

Submesoscale Horizontal Advection Drives High-Frequency Upper-Ocean Temperature Variability

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01 Motivation: The Energy Gap

Climate models overall capture **low-frequency** upper-ocean temperature variability across resolutions, but systematically underestimate it at **sub-seasonal periods** ($< O(100)$ days).

Q1. What drives higher-frequency upper-ocean temperature variability?

Q2. What processes may be underrepresented in coarse models?

02 Budget & Cross Spectral Density

$$\frac{\partial \langle T \rangle}{\partial t} = A_h + A_v + Q + R$$

$\langle T \rangle$ — layer averaged temperature

A_h — horizontal advection

A_v — vertical advection across the layer base

Q — surface flux retained within the layer

R — sub-grid mixing, parametrized processes

Terms are reconstructed offline from **hourly GEOS/ECCO C1440-LLC2160 (1/24°) output**.

We compute 3-D (t, x, y) **cross-spectral-densities (CSDs)** between T and each budget term over 1100 km sliding tiles and four independent 108-day windows.

$$\text{Var}(\langle T \rangle) = \text{CSD}(A_h, \langle T \rangle) + \text{CSD}(A_v, \langle T \rangle) + \text{CSD}(Q, \langle T \rangle) + R$$

SIGN CONVENTION

$\text{CSD}(\langle T \rangle, \text{term}) > 0 \rightarrow$ the term is a **source** of temperature variance at that scale.

$\text{CSD}(\langle T \rangle, \text{term}) < 0 \rightarrow$ the term is a **sink** of temperature variance at that scale.

03 Region of Interest: Kuroshio Extension & Northeastern Pacific

Region	domain	layer used for full budget	deepest MLD in the simulation
Kuroshio Extension	31–42°N, 147–163°E	0–400m	379 m
Northeastern Pacific	23–36°N, 152–138°W	0–200m	156 m

REFERENCES

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- [2] Martin, P. E. et al. (2021). Drivers of Atmospheric and Oceanic Surface Temperature Variance: A Frequency Domain Approach. *Journal of Climate*, 34(10), 3975–3990.
- [3] Menemenlis, D. et al (2026). The GEOS/ECCO C1440-LLC2160 Coupled Atmosphere-Ocean Simulation Dataset. *Scientific Data*.

04 Horizontal advection: The Dominant Source at (0.2-50) day and O(10-100) km

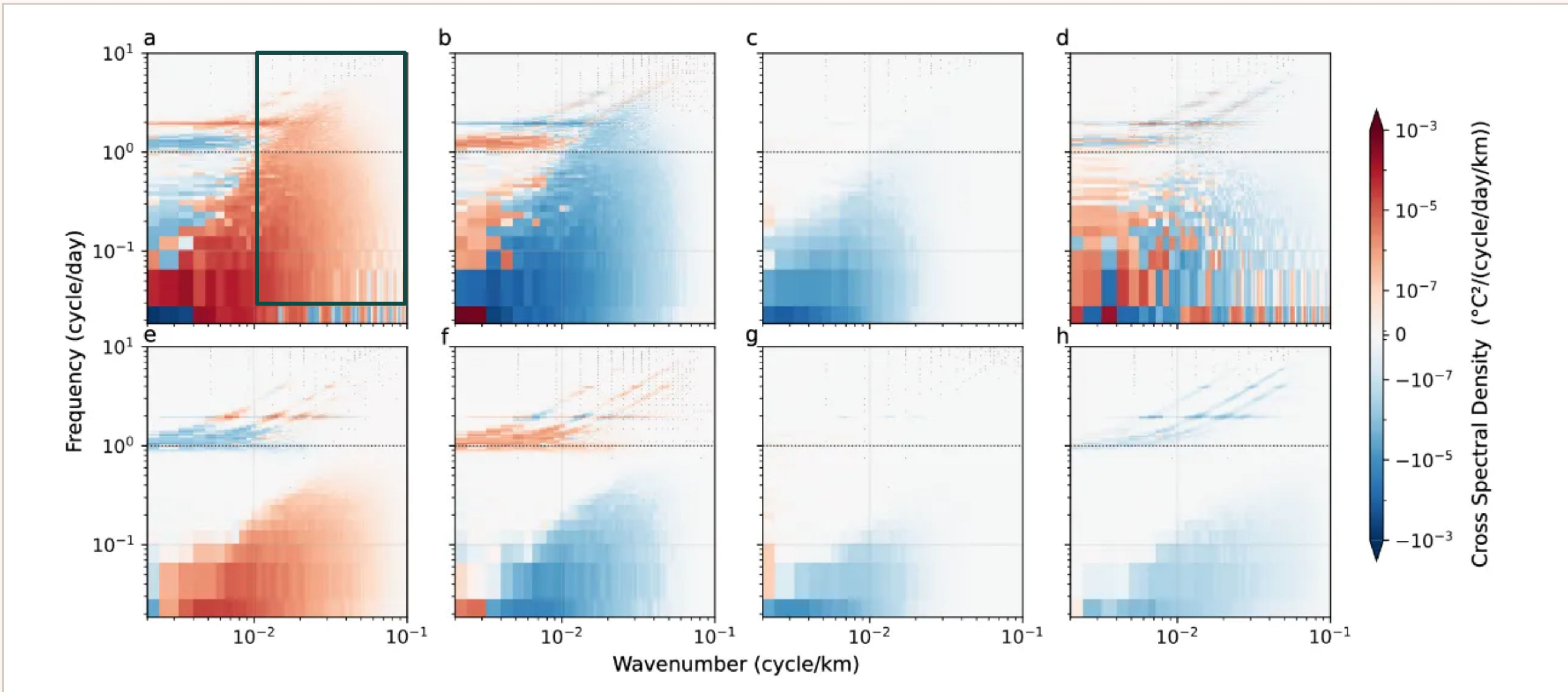


Fig. 1. Spectral contribution of each budget term to upper-ocean temperature variability across 8-500 km spatial scale and 0.2-54 day spatial scales, quantified by the CSD between each term and temperature.(a-d) CSD for horizontal advection (A_h), (b) vertical advection (A_v), (c) surface heat flux (Q), and (d) residual (R) in the Kuroshio Extension (KE). (e-h) Same as (a-d), but for the Northeastern Pacific (NEP). Positive CSD (red) indicates a source of temperature variance; negative CSD (blue) indicates a sink.

05 Horizontal advection: Important Source in Small Scales

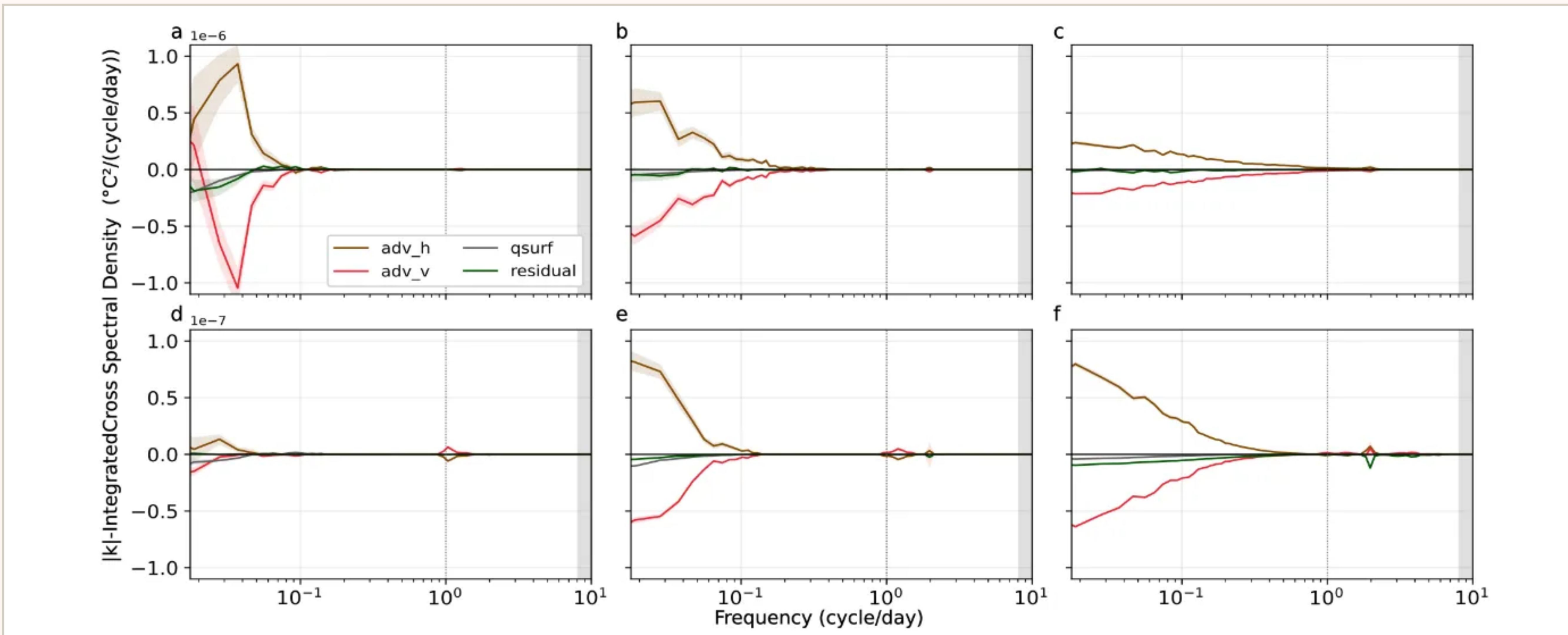


Fig. 2. Contribution of each budget term to upper-ocean temperature variability as a function of frequency for three spatial-scale bands. CSD is integrated over wavelengths of (a) > 200 km, (b) $100\text{--}200$ km, and (c) $10\text{--}100$ km in the KE. (d–f) Same as (a–c), but for the NEP. Positive CSD indicates a source of temperature variance; negative indicates a sink.

06 Horizontal Advection at O(10-100) km :

Dominant Driver of High-Frequency Upper-ocean Temperature Variability

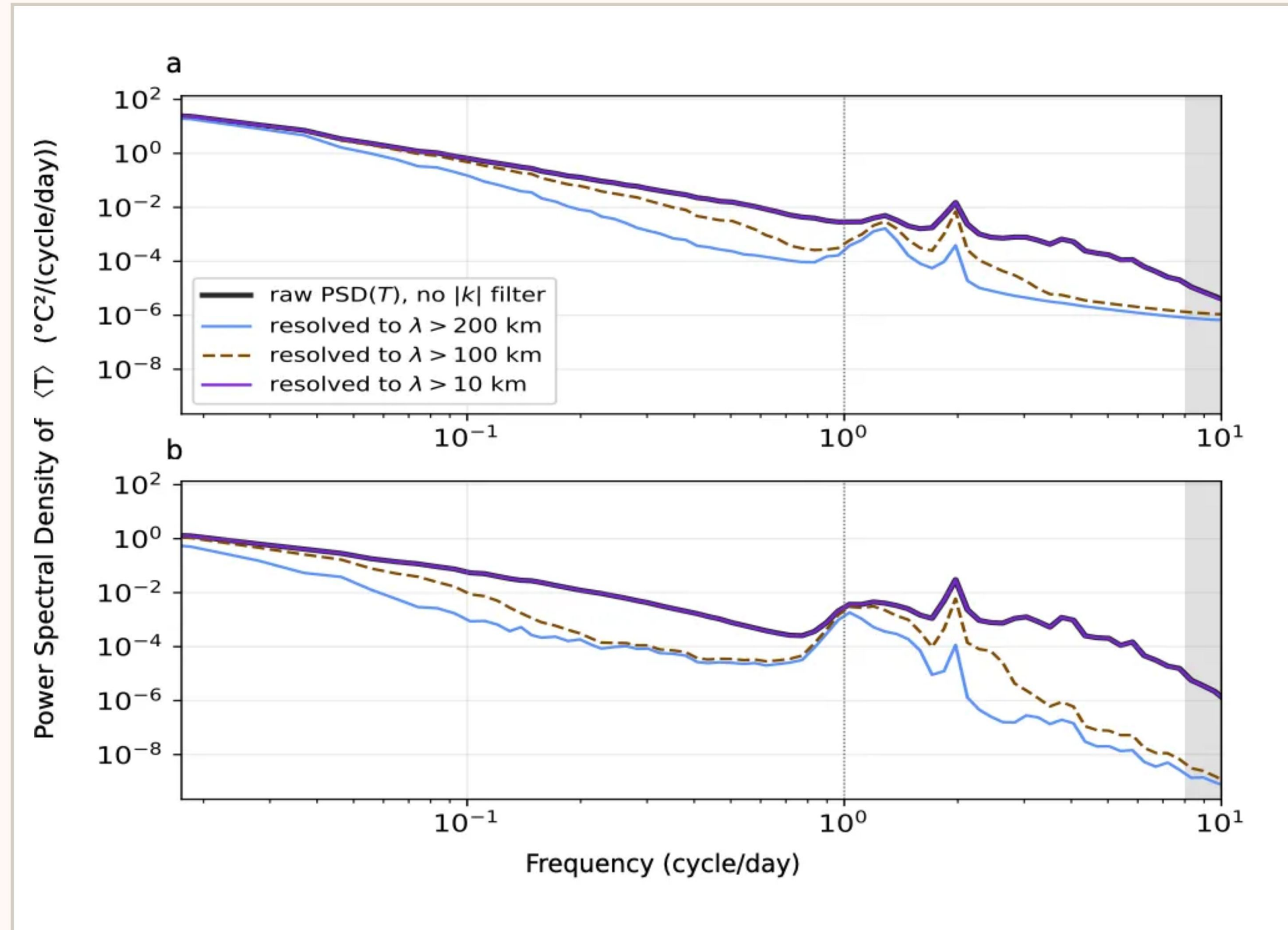


Fig. 3. Upper-ocean temperature variability as a function of frequency for progressively finer minimum resolved spatial scales. Temperature variability is calculated using budget terms resolved at wavelengths of (blue) > 200 km, (dashed brown) > 100 km, (purple) > 10 km, and (black) > 8 km, illustrating the increase in variability as smaller spatial scales are resolved (a) in the KE. (b) Same as (a), but for the NEP.

07 Key Findings & Implications

- 1 We first quantify how upper-ocean temperature budget terms contribute to temperature variance across spatial and temporal scales in a fully coupled, 1/24° simulation.
- 2 Horizontal advection is the only budget term that acts as a source of temperature variance across 10-100 km and 0.2-54 days—scales not resolved by a $\sim 1^\circ$ ocean model.
- 3 Unresolved 10-100 km horizontal advection provides a potential mechanism for the sub-seasonal temperature variance deficit in coarse-resolution models.

Acknowledgements

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