

Bathymetry-aware mesoscale eddy parameterizations

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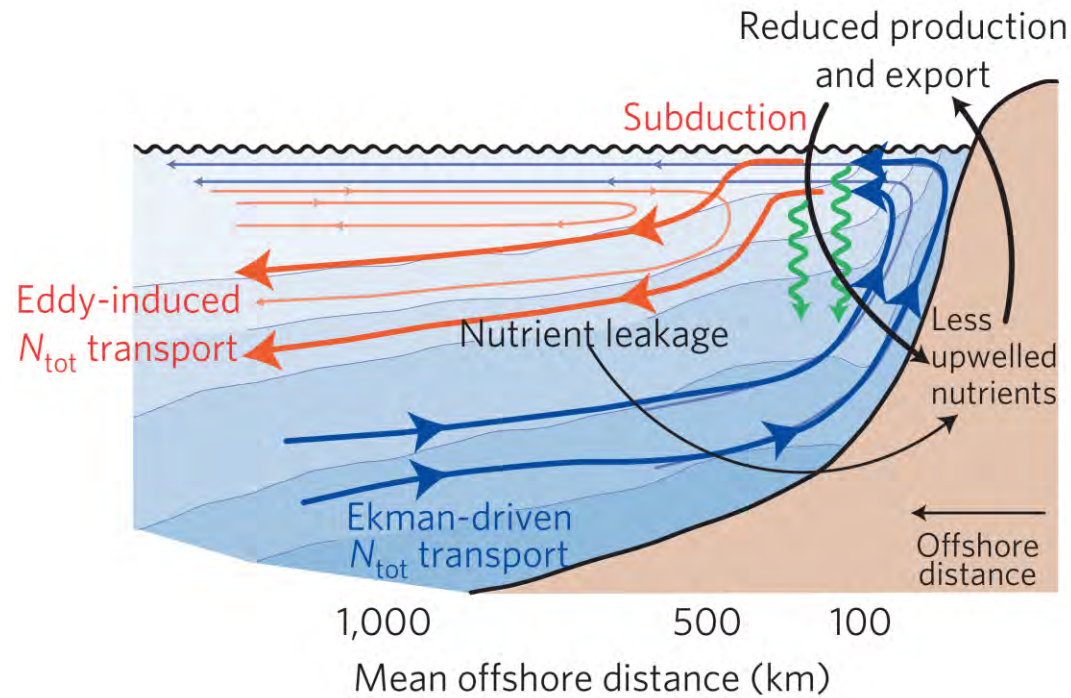
Joint Work with:

Huaiyu Wei (former student@HKUST; now postdoc@UCLA),

Chenyue Xie (former postdoc@HKUST; now faculty@USTC),

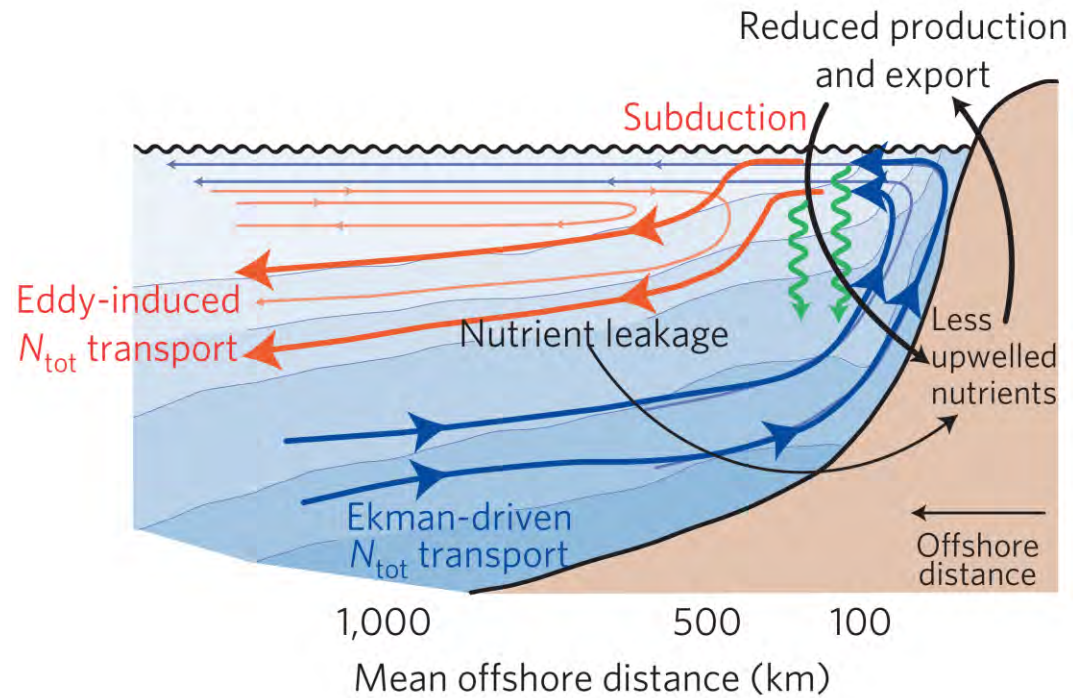
Julian Mak (HKUST), and Andrew Stewart (UCLA)

Eddies across continental margins



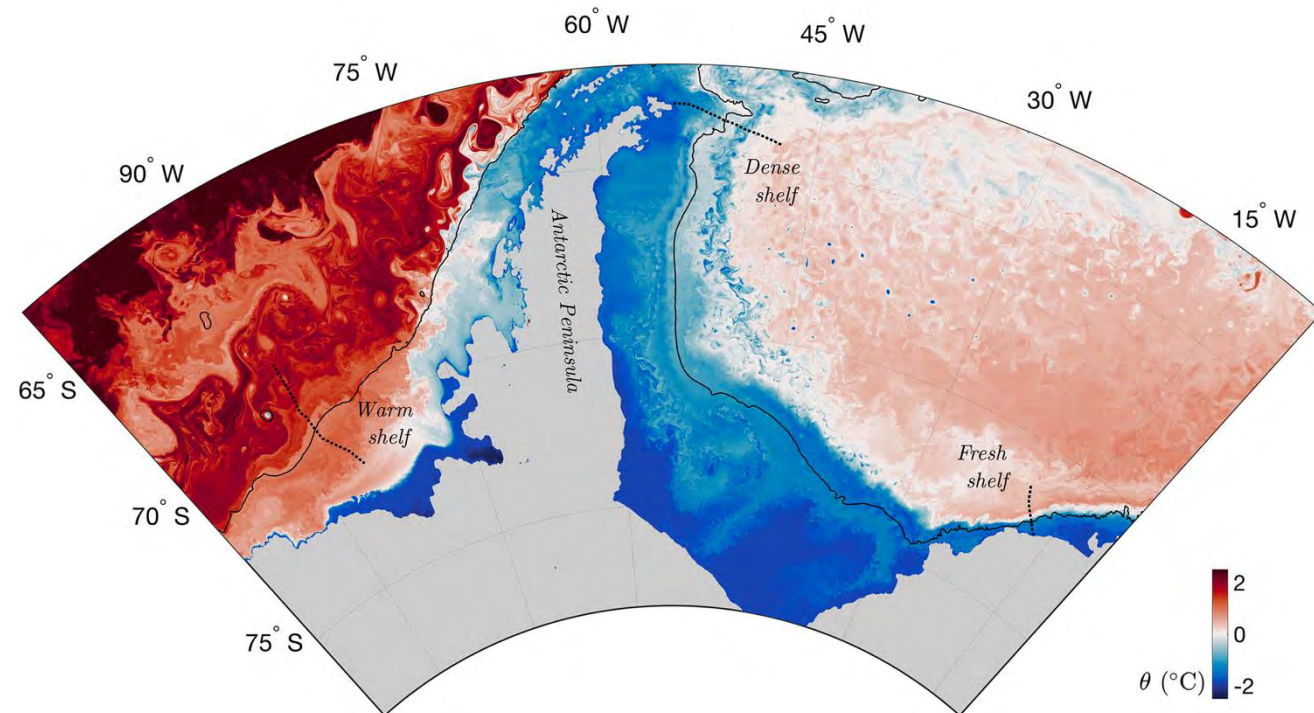
Eddies export nutrients off the Eastern Boundary Upwelling System, thus limiting marine productivity (Gruber et al. 2011).

Eddies across continental margins

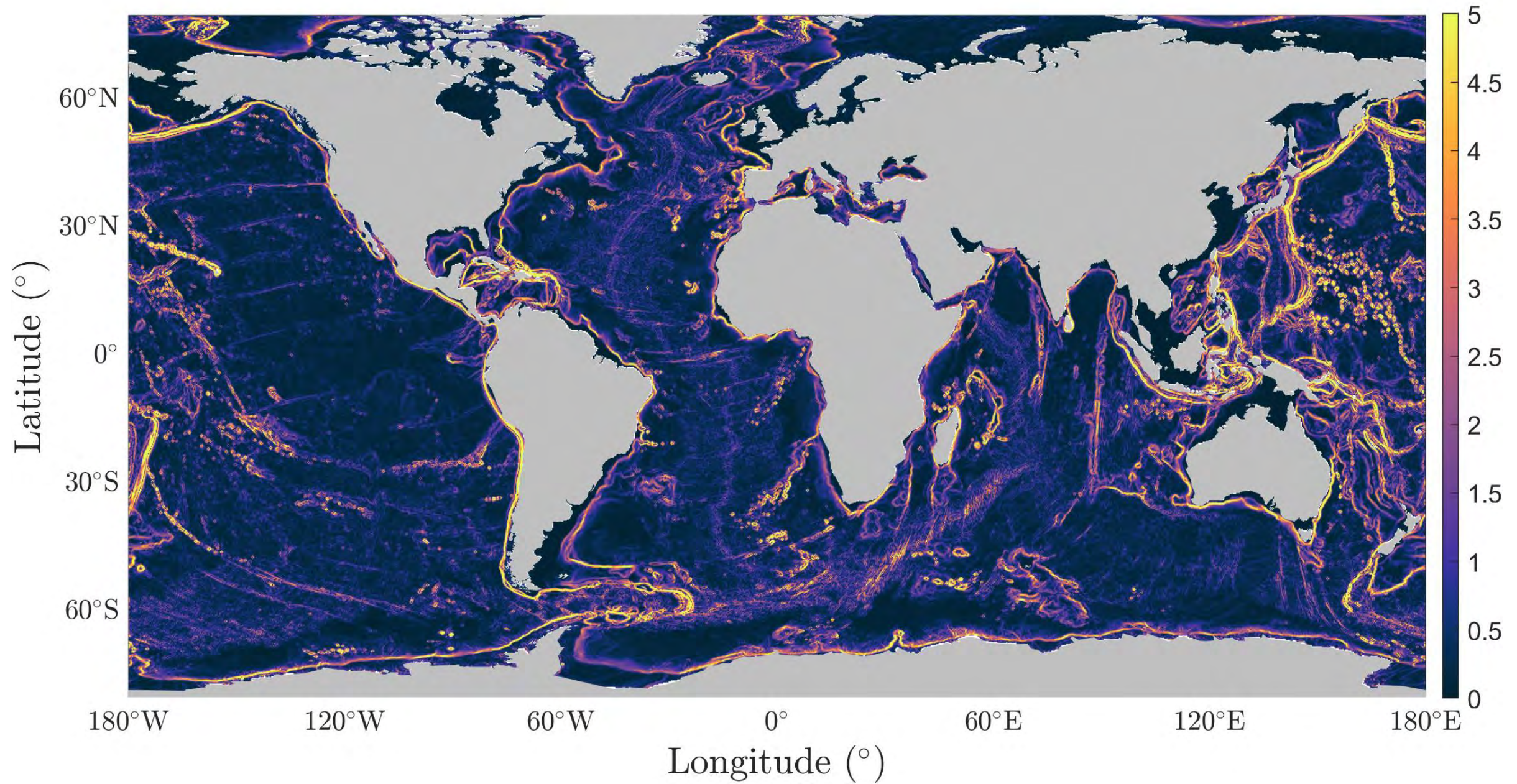


Eddies drive heat fluxes toward the ice shelves of Antarctica, shaping the abyssal MOC (Thompson et al. 2018).

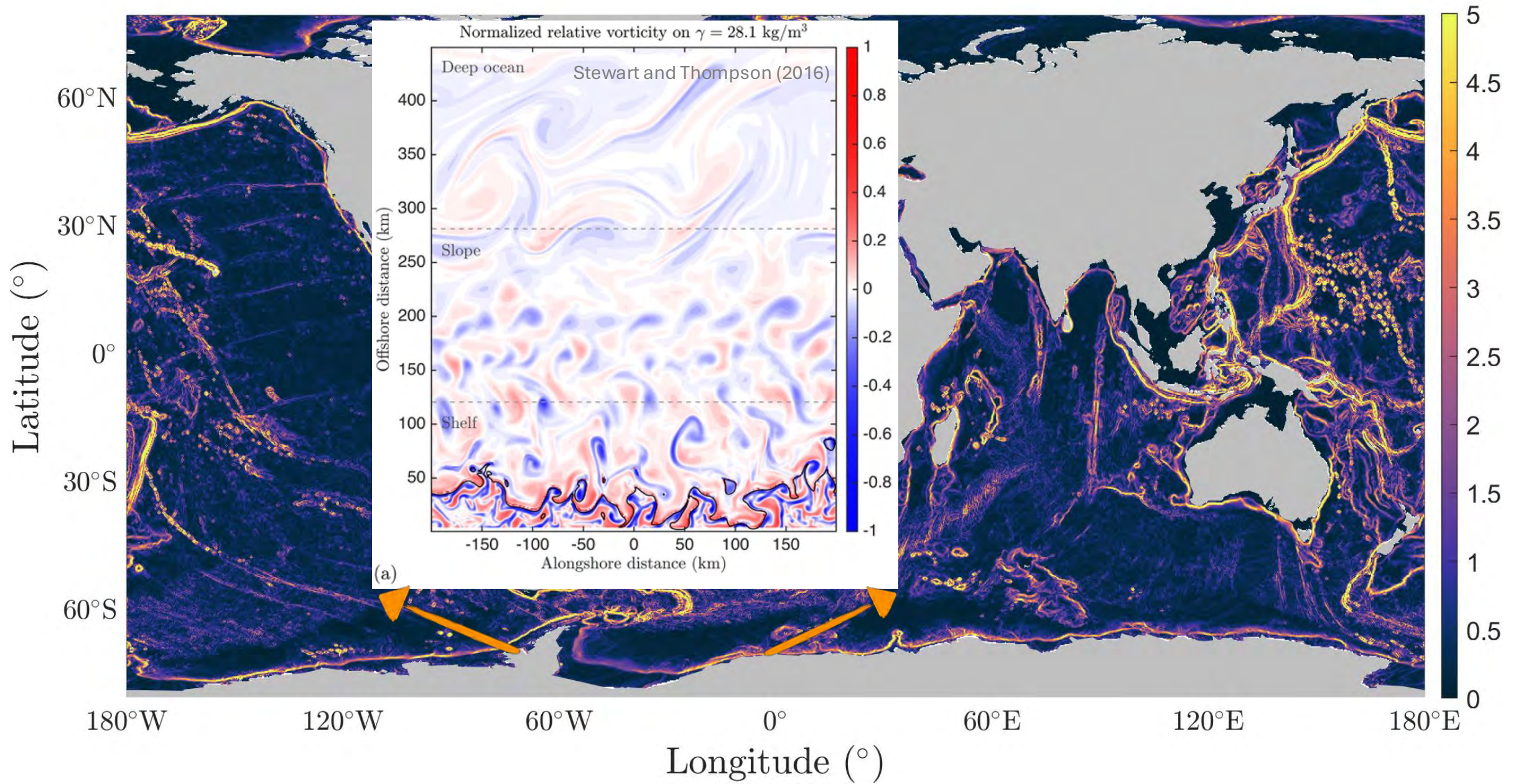
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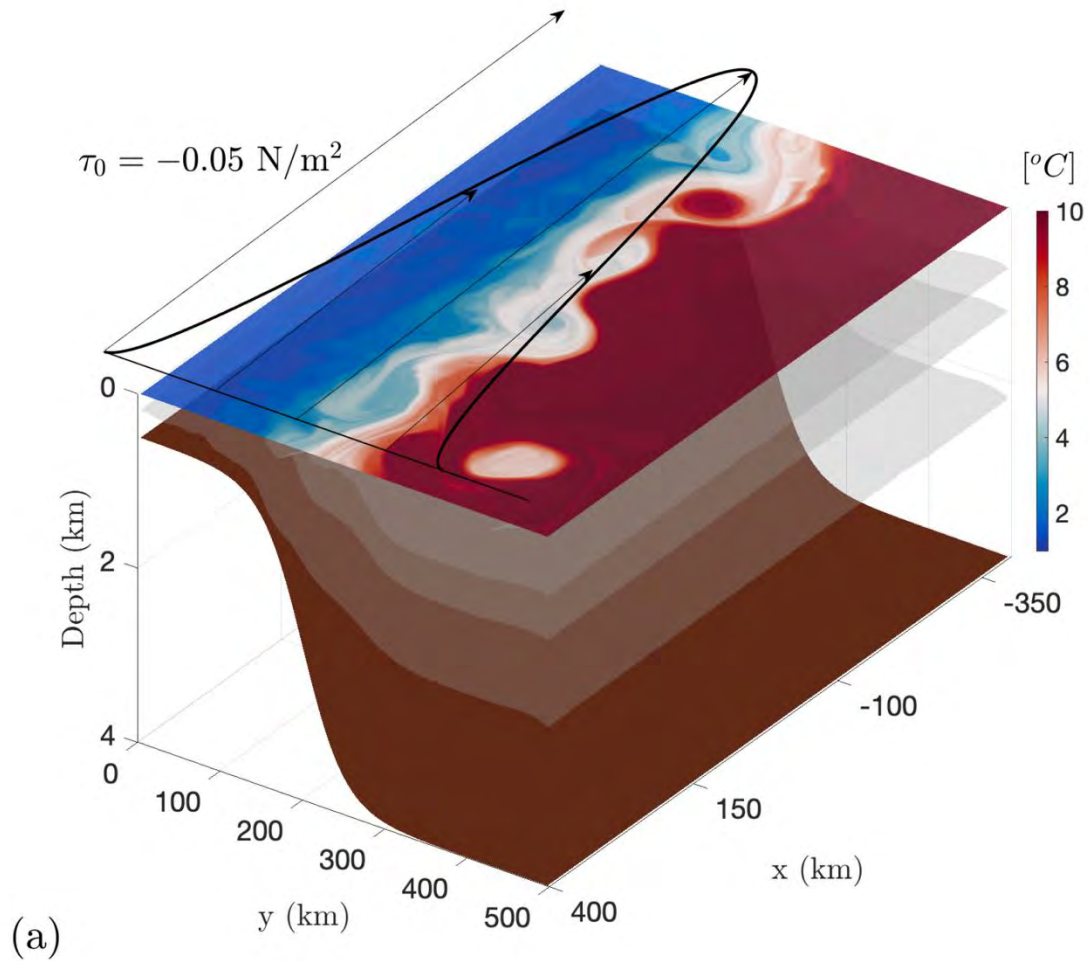
Eddies across continental margins



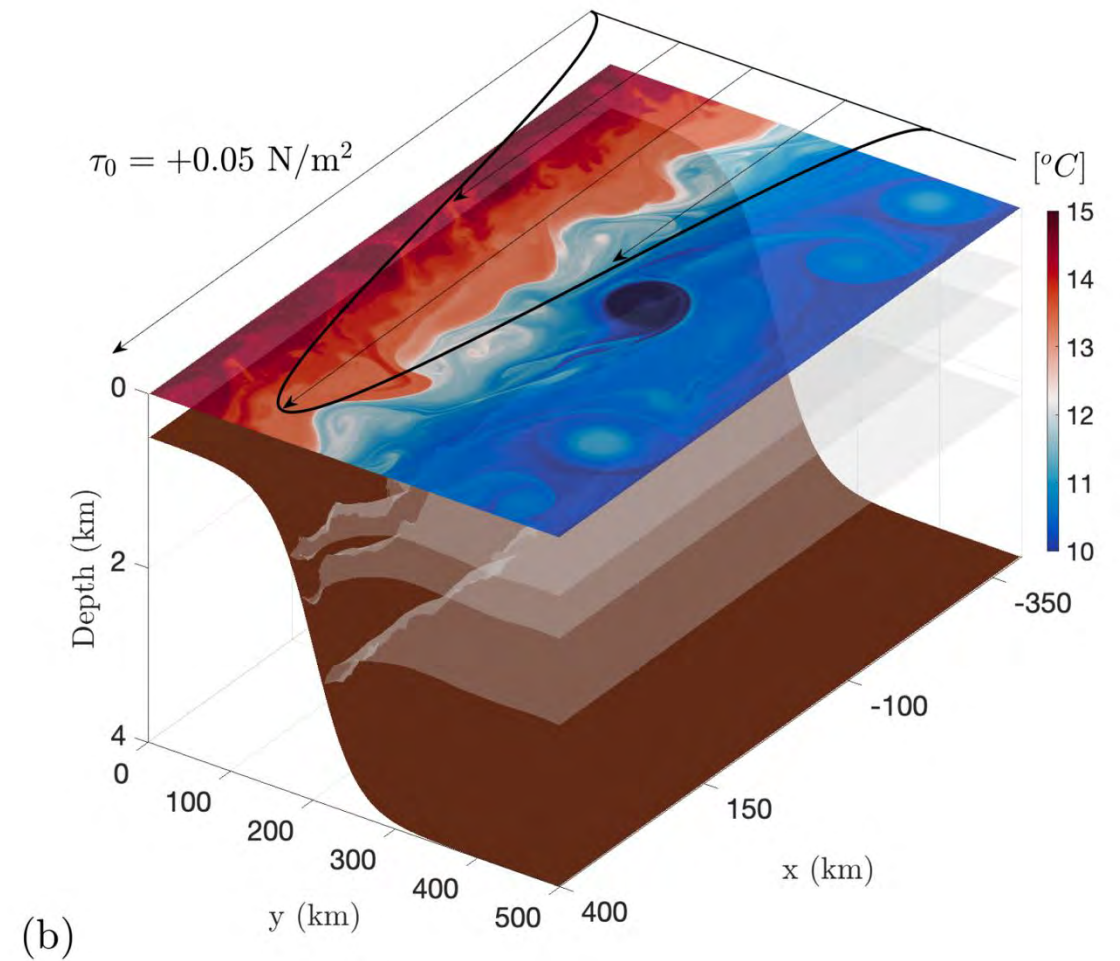
Eddies across continental margins



Retrograde versus Prograde

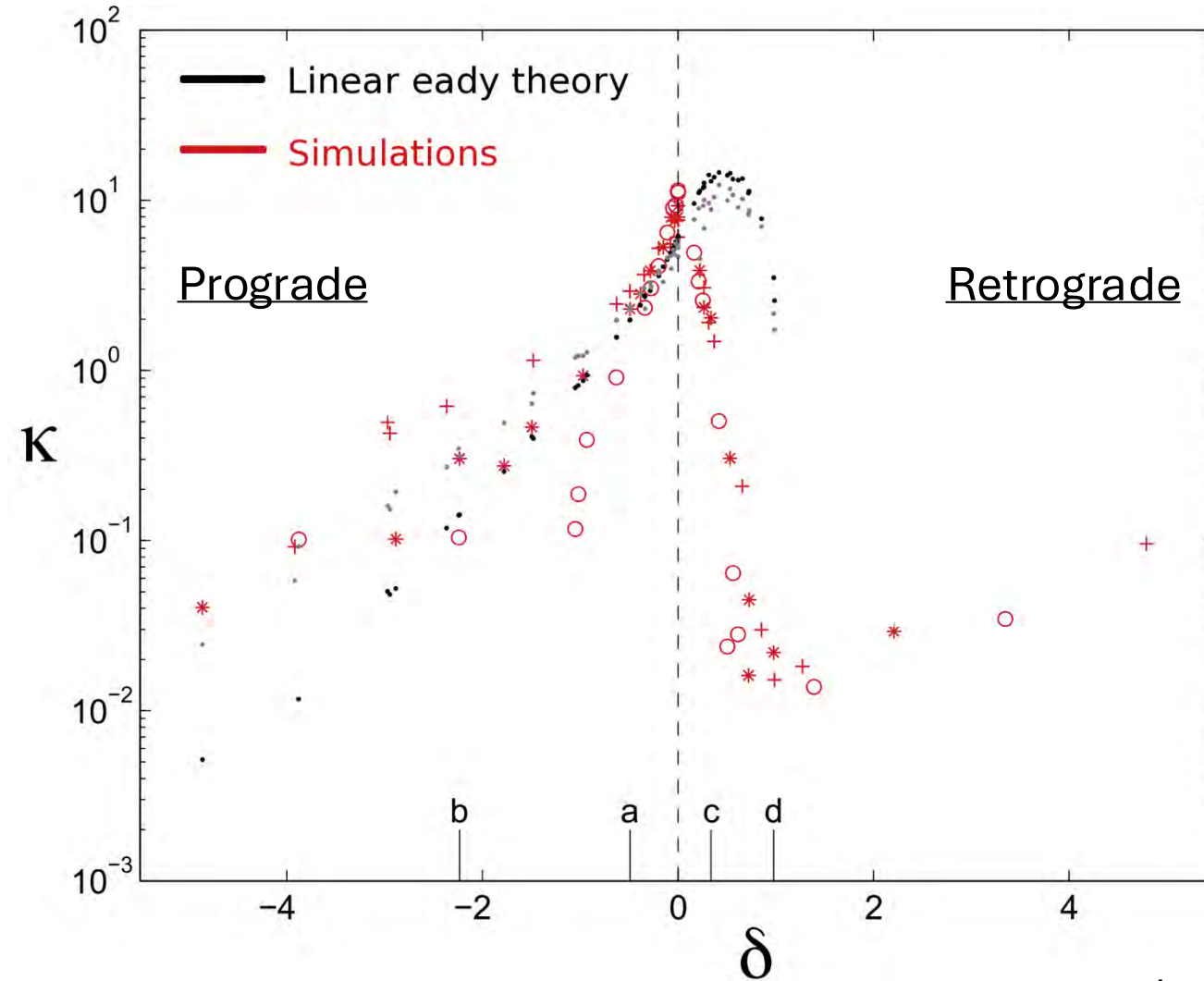


Retrograde or upwelling

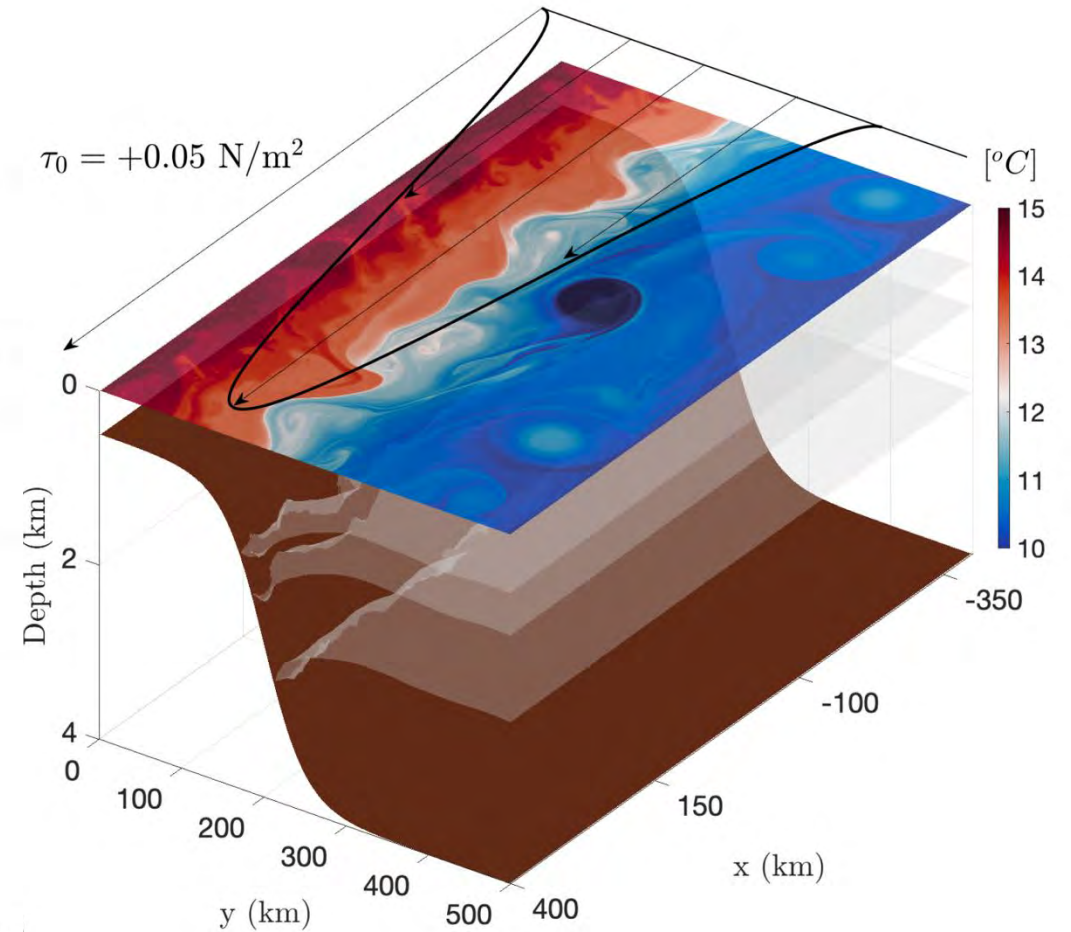
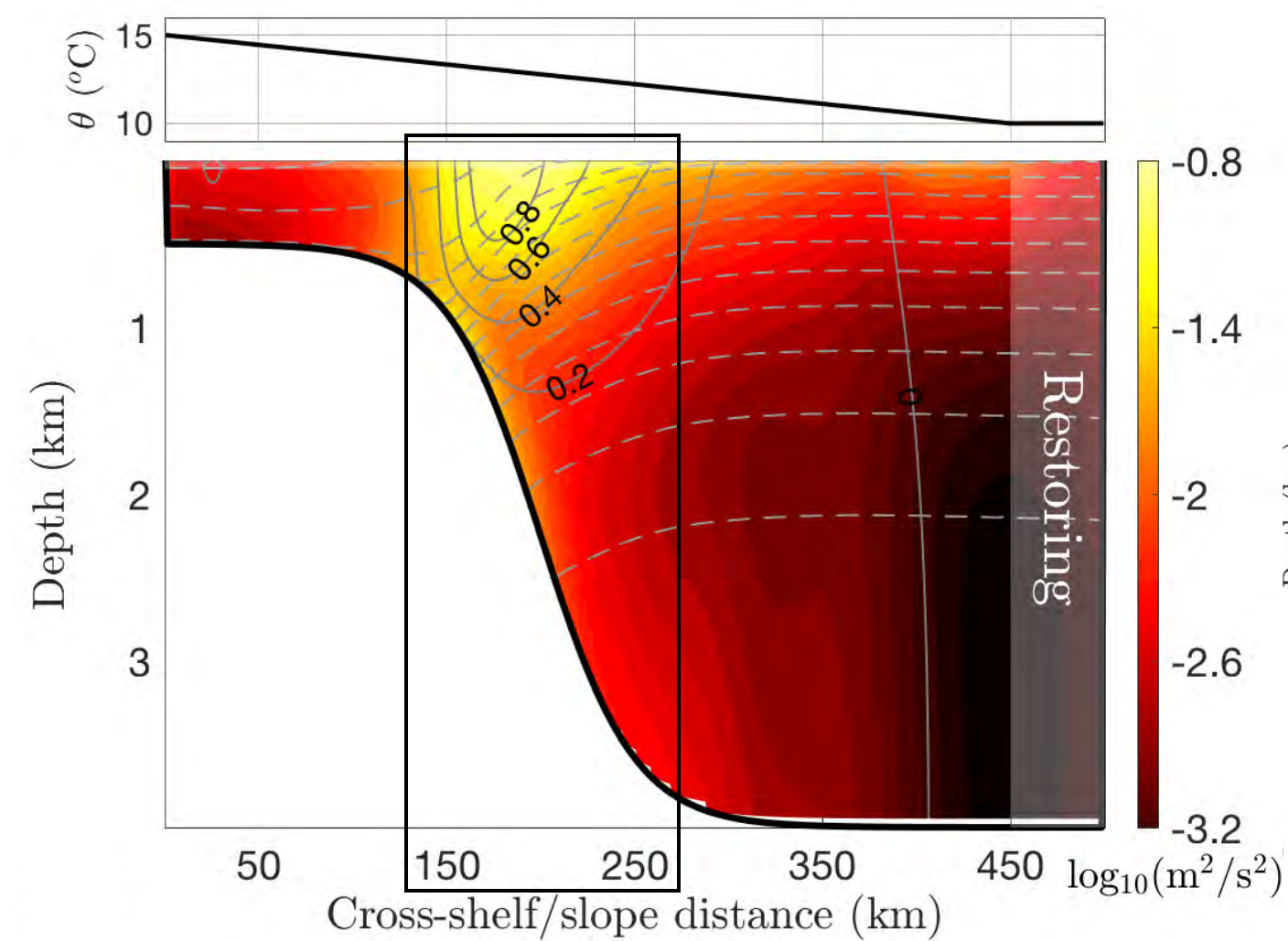


Prograde or downwelling

Retrograde versus Prograde



Eddy and mean properties: Prograde



Prograde or downwelling

Scaling of eddy buoyancy diffusivity: Prograde

$$K_{\text{GEOM}} = \alpha \frac{\sqrt{Ri}}{f_0} E,$$

The “GEOMETRIC” theory
(e.g. Marshall et al. 2012; Mak et al. 2022)

Scaling of eddy buoyancy diffusivity: Prograde

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Nonlinear eddy growth rate

$$\frac{f_0^2}{Ri} \cdot K_{\text{GEOM}} = \alpha \cdot \boxed{\sigma_E} \cdot E \equiv \boxed{2\sigma} \cdot E,$$

Normalized Eady growth rate: $\sigma_E = f_0 / \sqrt{Ri}$

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$$\sigma / [0.31 \cdot \sigma_E] \sim \mathcal{F}_{\text{GEOM}}(S)$$

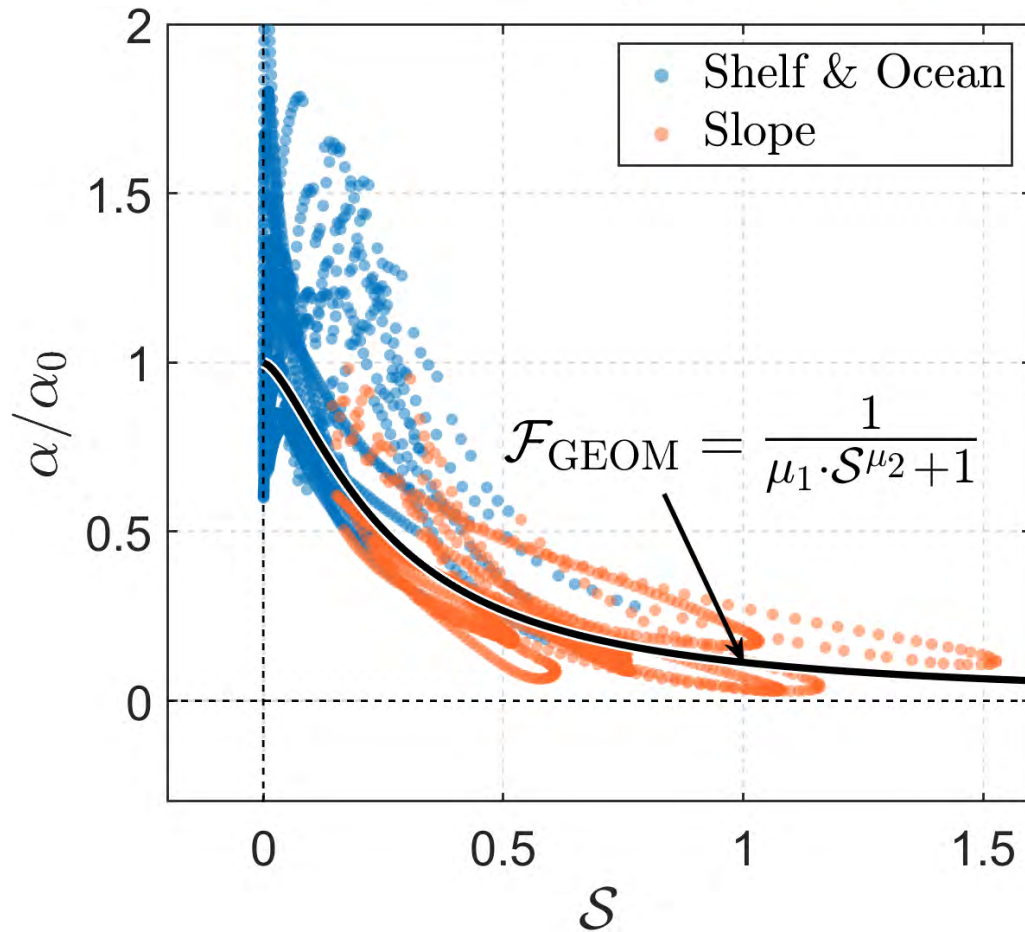
$$S = \left\langle \left| \frac{\partial H}{\partial y} \right| \frac{1}{|H|} \int_{-|H|}^0 \frac{N}{f_0} dz \right\rangle$$

The slope Burger number dependence
(e.g. Brink, 2012; 2016; Brink and Cherian, 2013; Hetland, 2017; Chen et al, 2020).

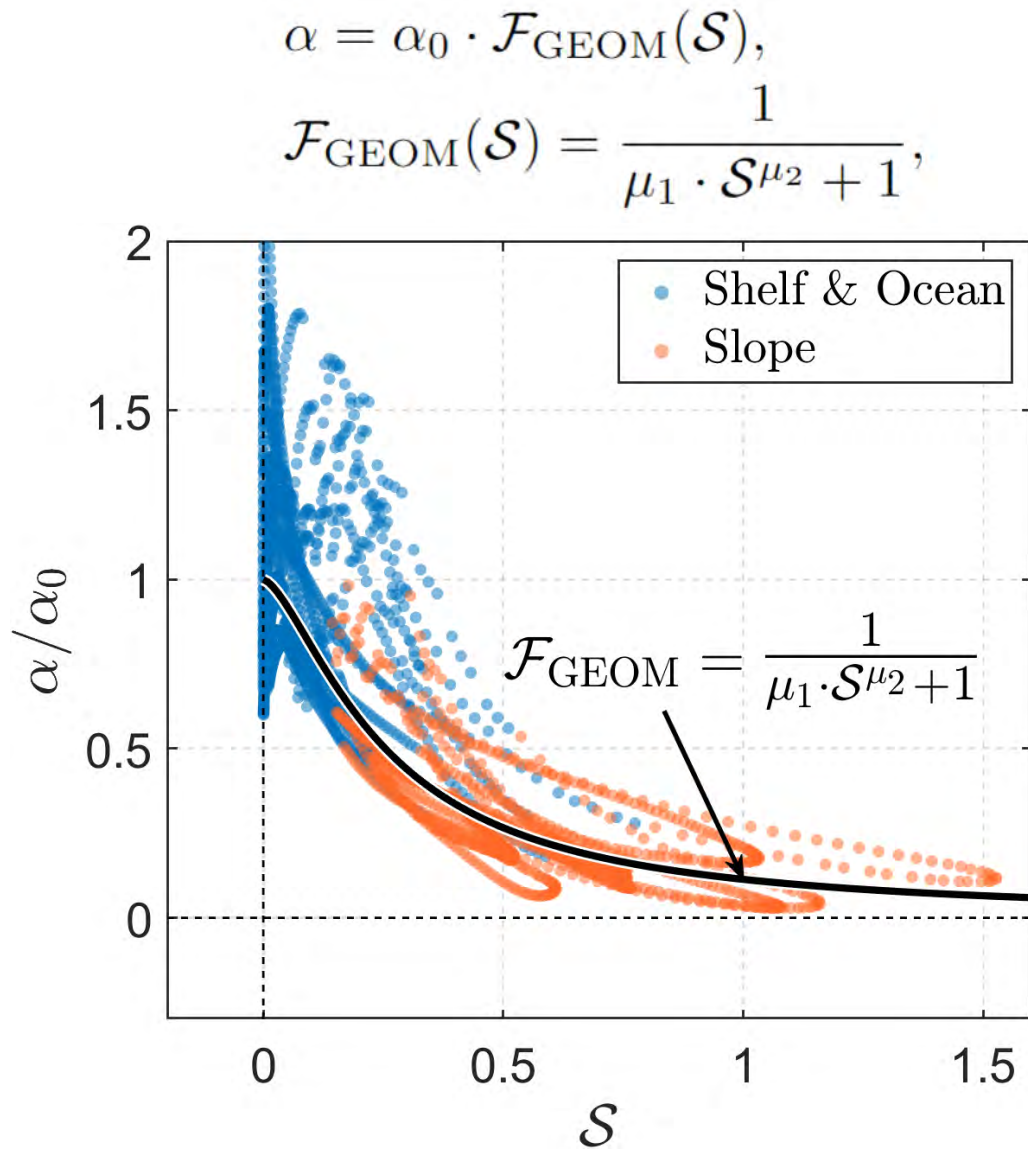
Scaling of eddy buoyancy diffusivity: Prograde

$$\alpha = \alpha_0 \cdot \mathcal{F}_{\text{GEOM}}(\mathcal{S}),$$

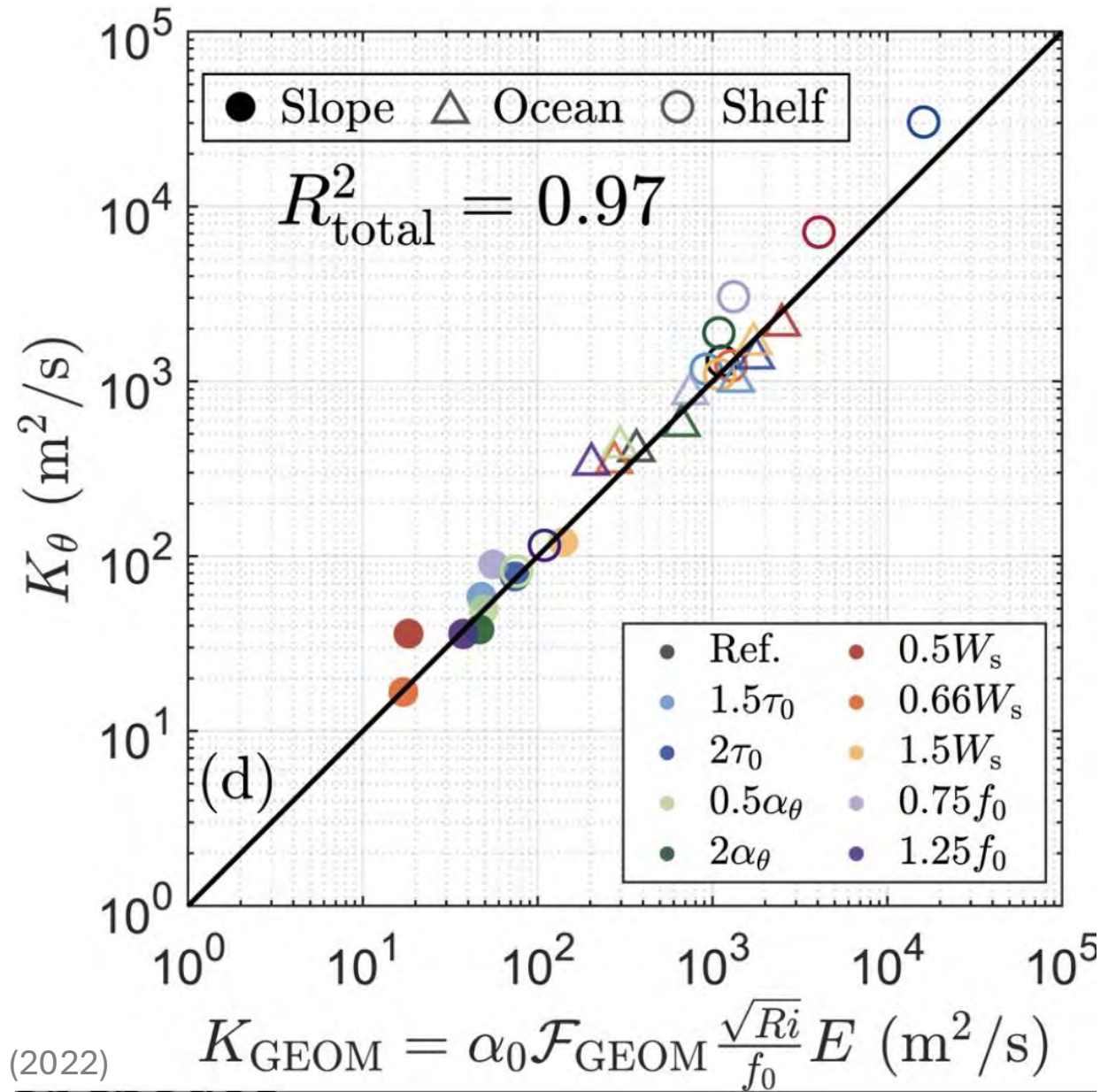
$$\mathcal{F}_{\text{GEOM}}(\mathcal{S}) = \frac{1}{\mu_1 \cdot \mathcal{S}^{\mu_2} + 1},$$



Scaling of eddy buoyancy diffusivity: Prograde



Wei, Wang, Mak, and Stewart (2022)



Convert scaling into eddy closure: Prograde

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial (\langle \mathbf{v} \rangle \hat{E})}{\partial y}}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}},$$

Convert scaling into eddy closure: Prograde

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \underbrace{\mathcal{K}_{\text{gm}} \frac{M^4}{N^2}}_{\text{source}} dz}_{\text{source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2}{\partial y^2} \hat{E}}_{\text{diffusion}},$$

$$\mathcal{K}_{\text{GEOM}}^{\text{slope}} = \alpha_{\text{geom}} \frac{\hat{E}}{\int_{-|H|}^0 \sigma_E dz},$$

$$\alpha_{\text{geom}} = \alpha_0 \mathcal{F}_{\text{GEOM}} = \alpha_0 \frac{1}{\mu_1 S^{\mu_2} + 1},$$

Convert scaling into eddy closure: Prograde

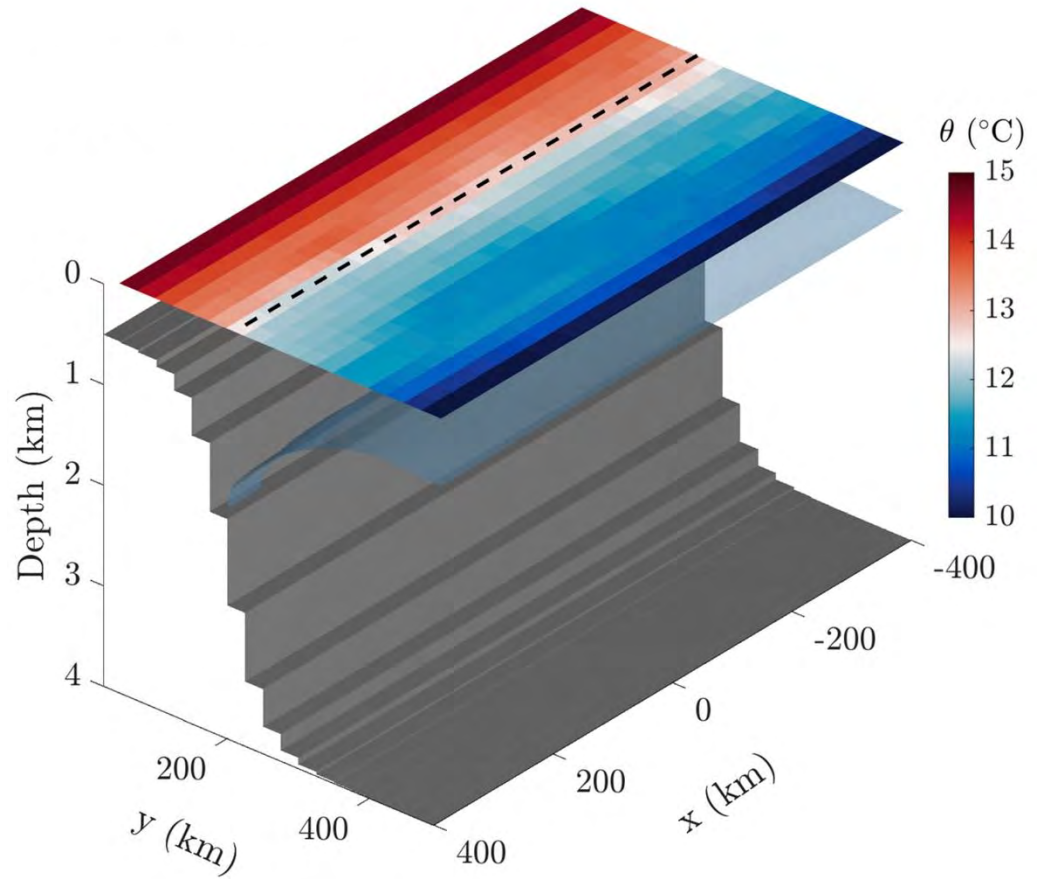
$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \underbrace{\kappa_{\text{gm}} \frac{M^4}{N^2}}_{\text{source}} dz}_{\text{source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}},$$

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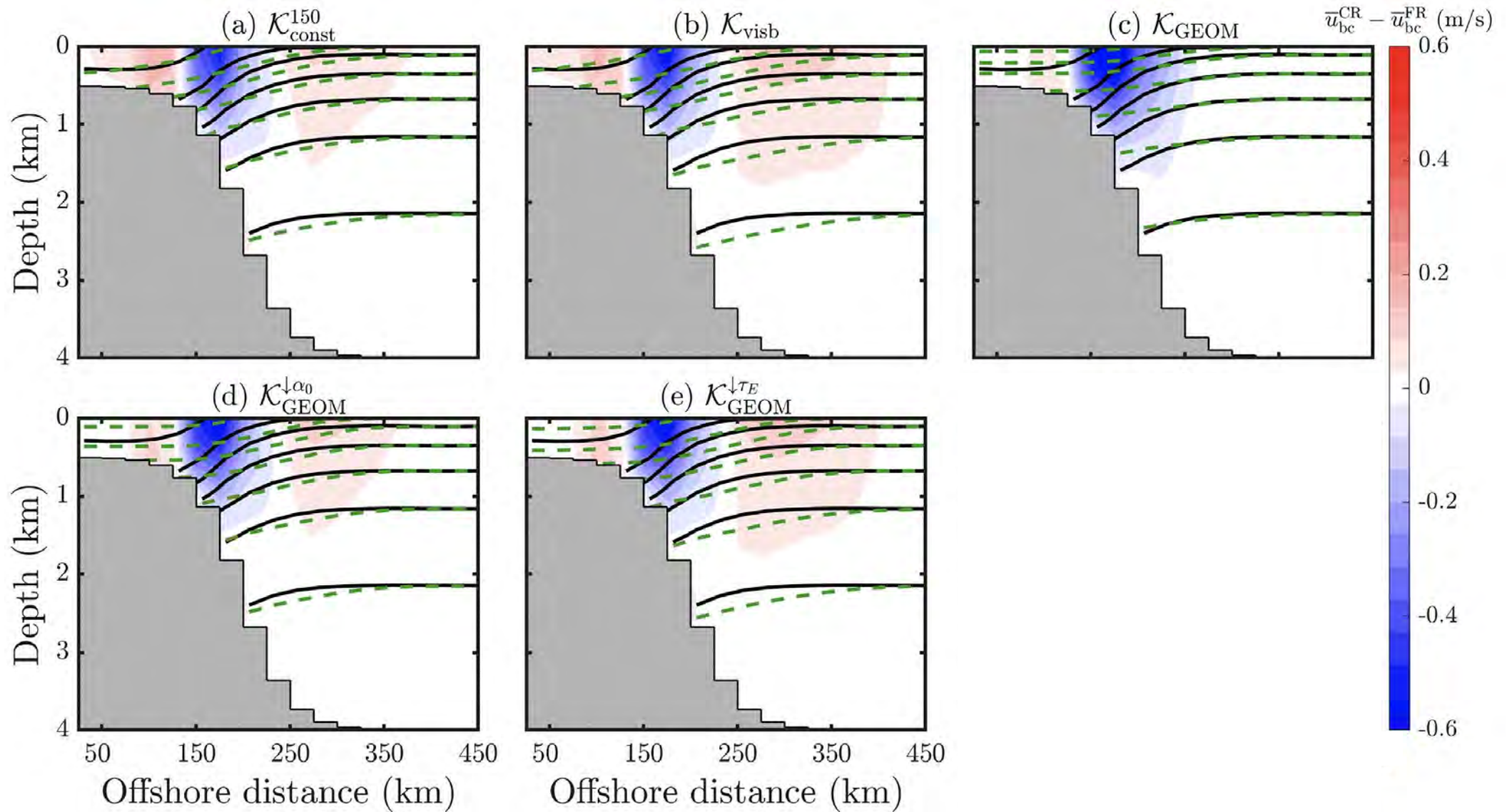
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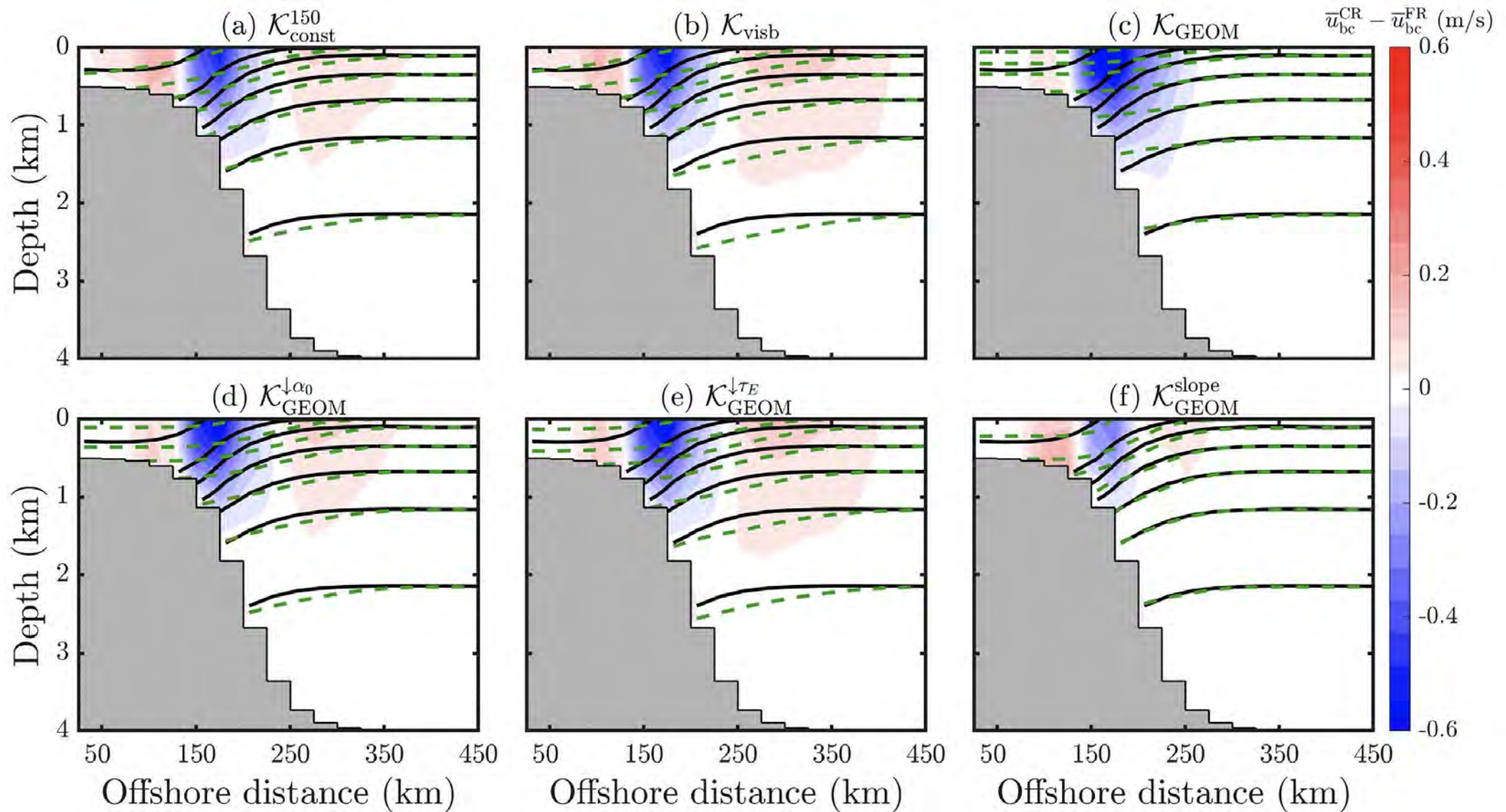
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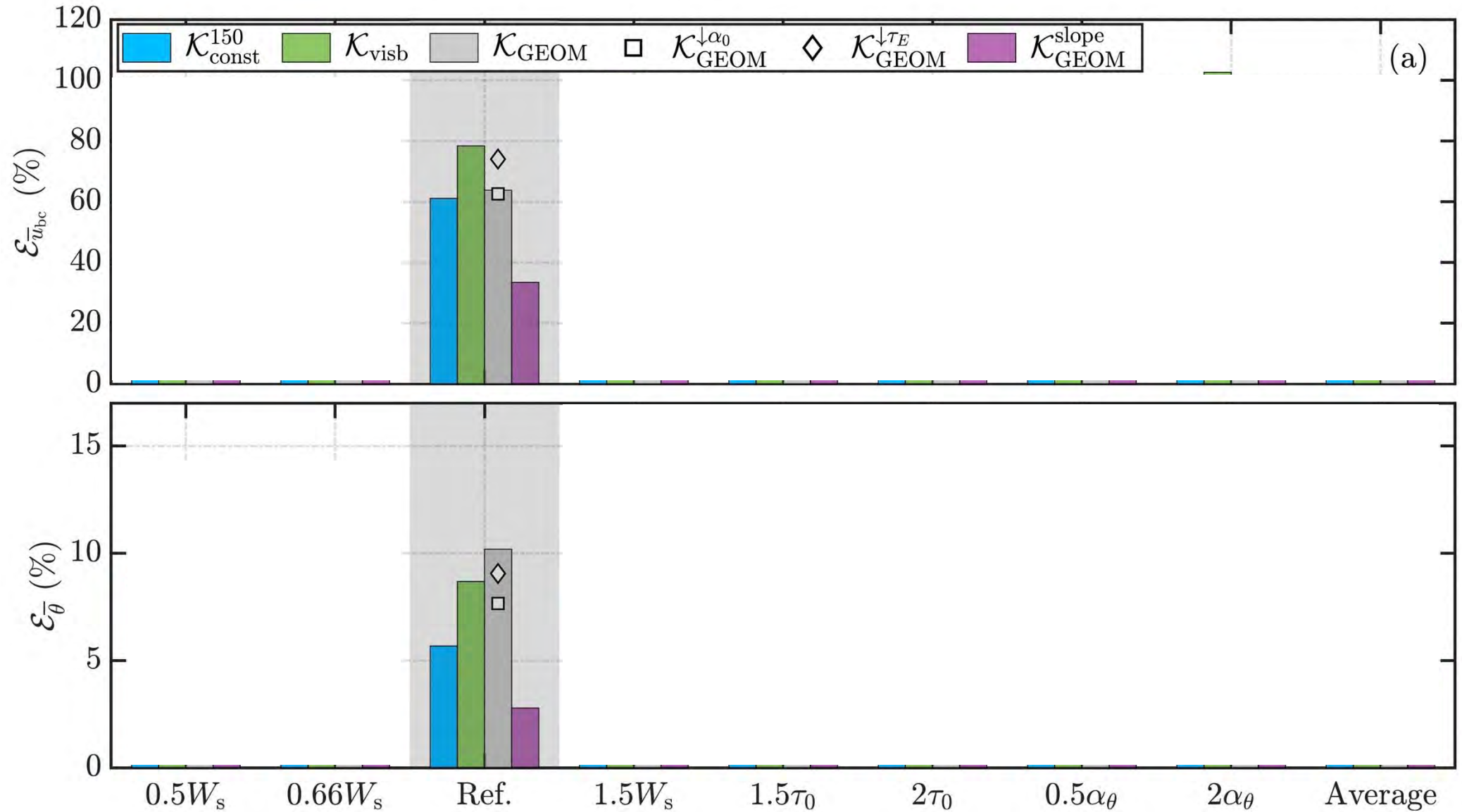
Parameterization: Prograde



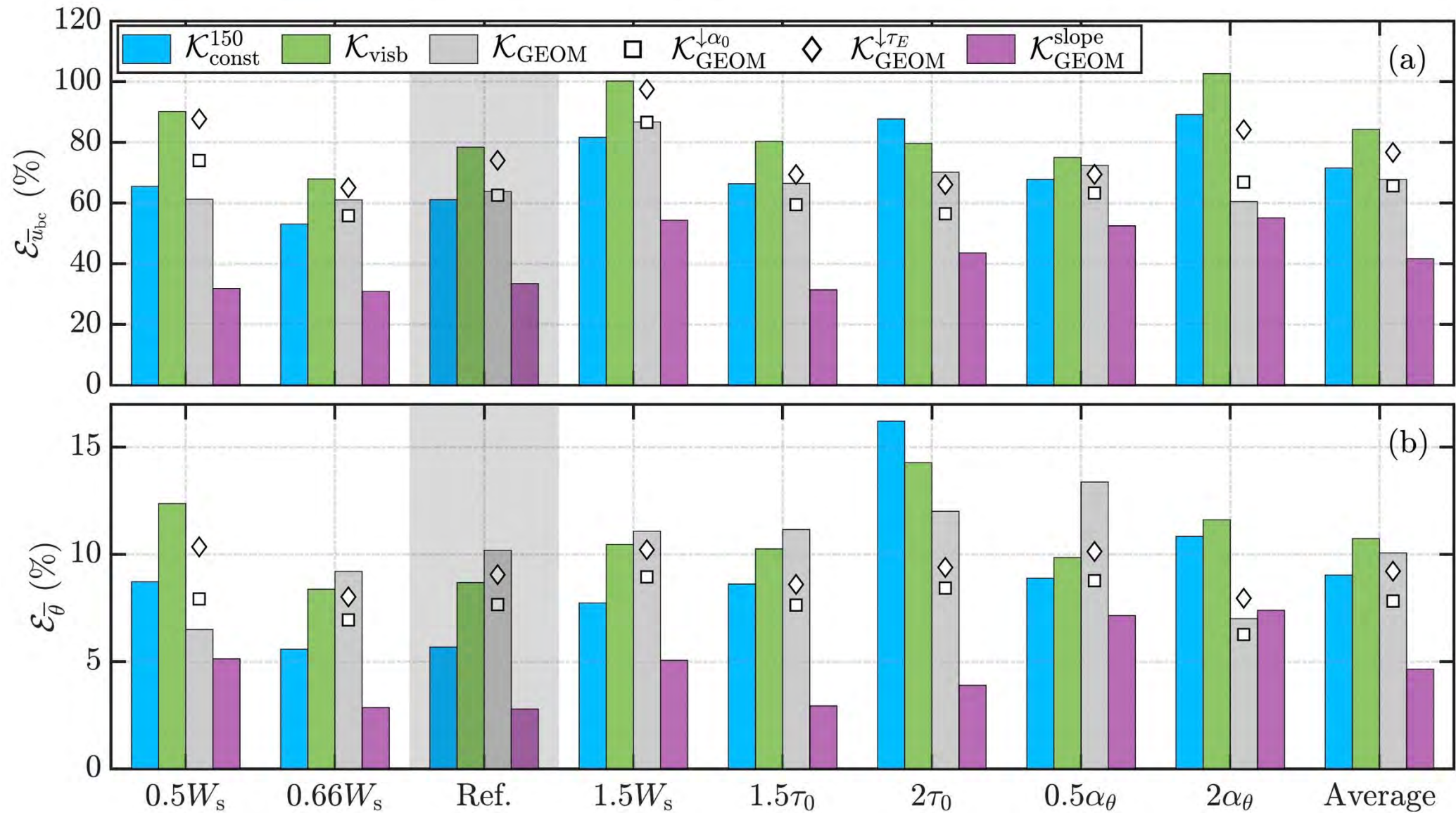
Parameterization: Prograde



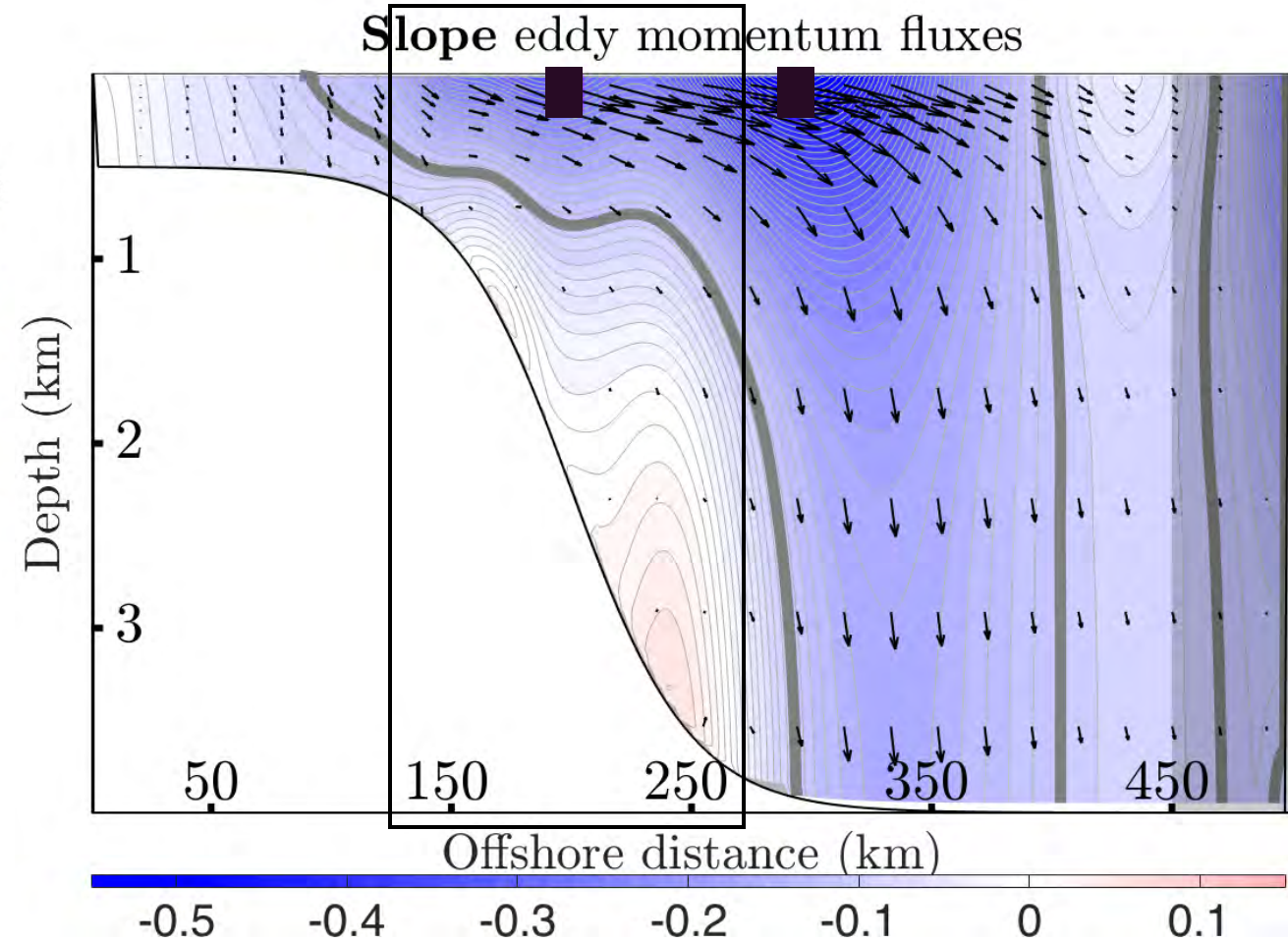
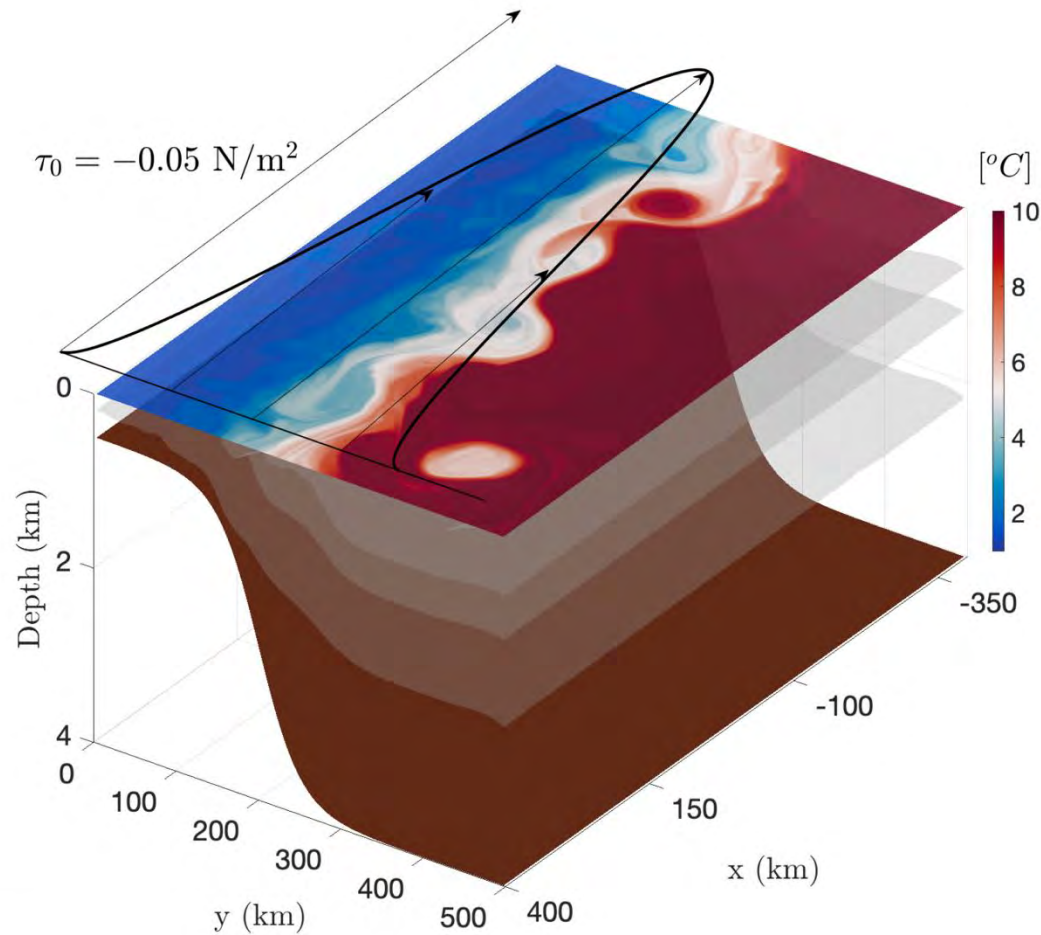
Parameterization: Prograde



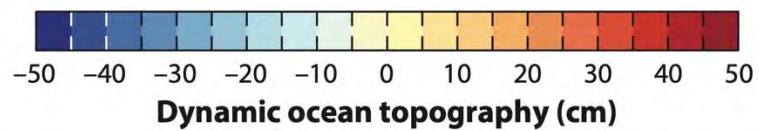
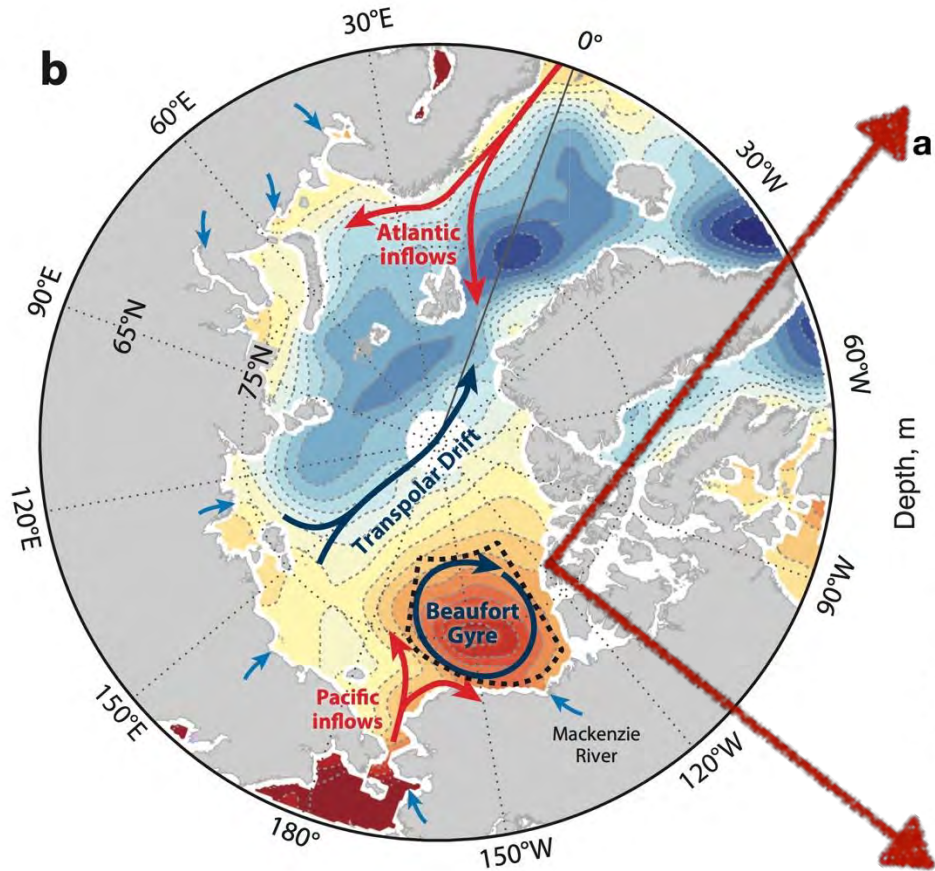
Parameterization: Prograde



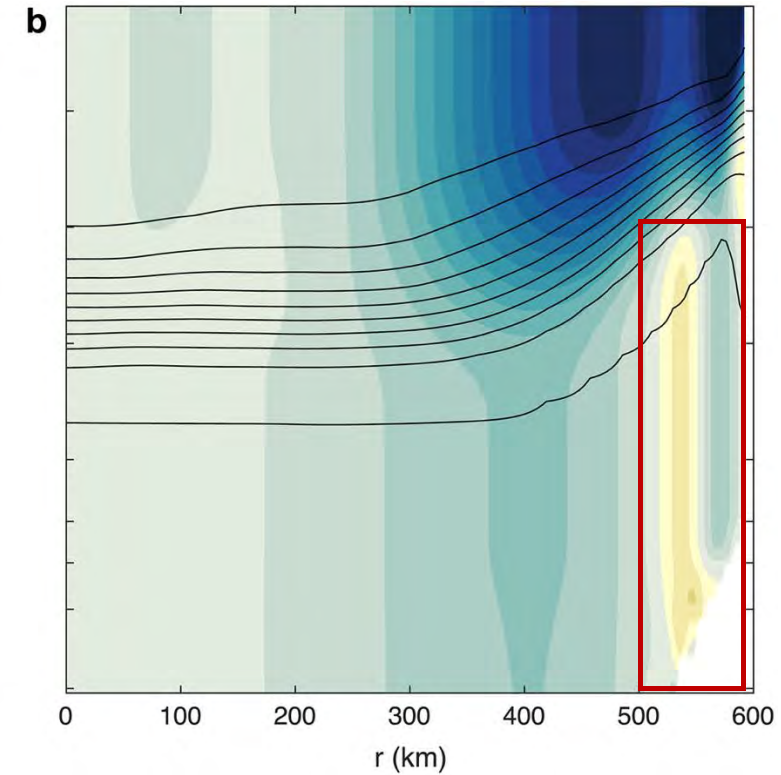
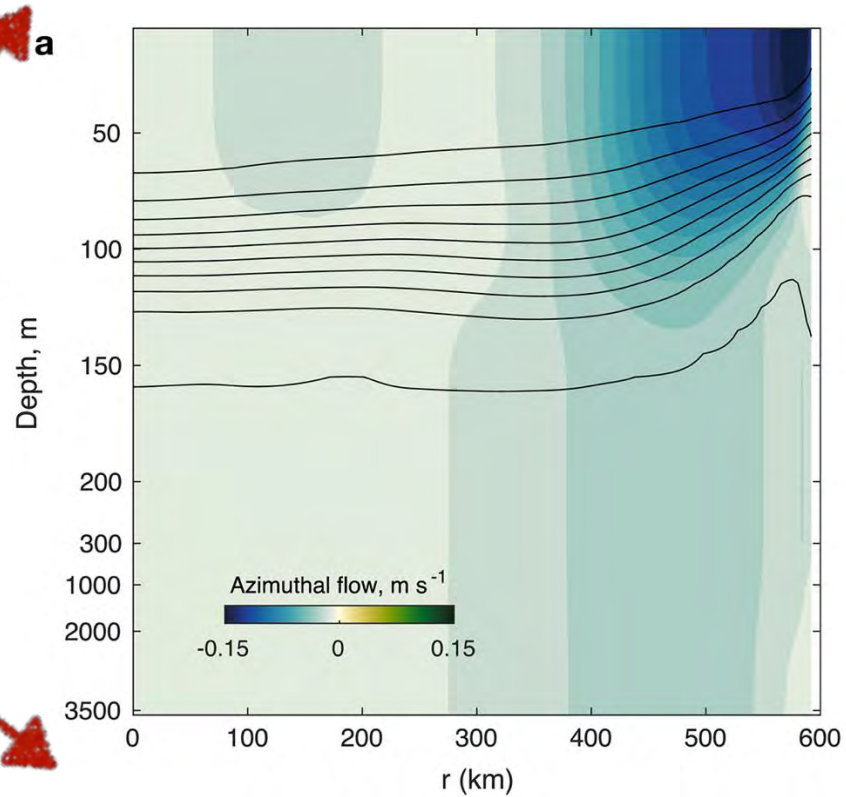
Eddy impact on the mean: Retrograde



Eddy impact on the mean: Retrograde

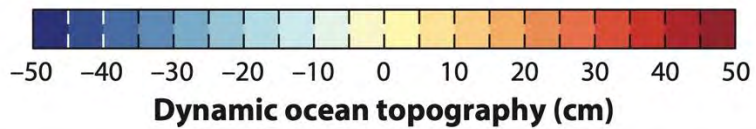
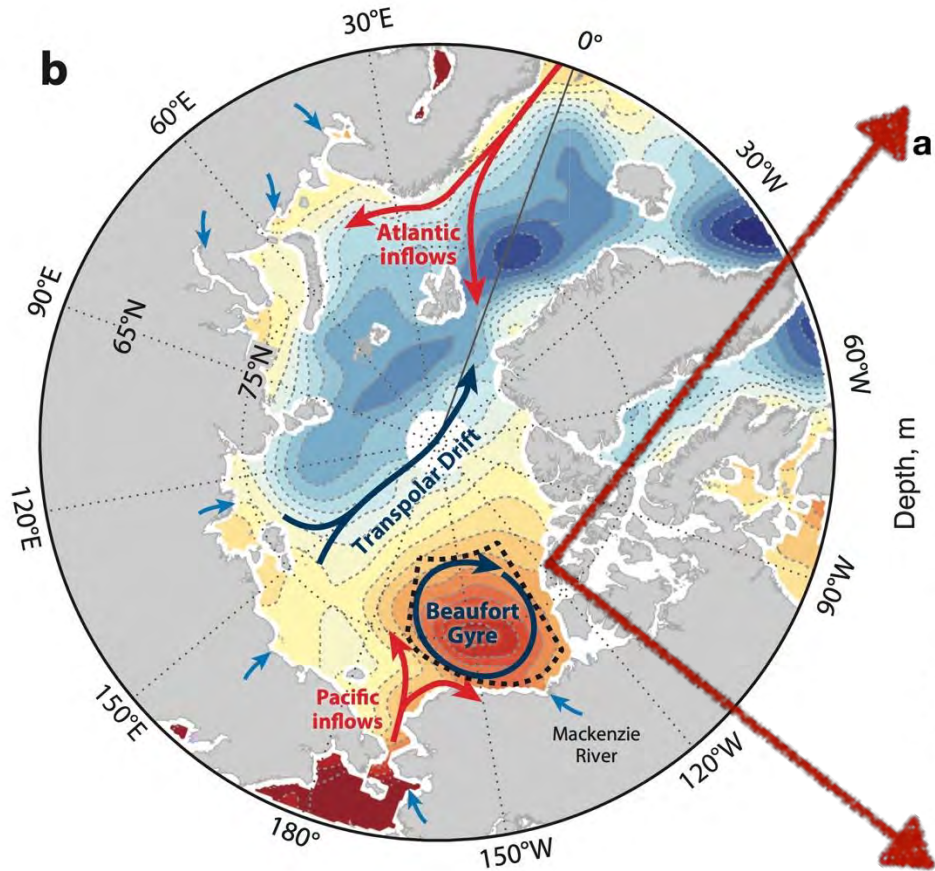


Timmermans and Toole (2019)



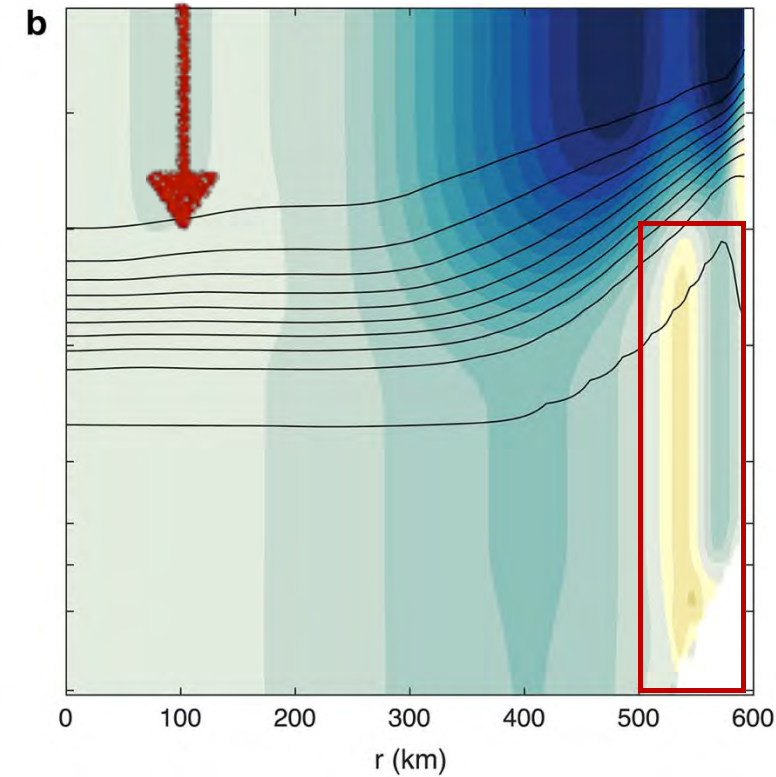
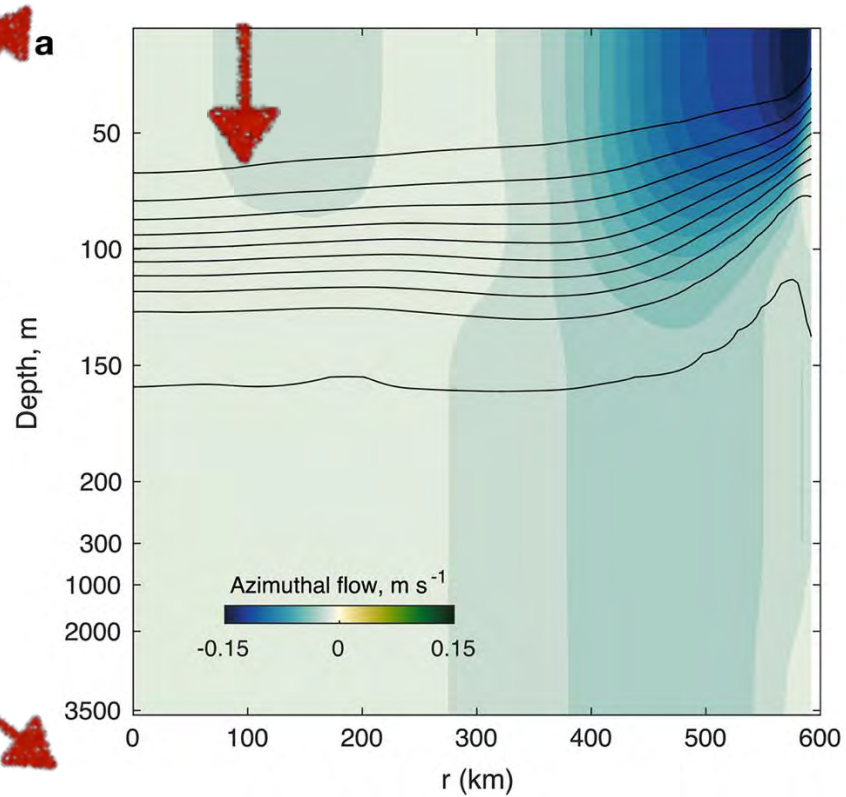
Manucharyan and Isachsen (2019)

Eddy impact on the mean: Retrograde



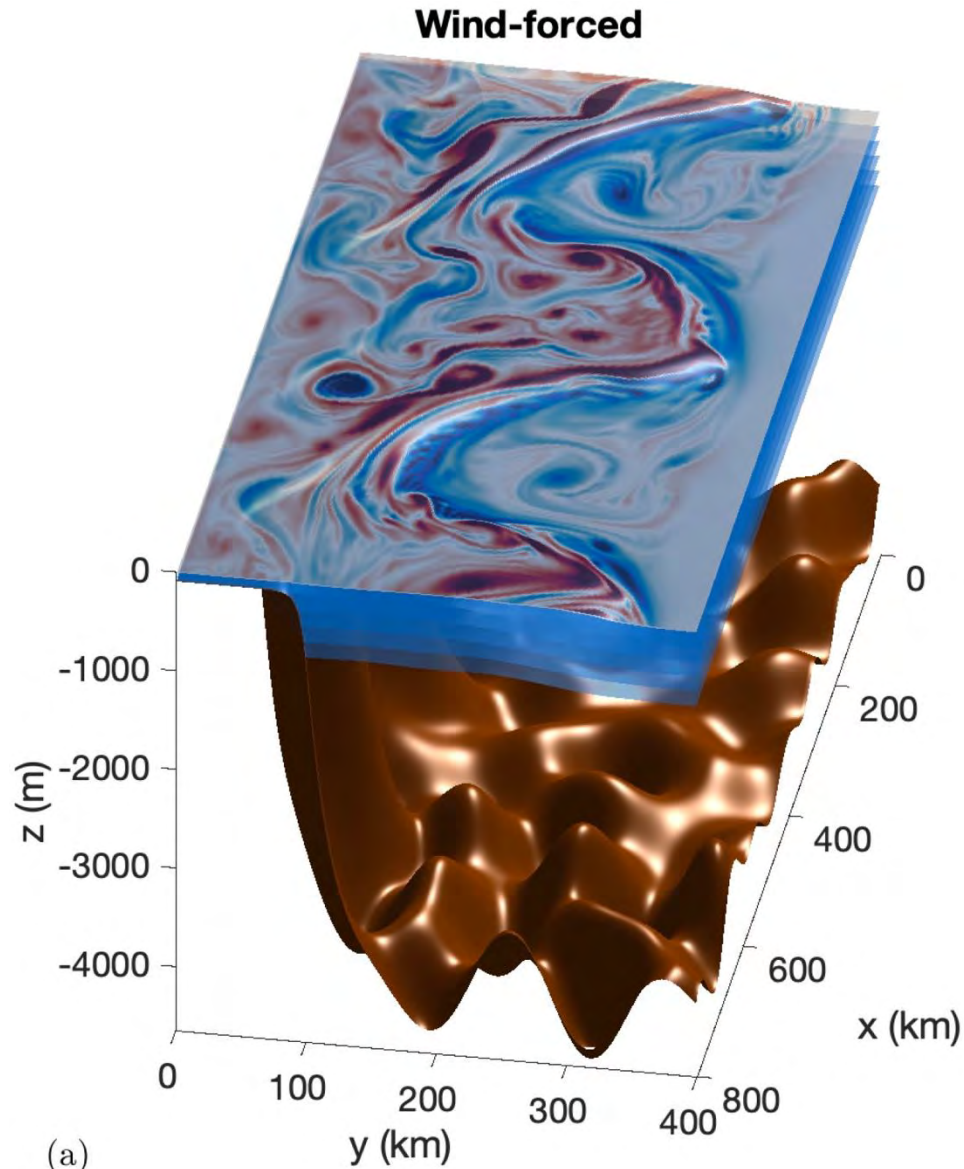
Timmermans and Toole (2019)

Freshwater content difference

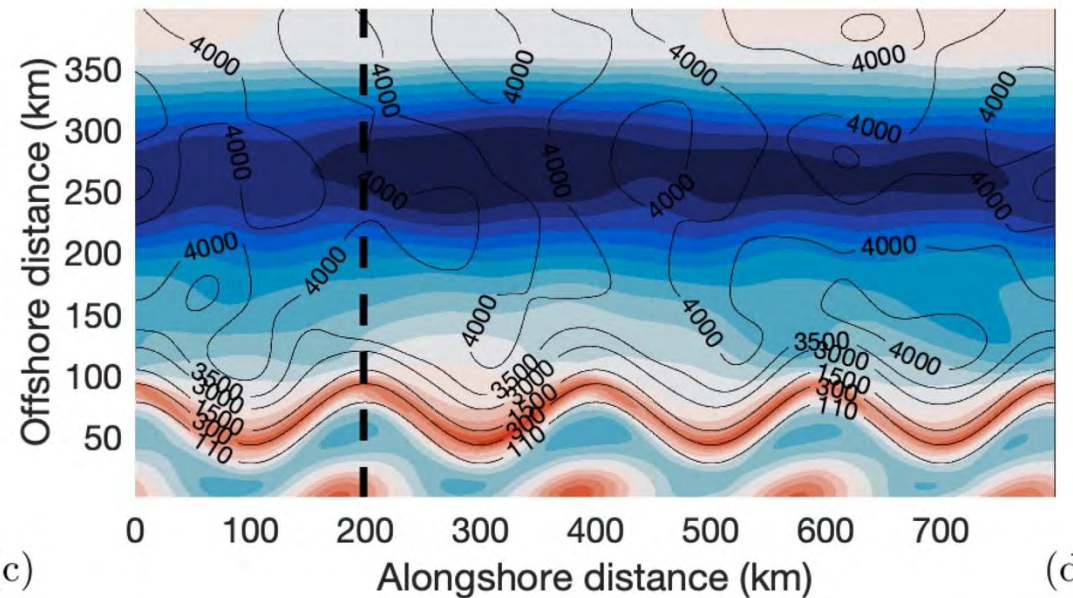
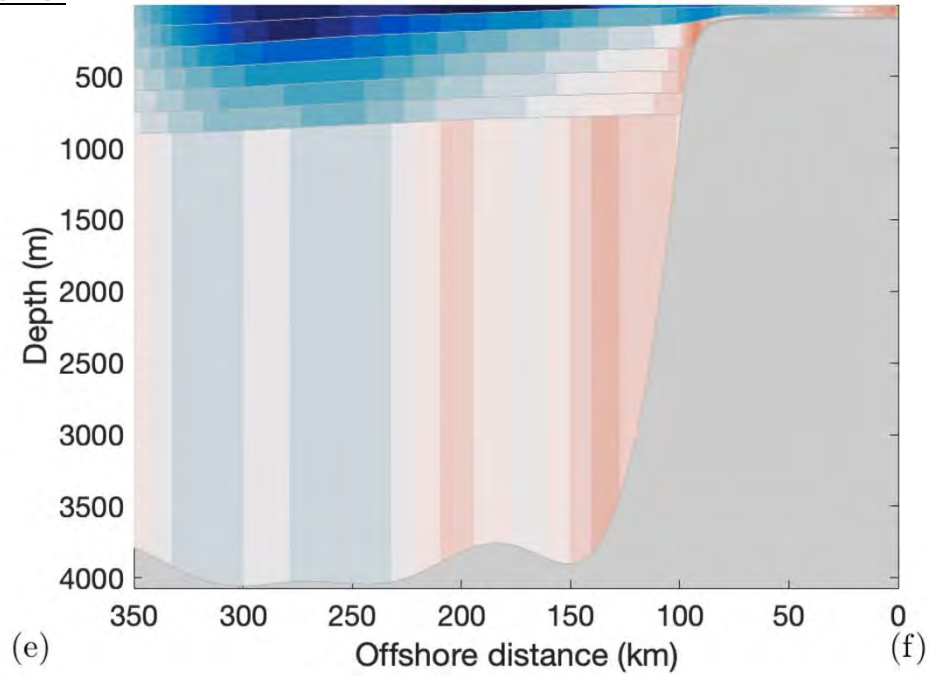


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Eddy impact on the mean: Retrograde

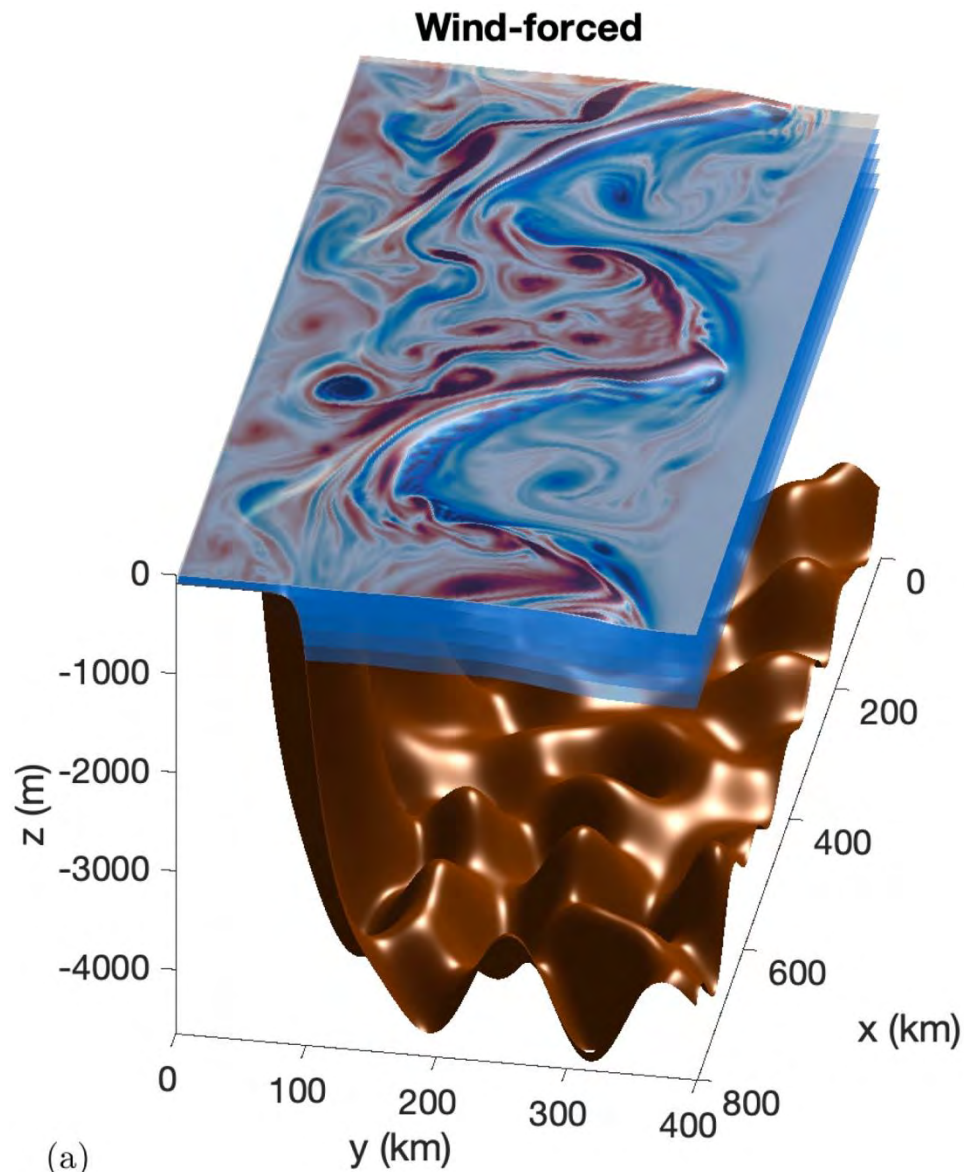


Stewart et al. (2024)

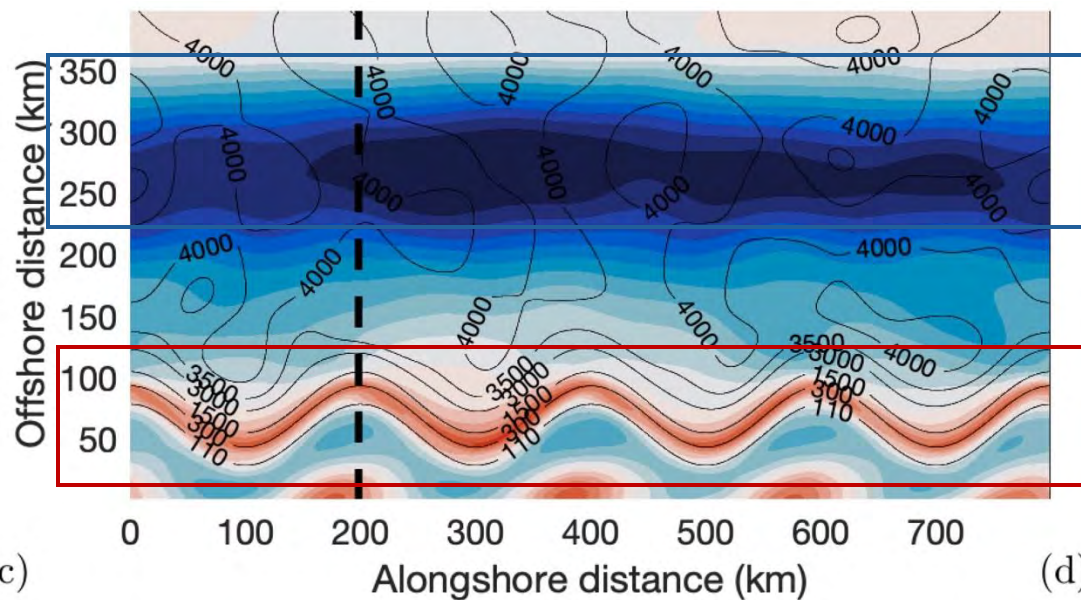
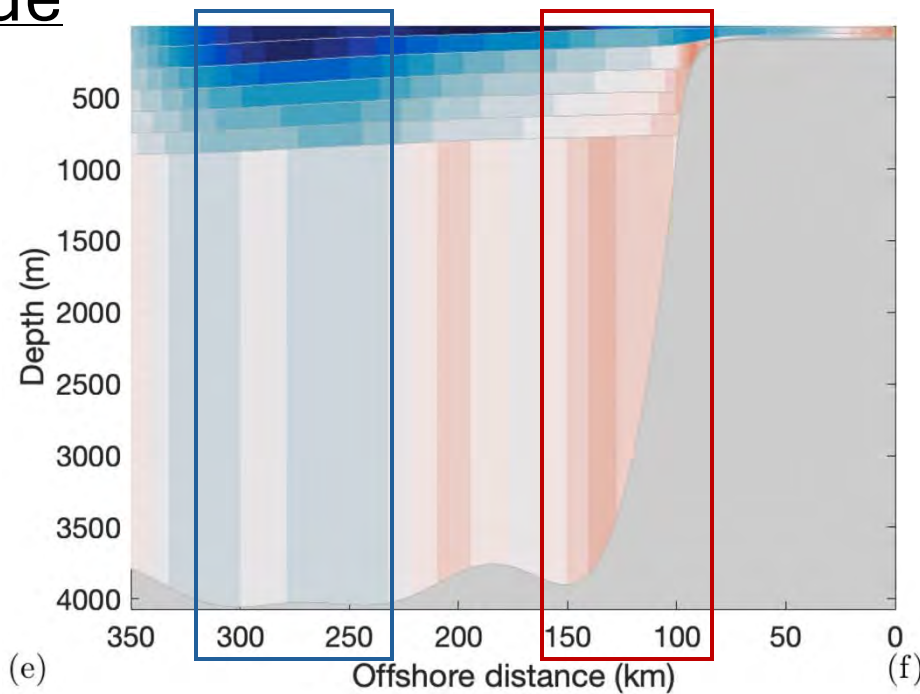


(c)

Eddy impact on the mean: Retrograde



Stewart et al. (2024)



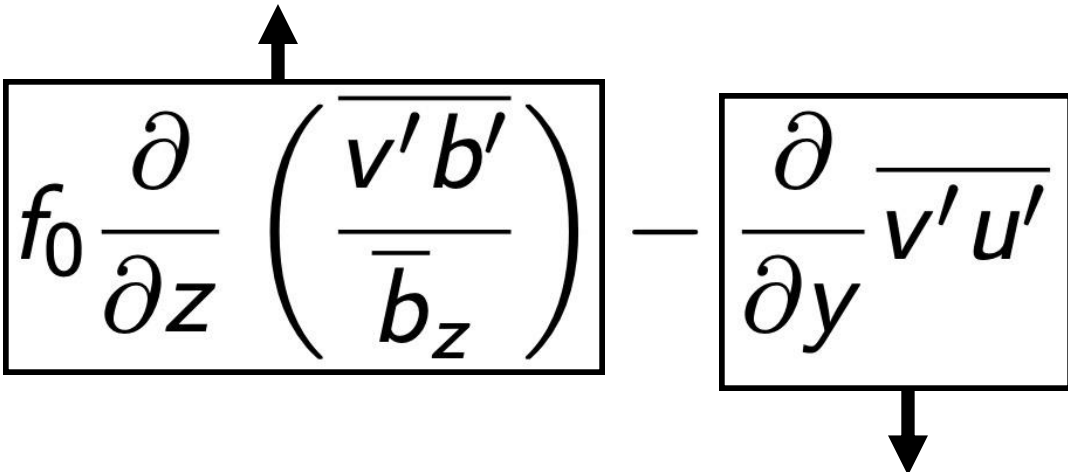
(d)

Eddy impact on the mean: Retrograde

$$\frac{\partial}{\partial t} \bar{u} \simeq \text{Forcing} + f_0 \frac{\partial}{\partial z} \left(\frac{\overline{v' b'}}{\bar{b}_z} \right) - \frac{\partial}{\partial y} \overline{v' u'}$$

Eddy impact on the mean: Retrograde

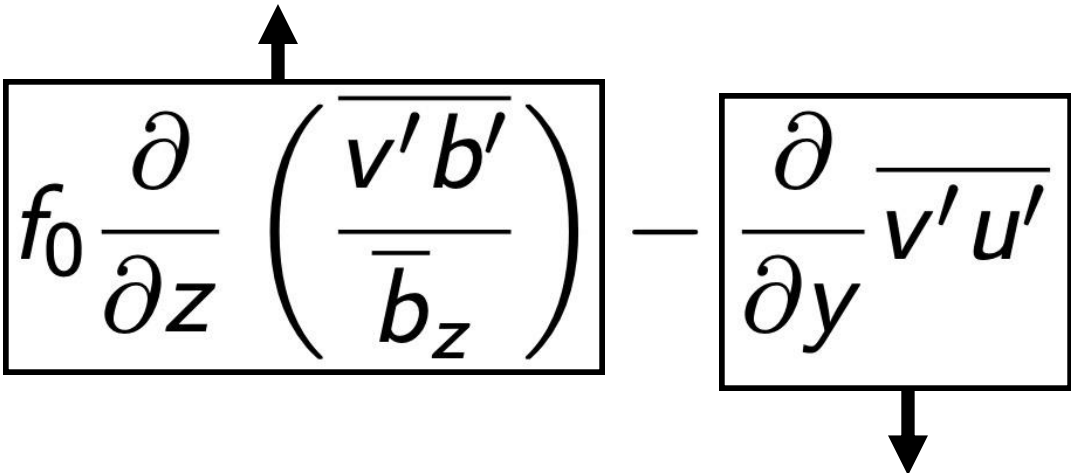
Eddy form stress transfers momentum *downward*

$$\frac{\partial}{\partial t} \bar{u} \simeq \text{Forcing} + \boxed{f_0 \frac{\partial}{\partial z} \left(\frac{\overline{v' b'}}{\bar{b}_z} \right)} - \boxed{\frac{\partial}{\partial y} \overline{v' u'}}$$


Eddy Reynolds stress transfers momentum *offshore*

Eddy impact on the mean: Retrograde

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Depth-integral gives (under equilibrium)

Eddy Reynolds stress transfers momentum *offshore*

$$\psi_{\text{EMF}} + \psi_{\text{GM}} \simeq \psi_{\text{Ekman}} + \psi_{\text{Residual}}$$

Eddy impact on the mean: Retrograde

Eddy form stress transfers momentum *downward*

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Eddy Reynolds stress transfers momentum *offshore*

Depth-integral gives (under equilibrium)

$$\frac{1}{f_0} \int_{\hat{z}}^0 \frac{\partial}{\partial y} \overline{v' u'} dz$$

$$\psi_{\text{EMF}}$$

Jet-shifter
(e.g. Stewart et al, 2024)

Two-way balance of the ACC
(e.g. Marshall and Radko, 2003)

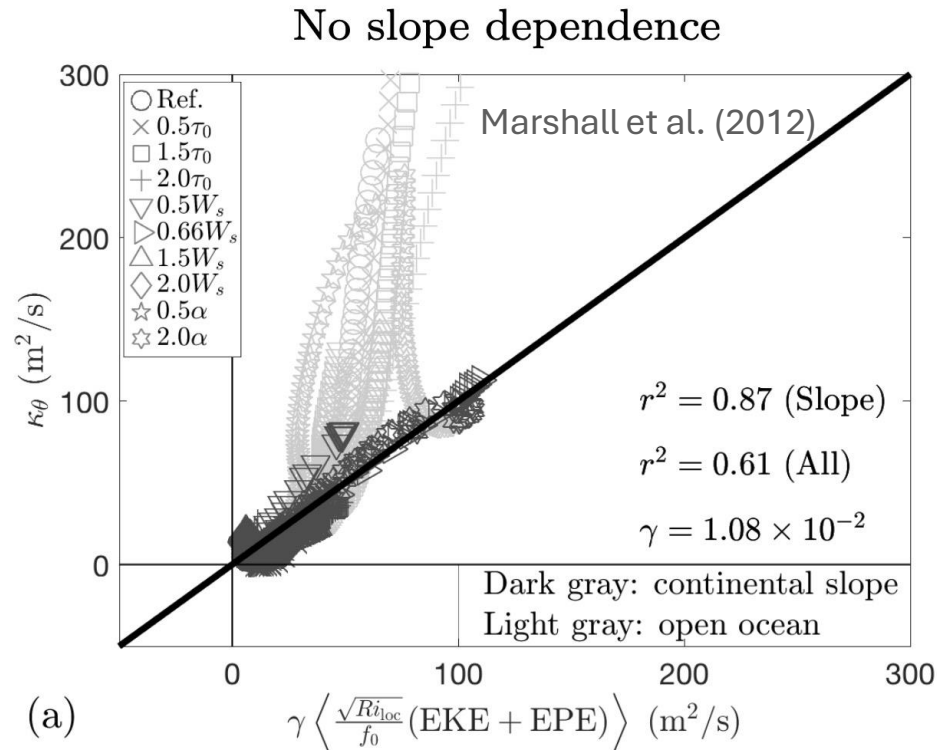
$$\psi_{\text{GM}} \simeq \psi_{\text{Ekman}}$$

$$\left. \frac{\overline{v' b'}}{\bar{b}_z} \right|_{z=\hat{z}} \quad \frac{\tau}{\rho_0 f_0}$$

$$+ \psi_{\text{Residual}}$$

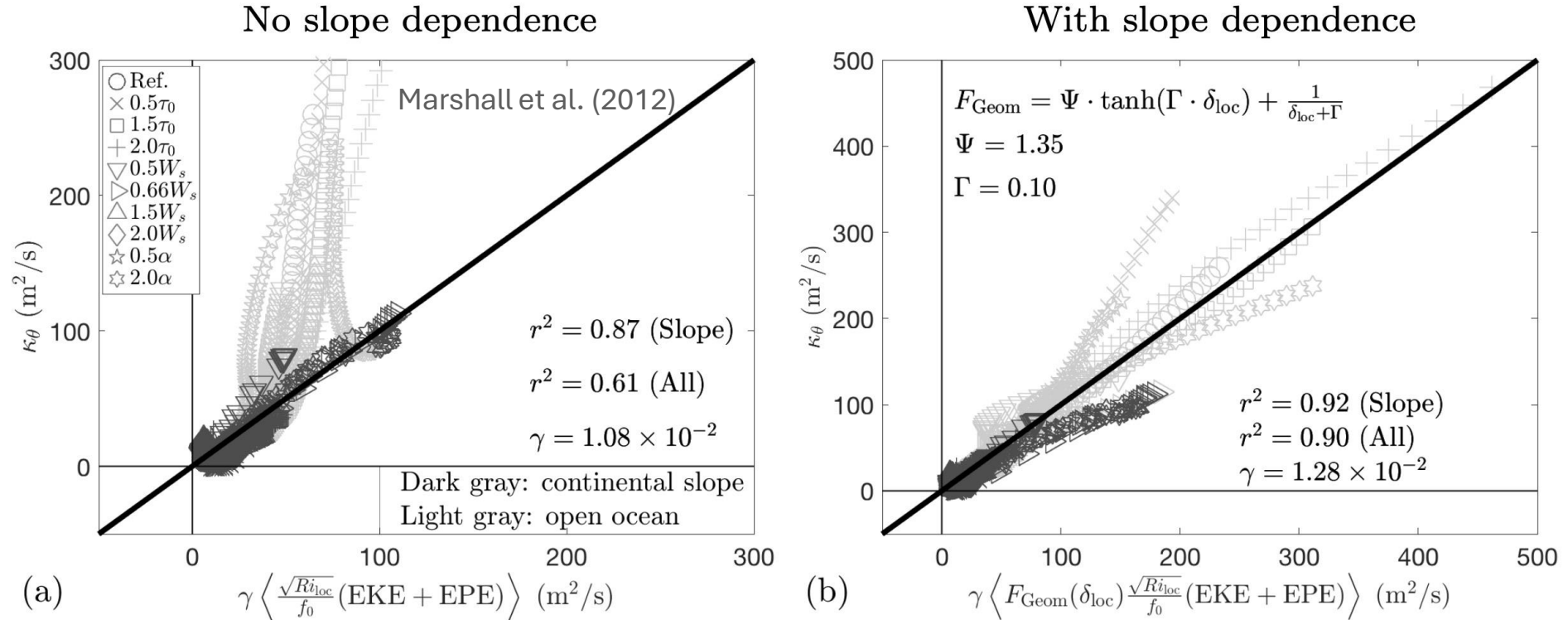
Presumably adiabaticity

Scaling(s) of eddy buoyancy diffusivity: Retrograde



$$\kappa_{\text{Geom}} = \gamma_{\text{Geom}} \frac{\sqrt{Ri_{loc}}}{f_0} (\text{EKE} + \text{EPE}),$$

Scaling(s) of eddy buoyancy diffusivity: Retrograde



(a) Original and (b) slope-aware forms of the GEOMETRIC scaling vs κ_θ .

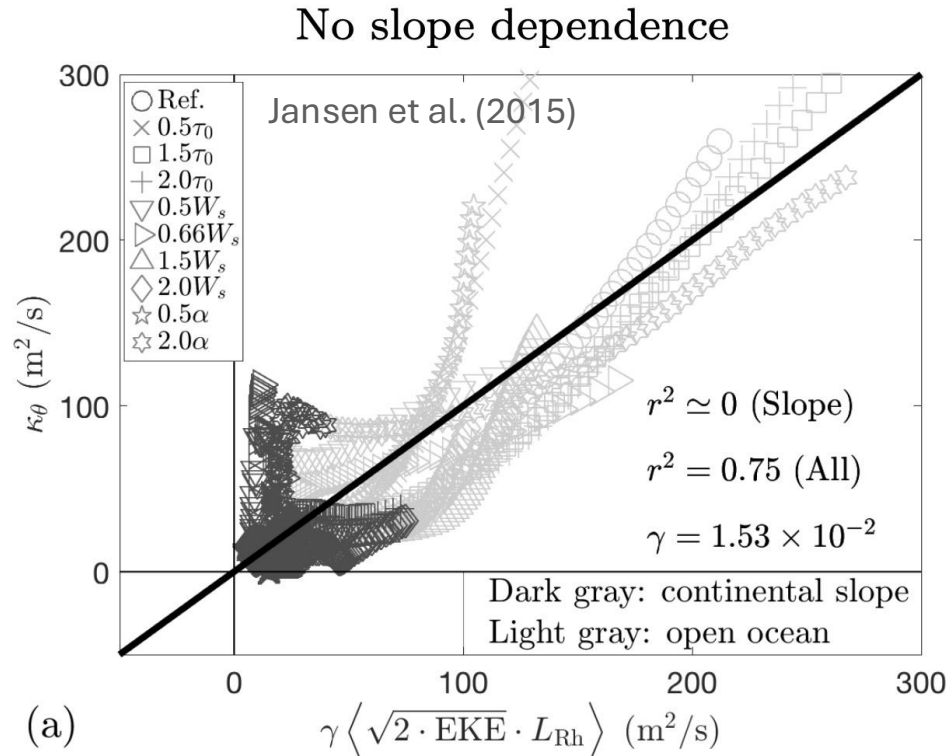
$$\kappa_{Geom} = \gamma_{Geom} \frac{\sqrt{Ri_{loc}}}{f_0} (EKE + EPE),$$

$$\delta = \frac{\text{bottom slope}}{\text{isopycnal slope}}$$

(e.g. Isachsen, 2011)

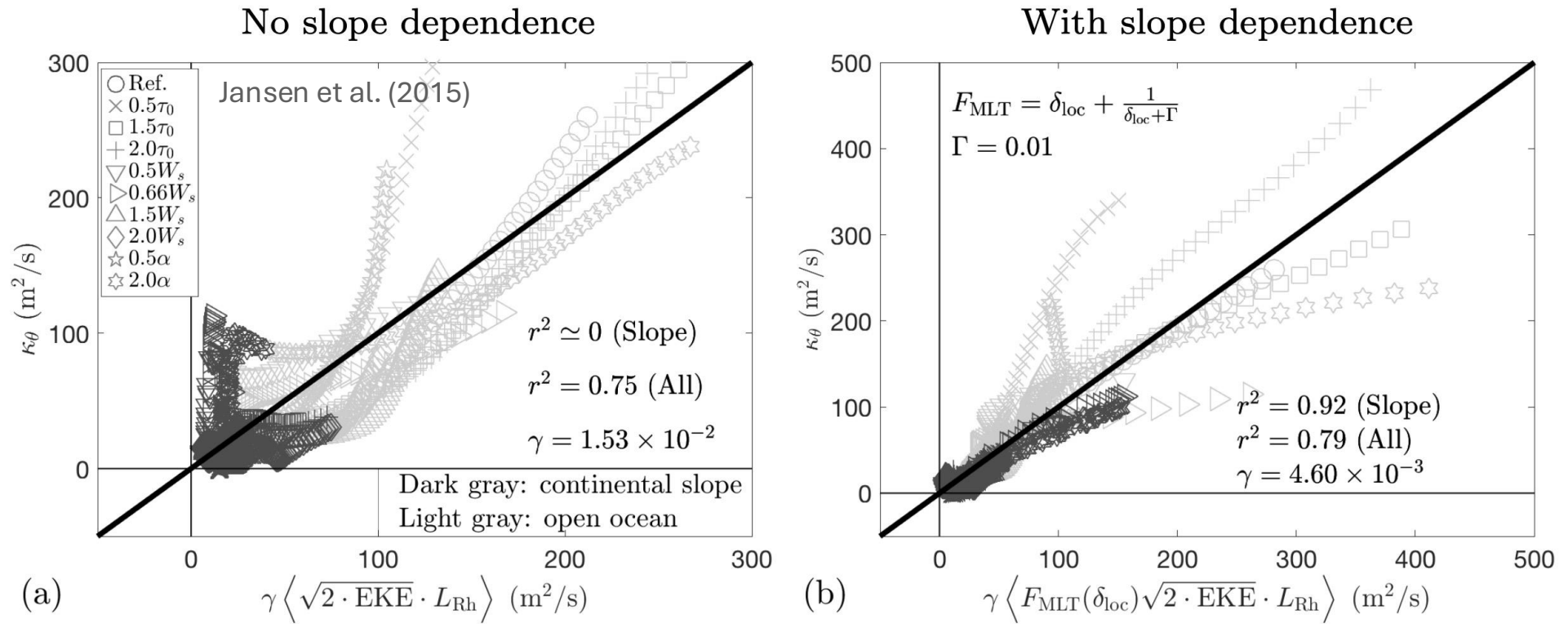
$$\gamma_{Geom} = \gamma \left[\Psi \cdot \tanh(\Gamma \cdot \delta_{loc}) + \frac{1}{\delta_{loc} + \Gamma} \right]$$

Scaling(s) of eddy buoyancy diffusivity: Retrograde



$$\kappa_{\text{MLT}} = \gamma_{\text{MLT}} \sqrt{2 \cdot \text{EKE}} \cdot L_{\text{Rh}},$$

Scaling(s) of eddy buoyancy diffusivity: Retrograde



(a) Original and (b) slope-aware forms of the MLT-based scaling vs κ_θ .

$$\kappa_{\text{MLT}} = \gamma_{\text{MLT}} \sqrt{2 \cdot \text{EKE}} \cdot L_{\text{Rh}},$$

$$\delta = \frac{\text{bottom slope}}{\text{isopycnal slope}}$$

$$\gamma_{\text{MLT}} = \gamma \left[\delta_{\text{loc}} + \frac{1}{\delta_{\text{loc}} + \Gamma} \right]$$

(e.g. Isachsen, 2011)

Scaling of eddy momentum fluxes: Retrograde

$$\frac{\partial}{\partial y} \overline{v'u'} \sim \mathcal{K}_q \cdot \beta_{\text{topog.}}, \quad \mathcal{K}_q \equiv C_{\text{eddy}} \frac{\text{EKE}}{f_0} \Big|_{\text{Barotropic}}$$

Scaling of eddy momentum fluxes: Retrograde

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↓

$$\overline{v'u'} \sim \text{EKE}$$

Scaling of eddy momentum fluxes: Retrograde

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↓

$$\overline{v'u'} \sim \text{EKE}$$

Consistent with barotropic “GEOMETRIC”
(Hoskins et al., 1983; Marshall et al., 2012):

$$M = \frac{\overline{v'^2 - u'^2}}{2}, \quad N = \overline{u'v'}$$

$$M^2 + N^2 \leq \text{EKE}$$

$$N = \gamma_m E \sin 2\phi_m \cos^2 \lambda,$$

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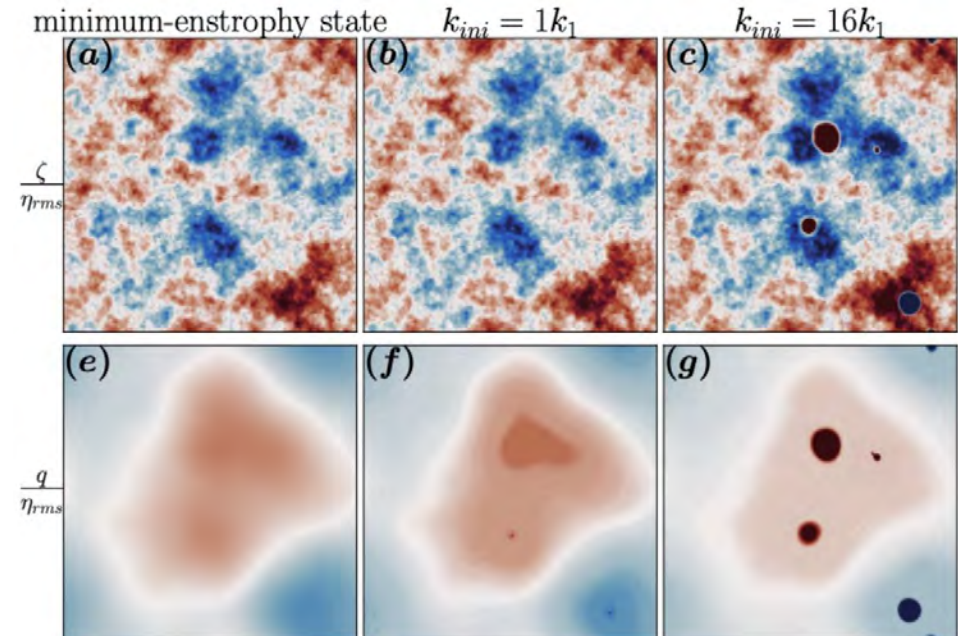
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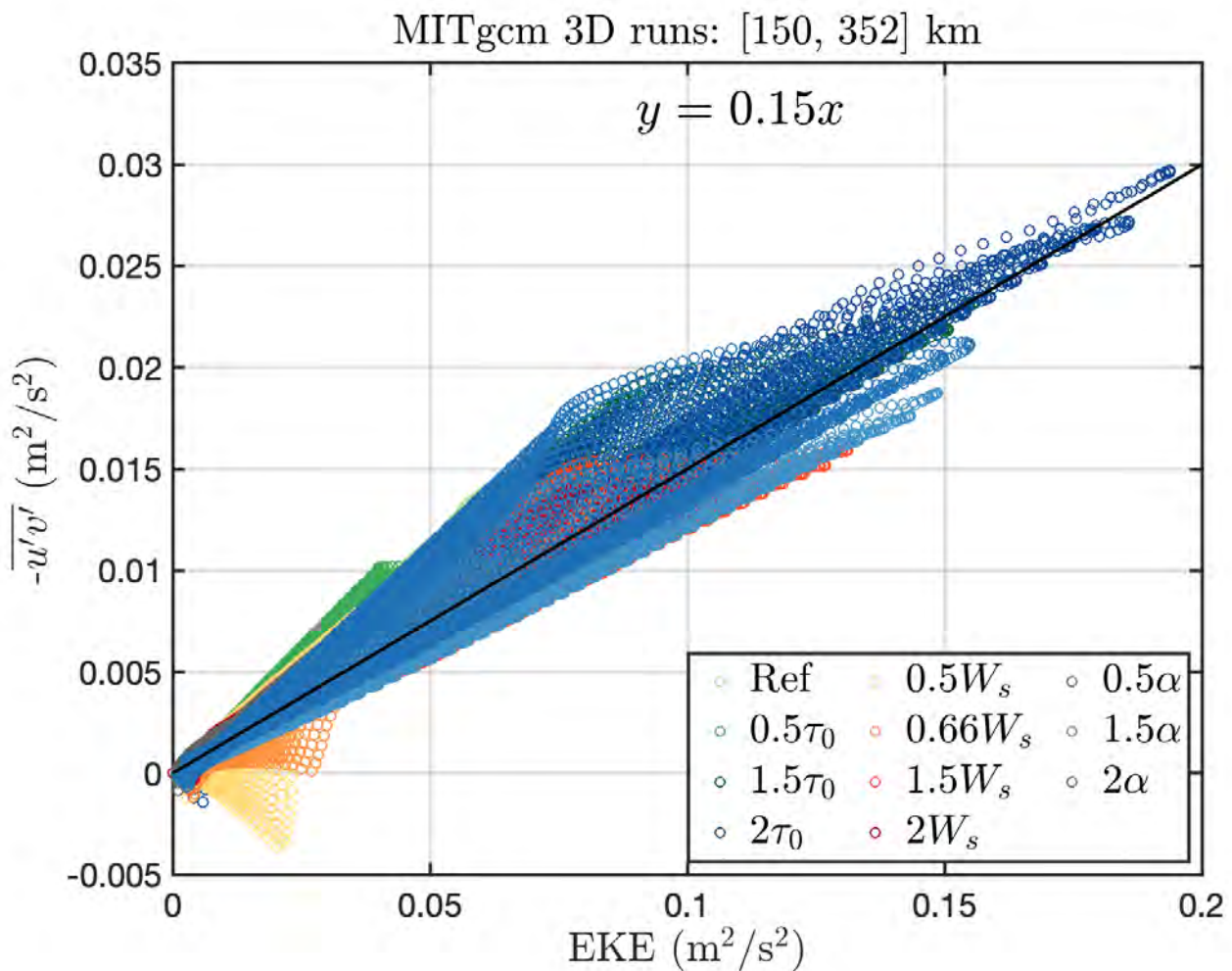
$$N = \gamma_m E \sin 2\phi_m \cos^2 \lambda,$$

Resembling “selective decay” of 2D topographic
turbulence (Bretherton and Haidvogel, 1976):



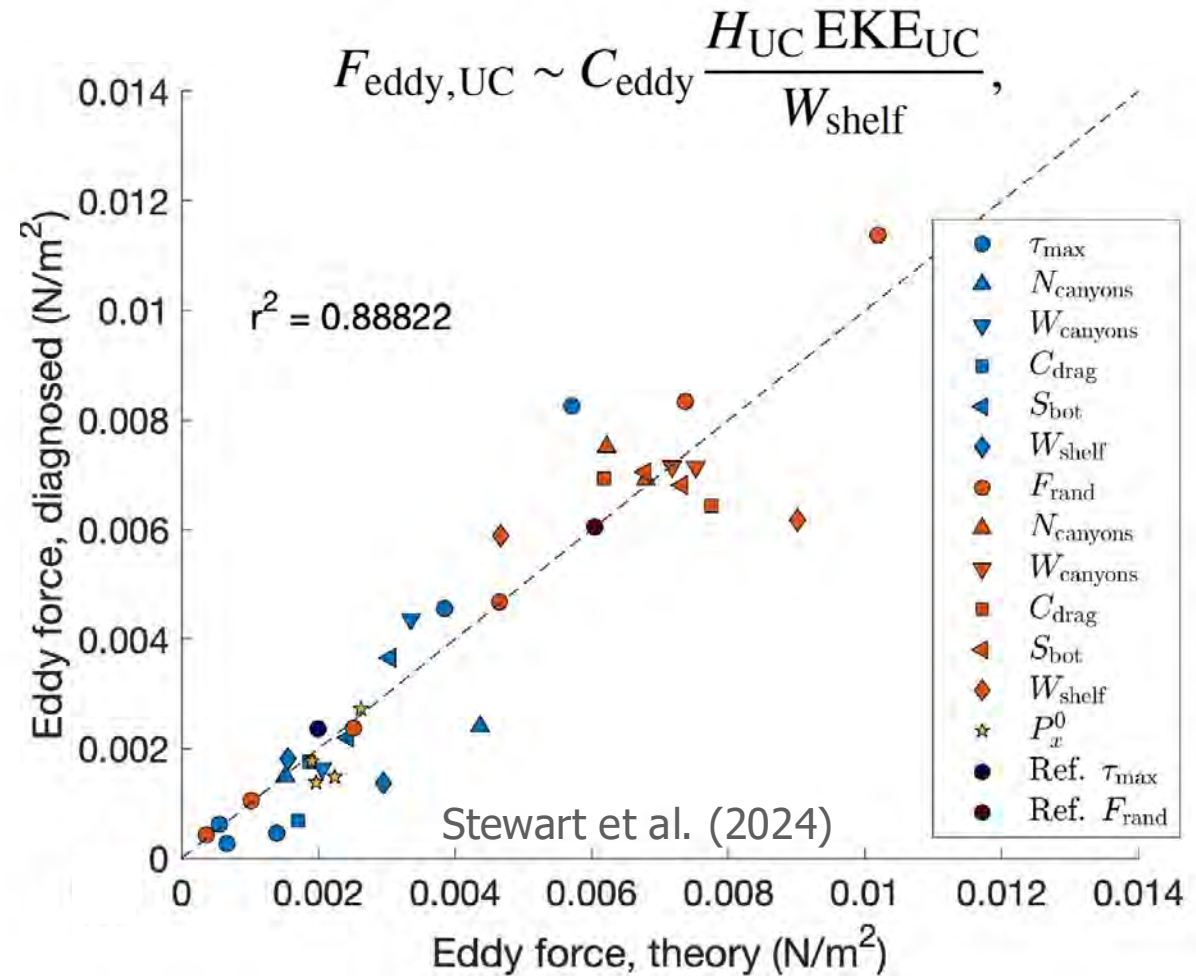
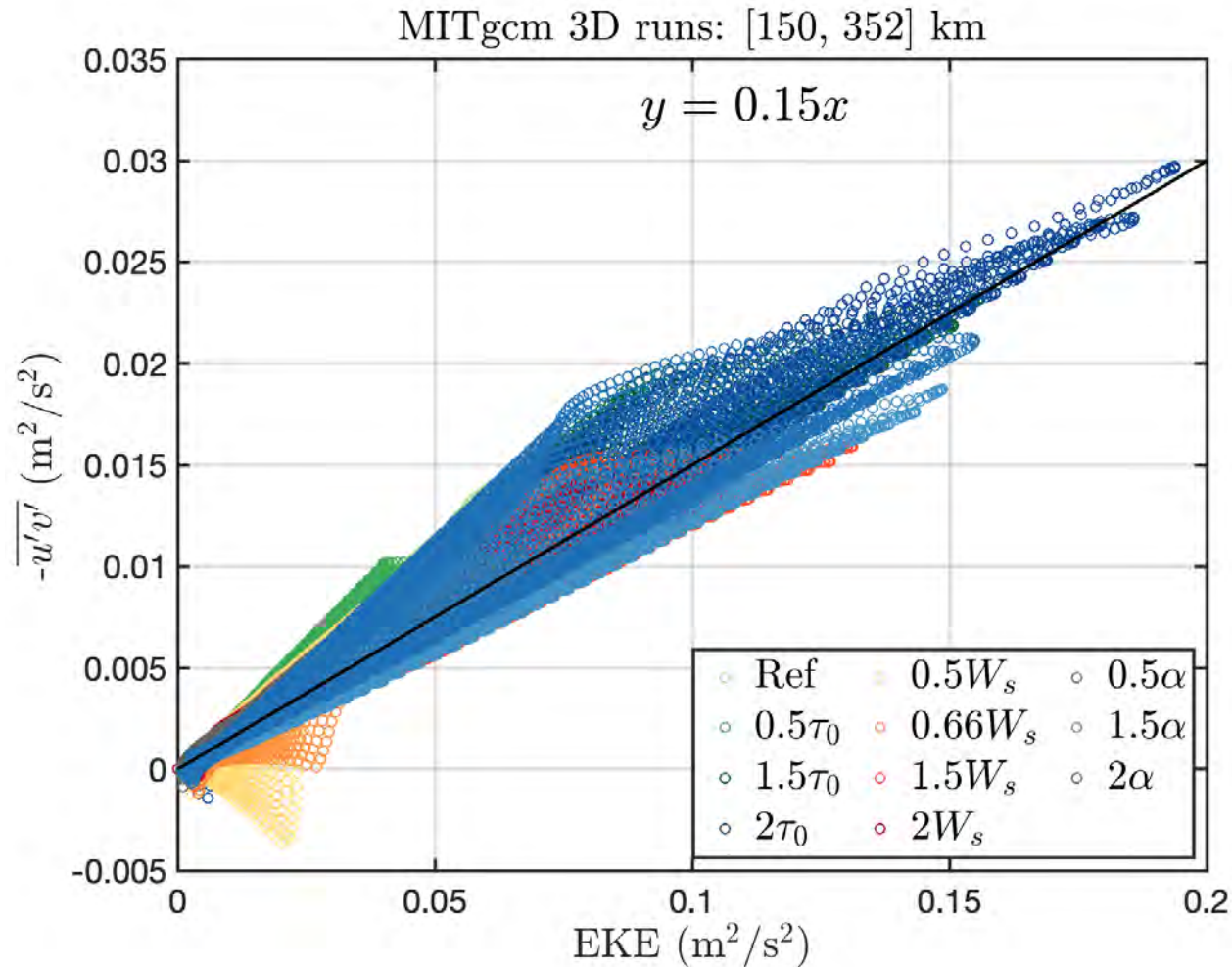
Scaling(s) of eddy momentum fluxes: Retrograde

$$\overline{v'u'} \sim \text{EKE}$$

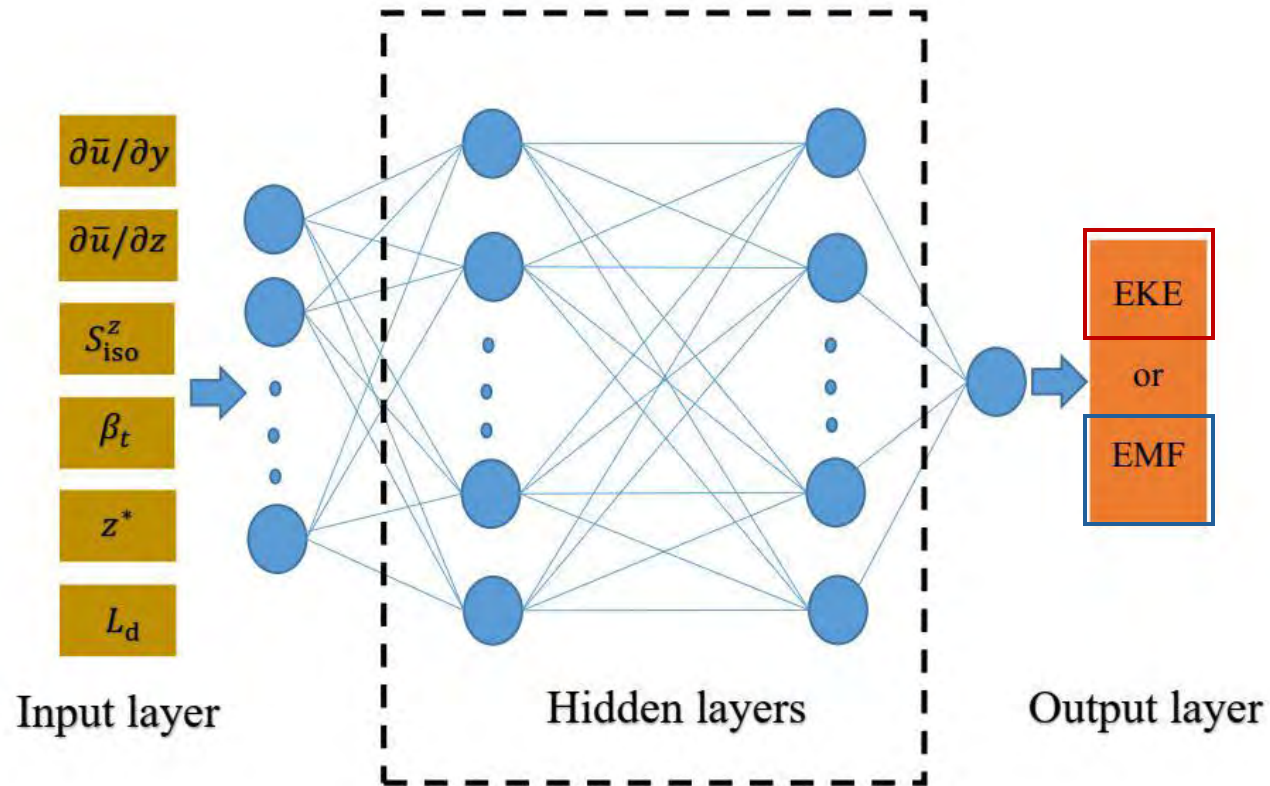


Scaling(s) of eddy momentum fluxes: Retrograde

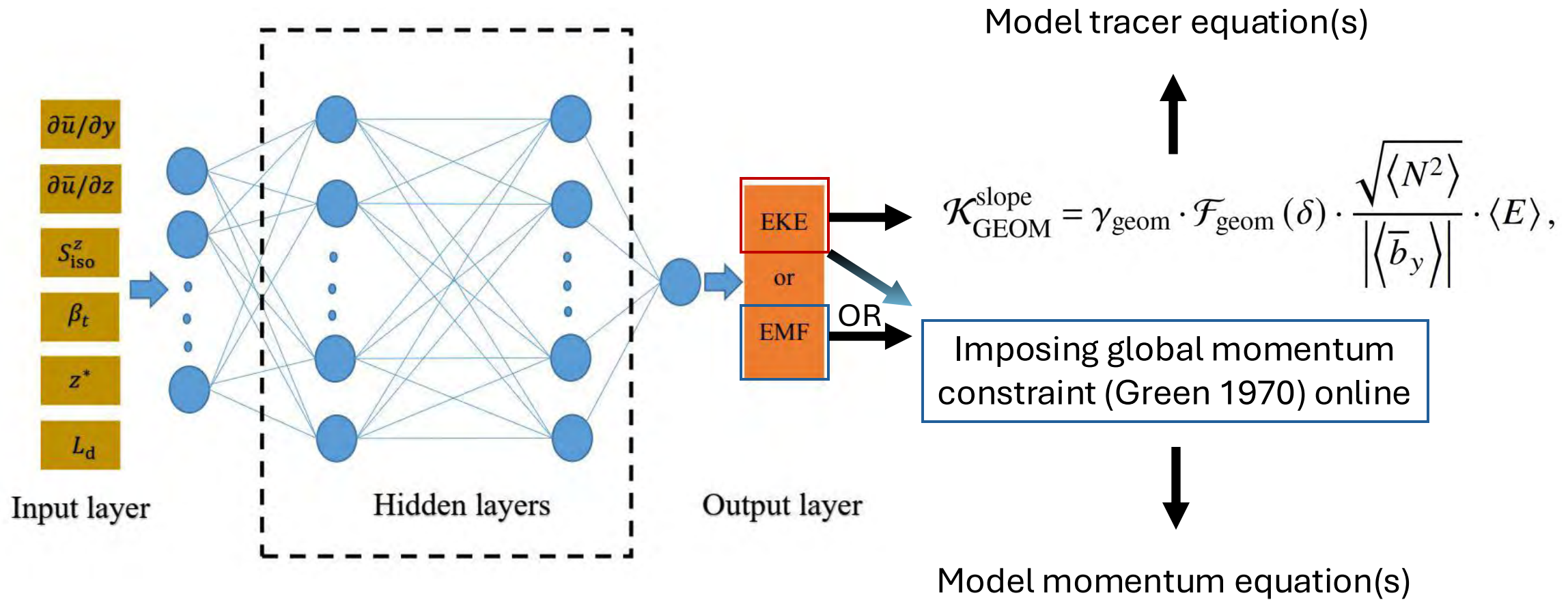
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Convert scaling(s) into eddy closure(s): Retrograde



Convert scaling(s) into eddy closure(s): Retrograde

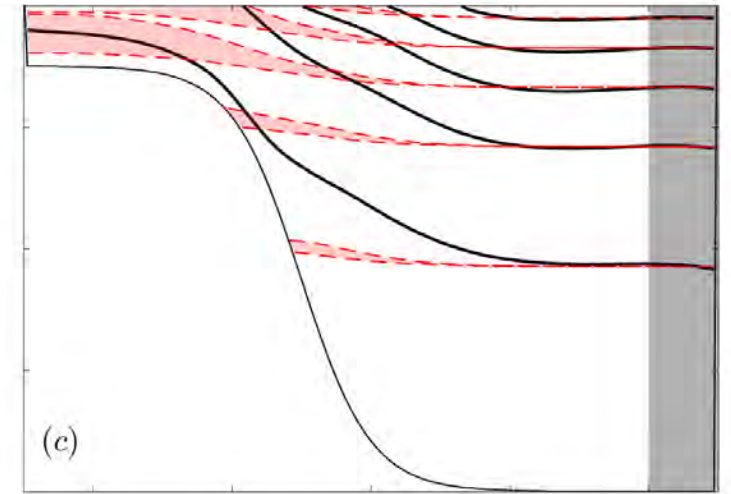
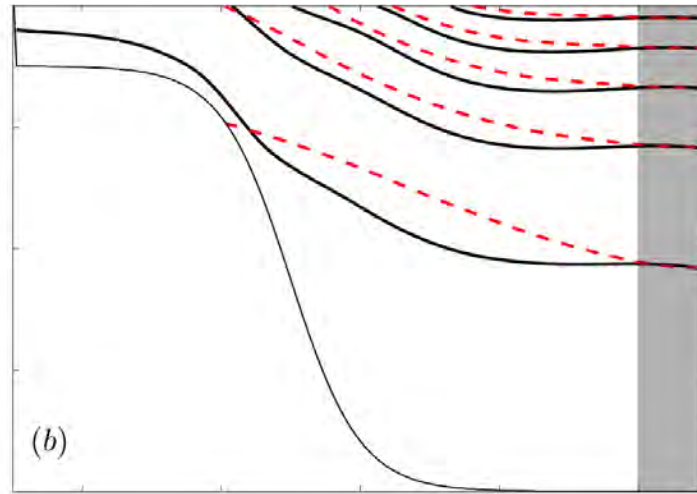
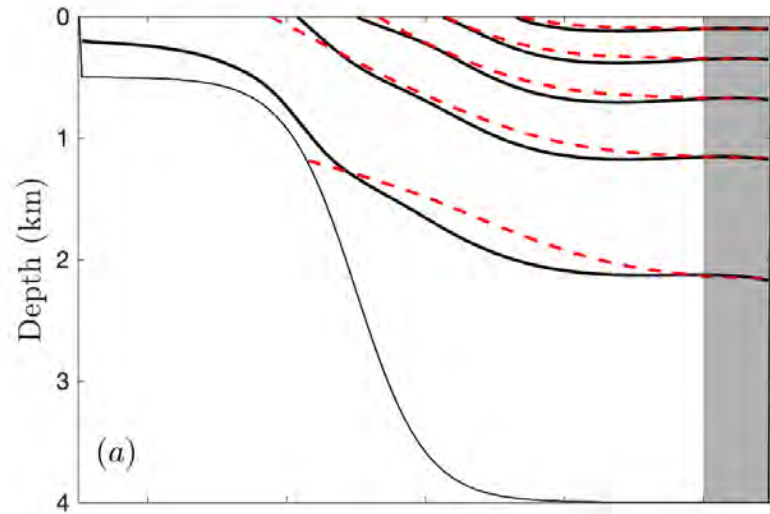


Parameterization: Retrograde

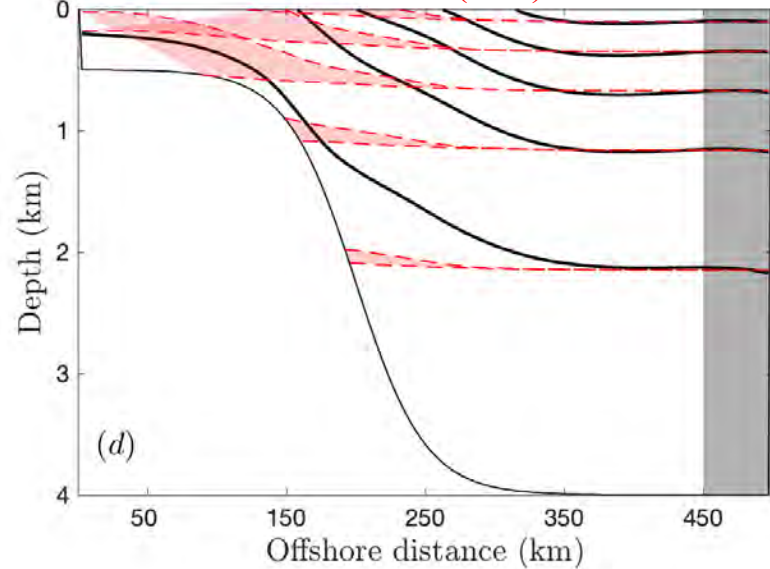
Optimized constant GM

Visbeck 97 scheme

Rhines scale-based (raw)



GEOMETRIC (raw)

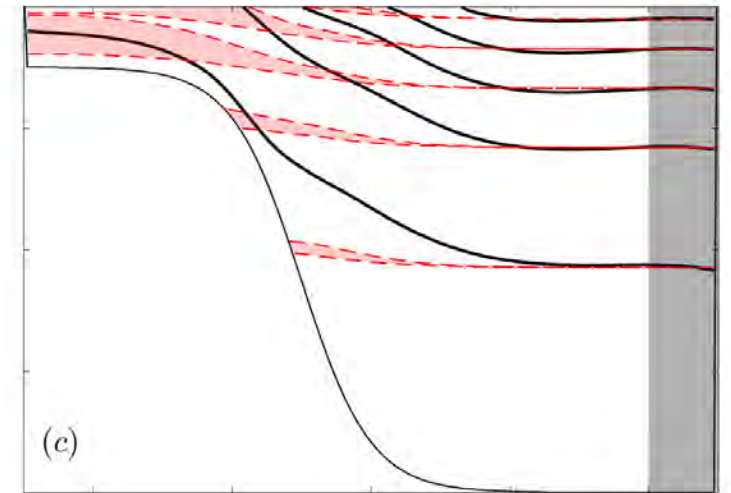
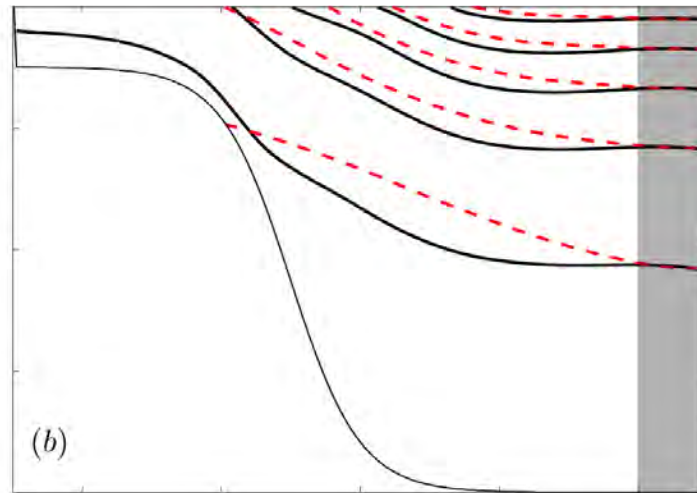
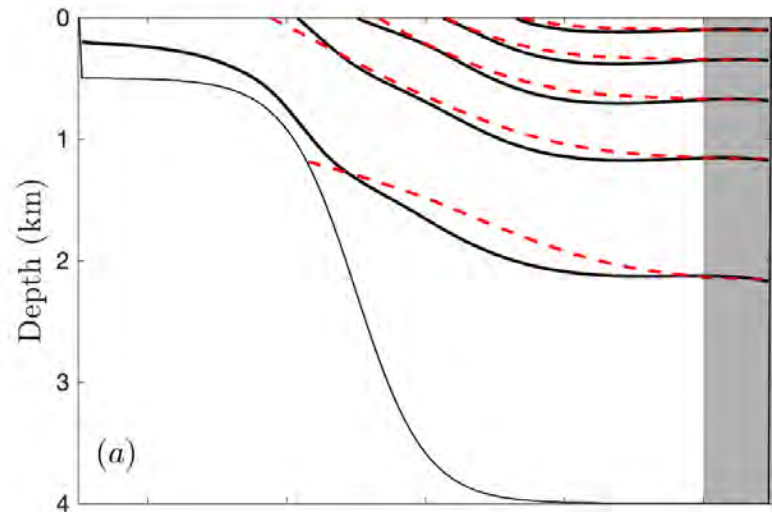


Parameterization: Retrograde

Optimized constant GM

Visbeck 97 scheme

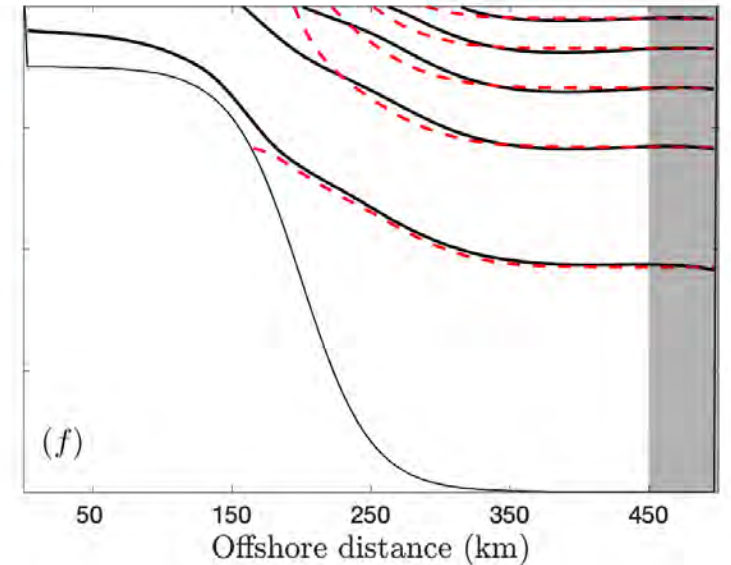
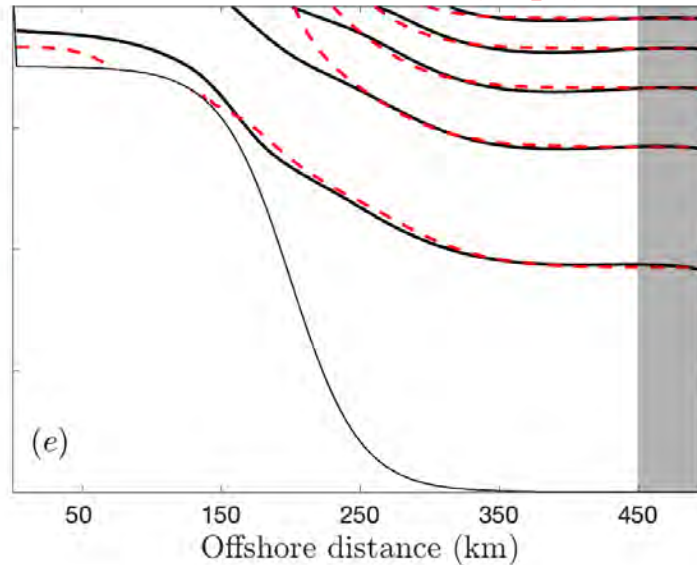
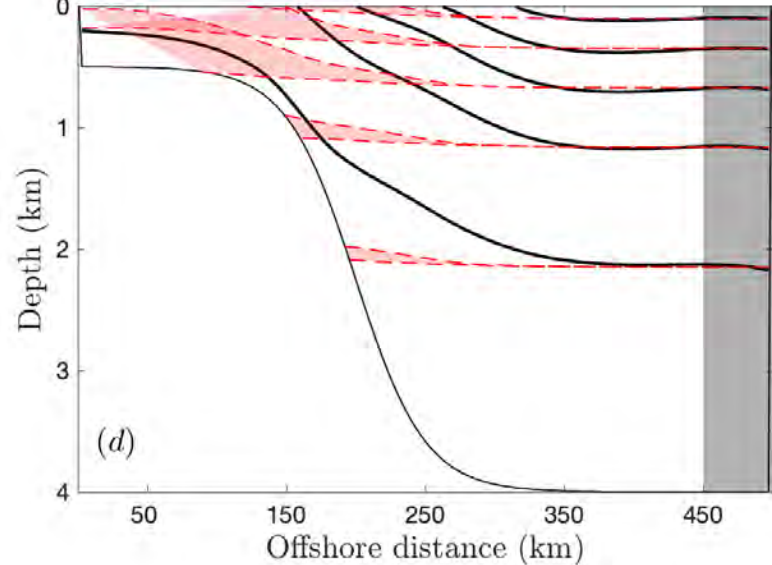
Rhines scale-based (raw)



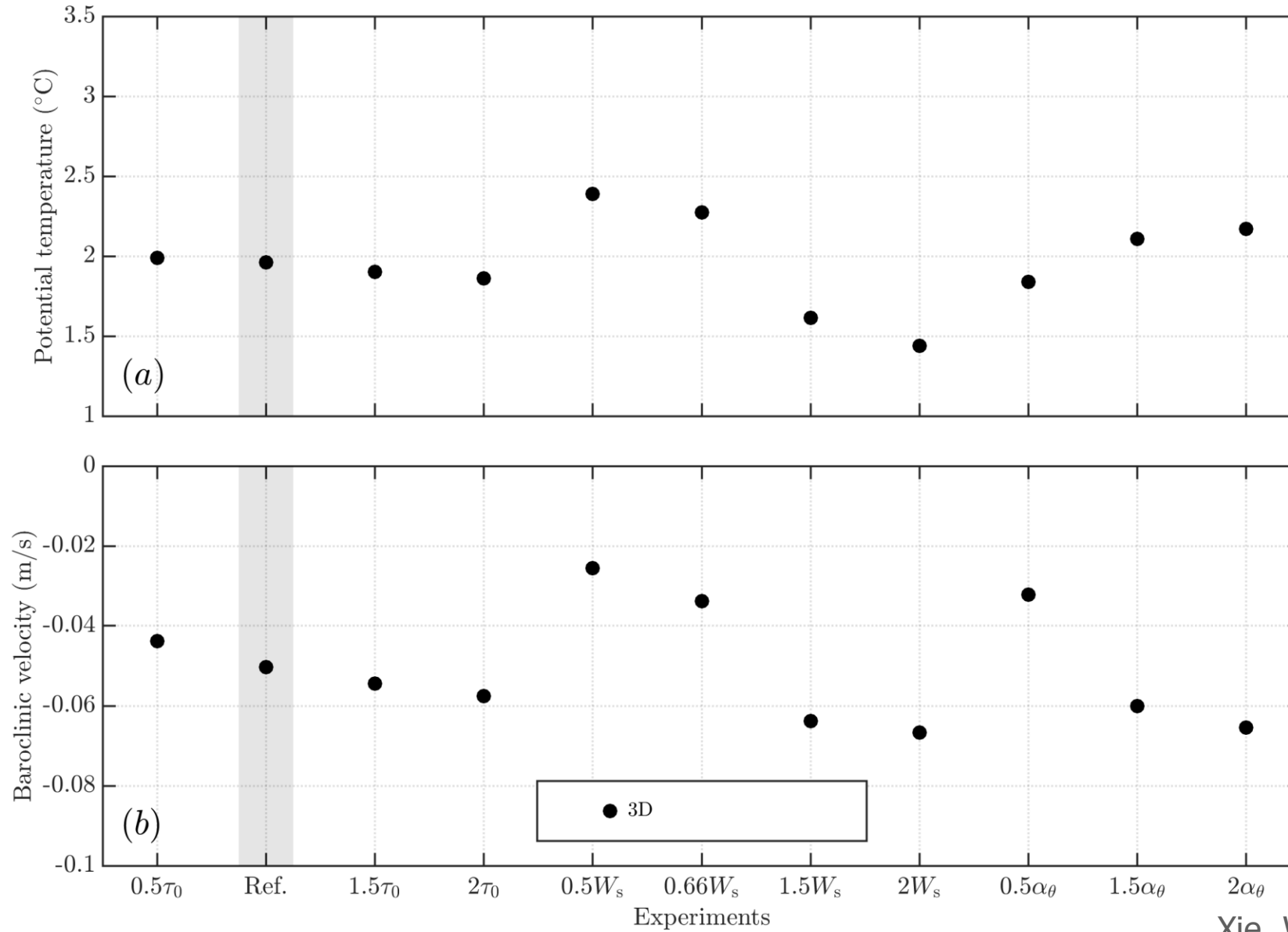
GEOMETRIC (raw)

Rhines-scale based (slope)

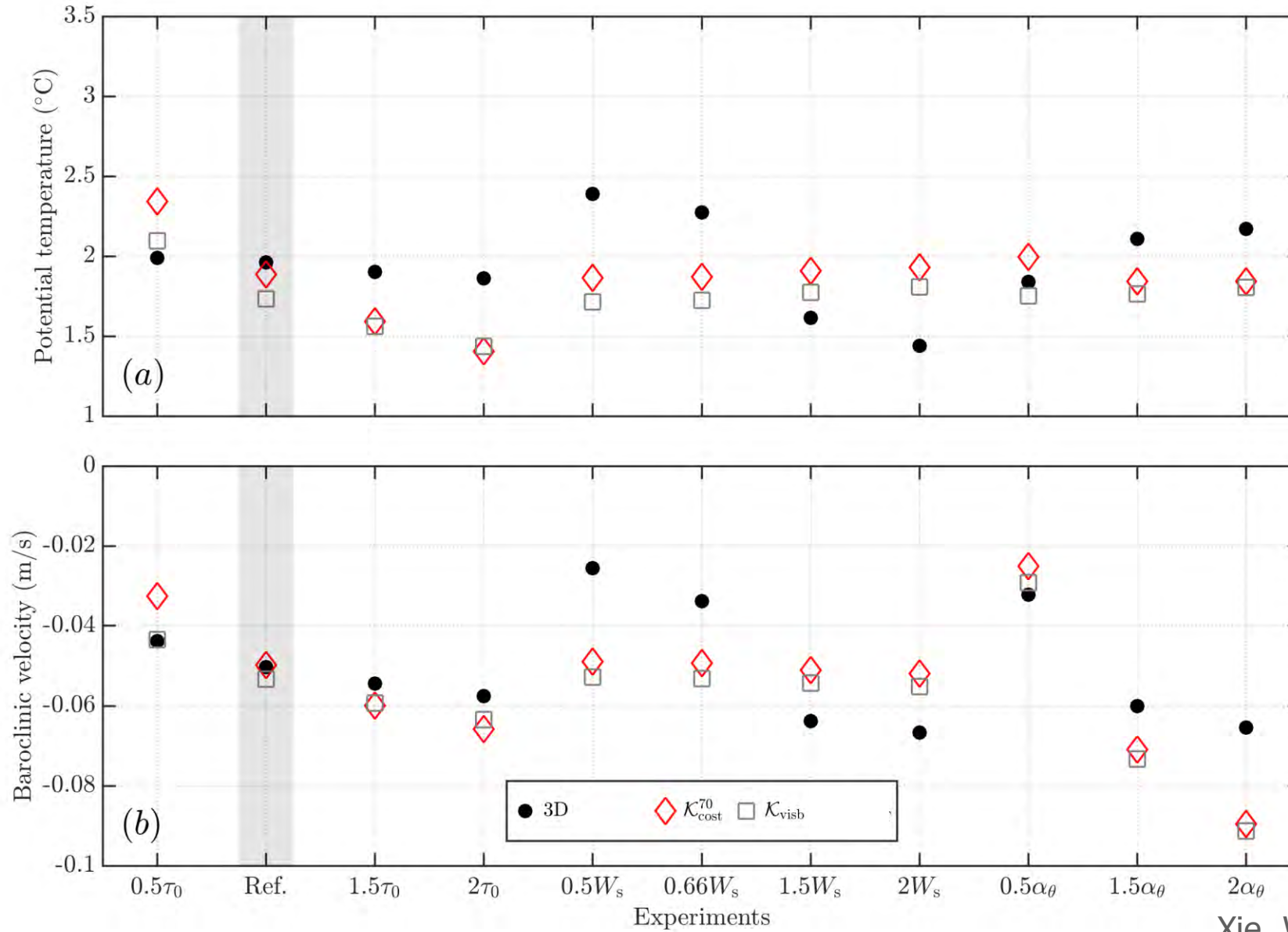
GEOMETRIC (slope)



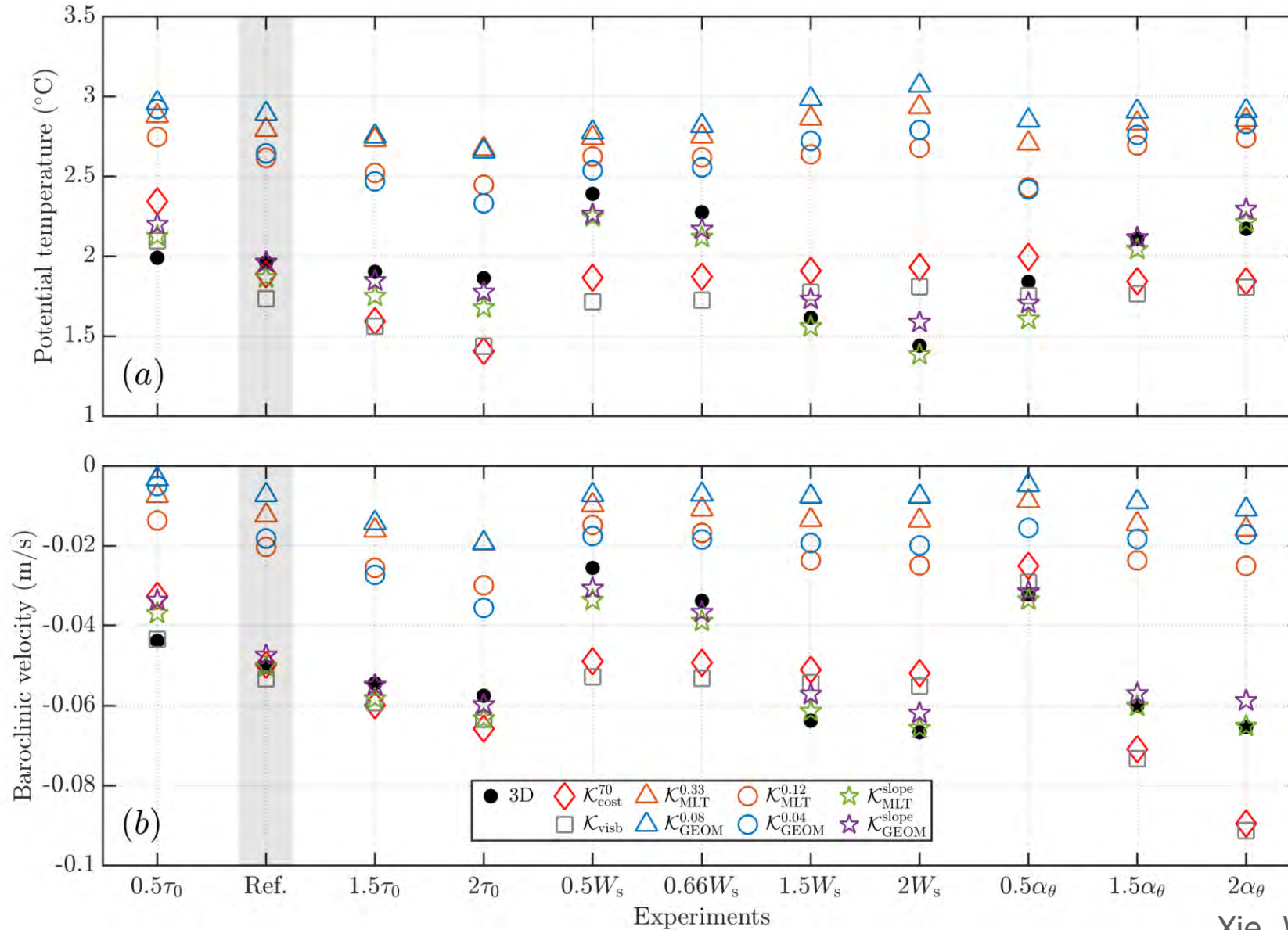
Parameterization: Retrograde



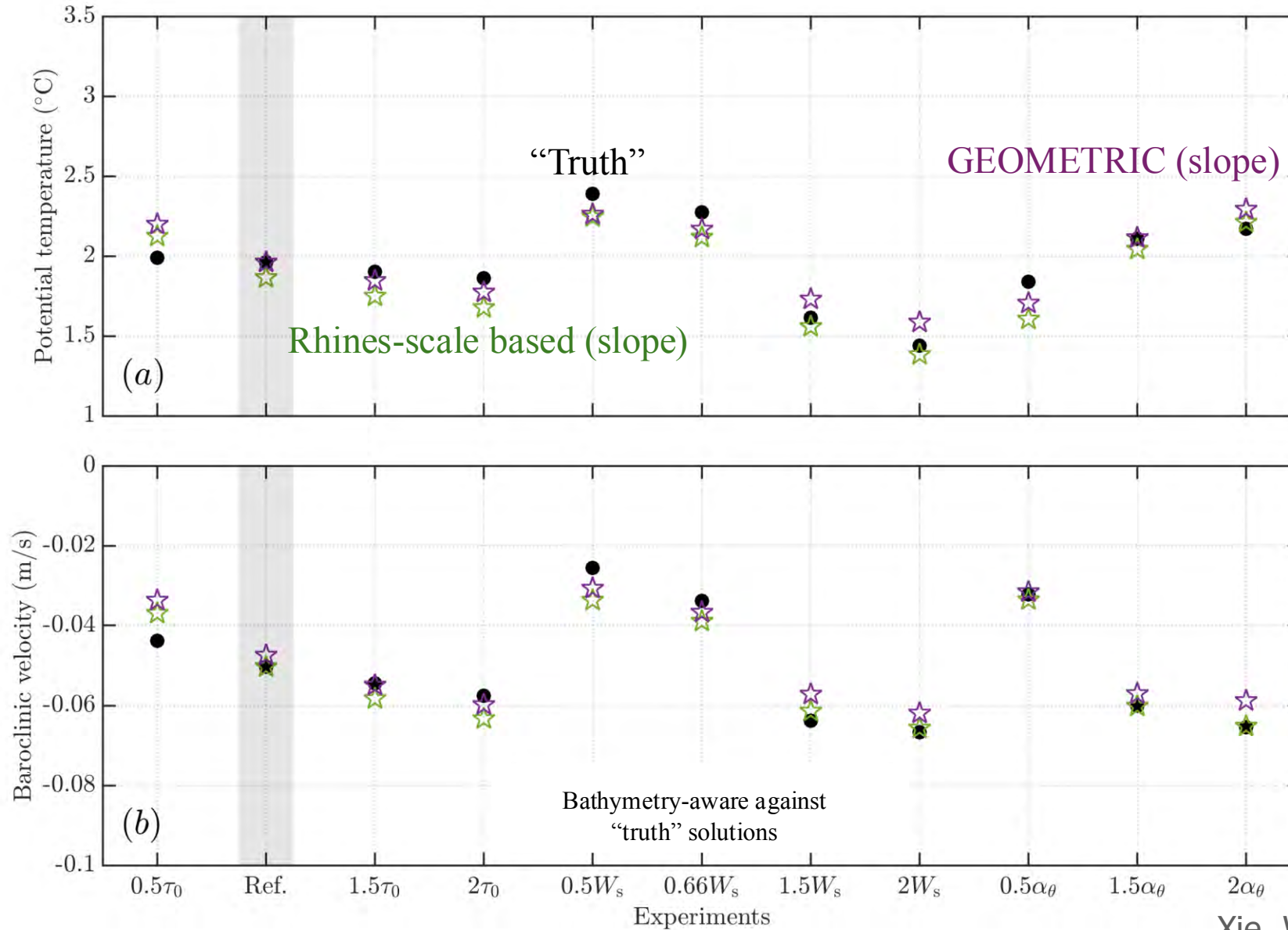
Parameterization: Retrograde



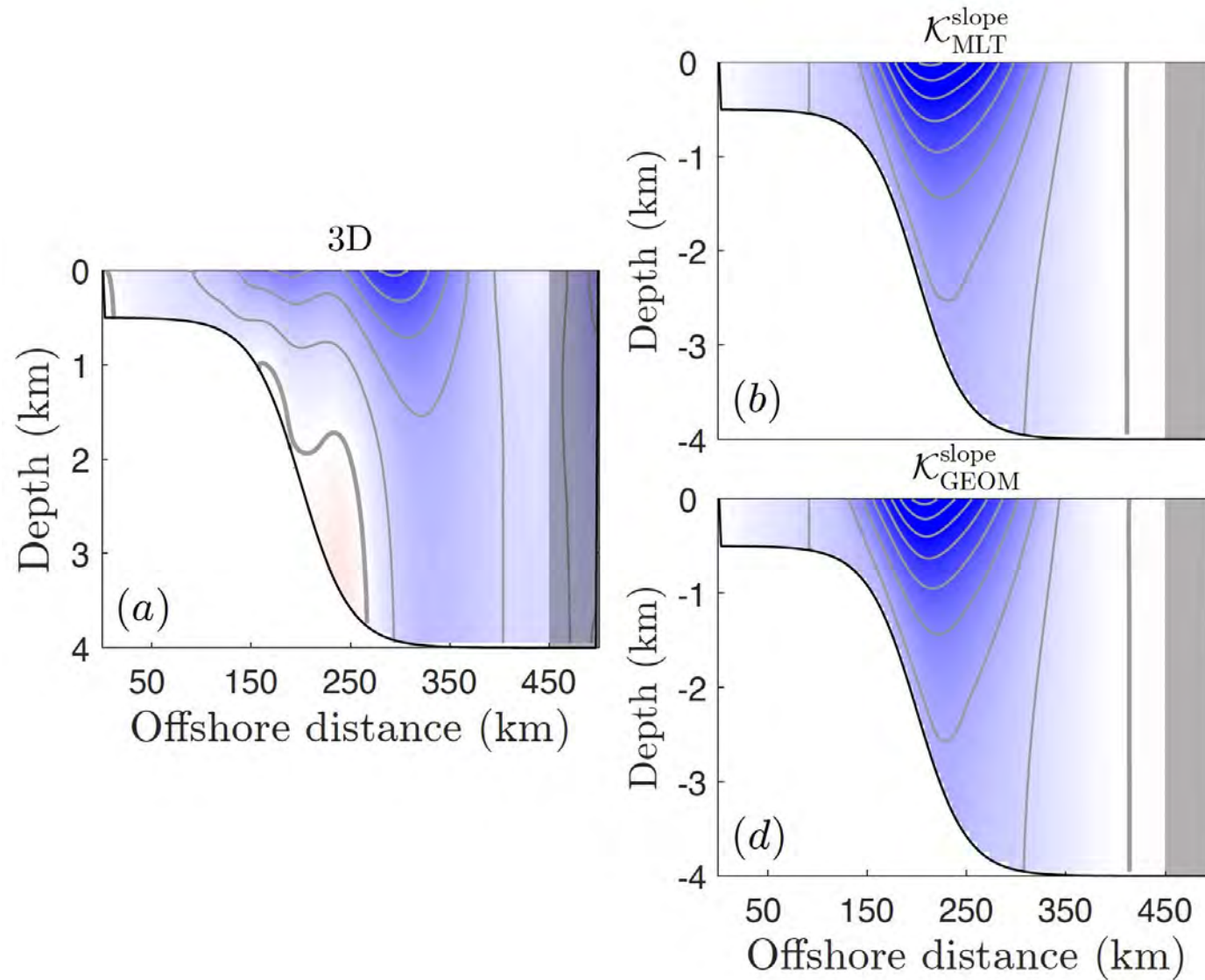
Parameterization: Retrograde



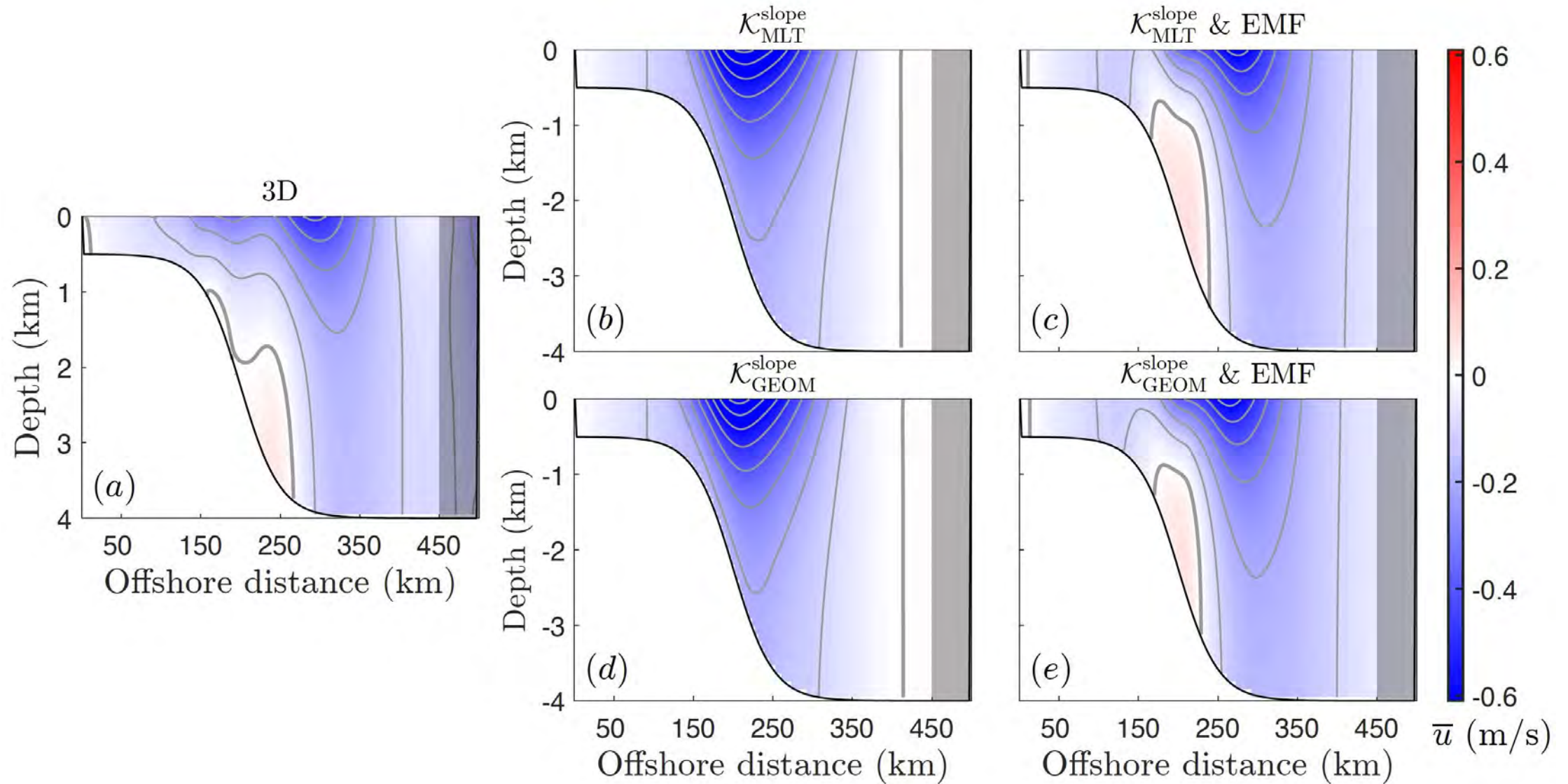
Parameterization: Retrograde



Impact of parameterized eddy momentum forcing: Retrograde



Impact of parameterized eddy momentum forcing: Retrograde



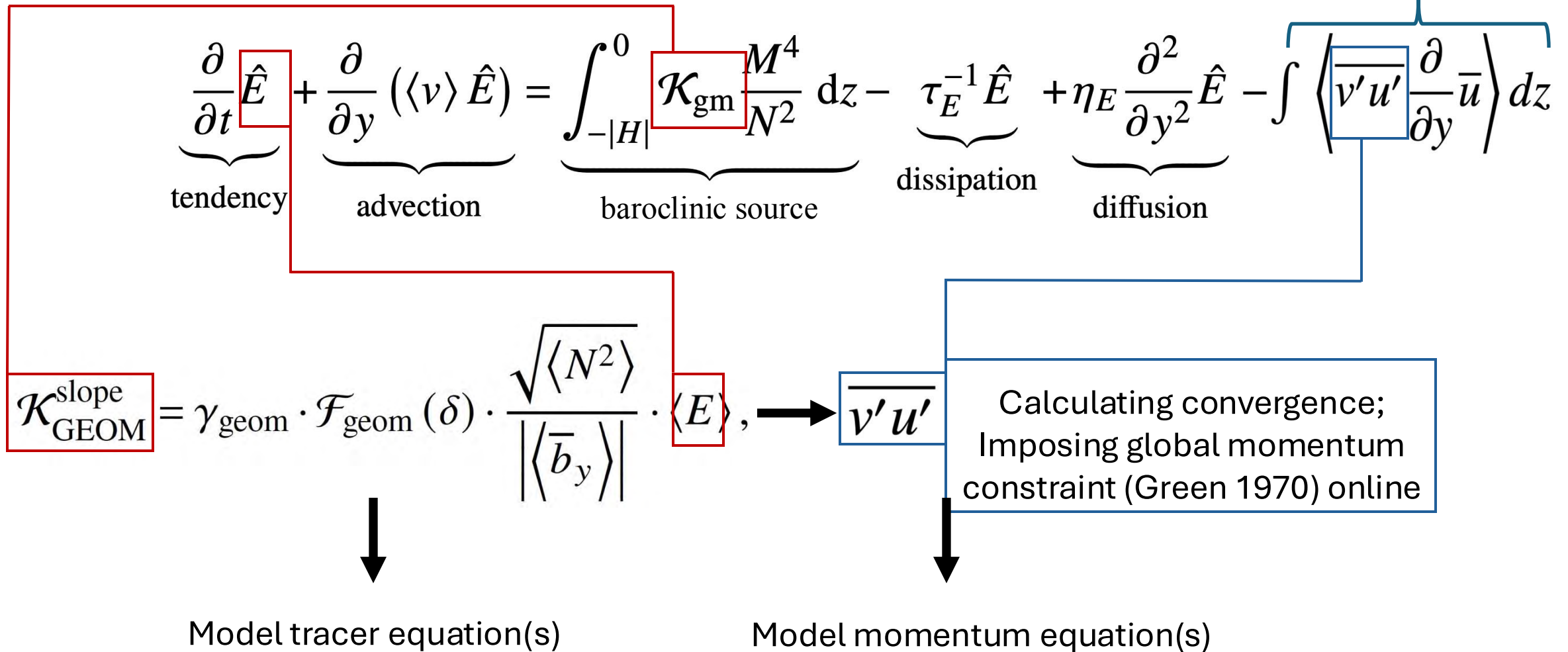
Physics-based eddy closure(s): Retrograde

$$\underbrace{\frac{\partial}{\partial t} \hat{E}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{baroclinic source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2}{\partial y^2} \hat{E}}_{\text{diffusion}}$$

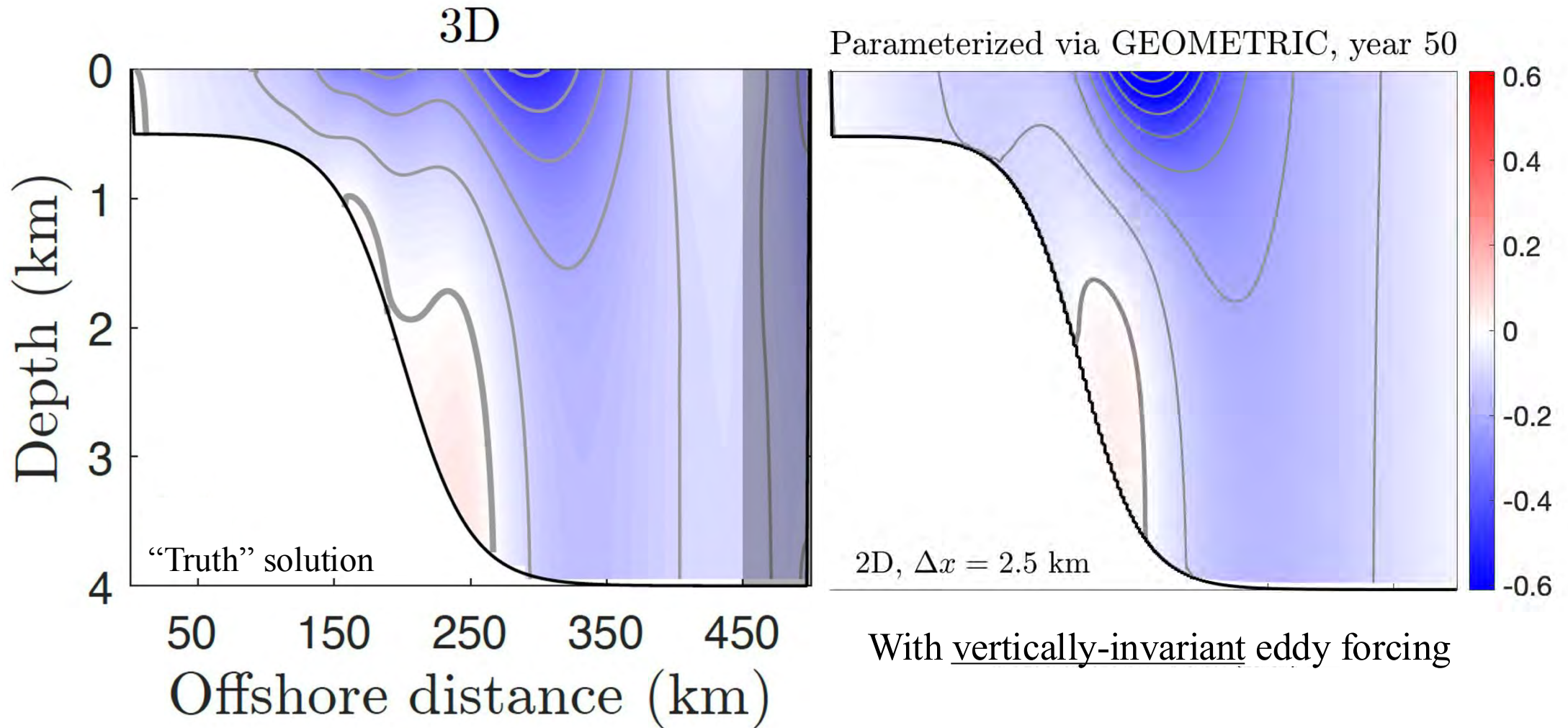
$$\mathcal{K}_{\text{GEOM}}^{\text{slope}} = \gamma_{\text{geom}} \cdot \mathcal{F}_{\text{geom}}(\delta) \cdot \frac{\sqrt{\langle N^2 \rangle}}{|\langle \bar{b}_y \rangle|} \cdot \langle E \rangle, \longrightarrow \overline{v' u'}$$

Calculating convergence;
 Imposing global momentum
 constraint (Green 1970) online

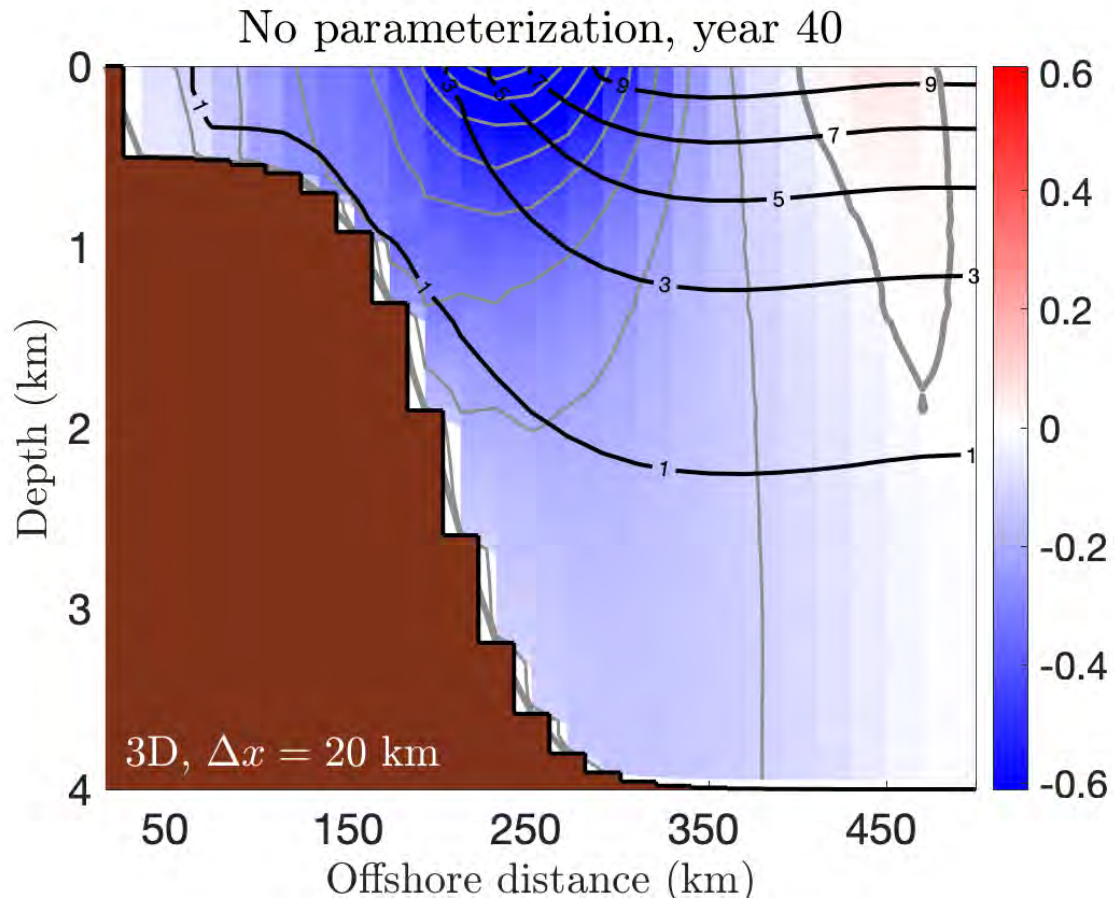
Physics-based eddy closure(s): Retrograde



Physics-based eddy closure(s): Retrograde 2D runs

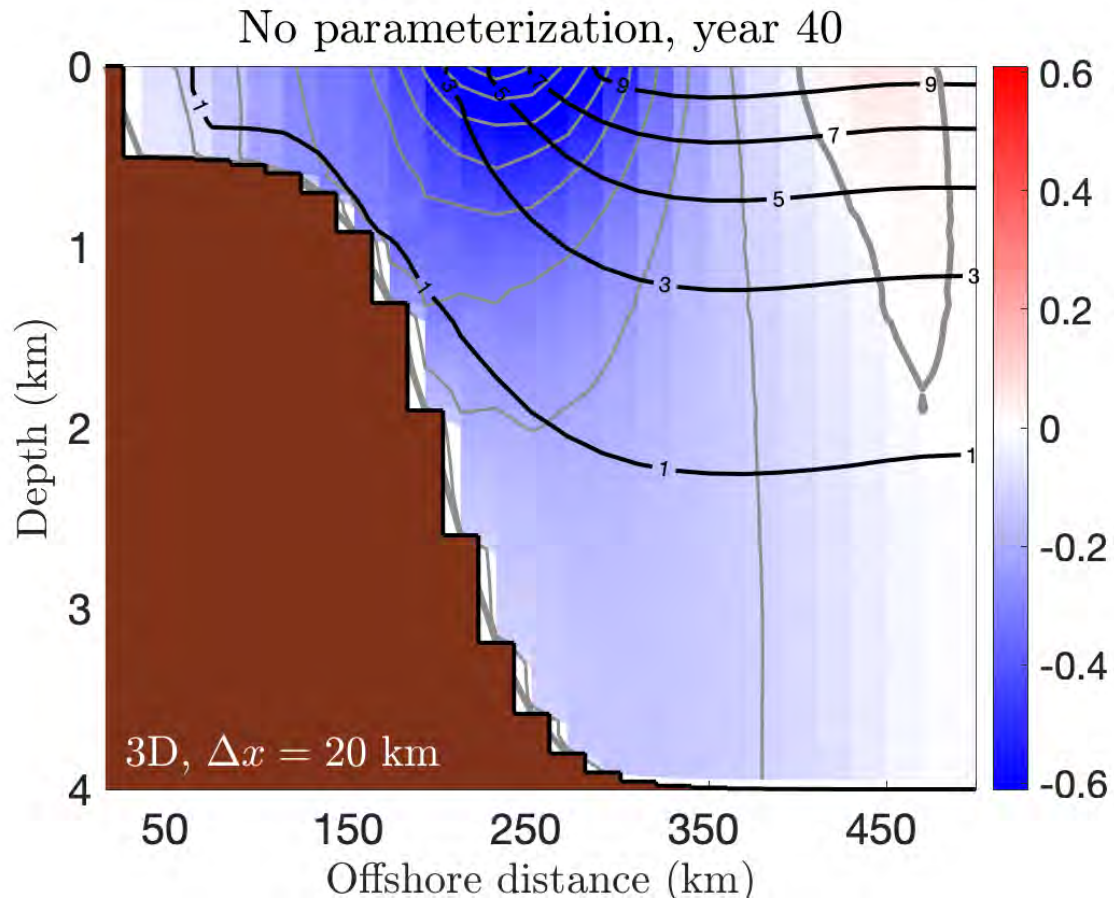


Physics-based eddy closure(s): 3D coarse-grid “raw” flow

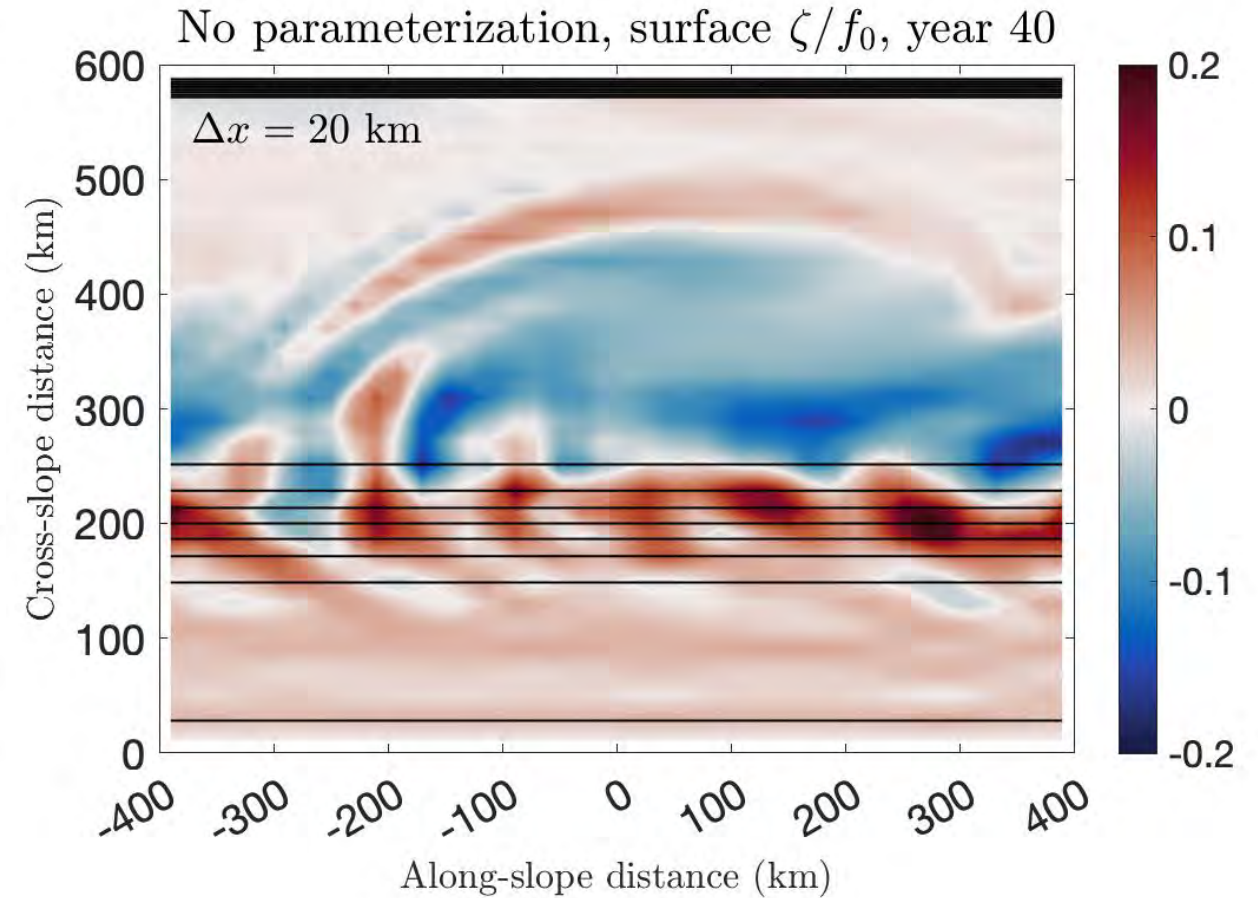


Acceptable stratification, wrong jet structures.

Physics-based eddy closure(s): 3D coarse-grid “raw” flow



Acceptable stratification, wrong jet structures.



Ongoing work with Julian Mak and Andrew Stewart

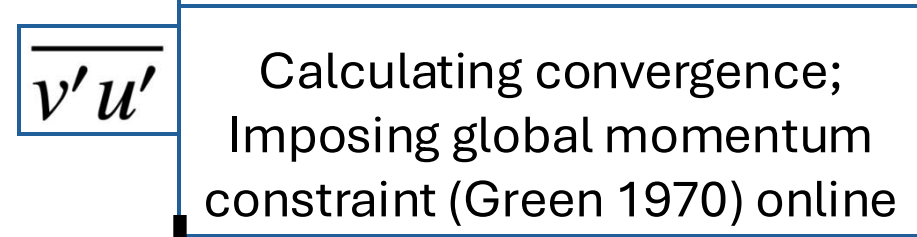
Physics-based eddy closure(s): Retrograde 3D runs

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{baroclinic source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}} - \int \left\langle \overline{v' u'} \frac{\partial \bar{u}}{\partial y} \right\rangle dz$$

Physics-based eddy closure(s): Retrograde 3D runs

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{baroclinic source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}} - \int \left\langle \overline{v' u'} \frac{\partial \bar{u}}{\partial y} \right\rangle dz$$

barotropic source or sink



Model momentum equation(s)

Physics-based eddy closure(s): Retrograde 3D runs

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{baroclinic source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}} - \int \left\langle \overline{v' u'} \frac{\partial \bar{u}}{\partial y} \right\rangle dz$$

barotropic source or sink

Subgrid entry as energy source

Grid-dissipated KE of alongshore anomalous flow

$\overline{v' u'}$

Calculating convergence; Imposing global momentum constraint (Green 1970) online

Model momentum equation(s)

Physics-based eddy closure(s): Retrograde 3D runs

$$\underbrace{\frac{\partial \hat{E}}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\partial}{\partial y} (\langle v \rangle \hat{E})}_{\text{advection}} = \underbrace{\int_{-|H|}^0 \mathcal{K}_{\text{gm}} \frac{M^4}{N^2} dz}_{\text{baroclinic source}} - \underbrace{\tau_E^{-1} \hat{E}}_{\text{dissipation}} + \underbrace{\eta_E \frac{\partial^2 \hat{E}}{\partial y^2}}_{\text{diffusion}} - \int \left\langle \overline{v' u'} \frac{\partial \bar{u}}{\partial y} \right\rangle dz$$

barotropic source or sink

Subgrid entry as energy source

Grid-dissipated KE of alongshore anomalous flow

Along-isobath averaging and calculating energetic rate online;

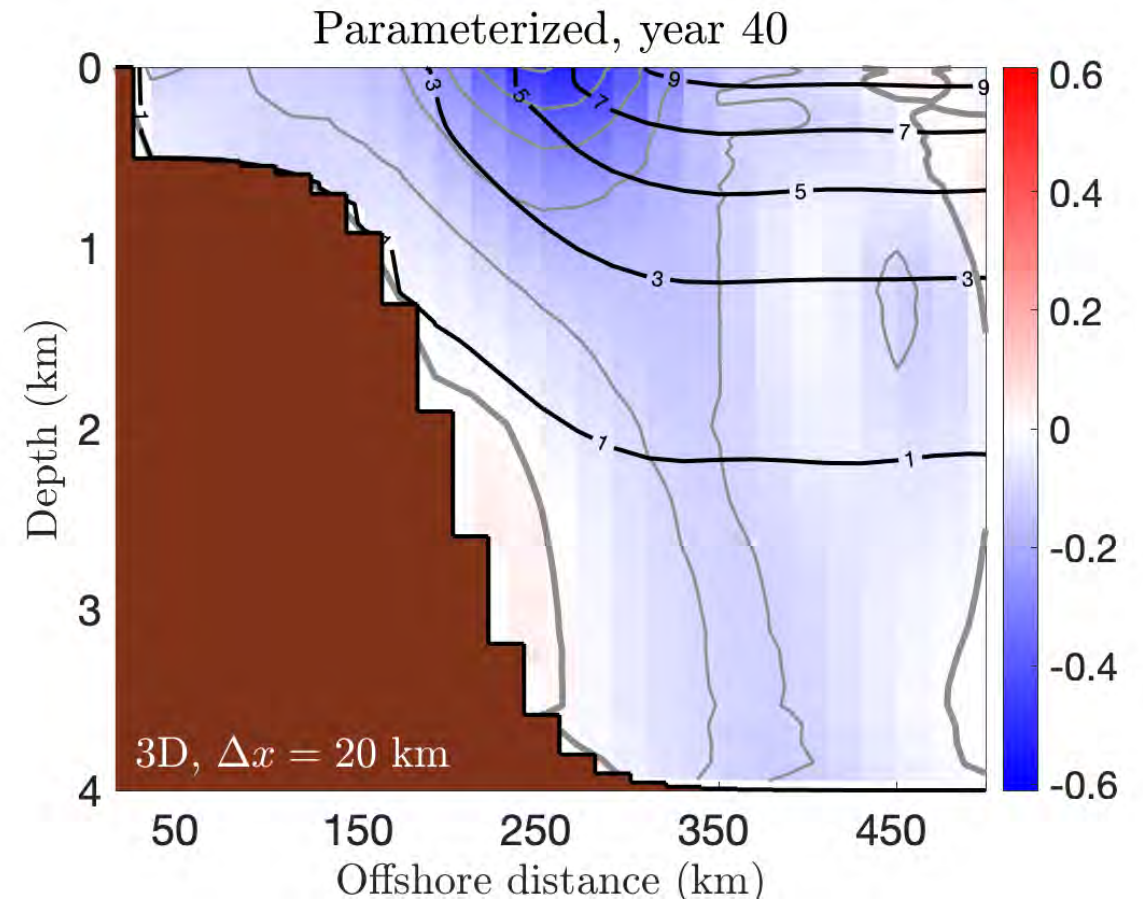
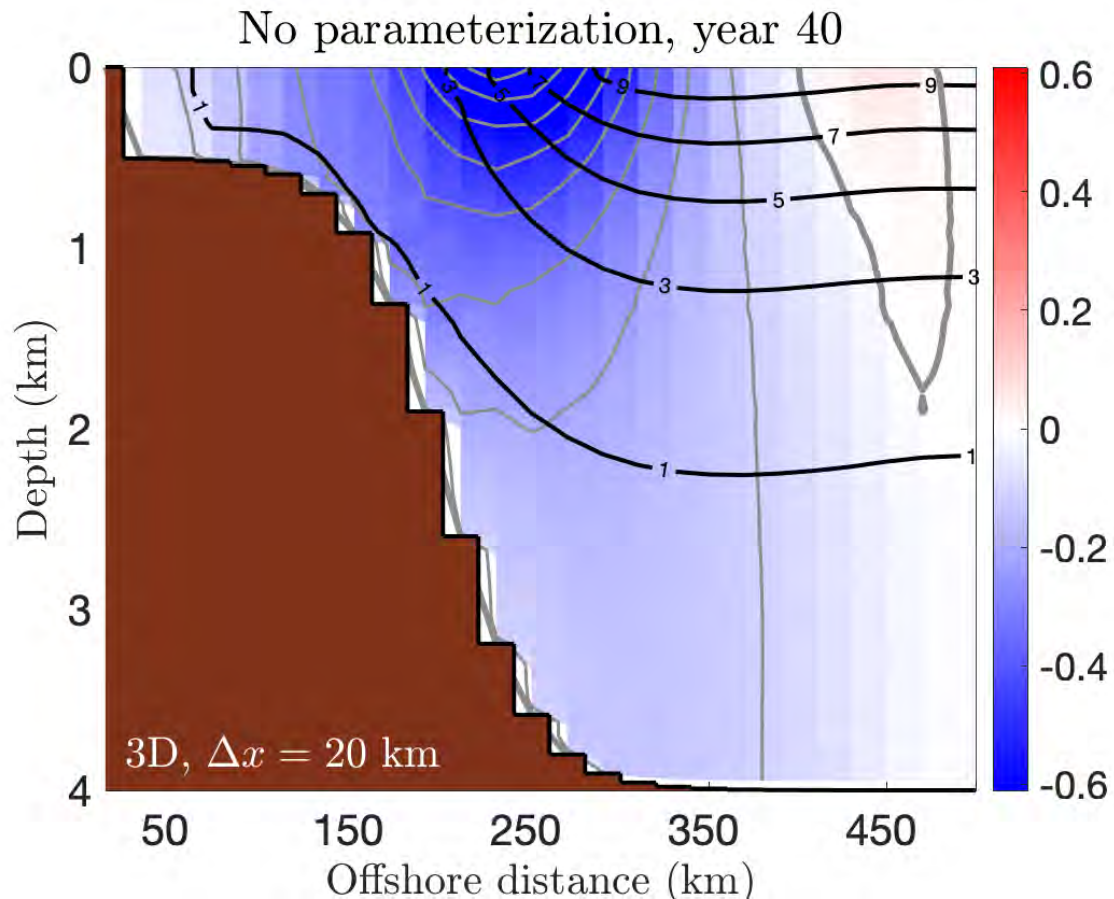
A special form of eddying/mean “flow-splitting” (Mak et al. 2023).

$\overline{v' u'}$

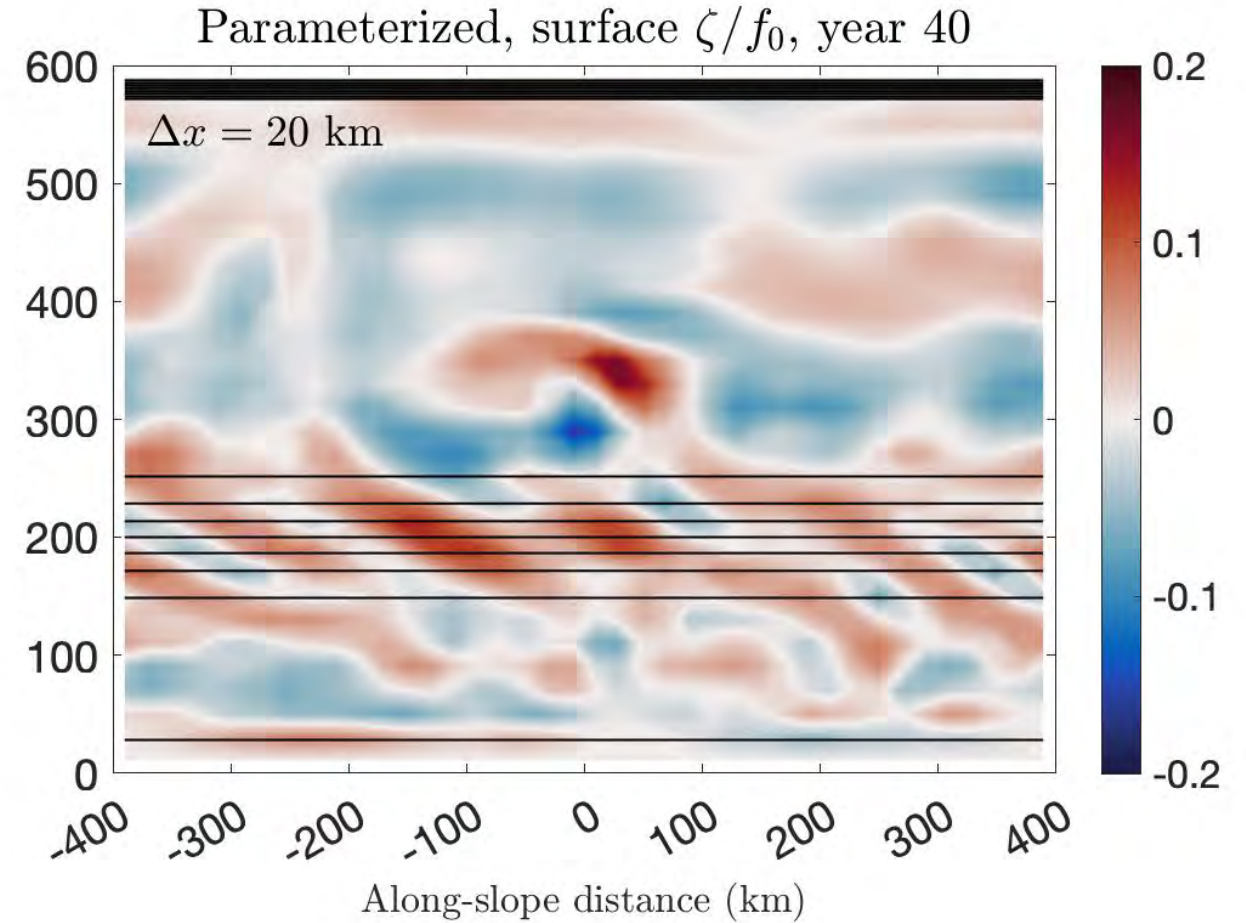
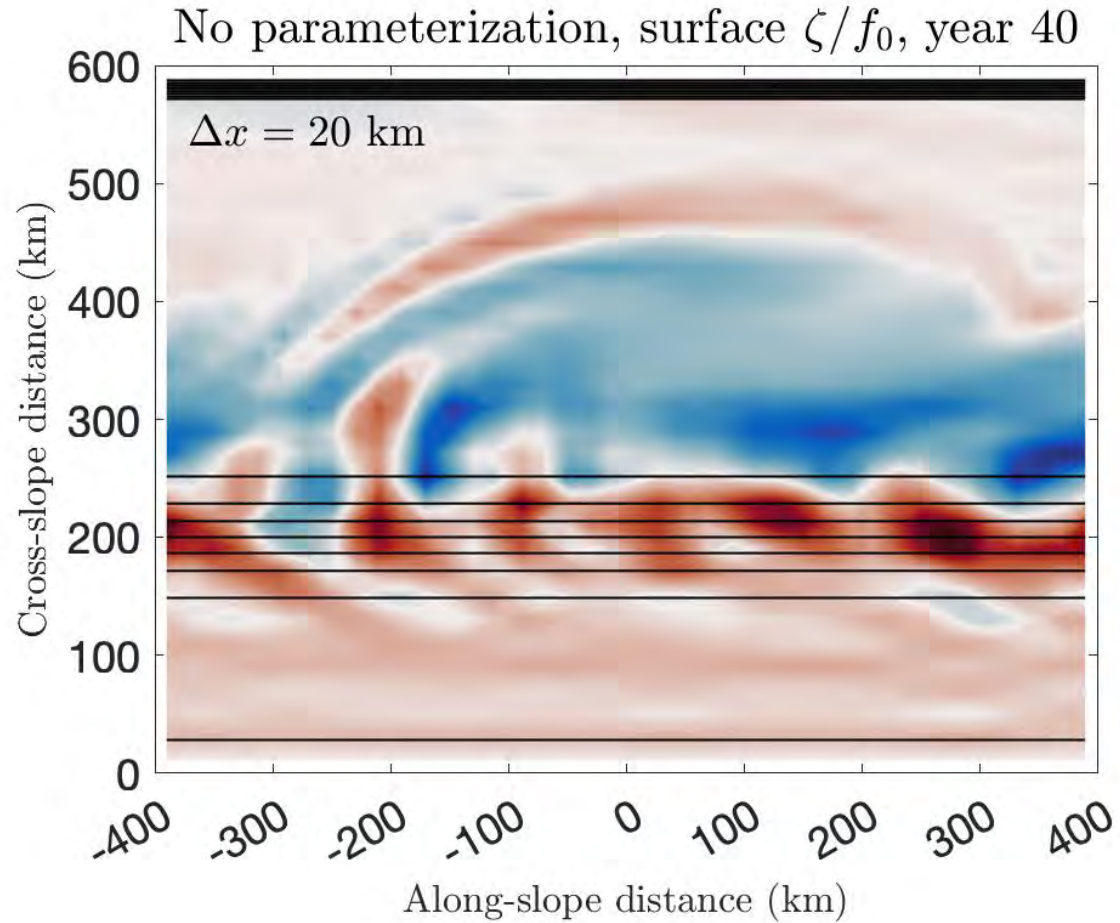
Calculating convergence; Imposing global momentum constraint (Green 1970) online

Model momentum equation(s)

Physics-based eddy closure(s): 3D coarse-grid parameterized flow



Physics-based eddy closure(s): 3D coarse-grid parameterized flow



Summary

- Prograde frontal systems are theorized and parameterized via GEOMETRIC adapted by an analytical function of the slope Burger number controlling efficiency of eddy buoyancy fluxes.
- Retrograde frontal systems are forced jointly by eddy buoyancy and momentum fluxes; the former can be quantified using a range of GM-based scalings, adapted via analytical functions of the topographic slope parameter.
- Eddy momentum fluxes across retrograde fronts depend linearly on eddy energy, and echo with barotropic eddy PV fluxes theorized in 2D topographic turbulence, driving prograde undercurrents.
- Machine learning approaches can augment physics-based, bathymetry-aware mesoscale eddy parameterizations by constraining eddy energy or/and forcing online.
- Scale separation may exist between eddy buoyancy and momentum forcing, which alludes to muting GM-based schemes but utilizing numerically-dissipated energy for driving subgrid-scale eddy momentum forcing in eddy permitting regimes across continental margins (*ongoing*).

Bathymetry-aware recipe (to be cont.):

[GM] Y. Wang, and A. L. Stewart, 2020, “Scalings for eddy buoyancy transfer across continental slopes under retrograde winds”, *Ocean Modelling*, 147, 101579.

[Redi] H. Wei, and Y. Wang, 2021, “Full-depth scalings for isopycnal eddy mixing across continental slopes under upwelling-favorable winds”, *Journal of Advances in Modeling Earth Systems*, 13, e2021MS002498.

[GM] H. Wei, Y. Wang, A. L. Stewart, and J. Mak, 2022, “Scalings for eddy buoyancy fluxes across prograde shelf/slope fronts”, *Journal of Advances in Modeling Earth Systems*, 14, e2022MS00322.

[Redi] C. Xie, H. Wei, and Y. Wang, 2023a, “Impact of parameterized isopycnal diffusivity on shelf-ocean exchanges under upwelling-favorable winds: offline tracer simulations augmented by artificial neural network”, *Journal of Advances in Modeling Earth Systems*, 15, e2022MS00342.

[GM + PV] C. Xie, H. Wei, and Y. Wang, 2023b, “Bathymetry-aware mesoscale eddy parameterizations across upwelling slope fronts: A machine learning-augmented approach”, *Journal of Physical Oceanography*, 53, 2861–289.

[GM] H. Wei, Y. Wang, and J. Mak, 2024, “Parameterizing eddy buoyancy fluxes across prograde shelf/slope fronts using a slope-aware GEOMETRIC closure”, *Journal of Physical Oceanography*, 54, 359–37.

[PV] J. He, and Y. Wang, In Press, “Multiple states of two-dimensional turbulence above topography”, *Journal of Fluid Mechanics*, DOI: 10.1017/jfm.2024.633.

[GM + Energy] P. Deng, and Y. Wang, In Press, “Distinct impacts of topographic versus planetary PV gradients on baroclinic turbulence”, *Journal of Physical Oceanography*, Early Online Release: <https://doi.org/10.1175/JPO-D-24-0014.1>.