

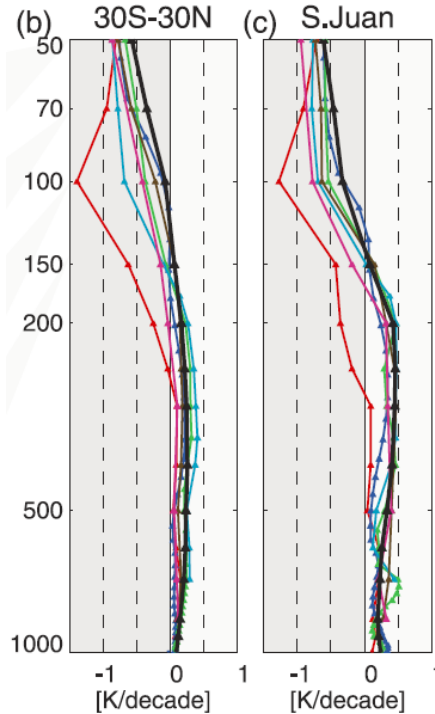
Impact of stratospheric temperature on hurricane intensity: An idealized modeling study

Shuguang Wang, Suzana Camargo,
Adam Sobel, Lorenzo Polvani

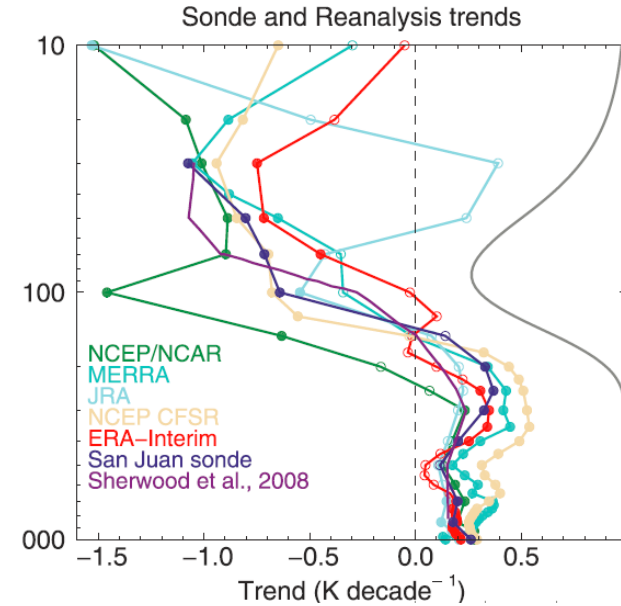
Columbia University

Trend in lower stratosphere and upper troposphere temperature

Randel et al 2009,
Emanuel 2011, Emanuel
et al 2013, Vecchi et al
2013



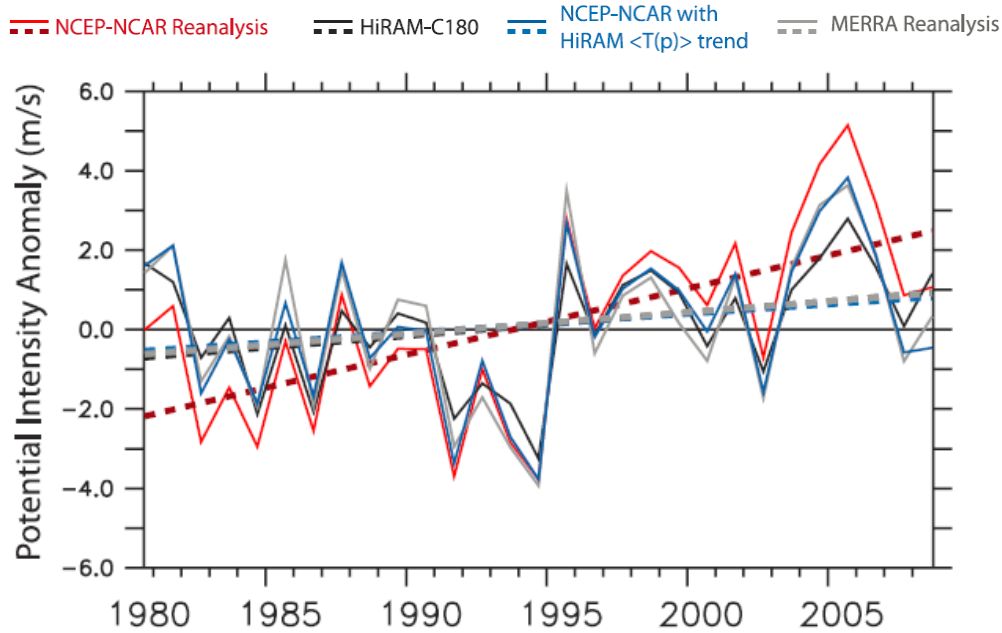
Vecchi et al 2013



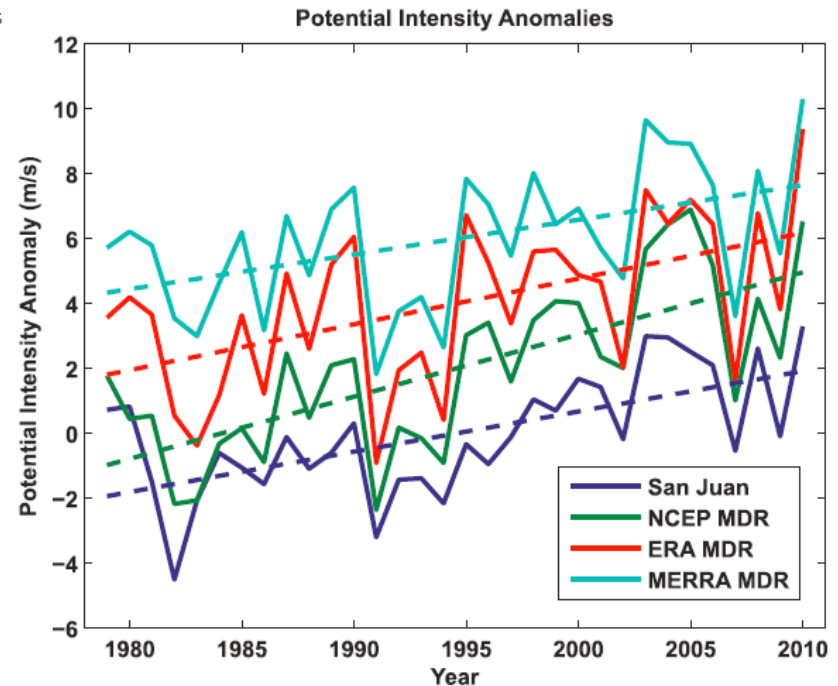
Emanuel et al 2013

Cooling trend in stratosphere temperature

Potential intensity at Atlantic MDR



Vecchi et al 2013



Emanuel et al 2013

~ 1 m/s per degree cooling (except NCEP-NCAR reanalysis)

=> Using WRF to better understand this issue

WRF model V3.0

- $f=5 \times 10^{-5} \text{ s}^{-1}$, two nested domain: DX=12 km, 4 km, 50 levels, no convective parameterization – Nonhydrostatic modeling

-Simple radiation: -1.25 K/day in troposphere (Pauluis and Garner 2006)

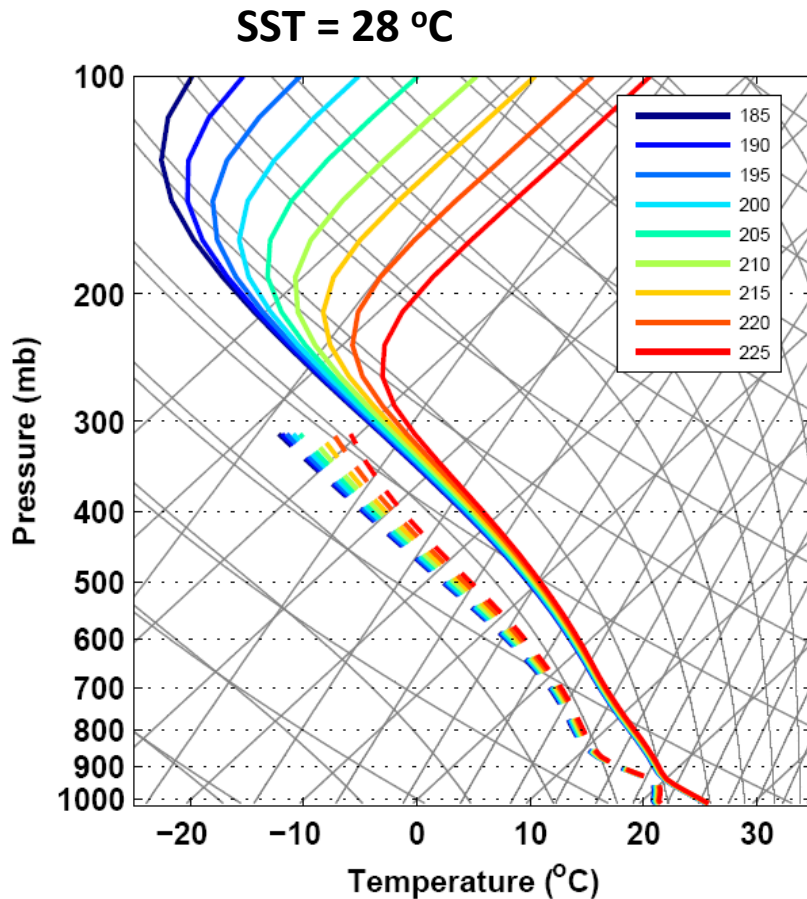
$$Q|_{rad. \text{ cooling}} = \begin{cases} -1.2 \text{ K} \cdot \text{day}^{-1} & \text{for } T > 6 + T_s \\ \frac{T_s - T}{5 \text{ days}} & \text{for } T \leq 6 + T_s \end{cases}$$

-Isothermal stratosphere, $T_s = 190 - 215 \text{ K}$

-Homogeneous sea surface temperature: 26, 28, or 30 degree

-Physics schemes: PLin et al. microphysics, YSU PBL, surface fluxes, Smagorinsky sub-grid model - mixing length proportional to grid spacing

The RCE soundings (~ 400 km domain)

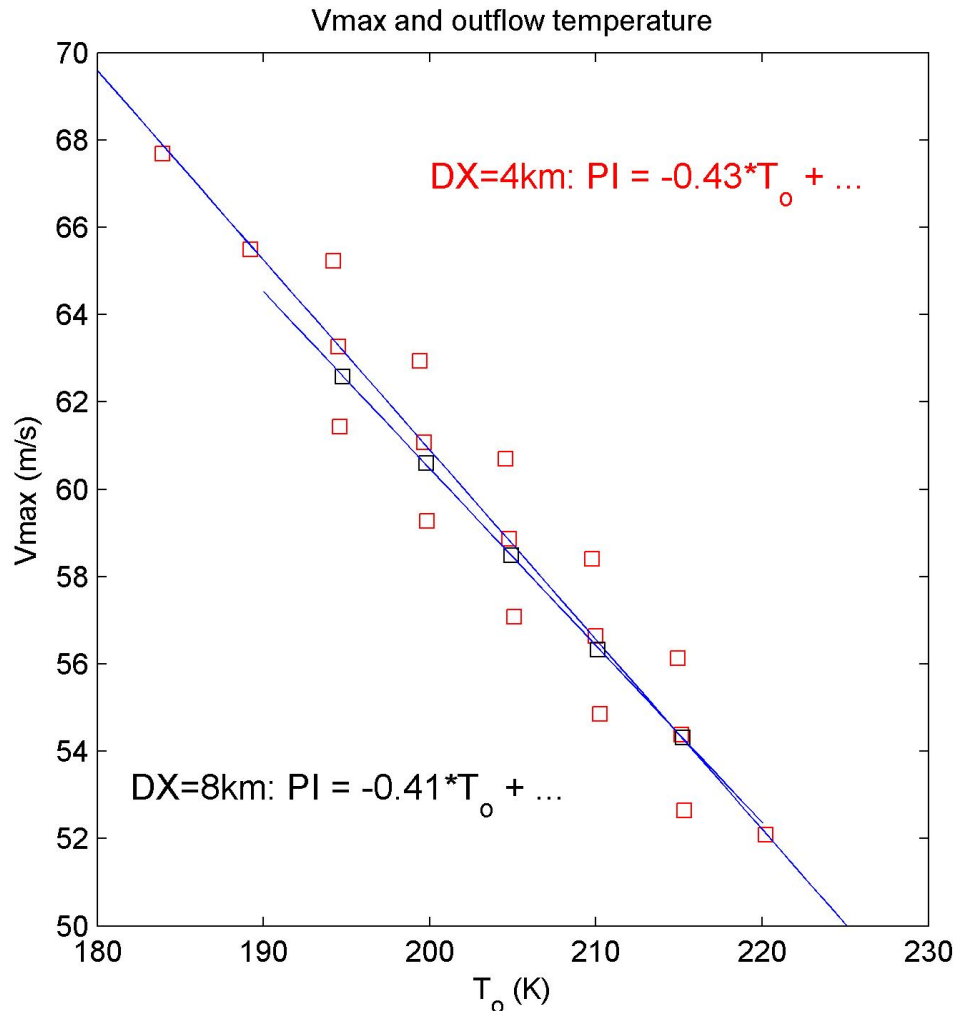


A set of initial sounding
with different
stratospheric temperature
In RCE (at $Dx=12$ km)
At different SST:
26, 28, and 30

Radiative cooling and
surface fluxes are nearly
fixed ~ 120 W/m²

Potential Intensity from the RCE soundings

Using assumptions in PI: $Ck/Cd=0.5$, Pseudo-adiabat, no dissipative heating



Less than 0.5 K
decrease in T_s

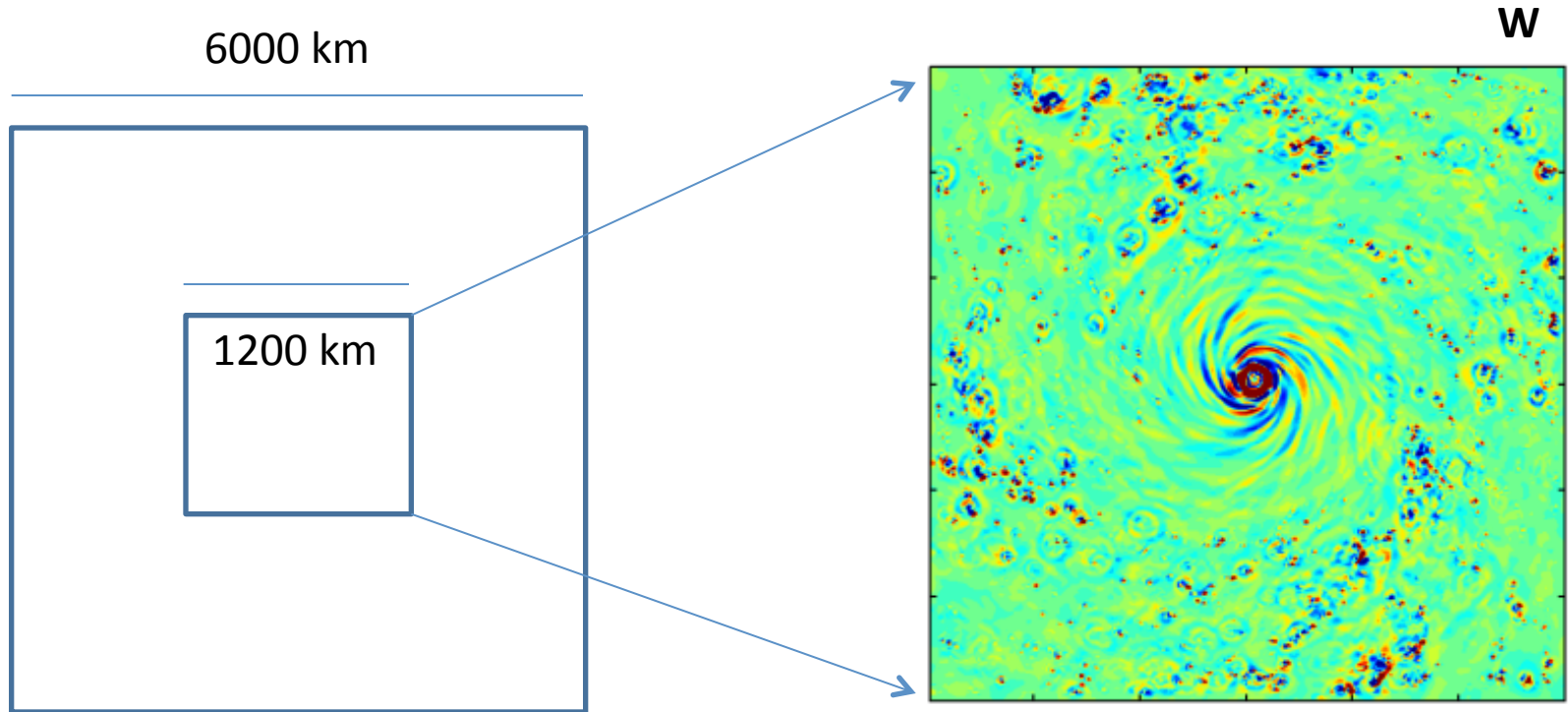


1 m/s in PI

The slope varies with
different Ck/Cd . E.g.,
 $Ck/Cd=1$, reversible-adiabat,
the slope is close to 1

WRF hurricane simulations

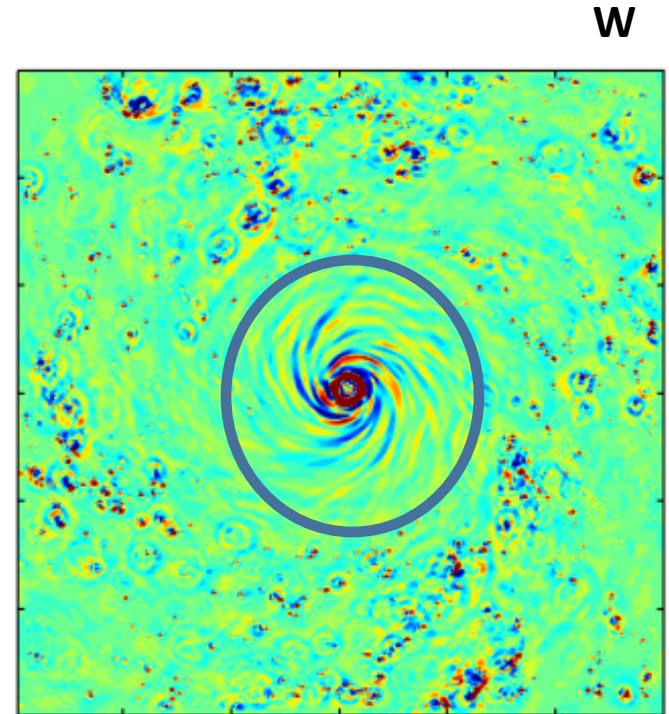
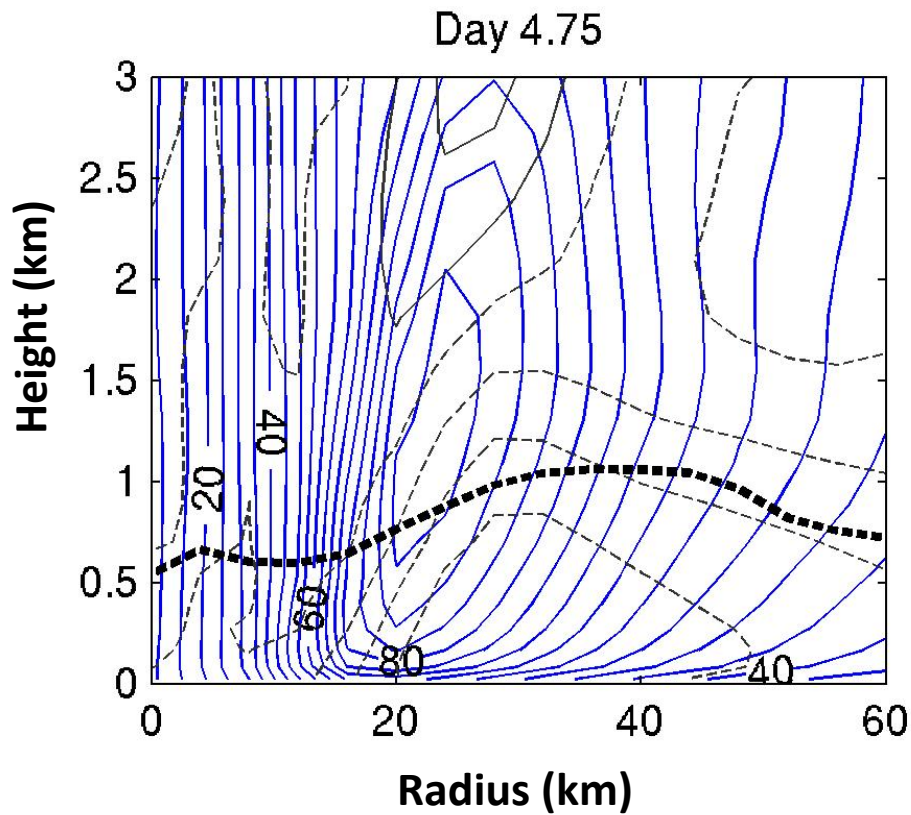
- Same physics as RCE, but in a larger, nested domain



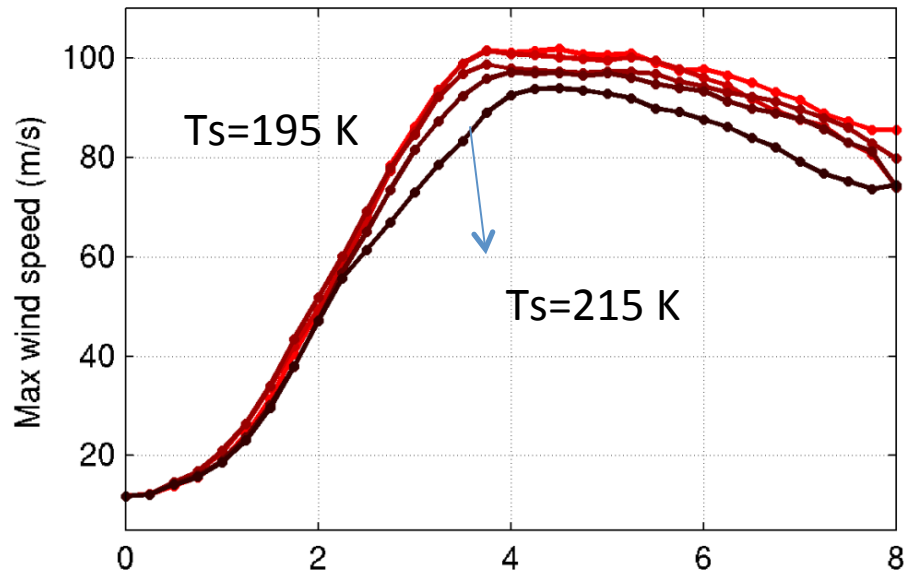
- Idealized hurricane simulations with an initial vortex (Rotunno and Emanuel 1987), but no wind shear, no interactive radiation, no ocean mixed layer etc.

WRF hurricane simulations

- Axis symmetric view of tangential wind, radial wind, and boundary layer height

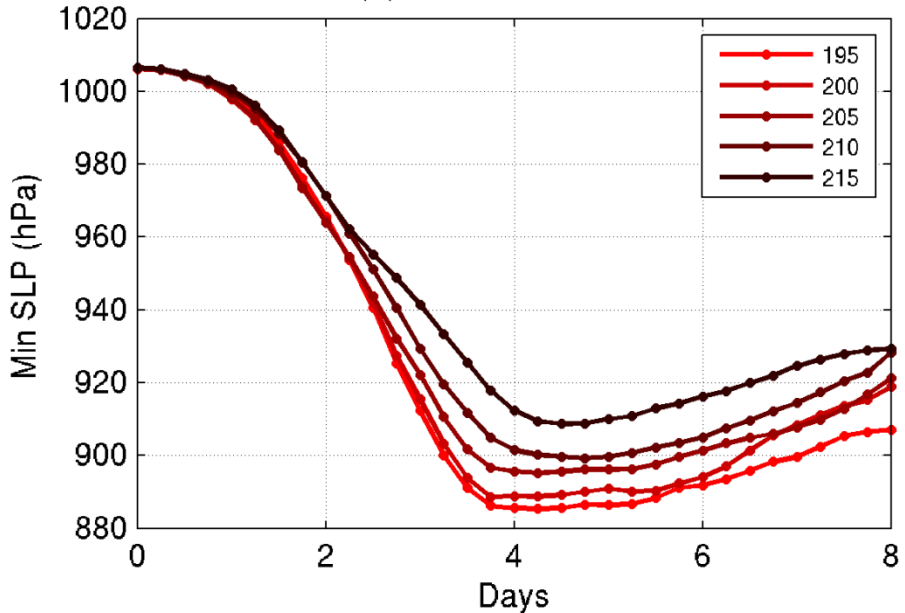


(b) SST=28, Max. Tangential Wind



Model simulated maximum azimuthal wind speed reaches ~ 100 m/s

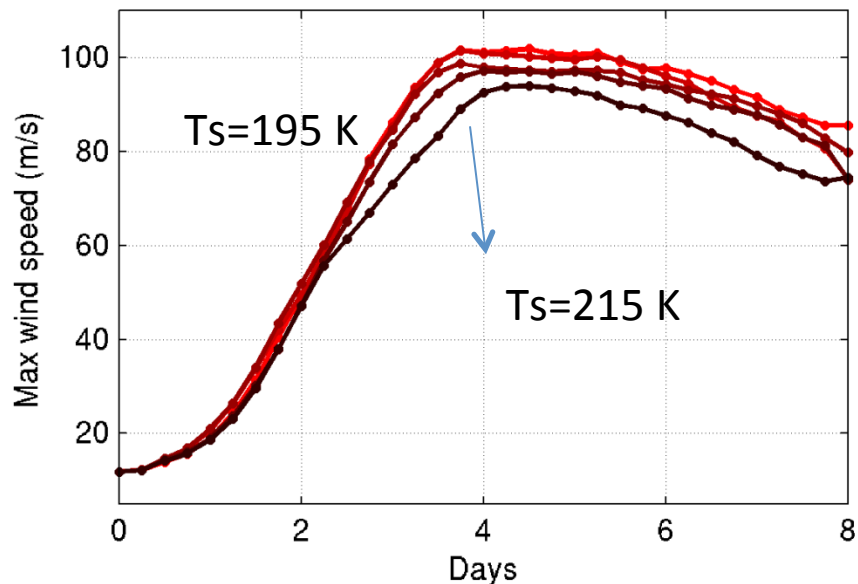
(a) SST=28, Min. SLP



Minimum sea level pressure

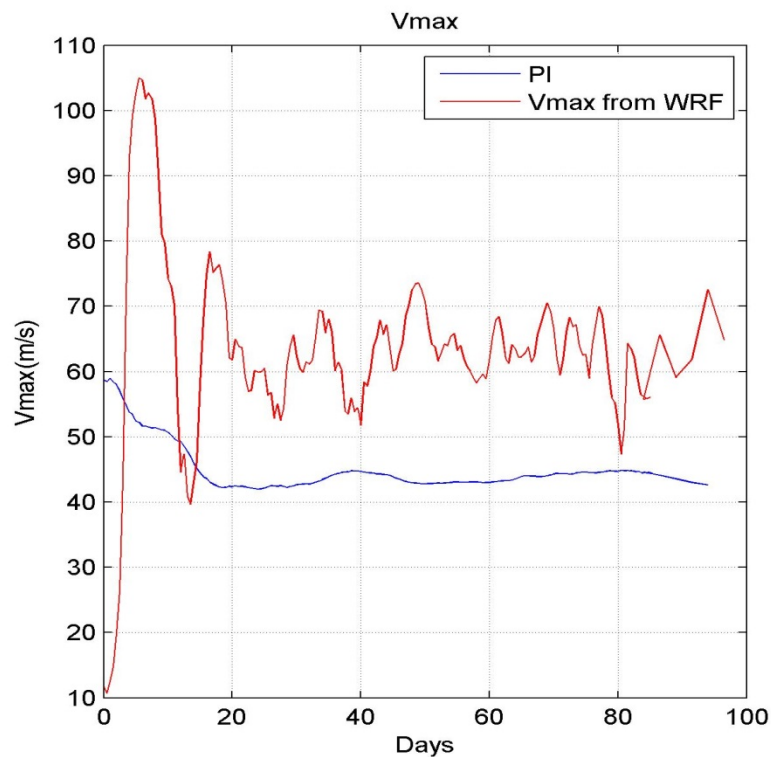
(c) SST=30, Min. SLP

(b) SST=28, Max. Tangential Wind



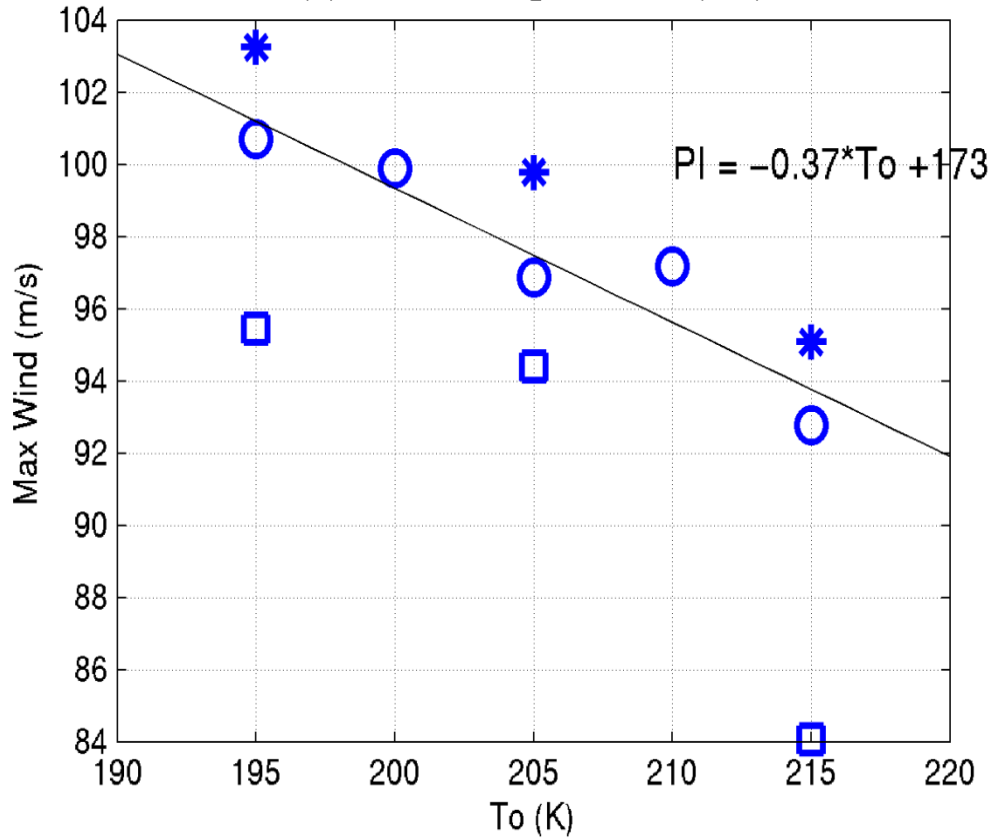
Model simulated maximum azimuthal wind speed reaches $\sim 100\text{ m/s}$

Model runs for longer time (lower resolution, smaller domain, no nesting)

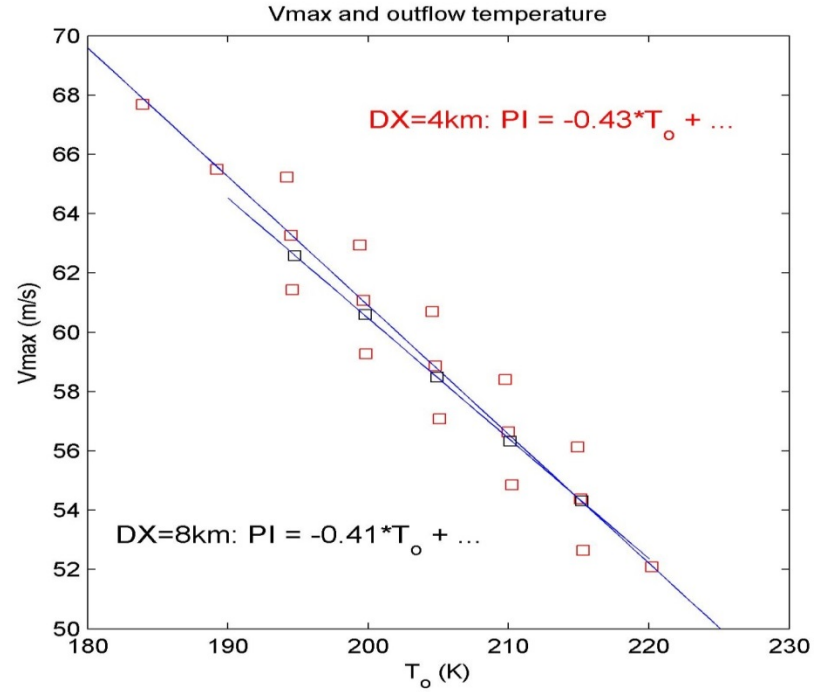


WRF 3D run

(a) Maximum tangential wind (m/s)

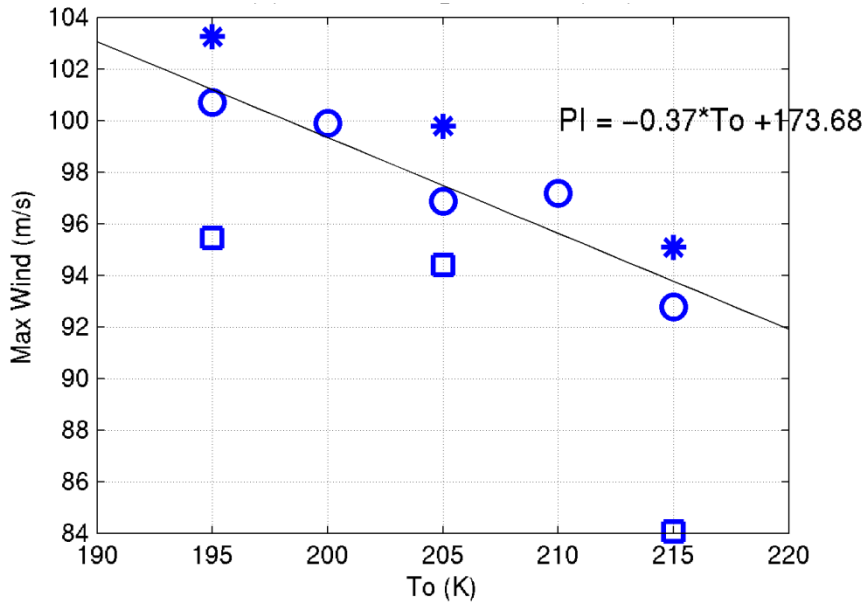


PI calculation

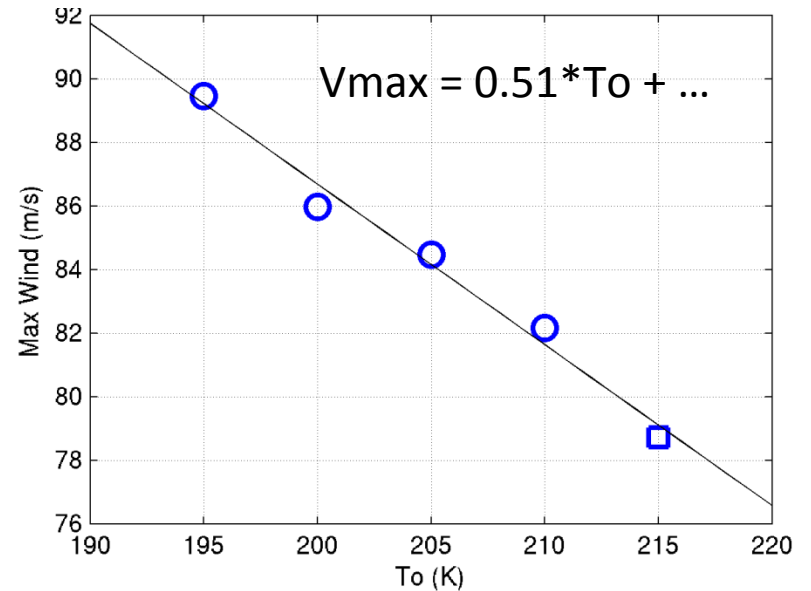


Resolution dependence

Dx=4 km, Slope ~ 0.37



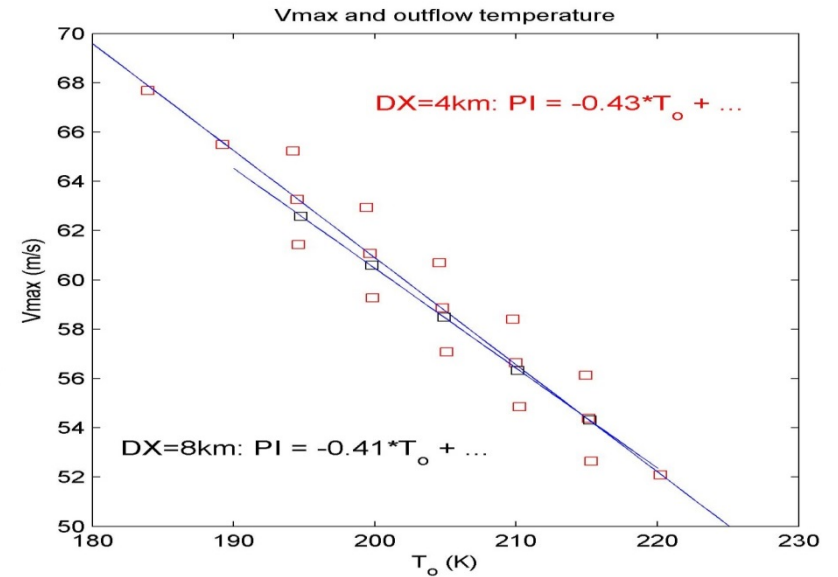
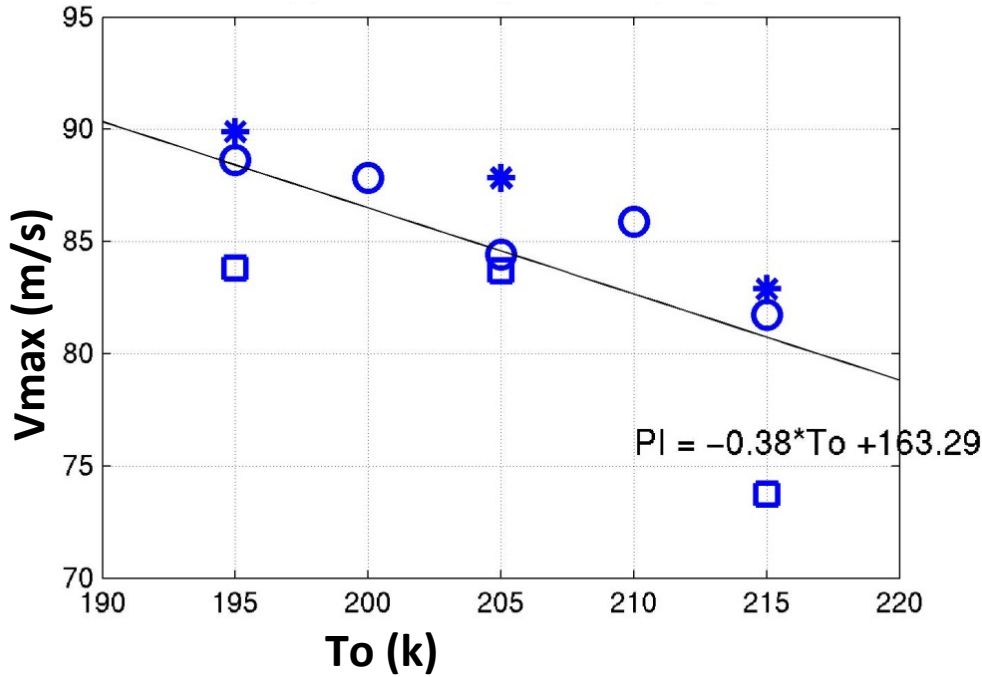
Dx=8 km, Slope ~ 0.51



higher resolution, smaller slope

WRF maximum gradient wind

PI calculation



$$v_g = -\frac{fr}{2} + \left(\frac{f^2 r^2}{4} + r c_p \theta_v \frac{\partial \pi}{\partial r} \right)^{1/2}$$

Still significantly larger than PI values

Hurricane intensity in high-resolution models

- Hakim 2011, 2D model, initial strong wind ~ 120 m/s, then 66 m/s at steady state.
- Hausman et al 2006, > 90 m/s
- Bryan and Rotunno 2009, PI ~ 60 m/s for initial sounding, substantial sensitivity of V_{\max} to turbulence length scale, V_{\max} from 30 -108 m/s, and sensitivity to moist adiabats, 40-104 m/s,
- Persing and Montgomery 2003,
- Yang et al., 2007, $c_d/c_k \sim 0.55$, PI=58 m/s, $V_{\max}=88$ m/s in 2D model, 78 m/s in 3D model
- Nolan et al 2007, WRF model, surface wind $\sim 70\%$ of PI
- Wang and Xu 2010, 3D model, wind about 25-40% stronger than PI (40-50 m/s)

Surface fluxes for PI

$$V^2 \sim \frac{T_0 - T_s}{T_0} \frac{F_s}{C_d \rho |V_s|} \quad F_s: \text{ surface fluxes, } V_s: \text{ surface wind speed}$$

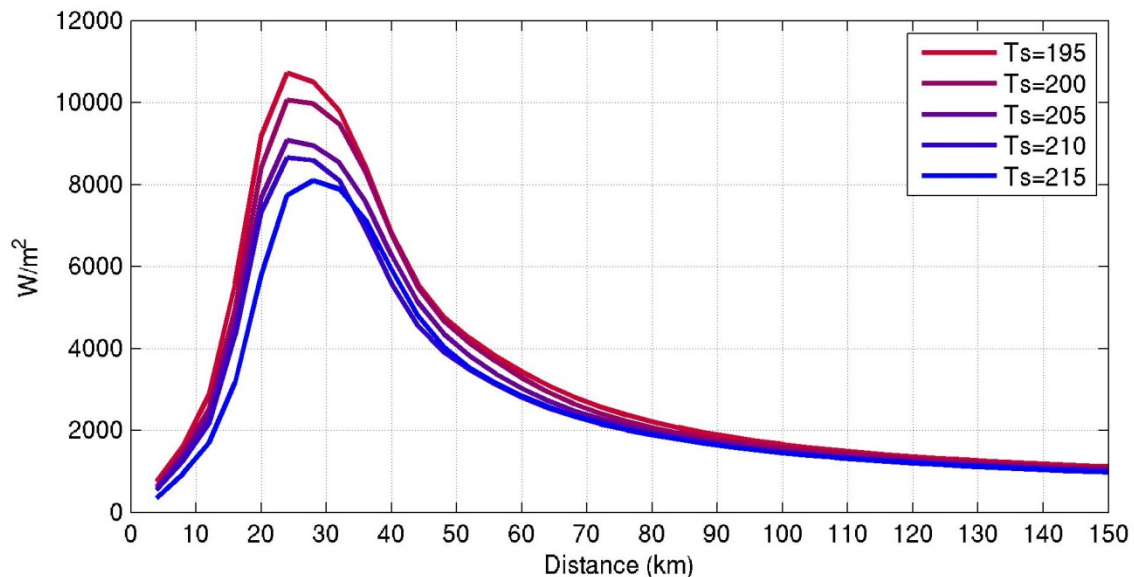
RCE: $F_s = \text{Radiative Cooling} = \sim 120 \text{ W/m}^2$

Assume that $V_s \sim 70\%$ of V_{max} , estimate of V_{max} can be made:

$V_{\text{max}} \sim 40 \text{ m/s}$, varies little with T_s

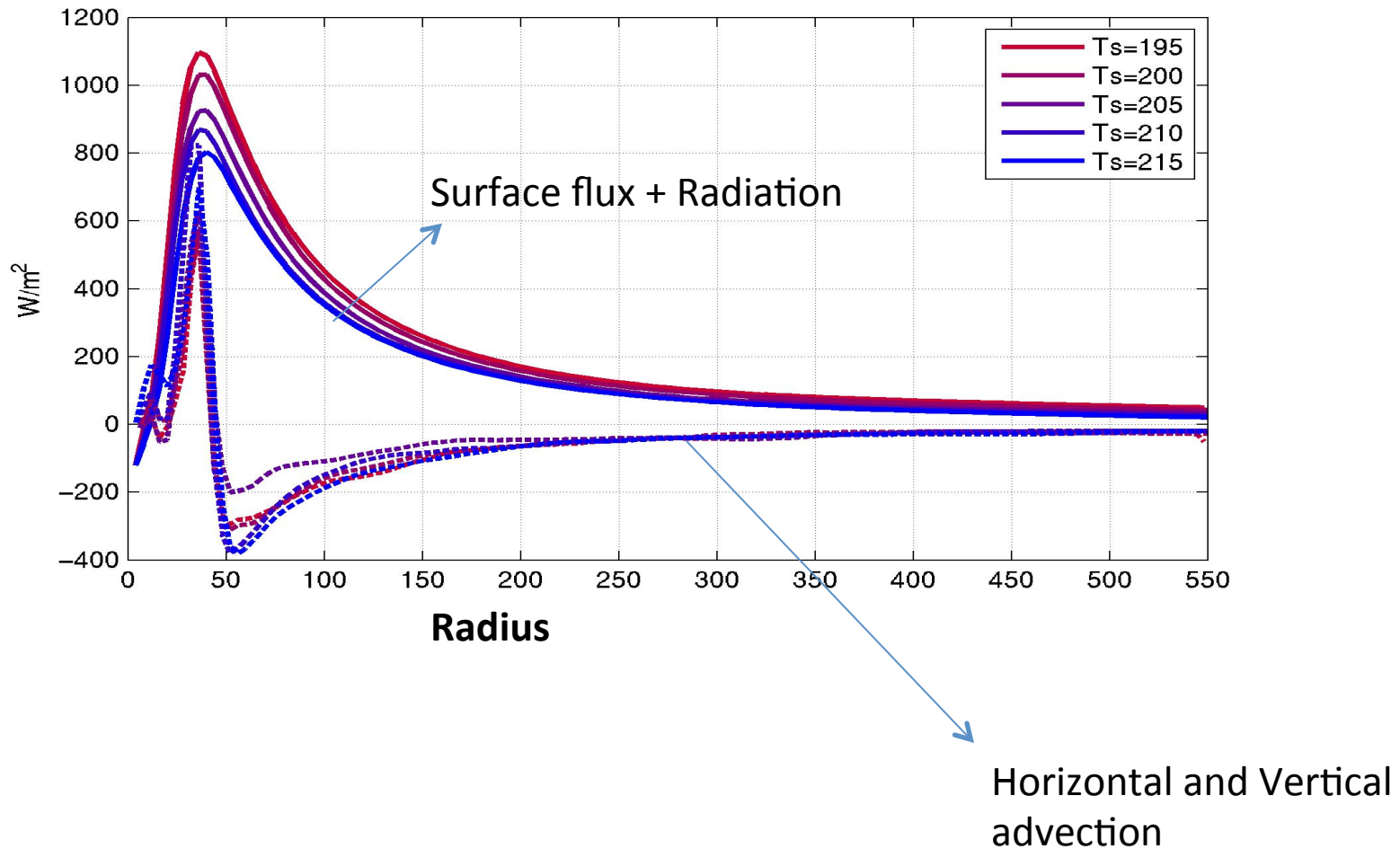
Surface fluxes

$$V^2 \sim \frac{T_0 - T_s}{T_0} \frac{F_s}{C_d \rho |V_s|} \quad F_s: \text{ surface fluxes, } V_s: \text{ surface wind speed}$$



If we use these numbers, V_{max} will be too large: 140 m/s

Surface fluxes + radiative cooling = Horizontal + vertical advection of moist static energy



Conclusion

- ❑ WRF simulated RCE soundings: ~ 0.5 m/s per degree cooling in the stratosphere, but this is sensitive to assumptions in the PI code
- ❑ Similar results from 3D simulations of transient hurricanes using the RCE soundings as initial state, but much stronger hurricanes
- ❑ Issues in model simulations: resolution, domain size, no dissipative heating, fixed radiation etc.

