

# Disentangling the causes of observed changes in biogeochemical variables along ventilation pathways



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## Results: Pacific SAMW

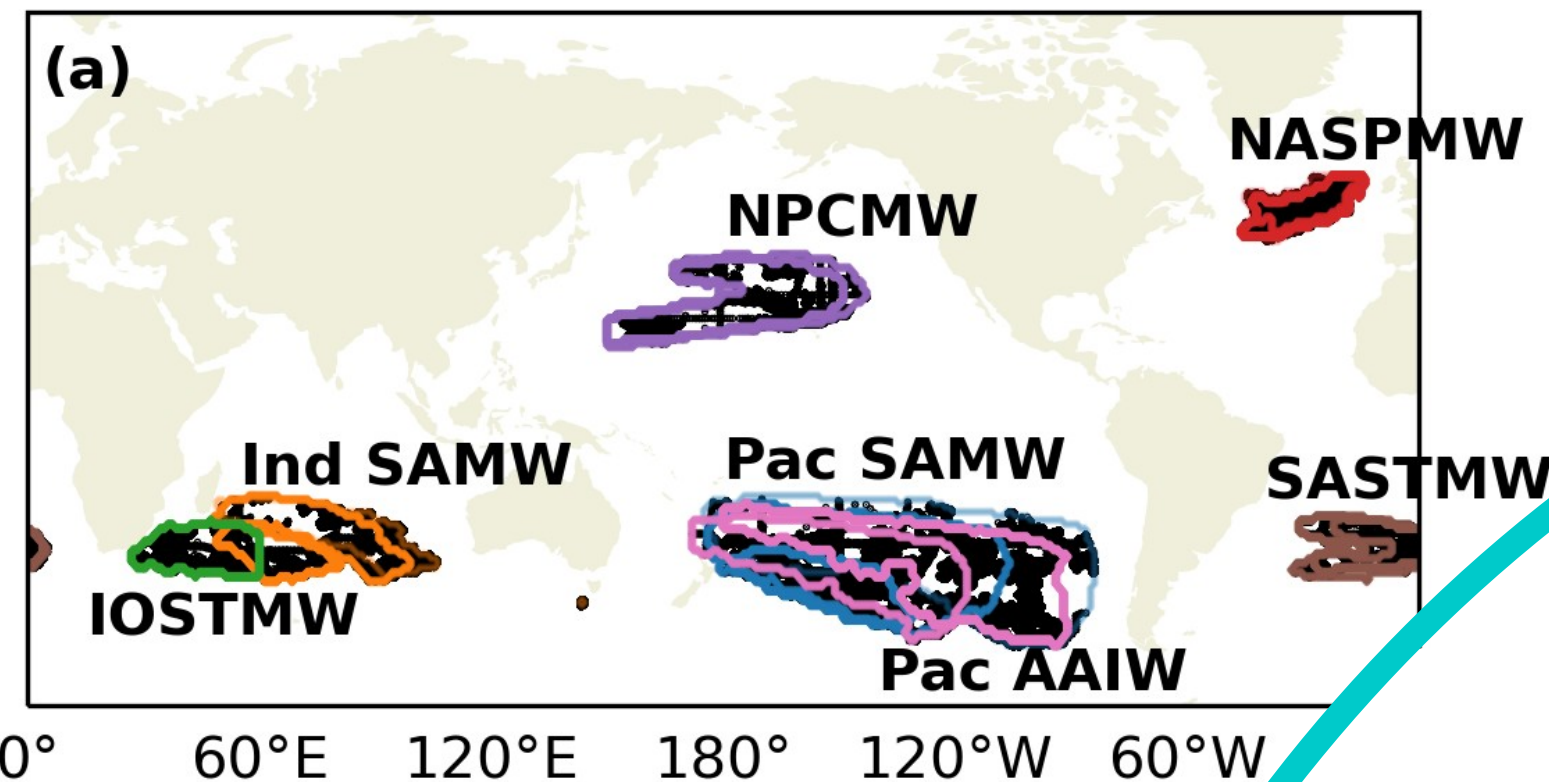
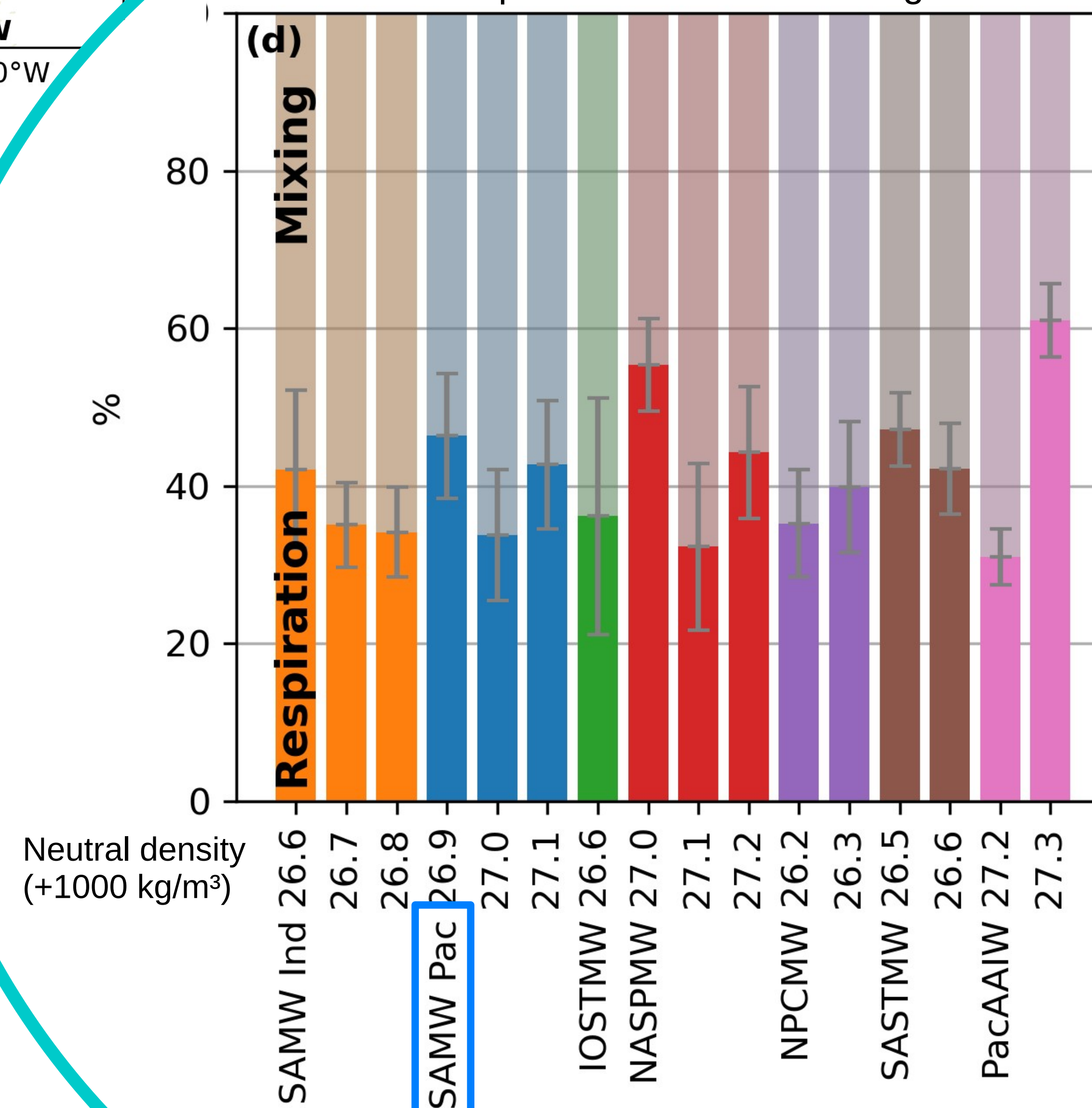
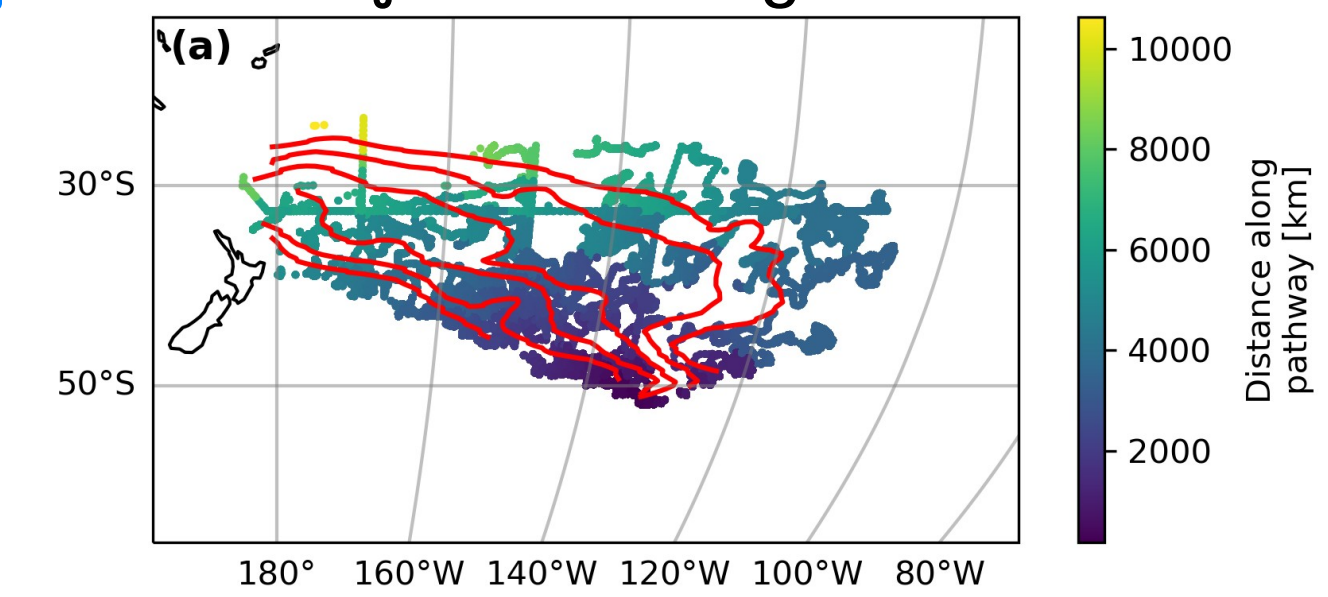


Fig. 1: % of  $\Delta O_2$  and  $\Delta NO_3$  due to either local respiration or vertical mixing



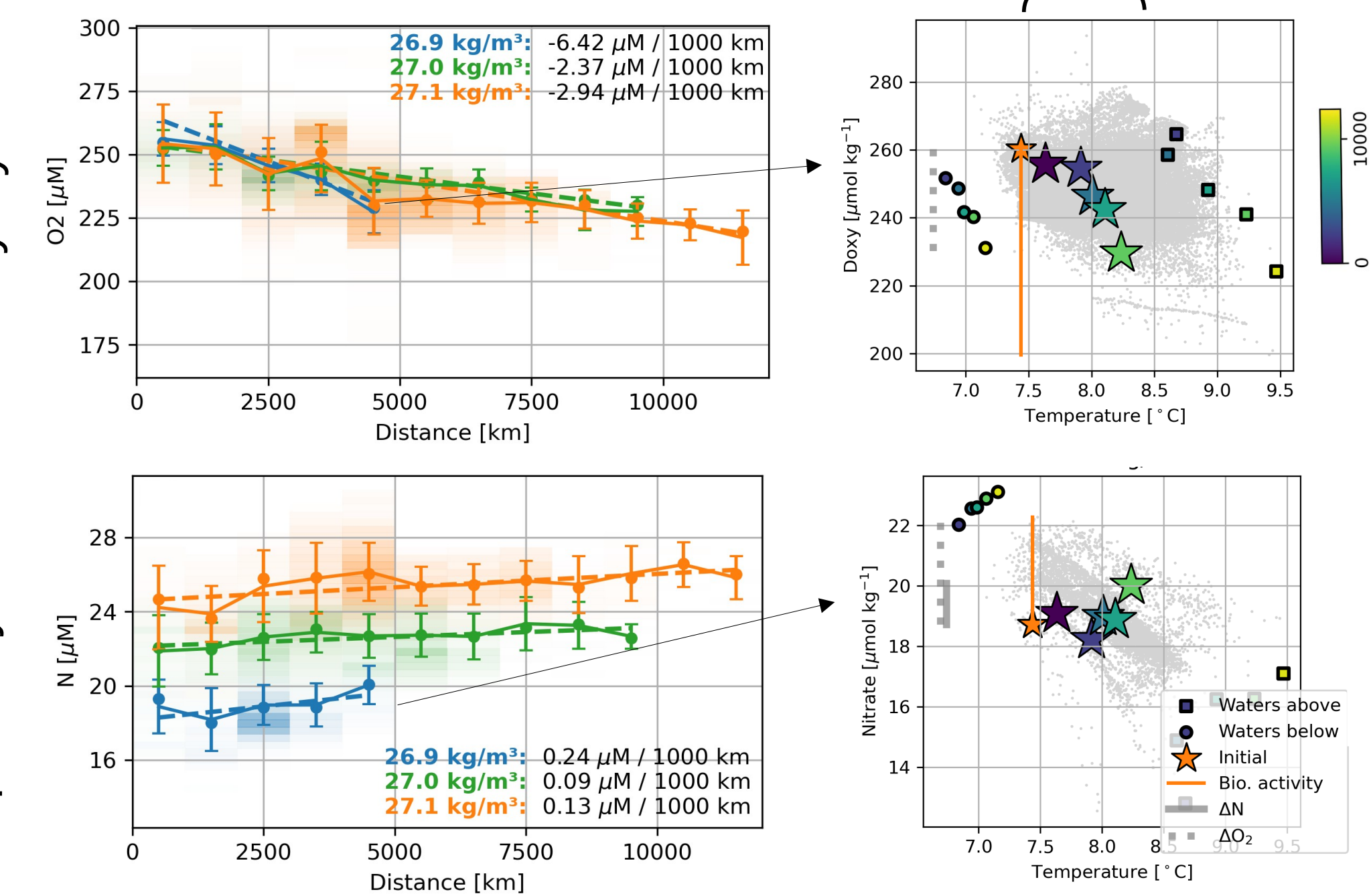
Mixing contributes to more than half of the changes in BGC variables

Fig. 3: Pathway for SAMW<sub>Pac</sub>  
 $\gamma_n = 1027.0 \text{ kg/m}^3$



Notice the change in temperature, which can only be explained by mixing

Fig. 4: Evolution along the pathway, for three density layers



## Conclusions

- Along ventilation pathways, mode waters [O<sub>2</sub>] decrease at a rate between 2 and 20 μmol/kg / 1000 km, and [NO<sub>3</sub>] at a rate between 0.2 and 2.6 μmol/kg / 1000km (Fig. 4).
- In mode waters, mixing contributes to more than half of the changes in BGC variables (Fig. 1)
- This proportion is relatively uniform across mode waters (Fig. 1)

## Future work

- How does the local respiration compare to the total respiration (including that in the waters that mix in)
- What are the implications for AOU and difference to pre-formed nitrate, and hence for calculations of respiration and remineralization rates, carbon export, and C<sub>ant</sub>.

## Background

- Mode waters transport oxygen and carbon into the ocean's interior
- BGC variables are affected by bio activity and mixing
- Mixing is not considered in the commonly used AOU and difference to pre-formed nitrate [1,2]

Goal: Quantify the role of mixing in setting BGC properties along ventilation pathways

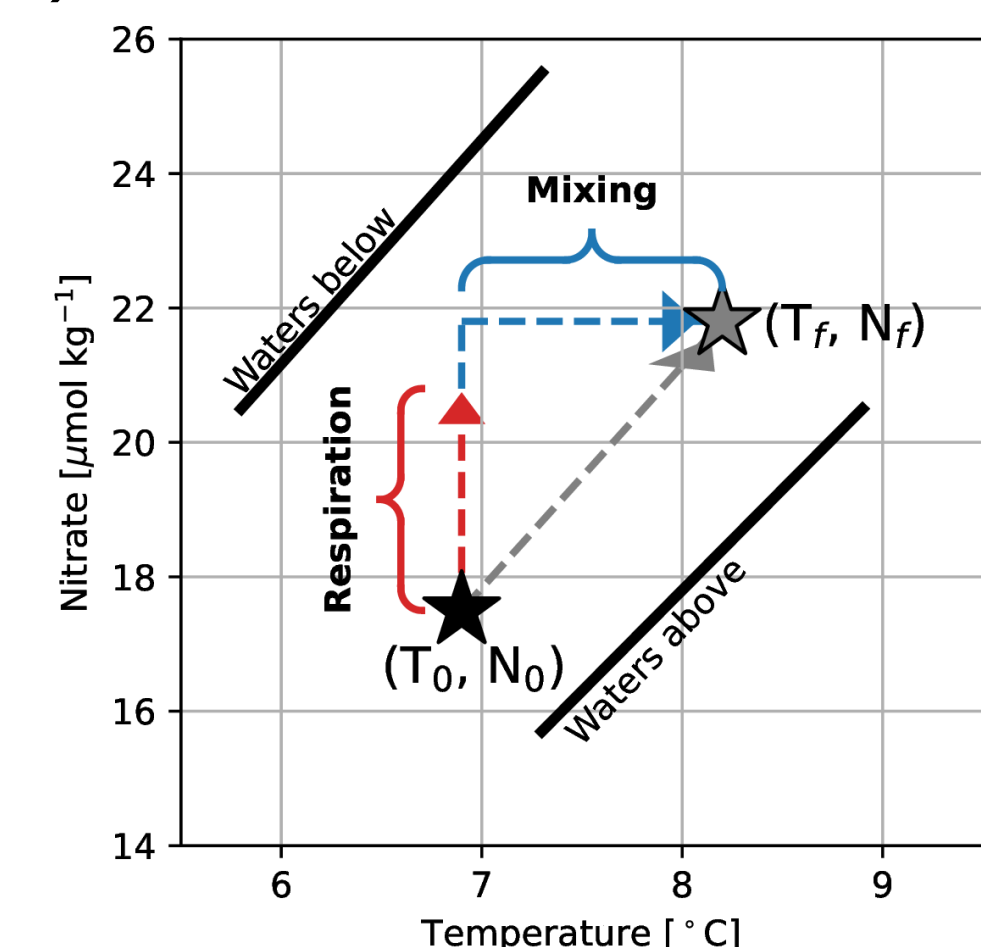
## Method

- 1) Identify the ventilation pathways
  - Core Argo-derived geostrophic streamlines [3]
- 2) Evaluate the evolution of O<sub>2</sub> and NO<sub>3</sub><sup>-</sup>
  - Average the Argo-BGC data along the pathways
- 3) Solve mixing equations, for distance bins along the pathways

$$\begin{aligned}
 f_0 C_0 + f_{above} C_{above} + f_{below} C_{below} &= C_i \\
 f_0 O_{2,0} + f_{above} O_{2,above} + f_{below} O_{2,below} + R &= O_{2,i} \\
 f_0 N_0 + f_{above} N_{above} + f_{below} N_{below} - r_{N:O_2} R &= N_i \\
 f_0 + f_{above} + f_{below} &= 1
 \end{aligned}$$

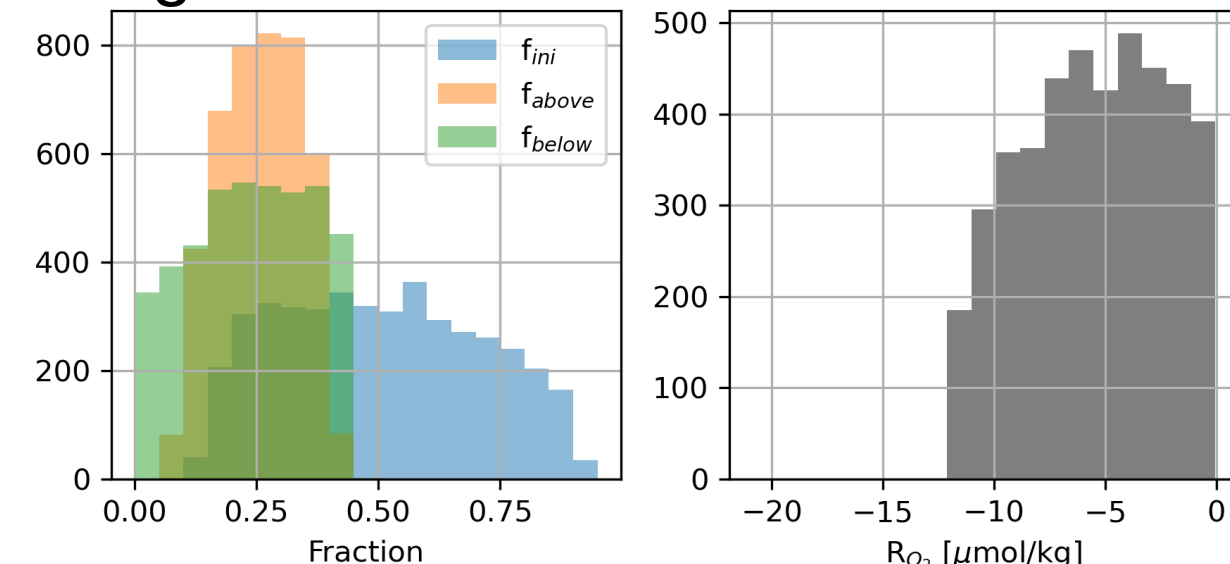
- Monte Carlo approach
- Outputs: fractions (*f*) and respiration (*R*)

- 4) Reconstruct the role of respiration and mixing



$$\begin{aligned}
 \Delta O_2^{respiration} &= R \\
 \Delta O_2^{mixing} &= f_{above} O_{2,above} + f_{below} O_{2,below} - (f_{above} + f_{below}) O_{2,i}
 \end{aligned}$$

Fig. 2: Solution for one distance bin



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## References

[1] Guo, H., Kriest, I., Oschlies, A., and Koeve, W. (2023). Can Oxygen Utilization Rate Be Used to Track the Long-Term Changes of Aerobic Respiration in the Mesopelagic Atlantic Ocean? *Geophysical Research Letters*, 50(13):e2022GL102645.

[2] Ono, T., Midorikawa, T., Watanabe, Y. W., Tadokoro, K., and Saino, T. (2001). Temporal increases of phosphate and apparent oxygen utilization in the subsurface waters of western subarctic Pacific from 1968 to 1998. *Geophysical Research Letters*, 28(17):3285–3288.

[3] McDougall, T. J., and A. Klocker. (2010): An approximate geostrophic streamfunction for use in density surfaces. *Ocean Modelling*, 32, 105– 117.