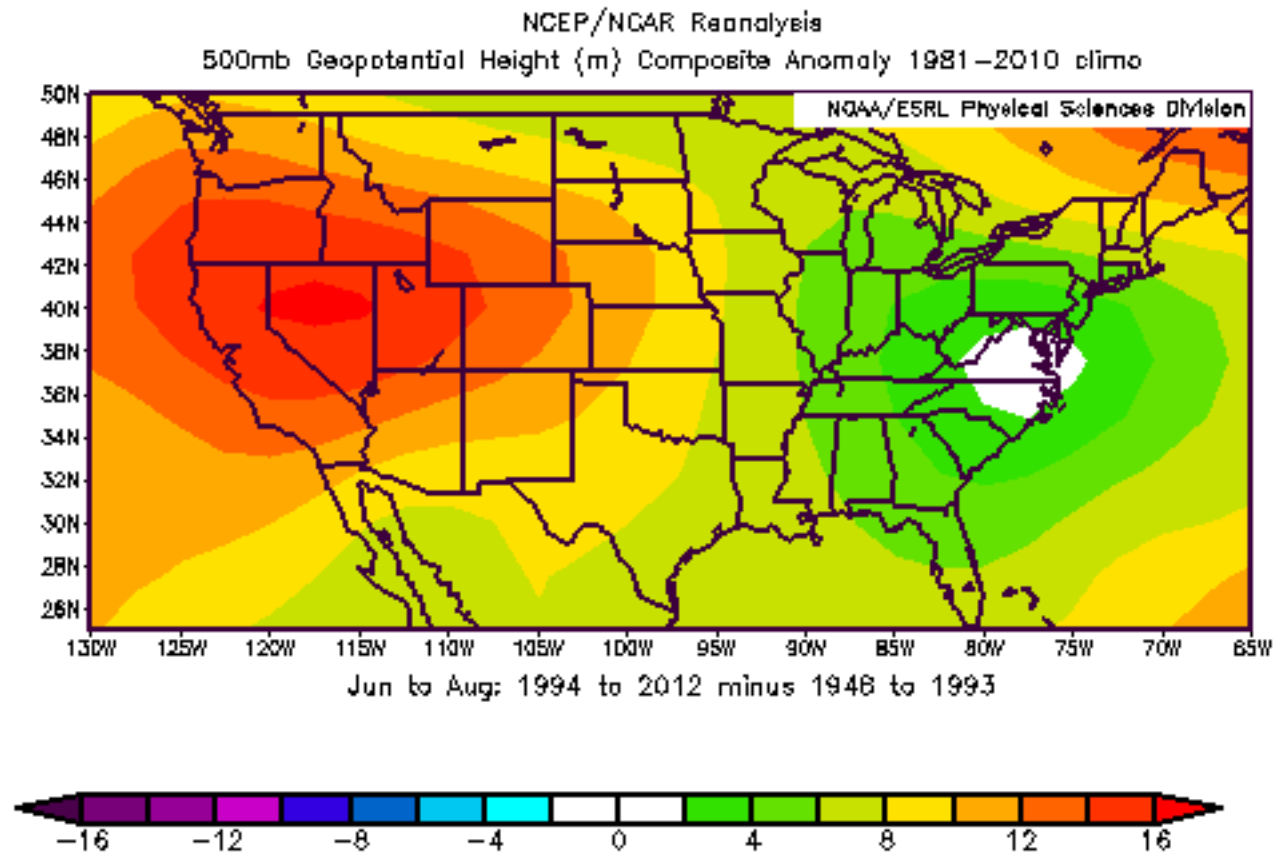
Southwest Extremes in PDSI (Step 3) Summer (June-August) 1910-2012

Much Below Normal 9-Point Binomial Filter Mean Much Above Normal



Summer Circulation Change since mid-1990s

500 hPa anomaly (post-1994 minus pre-1994)



Synoptic Evolution of Extreme Cold-Air Outbreaks (CAOs)

Midwestern CAOs

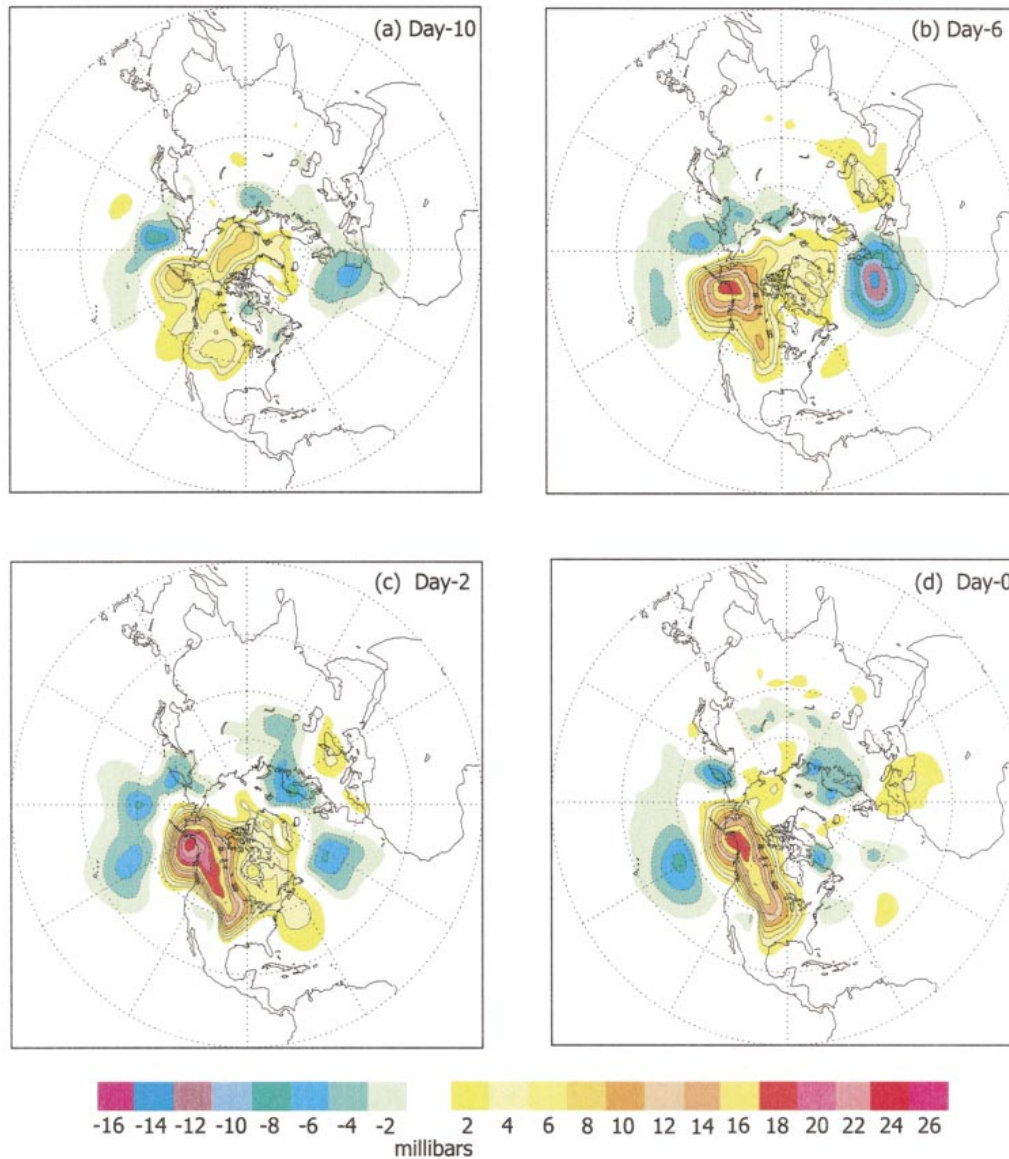
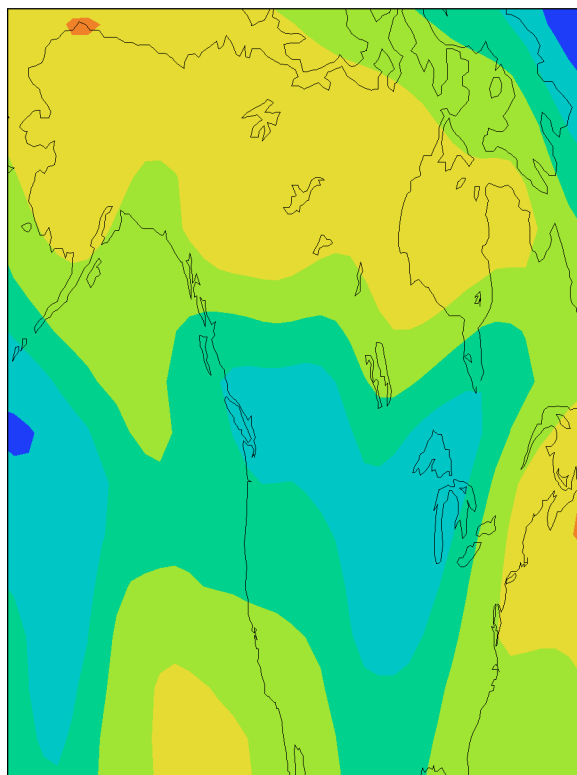


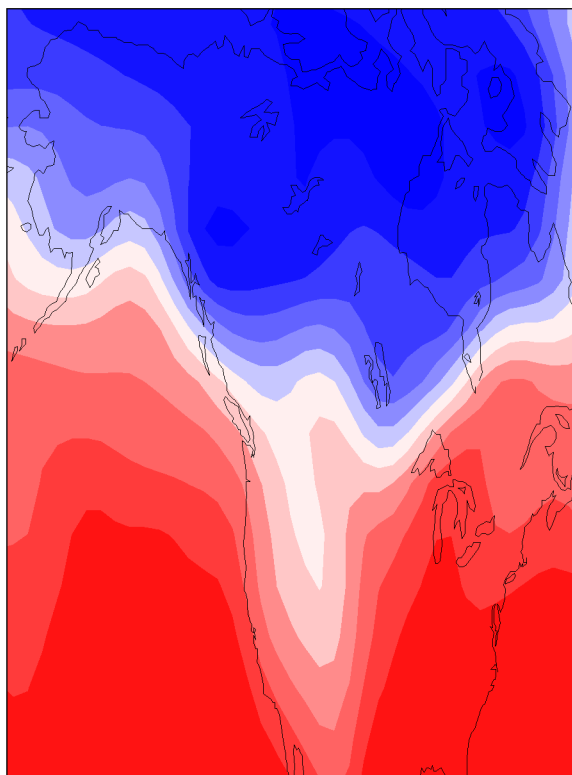
FIG. 7. SLP anomaly composites for (a) 10 days, (b) 6 days, (c) 2 days, and (d) 0 days prior to the dates of the 10 cold events in the Midwest (MW) region. Numbers below color bar denote the lower limit of each 2-mb range of magnitudes. Unshaded areas have magnitudes smaller than 2 mb.

CSM Simulation of a CAO (February 24 - March 8)

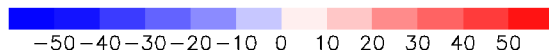
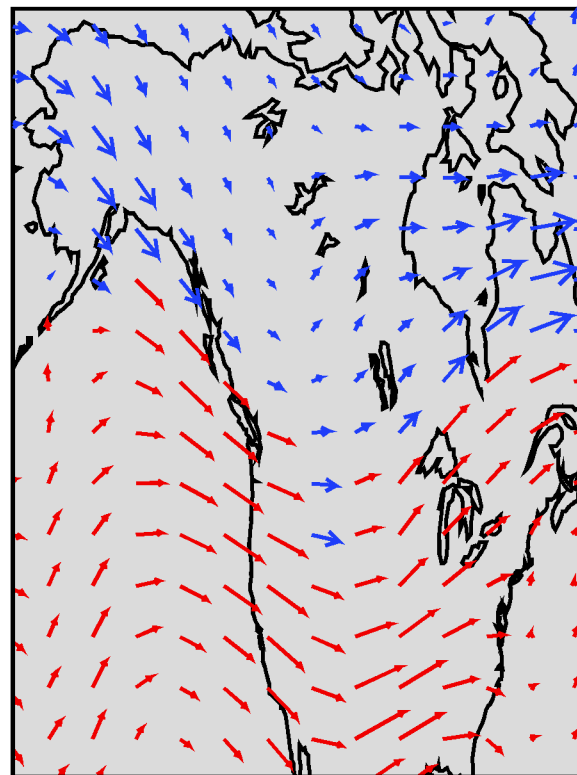
Day 1 SLP



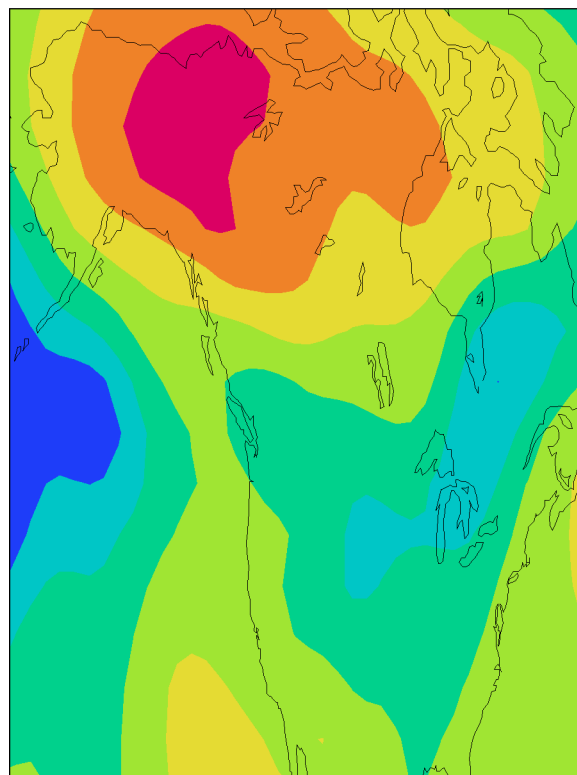
Day 1 Temperature



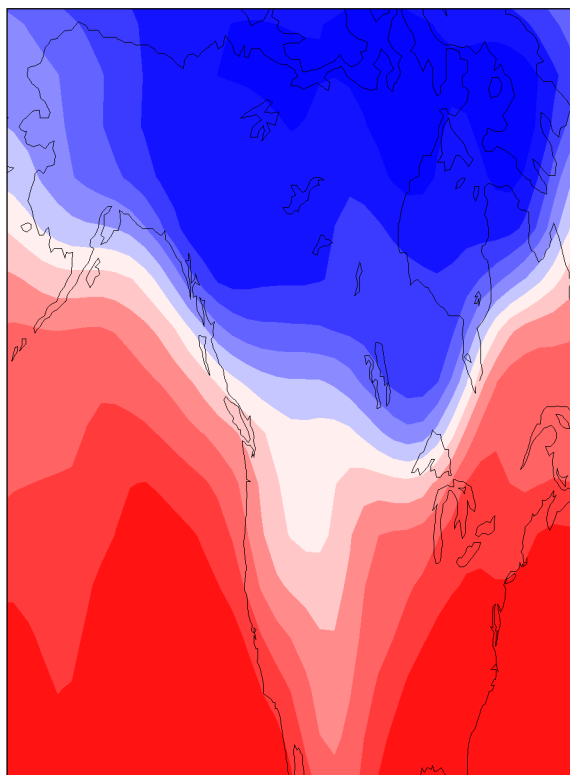
Day 1 500 hPa Wind



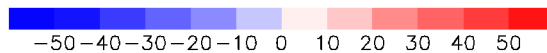
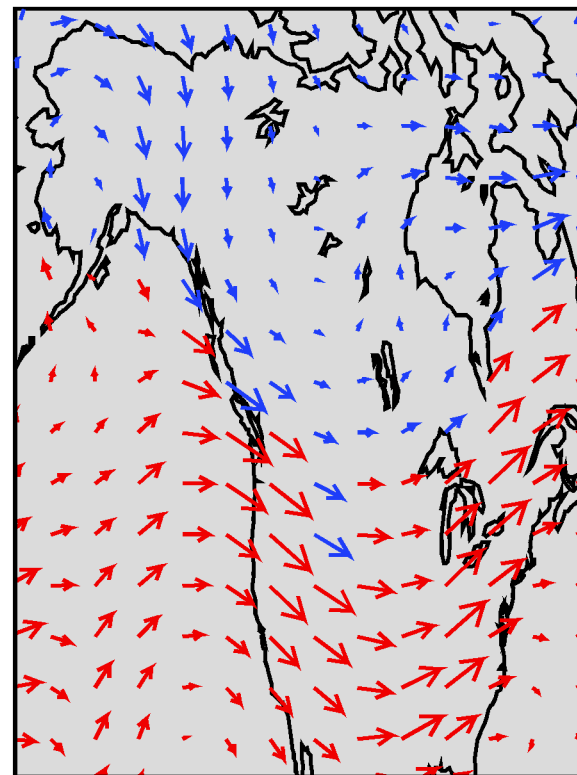
Day 2 SLP



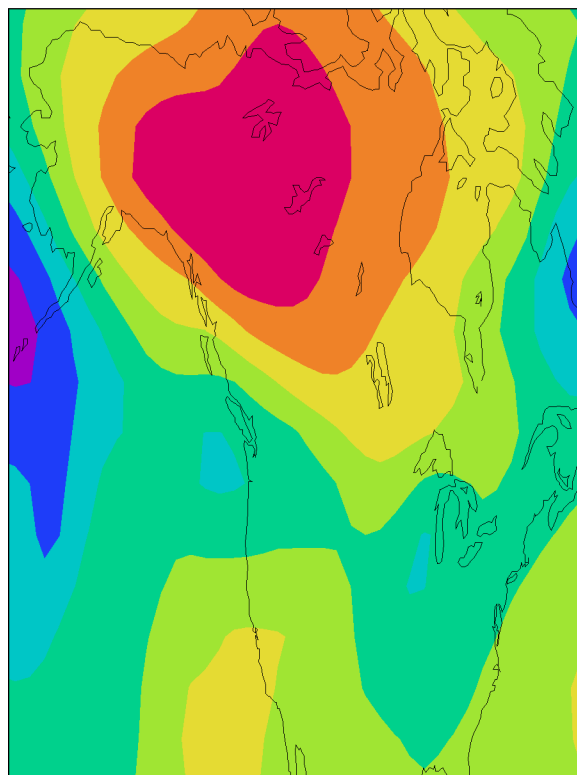
Day 2 Temperature



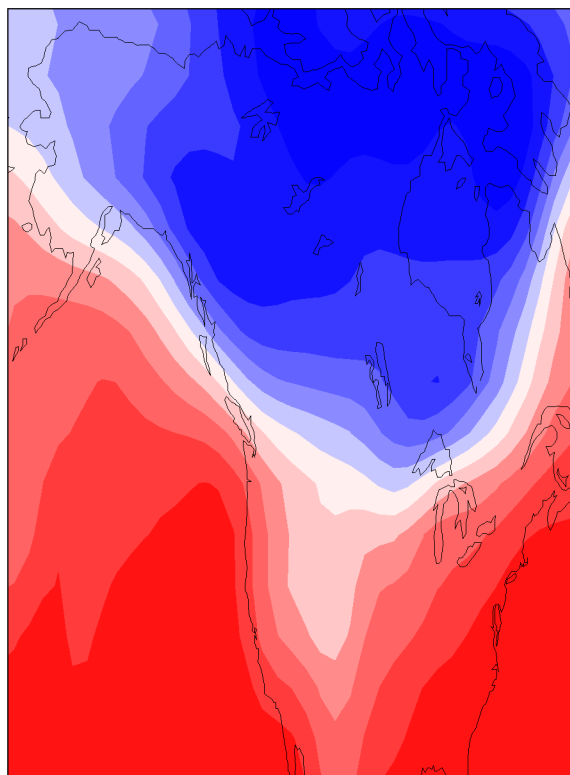
Day 2 500 hPa Wind



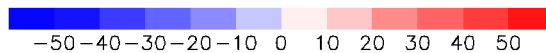
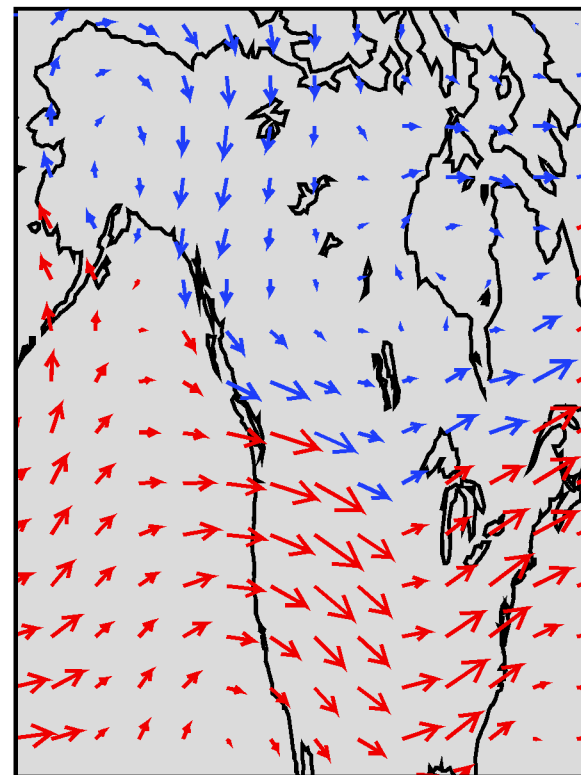
Day 3 SLP



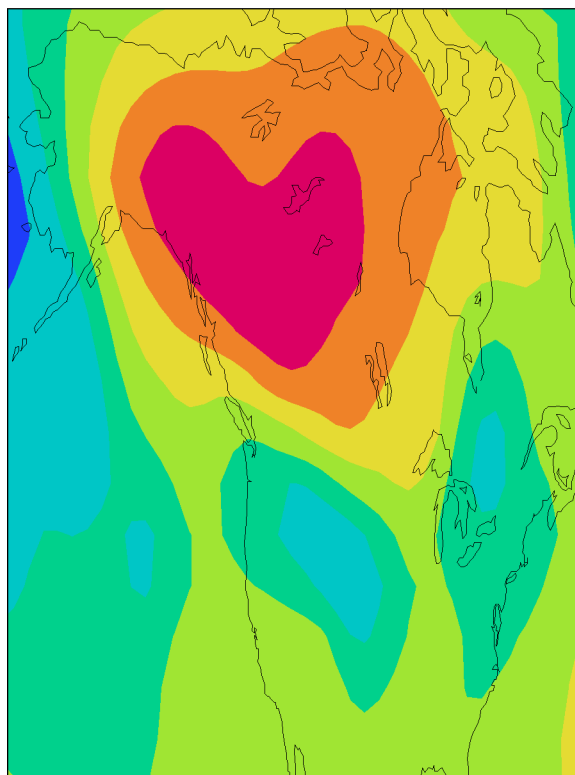
Day 3 Temperature



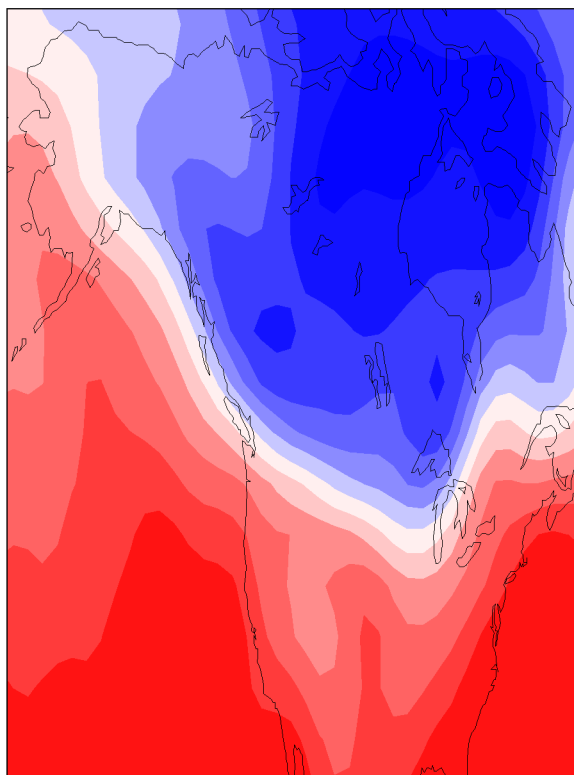
Day 3 500 hPa Wind



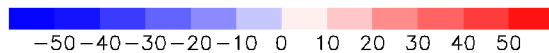
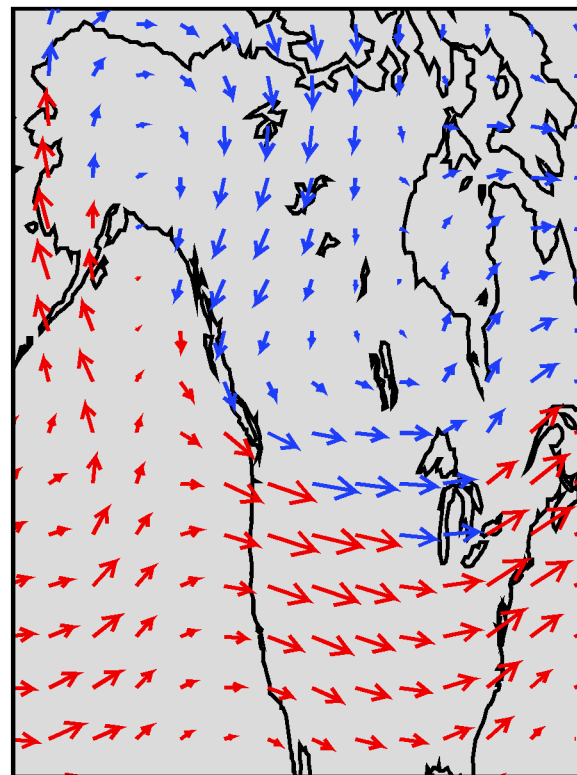
Day 4 SLP



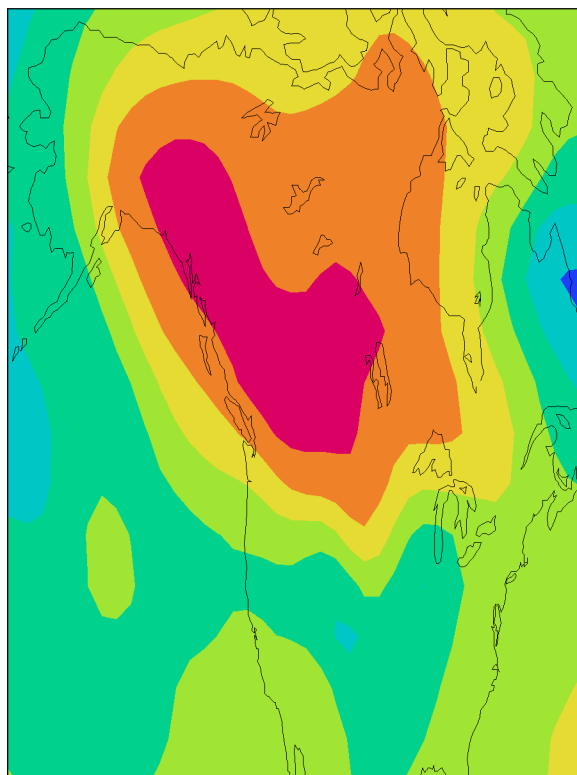
Day 4 Temperature



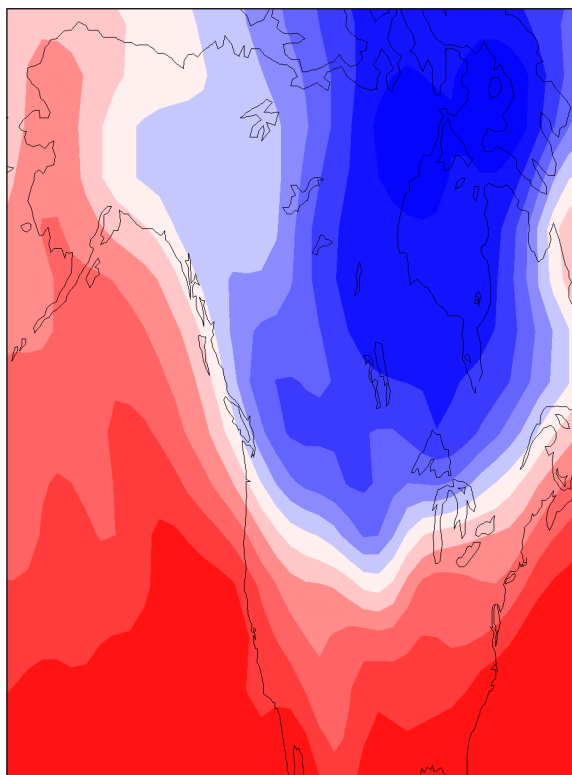
Day 4 500 hPa Wind



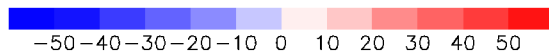
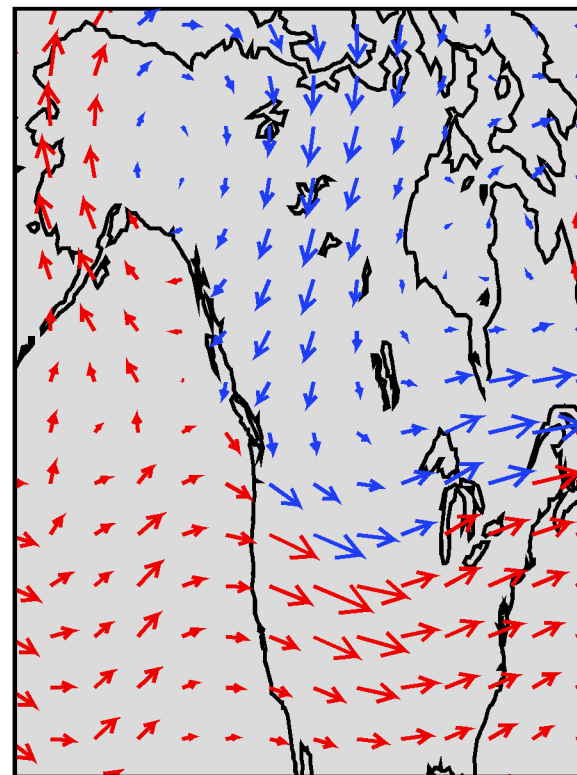
Day 5 SLP



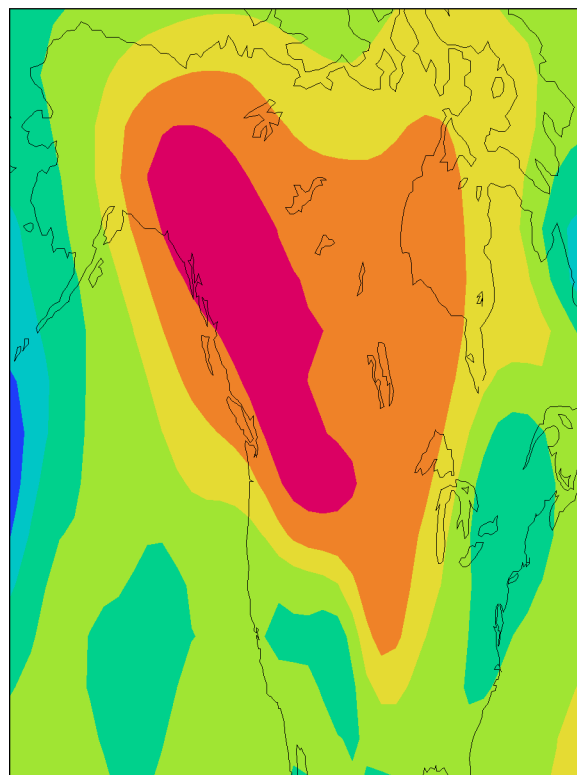
Day 5 Temperature



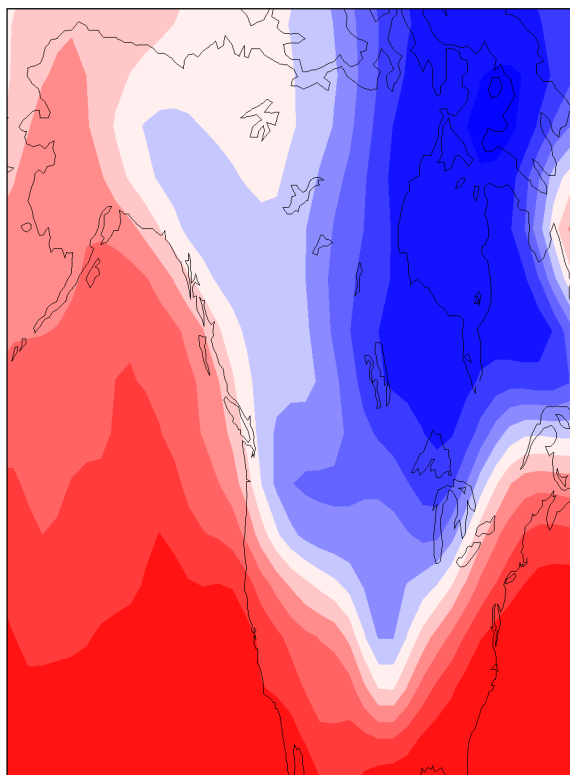
Day 5 500 hPa Wind



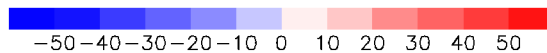
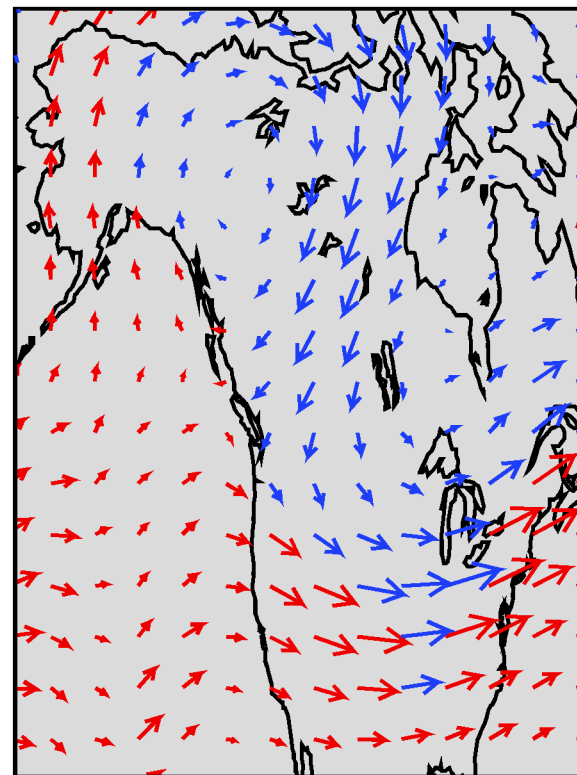
Day 6 SLP



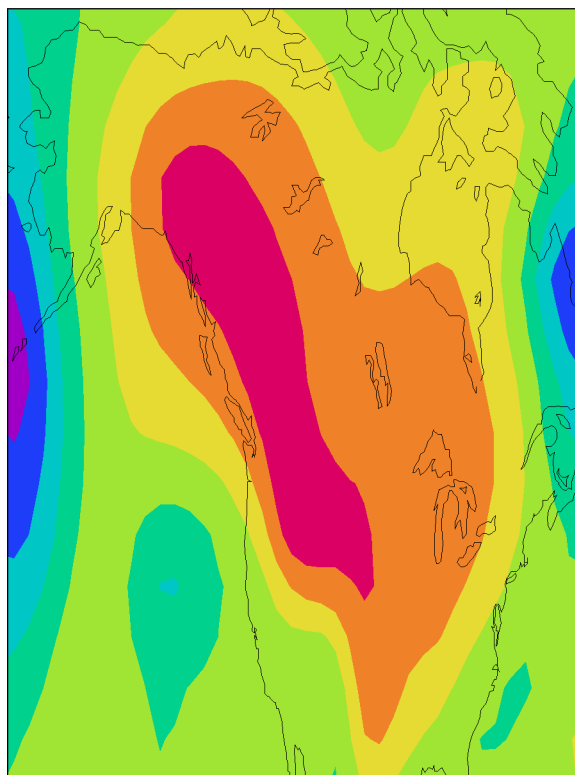
Day 6 Temperature



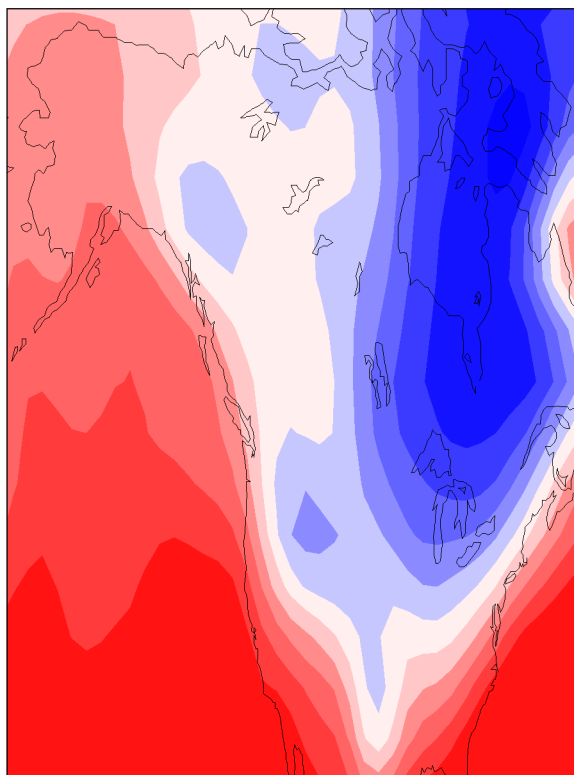
Day 6 500 hPa Wind



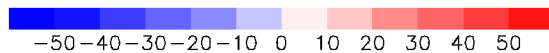
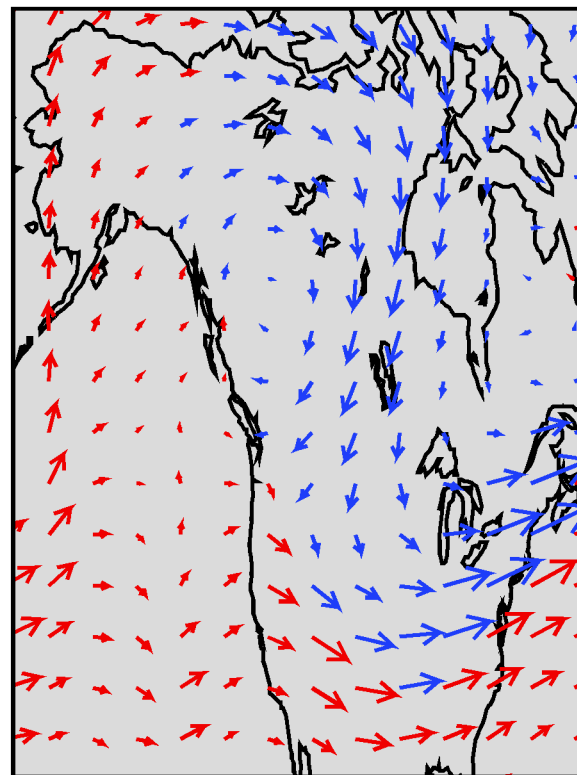
Day 7 SLP



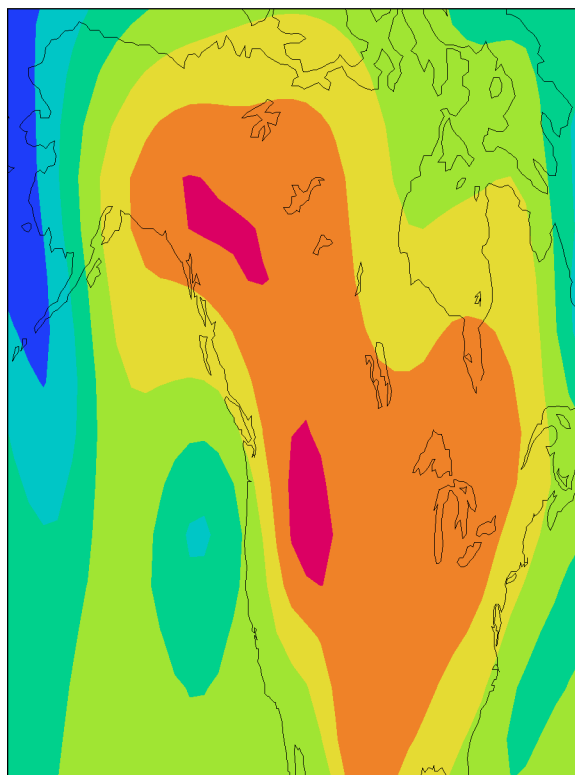
Day 7 Temperature



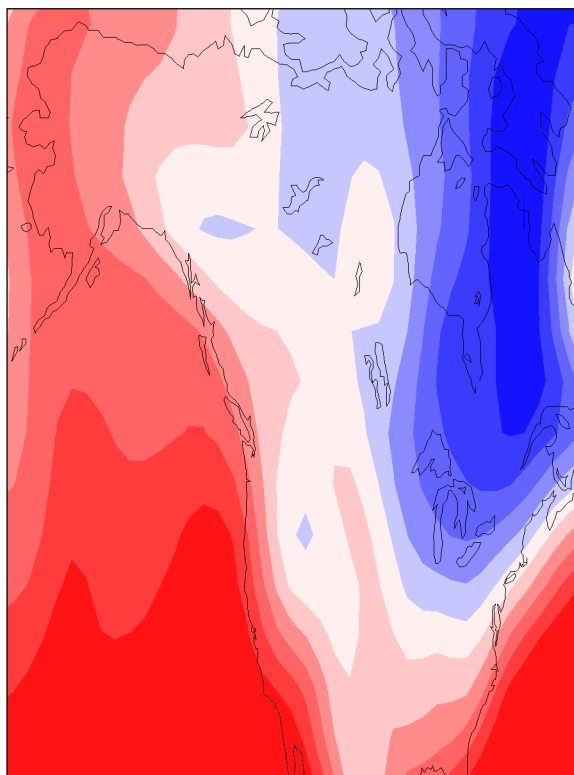
Day 7 500 hPa Wind



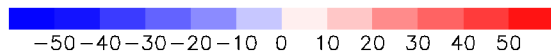
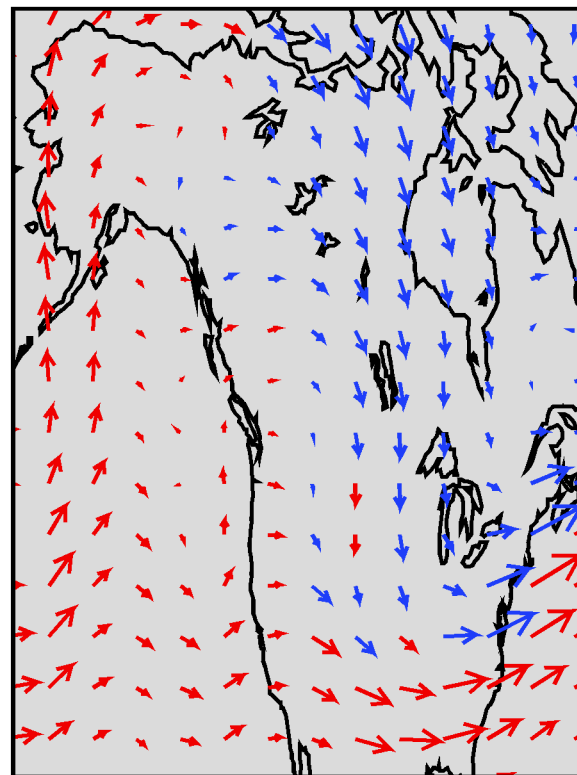
Day 8 SLP



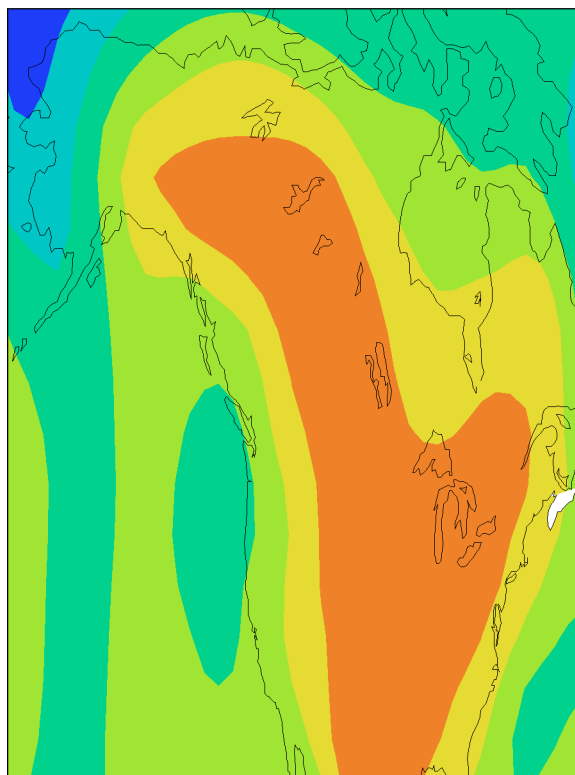
Day 8 Temperature



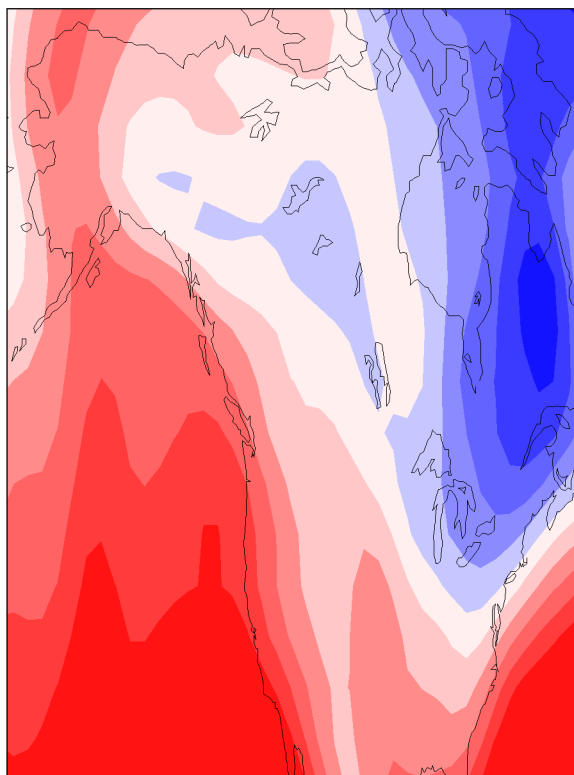
Day 8 500 hPa Wind



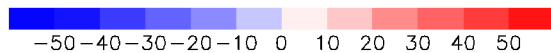
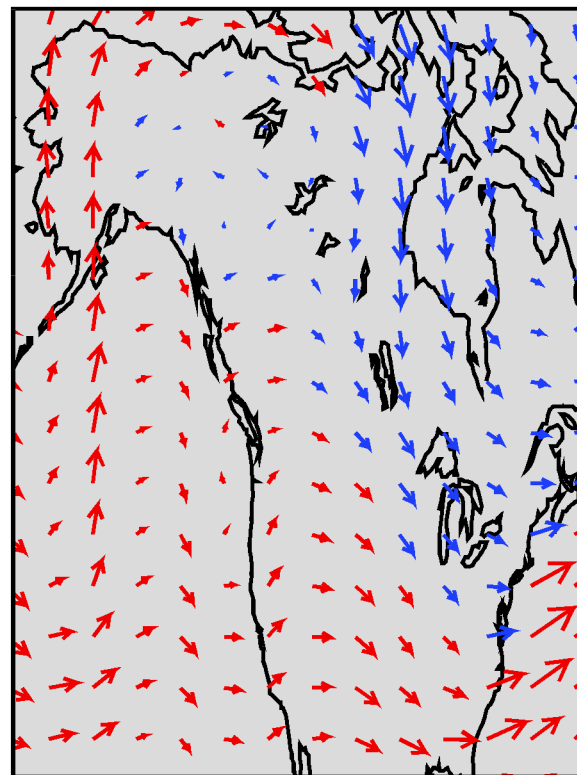
Day 9 SLP



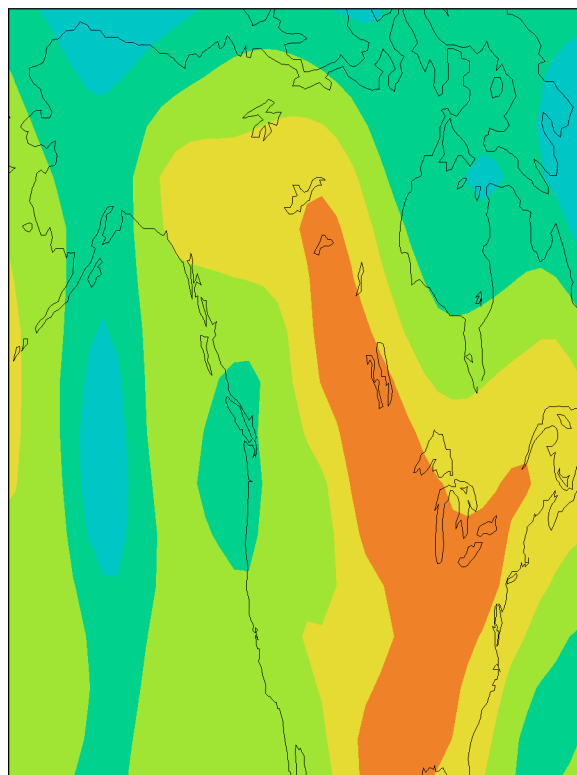
Day 9 Temperature



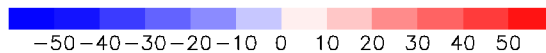
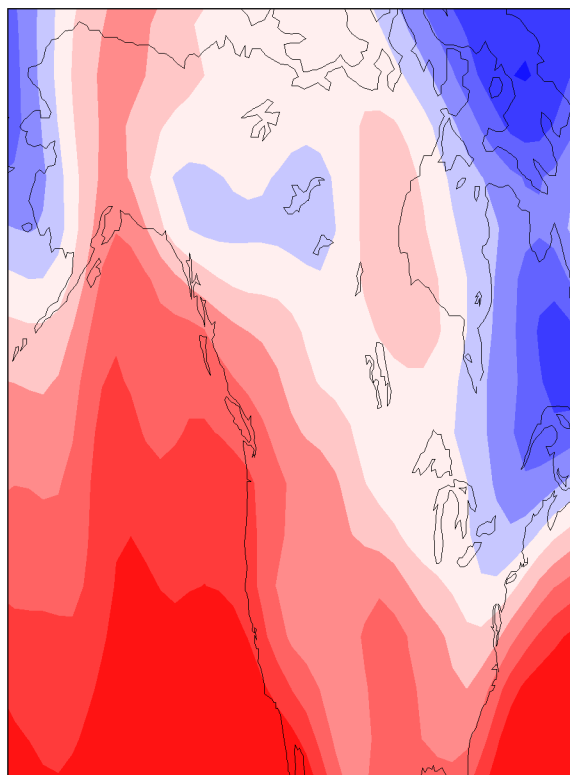
Day 9 500 hPa Wind



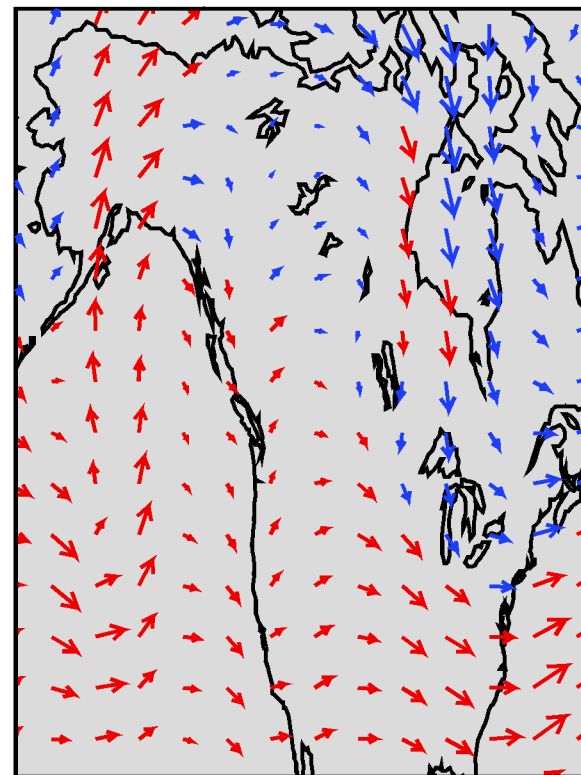
Day 10 SLP



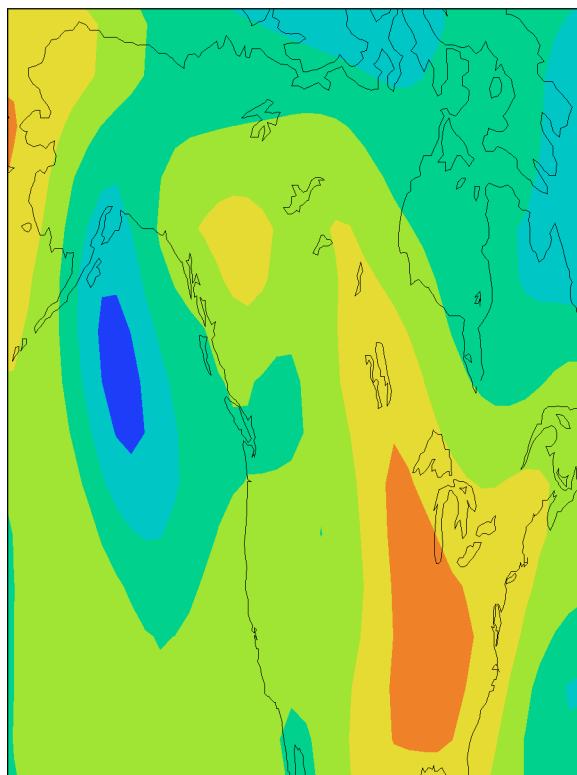
Day 10 Temperature



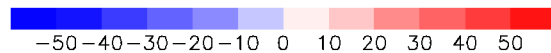
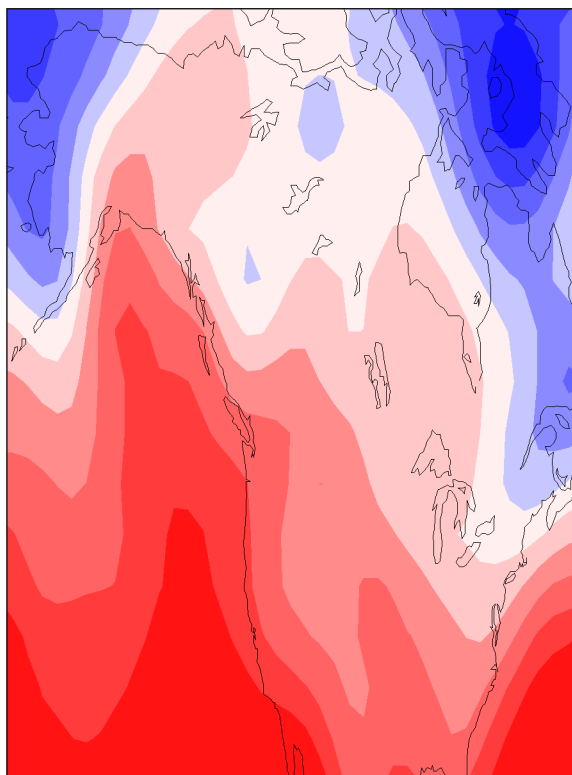
Day 10 500 hPa Wind



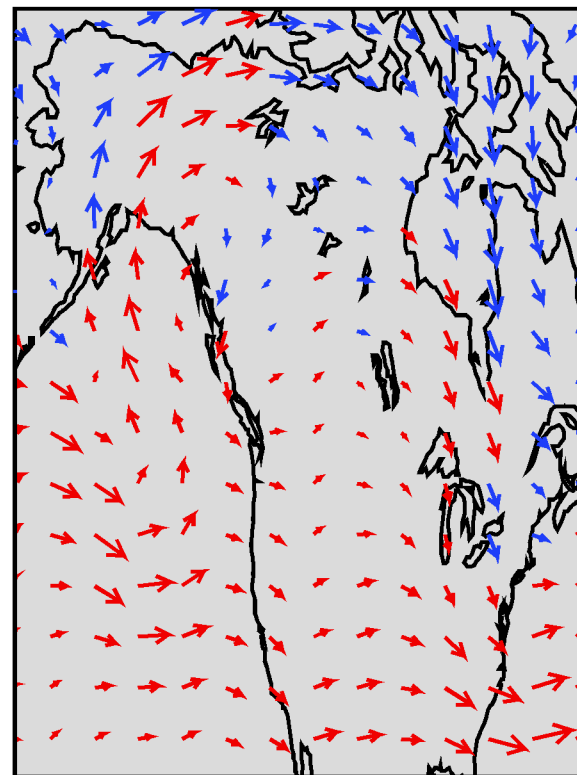
Day 11 SLP



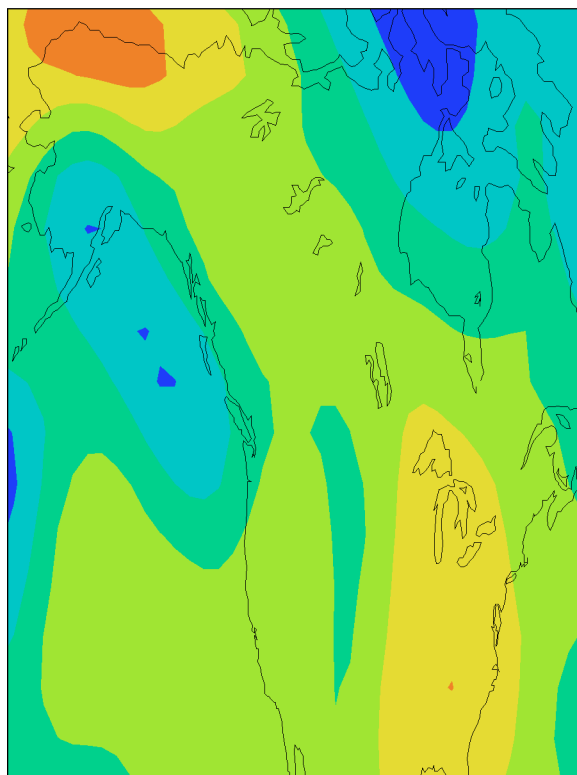
Day 11 Temperature



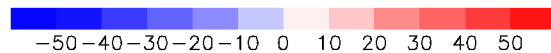
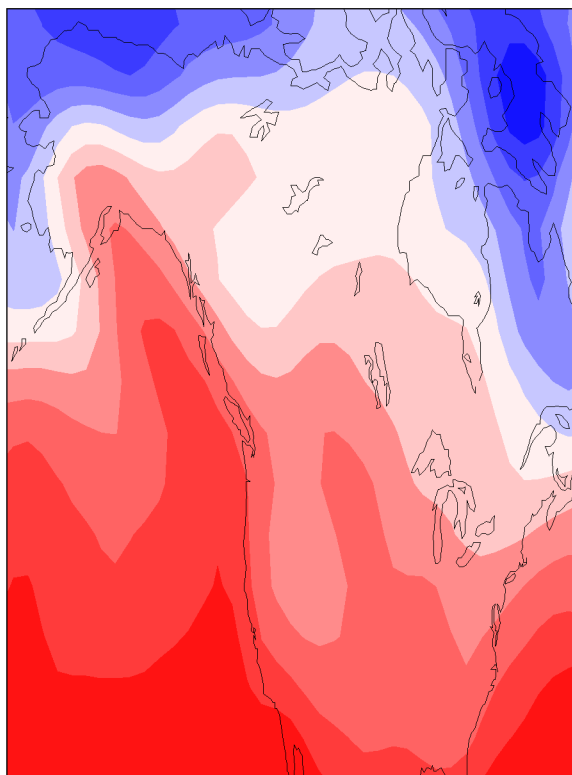
Day 11 500 hPa Wind



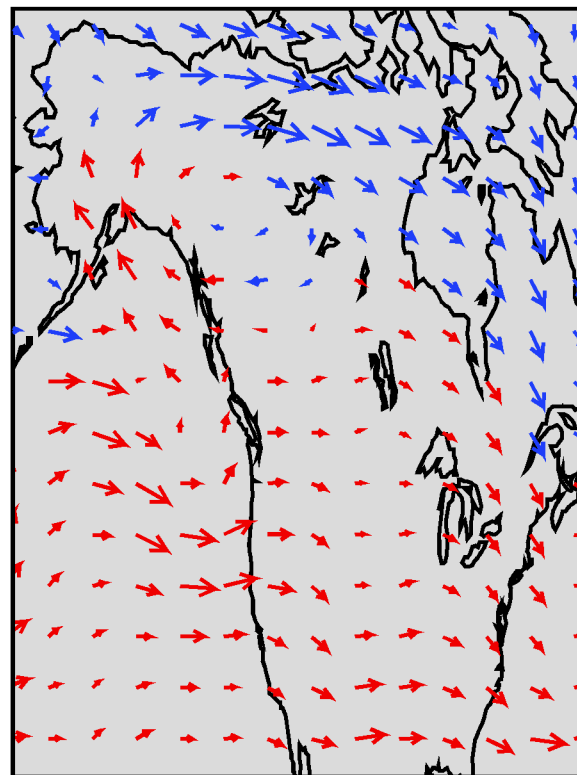
Day 12 SLP



Day 12 Temperature

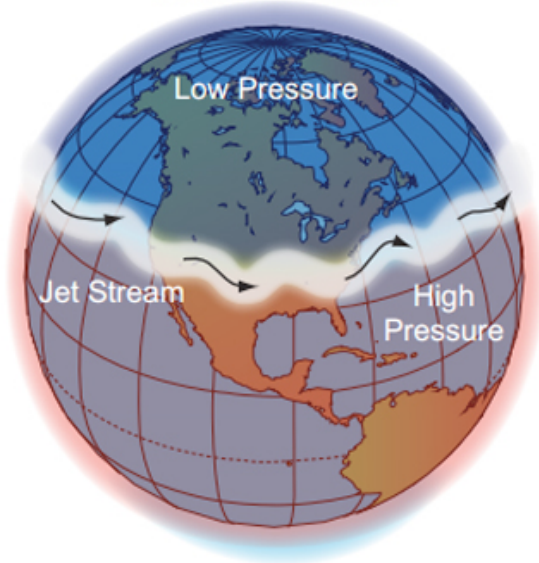


Day 12 500 hPa Wind

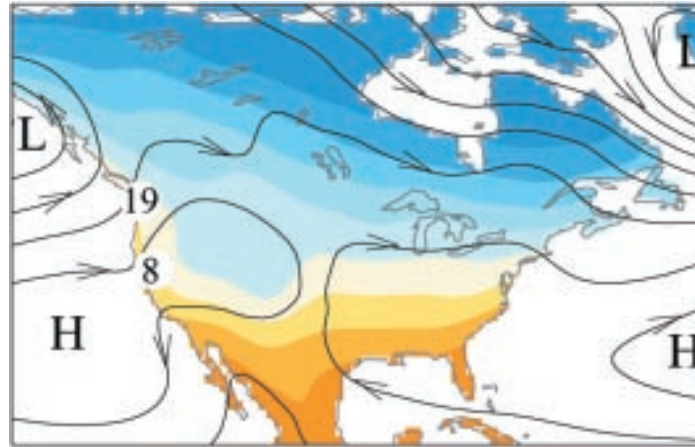


Role of the Arctic Oscillation in Extreme Cold-air Outbreaks

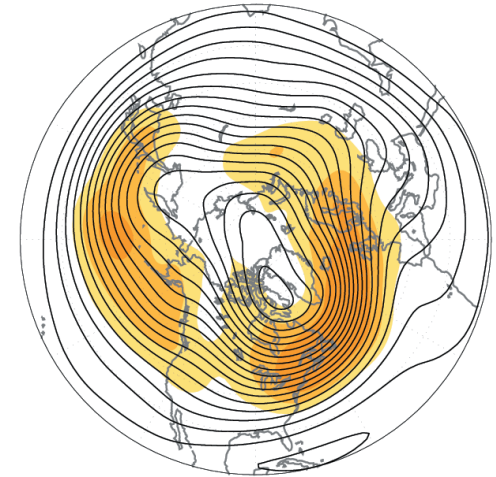
Positive Phase



Surface Response in + AO (SLP)

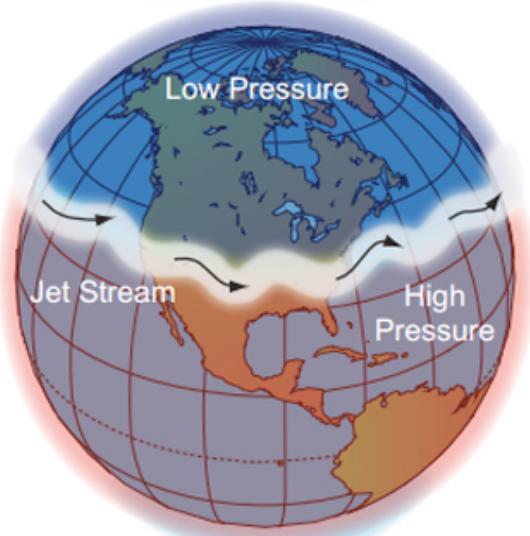


Response Aloft in + AO (Z500)

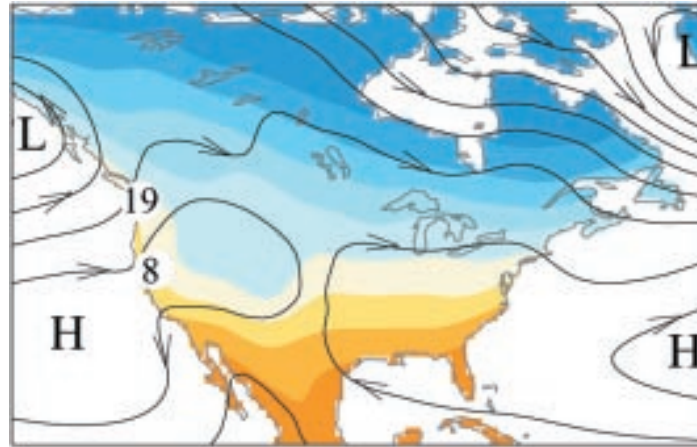


Role of the Arctic Oscillation in Extreme Cold-air Outbreaks

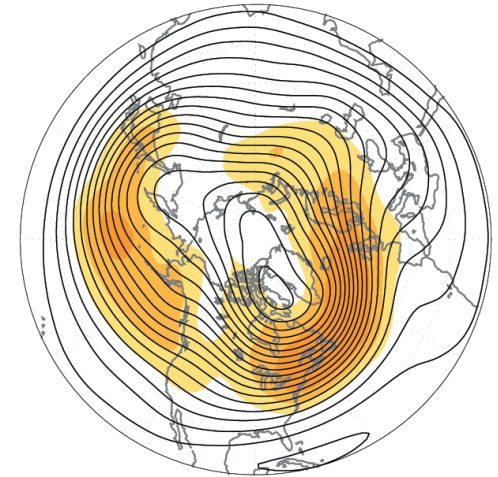
Positive Phase



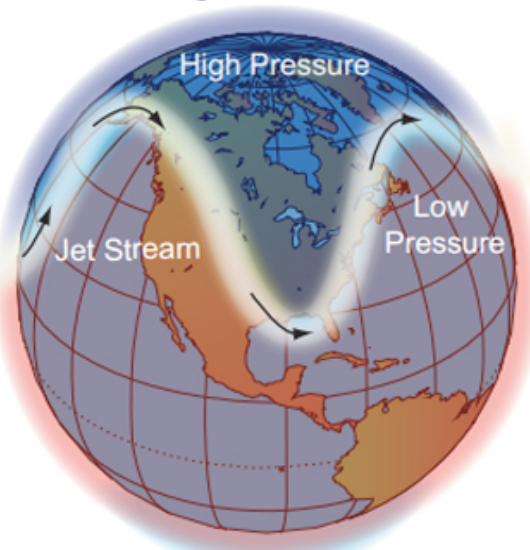
Surface Response in + AO (SLP)



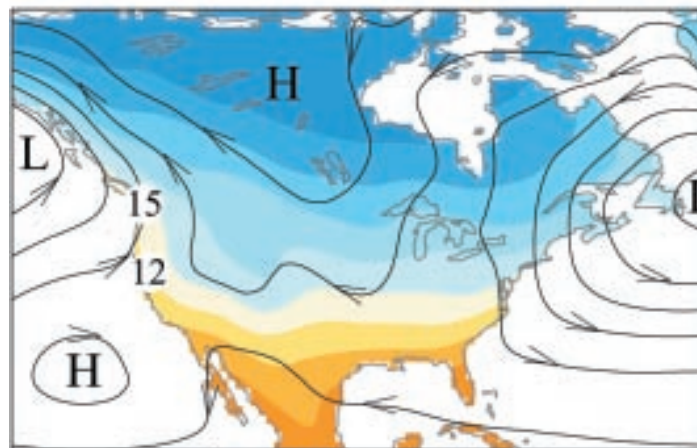
Response Aloft in + AO (Z500)



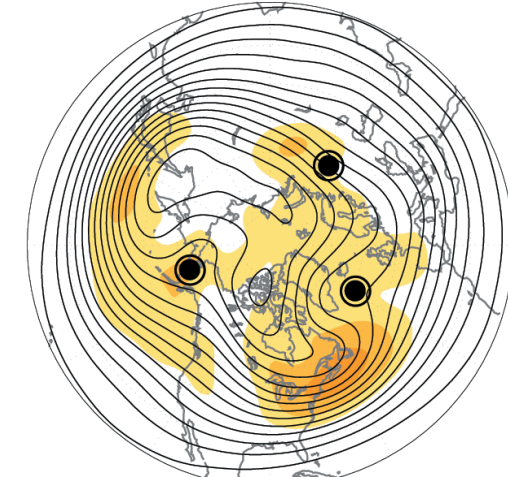
Negative Phase



Surface Response in - AO (SLP)

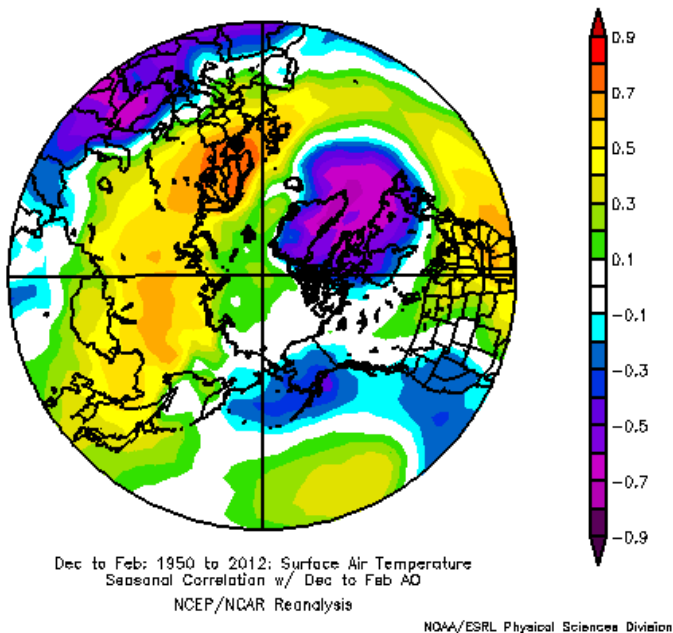


Response Aloft in - AO (Z500)



Arctic Oscillation Effect on Winter Temperature

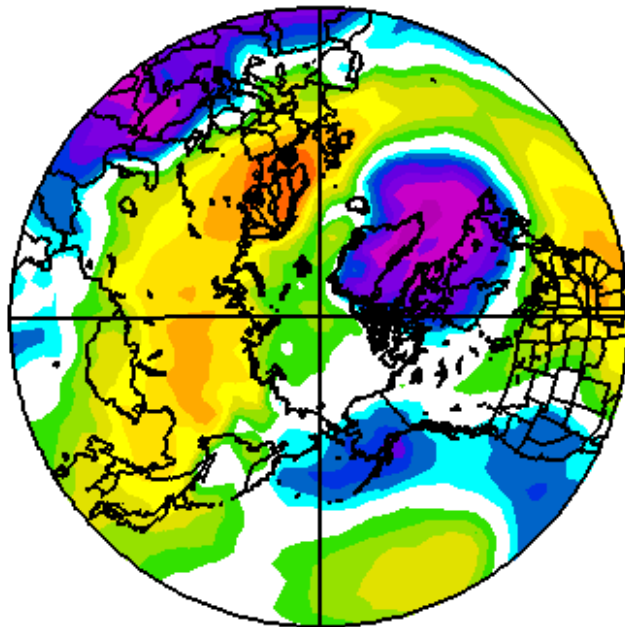
*Mean Surface
Temperatures*



*Cold winters
typical in AO
negative phase
over mid-high
latitude land*

Arctic Oscillation Effect on Winter Temperature

*Mean Surface
Temperatures*

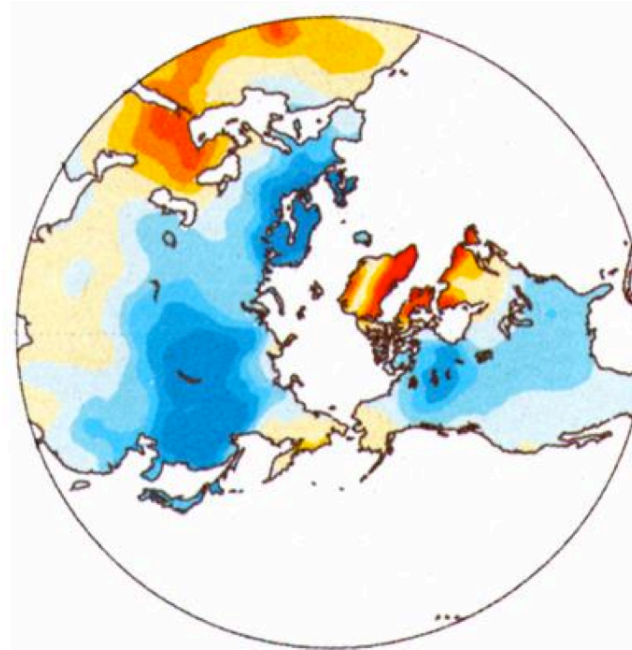


Dec to Feb: 1950 to 2012: Surface Air Temperature
Seasonal Correlation w/ Dec to Feb AO

NCEP/NCAR Reanalysis

NOAA/ESRL Physical Sciences Division

*Extreme Surface
Temperatures*



Daily Cold
Extremes:
Neg AO/Pos AO

*Cold winters typical in
AO negative phase over
mid-high latitude land*

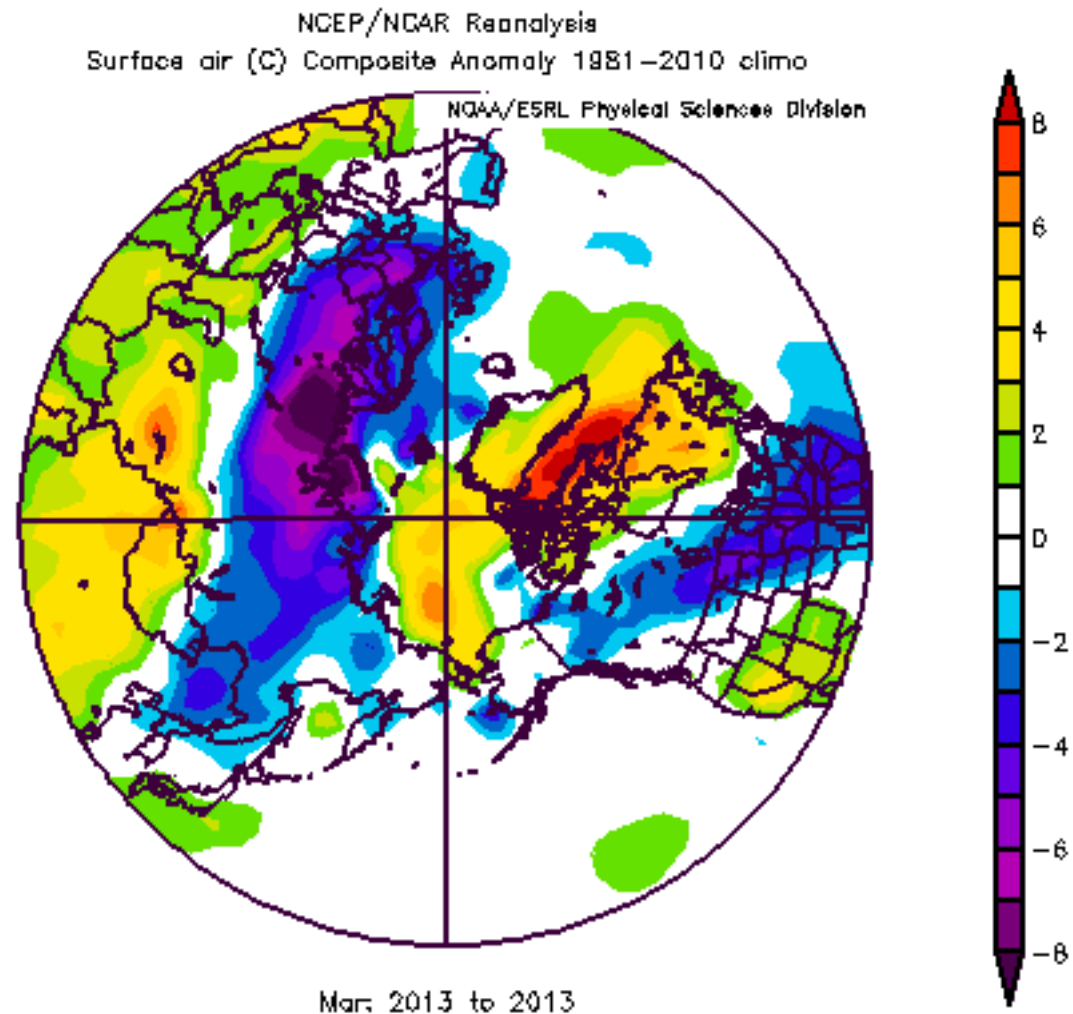
Thompson and Wallace
Science, 2001

Fig. 4. The signature of the NAM in the frequency of occurrence of cold events (daily minimum temperature < 1.5 SD below the climatological mean) based on 6-hourly JFM data, 1958–1997, from the NCEP/NCAR Reanalysis (10). (Top) The ratios of the number of cold events in the low NAM-index sample to the number of cold events in the high NAM-index sample. (Bottom) As in the top panel, but for the ratios attributable to shifts in mean temperature (19). Shading is drawn for ratios $\pm 2:1$, $3:1$, $4:1$, and $6:1$. Blue shading indicates more cold events under low index conditions.

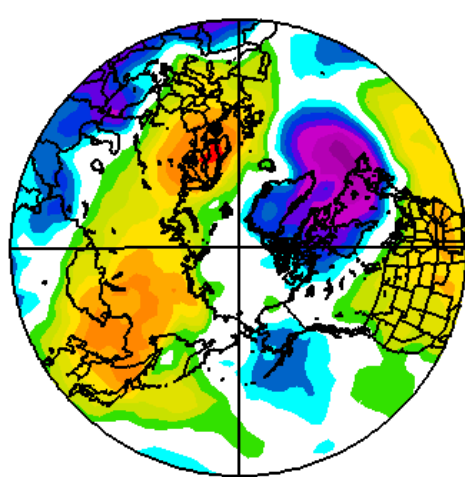
Figure 1 is a line graph showing the AO Index (Y-axis, ranging from -6 to 4) over time (X-axis, labeled 'Month', spanning from October 2012 to April 2013). The graph displays a highly variable time series. A horizontal blue line is drawn at AO Index = 0. A vertical yellow shaded region highlights the period from late February to mid-March 2013, corresponding to the time of the 2013 Superstorm. During this period, the AO Index drops sharply, reaching a minimum of approximately -5.5 in early March 2013, before recovering to around -1.0 by mid-March.

Record low March
AO index (-3.2)

March 2013 Surface Temperature Anomalies

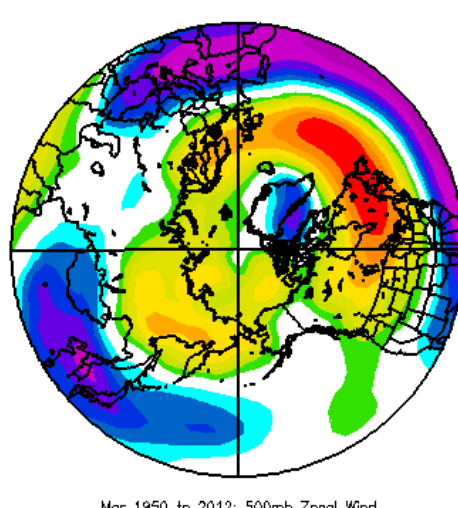


March AO Correlations with Tsfc, 500mb U, 500mb V



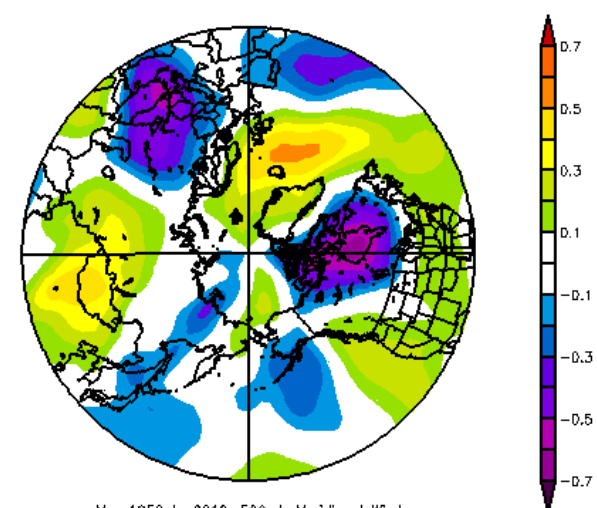
Mar 1950 to 2012: Surface Air Temperature
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

NOAA/ESRL Physical Sciences Division



Mar 1950 to 2012: 500mb Zonal Wind
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

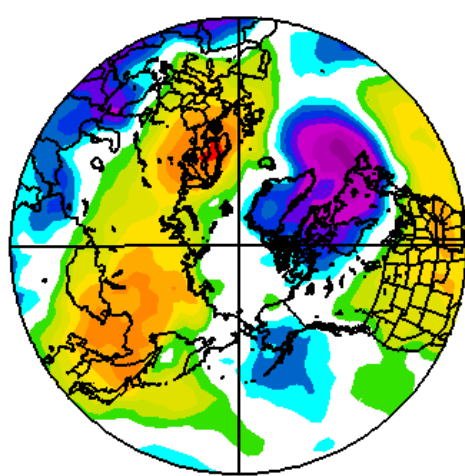
NOAA/ESRL Physical Sciences Division



Mar 1950 to 2012: 500mb Meridional Wind
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

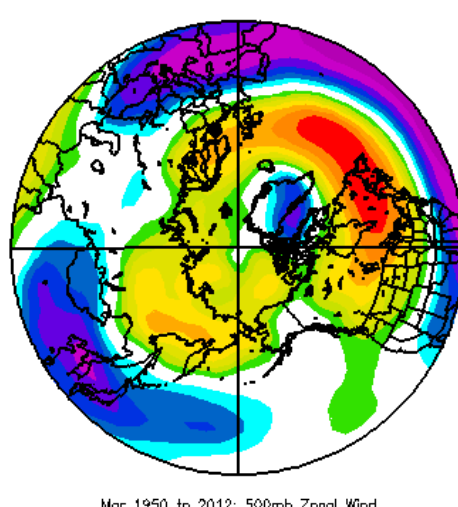
NOAA/ESRL Physical Sciences Division

March AO Correlations with Tsfc, 500mb U, 500mb V



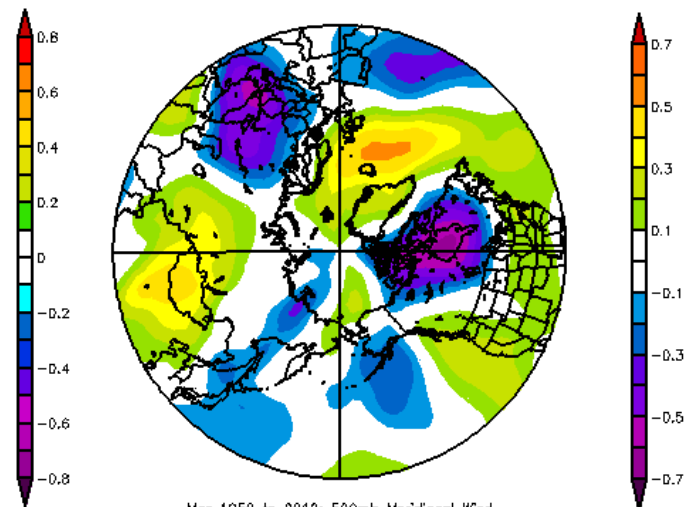
Mar 1950 to 2012: Surface Air Temperature
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

NOAA/ESRL Physical Sciences Division



Mar 1950 to 2012: 500mb Zonal Wind
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

NOAA/ESRL Physical Sciences Division

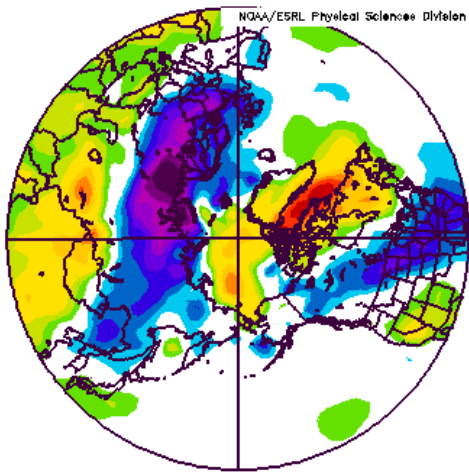


Mar 1950 to 2012: 500mb Meridional Wind
Seasonal Correlation w/ Mar AO
NCEP/NCAR Reanalysis

NOAA/ESRL Physical Sciences Division

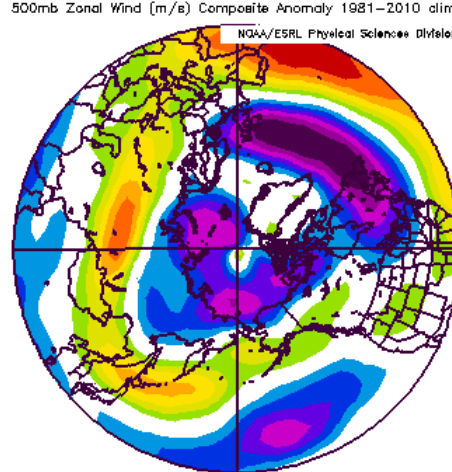
March 2013 anomalies of Tsfc, 500mb U, 500mb V

NCEP/NCAR Reanalysis
Surface air (C) Composite Anomaly 1981–2010 clima
NOAA/ESRL Physical Sciences Division



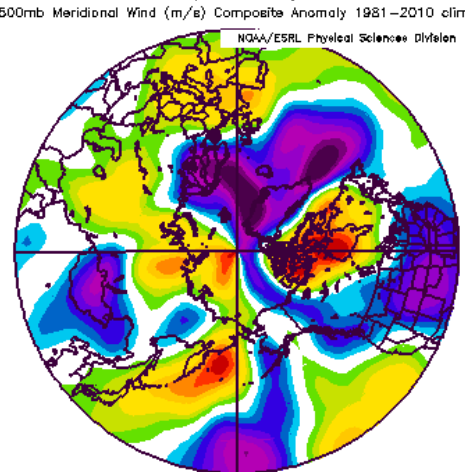
Mar: 2013 to 2013

NCEP/NCAR Reanalysis
500mb Zonal Wind (m/s) Composite Anomaly 1981–2010 clima
NOAA/ESRL Physical Sciences Division



Mar: 2013 to 2013

NCEP/NCAR Reanalysis
500mb Meridional Wind (m/s) Composite Anomaly 1981–2010 clima
NOAA/ESRL Physical Sciences Division

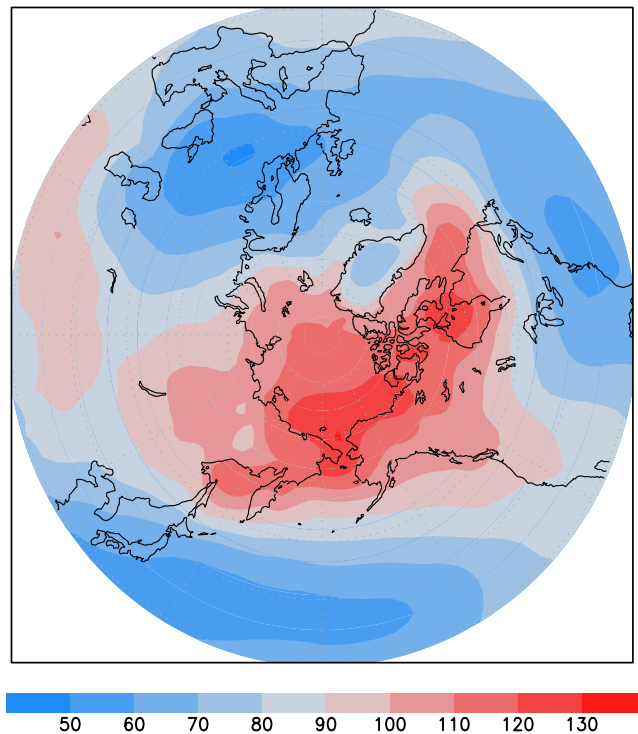


Mar: 2013 to 2013

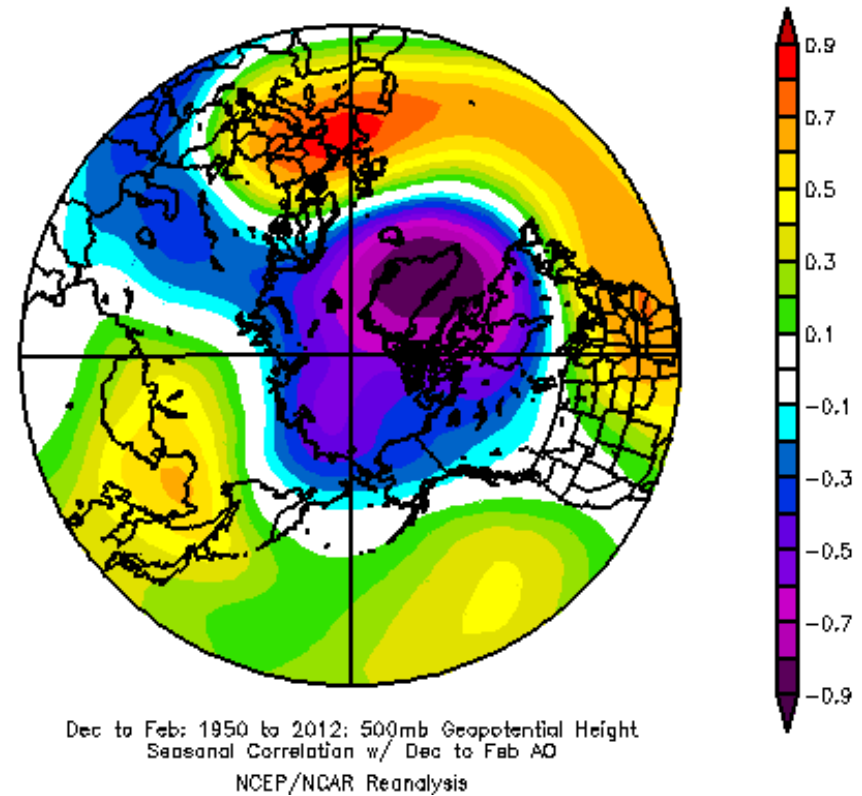
Projected Winter Changes resemble Negative Phase of AO

Future Height Changes Aloft

Change in 500 hPa Heights DJF 4xCO₂



Present-day AO Relationship (500 hPa Heights)

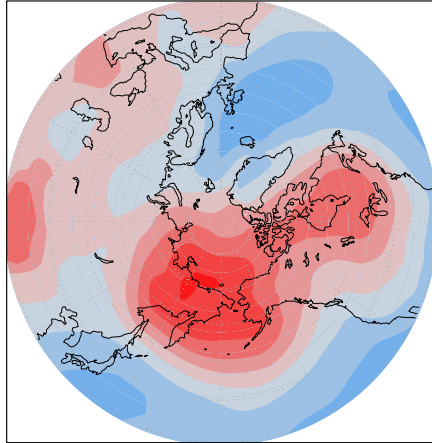


Should we expect slower and more meandering (AO-like) circulation pattern in future?

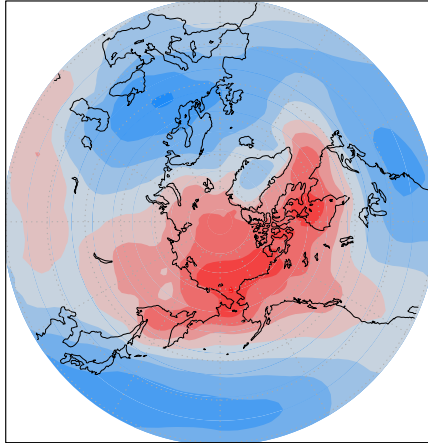
Upper-air Circulation Response in CCSM4's 4xCO₂ Simulation

500 hPa Heights by Season

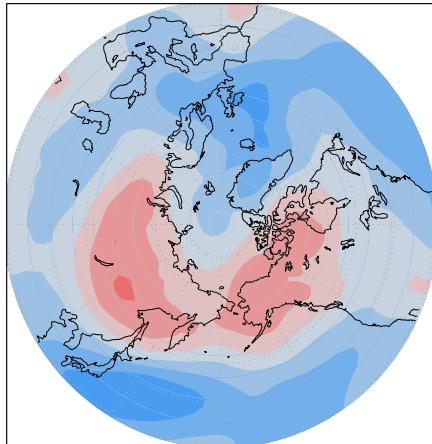
Change in 500 hPa Heights SON 4xCO₂



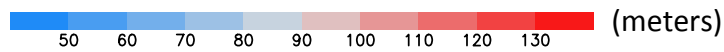
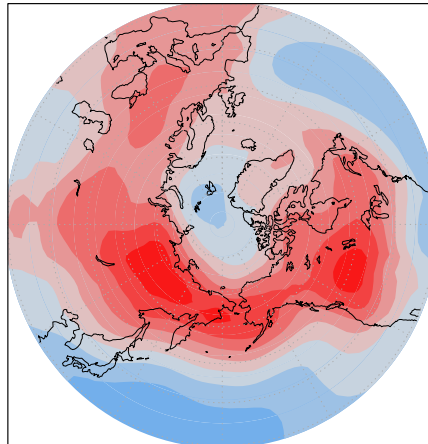
Change in 500 hPa Heights DJF 4xCO₂



Change in 500 hPa Heights MAM 4xCO₂



Change in 500 hPa Heights JJA 4xCO₂



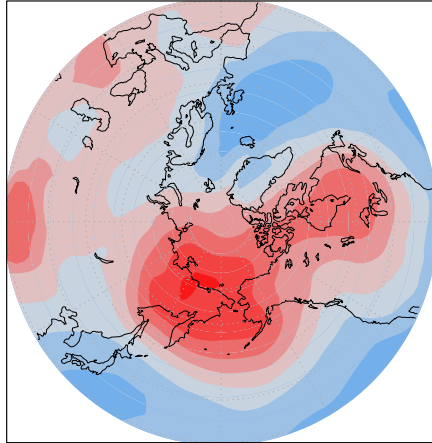
Autumn-Winter ridging
largely over Arctic Ocean
(sea ice loss)

Spring-Summer ridging
over high-latitude land
(snow cover loss,
lower land heat capacity)

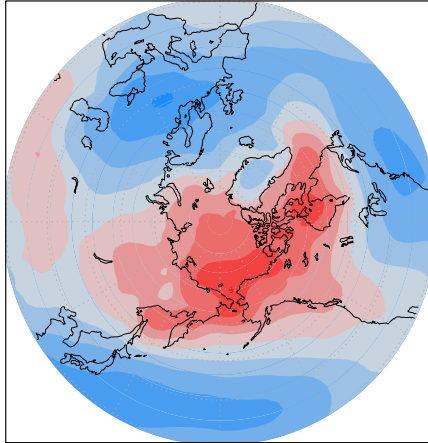
Upper-air Circulation Response in CCSM4's 4xCO₂ Simulation

500 hPa Height Changes by Season

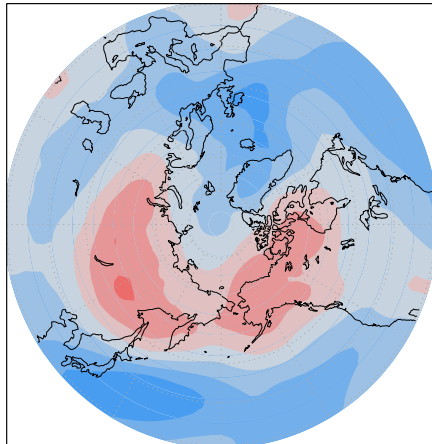
Change in 500 hPa Heights SON 4xCO₂



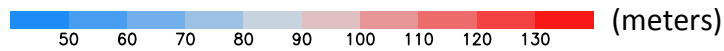
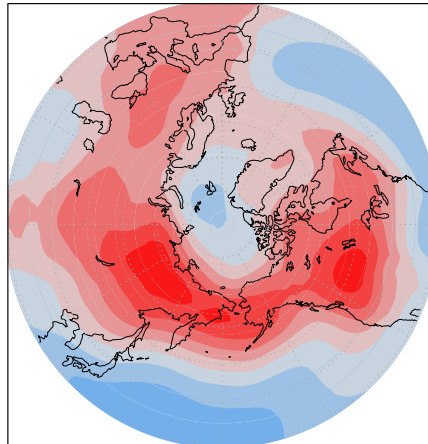
Change in 500 hPa Heights DJF 4xCO₂



Change in 500 hPa Heights MAM 4xCO₂

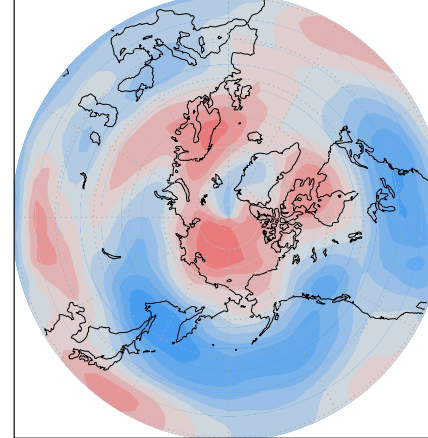


Change in 500 hPa Heights JJA 4xCO₂

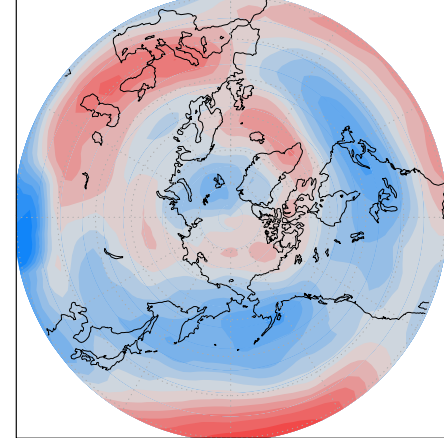


500 hPa Zonal Wind Changes by Season

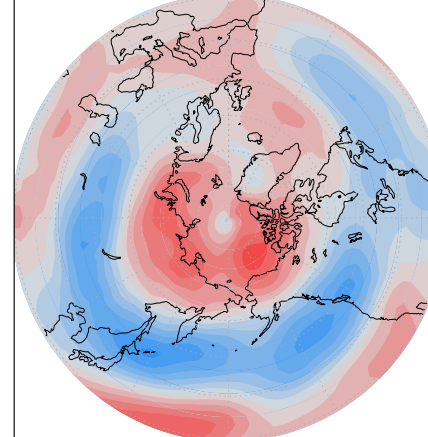
Change in 500 hPa Zonal Wind SON 4xCO₂



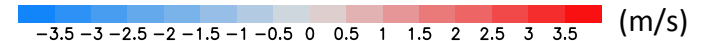
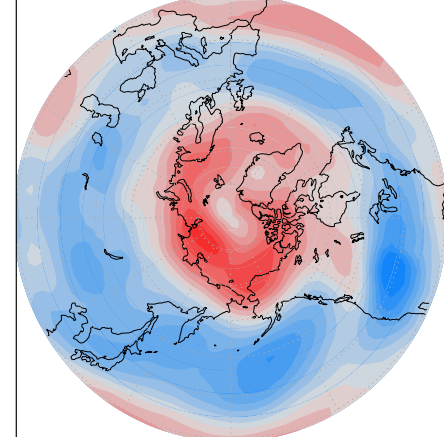
Change in 500 hPa Zonal Wind DJF 4xCO₂



Change in 500 hPa Zonal Wind MAM 4xCO₂



Change in 500 hPa Zonal Wind JJA 4xCO₂



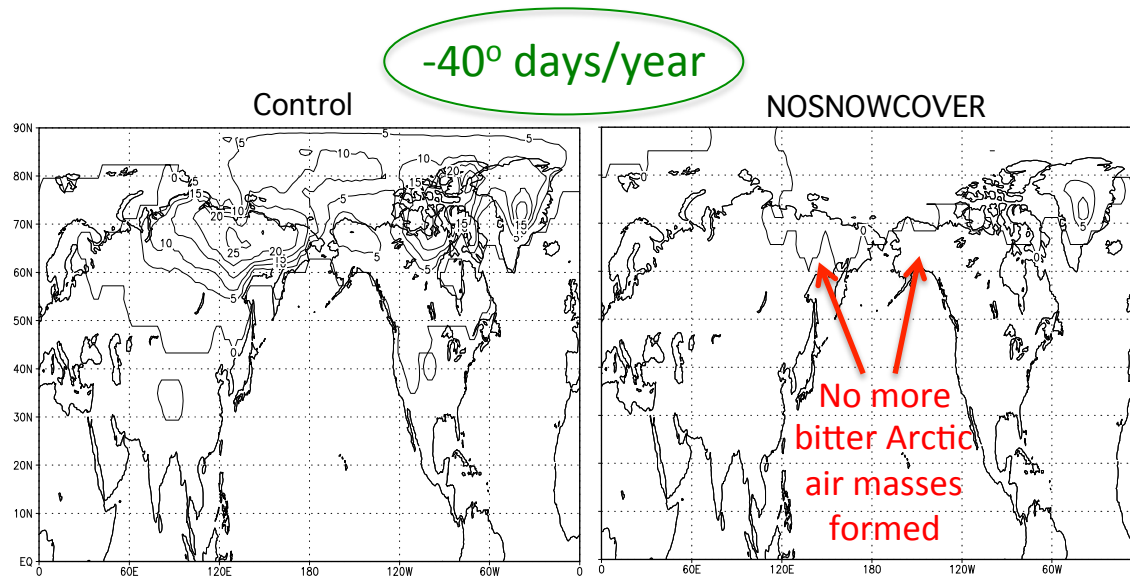
Weakened westerlies on
equatorward side of ridging

Large-Scale Meteorological Patterns (LSMPs) are clearly important for regulating extreme weather, but they aren't the only show in town.

We need to also recognize the strong thermodynamic role played by surface boundary conditions, such as snow cover and soil moisture. . . .

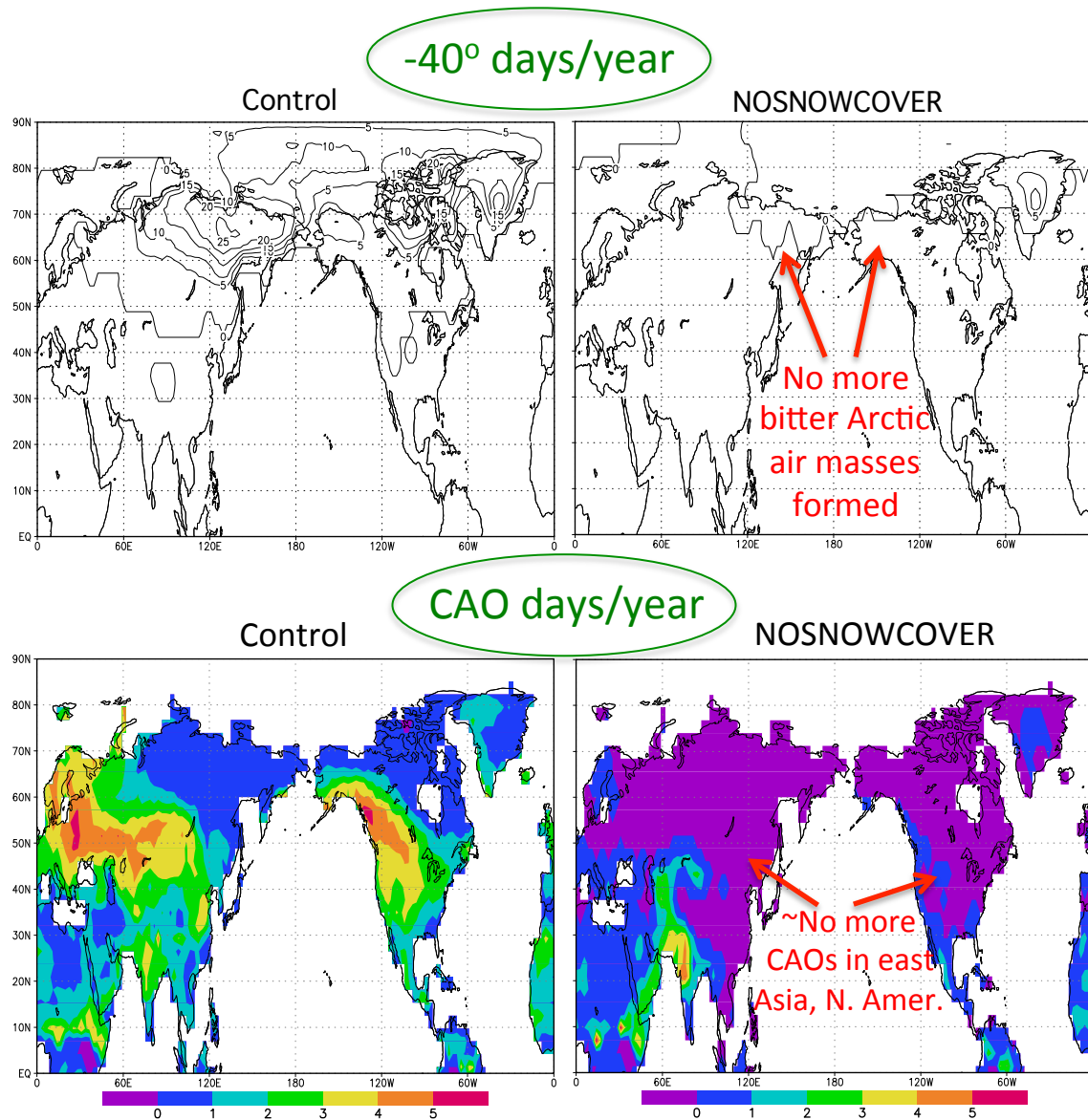
Role of Snow Cover in Extreme Cold-Air Outbreaks (CAOs)

CCSM3 Simulation with terrestrial snow cover suppressed



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CCSM3 Simulation with terrestrial snow cover suppressed



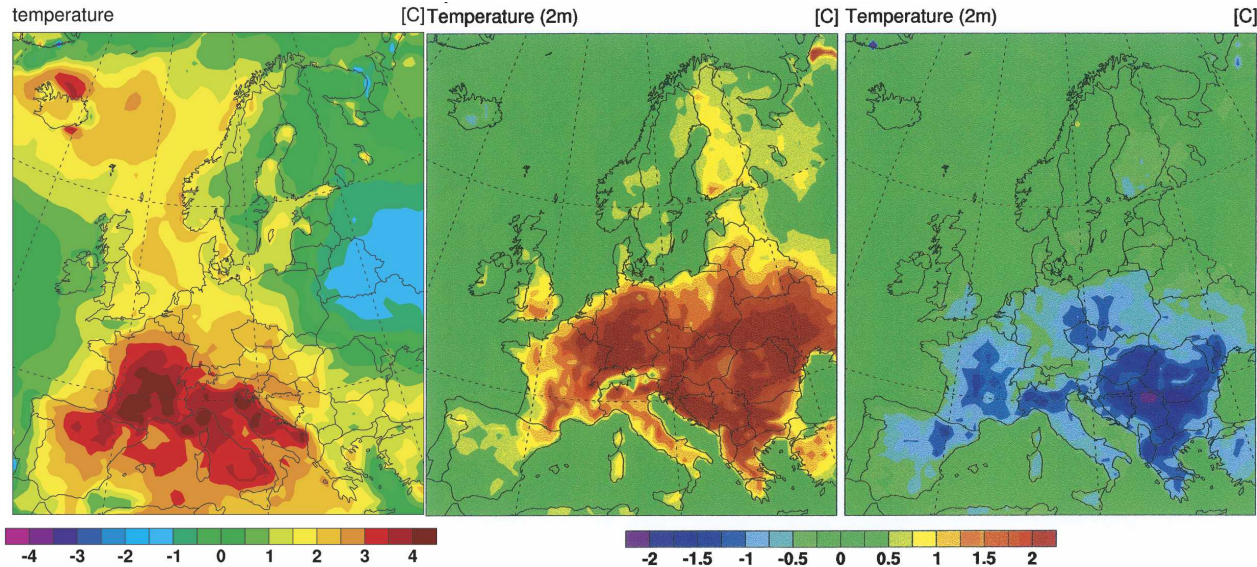
Role of Soil Moisture in Extreme Heat Wave: Europe 2003

Control Run

-25% Spring Soil Moisture

+25% Spring Soil Moisture

Surface
Temperature
Anomalies



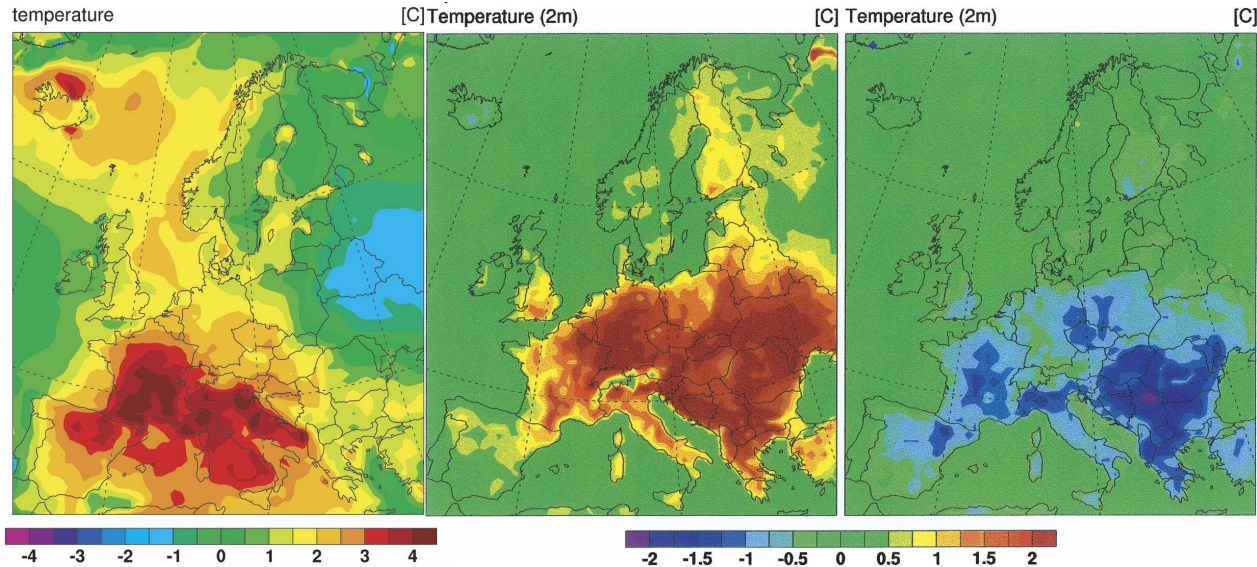
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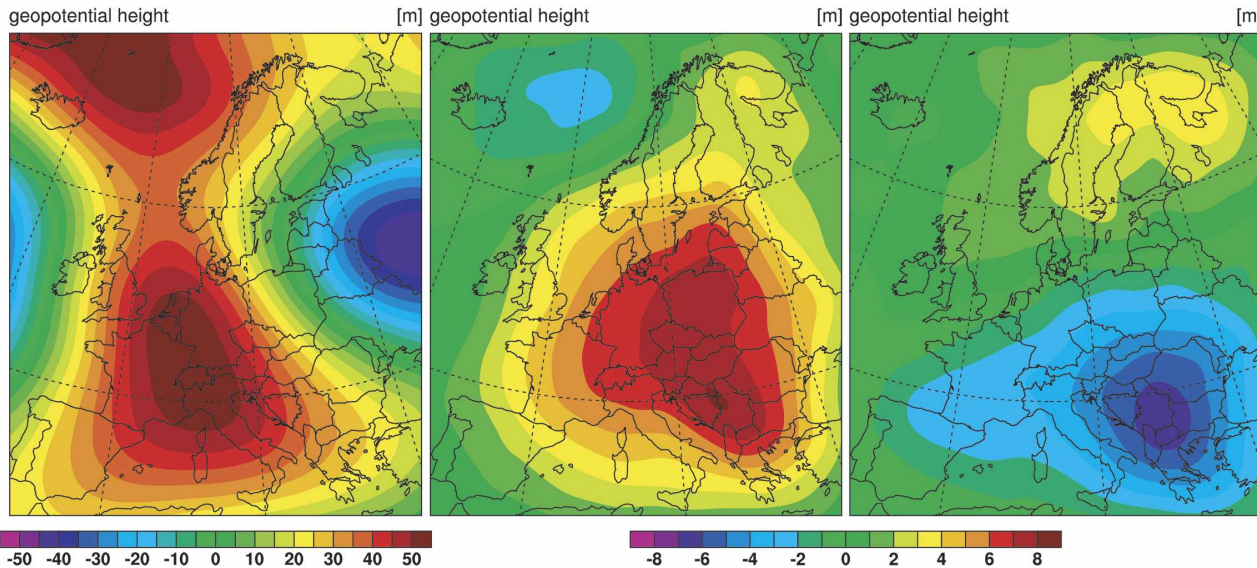


Control Run

-25% Spring Soil Moisture

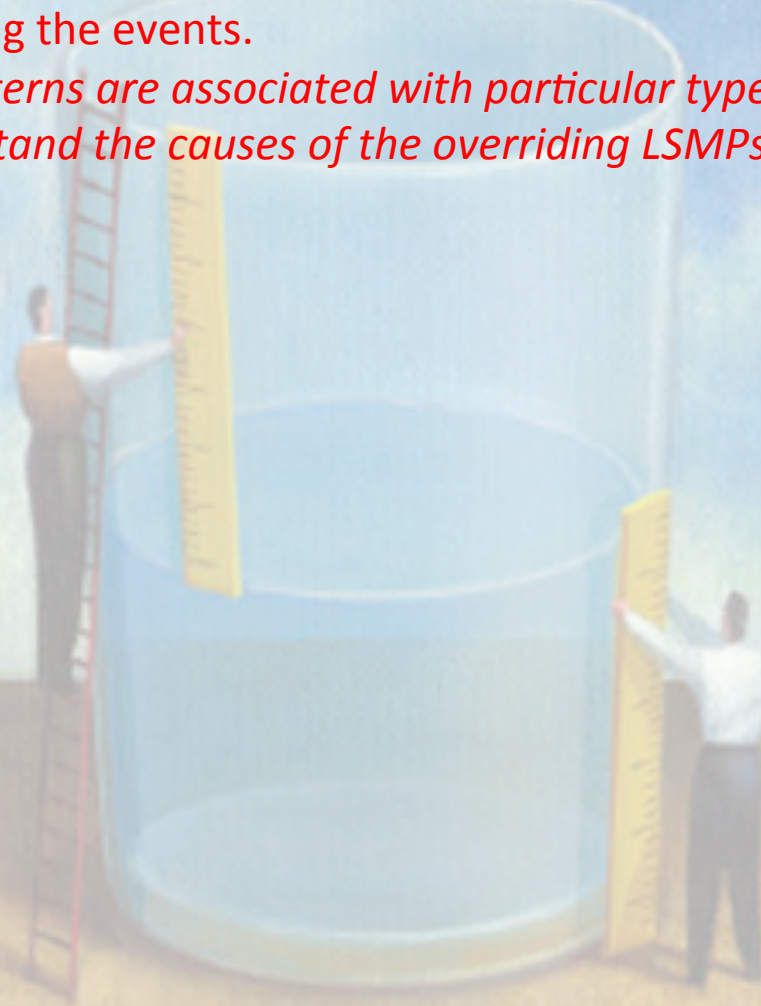
+25% Spring Soil Moisture

500 hPa
Height
Anomalies



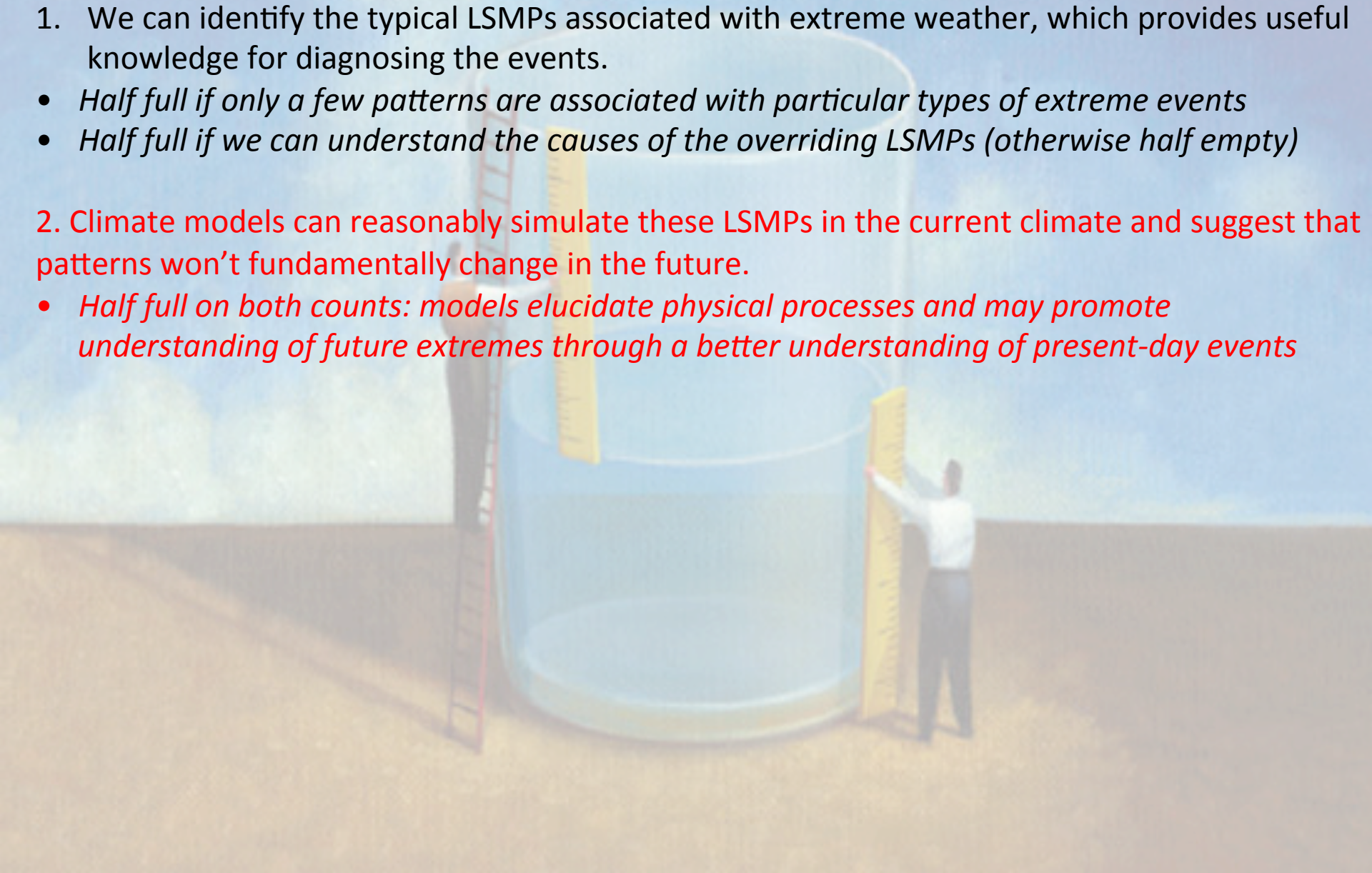
Relating Extreme Weather Events to LSMPs: Is the Glass half full or half empty?

1. We can identify the typical LSMPs associated with extreme weather, which provides useful knowledge for diagnosing the events.
 - *Half full if only a few patterns are associated with particular types of extreme events*
 - *Half full if we can understand the causes of the overriding LSMPs (otherwise half empty)*



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2. Climate models can reasonably simulate these LSMPs in the current climate and suggest that patterns won't fundamentally change in the future.
 - *Half full on both counts: models elucidate physical processes and may promote understanding of future extremes through a better understanding of present-day events*



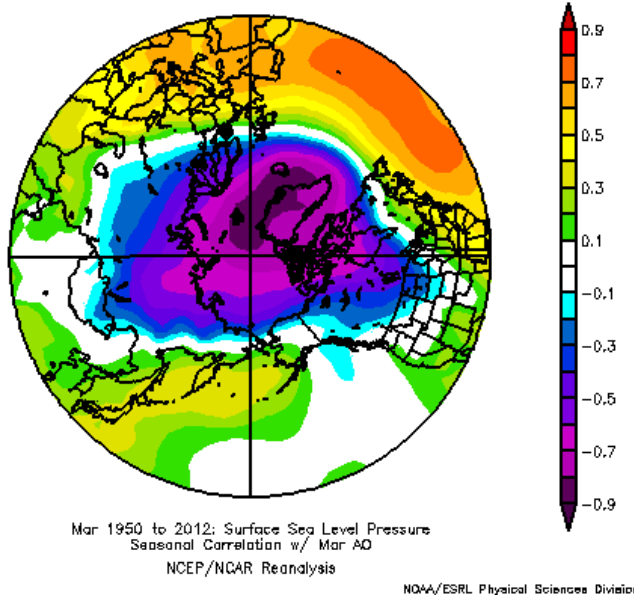
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 - *Half full on both counts: models elucidate physical processes and may promote understanding of future extremes through a better understanding of present-day events*
3. Mean states and extreme events are related, so explaining causes of extreme weather may be facilitated by better understanding of means and modes of variability.
 - *Half full to the extent that this is true (e.g., negative AO winters promoting colder mean states and CAOs)*
 - *Half empty, in that short-lived but high-impact extreme events and circulation anomalies are sometimes masked in monthly or seasonal averages*

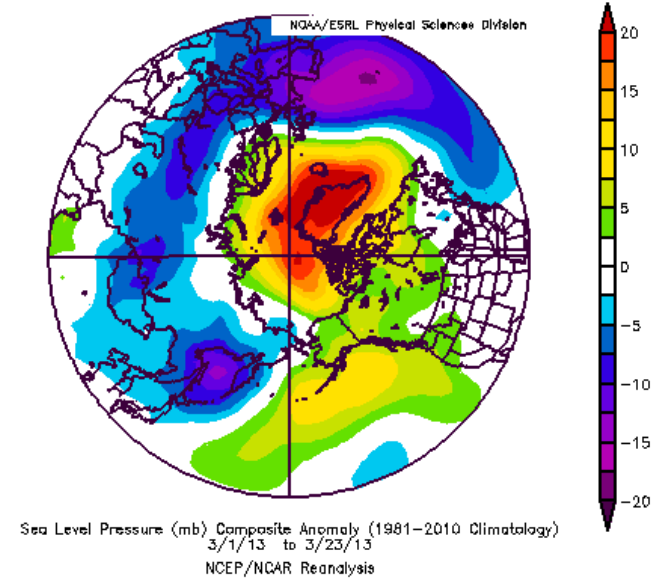
Relating Extreme Weather Events to LSMPs: Is the Glass half full or half empty?

1. We can identify the typical LSMPs associated with extreme weather, which provides useful knowledge for diagnosing the events.
 - *Half full if only a few patterns are associated with particular types of extreme events*
 - *Half full if we can understand the causes of the overriding LSMPs (otherwise half empty)*
2. Climate models can reasonably simulate these LSMPs in the current climate and suggest that patterns won't fundamentally change in the future.
 - *Half full on both counts: models elucidate physical processes and may promote understanding of future extremes through a better understanding of present-day events*
3. Mean states and extreme events are related, so explaining causes of extreme weather may be facilitated by better understanding of means and modes of variability.
 - *Half full to the extent that this is true (e.g., negative AO winters promoting colder mean states and CAOs)*
 - *Half empty, in that short-lived but high-impact extreme events and circulation anomalies are sometimes masked in monthly or seasonal averages*
4. Other (non-dynamical) factors, such as local snow cover and soil moisture anomalies can also shape weather extremes.
 - *Half empty, except that LSMPs influence the formation and persistence of these surface-based anomalies and can in turn be significantly affected by them*

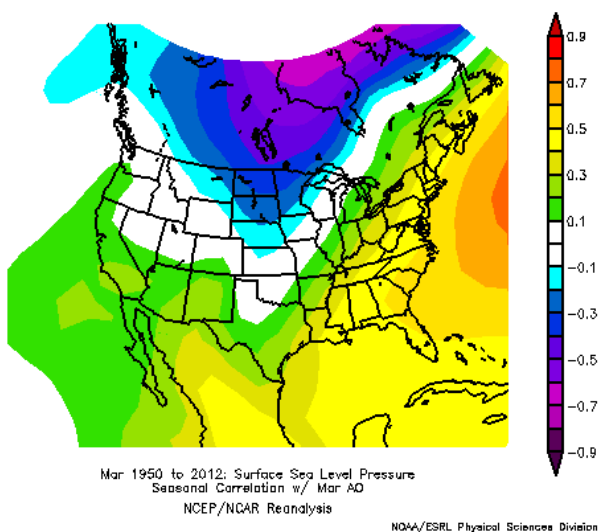
March AO Correlations with SLP



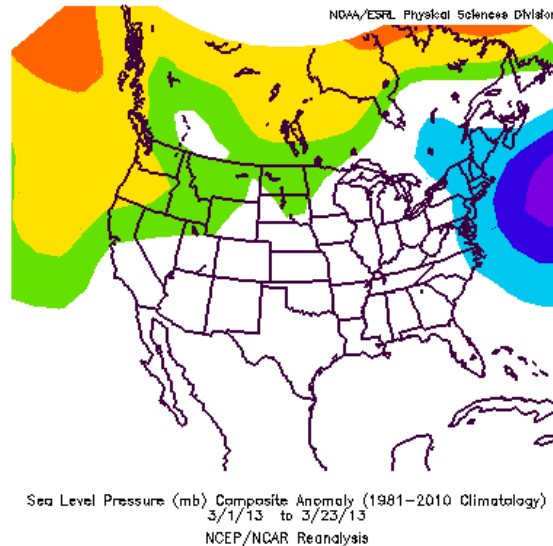
March 2013 SLP Anomaly



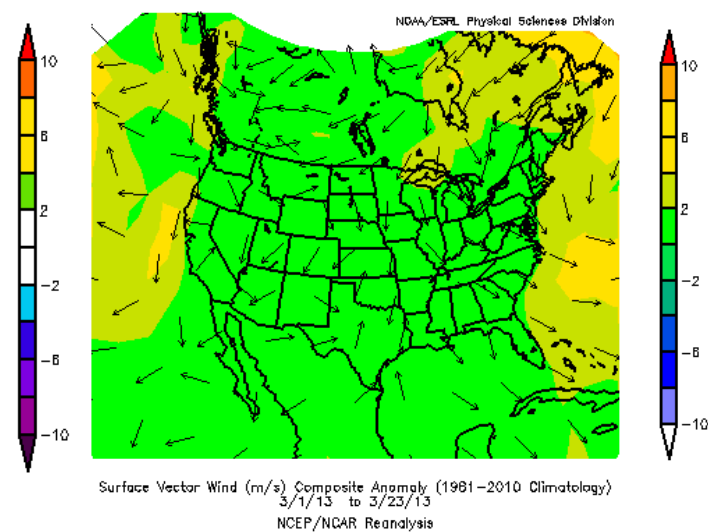
March AO Correlations with SLP



March 2013 SLP Anomaly

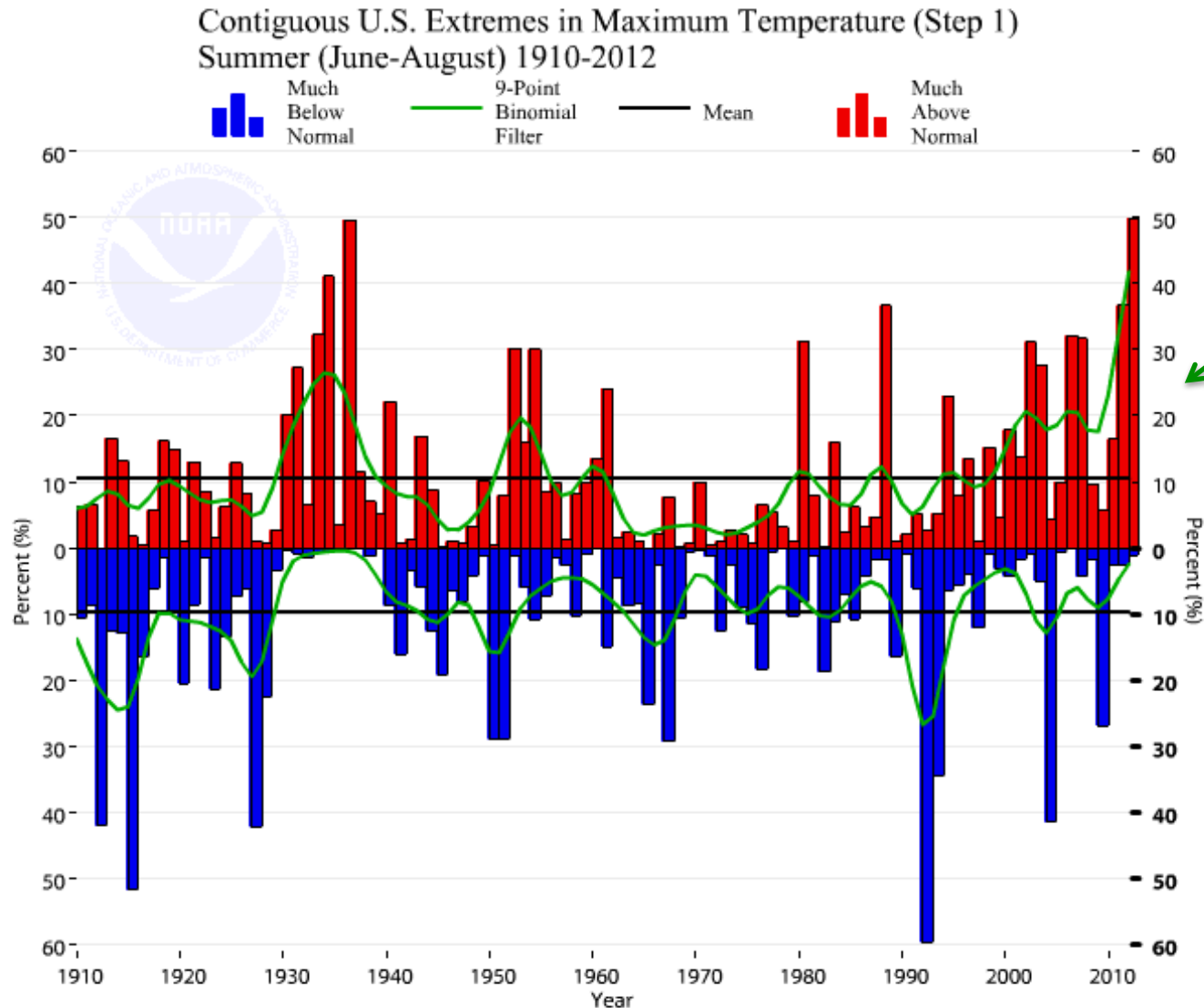


March 2013 Sfc Wind Anomaly



U. S. Climate Extremes Index

Summer Daily Maximum Temperature (Continental U. S.)

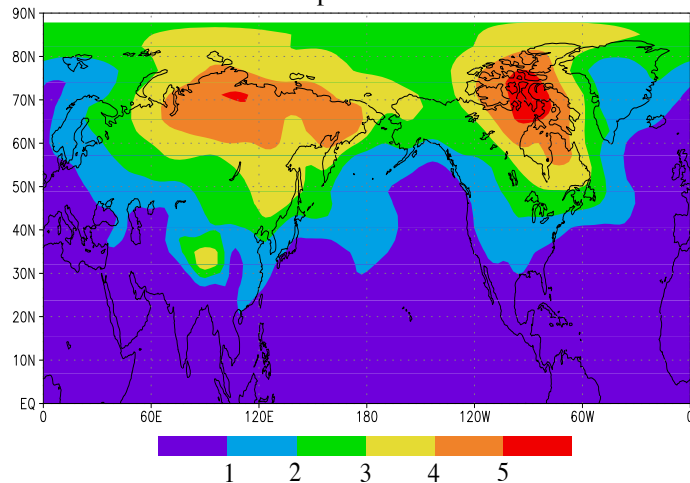


Role of Snow Cover in Extreme Cold-Air Outbreaks (CAOs)

CCSM3 Simulation with terrestrial snow cover suppressed

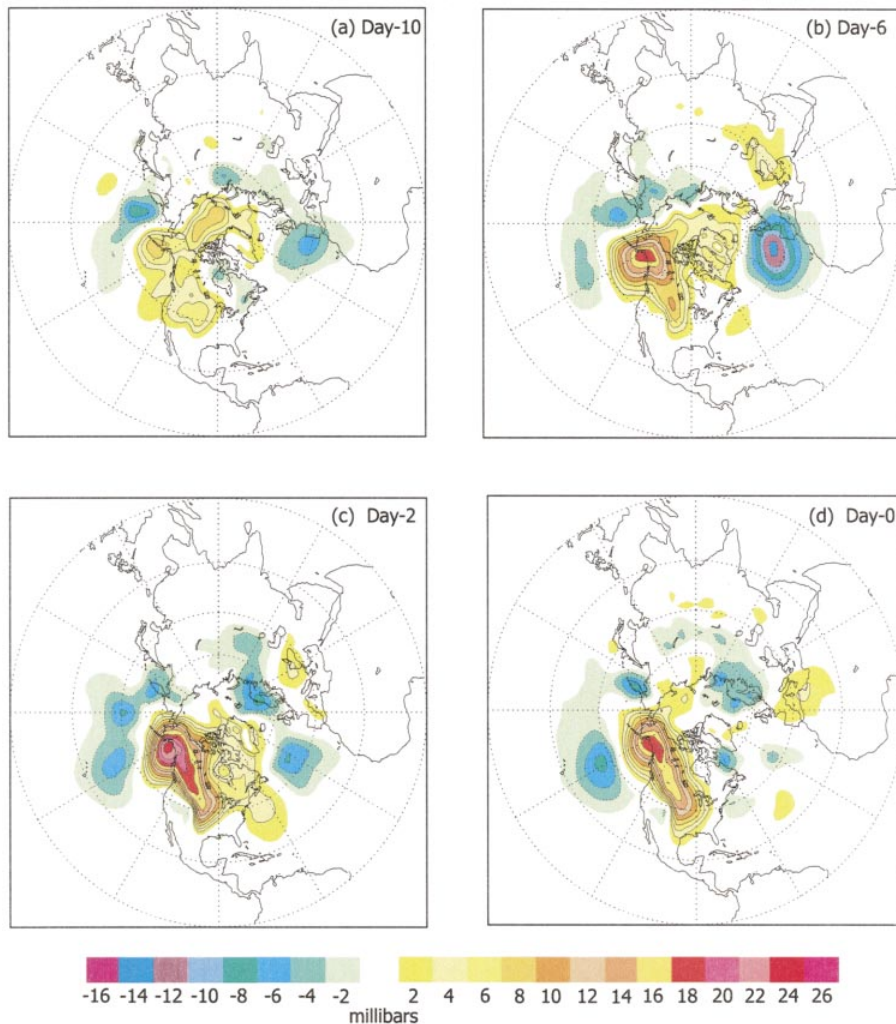
Annual Average

Increase in Air Temperature without Snow Cover



Synoptic Evolution of Extreme Cold-Air Outbreaks (CAOs)

Midwestern CAOs



East Coast CAOs

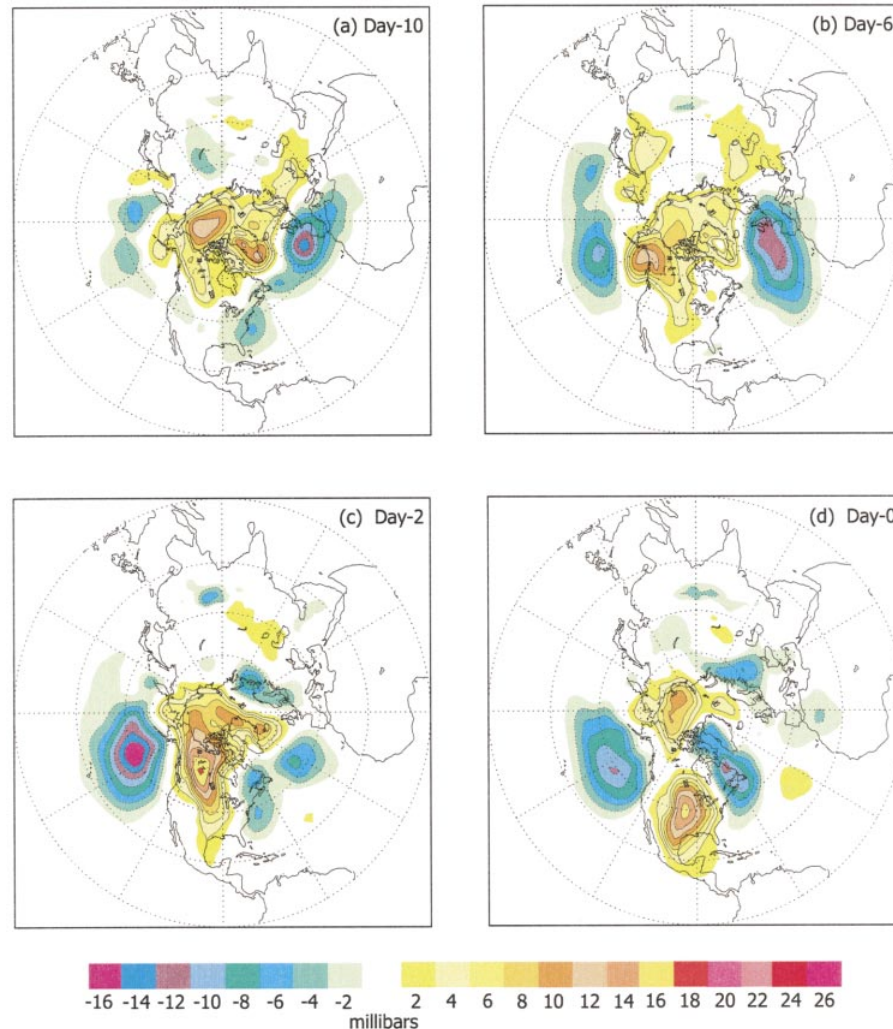


FIG. 7. SLP anomaly composites for (a) 10 days, (b) 6 days, (c) 2 days, and (d) 0 days prior to the dates of the 10 cold events in the Midwest (MW) region. Numbers below color bar denote the lower limit of each 2-mb range of magnitudes. Unshaded areas have magnitudes smaller than 2 mb.