

Response of the Atlantic Ocean Circulation to Greenland Ice Sheet Melting

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Weijer et al., GRL 2012

Atlantic Meridional Overturning Circulation (AMOC)

- How sensitive is AMOC to GrIS freshwater input?
- 2007 mass deficit: 267 ± 38 Gt/yr (Rignot et al. 2008)
 - ~ 0.01 Sv
 - Freshwater input expected to increase
 - Surface melting
 - Precipitation
 - Calving
- Studied exclusively in low-resolution ocean models
 - No explicit eddies, effects parameterized
 - Sluggish, broad boundary currents
 - Freshwater input over large area Greenland Ice Sheet

Research Question

- **Does enhanced model resolution modify AMOC sensitivity?**
 - More accurate representation of ocean transports
 - Boundary currents
 - Eddy fluxes
 - Less reliance on parameterizations
- **Does result depend on spatial pattern of freshwater forcing?**
- **Approach**
 - Perform freshwater sensitivity studies in hierarchy of models
 - 1 degree and 0.1 degree
 - Different freshwater flux patterns
 - Realistic distribution of freshwater around Greenland
 - Hosing

The Model

- **Los Alamos' Parallel Ocean Program (POP 2.0)**
 - Global domain
 - Ocean-only
- **2 configurations**
 - 1° ("by-one")
 - non-eddying
 - IPCC class
 - 40 levels
 - Dipole, displaced pole over Greenland
 - 0.1° ("point-one")
 - strongly eddying
 - 42 levels
 - tripole

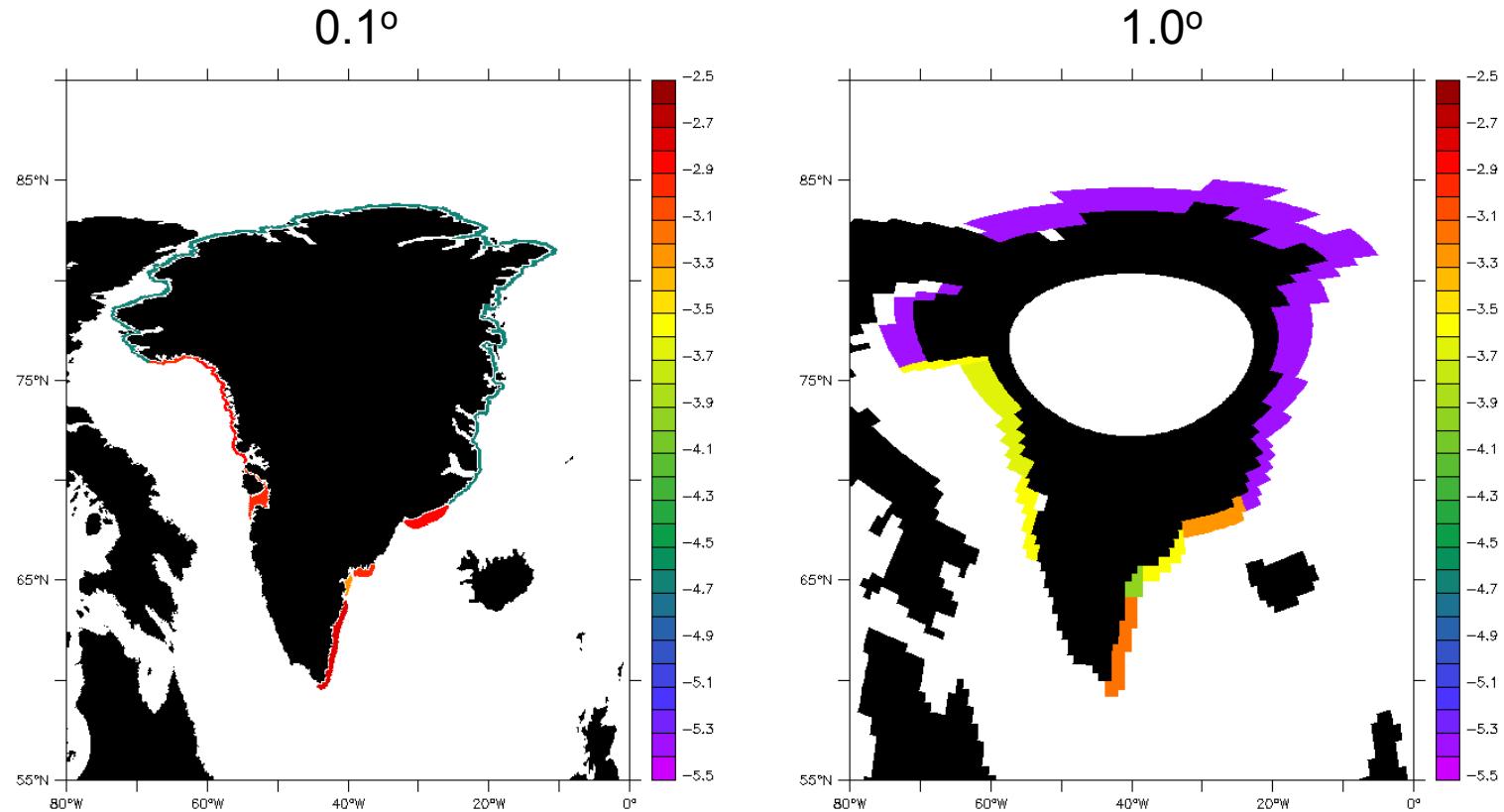
Forcing

- **CORE monthly-mean normal-year atmospheric climatology**
 - Large & Yeager (2004) data
 - Standard bulk formulae
- **SST**
 - Latent heat flux
 - Sensible heat flux
 - Radiation
- **SSS**
 - Precipitation
 - Evaporation
 - Run-off
 - Flux correction
 - Mixed boundary conditions
 - No relaxation

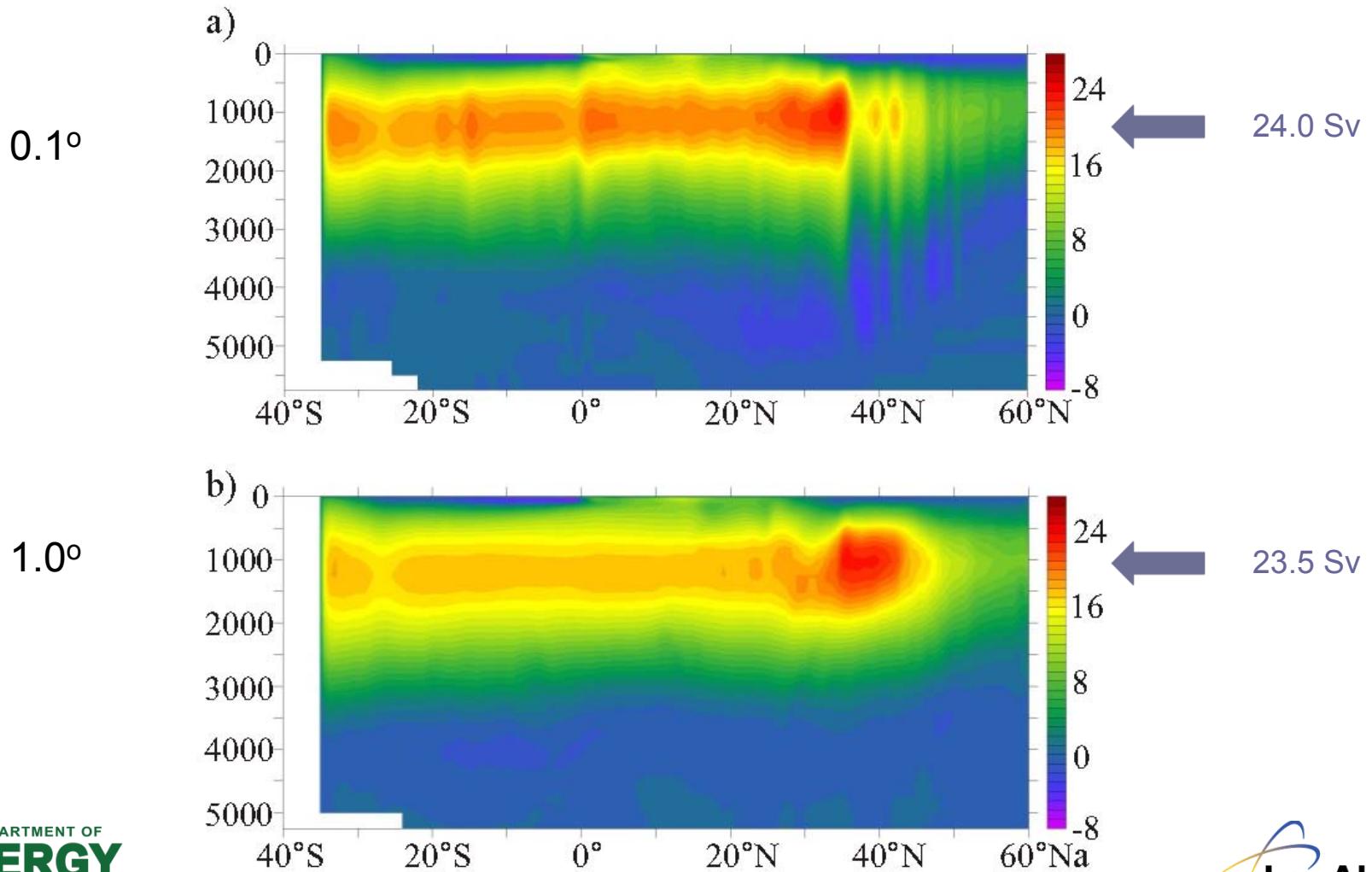
Experimental Procedure: Spin-up + 3 Branches

- **C-Mixed: Control Integration**
- **E-Greenland: Greenland Freshwater**
 - ‘Realistic’ near-coastal distribution around Greenland (*Rignot & Kanagaratnam 2006*)
 - Time-mean amplitude 0.1 Sv (*Gerdes et al. 2006*)
 - Seasonal variation
- **E-Hosing: Traditional Hosing**
 - Distribution North Atlantic 50°-70°N (CMIP)
 - Constant 0.1 Sv flux

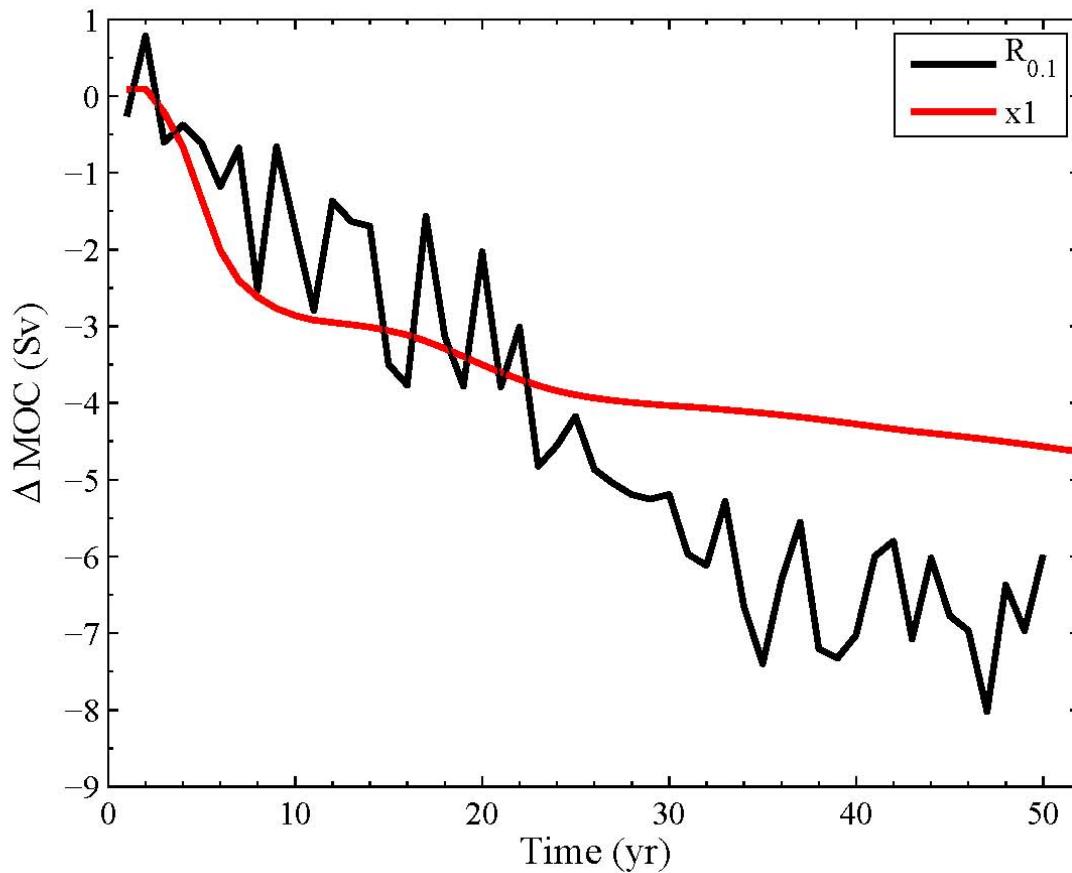
E-Greenland Freshwater Perturbation



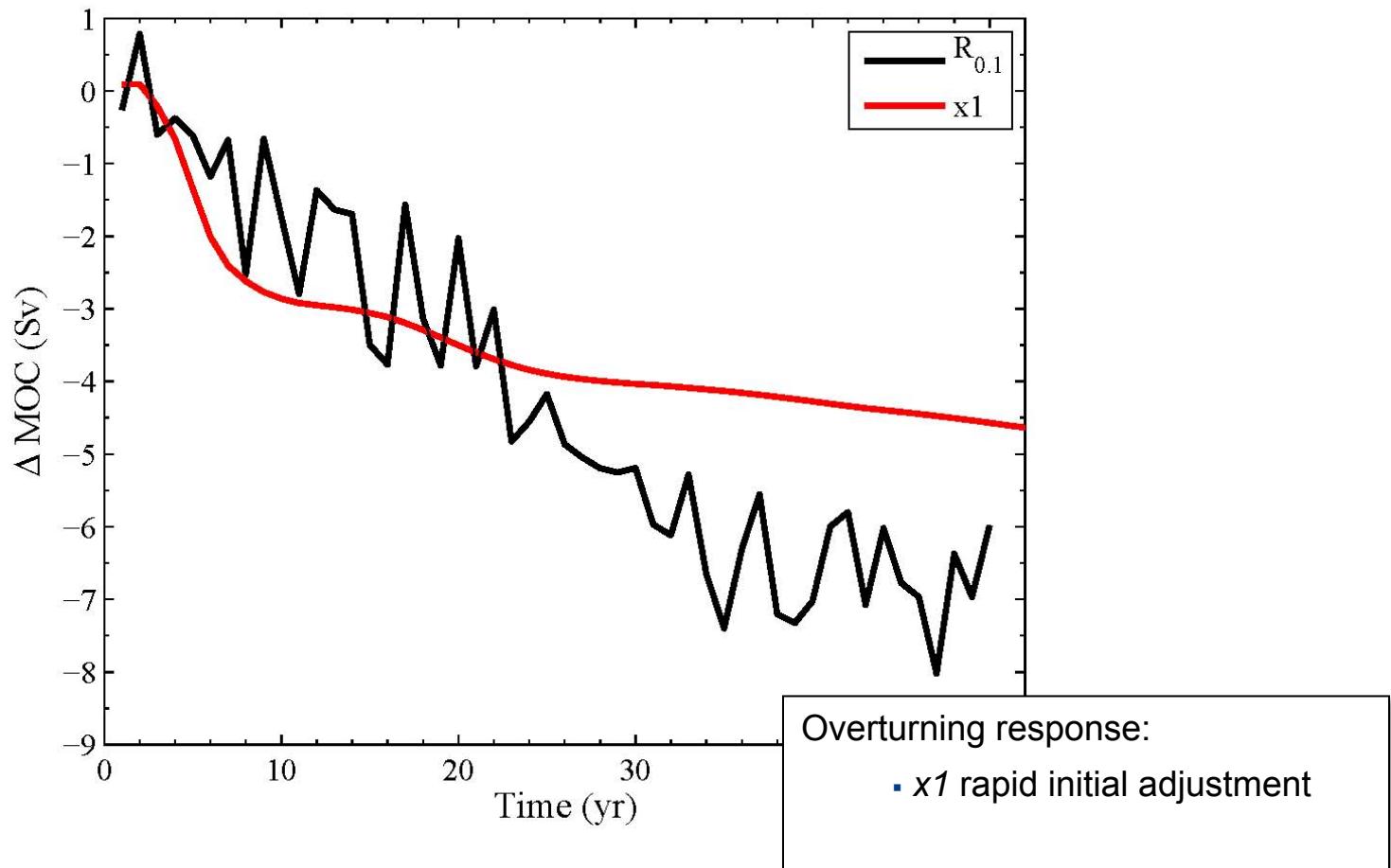
OVERTURNING STREAMFUNCTION



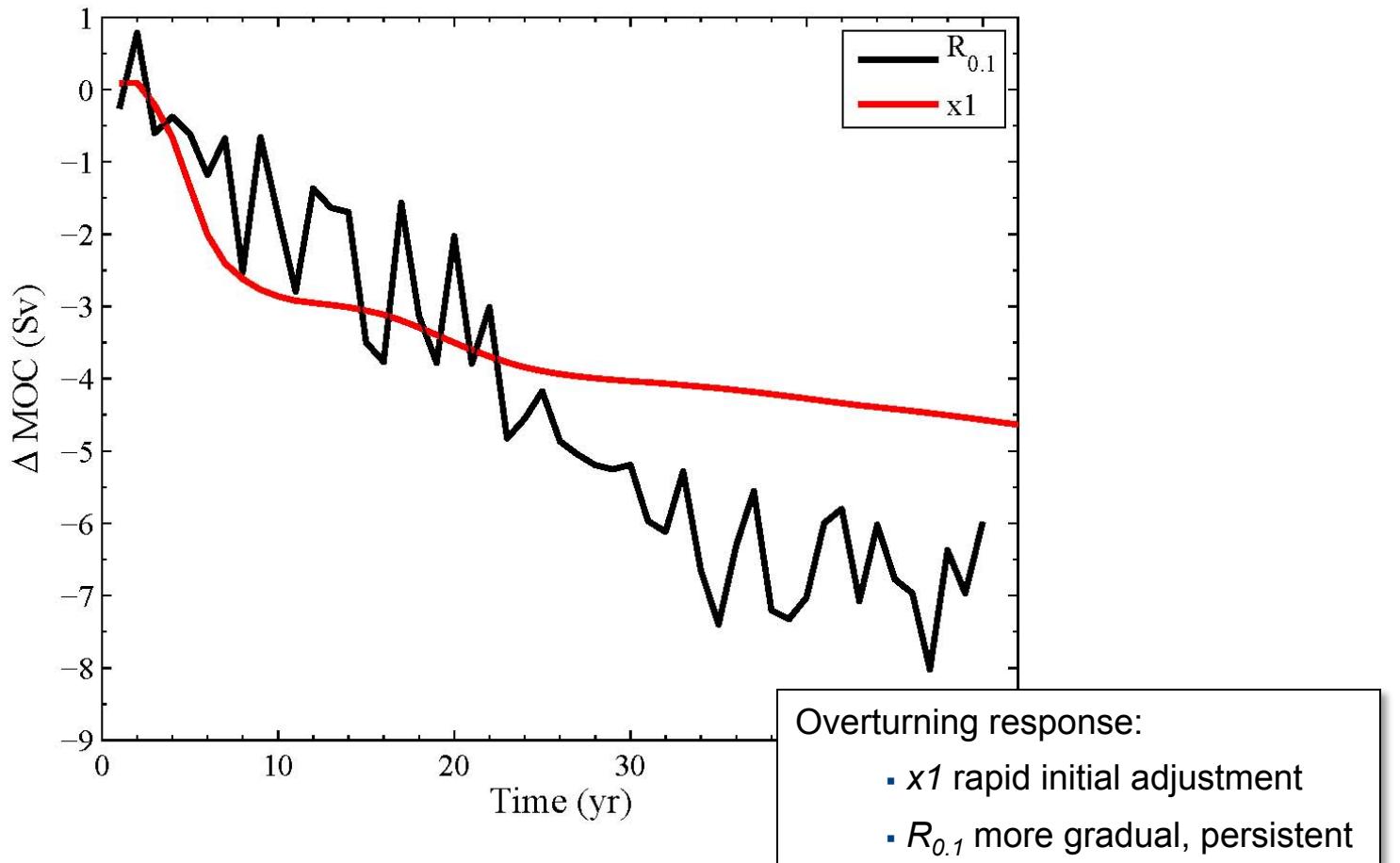
OVERTURNING RESPONSE: E-GREENLAND



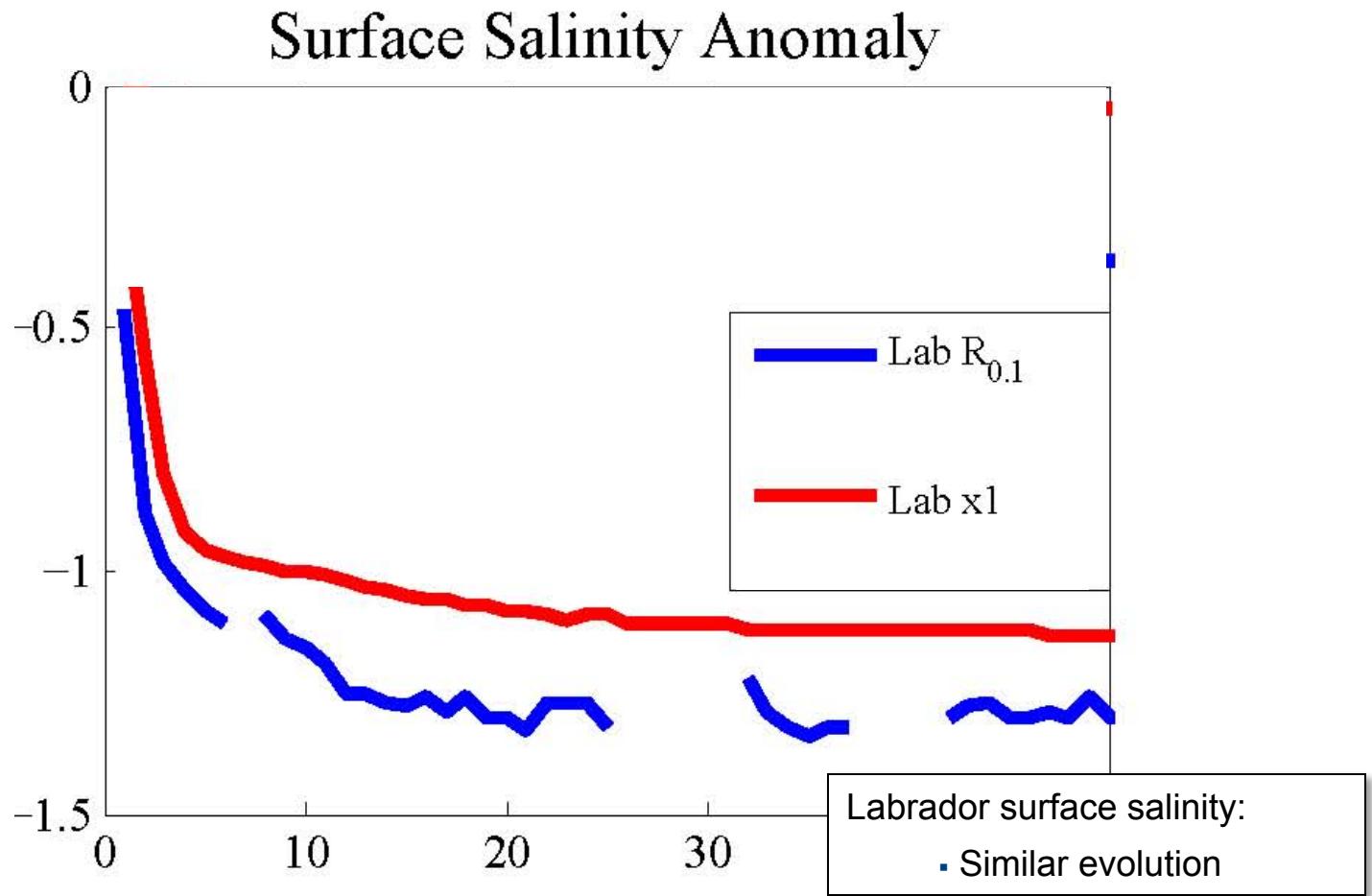
OVERTURNING RESPONSE: E-GREENLAND



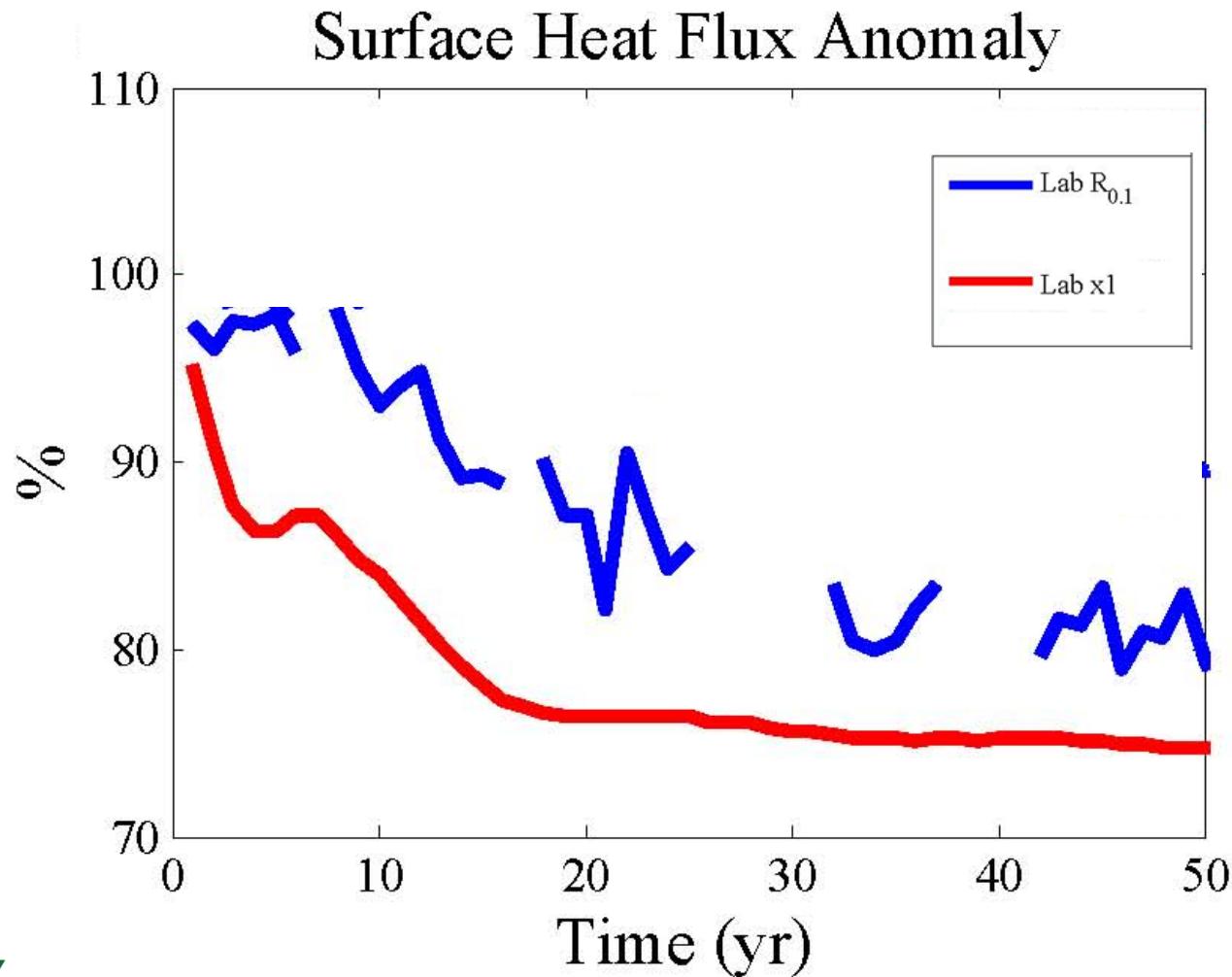
OVERTURNING RESPONSE: E-GREENLAND



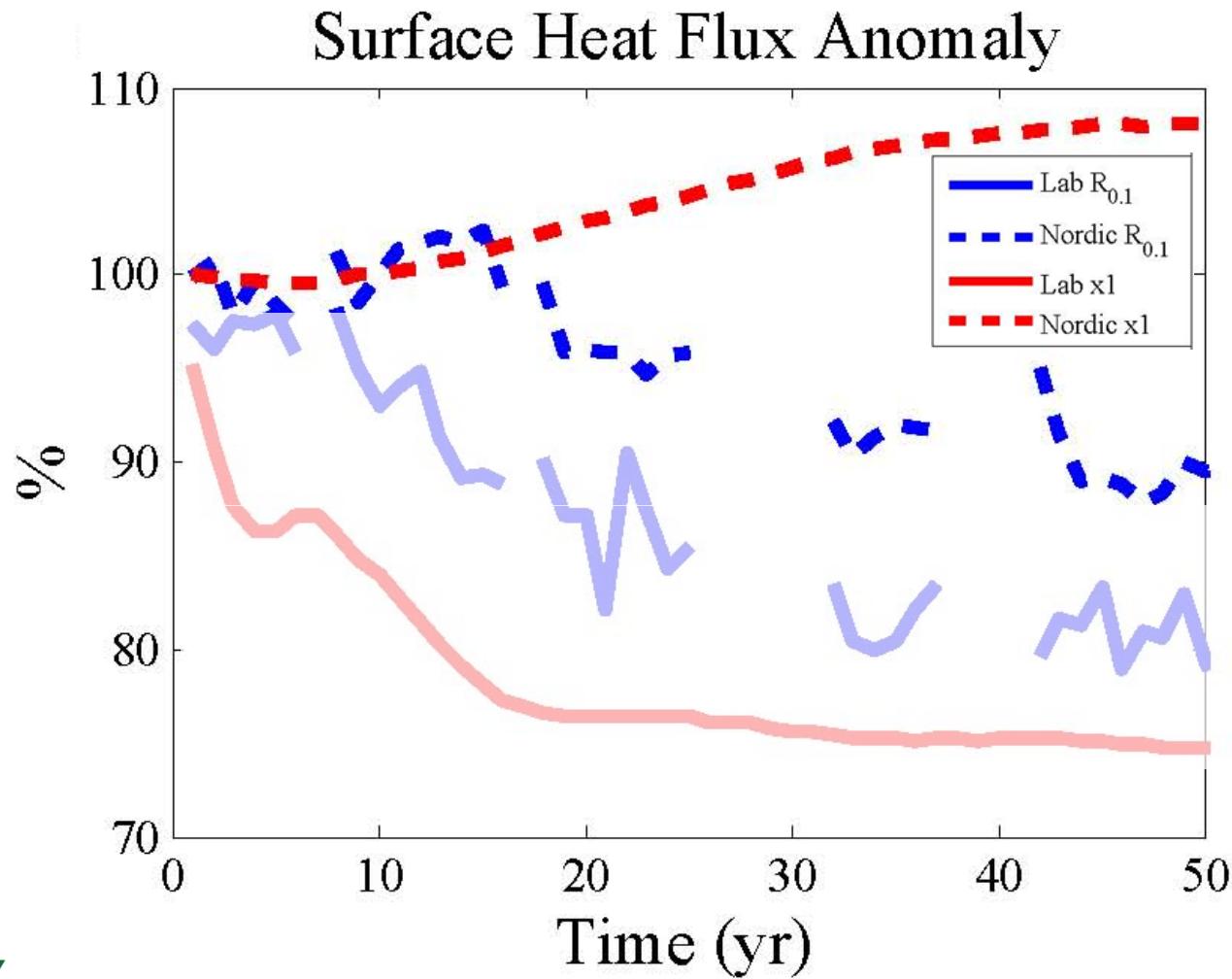
Surface Salinity: E-Greenland



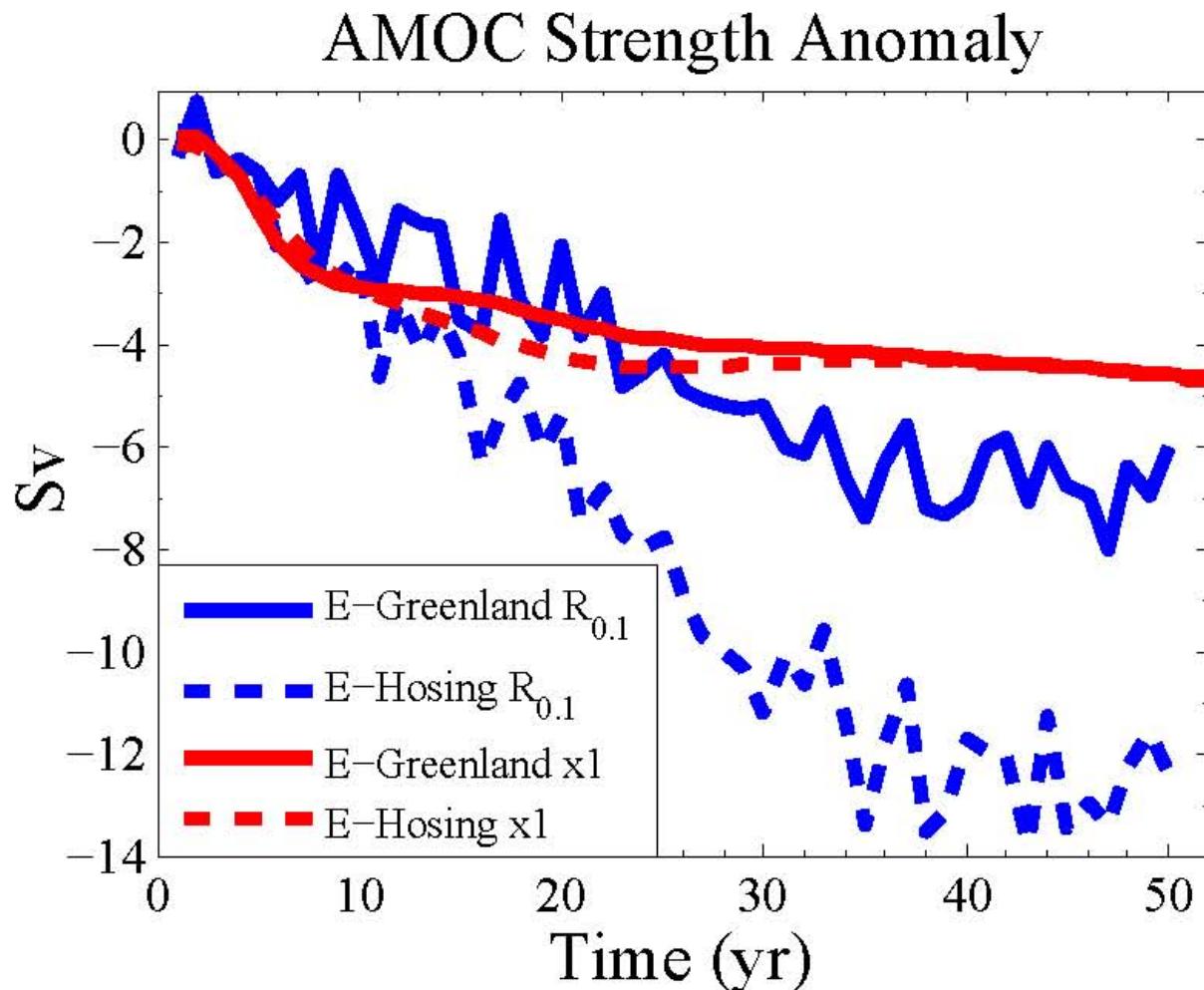
Surface Heat Flux: E-Greenland



Surface Heat Flux: E-Greenland



OVERTURNING RESPONSE: HOSING



New Results: 0.5 Sv Perturbation

Den Toom et al. (in prep.)

$$\psi(\varphi, b, t) = -r_0 \cos \varphi \int_{\lambda_w}^{\lambda_E} \int_{-H}^{\xi(\lambda, \varphi, b, t)} v(\lambda, \varphi, z, t) dz d\lambda \quad \text{Streamfunction}$$

$$V(\varphi, b, t) = r_0 \cos \varphi \int_{\varphi}^{90N} \int_{\lambda_w}^{\lambda_E} \int_{-H}^{\xi(\lambda, \varphi, b, t)} dz d\lambda d\varphi \quad \text{Volume}$$

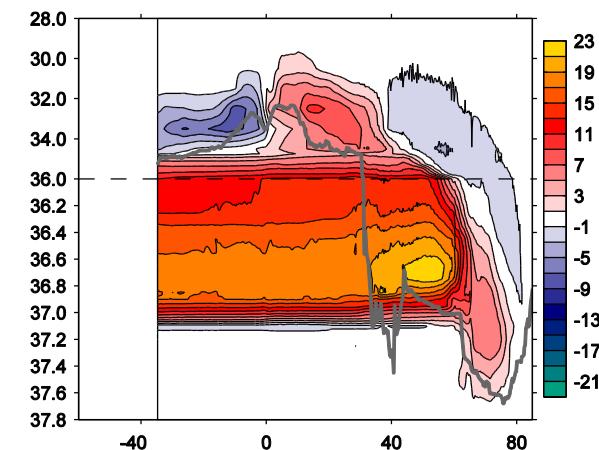
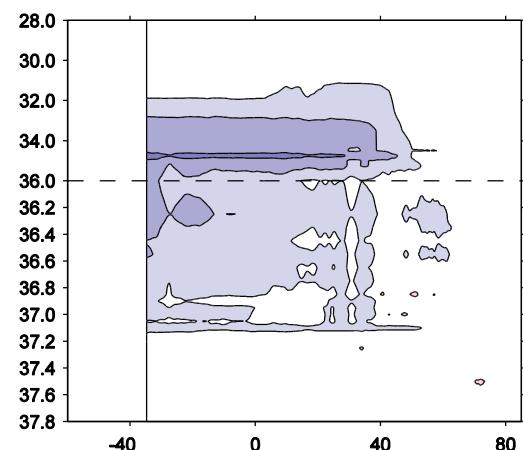
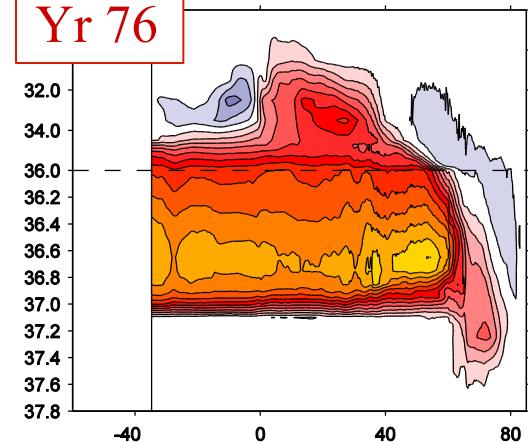
$$G(\varphi, b, t) = \psi(\varphi, b, t) + \frac{\partial V(\varphi, b, t)}{\partial t} \quad \text{Transformation streamfunction}$$

e.g., Marsh et al. (2000); Wolfe & Cessi (2011)

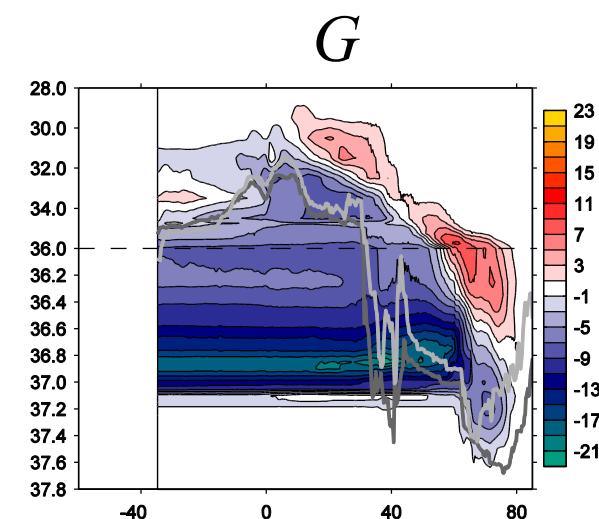
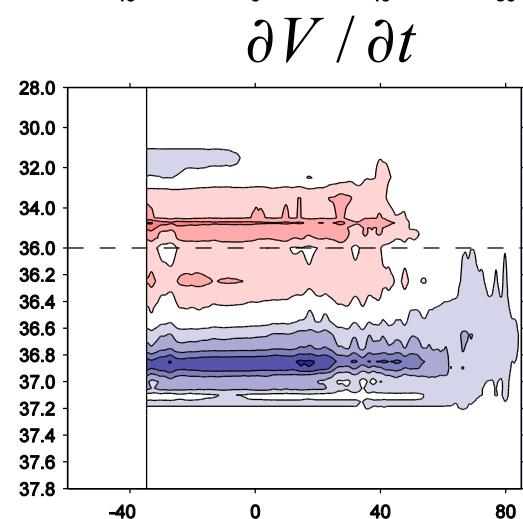
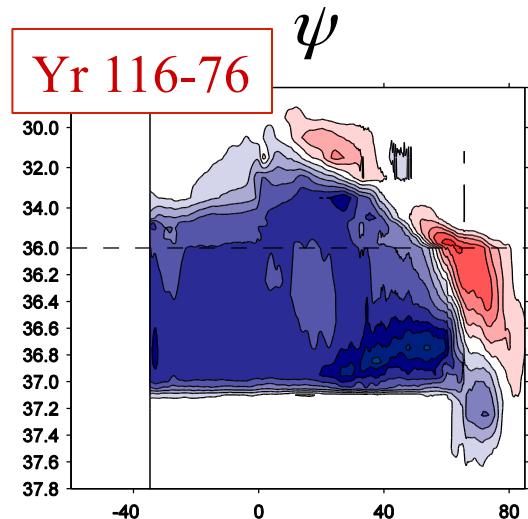
New Results: 0.5 Sv Perturbation

Den Toom et al. (in prep.)

Yr 76



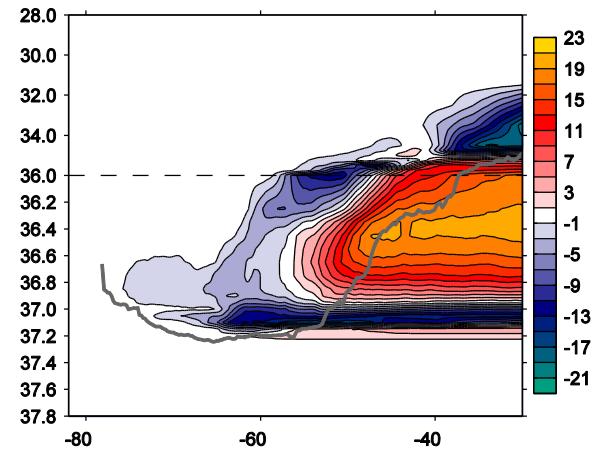
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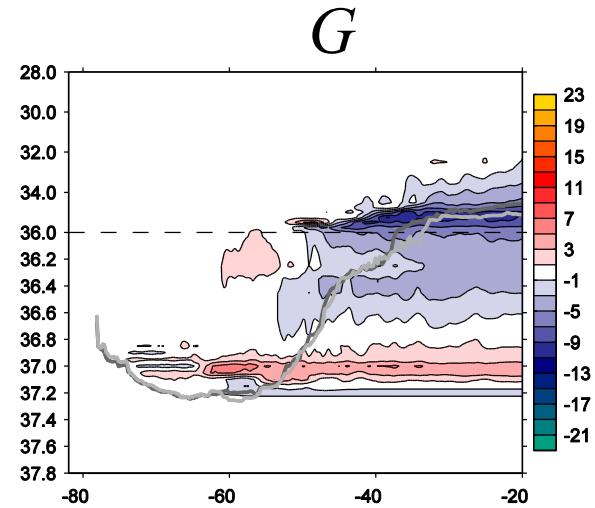
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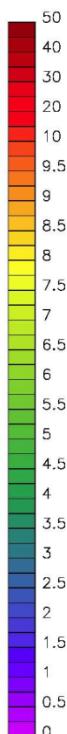
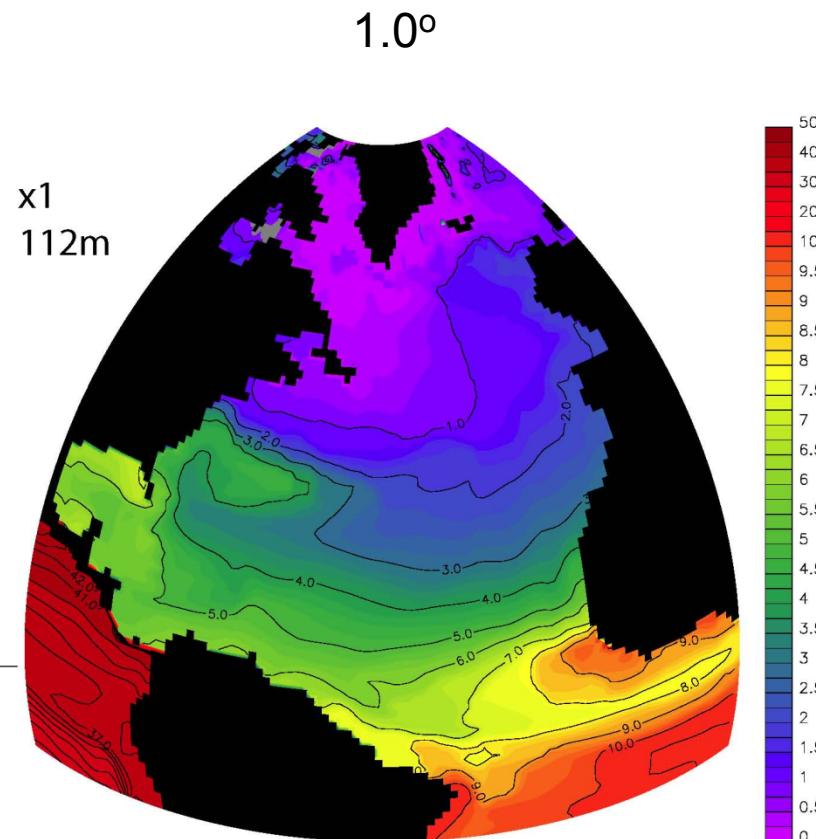
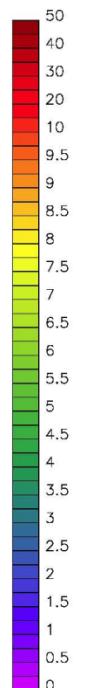
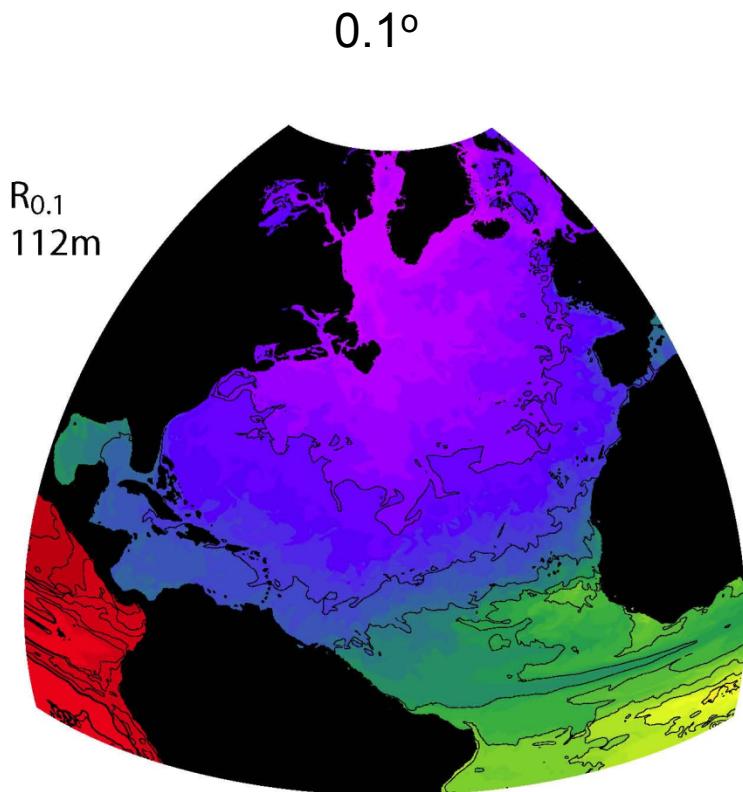
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Conclusions

- **Greenland freshwater flux perturbations**
 - No quantitative difference in MOC response
 - Qualitative differences
 - Rapid initial adjustment in $x1$
 - Gradual, more persistent in $R_{0.1}$
 - Deep convection more resilient to freshwater perturbation
 - Increase of Nordic Seas ventilation in $x1$
- **E-Greenland vs. E-Hosing**
 - No difference in $x1$
 - $R_{0.1}$ much more sensitive to *spatial distribution*
- **Adiabatic pole-to-pole: limited predictive value for transient response**

Dye Arrival Time (112 m)



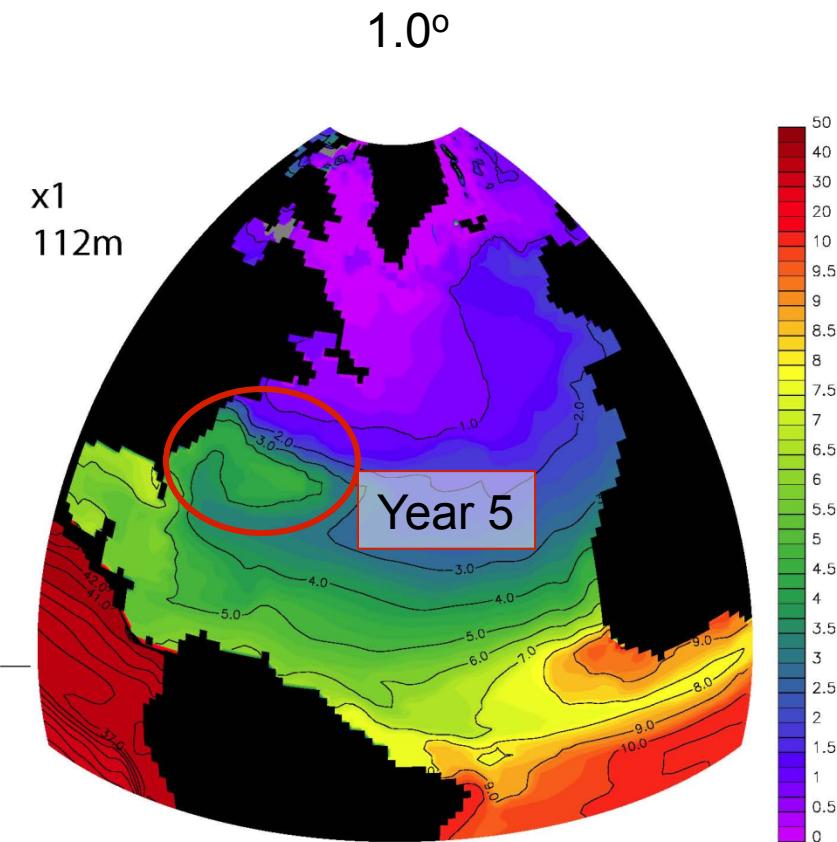
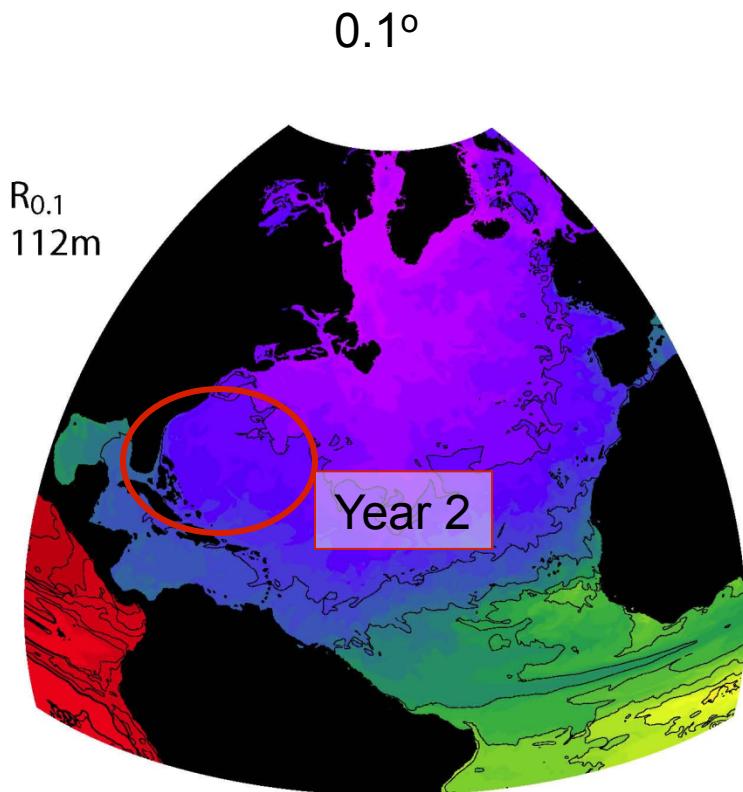
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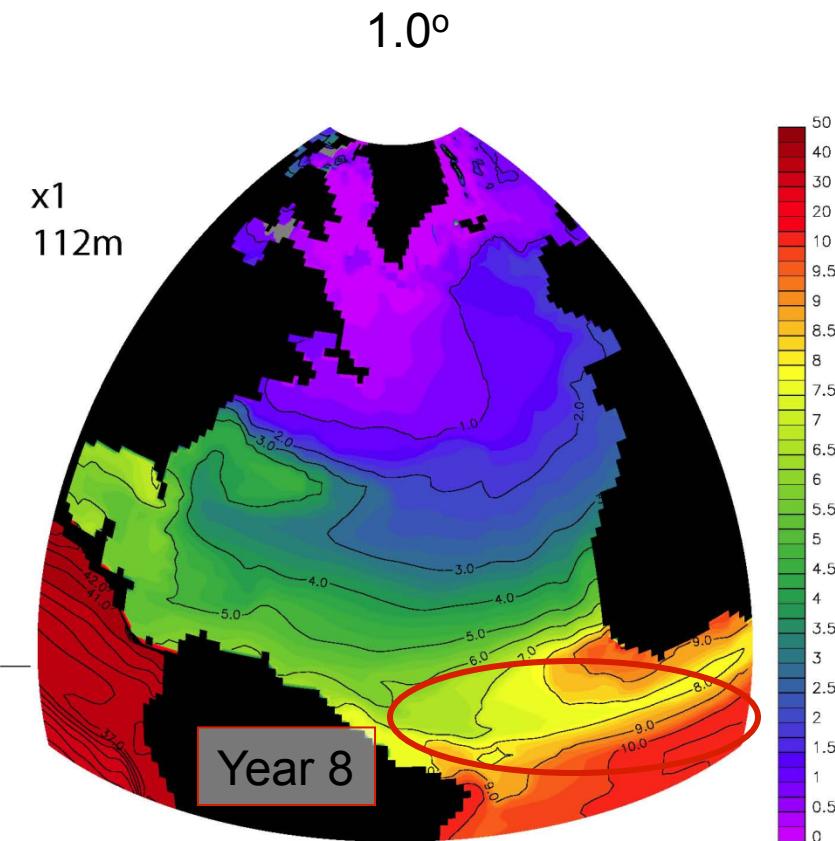
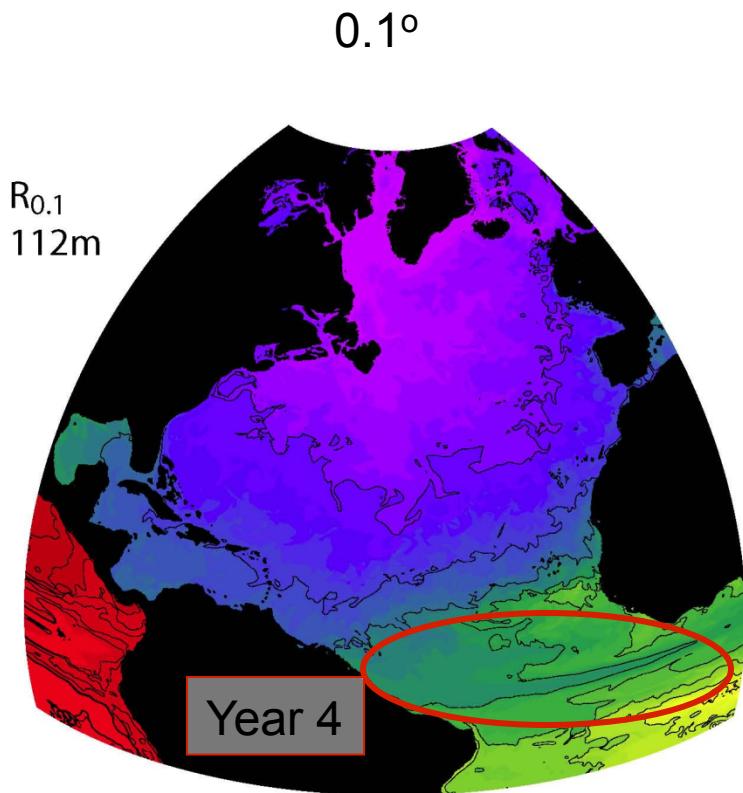
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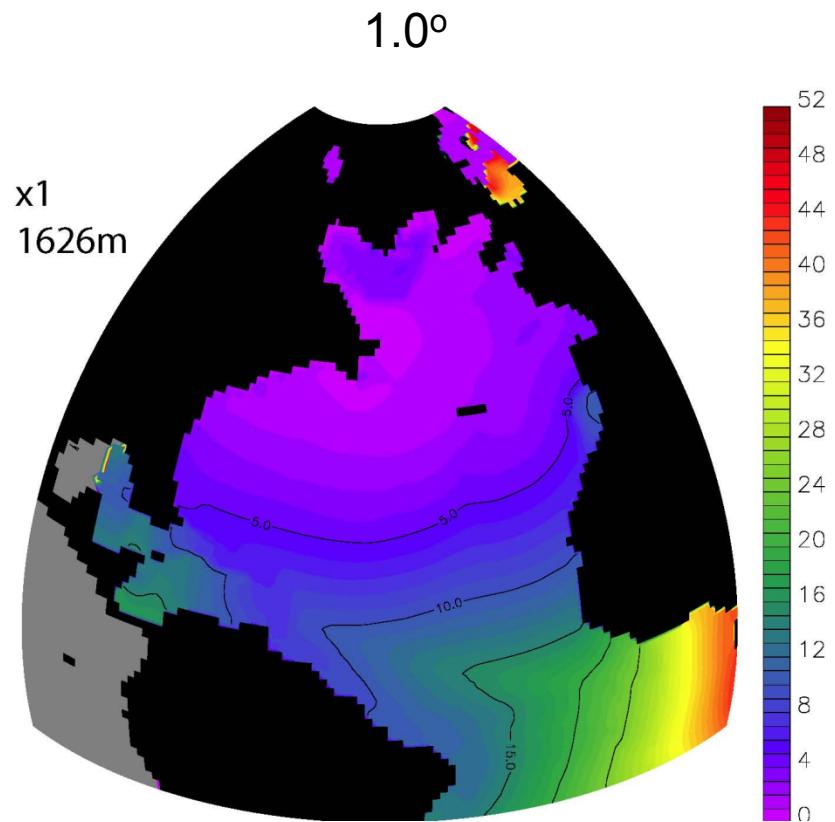
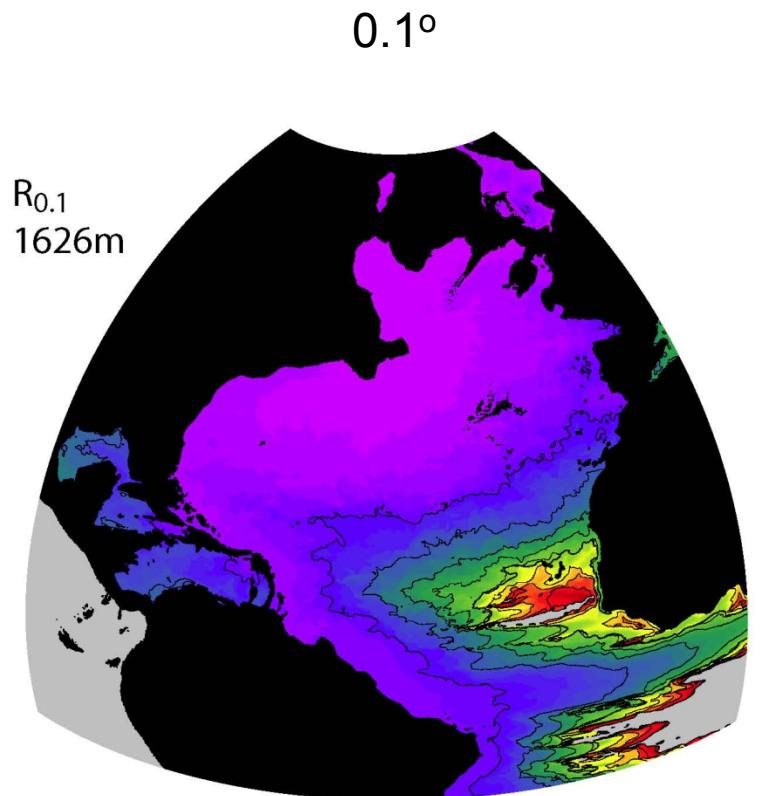
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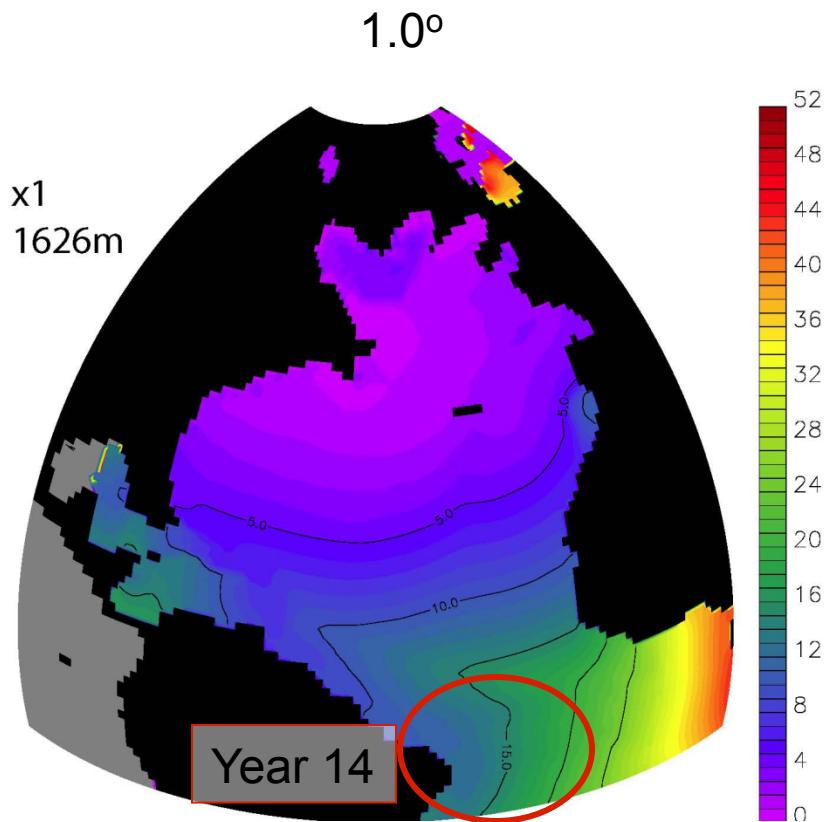
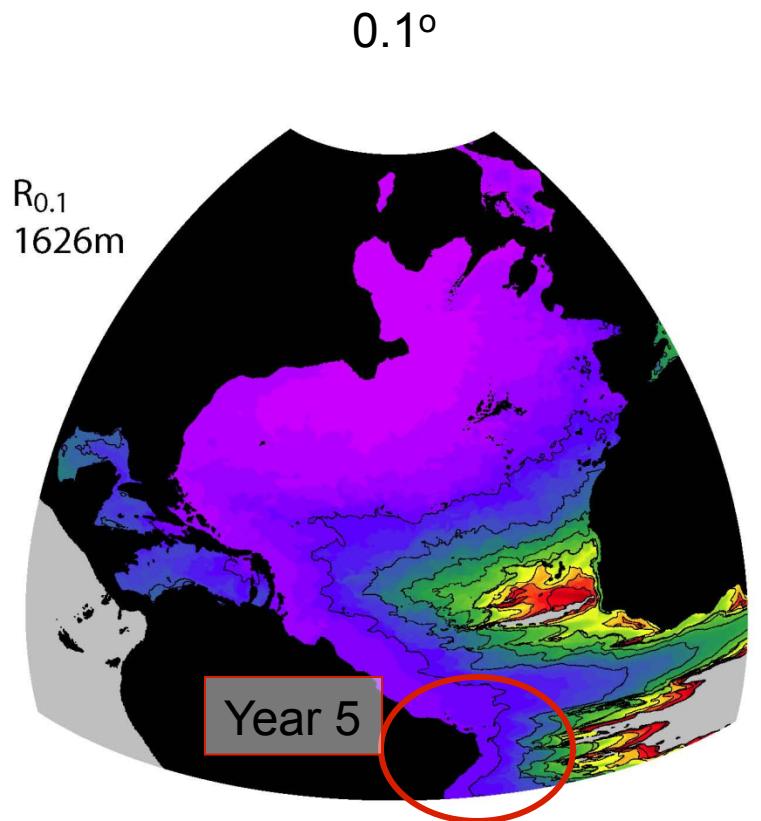
Dye Arrival Time (1626 m)



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