## Arctic seasonal variability and extremes, and the role of weather systems in a changing climate

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## BACKGROUND

The Arctic climate is changing dramatically. Long-term trends in surface temperature and sea ice extent caused by anthropogenic warming strongly affect the Arctic atmosphere [1,2]. Air-sea-ice feedback processes additionally amplify these trends [3,4].

The Arctic region further exhibits a strong variability on time scales from days to seasons. Synoptic weather systems such as cyclones and anticyclones are related to air mass exchanges with the mid-latitudes [5,6] and air mass transformations within the Arctic [7,8], and are thus key drivers of Arctic variability.

In this study it is our goal to analyze the change of Arctic seasonal variability as well as the role of weather systems in a warming climate by using large-ensemble climate model data.

## Climate change in the *T-P* phase space

### Seasonal mean

- Substantial regional and seasonal differences in magnitude of overall warming and wettening
- Particularly large T increase in seasons and regions with strong sea ice loss
- Extremely warm seasons in S2000 rank among the coldest seasons / become unrealistic in S2100

## **Inter-annual variability**

- Increase in *P* variability over Arctic Ocean and reduction in T variability in the Kara and Barents Seas in winter
- Changes related to increase in seasonal cycle and interannual variability of sea-ice concentration as well as increasing moisture transport towards high latitudes [11]

# 1990-2000 (S2000) 2031-2040 (\$2040) 2061-2070 (\$2070) 2091-2100 (S2100)

## **Application of large-ensemble data: changes in synoptic processes**

## Role of weather systems for extreme seasons (CESM1) Arctic moisture transport (CESM2)

- extremes driven by colocated anomalies in • **P** cyclone/anticyclone occurrence, while **T** extremes driven by large-scale dipole patterns
- Large-scale patterns largely unaffected by global warming



Figure 3: Differences in seasonalmean cyclone frequency anomaly composites for DJF in the ice-free part of the Kara and Barents Seas (black-white contour) for extremely (a, c) wet  $(P_{+}) - dry (P_{-})$ and (b, d) warm  $(T_+)$  – cold  $(T_-)$ seasons in (a, b) S2000 and (c, d) S2100.



Master thesis project by Sven Voigt, ETH Zurich (supervised together with Iris Thurnherr)



**Figure 4**: Mean moisture source areas for the Kara and Barents Seas (red contour) for DJF in (left) present-day and (right) future CESM2 simulations. Ensemble mean sea ice edge in each climate state is shown by solid green line.

- Increasing contribution of local moisture sources linked to sea ice retreat
- Increasing contribution of remote moisture sources (e.g., subtropical North Atlantic, Mediterranean) which can be associated with changes in cyclone and anticyclone frequency

## **METHOD**

For this analysis we use large-ensemble data of the CESM1 climate model [9,10]. Simulations are performed using the RCP8.5 emission scenario. 6-hourly output is available for four different periods:

**S2000**: 1990-2000 [**105** ensemble members]

**S2040**: 2031-2040 [**40** ensemble members]

**S2070**: 2061-2070 [**40** ensemble members]

To account for varying climatological and surface conditions, we differentiate between distinct geographic Arctic sub-regions and surface types (open ocean and sea ice). The seasonal variability in these sub-regions is analysed in the phase space spanned by the seasonal-mean anomalies of 2m-temperature (T) and precipitation (P).

- Seasons ranking among the warmest in present-day climate are soon projected to rank among the coldest or **become unrealistic**.
- Robust dynamical relationship between weather system frequency and seasonal extremes **persists in warming climate**.
- Compared to a range of CMIP6 models, **CESM1 performs well** regarding trends in seasonal-mean T and P, and variability thereof.

## Acknowledgements

The authors acknowledge Gary Strand and Clara Deser (both NCAR) for providing the CESM restart files and are thankful to Urs Beyerle (ETH Zurich) for performing the CESM simulations. This research has been supported by the H2020 European Research Council (grant no. 787652).

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## **S2100**: 2091-2100 [**105** ensemble members]

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