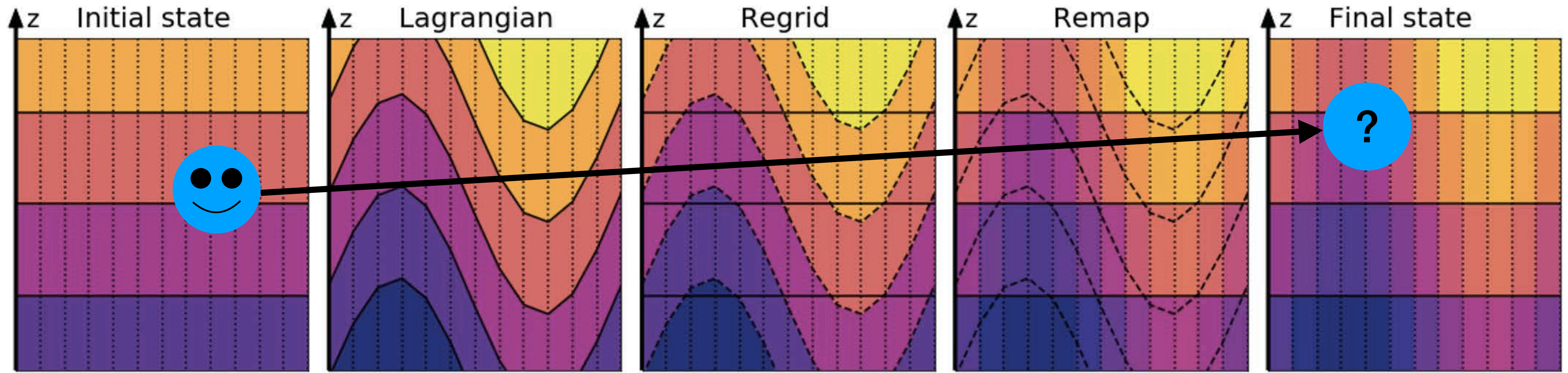


# Online Lagrangian Particle Advection in MOM6



**Spencer Jones (he/him)**<sup>1</sup>, Kaila Uyeda<sup>1,2</sup>, Luyu Sun<sup>3</sup>,  
Alistair Adcroft<sup>4</sup>, Matthew Harrison<sup>4</sup> and Stephen Griffies<sup>4</sup>

1. Texas A&M University, 2. UC, Irvine, 3. University of Maryland, 4. NOAA GFDL

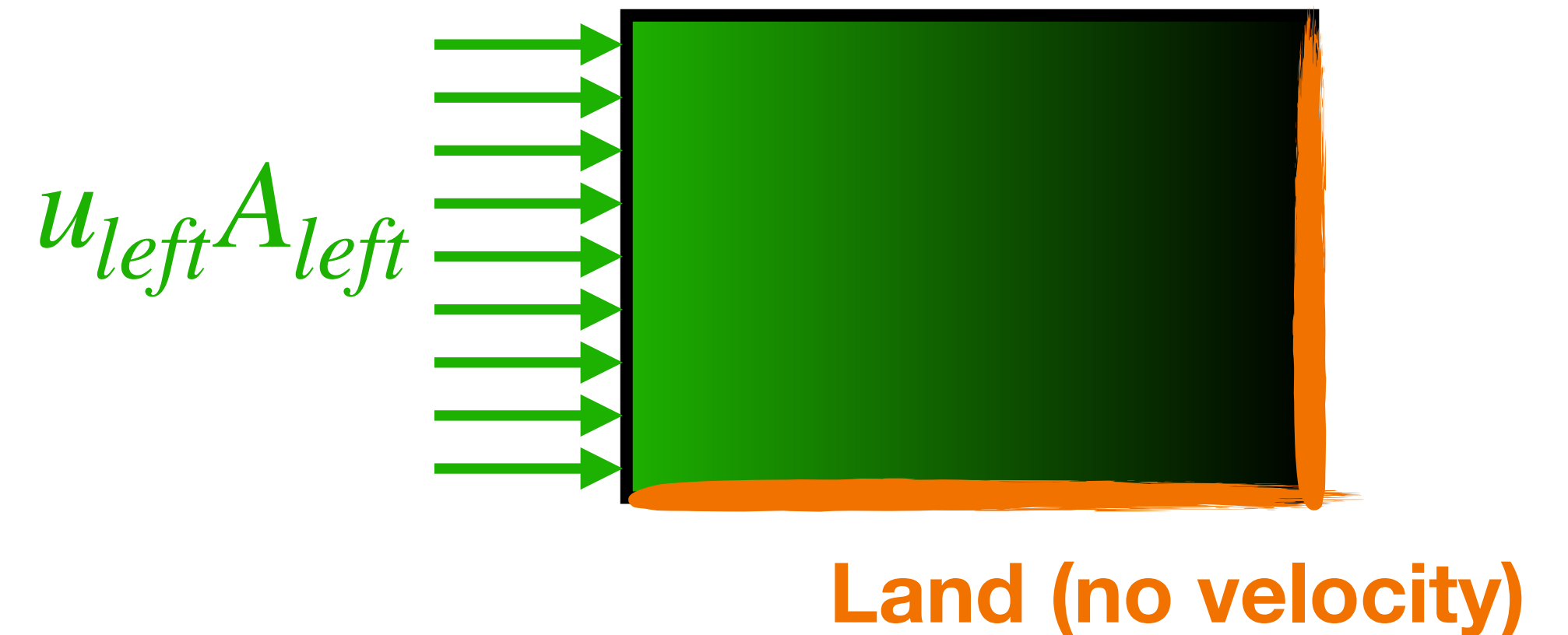


# Online particle advection has some advantages over offline advection

- Offline particle advection typically involves
  - saving time-averaged velocities
  - using them to advect virtual particles
- You can save more output to get more accurate trajectories, but this requires a lot of storage
- Online particle advection avoids this problem, because the particles are advected as part of the model run

# In a model where the grid is stationary (not true in MOM6)

$u$ ,  $v$ ,  $w$  are interpolated within each grid cell

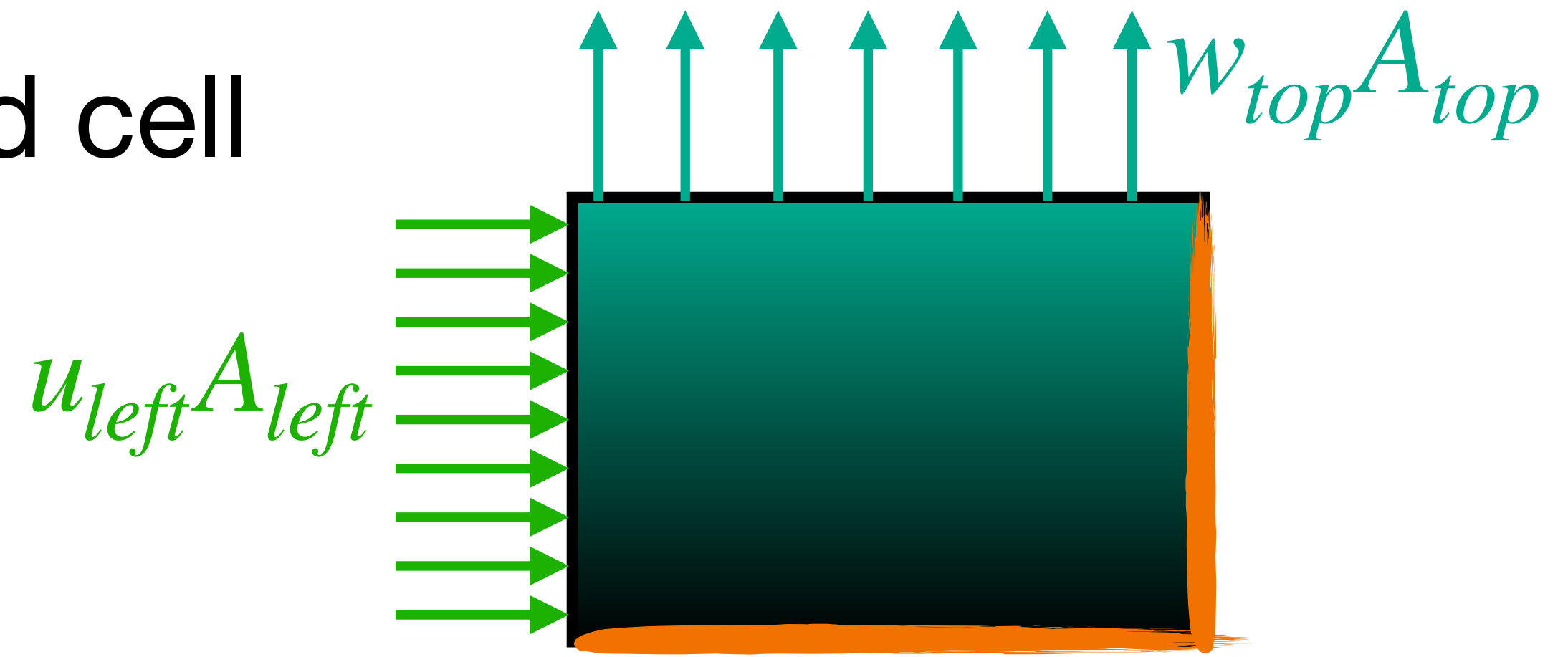


Most commonly,

- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y$ ,  $z$

# In a model where the grid is stationary (not true in MOM6)

$u$ ,  $v$ ,  $w$  are interpolated within each grid cell



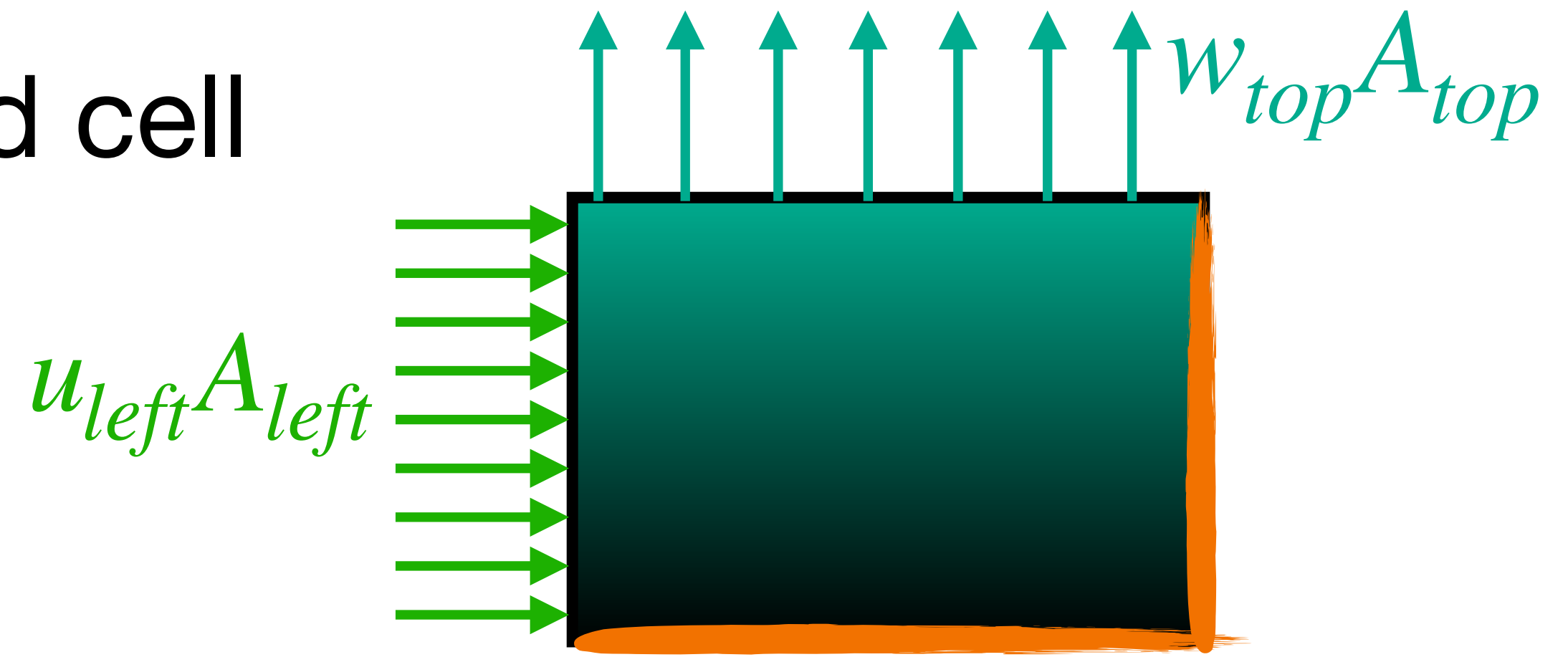
Most commonly,

- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y$ ,  $z$
- $v$  is interpolated linearly in the  $y$ -direction, and is constant in  $x$ ,  $z$
- $w$  is interpolated linearly in the  $z$ -direction, and is constant in  $x$ ,  $y$

Land (no velocity)

# In a model where the grid is stationary (not true in MOM6)

$u$ ,  $v$ ,  $w$  are interpolated within each grid cell



Most commonly,

- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y$ ,  $z$
- $v$  is interpolated linearly in the  $y$ -direction, and is constant in  $x$ ,  $z$
- $w$  is interpolated linearly in the  $z$ -direction, and is constant in  $x$ ,  $y$

This is a “mass conserving” scheme.

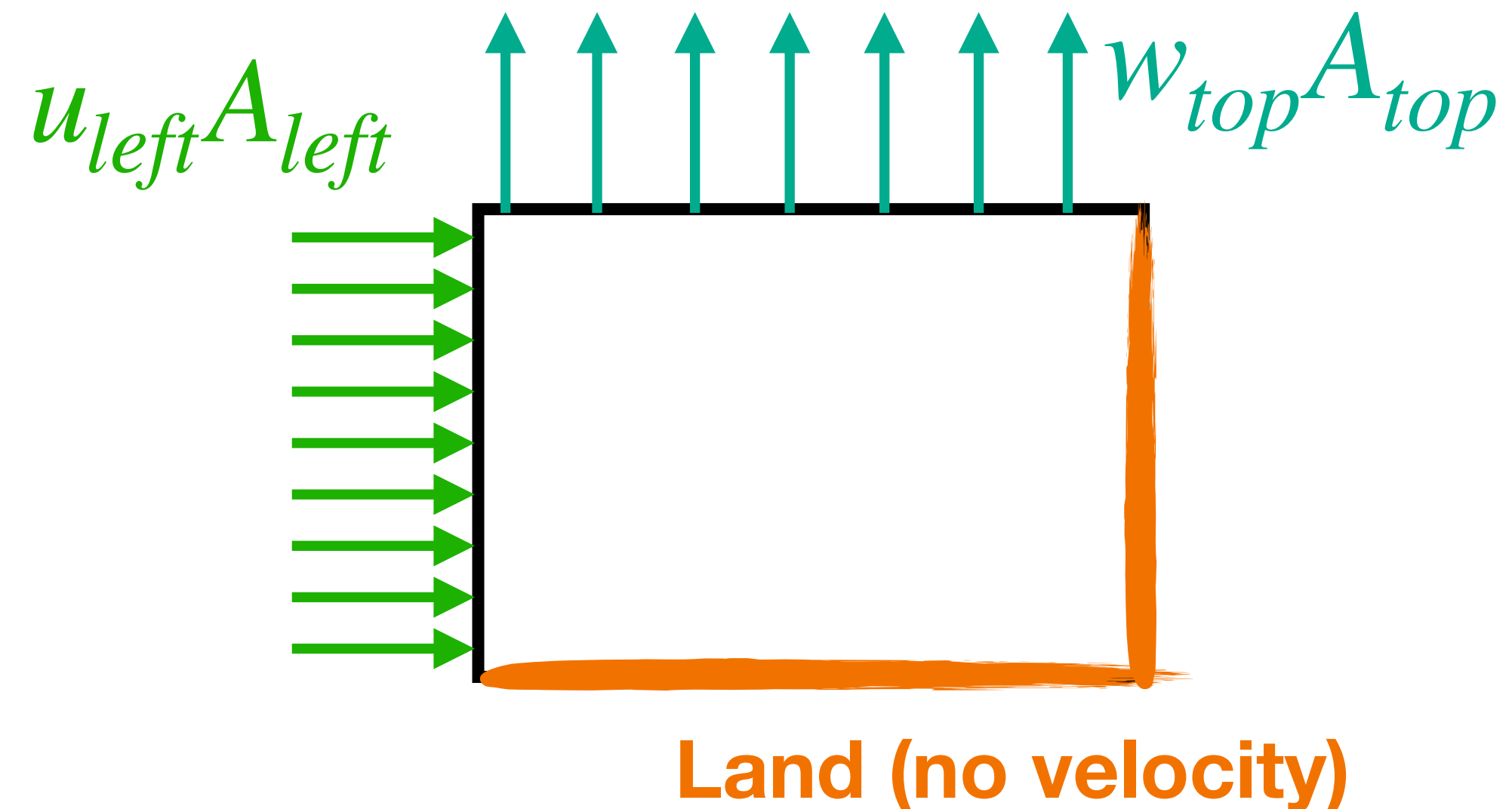
Land (no velocity)

# Defining a “mass conserving” scheme.

If

- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y, z$
- $v$  is interpolated linearly in the  $y$ -direction, and is constant in  $x, z$
- $w$  is interpolated linearly in the  $z$ -direction, and is constant in  $x, y$

**In the ocean model, inputs and outputs match for each grid cell**

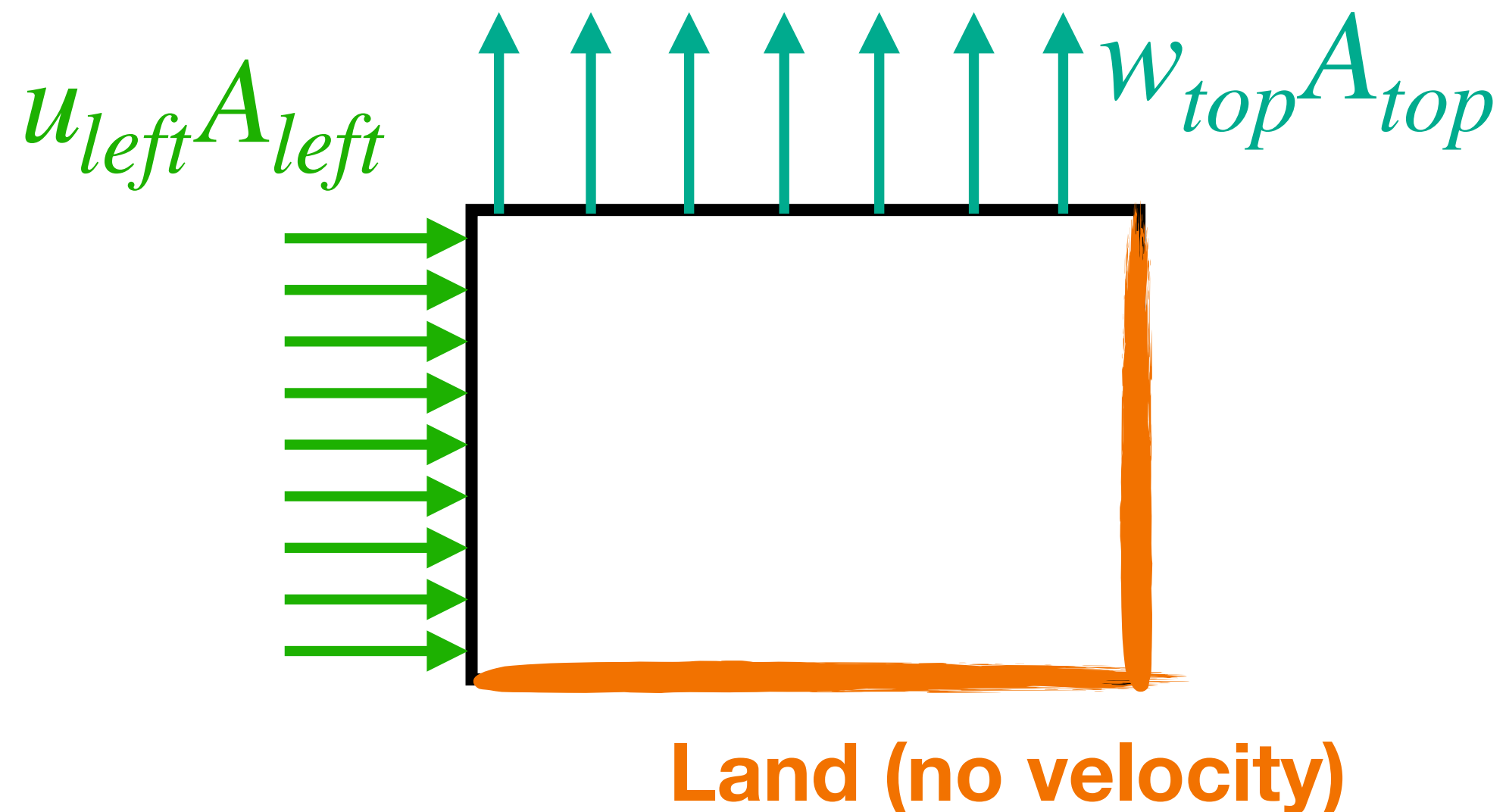


# Defining a “mass conserving” scheme.

If

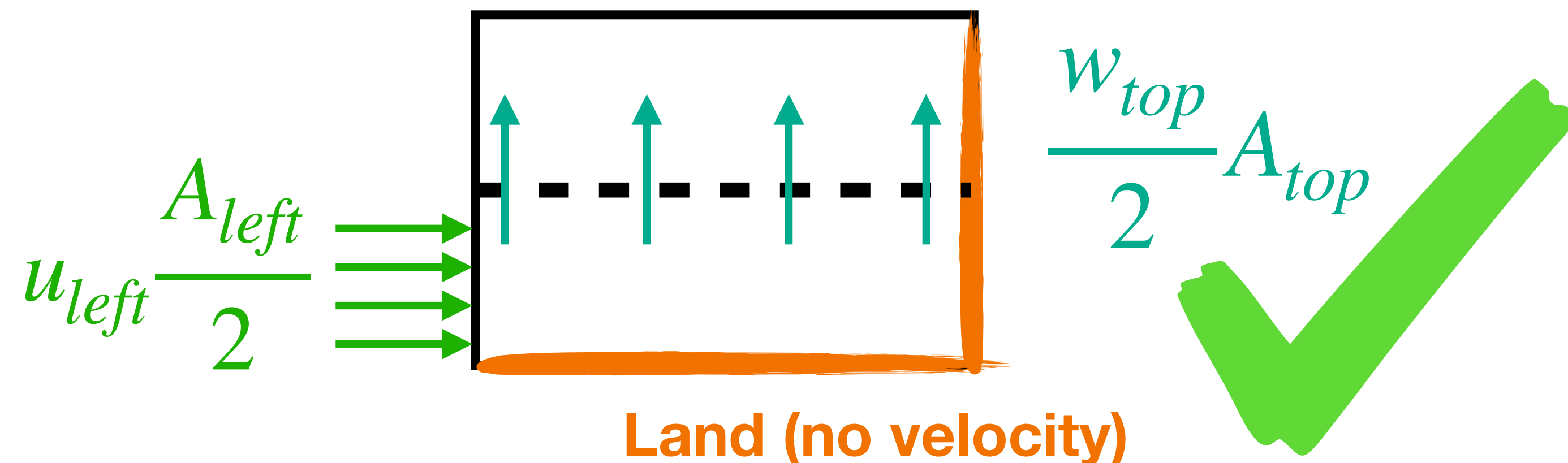
- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y, z$
- $v$  is interpolated linearly in the  $y$ -direction, and is constant in  $x, z$
- $w$  is interpolated linearly in the  $z$ -direction, and is constant in  $x, y$

In the ocean model, inputs and outputs match for each grid cell



Subdividing the grid cell does not create convergence of mass

e.g.



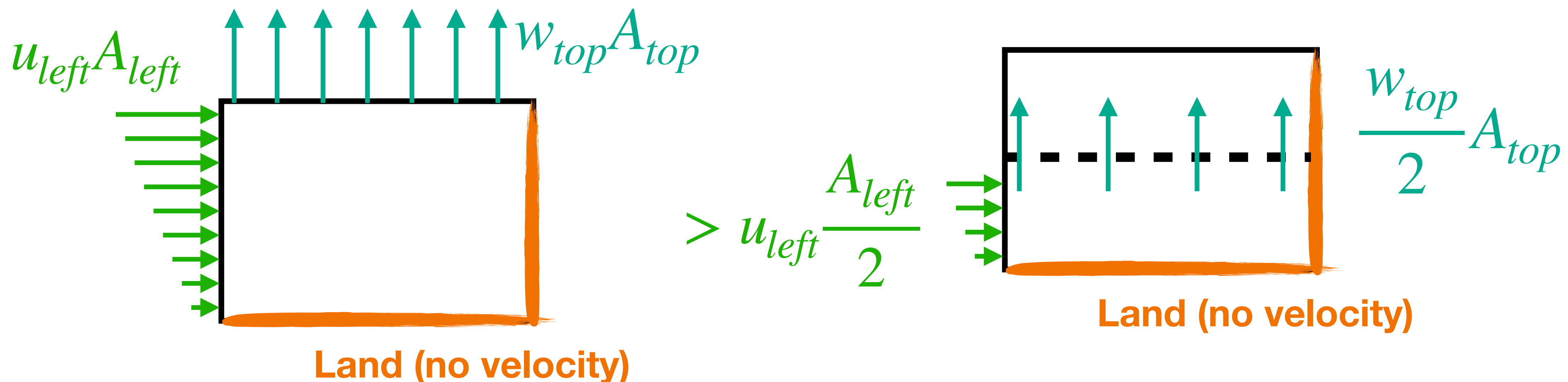
# Example of a scheme that doesn't conserve mass

- $u$  is interpolated linearly in the  $x$ -direction, is constant in  $y$  and varies linearly in  $z$
- $w$  is interpolated linearly in the  $z$ -direction, and is constant in  $x, y$

(Perhaps with the motivation of having  $u$  and  $v$  decrease towards the ocean floor)

**In the ocean model, inputs and outputs match for each grid cell**

**Convergences appear when we subdivide our grid cells**



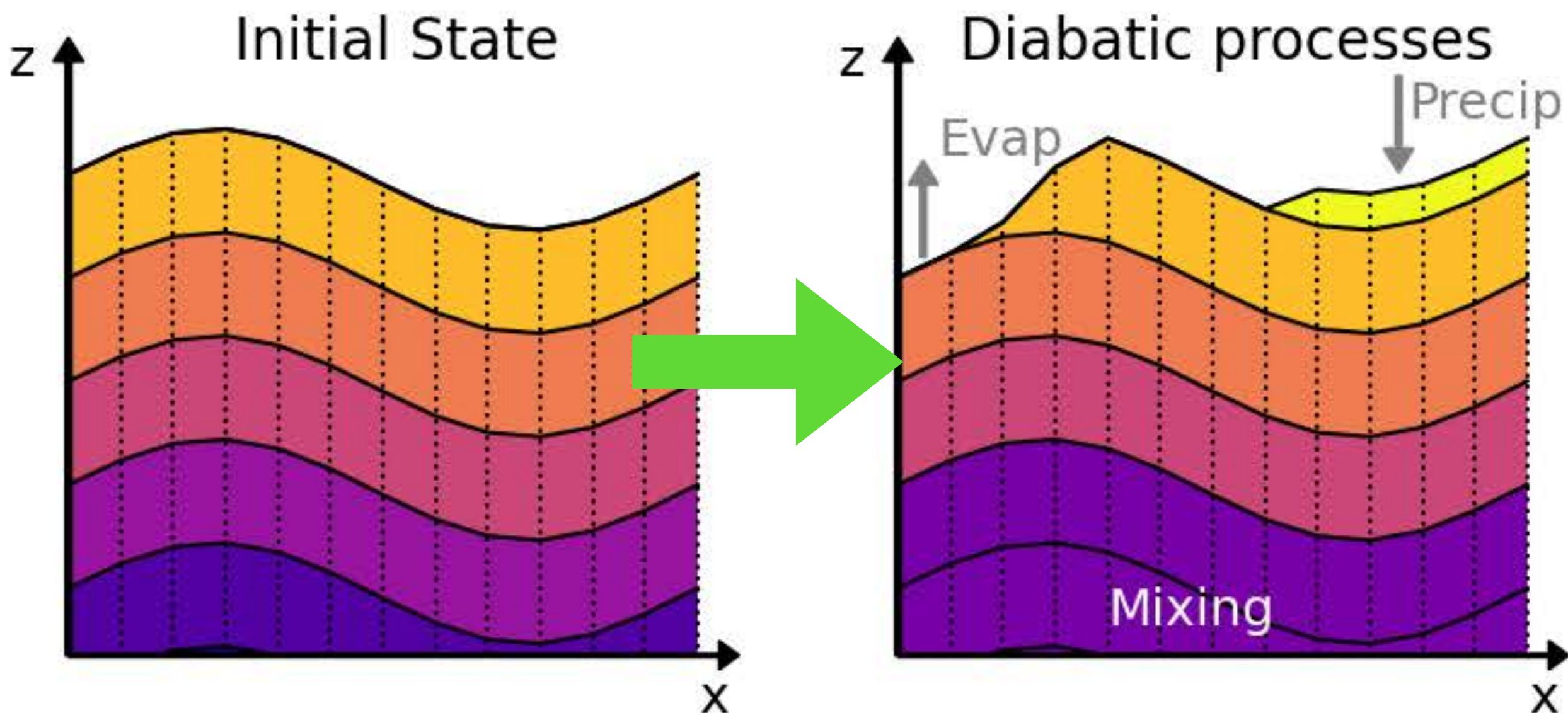


# Online in MOM6, we cannot use a traditional scheme

$w$  does not exist

During the dynamics step,  
the layers move together with  
the fluid

Order of operations usually  
proceeds like this:

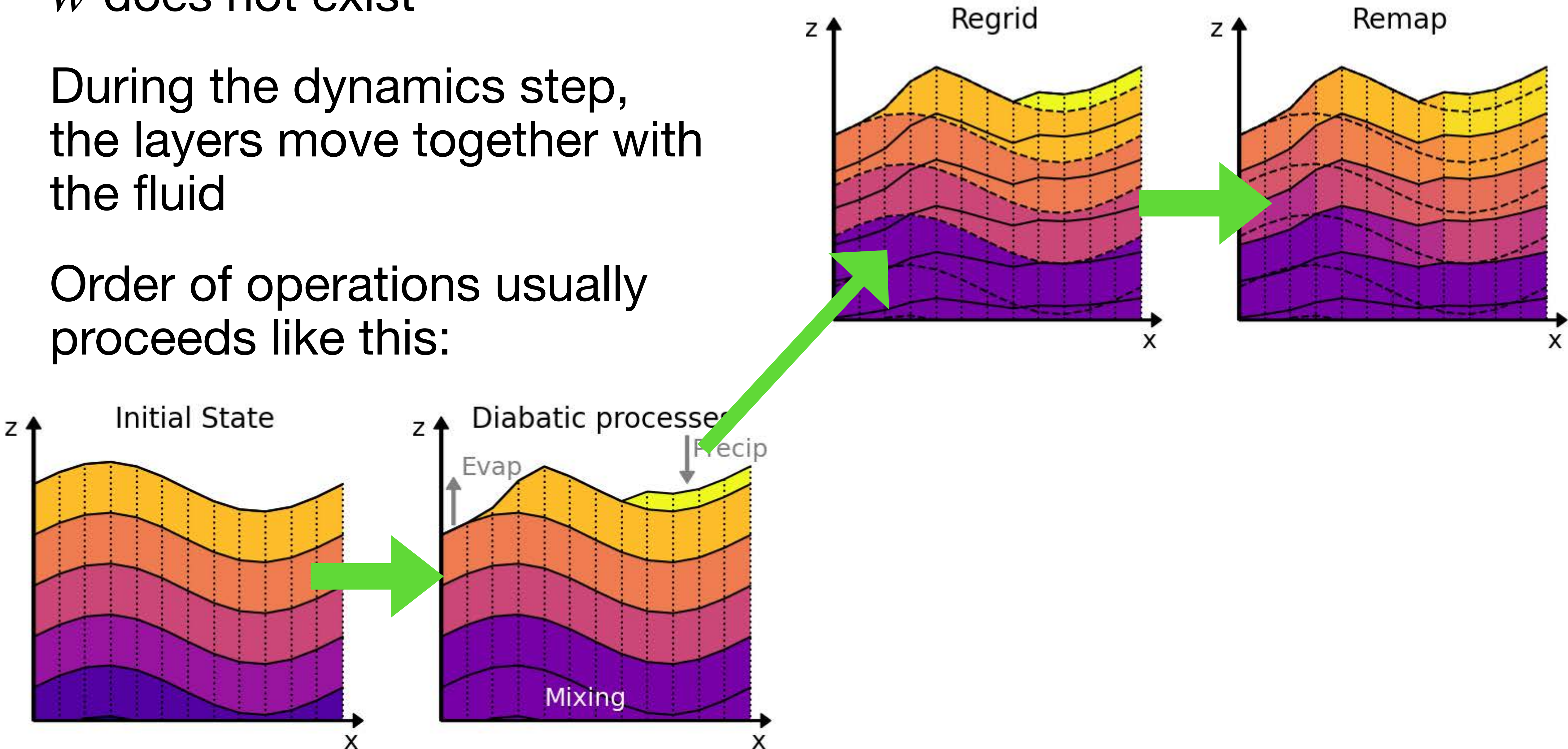


# Online in MOM6, we cannot use a traditional scheme

$w$  does not exist

During the dynamics step, the layers move together with the fluid

Order of operations usually proceeds like this:

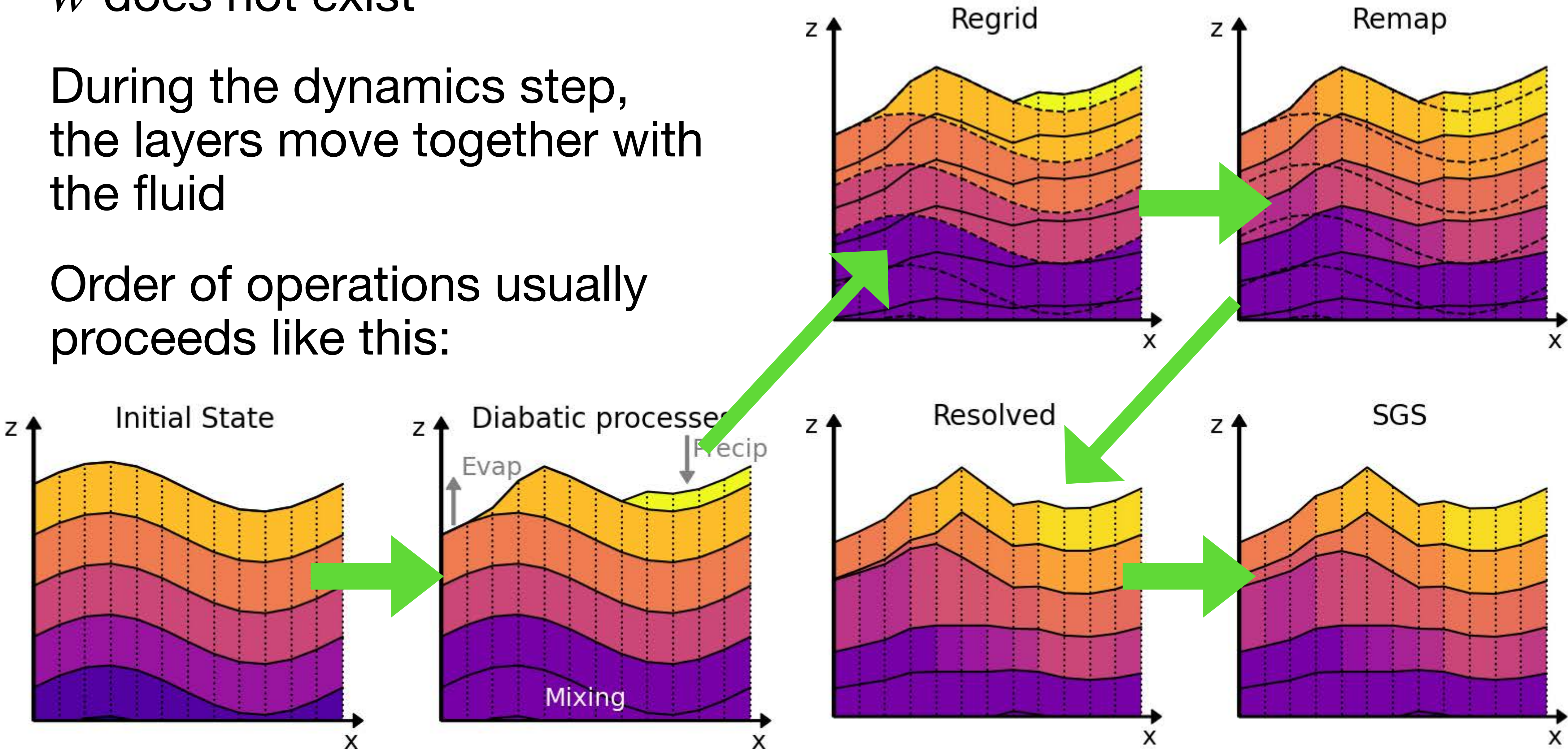


# Online in MOM6, we cannot use a traditional scheme

$w$  does not exist

During the dynamics step, the layers move together with the fluid

Order of operations usually proceeds like this:

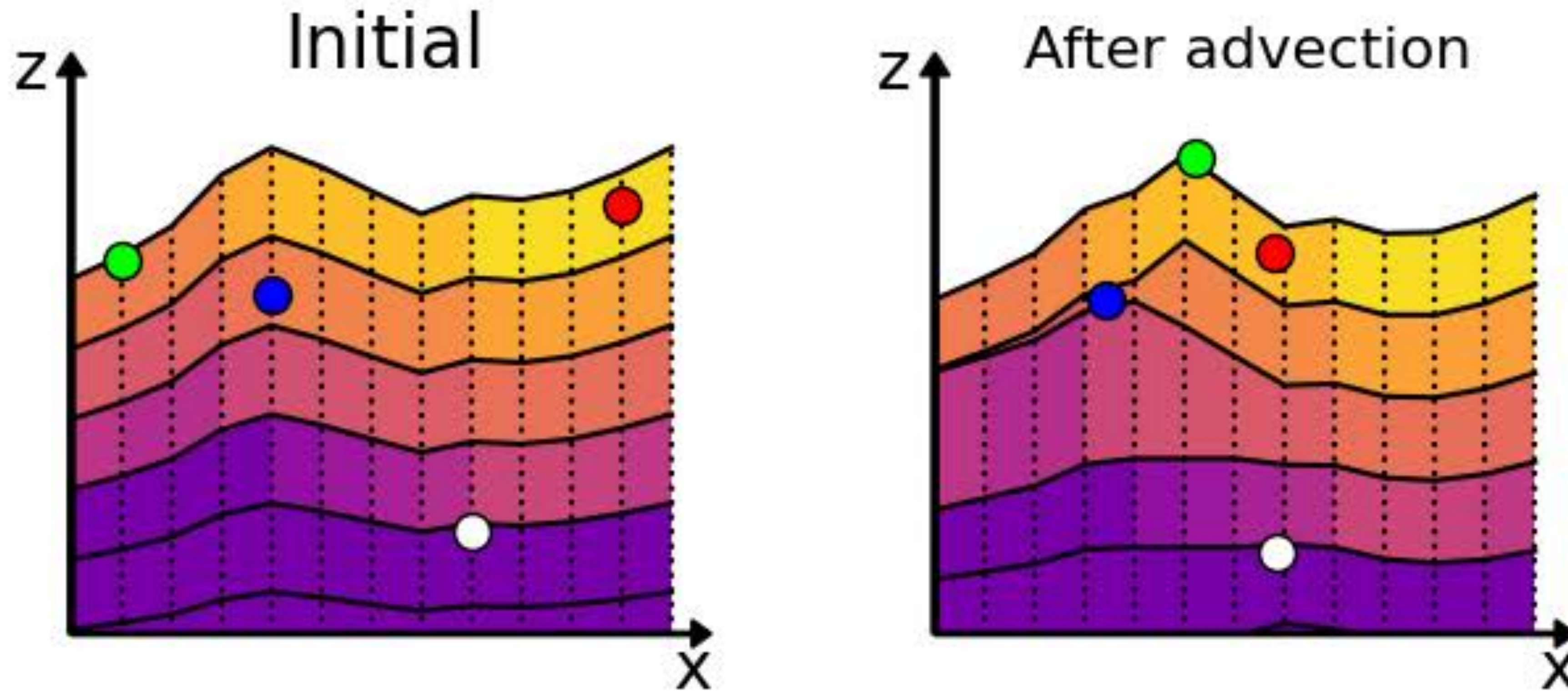


# Online in MOM6, we cannot use a traditional scheme

That's a lot of steps! But let's first just focus on the dynamics step, in which the water actually moves



# Particle advection during dynamics step



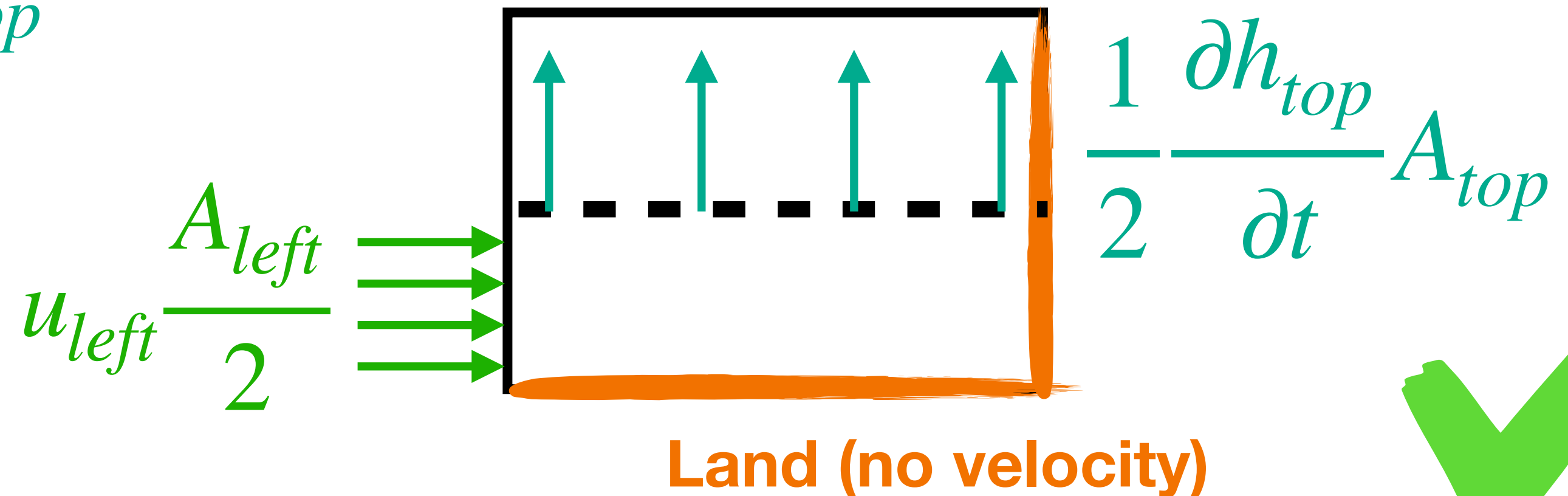
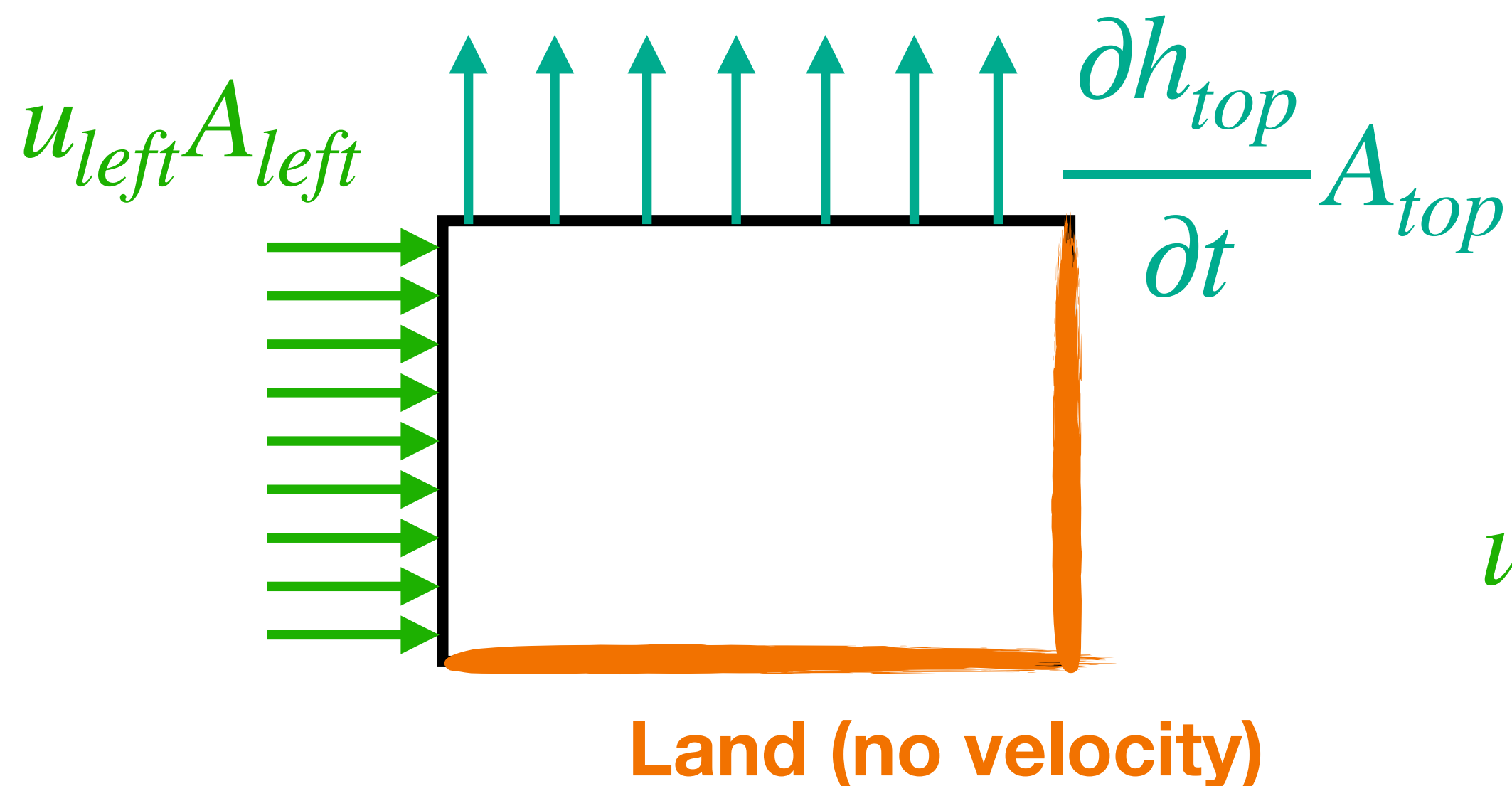
Particles are advected horizontally and maintain their fractional position within each layer

# Mass conservation in this new setup

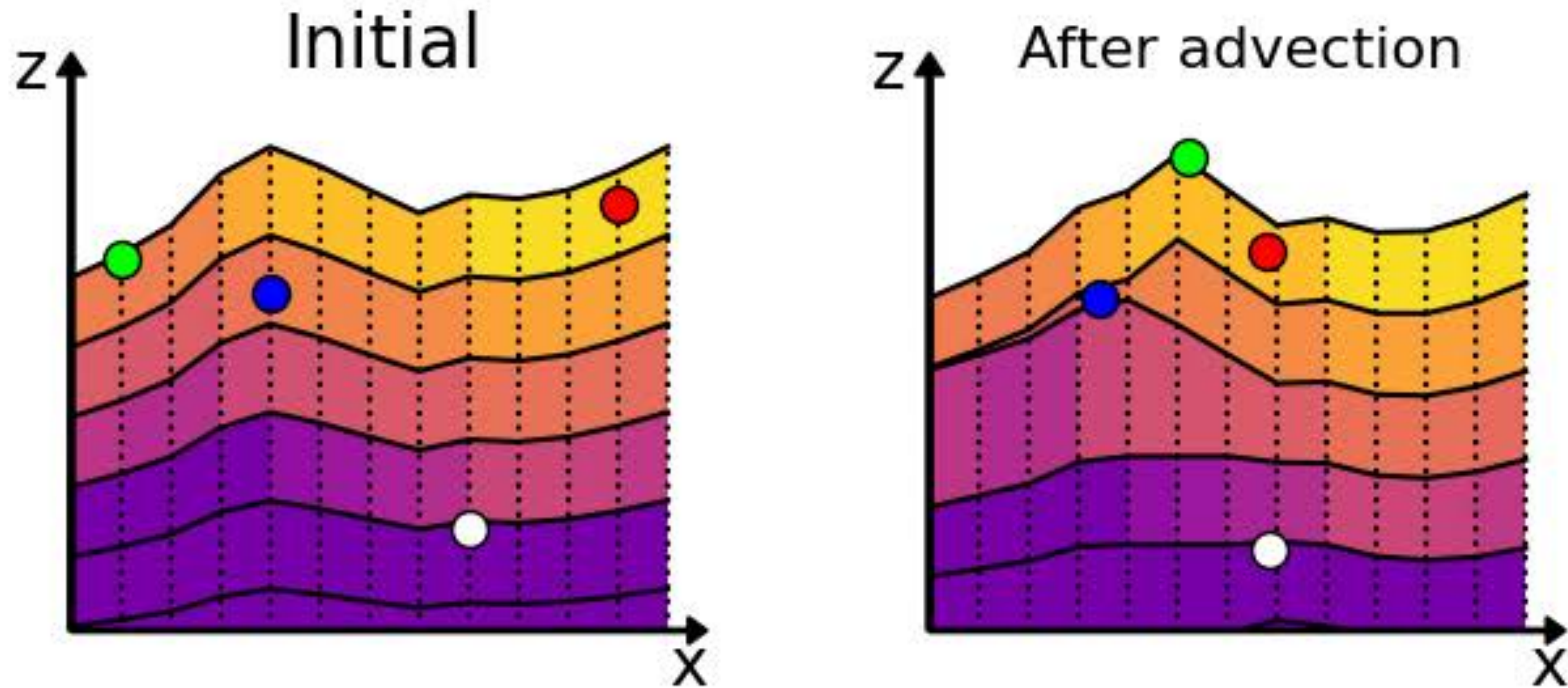
- $u$  is interpolated linearly in the  $x$ -direction, and is constant in  $y, z$
- $v$  is interpolated linearly in the  $y$ -direction, and is constant in  $x, z$
- Particle maintains its fractional position in the cell

**Volume is conserved because top interface moves**

**Subdividing the grid cell does not create convergence of mass, because “center” interface moves half the distance of top interface**



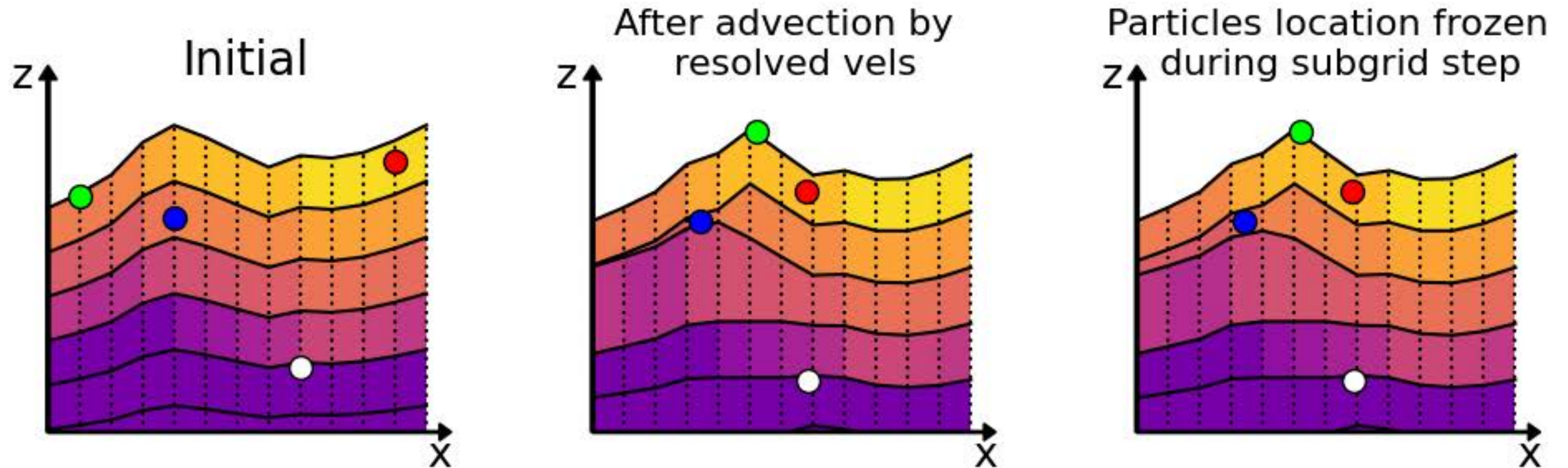
# Our code has 2 modes



Advect particles with resolved velocities

Advect particles with residual velocities

# Advection with resolved velocities only

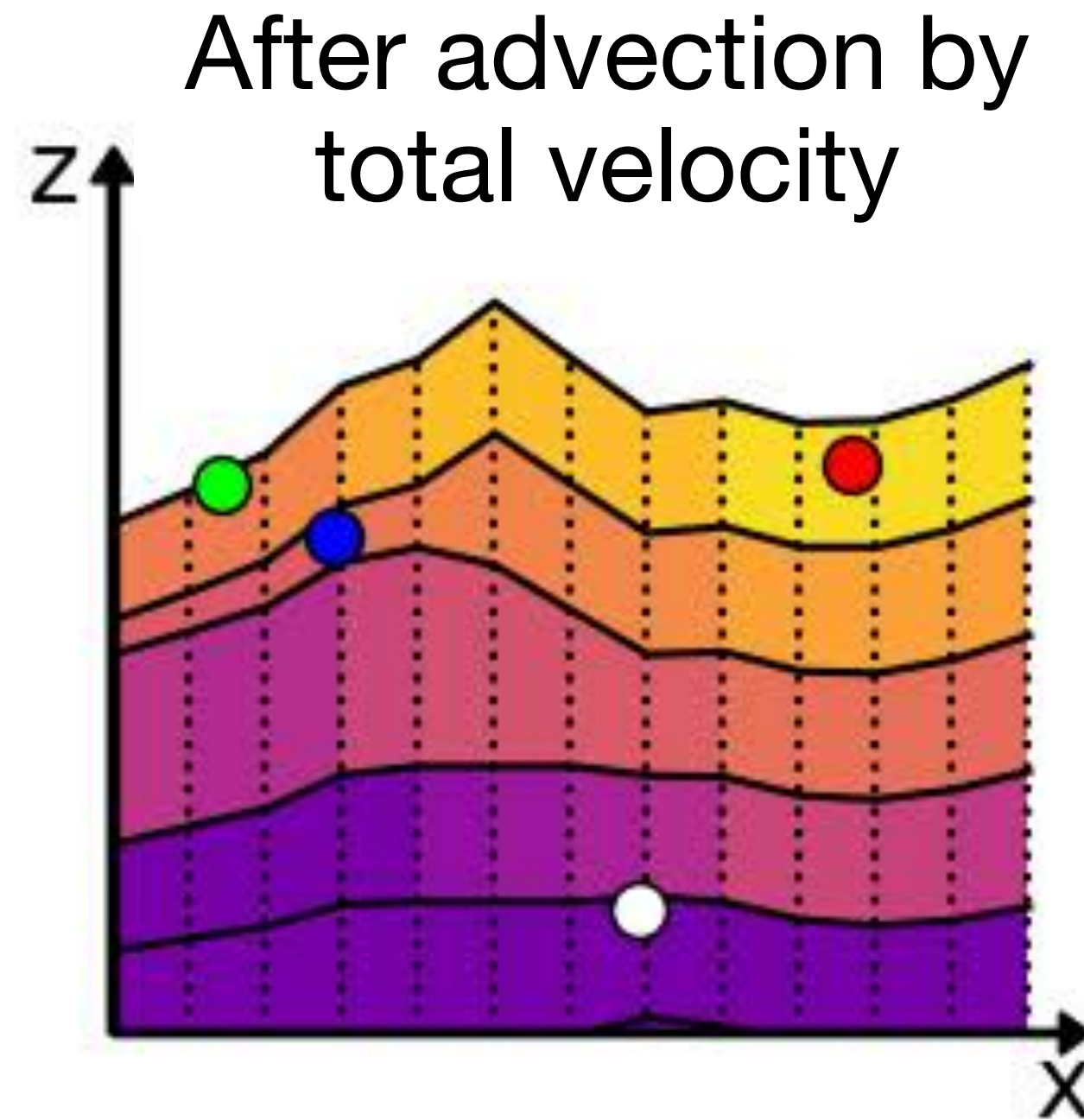
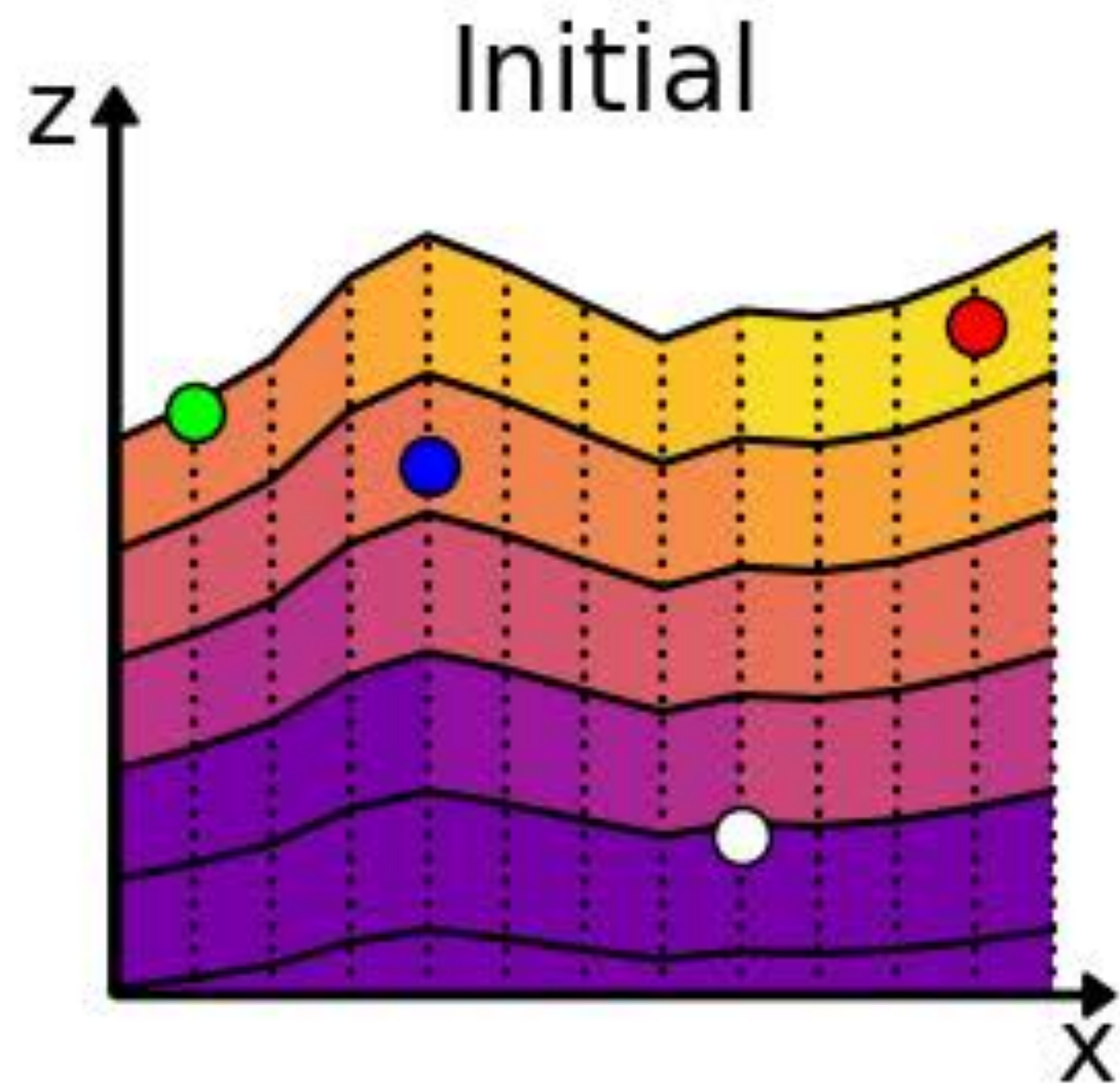


Particles are advected horizontally using the resolved velocities

Particles are frozen in depth and time during the part of the timestep where sub grid scale velocities are used



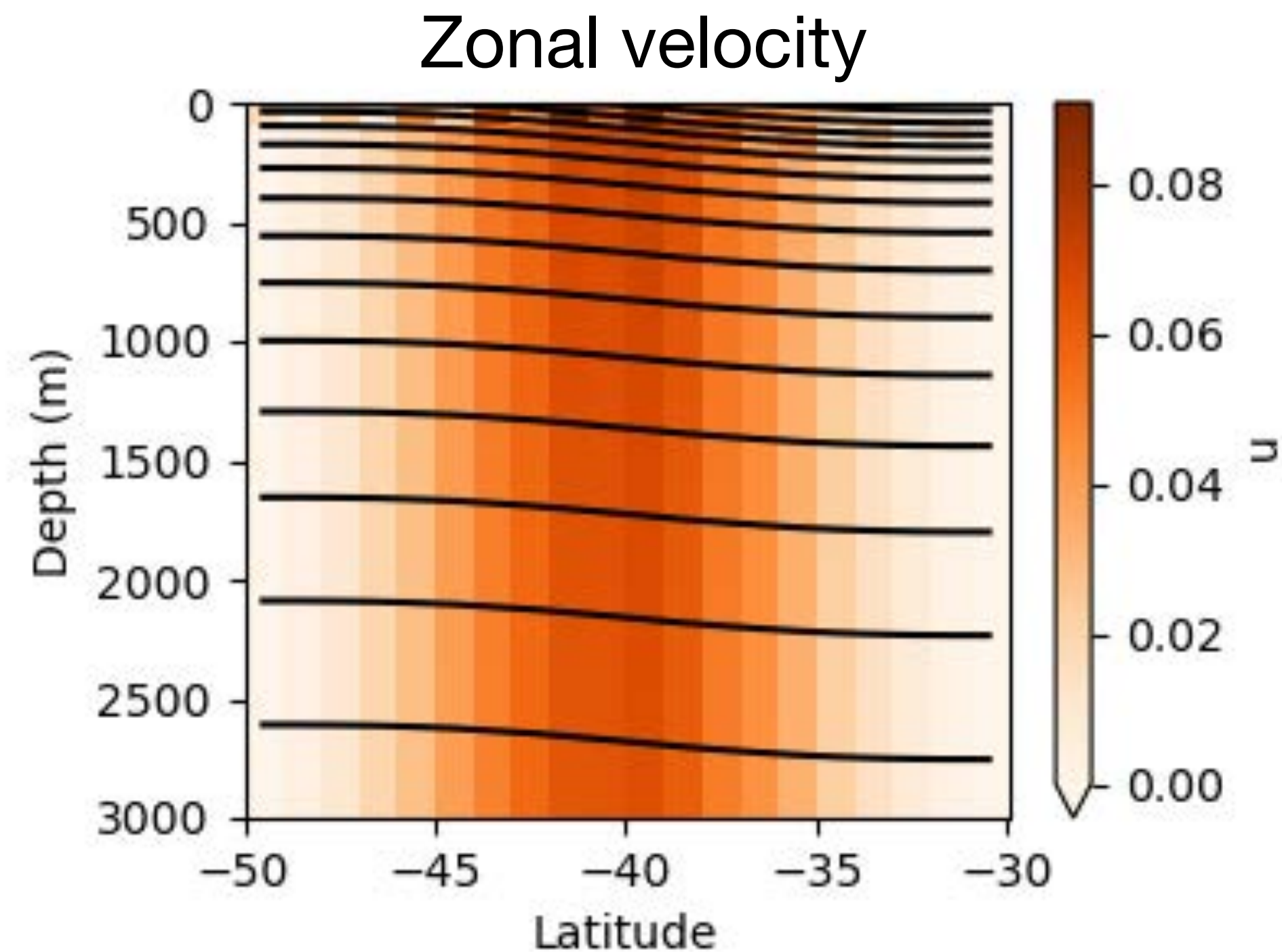
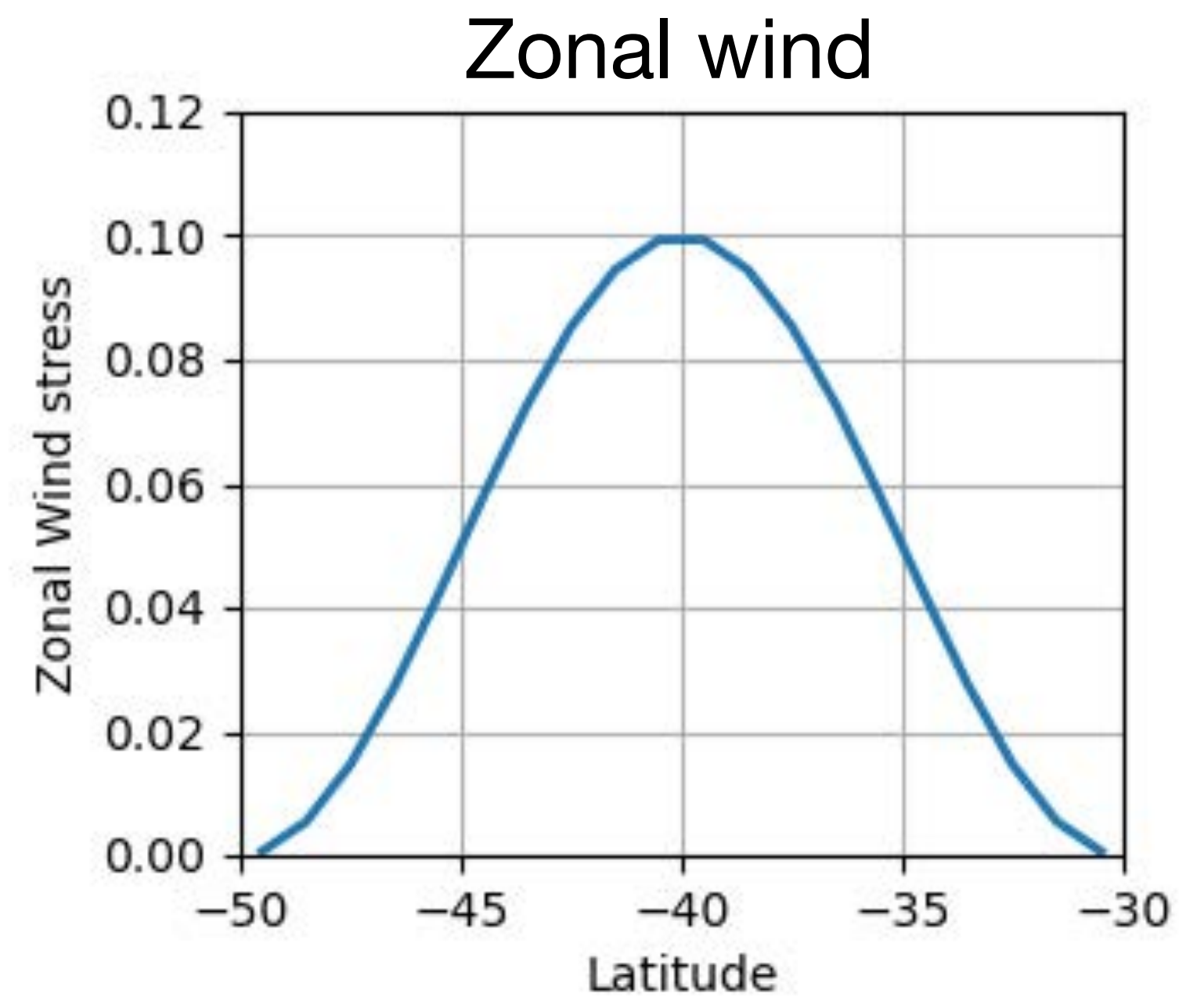
# Advection with full velocity field



In this case, we actually use  $\frac{\mathbf{u}h}{h}$  to advect the particles

# Examples in adiabatic “stacked shallow water” channel

First examples use 1 degree resolution and  
 $\kappa_{gm} = 8000m^2/s$

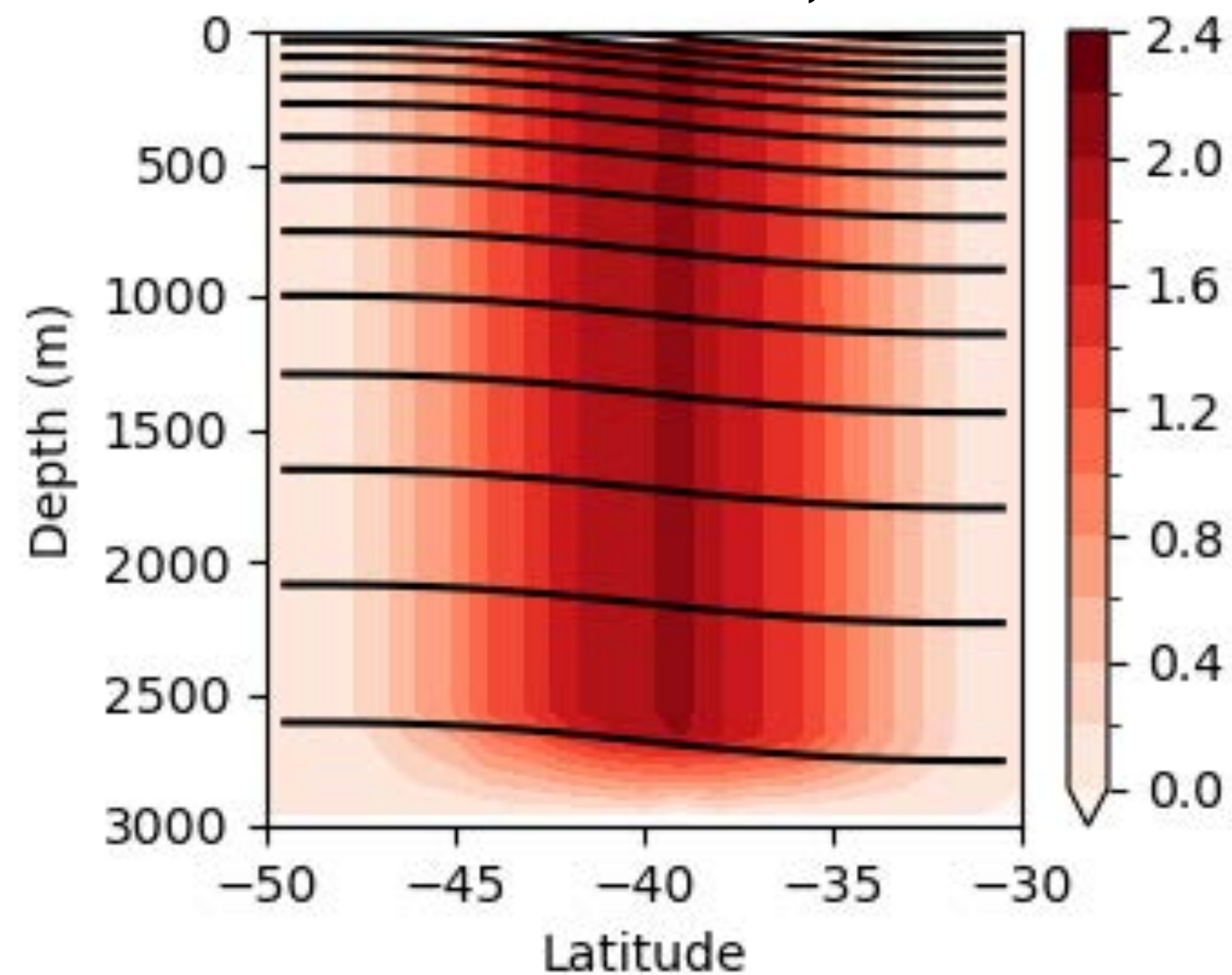


# Examples in adiabatic “stacked shallow water” channel

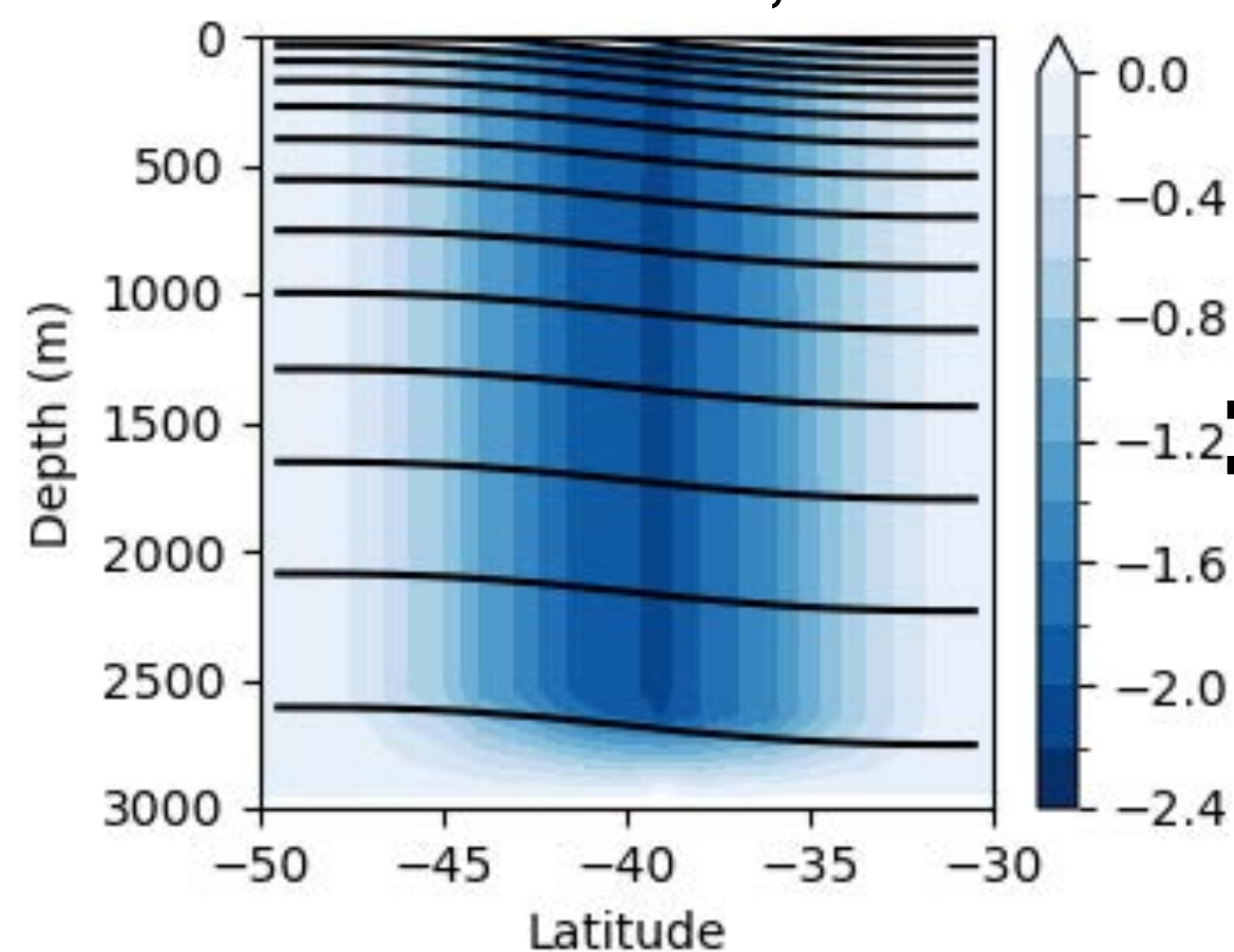
First examples use 1 degree resolution and  $\kappa_{gm} = 8000m^2/s$

In adiabatic mode, there are no buoyancy fluxes, so the residual circulation must be zero everywhere

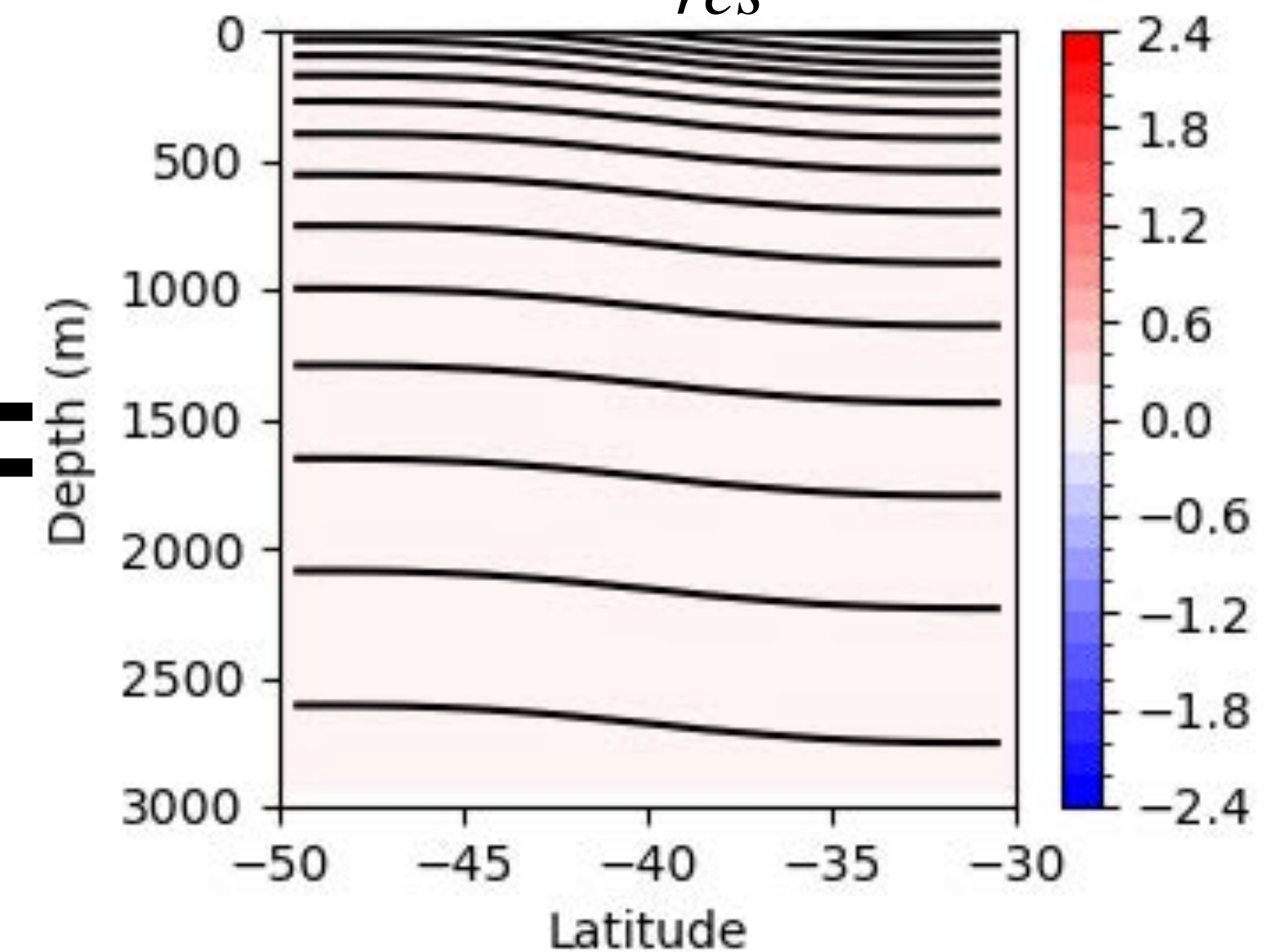
Overturning  
streamfunction for  
time-mean  
circulation,  $\Psi$



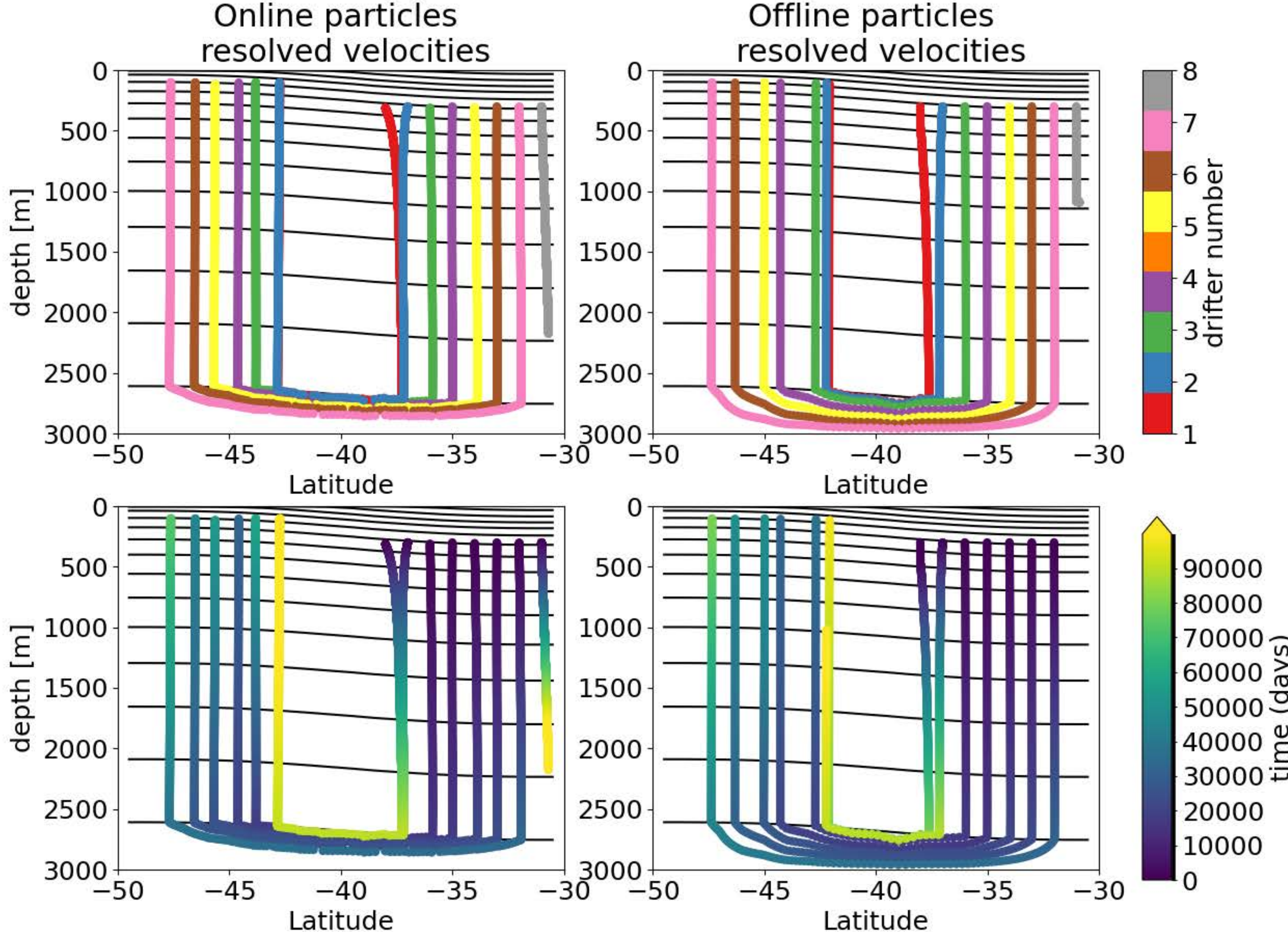
Overturning  
streamfunction for  
eddy-driven  
circulation,  $\Psi^*$



Overturning  
streamfunction for  
residual circulation,  
 $\Psi_{res}$



# Online vs offline trajectories in adiabatic channel (resolved)

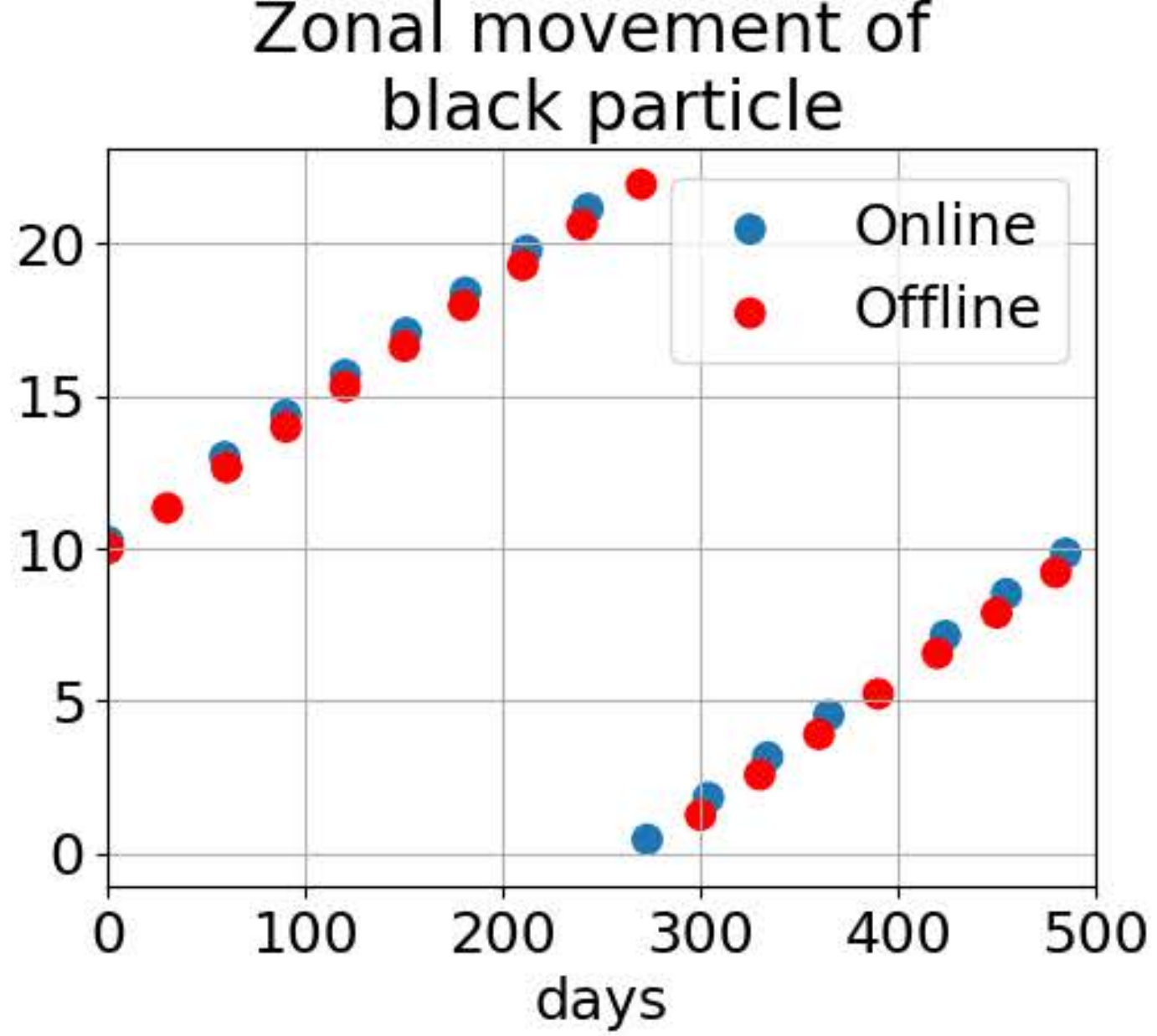
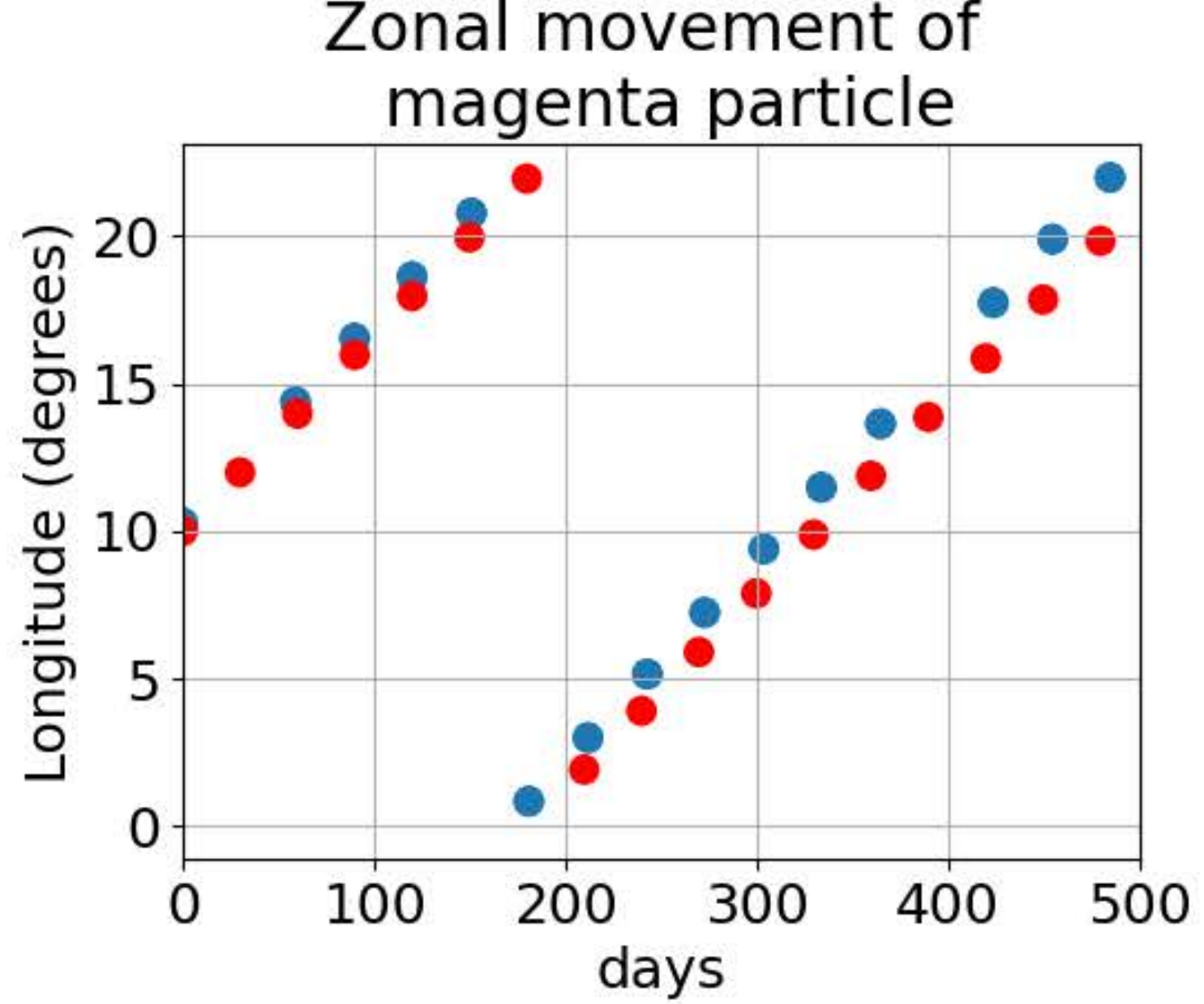
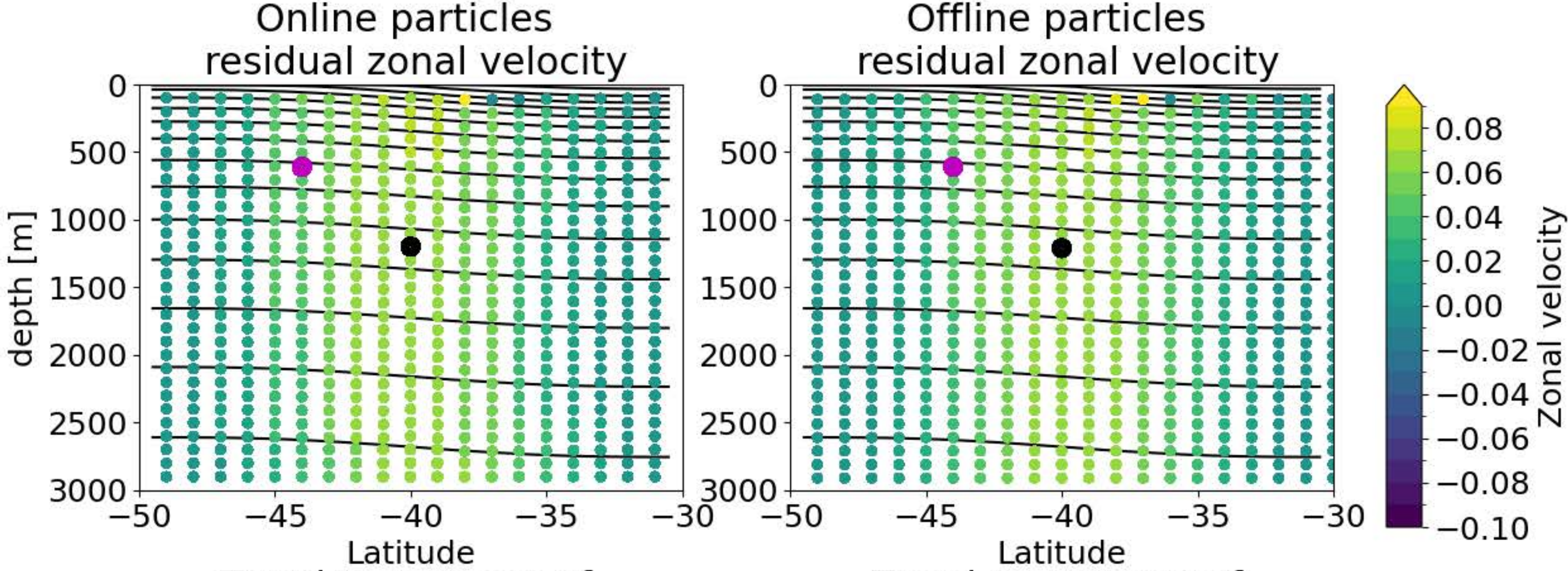


Offline particle advection performed using Parcels

Some small differences that I am still trying to understand

Overall, online particles are spread evenly in upward branch, suggesting mass conservation

# Online vs offline trajectories in adiabatic channel (residual)

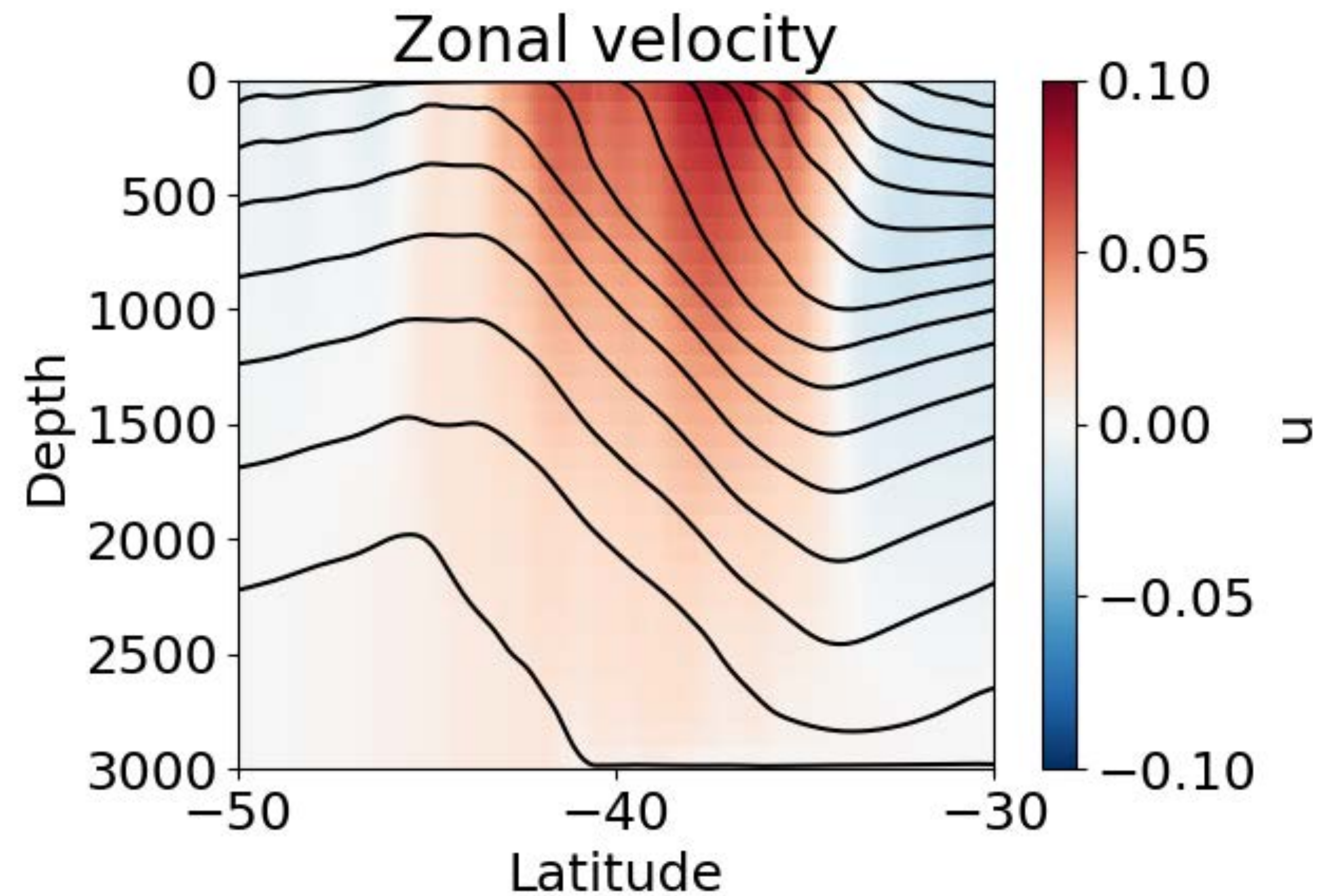


As expected, no flow in meridional or z directions

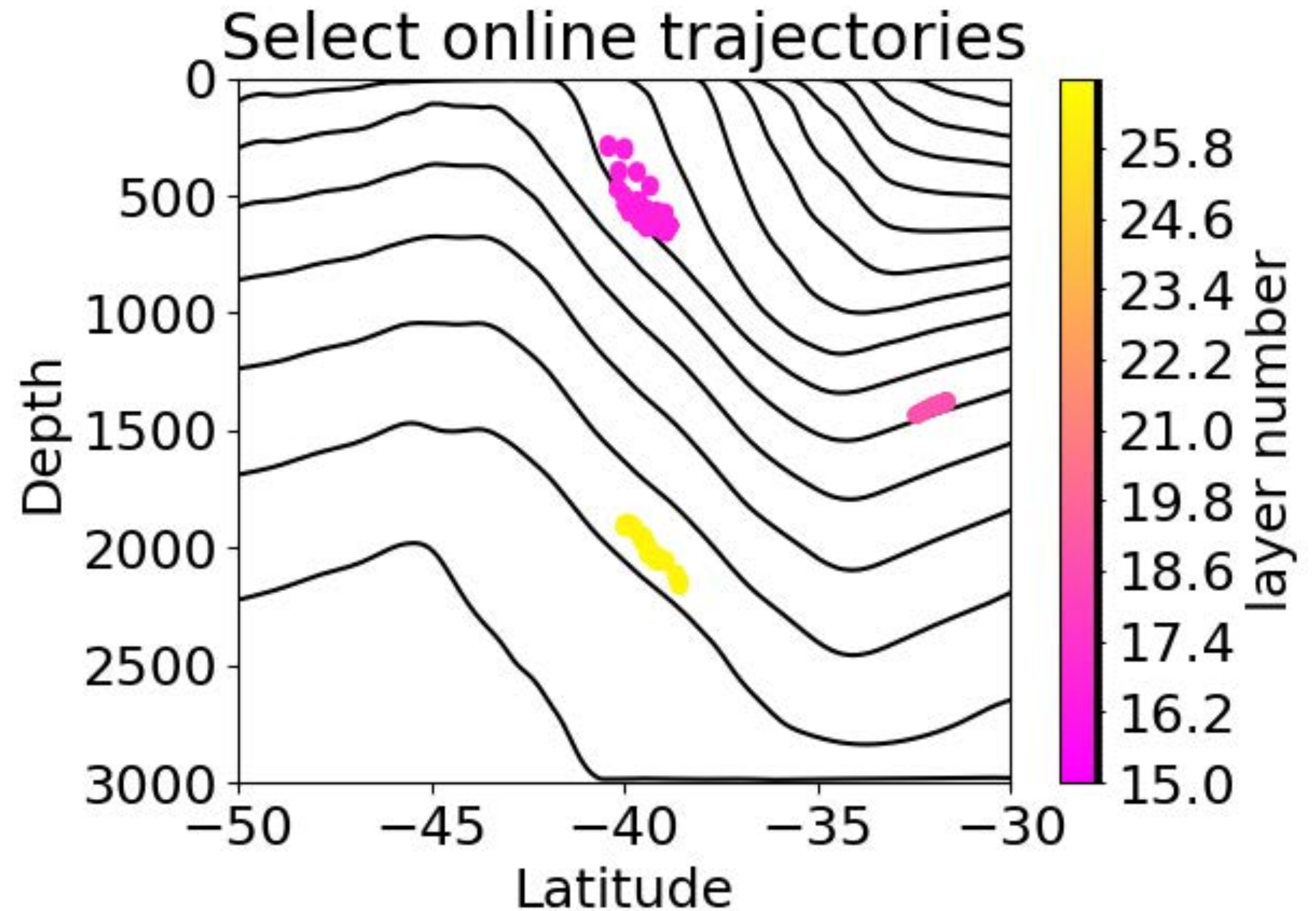
Particle trajectories match very closely in the zonal direction

# Examples in adiabatic “stacked shallow water” channel

Now we go to 0.1 degree resolution and switch off the eddy parameterization

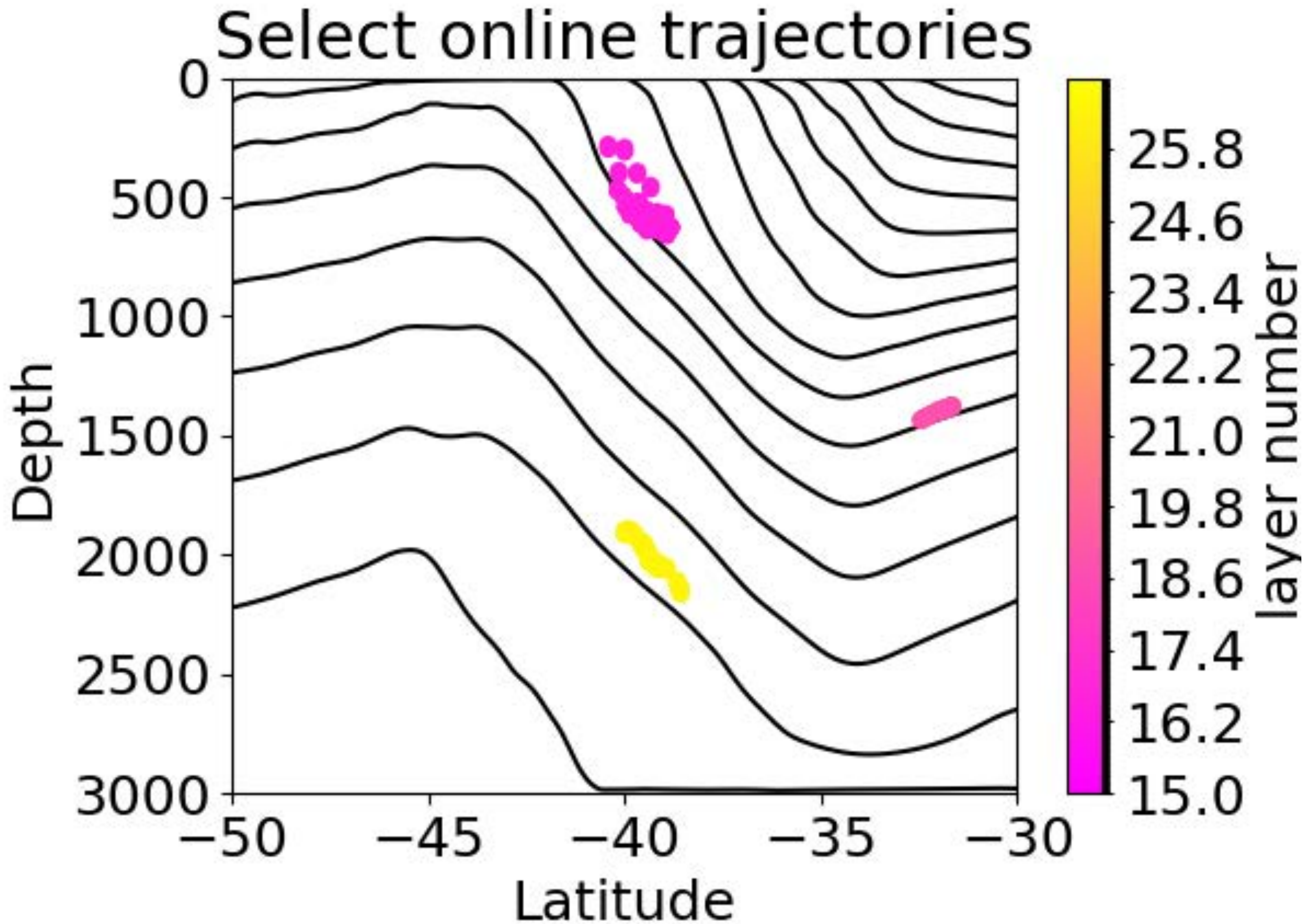


(This run is not in equilibrium yet)

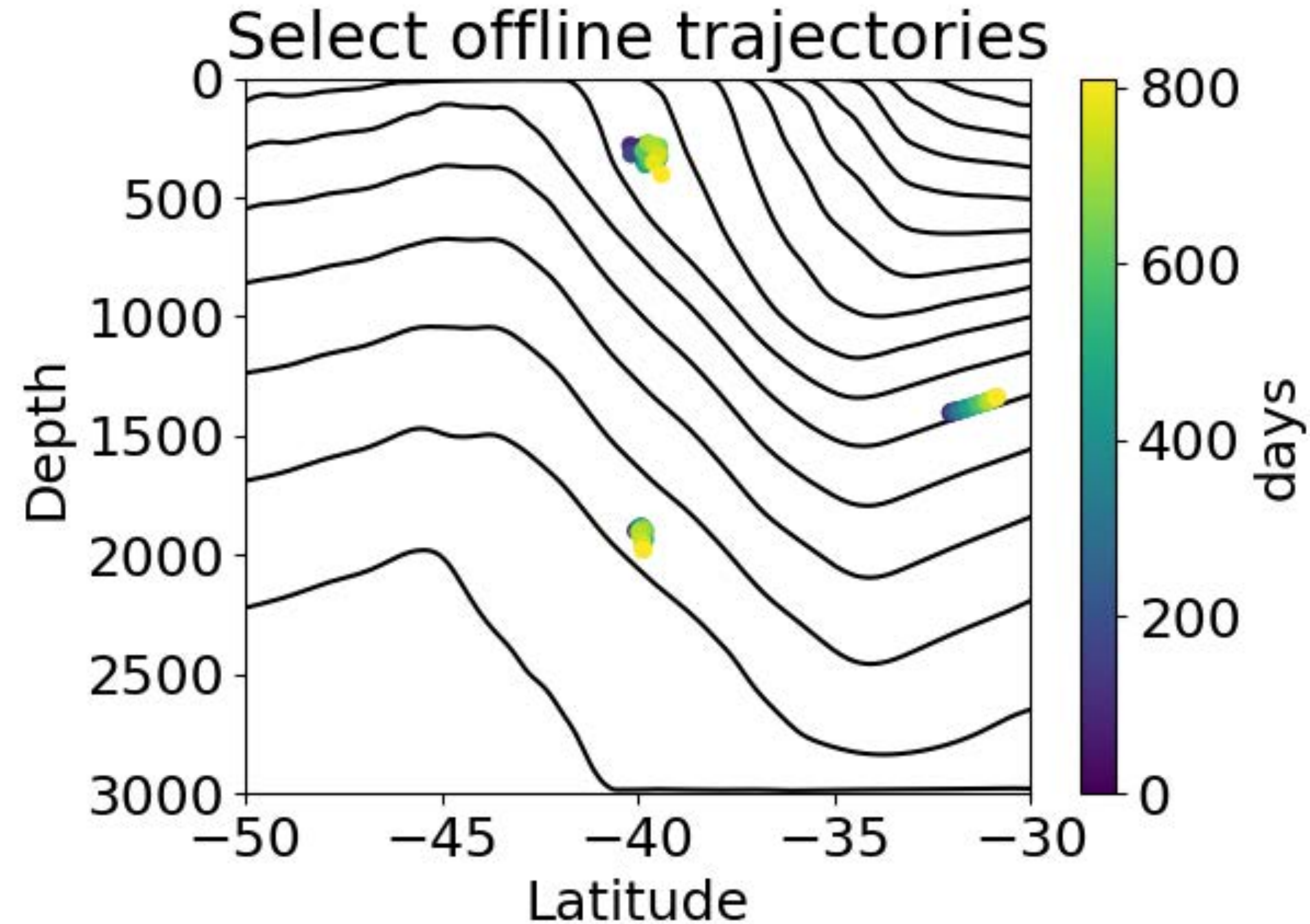


Online trajectories stay in the same layer

# Offline trajectories are sensitive to output frequency

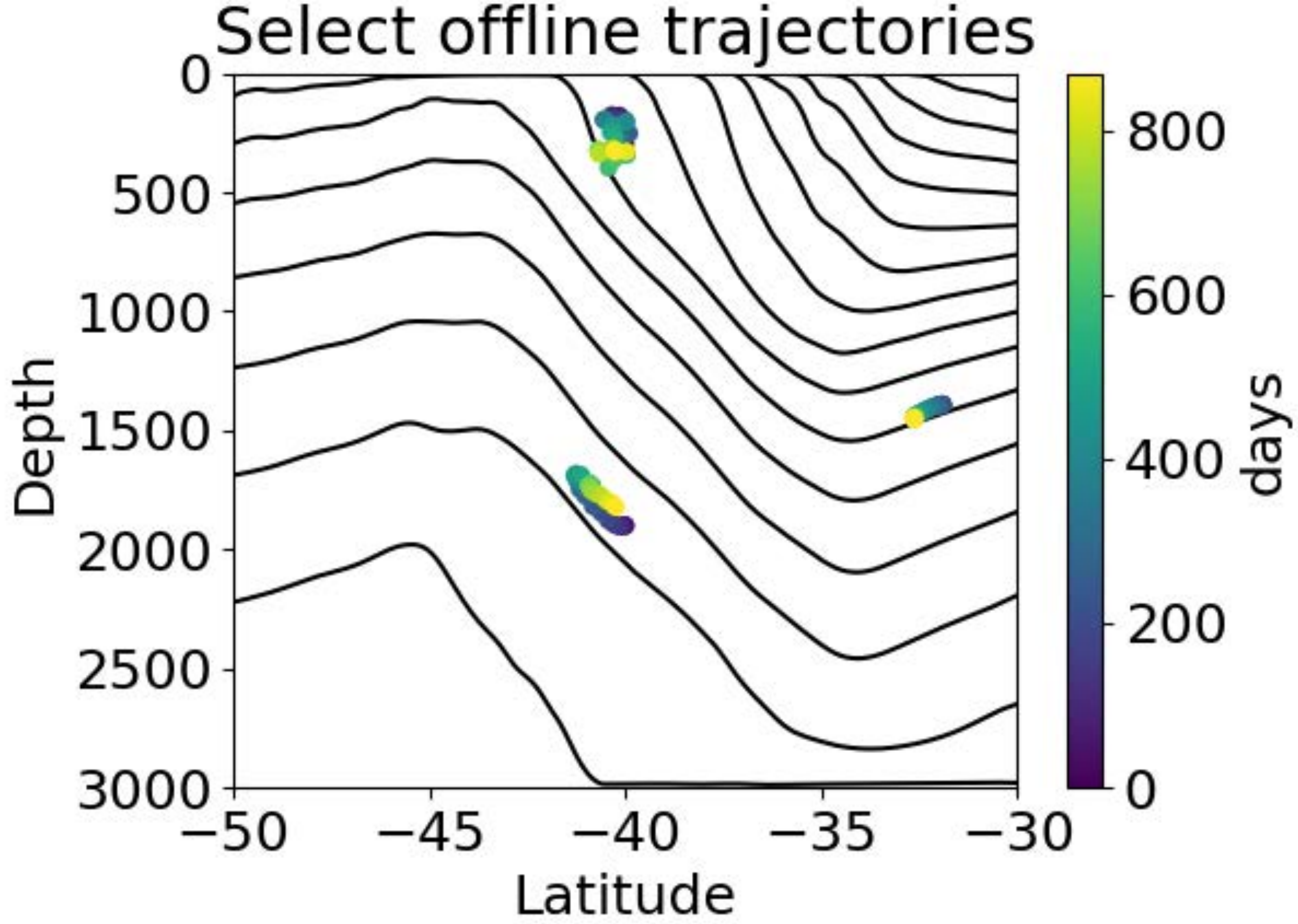
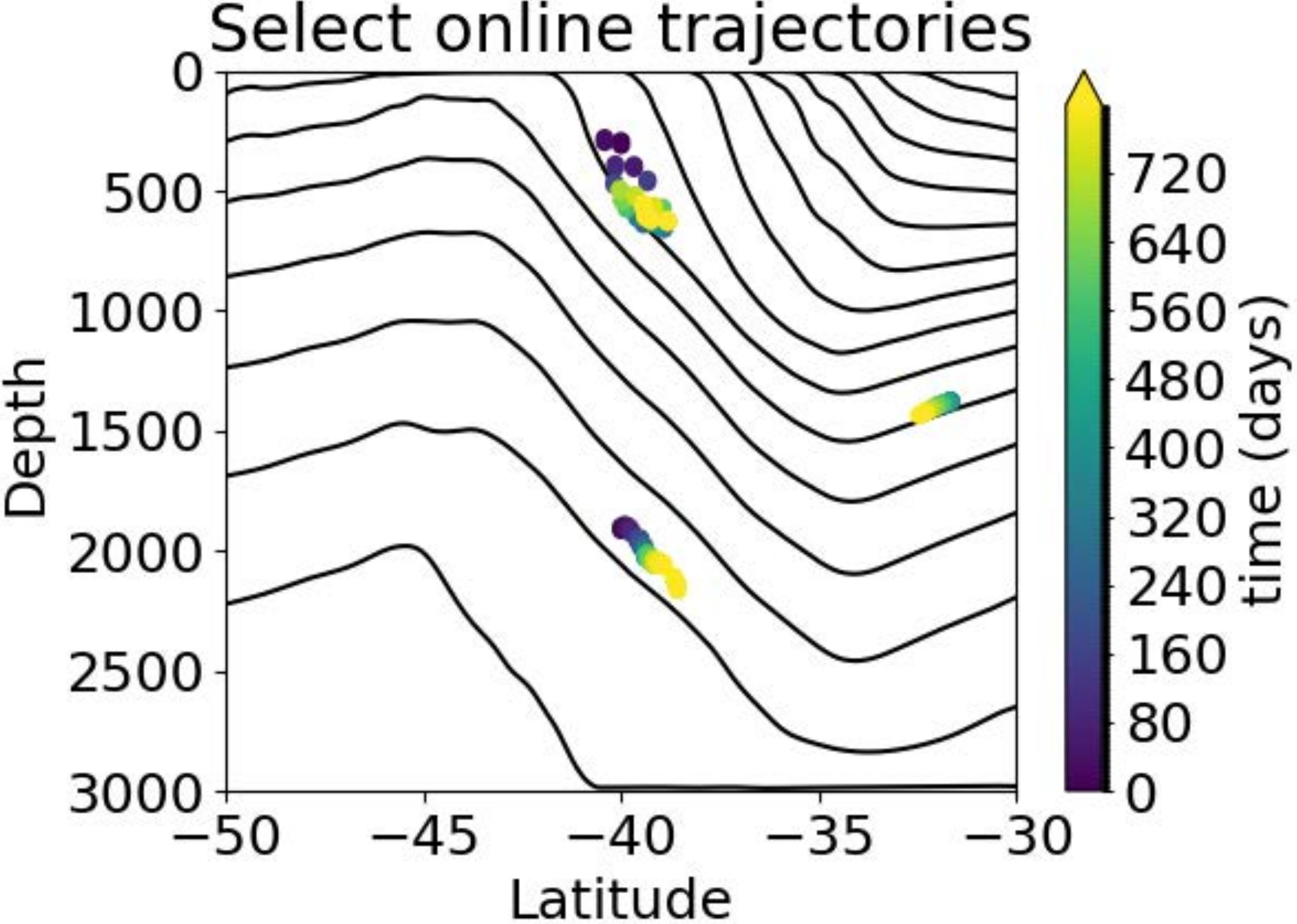


Online trajectories stay in the same layer



e.g. if velocities are averaged over 120 days, particles don't move much because this averages over the eddies

# Offline trajectories are sensitive to output frequency



Online trajectories stay in the same layer

If velocities are averaged over 1 day, offline trajectories are more similar to online trajectories



# Conclusions

- The traditional method for online particle advection will not work in models with a Lagrangian vertical coordinate
- The new method presented here conserves mass
- We can advect particles using both the resolved and the residual velocity field (though not at the same time)
- Online methods generally match offline methods at coarse resolution
- Online methods should have less spurious vertical movement for high resolution

[spencerjones@tamu.edu](mailto:spencerjones@tamu.edu)