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Dr. Fabricio Rodrigues Lapolli; Dr. Maximilian Witte; Dr. Jan Phillip Freese; Prof. Christopher Kadow; Dr. Habil. Peter Korn; Prof. Daniel Ruprecht



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- Climate models are comprised of two important components:
  - dynamical core and subgrid models
  - Dynamical core:
    - Grid resolving features
  - Subgrid Model:
    - Parametrization of unresolved features ((Sub)Mesoscale Parametrization, Ice)

- Physically, the ocean complex dynamics allows for the energy to be transferred for different scales
  - Then this subgrid parametrization will play a higher role



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- In order to improve accuracy of climate projections, finer mesh resolution is required
  - This in turn greatly increases computational cost
- Machine Learning has now been rapidly increasing and being investigated and being used in many climate areas
  - Parametrization (Previous talk Ernout N.)
  - PINN solve free surface (Previous talk Gorenstein I.)
  - Super-resolution (image, video processing)
    - This methodology uses Convolutional Neural, Generative Adversarial Network or Diffusion models
    - Super-resolution has been also successfully applied in climate science as image inpainting



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Ground Truth

Machine Learning Prediction



• Can we train a Machine Learning to help numerical models estimate a high resolution output from a low resolution output?

- Objective:
  - Run a shallow water global simulation (in ICON-O) in a coarse mesh while also correcting it with a ML-model trained with high resolution data



- ICON-O:
  - Dynamical Core:
    - Spherical Grid
    - Finite Volume (C-grid)
      - Based on Admissible Reconstructions (minimizes noise in triangular grids)
    - Triangular grid (icosahedral based)
    - Solves the invariant form equation

$$\underbrace{\frac{\partial \mathbf{v}}{\partial t} + \nabla \cdot (\mathbf{v} \otimes \mathbf{v}) + g\nabla(b+\eta)}_{\text{Advective form}} = \underbrace{\frac{\partial \mathbf{v}}{\partial t} + \omega \mathbf{v} + \nabla \frac{|\mathbf{v}|^2}{2} + g\nabla(b+\eta)}_{\text{Vector invariant form}} = 0$$



- Initial condition (Galewsky et al. 2014):
  - Geostrophic jet at the north hemisphere;
  - With an additional perturbation
  - Barotropic Instability
  - Maximum velocity 80 ms-1
    - Atmospheric initial condition, but the overall dynamics is present in the ocean







- Initial condition (Galewsky et al. 2014):
  - The training data (velocity fields):
    - Pair of 2.5km and 20km (with 12 hour integrated comparison output)
      - Velocity fields substituted HR -> LR
    - 10 days integration with 12 hour output
    - Jet location core at ~42 N +/- 5
    - 8 different perturbation location
    - Total: 23 different conditions with 12 hour output=383 snapshots
  - Coupled Run (Traditional Barotropic Instability):
    - At each 12 hours, the LR output is delivered to the trained ML for correction.



#### • Neural Network:

- The grid is divided into patches and interpolated into a regular high-resolution grid with 10% overlap between each patch;
- Each patch is set inside a U-Net structure (without biases)

$$\mathcal{L} = \mathcal{L}_{abs} + \gamma \mathcal{L}_{rel} \quad \mathcal{L}_{abs} = \frac{1}{N} \sum_{i}^{N} \left( |\bar{u}_i - u'_i| \right)^2 + \left( |\bar{v}_i - v'_i| \right)^2 \quad \mathcal{L}_{rel} = \frac{1}{N} \sum_{i}^{N} \left( \min\left[ 1, \frac{|\bar{u}_i - u'_i|}{u'_i + \epsilon} \right] + \min\left[ 1, \frac{|\bar{v}_i - v'_i|}{v'_i + \epsilon} \right] \right) \\ \min_{\theta} \mathcal{L} \left( \Theta(\mathbf{u}(t + \tau)), \mathbf{u}'_{hr}(t + \tau) \right)$$



• Result (10day integration):



- Maximum Norm error
  - The coupled run initially loses accuracy in the early time
  - As the instability grows the accuracy is improved
  - The same is observed in the height field (which was not corrected by ML)
  - Overall we achieve 10km in our run (2 times fold)
  - As the optimal (blue line) result indicates, there's still room for improvement in the run









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$10\mathrm{km}$	$\mathrm{ML}_{\mathrm{coupled}}$	$20\mathrm{km}$
$63\mathrm{s}$	$55\mathrm{s}$	$31\mathrm{s}$









- Vorticity (instability day 7):
  - The low resolution triggers the instability earlier due to unstructured grid error, show faster meanderings
  - The coupled run delays this instability with a lower meandering equivalent to the 10km run









#### • Energy conservation

- Our choice of assimilation with ML inputs energy into the system;
  - The growth does not seem constant as it decreases after day 6
  - The noise present in the vorticity is possibly being an added energy in the system
- Potential energy conservation is not violated with our methodology (given that the height is not changed)





#### • Spectral energy

- The energy spectrum is violated for lower wavenumbers than the other runs
  - The observed input of energy observed is mostly concerned in the smallest scales
  - It does not seem to leak into the largest scales of the grid
- The Enstrophy is not significantly improved, but it is not damaged by the assimilation
- Non-dissipative simulation
  - In a more real dissipative simulation, these noises may spuriously contribute to mixing







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#### Summary

Summary

- Our ML methodology for shallow water show potential use for increasing the accuracy of fluid simulations
- The use CNN (possibly due to its necessity of regular grids) provides spurious noise to the system
  - These can, however, be mitigated by low order pass filters
  - Use ML Transformers Network? (ongoing)
- The assimilation (of substituting variables) may also be improved by diminishing this noise and increase the accuracy (4Dnet, nudging, etc.)
- Future Work:
  - Investigate improvements in data-assimilation technique for the ML coupled runs (preliminary)
  - Investigate different frequencies of assimilation (ongoing)
  - Apply results in a real 3D ocean



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#### Corrected

#### Not Corrected

