Optimal Atmospheric Excitation of AMOC as Estimated via Application of the FDT

Grant Branstator, NCAR Andrey Gritsun, RAS

(Simplified, Quasi-gaussian) Fluctuation Dissipation Theorem (Leith, 1975; Deker&Haake, 1975; Risken, 1984)

Suppose have a discretized dynamical system with noise and a F-P eqn with unique solutions. Also assume gradients of the system PDF are well approximated by a Gaussian fit. Then the PDF-averaged response to weak forcing f is

$$r(t) = \int_{t_0}^{t} C(t - \tau)C^{-1}(0)f(\tau)d\tau$$
for $C(\tau) = \log \tau$ cov matrix

Sufficient data to find C

Atmospheric applications:

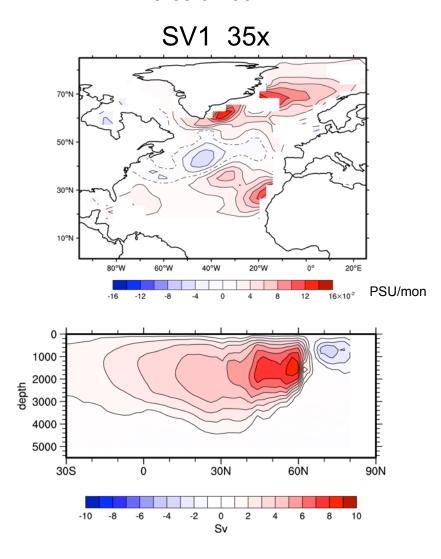
- * Gritsun, Branstator (2007)
- * Gritsun, Branstator, Majda (2008)
- * Liu et al. (2012)

CCSM4 T31, 3deg 8000 year control

$$\begin{bmatrix} T(t) \\ S(t) \\ u(t) \\ v(t) \end{bmatrix} = r(t) = \mathbf{M}_t \overline{f} = \mathbf{M}_t \begin{bmatrix} \dot{T} \\ \dot{S} \\ \dot{u} \\ \dot{v} \end{bmatrix}$$

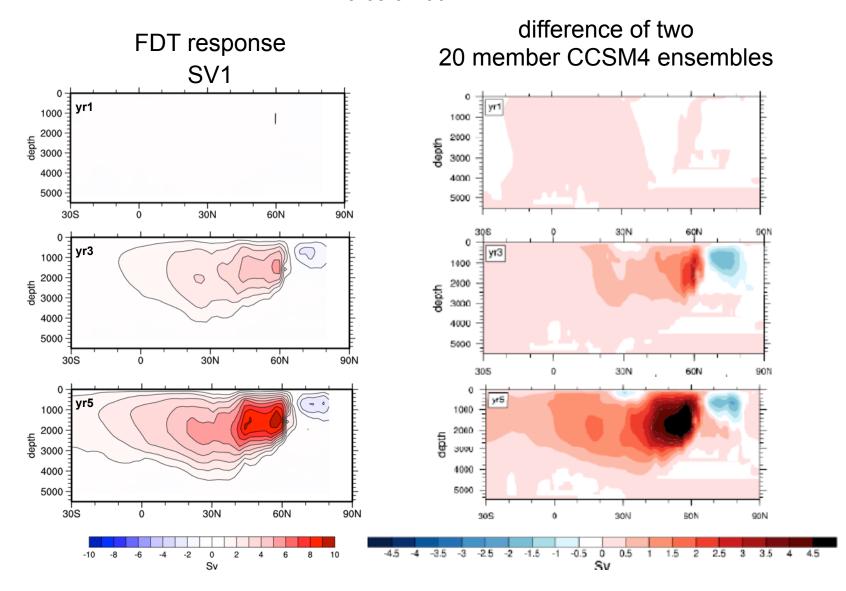
state of multivariate 3D fields is represented by 675 EOFs

Optimal N Atlantic Salinity Forcing of AMOC



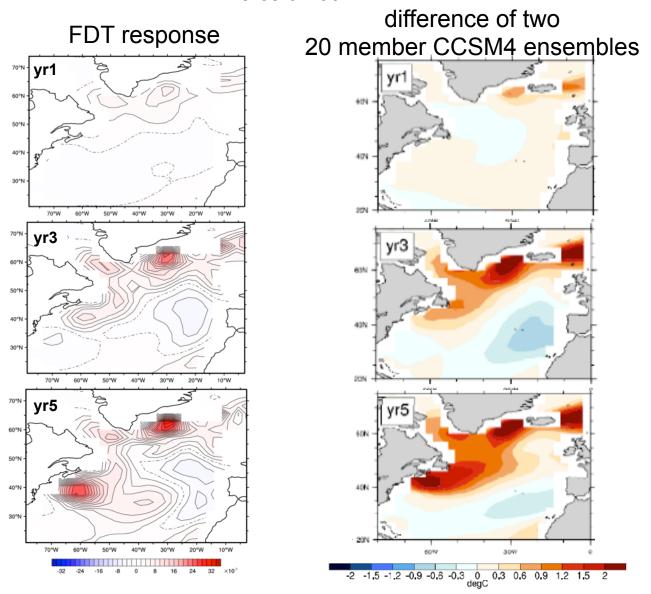
Optimal & CCSM4 Response of AMOC to Salinity Forcing

force 0-100m



FDT & CCSM4 SST Response to Optimal Salinity Forcing of AMOC

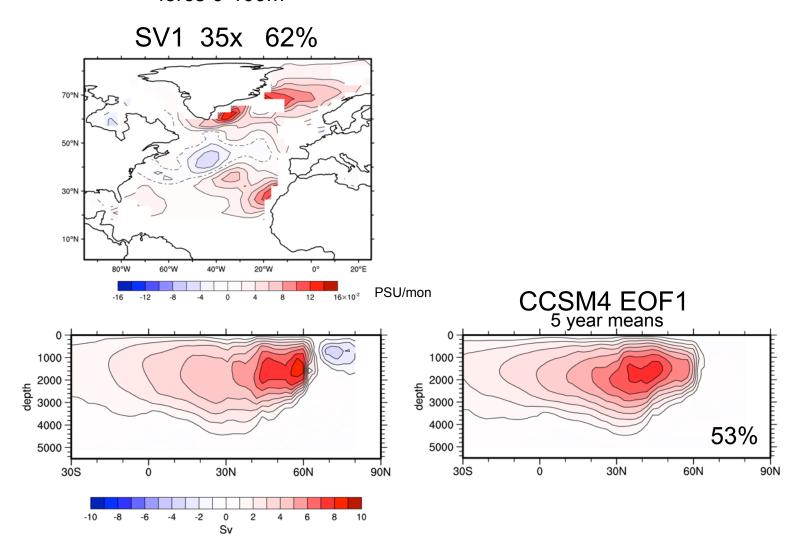
force 0-100m



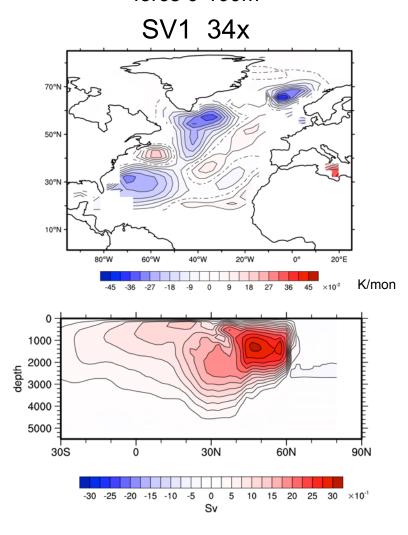
Remember

- FDT can be used to systematically study ocean response
- The leading patterns of AMOC variability do not depend on the existence of special atmospheric structures
- The effect of a given atmospheric pattern on AMOC depends sensitively on details of its structure and of ocean dynamics

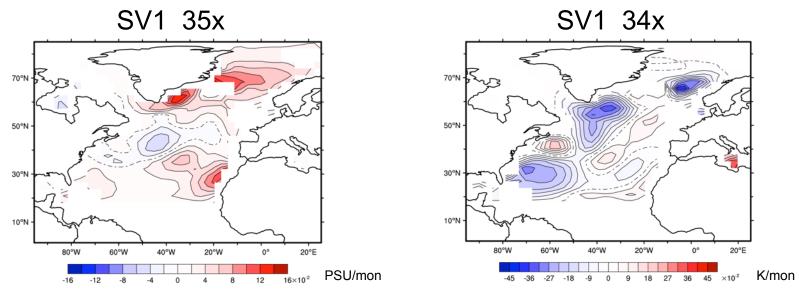
Optimal Salinity Forcing of AMOC

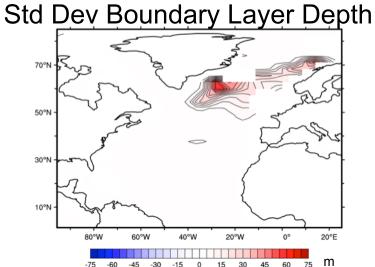


Optimal North Atlantic Temperature Forcing of AMOC

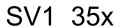


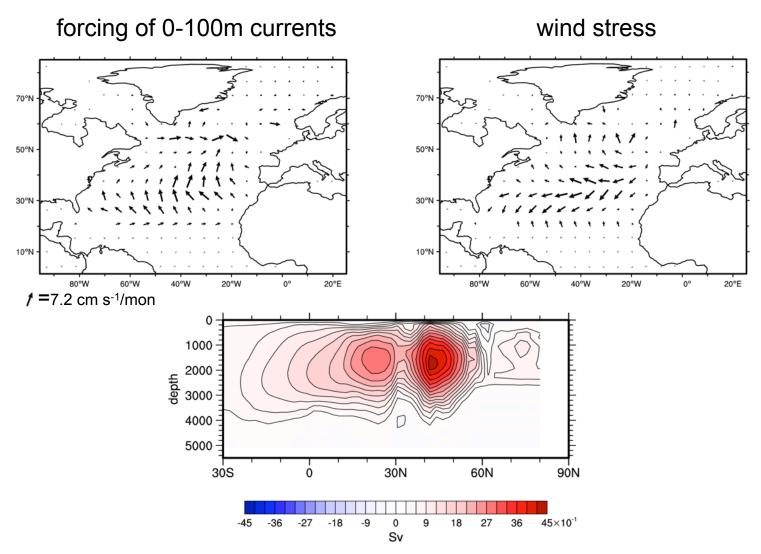
Optimal Salinity & Temperature Forcing of AMOC





Optimal Excitation of AMOC by Currents in the North Atlantic



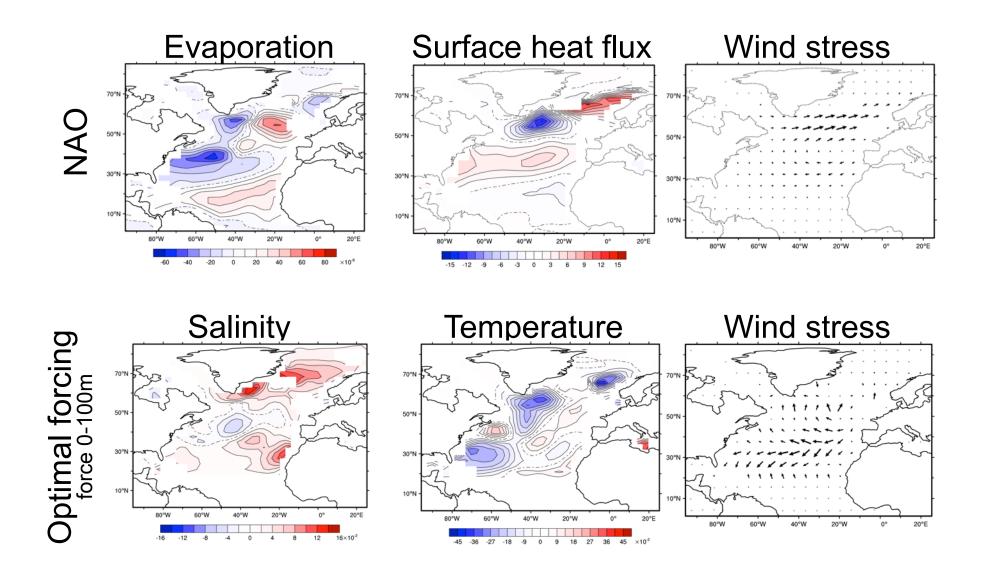


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Comparing NAO Fluxes & Optimal Forcing

5 year forcing; year 5 response



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