

Early Deglacial AMOC Decline and its Role for Atmospheric CO₂ Rise Inferred from Carbon Isotopes

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Questions:

1. Was AMOC reduced during the early deglacial Heinrich Stadial event 1 (HS1; 18-15 ka BP)?
2. Did AMOC decline contribute to rise in CO₂?

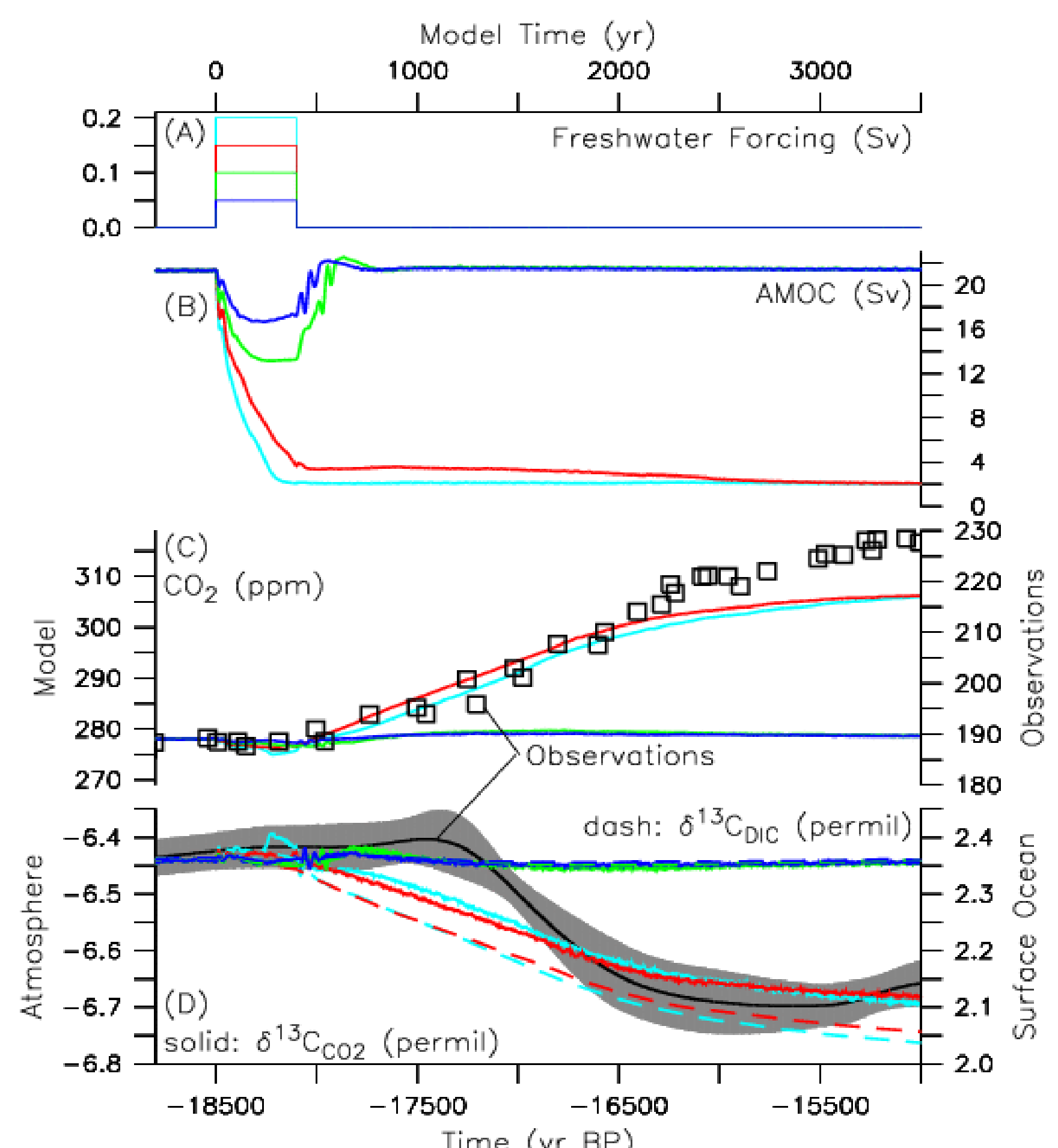


Fig. 1: Timeseries of forcing (A) and AMOC (B), CO₂ (C), δ¹³C_{CO2} (D) for four model experiments starting from pre-industrial conditions: FW0.05 (blue), FW0.1 (green), FW0.15 (red), and FW0.2 (cyan). Symbols and grey bar with black line show ice core data (bottom time scale). Note that absolute scales are different for model (left) and observations in (C), but the range is the same.

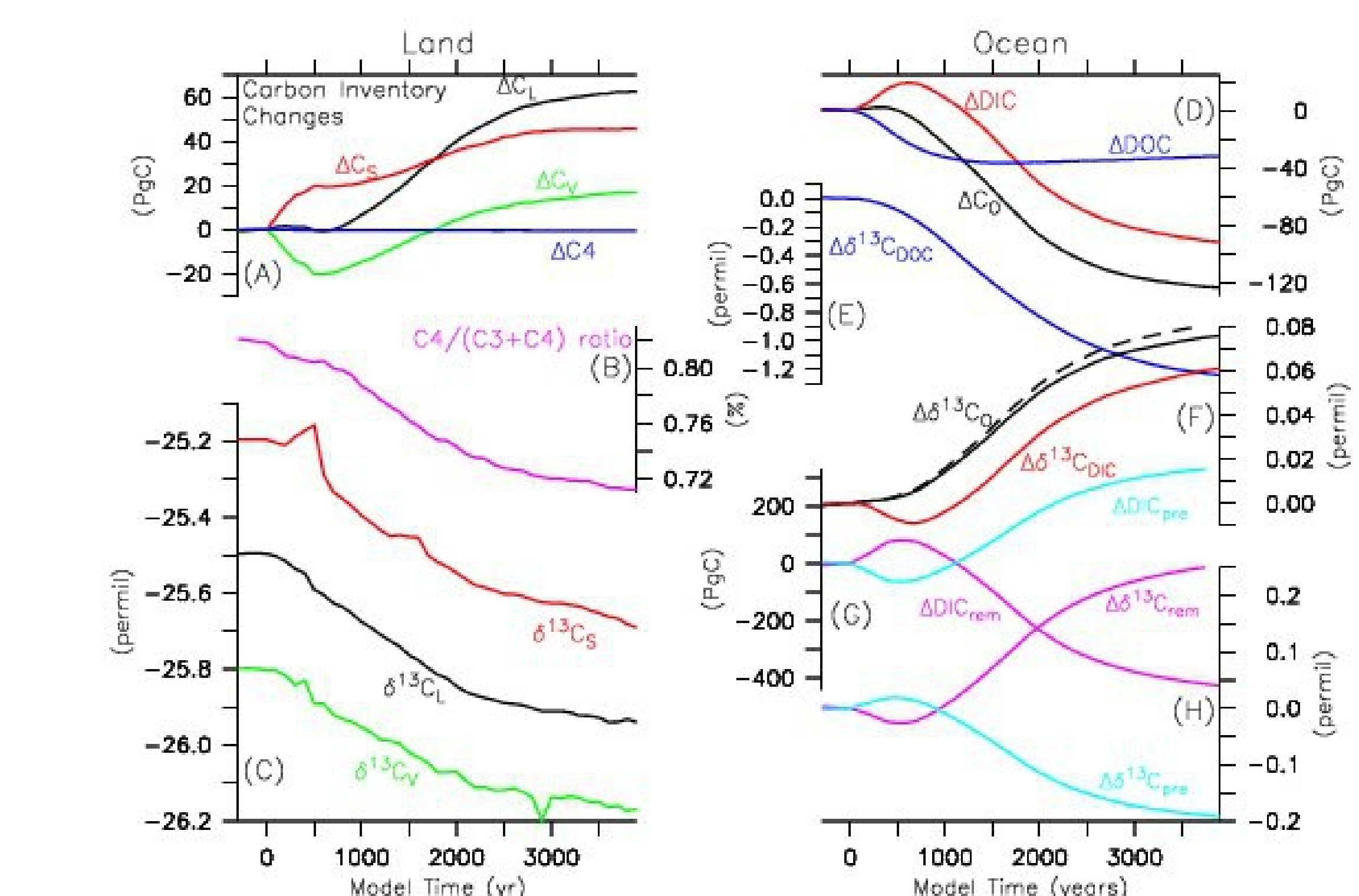


Fig. 2: Response of land and ocean carbon pools in experiment FW0.15.

Sediment Data: Compilation of 23 high-resolution carbon isotope records.

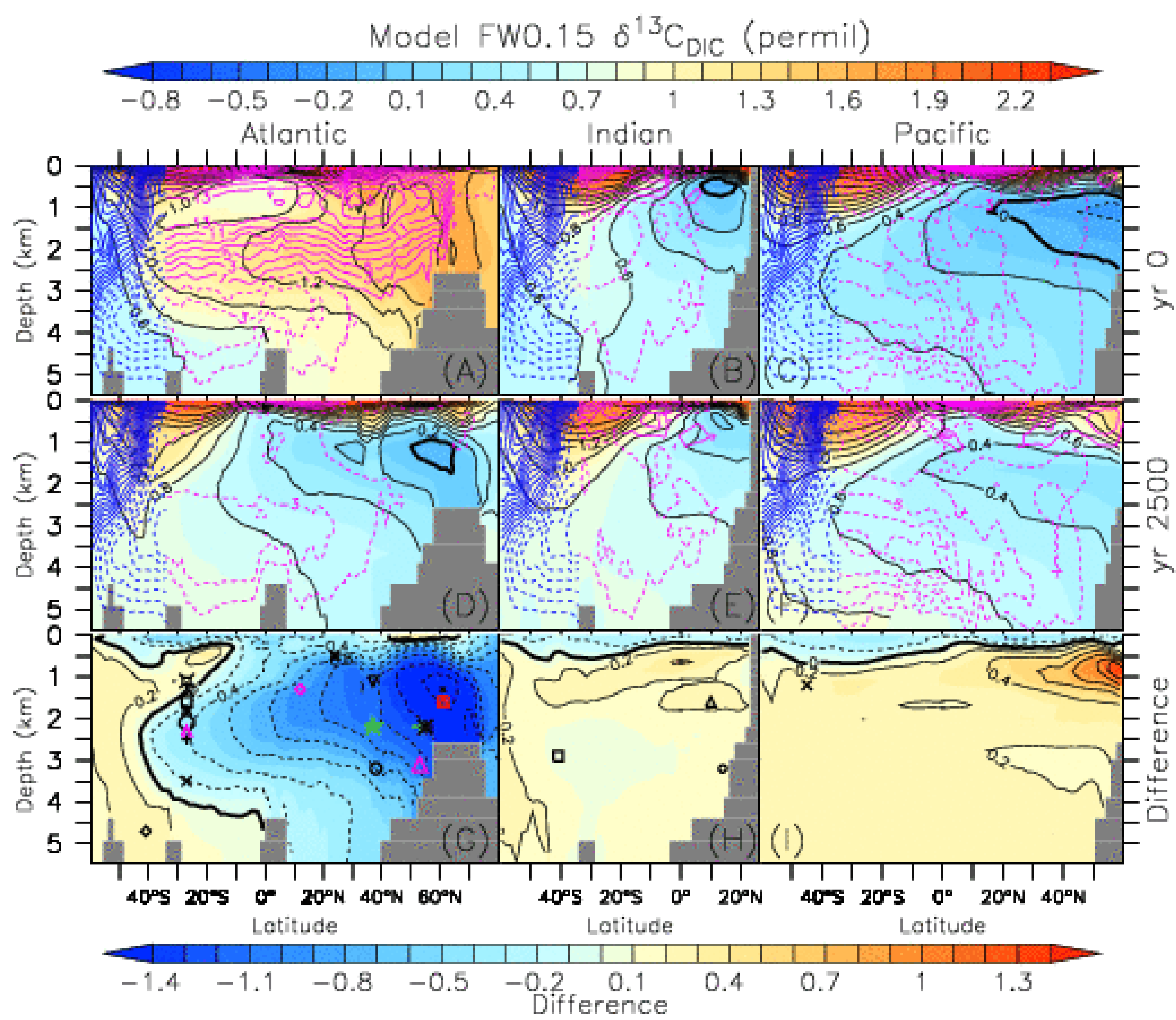


Fig. 4: δ¹³C_{DIC} distributions (color shading and black isolines) and MOC (purple and blue isolines) for model years 0 and 2,500 as well as their differences (bottom). The symbols in the bottom panels denote locations of high-resolution δ¹³C sediment data shown in Fig. 5.

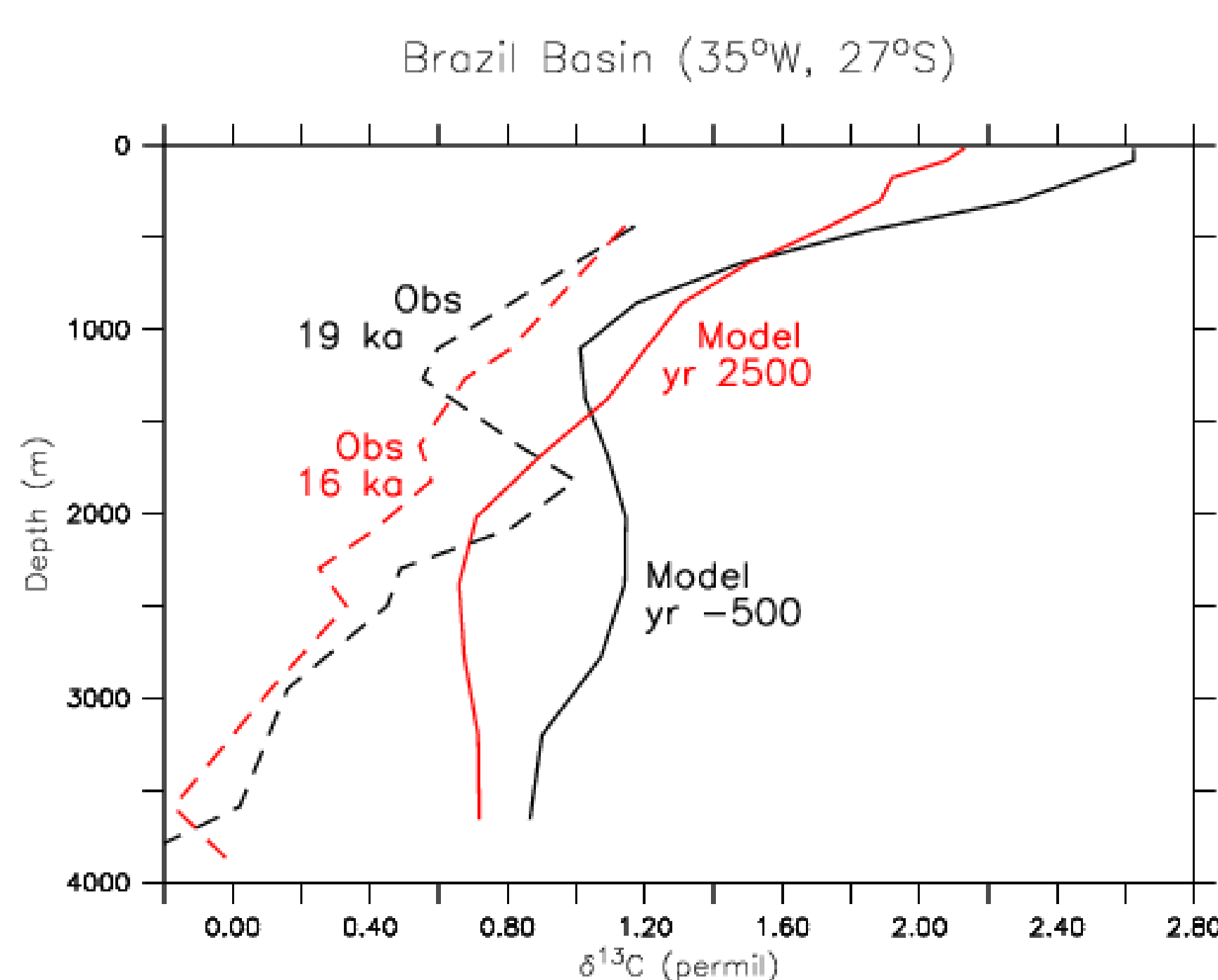


Fig. 7: Vertical profiles of carbon isotopes along the Brazil Margin from sediment data (dashed) and model FW0.15 (solid).

Conclusions:

1. AMOC was substantially reduced for a multi-millennial time period during HS1 compared to the prior LGM.
2. Early deglacial CO₂ rise and decrease in δ¹³C_{CO2} may have been caused by AMOC reduction.

Reference: Schmittner, A., and Lund, D. C. (2014) Carbon Isotopes Support Atlantic Meridional Overturning Circulation Decline as a Trigger for Early Deglacial CO₂ rise, *Climate of the Past Discussions*.

UVic Climate model:

3D ocean 1.8x3.6 deg, 19 levels
2D atmospheric Energy Moisture Balance
Land and ocean carbon cycle
Model of Ocean Biogeochemistry and Isotopes (MOBI 1.4) includes carbon isotopes in ocean, land, and atmosphere

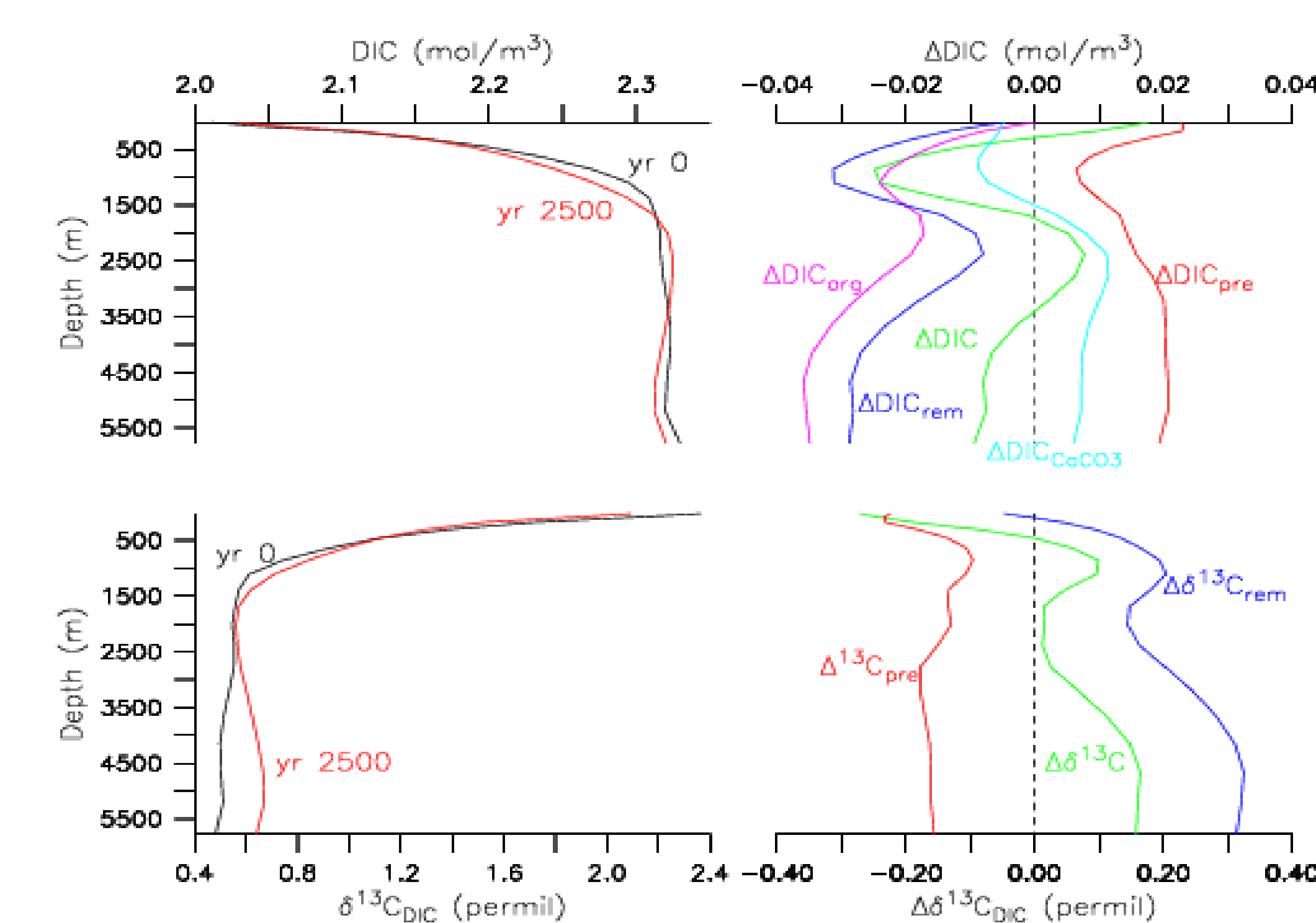
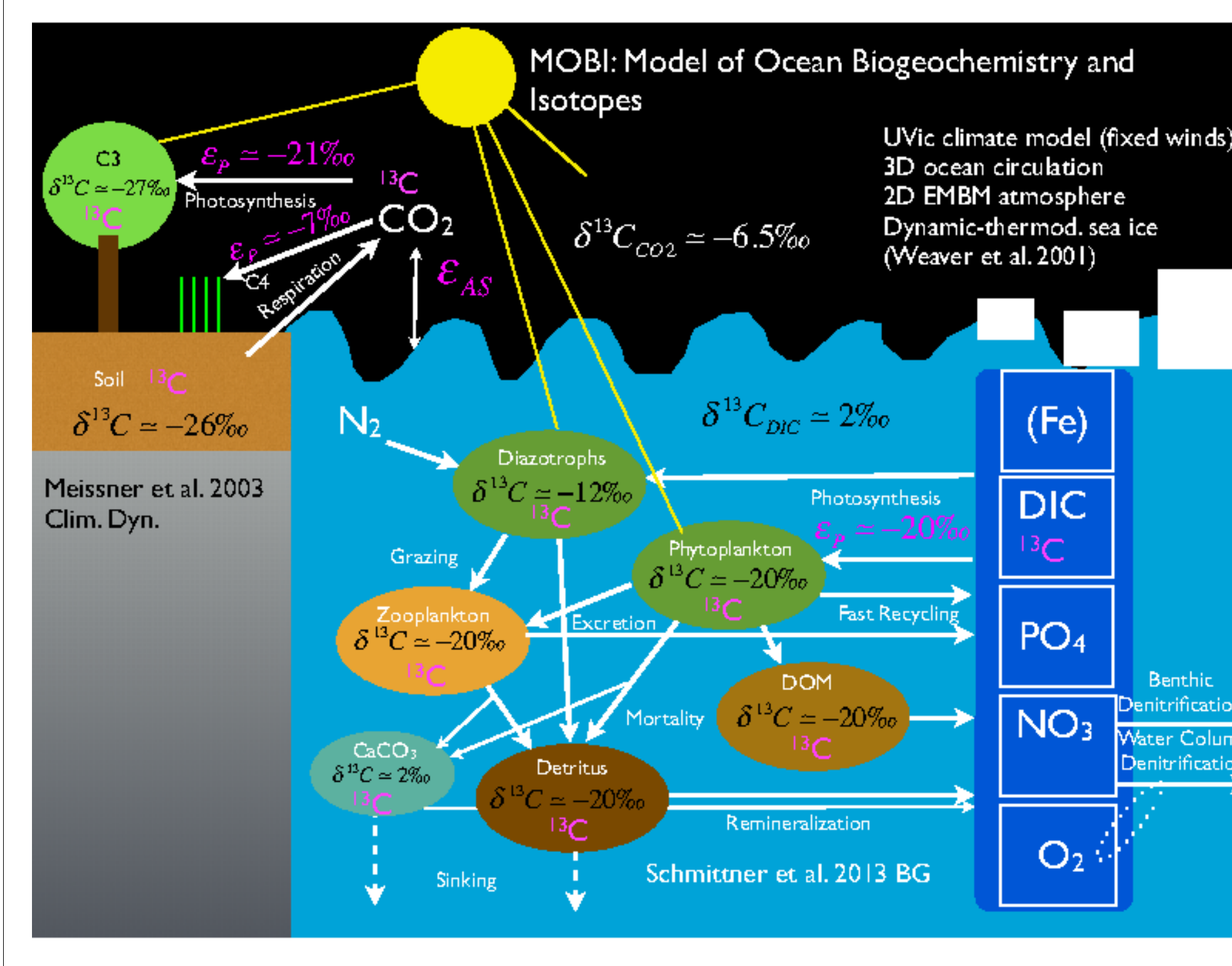


Fig. 3: Globally averaged vertical profiles at years 0 and 2,500 of DIC and δ¹³C_{DIC} for model experiment FW0.15 (left). Right panels show the differences as well as contributions from organic matter (org), CaCO₃, their sum (rem) and preformed (pre).

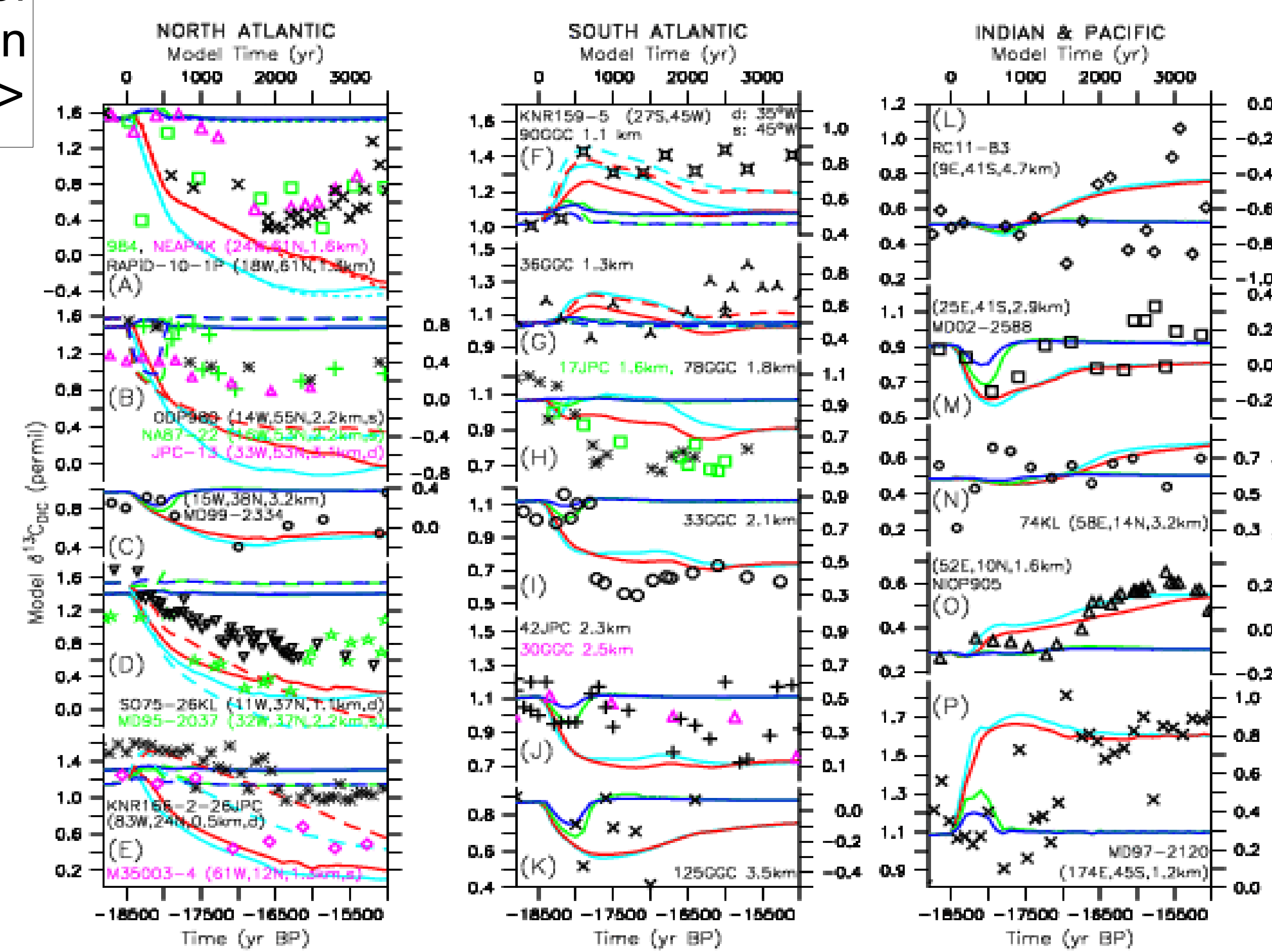


Fig. 5: Timeseries of sediment data (symbols) and model results at the corresponding grid points (color lines as in Fig. 1). See Fig. 4 for location of data.

Results:

Simulated AMOC collapse

- decreases efficiency of biological pump (Figs. 2, 3)
- causes CO₂ to increase, δ¹³C_{CO2} to decrease, consistent with ice core data (Fig. 1)
- has large effect on ocean δ¹³C distribution (Fig. 4)
- modeled changes agree with sediment data from LGM-HS1 (Figs. 5,6,7,8; Table)
- but model overestimates changes in North Atlantic. Why?
 - HS1 AMOC changes were smaller than simulated, or
 - different initial conditions (Fig. 7) could affect results

Table: Correlation coefficient (R) and root-mean-squared (RMS) error for model data comparison.

Model	R	RMS error
FW0.05	0.77	0.48
FW0.1	0.77	0.48
FW0.15	0.89	0.42
FW0.2	0.9	0.55

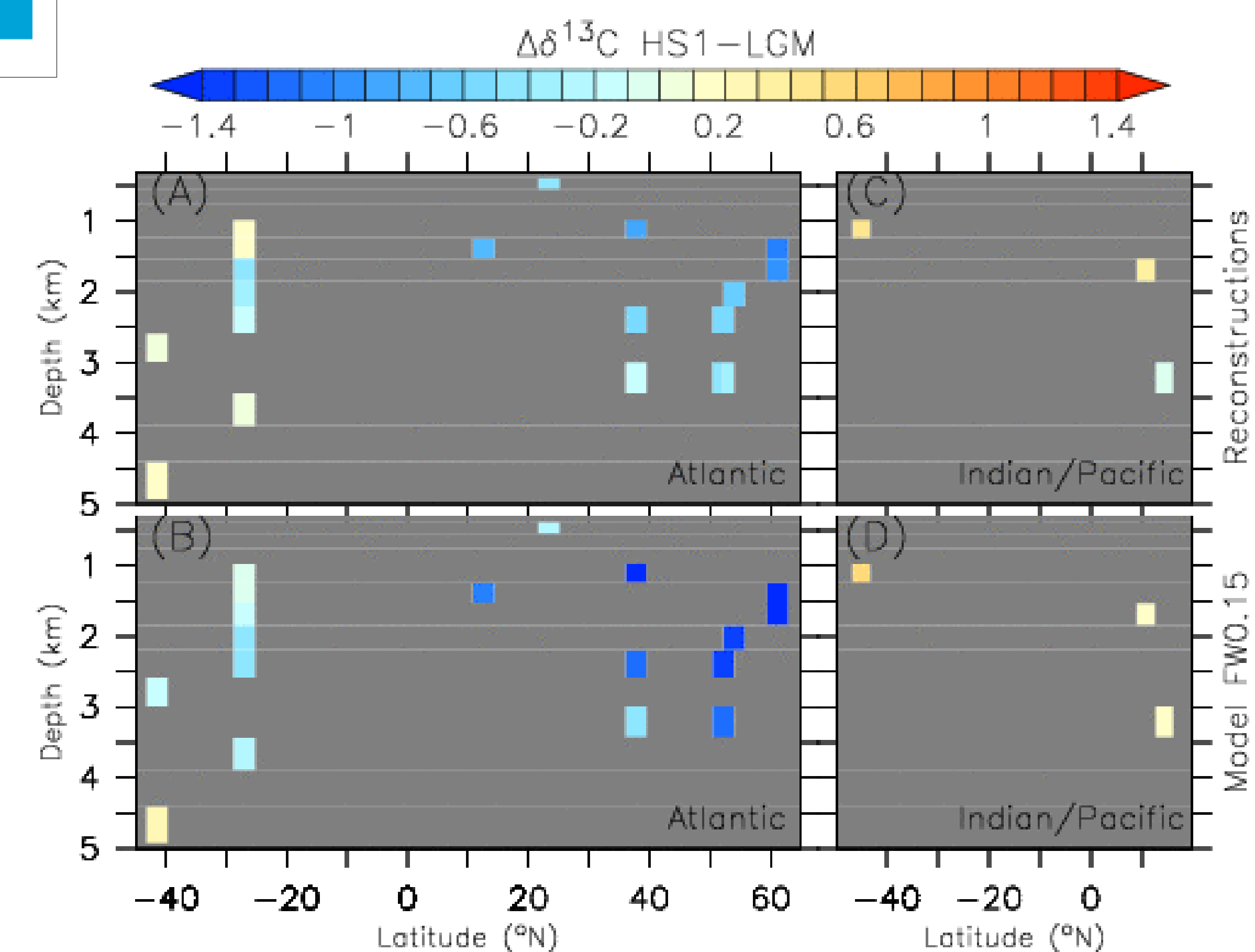


Fig. 6: Sections of changes in carbon isotopes between the Last Glacial Maximum (19-20 ka BP) and HS1 (15.5-16.5 ka BP) from sediment data (top) and experiment FW0.15 (bottom). Note that while the pattern of changes is similar the model overestimates the amplitude in the North Atlantic.

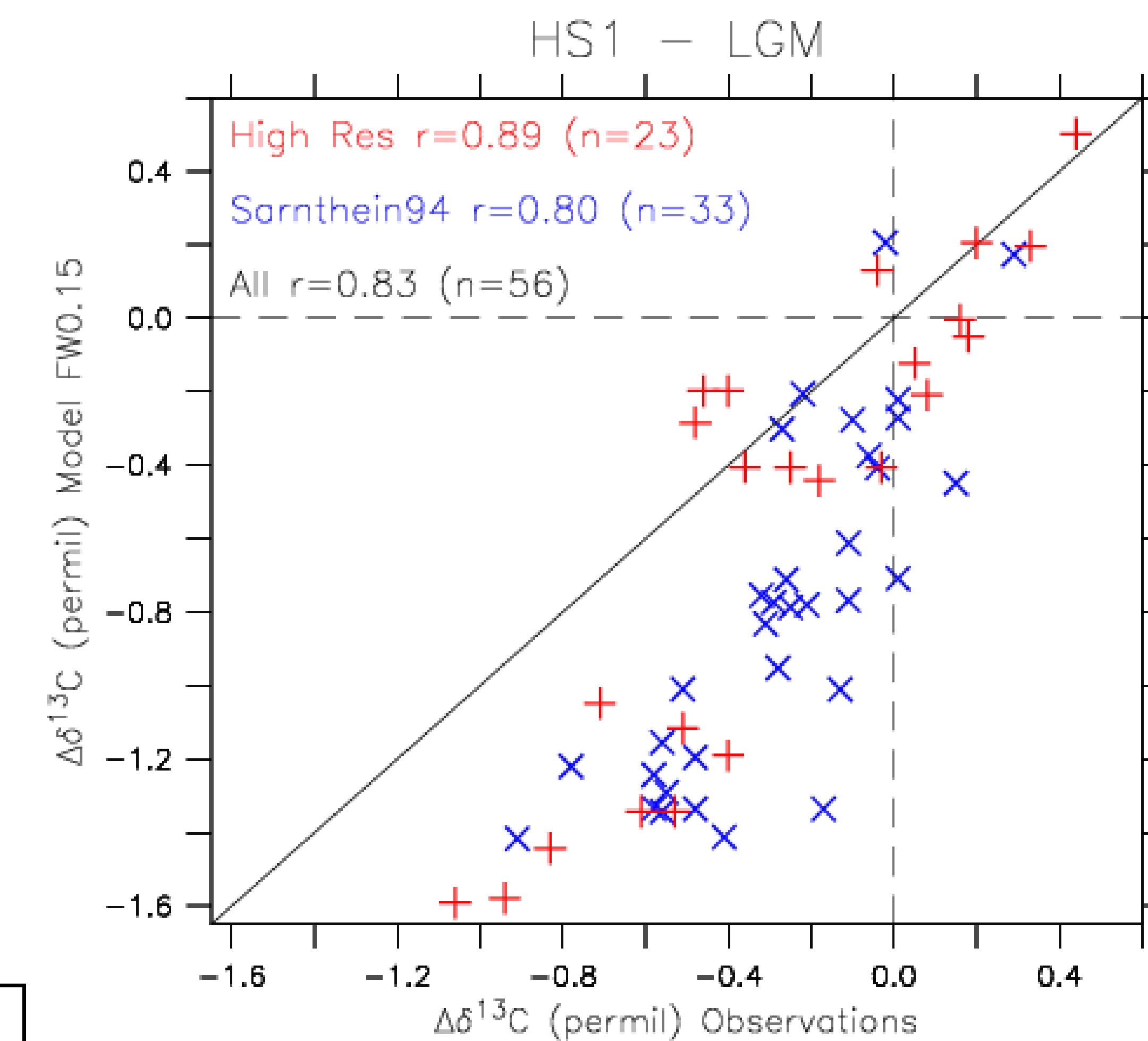


Fig. 8: Modeled vs observed carbon isotope changes from our high-resolution sediment data (red) and a previous study using lower resolution data (Sarnthein et al. 1994).

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