

A satellite image of a tropical cyclone, showing a well-defined eye and a dense, swirling cloud structure. The eye is a small, clear circular area in the center, surrounded by a thick, white ring of clouds. The outer clouds are more diffuse and spread out, creating a large, circular storm system. The colors range from light green and yellow to white and grey, indicating different cloud heights and densities.

On the relationship between PI and CAPE

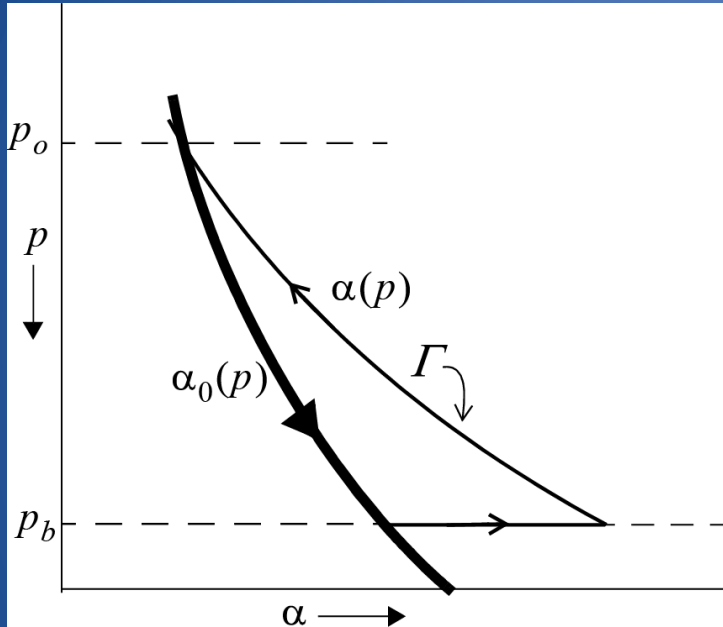
Steve Garner, GFDL

June 5, 2013

1. Approximations in the CAPE algorithm*
2. Approximations in the closed formulas
3. Updates to the closed formulas

*Bister and Emanuel *JGR* 2002

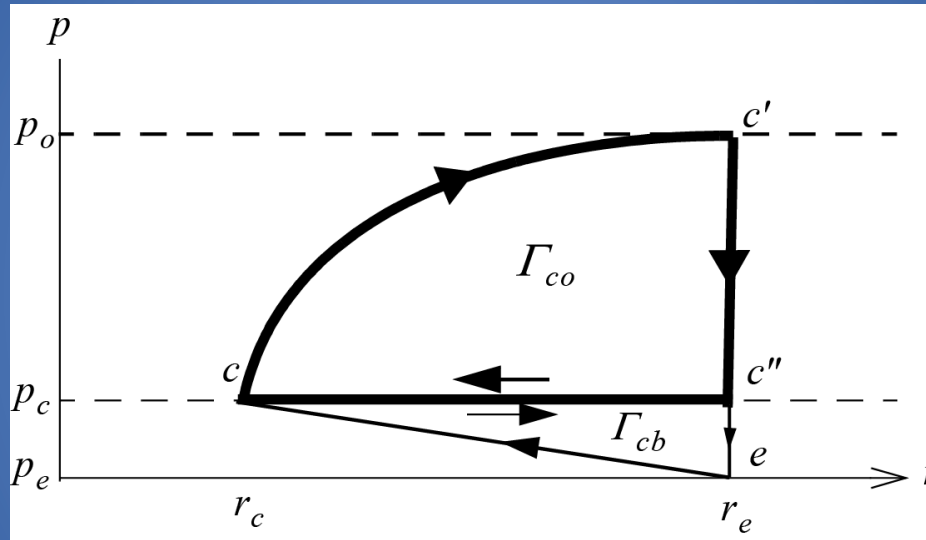
CAPE defined as a contour integral



$$CAPE = \int_{p_o}^{p_b} (\alpha - \alpha_0) dp = - \oint_{\Gamma} \alpha dp$$

“Hurricane CAPE” may have $\alpha \neq \alpha_0$ at bottom

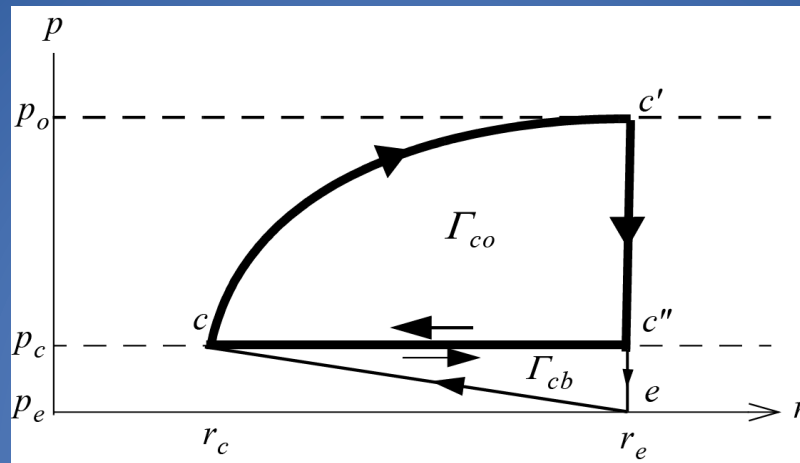
Vertical cross-section in pressure coordinates



Dynamical constraint:

$$-\alpha dp = d(v^2/2 + \Phi)$$

$$\Gamma \quad - \oint_{\Gamma} \alpha dp = -\frac{1}{2}(v^2 - v'^2) + R\overline{T_{vb}} \log\left(\frac{p_e}{p}\right)$$



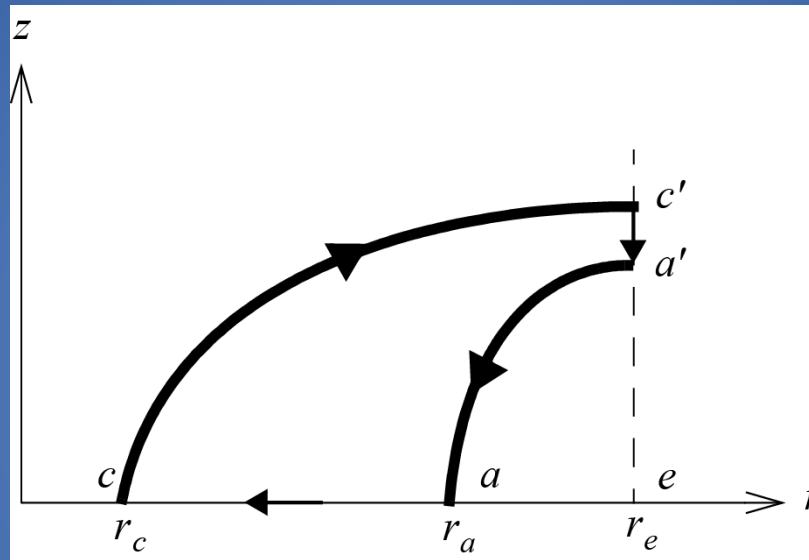
$$\Gamma_o \quad - \int_{\Gamma_o} \alpha dp = CAPE$$

$$\Gamma_b \quad - \int_{\Gamma_b} \alpha dp = R(\overline{T_{vb}} - \overline{T_{vo}}) \log\left(\frac{p_e}{p}\right)$$

$$R\overline{T_{v0}} \log\left(\frac{p_e}{p}\right) = CAPE + \frac{1}{2}(v^2 - v'^2)$$

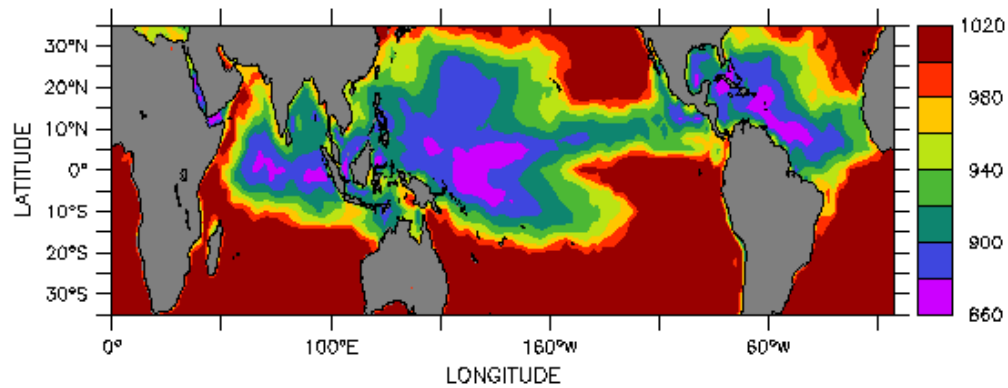
Holland 97
Bister-Emanuel 02

If SLP is known at a second overturning radius (a)

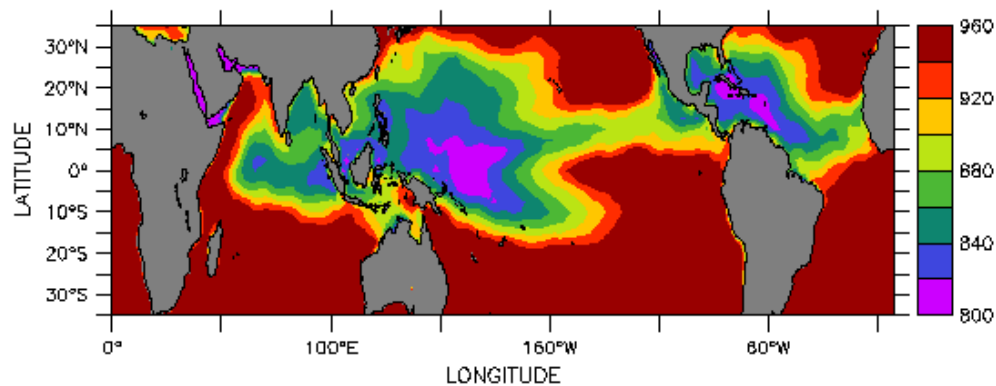


$$R\overline{T_{v0}} \log \left(\frac{p_a}{p_c} \right) = CAPE_c - CAPE_a + \frac{1}{2} (v_c^2 - v_a^2)$$

Emanuel 86,95
Emanuel code



IEW pressure PI (hPa, Emanuel-pseudo)



IEW pressure PI (hPa, Holland-pseudo)

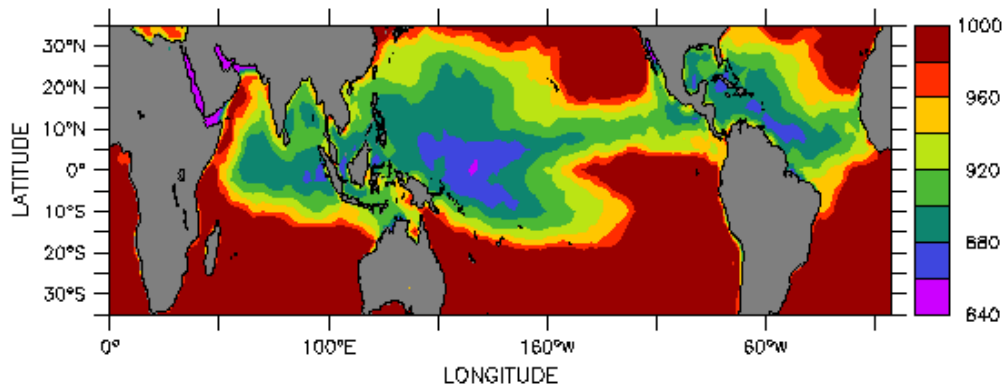
$$R\overline{T_{v0}} \log \left(\frac{p_a}{p_c} \right) = CAPE_c - CAPE_a$$

Take $p_a = p_e$; solve for p_c

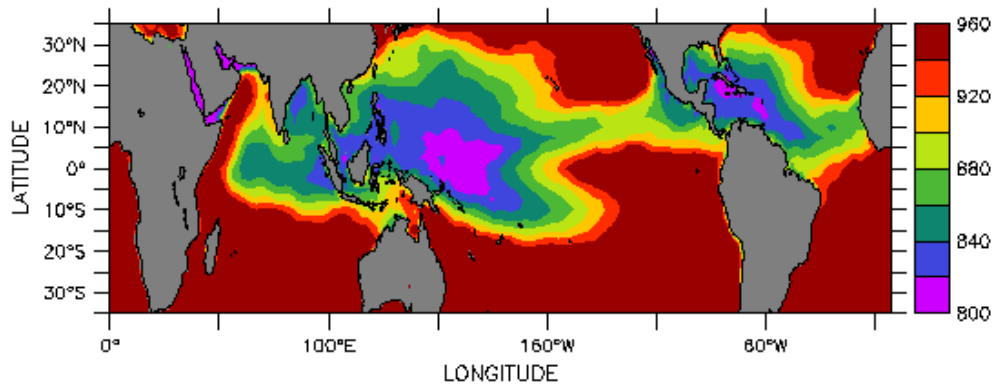
Pressure PI with $v^2 = 0$

$$R\overline{T_{v0}} \log \left(\frac{p_e}{p} \right) = CAPE$$

Solve for p



OEW pressure PI (hPa, E95-pseudo)



OEW pressure PI (hPa, BE02-pseudo)

$$p_a = p_e \quad (\text{code})$$


Pressure PI with $v^2 \neq 0$

$$p_a < p_e \quad (\text{BE 2002})$$

Definitions of PI

	IEW ($v^2 = 0$)	OEWS ($v^2 > 0$)
$p_a = p_e$	E86, E88	E95, code
$p_a < p_e$	H97	BE02

Velocity closure

$$-\alpha dp = T ds - dk$$

$$-\frac{1}{2}(v^2 - v'^2) + R\overline{T_{vb}} \log\left(\frac{p_e}{p}\right) = (\tilde{T}_b - \tilde{T}_0)(s - s_e)$$

Differentiate once wrt inflow radius:

$$-\left(\frac{1}{r^2} - \frac{1}{r_e^2}\right)M \frac{\partial M}{\partial r} = (T_b - T_o) \frac{\partial s}{\partial r}$$

Then use aerodynamic law:

$$v^2 = \frac{c_k}{c_d} (T_b - T_o)(s^* - s)$$

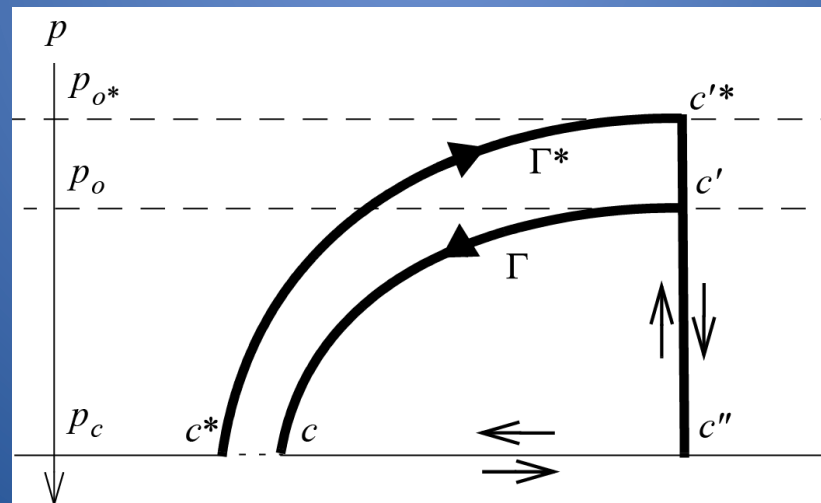
PI algorithm for velocity closure

$$CAPE = - \oint_{\Gamma_o} \alpha dp = (\tilde{T}_B - \tilde{T}_0) [s - s_0(p')]$$

$$\tilde{T}_B = [k - k_0(p')] / [s - s_0(p')]$$

$$CAPE^* = - \oint_{\Gamma_o^*} \alpha dp = (\tilde{T}_B - \tilde{T}_0) [s^* - s_0(p')]$$

$$\tilde{T}_B = [k^* - k_0(p')] / [s^* - s_0(p')]$$



$$CAPE^* - CAPE = (\tilde{T}_B - \tilde{T}_o)(s^* - s)$$

$$\tilde{T}_B = (k^* - k)/(s^* - s)$$

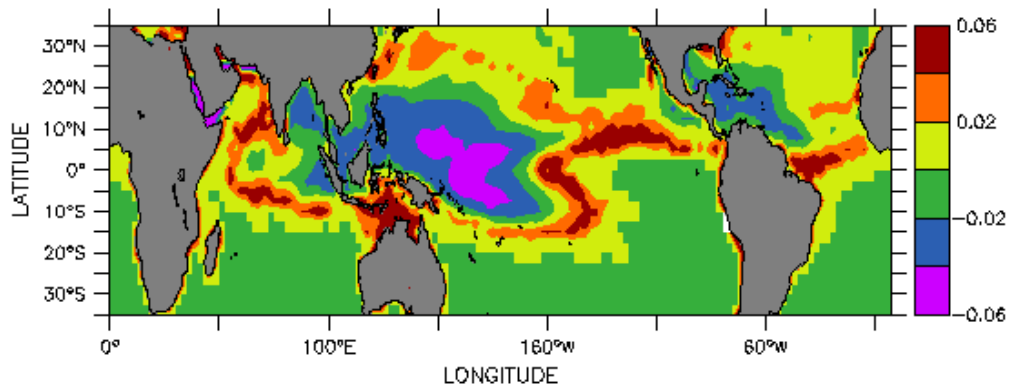
$$\tilde{T}_o = \int_s^{s^*} T_0 ds / (s^* - s)$$

Compare:

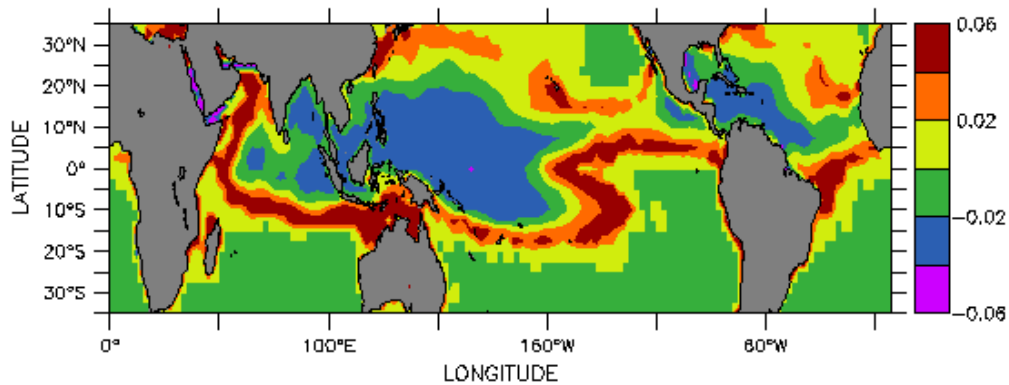
$$v^2 = \frac{c_k}{c_d} (T_b - T_o)(s^* - s)$$

$$\Rightarrow v^2 \approx \frac{c_k}{c_d} (CAPE^* - CAPE)$$

Bister-Emanuel 02



relative error due to averaging (revers)



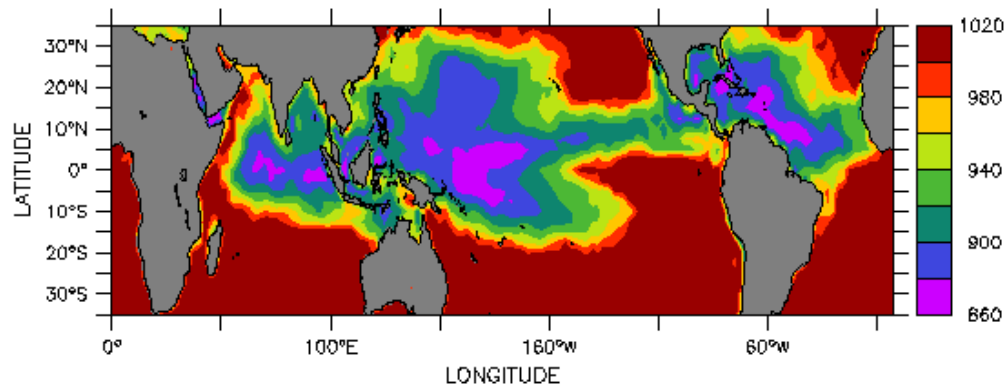
relative error due to averaging (pseudo)

Error from T averaging (%)

$$\tilde{T}_B - \tilde{T}_O \quad \text{vs} \quad T_b - T_o$$

2. Approximations in the thermodynamics: Fully moist vs. “moist-like”

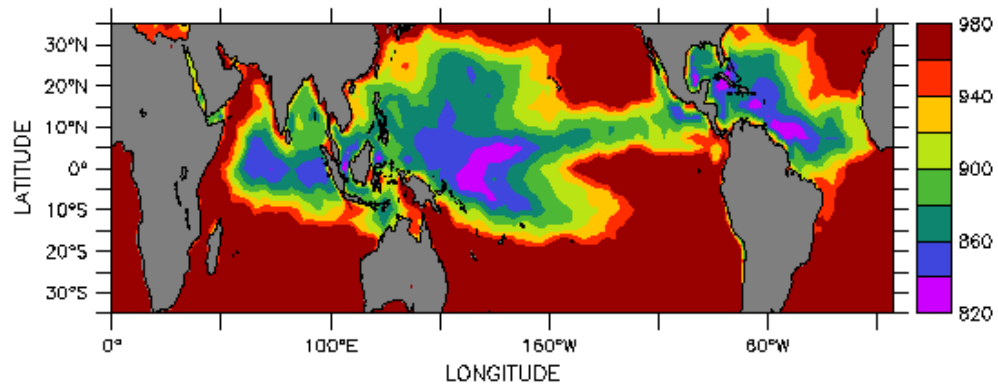
“Moist-like” neglects condensate heat capacity, loading



IEW pressure PI (hPa, pseudo)

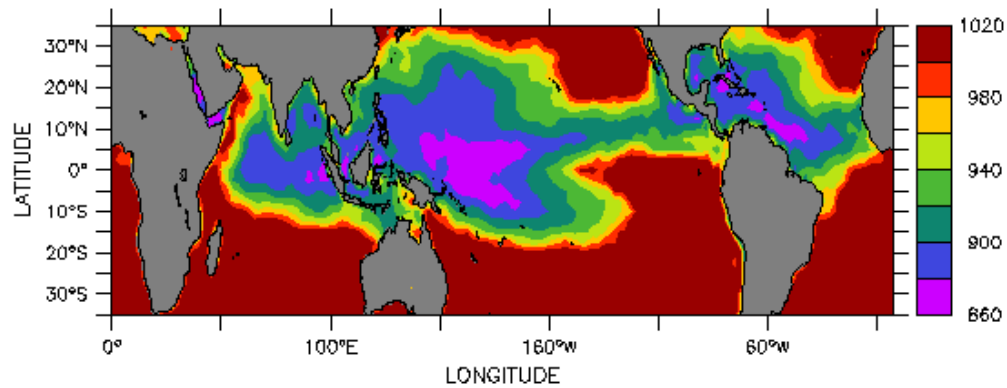
Full moisture (code)

Pressure PI with $v^2 = 0$

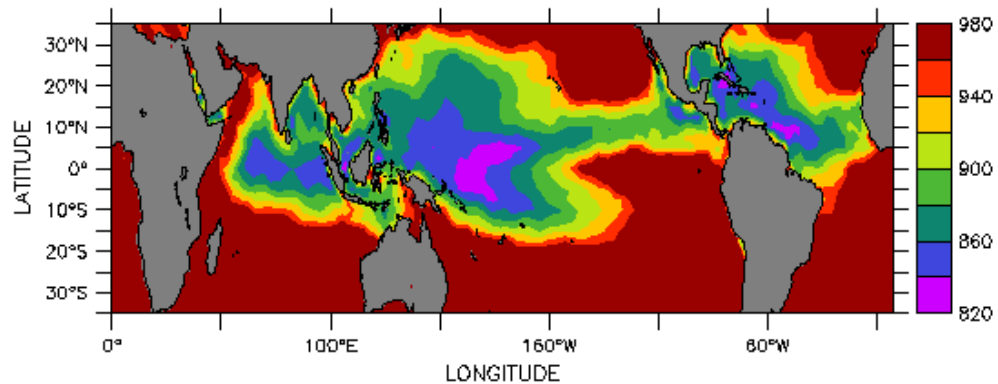


IEW pressure PI (hPa, dry-pseudo)

Moist-like



OEW pressure PI (hPa, pseudo)

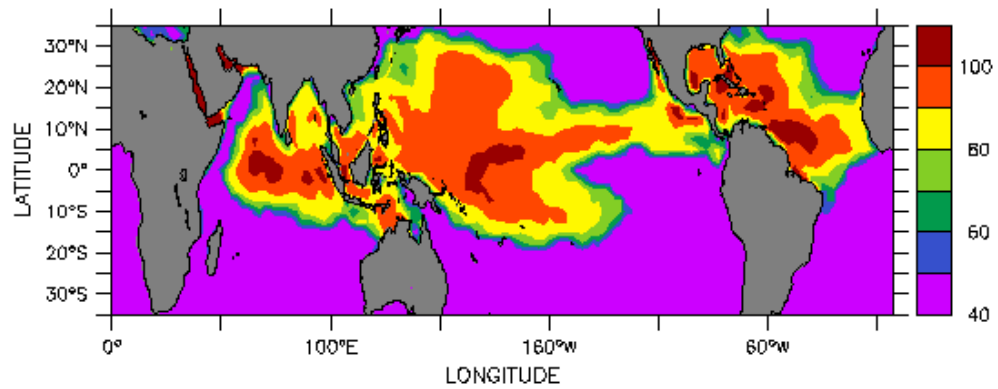


OEW pressure PI (hPa, dry-pseudo)

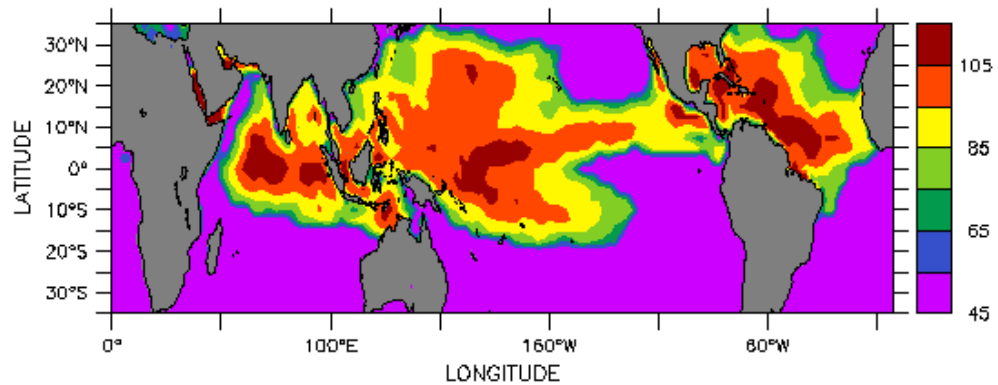
Full moisture (code)

Pressure PI with $v^2 \neq 0$

Moist-like



velocity PI (m/s, pseudo)



velocity PI (m/s, dry-pseudo)

Full moisture (code)

Velocity PI

Moist-like

3. Closed formulas for full moisture

Thermodynamics:

$$(c_p + wc_{pv} + lc_l)dT - (1 + Q)\alpha dp = -Ldw \quad 1^{\text{st}} \text{ law}$$

$$-\alpha dp = Tds - dk + Q\alpha dp + c_l T dQ \quad 1^{\text{st}} \text{ \& } 2^{\text{nd}} \text{ laws}$$

Dynamics:

$$-\alpha dp = d(v^2/2 + \Phi)$$

$$k = (c_p + c_l Q)T + Lw$$

$$Q = w + l$$

$$\Phi = gz$$

Irreversible case (full moisture)

Define:

$$\int_0^{\Phi_o} Q d\Phi = w_b \Phi_q \quad \int_{T_b}^{T_o} Q dT = w_b (\mathbf{T}_q - T_b) \quad \int_{0T_b}^{T_o} Q d\log T = w_b (\Lambda_q - \Lambda_b)$$

$$\overline{(1 + w_b)RT_{vb} \log\left(\frac{p_a}{p_c}\right)} = D_{a,c} + (1 + \tilde{w}_c) \frac{v_c^2}{2} - (1 + \tilde{w}_a) \frac{v_a^2}{2}$$

$$D_{a,c} = (\tilde{T}_b - \tilde{T}_o)(s_c - s_a) + [(G_b - \mathbf{G}_q)w_b]_a^c$$

$$G = \Phi + c_l(T - \tilde{T}_o \log T)$$

Velocity PI for irreversible full moisture case

$$v^2 = \frac{c_k}{c_d} (1 + w)^{-1} [(T_b - T_o)(s^* - s) + (G_b - \mathbf{G}_q)(w^* - w)]$$

Effective outflow geopotential Φ_q (pseudoadiabatic)

Environment

Eyewall

Saturated
Eyewall

