Intense precipitation events associated with landfalling tropical cyclones in a warmer climate

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BACKGROUND & MOTIVATION

In the past a number of investigations have been done to explore the **effects of global warming on the precipitation associated to Tropical Cyclone (**e.g.,Gualdi et al. 2008, Knutson et al. 2004; Hasegawa and Emori, 2005). Despite the agreement in claiming that TC-related rainfall rates are likely to increase with greenhouse warming, the range of projection for the late twenty-first century between existing studies is very large (+3 to +37%, Knutson et al. 2010) and depending on the different used methodologies.

Rainfall is projected to increase over land, **both in terms of average and extremes** (Liu et al. 2009, Chou et al. 2009). Since there is a **large spatial variability** (Trenberth et al. 2011, Scoccimarro et al. 2013) associated to changes in projected rainfal amount , we would **verify and quantify the LANDFALLING TCs contribution region by region.**

The availability of a set of high resolution GCMs and simulations following common protocols (the HWG ones), should strengthen the projected information.





MOTIVATION

CMIP5 (20 models) difference between 99th and 90th percentiles (99p-90p) [mm/d] in 1997-2005



Scoccimarro, E., S. Gualdi, A. Bellucci, M. Zampieri, A. Navarra (2013): Heavy precipitation events in a warmer climate: results from CMIP5 models, Journal of Climate, DOI: 10.1175/JCLI-D-12-00850.1

MOTIVATION

CMIP5 (20 models) total precipitation, 90p and 99p-90p

CHANGES IN A WARMER CLIMATE

(% increase in RCP8.5 2061-2100 wrt 1966-2005) 99p-90p incr.

90p incr.

total precip.

JJA

75[°] N 75[°] N 75[°] N 60[°] N 60[°] 60[°] N N 45[°] N 45[°] N 45[°] N 30[°] N 30[°] N 30[°] N DJF 15[°] N 15[°] N 15[°] N ິ0 0° 0° 15[°] S 15[°] S 15[°] S 30[°] S 30[°] S 30[°] S 45[°] S 45[°] S 45[°] S 60[°] S 4120[°] W60[°] W 0[°] 60° S 180° W120° W60° W 0° 60° S 180° W120° W60° W 0° 60[°] E120[°] E180[°] E 60[°] E120[°] E180[°] E 60[°] E120[°] E180[°] E 75[°] N 75[°] N 75[°] N 60[°] N 60[°] N 60[°] N 45[°] N 45[°] N 45[°] N 30[°] N 30[°] N 30[°] N 15[°] N 15[°] N 15[°] N 0° 0° °0 15[°] S 15[°] S 15[°] S 30[°] S 30[°] S 30[°] S 45[°] S 45[°] S 45[°] S 60° S 180° W120° W60° W 0° 60[°] S 180[°] W120[°] W60[°] W 0[°] 60° S 180° W120° W60° W 0° 60[°] E120[°] E180[°] E 60[°] E120[°] E180[°] E 60[°] E120[°] E180[°] E [%] 20 -20 40 -60 60 -40 0







DATA and METHODS Informations about HWG models used

	GFDL	СМСС
	HIRAM	ECHAM5
Spatial Resolution	50 km	80 km
Vertical levels	32	31
Precipitation param.	Bretherton et. al 2004, Rotstayn 1997, Rotstayn 2000.	Tiedtke 1989, modified following Nordeng, 1994.
Reference	Zhao, M., I.M. Held, S-J. Lin, and G.A. Vecchi, 2009: Simulations of Global Hurricane Climatology, Interannual Variability, and Response to Global Warming Using a 50km Resolution GCM. J. Climate, 33, 6653-6678.	Roeckner, E., and Coauthors, 2003: The atmospheric general circulation model ECHAM5. Part I: Model description. MPI Rep. 349, 127 pp.

DATA and METHODS TC DETECTING algorithm used

	GFDL	СМСС
DETECTING	 Relative vorticity at 850 hPa larger than 1.6 x 10⁻⁵ s⁻¹ is found within areas of 6°x6° latitude and longitude. The local minimum of sea level pressure (within a distance of 2° latitude or longitude from the vorticity maximum) is defined as the center of the storm. The local maximum temperature averaged between 300 and 500 hPa is defined as the center of the warm core. The distance of the warm-core center from the storm center must not exceed 2°. The warm-core temperature must be at least 1°K warmer than the surrounding local mean. 	 1) Relative vorticity at 850 hPa is larger than 1 x 10⁻⁵ s⁻¹ 2) There is a relative surface pressure minimum, and the surface pressure anomaly, compared to a surrounding area with a radius of 350 km, is larger than 2 hPa 3) In a region with a radius of 350 km around the grid point considered, there is a grid point where the maximum surface wind velocity is larger than 15.5 ms^{-1.} 4) Wind velocity at 850 hPa is larger than wind velocity at 300 hPa 5) The sum of temperature anomalies at 700, 500, and 300 hPa is larger than 1° K, where anomalies are defined as the deviation from a spatial mean com- puted over a region with a radius of 350 km.

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DATA and METHODS TC TRACKING algorithm used

	GFDL	СМСС
TRACKING	 4) For each detected snapshot, a check is performed to see if there are storms during the following 6-h time period within a distance of 400 km. 5) If there are none, the trajectory is considered to have stopped. If there are some, the closest storm is chosen as belonging to the same trajectory as the initial storm. If there is more than one possibility, preference is given to storms that are to the west and poleward of the current location. 6) To qualify as the model storm trajectory, a trajectory must last at least 3 days, and have a maximum surface wind speed greater than 17 ms⁻¹ during at least 3 days (not necessarily consecutive) 	 6)) For each detected snapshot, a check is performed to see if there are storms during the following 6-h time period within a distance of 300 km. 7)) If there are none, the trajectory is considered to have stopped. If there are some, the closest storm is chosen as belonging to the same trajectory as the initial storm. If there is more than one possibility, preference is given to the nearest storm. 8) To qualify as the model storm trajectory, a trajectory must last at least 1 day (corresponding to four time steps of the model output)



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Fraction of precipitation associated with TCs in the CONTROL (10y) experiment [% wrt the total]





RESULTS GFDL model TC genesis point in HWG experiments



Changes in the total amout of water associated to TCs region by region in the HWG experiments

Changes in the total amout of water associated to TCs, NORMALIZED BY TCs DAYS

CMCC TCprecip/TCdays increase [% wrt CONTROL]

Changes in the total amout of water associated to TCs OVER LAND in the HWG experiments

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Changes in the total amout of water associated to TCs OVER LAND, NORMALIZED BY TCs DAYS

CONCLUSIONS

- 1. Global patterns of **precipitation associated with TC activity** are reasonably well represented by the considered GCMs. GFDL model (50 Km) performs better than CMCC model (80Km).
- 2. In terms of amount of water associated to TCs at global scale, in 2xCO2, p2K and p2K2xCO2 experiments, the CMCC model shows more pronounced changes wrt GFDL model (resolution?, different precipitation parameterization?).
- 3. Changes in the amount of water associated to LANDFALLING TCs are more pronounced if compared to changes in the corresponding basin. The obtained results are not coherent between models.

NEXT STEP:

- Improve the statistics: period 10y ->20y
- - include all available HWG GCMs precipitation (daily) res.
- Verify the relationship between precipitation results and the general circulation associated to each model/experiment (CAPE, low level vort., winds, etc.).

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thank you!

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Precipitation associated with TCs in the CONTROL (10y) experiment [mm/y]

Tropical Cyclones as simulated by CMCC-INGV fully coupled climate models

COUPLED MODELS at CMCC

Tropical Cyclones as simulated by CMCC-INGV fully coupled climate models: INGV-SXG (atm: echam4, T106) and CMCC_MED (atm: echam5, T159)

COUPLED MODELS at CMCC

-Changes in Tropical Cyclone Activity due to Global Warming: Results from a High-Resolution CGCM. Gualdi et al. 2008 - J. of Climate, Vol. 21, pp. 5204-5228

-Effects of Tropical Cyclones on Ocean Heat Transport in a High Resolution CGCM. Scoccimarro et al. 2011 - J. of Climate, in press. Doi: 10.1175/2011JCLI4104.1

Tropical Cyclones as simulated by CMCC-INGV fully coupled climate models: COUPLED INGV-SXG (T106) and CMCC-CM (T159) MODELS at CMCC

Annual TCs number and variability

Gualdi et al. 2008 - J. of Climate, Vol. 21, pp. 5204-5228

-Effects of Tropical Cyclones on Ocean Heat Transport in a High Resolution CGCM. Scoccimarro et al. 2011 – *J of Climate*, in press. Now moving from MPI ECHAM based atmospheric model toCOUPLEDNCR/UCAR-Community Earth System Model (CESM):MODELS at CMCCCMCC developed the coupling with the Nucleus for European Modelling of the OceanNEMO) GCM.

