Modelling the dynamics of tidewater outlet glaciers: approaches, issues and perspectives

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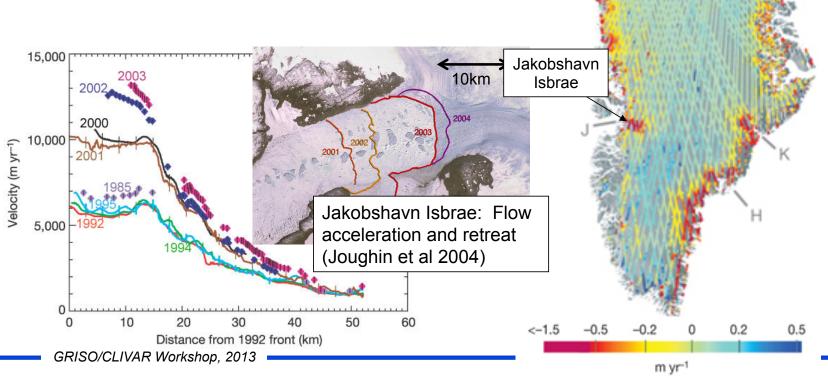
Thanks to Faezeh Nick,...

- Introduction

Tidewater outlet glaciers

 Retreat, acceleration and thinning

Not reproduced/predicted by ice sheet models (IPCC 2007)



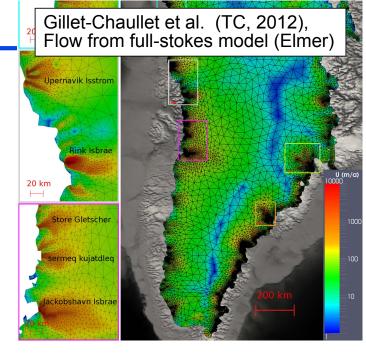
The modelling issue

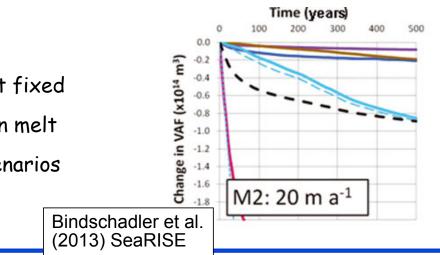
Ice sheet models do well for:

- Ice flow (SIA, higher order,...) ✓
- Surface evolution \checkmark
- Numerical methods (1D-3D, finite Differences/Elements,...) ✓
- Grounding line motion (\checkmark)

But poor representation of processes at outlet-ocean boundary

- Spatial resolution: few km wide
- Calving/terminus dynamics: front fixed
- Forcing crude: e.g. indirect ocean melt
- Prescribed 'what if...' forcing scenarios (no retreat feedbacks)





— Models of tidewater outlet glacier dynamics

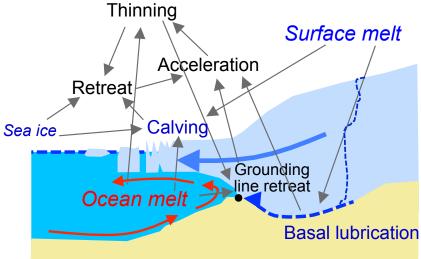
Ice sheet models IPCC/Sea-level

- Predictions (sea-level, IPCC)
- Ice sheet scale

Minimal fully dynamic outlet glacier model (1D, SSA,...)

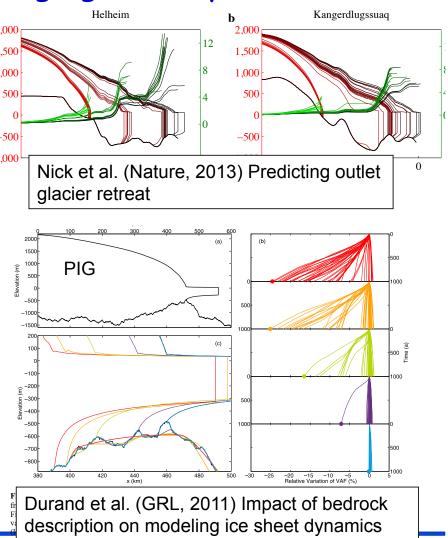
- Moving calving front
- Couple forcing to dynamic response
- Explore/illustrate dynamic feedbacks modelling issues





Sensitivity to fjord/trough geometry

- Retreat rate strongly depends^{2,00}_{1,50} on fjord/trough geometry $\frac{1}{5}$
 - Water depth/Trough width 🛓
 - Non-linear with depth (h^{~4}, Schoof 2007)
 - P_{eff}-sliding (Pfeffer, 2007)
- Threshold behaviour
- Variations/details of bed important (uncertainty)
- Requires high spatial model resolution (adaptive)

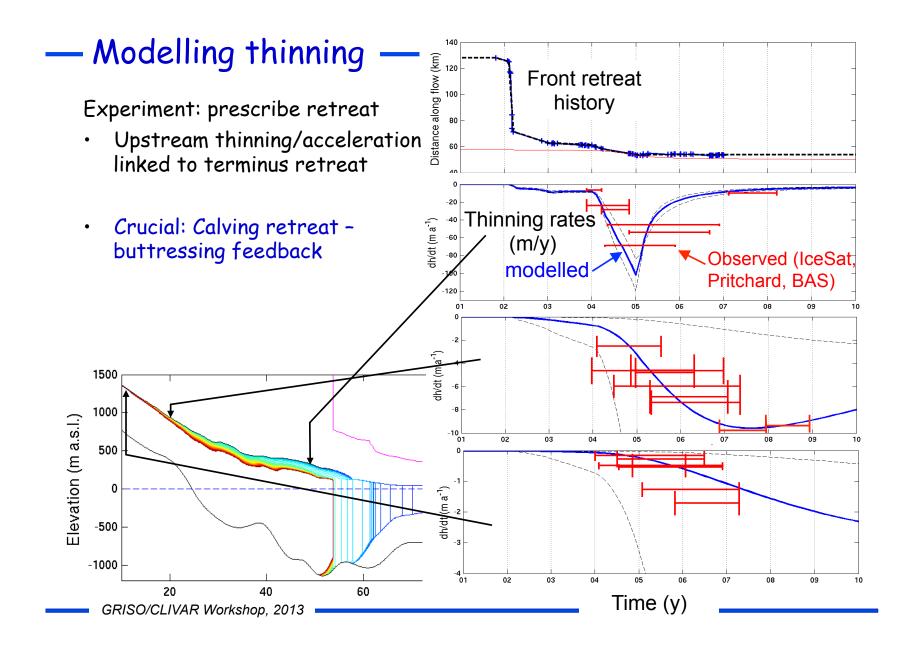


- Terminus/upstream dynamics

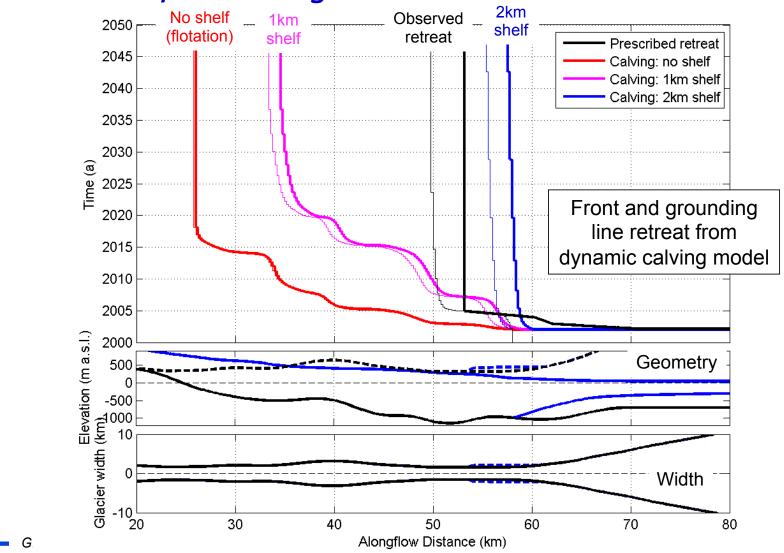
Experiment: collapse of Larsen B ice shelf 2002 Effect on tributary glacier Crane **Crane Glacier** ٠ Larsen B -2003 5km 2005-2012 MSC A Crane Glacier 70km 31. Jan. 2002 2004 2002 Flow speed of Crane Glacier (c)2000 2005 From Rignot et al. 2004/06 2003 Terminus retreat of 1500 2002 $V(\mathrm{m\,yr^{-1}})$ Crane glacier 1000 1996 500 2000 -15 -10-5 0 5

- Terminus/upstream dynamics

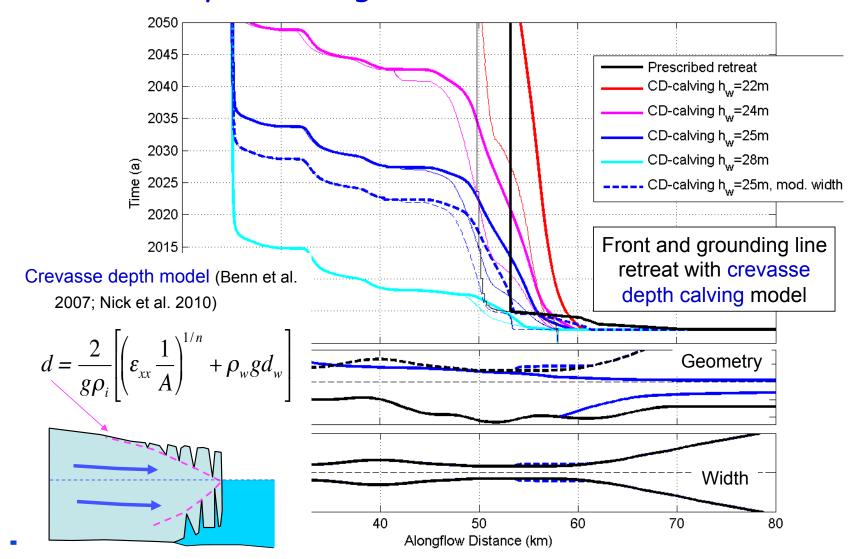
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- Sensitivity to calving model I



— Sensitivity to calving model II



Response to external forcing: ocean melt

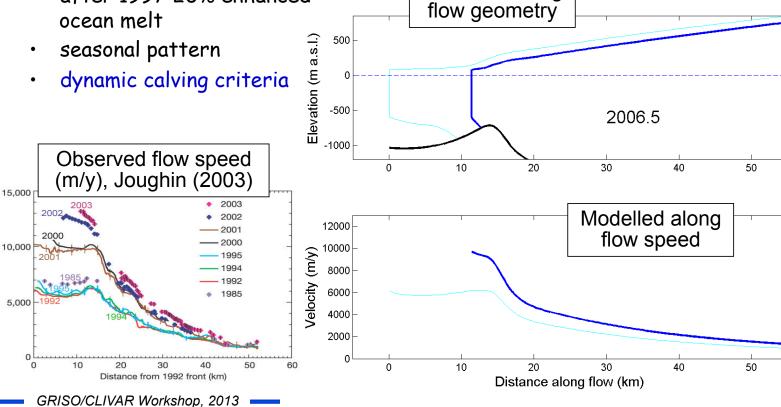
Jakobshavn Isbrae

Ocean melt: m/d (Motyka 2010)

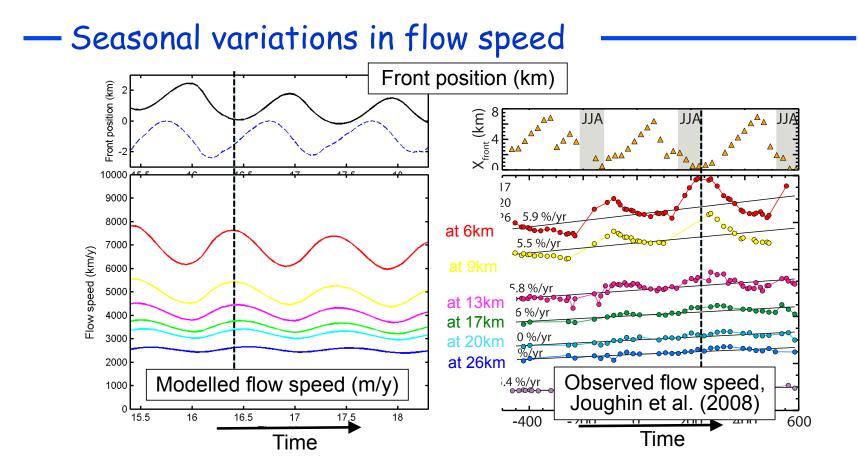
Experiment:

Velocity (m yr⁻¹)

- after 1997 20% enhanced ocean melt
- ٠



Modelled along

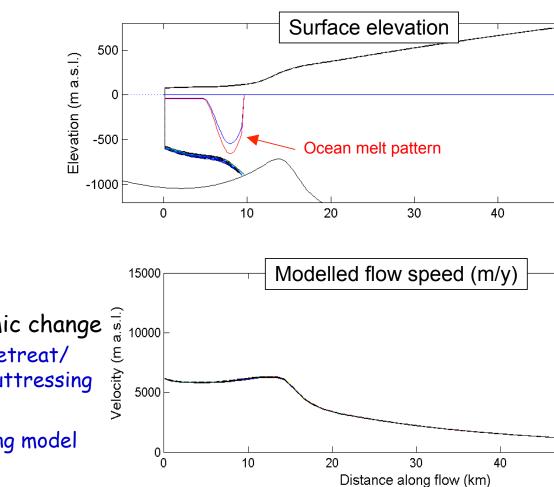


- Reproduce upstream variations in speed
- Changes from terminus



Experiment:

- 20% enhanced melt
- Fixed calving front



Trigger:

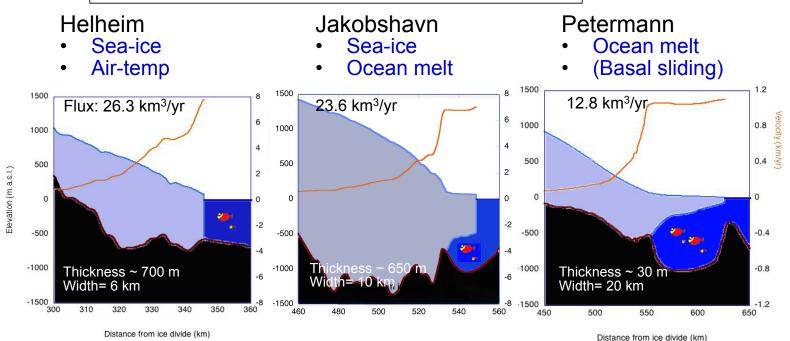
• Ocean melt

Mechanism for dynamic change

- Feedback between retreat/ calving and loss of buttressing
- Dependence on calving model

— Sensitivity to forcing type

Modelling outlet glacier dynamics of 4 major outlets (Nick et al. Nature, 2013)



Variable sensitivity to forcing

- depending on glaciers/terminus type/climate
- can change during retreat

Ocean melt forcing and calving

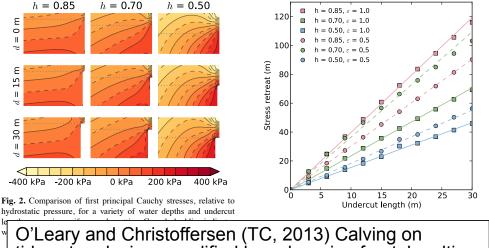
Grounded terminus (instead of floating)?

- Oversteepen/weaken calving front
 - \rightarrow enhance calving

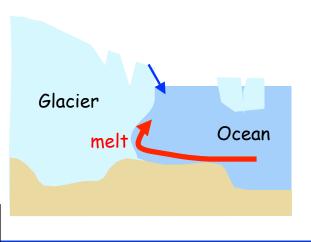
Numerical modelling issues

- Difficult to implement at vertical ice front
- Coupling of ocean melt to ice models

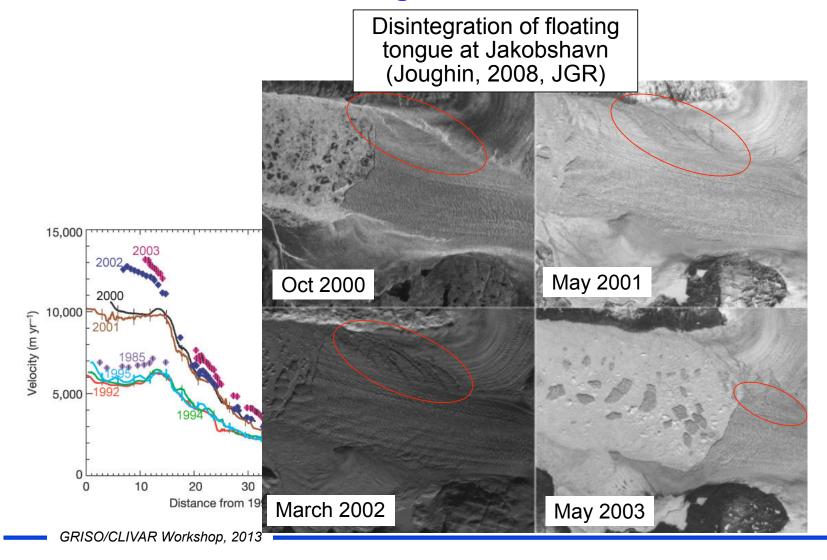


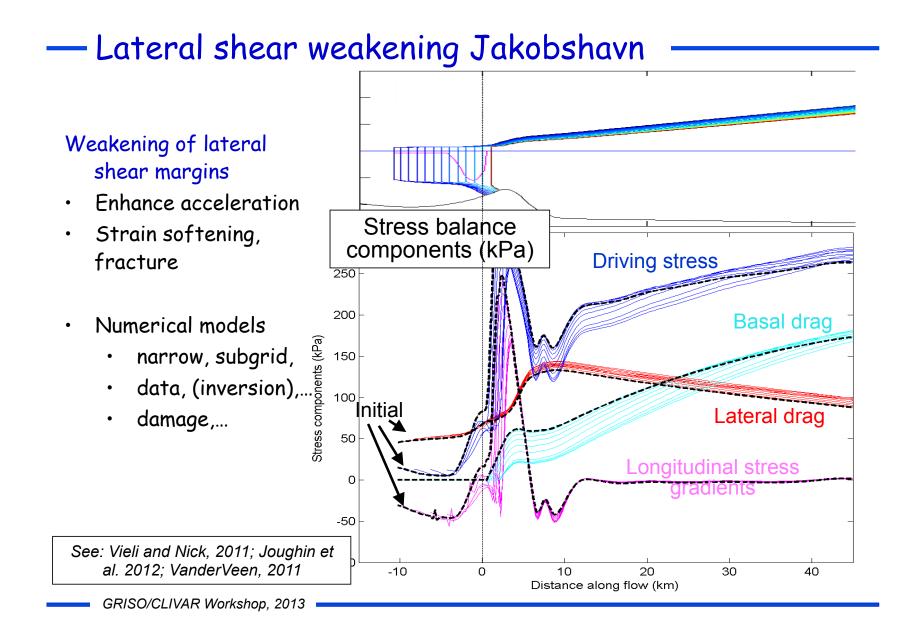


O'Leary and Christoffersen (TC, 2013) Calving on tidewater glaciers amplified by submarine frontal melting

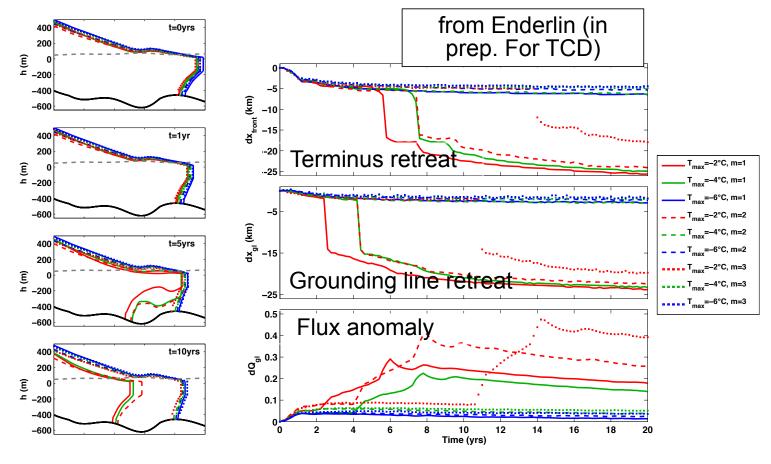


— Lateral shear weakening: Jakobshavn Isbrae





Sensitivity to ice rheology/sliding exponent



• Sensitivty/uncertainty from internal parameters

— Implications from modelling

- Terminus dynamics controls upstream dynamics (mass loss)
 - Upstream propagation: ok
 - Calving model crucial robust/validation
- Forcing
 - Calving linked to forcing: ocean melt, air temp., sea-ice... (How?)
 - Not just ONE forcing important (terminus type/climate)
- High sensitive to fjord/trough geometry (bed/width)
 - Threshold/non-linear behaviour (predictions?)
 - Need accurate topography
- Lateral shear softening, rheology?
- Only 'simple' fully dynamic outlet glacier model (SSA, flowline,...)

- How to improve models

Technical/numerical development

- Grid resolution, 1D to 3D
- Moving boundaries (calving front: resolution, adaptive grids,..)
- Flow approximation: higher-order, full stokes

Process representation

- Calving models (3D, damage, statistical,...)
- Coupling of calving to forcing (ocean melt models...)
- Initial conditions (transient?)

Validation/calibration: Benchmark data sets

- Geometry data (bed,...)
- Response and forcing data (e.g. ocean, ...)
- Range of time scales: in particular long-term
- model-model intercomparison

— Final thoughts

Rethink expectation on models

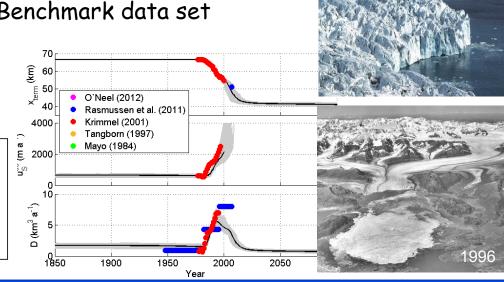
- Not predict exact timing of single retreat-event
- 'Average retreat trend' (ensemble of glaciers/responses, decades/ century)

Not just large Greenland ice-sheet outlets

- Smaller tidewater glaciers (Alaska, Svalbard,...)
- Extensive dataset Benchmark data set
- Easier accessible

Apply models!!!

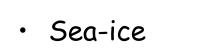
Colgan et al. (TC, 2012) Monte Carlo ice flow modeling projects a new stable configuration for Columbia Glacier, Alaska, c. 2020

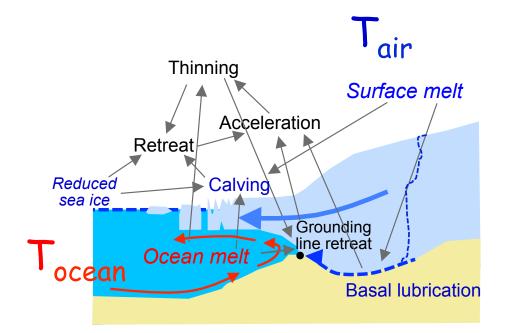






- Surface melt
- Ocean melt





• Dynamic response coupled to forcing

- Summary: terminus/upstream dynamics

- Terminus dynamics determines upstream dynamics
 - thinning, acceleration, mass loss
- Inland propagation \checkmark
- Strong dependence on calving model/parametrization
- Influence of channel topography

Ocean melt forcing and calving

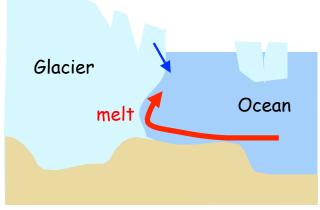
Grounded terminus (instead of floating)

- Oversteepen/weaken calving front
 - \rightarrow enhance calving

Numerical modelling issues

- Ocean forcing records: temperature/ salinity
- Coupling of ocean melt to ice models
- Difficult to implement at vertical ice front





Hansbreen Svalbard (Vieli et al 2002)

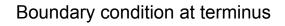
— A minimal fully dynamic flowline model (1D)

• Dynamic treatment of calving: front position criteria -

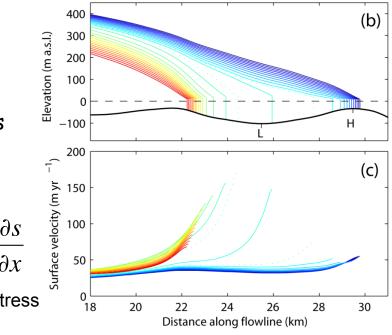
Explore/illustrate

Mostly confined in trough/fjord: 1Dflowline/band

- Couple forcing to dynamic response
- Role of terminus dynamics/calving on upstream dynamics



$$\frac{\overline{\partial u}}{\partial x}\Big|_{front} = A \cdot \left[\frac{1}{4}\rho_i g\left(1 - \frac{\rho_i}{\rho_w}\right)\right]^n \cdot h_f^n$$



- Basal, lateral and longitudinal stresses
- Moving grounding line/front stretching grid Stress balance

$$2\frac{\partial}{\partial x}h\nu\frac{\partial u}{\partial x} - \beta u^{1/m} - \frac{h}{W}\left(\frac{5u}{2AW}\right)^{\frac{1}{m}} = \rho_i gh\frac{\partial}{\partial x}$$

Longitud. Basal Lateral Driving stress

External forcing: ocean melt Experiment: Modelled surface prescribed ocean melt pattern • Elevation (m a.s.l.) 20% enhanced melt (+1°C) Dynamic calving criteria -500 Ocean melt pattern (crevasse depth model) -1000 Observed flow speed (m/y) 15,000 Modelled flow speed (m/y) 10,000 Velocity (m yr⁻¹) Velocity (m a.s.l.) 5,000 ŏ Distance along flow (km) GRISO/CLIVAR Workshop, 2013